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To cite this version:

G. Lamble, R. Reeder, P. Northrup. Characterization of Heavy Metal Incorporation in Calcite by XAFS Spectroscopy. Journal de Physique IV Proceedings, 1997, 7 (C2), pp.C2-793-C2-797. 10.1051/jp4:1997238 . jpa-00255317

HAL Id: jpa-00255317 <https://hal.science/jpa-00255317v1>

Submitted on 4 Feb 2008

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Colloque C2, Supplément au Journal de Physique III d'avril 1997

Characterization of Heavy Metal Incorporation in Calcite by XAFS Spectroscopy

G.M. Lamble***, R.J. Reeder*** and P.A. Northrup***

* *Brookhaven National Laboratory, Upton, NY 11973, U.S.A.*

** *Present address, Wrence Berkeley National Laboratory, Berkeley, California 94720, U.S.A.*

*** *Department of Earth and Space Sciences, State University of New York, Stony Brook, NY, 11794-2100, U.S.A.*

Abstract: Zn and Co K-edge XAFS have been analyzed for spectra obtained from synthetic calcite that incorporated trace amounts of zn2+ and co2+ during growth from aqueous solutions at room temperature. **Zn** *and* Co concentrations in the calcite ranged from 250 to 2150 ppm. Analysis of the first shells indicates that the immediate coordination of both metal ions is approximately six oxygens at 2.11-2.12 **A.** These results are consistent with octahedral coordination at the unique Ca site in calcite. The observation that Co and Zn attain similar coordination in the bulk following incorporation suggests that their strongly contrasting surface site preferences are not expressed after being buried by subsequent growth layers.

1. INTRODUCTION

One of several important processes controlling the distribution and fate of heavy metal species dissolved in surface waters is their coprecipitation with authigenic minerals. Among these minerals, calcite $(CaCO₃)$ has attracted much interest because of the relative ease with which it forms in a broad range of environments and its tendency to incorporate a wide variety of divalent heavy metals. Coprecipitation results in the formation of dilute solid solutions, and, in sparingly soluble phases like calcite, heavy metals that have been sequestered by coprecipitation may remain isolated from further reaction. Relative incorporation preferences of divalent metal ions in calcite have been known qualitatively for some time [I], and are generally represented by conventional distribution coefficients [2, 3], where $K_d > 1$ corresponds to preferential enrichment in the precipitating solid (compatible behavior) and K_d < 1 corresponds to discrimination by the solid (incompatible behavior). In general, differences in K_d values among metal ions of like charge are often explained on the basis of ion size relative to that of the host metal in the bulk structure (Ca for calcite). However, such ion size correlations for divalent metals incorporated in calcite are inconsistent, and similarly sized metal ions may exhibit contrasting behavior $(e.g., K_{\text{Sr}} < 1$ and $K_{\text{Pb}} > 1)$, suggesting that electronic configuration plays a role. Furthermore, the range of reported K_d values for an individual divalent metal ion may vary up to an order of magnitude. It has been shown that an additional factor influencing the extent of metal ion uptake is the growth rate [4, 5], and some of the variability in reported K_d values apparently reflects such differences.

Recent experimental work has demonstrated that the type and distribution of different surface sites on calcite play a major role in determining the relative extent of metal incorporation during growth [6, 7]. Structurally nonequivalent growth steps on calcite (1014) faces-moving in symmetrically nonequivalent directions on the face-have been shown to incorporate metal impurities in different amounts. In some cases the incorporation difference is as large as a factor of ten. The nonequivalent growth steps occur on the vicinal faces of polygonized growth hillocks, which indicate growth **occurred** by the spiral mechanism. Kink sites exposed within nonequivalent growth steps have different local coordination depending on the step and kink orientation. Differences in coordination of such kink sites have been identified as the cause for different incorporation preferences for impurity metals. In situ **AFM** studies of layer growth on calcite (1014) surfaces also support the existence of structurally distinct growth steps and incorporation sites within them **[8,** 91. Consequently, surface site structure and distribution are additional important factors influencing the extent of metal uptake during crystallization.

