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# Training Local Models from Reanalysis Data to Estimate Reference Evapotranspiration with Fewer Onsite Sensors, an Evaluation in West Africa

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# 1 Title

2 Training local models from reanalysis data to estimate reference evapotranspiration with fewer  
3 onsite sensors, an evaluation in West Africa

## 4 Authorship

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## 11 Abstract

12 This study addresses the critical need for accurate reference evapotranspiration (ET<sub>0</sub>)  
13 estimation in data-scarce environments such as West Africa, where rapid population growth  
14 and climate change intensify water resource challenges. Traditionally, computing ET<sub>0</sub> with the  
15 FAO-56 Penman-Monteith method requires multiple meteorological inputs—such as solar  
16 radiation, humidity, and wind speed—collected from costly, fully instrumented weather  
17 stations. However, the availability and maintenance of such equipment can be prohibitive in  
18 remote regions. To overcome these constraints, we explore the potential of machine learning  
19 (ML), specifically XGBoost (XGB) models, trained on NASA Power climate reanalysis  
20 datasets. Our approach relies on a limited subset of easily measured in situ variables—daily  
21 minimum and maximum temperatures and rainfall—to estimate ET<sub>0</sub>. Departing from standard  
22 ML practices that depend on short-term, site-specific data, we leverage the extensive historical

23 depth and broad spatial coverage of reanalysis products. We trained and validated twenty  
24 locally adapted XGB models using measurements from twenty diverse West African weather  
25 stations. Our results show that certain XGB model configurations, notably those incorporating  
26 temperature and rainfall data, can approximate ET<sub>0</sub> estimates from the FAO-56 Penman-  
27 Monteith equation with median RMSE values frequently below 1 mm/day—levels comparable  
28 to commonly employed empirical formulas. This finding demonstrates that minimal on-site  
29 instrumentation, combined with ML and reanalysis data, can effectively support irrigation  
30 scheduling and enhance water-use efficiency under varying agro-ecological conditions. To  
31 foster broader implementation, we have released our XGB-ET<sub>0</sub> code under the GPLv3 licence  
32 (<https://github.com/SARRA-cropmodels/RF-ET0>), enabling researchers and practitioners to  
33 locally train and deploy these models, thereby improving ET<sub>0</sub> estimation and sustainable  
34 agricultural water management in resource-limited settings.

## 35 Keywords

- 36 • Evapotranspiration estimation
- 37 • Machine learning
- 38 • Precision irrigation
- 39 • XGBoost models
- 40 • Reanalysis datasets
- 41 • FAO-56 Penman-Monteith
- 42 • West Africa

## 43 1. Introduction

44 Accurately estimating actual evapotranspiration (ET<sub>a</sub>) is critical for optimizing irrigation  
45 practices and promoting sustainable agricultural water management. This need is particularly  
46 acute in West Africa, where growing population pressures and increasing variability in water

47 availability caused by climate change (Arnell and Lloyd-Hughes, 2014; Ayers et al., 2023; Brouillet  
48 and Sultan, 2023; Satgé et al., 2020) have intensified demands on limited water resources.  
49 Precision irrigation (PI), which tailors water inputs to compensate for ET<sub>a</sub>, offers a promising  
50 solution to reduce excessive irrigation, even in humid environments (Adeyemi et al., 2017;  
51 González Perea et al., 2018). Beyond irrigation scheduling, reliable ET<sub>a</sub> estimates play a  
52 pivotal role in understanding the water cycle, modeling water availability, and predicting  
53 changes in hydrological conditions. ET<sub>a</sub> is also a key input in process-based crop models,  
54 helping farmers anticipate the impact of different irrigation strategies on yields and overall  
55 water usage (Gu et al., 2020; Lopez-Jimenez et al., 2022; Pereira et al., 2020). Consequently,  
56 improving ET<sub>a</sub> estimation methods under data-scarce conditions is essential for both  
57 advancing scientific understanding and informing practical water management decisions in the  
58 region.

59         However, direct experimental methods for estimating ET<sub>a</sub> often require complex  
60 instrumentation, careful handling, and trained personnel, making them difficult to deploy and  
61 maintain (Li et al., 2009; Valayamkunnath et al., 2018). As a result, indirect ET<sub>a</sub> estimation  
62 methods have been developed, including several mathematical formulations based on  
63 meteorological variables (compiled in (McMahon et al., 2013)). Some of these methods rely  
64 on the concept of reference crop evapotranspiration (ET<sub>0</sub>), which represents the amount of  
65 water lost to the atmosphere by a reference healthy, well-watered grass or alfalfa cover; this  
66 ET<sub>0</sub> is further modulated by a crop and growth stage-specific coefficient (crop coefficient,  
67 symbolised as K<sub>c</sub>), to provide an estimate of actual evapotranspiration (noted ET<sub>a</sub>). This ET<sub>0</sub>,  
68 which represents the climate forcing component and is a widely used variable in crop water  
69 modelling and management, can be computed based on air temperature, solar radiation,  
70 relative humidity, and wind speed using the FAO-56 standardized Penman-Monteith equation  
71 (FAO-56-PM) (Richard G. Allen et al., 1998), the most widely-used procedure for estimating  
72 the ET<sub>0</sub> value.

73         However, obtaining all the meteorological variables required by the FAO-56-PM  
74 equation can be challenging. The cost, maintenance, and limited availability of fully

75 instrumented weather stations restrict the widespread calculation of  $ET_0$ , especially in  
76 developing countries (Djaman et al., 2016, 2015; Fisher and Pringle lii, 2013; Valipour, 2015).  
77 To circumvent these constraints, a range of empirical  $ET_0$  estimation models requiring fewer  
78 input variables has emerged (McMahon et al., 2013). While these simplified approaches can  
79 reduce the need for complex instrumentation, their accuracy varies markedly depending on  
80 local agroclimatic conditions (Djaman et al., 2015; Tabari and Talaei, 2011; Xystrakis and  
81 Matzarakis, 2011). Consequently, selecting and calibrating an appropriate empirical model for  
82 a given site remains a non-trivial, expert-dependent process that complicates routine  $ET_0$   
83 estimation and its integration into decision support services.

84 This study explores whether leveraging climate reanalysis datasets can help overcome  
85 these challenges. We hypothesize that machine learning (ML) models trained exclusively on  
86 local and extensive historical coverage of reanalysis data, and used to predict  $ET_0$  from a  
87 limited set of weather variables (from only a few meteorological instruments) can achieve  $ET_0$   
88 estimation accuracy comparable to established empirical formulas. If confirmed, this approach  
89 could significantly simplify  $ET_0$  estimation by reducing instrumentation requirements and  
90 expanding applicability in data-scarce contexts.

91 Recent ML advancements have shown promise for deriving agroclimatic indicators,  
92 including  $ET_0$ , from limited meteorological inputs. Such estimations have been performed  
93 using methods like support vector machines (SVM), random forests (RF), Gaussian process  
94 regression (GPR), and artificial neural networks (ANN) (Chia et al., 2020; Hebbalaguppa  
95 Krishnashetty et al., 2021). Yet, most existing ML-based  $ET_0$  models rely on short-term, site-  
96 specific datasets, limiting their generalizability