Surface growth steplsite preferences for heavy metal incorporation into calcite show some surprising inconsistencies among different ions. A notable example is the strongly contrasting behavior shown by Co^{2+} and Zn^{2+} . Despite nearly identical sizes ($r(Co^{2+}) = 0.745 \text{ Å}$; $r(Zn^{2+}) = 0.74 \text{ Å}$), these metal ions show exactly opposite surface step/site preferences for incorporation. Co exhibits a preference similar to divalent ions including Mg, Mn, and Fe, which are smaller than host Ca. In contrast, Zn exhibits a preference that is similar to those shown by larger divalent ions such as Sr, Ba, and Pb. The presence of multiple, structurally distinct surface sites on calcite, combined with observed differences in step/site selectivity, raises fundamental questions concerning the mechanism of impurity metal discrimination and incorporation. One important aspect of this relates to the local coordination of the metal impurities within the bulk crystal following incorporation. More specifically, one can ask: do Co and Zn have similar coordination in the bulk? The present work addresses this particular aspect using XAFS spectroscopy. The initial results reported here are limited to Co and Zn because of their strongly contrasting incorporation preferences. Future work addresses Ba and Pb incorporation.

The calcite structure has a single metal site normally occupied by Ca. Its first coordination shell consists of six oxygens, all at 2.36 Å, in the form of an octahedron, only very slightly trigonally elongated. The CaO₆ octahedra share corners via planar CO₃ groups, which are arranged in layers separating cation layers along the three-fold axis. Size considerations alone would favor both Co and Zn in octahedral coordination, substituting for Ca. And $CoCO₃$ and $ZnCO₃$ compounds have the calcite structure, although equilibrium solubilities of $CoCO₃$ and $ZnCO₃$ in calcite are extremely limited. However, because both Co and Zn exhibit compatible incorporation behavior *(i.e.,* $K_d > 1$), natural calcites may contain as much as several thousand pprn Co or Zn.

2. METHODS

2.1. Coprecipitation Experiments

Calcite single crystals (up to 500 μ m) were grown from Ca-NH₃-CO₃-Cl aqueous solutions at room temperature according to the method described by Paquette and Reeder [6]. Solutions were doped with individual divalent metal species (Co or Zn), which were incorporated into the calcite during growth. Crystals were removed from growth solutions, rinsed with deionized water to remove any metals adsorbed to the surface, and dried. Concentrations and distributions of divalent metals were analyzed by synchrotron X-ray fluorescence microanalysis at NSLS (beamline X-26A) and by EPMA. Two Codoped calcites selected for XAFS work contained 375 and 2150 ppm Co, respectively. In addition two Zn-doped samples were selected, containing 250 and 1200 pprn Zn, respectively. Crystals were ground in a pestle and mortar in alcohol or acetone and allowed to dry.

2.2. Other Samples and Reference Compounds

A natural Zn-containing calcite from Franklin, NJ, USA, was used for comparison with the synthetic calcite. Reagent ZnO and natural $ZnCO₃$ (smithsonite) from Kelly Mine, NM, USA, were used as reference compounds for Zn^{2+} . CoO was used as a reference compound for Co^{2+} . At the time of writing, a suitable $CoCO₃$ sample was not yet available, and a natural Cocontaining calcite from Katanga, Zaire, was incompletely analyzed.

2.3. XAFS Experiments and Analysis

Co and Zn K-edge XAFS spectra were collected at beamline X-11A of the National Synchrotron Light Source, Brookhaven National Laboratory, USA. Powder samples were pressed into aluminum sample holders, which were mounted on a cold finger cooled with liquid nitrogen. Si(ll1) crystals were used in the monochromator, which was calibrated using the first inflection point in edge scans of Co or Zn metal foils. The energy resolution at the edges of interest was approximately 1-1.5 eV. Higher order harmonics were rejected by slightly detuning the angle of one monochromator crystal with respect to the other until the harmonic contribution was negligible. One data set was obtained using a focusing-crystal arrangement that increased the flux at the sample [lo]. Absorption measurements on the dilute heavy metal-containing calcite samples were made by monitoring the fluorescence yield above the k-edge of interest using a Canbena 13-element *Ge* solid-state detector. Spectra were obtained for reference compounds in transmission mode.

After background subtraction utilizing a modified cubic spline method, data were normalized and converted to k-space. Background-subtracted, normalized **XAFS** data were fined in R-space using the program **FEFFIT** [l 11 with model calculations from FEFF6 [12,13]. Fitting used both the imaginary and real parts of the Fourier transforms and minimized the difference between the background-subtracted, normalized data and the model $\chi(R)$, where $\chi(R)$ is the Fourier transform calculated from $\chi(k)$ by FEFF6. E_0 was treated as a variable parameter during fitting.

3. RESULTS AND DISCUSSION

3.1. Coordination of Zn in Calcite

Both Zn-doped synthetic calcite samples (250 pprn and 1200 ppm) showed essentially identical absorption spectra. In addition, these spectra are nearly identical to the natural Zn-containing calcite. Figure 1 shows the background-subtracted raw XAFS data for both synthetic calcite samples, the natural calcite, and the $ZnCO₃$ reference compound. Although some of the features seen in the Zn-containing calcite spectra resemble those in $ZnCO₃$, the latter spectrum is clearly different. Figure 2 shows the corresponding Fourier transform magnitudes for the background-subtracted raw data shown in Figure 1. The similarity among the Zn-containing calcite samples is evident, and they are distinct from $ZnCO₃$.

Single-shell fitting of the Zn-containing samples was chosen to be over the R-space range 0.8 to 2.2 Å (not phase-shift corrected) and over the k-space range 2.0 to 12.0 Å⁻¹. Figure 3 is a representation in k-space of the first-shell contribution for the low-concentration (250 ppm) synthetic Zn-containing calcite compared with the corresponding fit using **FEFFIT.** Fits for the synthetic Zn-containing calcites yielded first shells of 6.8 (\pm 1) oxygens at 2.12 (\pm 0.02) Å (Debye-Waller type factor σ^2 = 0.003 Å²) for the lower concentration and 6.6 (±1) oxygens at 2.12 (±0.02) Å (σ^2 = 0.003 Å²) for the higher concentration sample. Fitting for the natural Zn-containing calcite gave 6.8 (± 1) oxygens at 2.12 (± 0.02) Å ($\sigma^2 = 0.004$ Å²). These values are consistent with Zn in octahedral coordination. The 2.12 A distance compares well with the Zn-0 distance of 2.111 Å determined for ZnCO₃ by X-ray diffraction methods [14]. Multishell fits based on a model with Zn in the Ca position of calcite yielded good agreement with the observed spectrum; further details will be given elsewhere.

Figure 1: The k-weighted, background-subtracted raw XAFS data for Zn-containing samples: (a) natural Zn-containing calcite, (b) synthetic Zn-containing calcite (250 ppm Zn), (c) synthetic Zn-containing calcite (1200 ppm Zn), and (d) ZnCO₃ reference compound.

Figure 2: Fourier transform magnitudes of the background-subtracted raw data shown in Figure 1. Sample labels are given in Figure 1 caption.

Figure 3: The k-space representation of the first-shell contribution of the low-concentration (250 ppm) Zn-containing calcite (solid line) compared with that of the fit obtained in R-space using **FEFmT** (dashed line).

3.2. Coordination of Co in Calcite

Results obtained for the synthetic Co-containing calcite samples are remarkably similar to those for the synthetic Zn-doped calcite. The two synthetic Co-doped samples produced similar spectra. A single-shell fit yielded a first shell of 5.6 (± 1) oxygens at 2.11 (\pm 0.02) Å (σ ² = 0.006 Å²). The agreement with the Co-O distance determined by X-ray diffraction, 2.110 Å [15], is quite good. A multishell fit provides further detail of the local coordination and will be reported elsewhere, along with results for a natural Co-containing calcite.

3.3. Implications

These results confirm that both Co and Zn impurities are present in calcite in sixfold coordination with oxygen atoms, consistent with their substitution for Ca in the unique octahedral site. The Co-0 and Zn-0 distances determined here very closely match values found by X-ray structure refinements for CoCO₃ and ZnCO₃. Although preliminary, this observation suggests that a metal ion substituting in a site normally occupied by a considerably larger ion still maintains the characteristic M-0 distance observed in compounds where a site is filled entirely by that metal. The structural topology of calcite apparently does not hold the metal site open when accommodating smaller metal ions. Such behavior is important for understanding both structural and thermochemical properties of dilute solid solutions. Further examination of this behavior over a greater range of substitution is certainly warranted.

The observation that Co and Zn impurities have essentially identical coordination in the bulk calcite does little to explain their strongly contrasting preferences for different surface sites. Apparently, both ions "see" the same types of surface sites during growth, and despite opposite preferences for the differing coordination among them, their final state after incorporation is the same. It seems that upon "burial" by additional layers, any differences in surface conformation between Zn and Co fade or become overwhelmed once within the bulk. Similar results might be expected for other small ions. However, the question remains as to whether larger ions, such as Ba and Pb, achieve similar coordination. An XAFS study of Sr incorporation in calcite indicates that octahedral coordination is indeed possible for that ion [16].

Figure 4: The k-space representation of the first-shell contribution of the high-concentration (2150 ppm) Co-containing calcite (solid line) compared with that of the fit obtained in R-space using FEFFIT (dashed line).

Acknowledgments

This work was performed at beamline X-11A at NSLS, which is supported by the US DOE Office of Basic Energy Sciences under contracts DE-AC02-76-CH00016 and DE-ASO5-80-ER10742.

References

- [I] Veizer, J., Carbonates: Mineralogy and Chemistry *Rev. in Mineral.* **11** (1983) 265-299.
- [2] Morse, J. W. and Bender, M. L., *Chem. Geol.* 82 (1990) 265-277.
- [3] Mucci, A. and Morse, J.W., *Aquatic Sciences* 3 (1990) 217-254.
- [4] Lorens, R.B., *Geochim. Cosmochim. Acta* **45** (1981) 553-561.
- [S] Dromgoole, E.L. and Walter, L.M., *Chem. Geol.* **81** (1990) 311-336.
- [6] Paquette, J. and Reeder, R.J., *Geochim. Cosmochim. Acta 59* (1995) 735-749.
- [7] Reeder, R.J., *Geochim. Cosmochim. Acta* **60** (1996) 1543-1552.
- [8] Gratz, A.J., Hillner, P.E., Hansma, P.K., *Geochim. Cosmochim. Acta* **57** (1993) 491-497.
- [9] Stipp, S.L., Eggleston, C.M., Nielsen, B.S., *Geochim. Cosmochim. Acta* **58** (1994) 3023-3033.
- [10] Lamble, G.M., and Heald, S.M., *Rev. Sci. Instrum.* **63** (1992) 880-884.
- [ll] M. Newville, Dept. of Physics, Univ. of Washington, Seattle, WA, 98195, USA, See also Stern, E.A., *Physica* B **2081209** (1995) 117.
- [12] Rehr, J.J., Mustre de Leon, J., Zabinsky, S.I., Albers, R.C., *3. Amer. Chem. Soc.* **113** (1991) 5135-5140.
- [13] Mustre de Leon, J., Rehr, J.J., Zabinsky, S.I., Albers, R.C., *Phys. Rev.* **B44** (1991) 4146-4156.
- [14] Effenberger, H., Mereiter, K., Zemann, J., *Zeit.* **j** *Kristallogr.* **156** (1981) 233-243.
- [15] Pertlik, F. *Acta Cryst.* C42 (1986) 4-5.
- [16] Pingitore, N.E., Lytle, F.W., Davies, B.M., Eastman, M.P, Eller, P.G., Larson, E.M., *Geochim. Cosmochim. Acta* **56** (1992) 1531-1538.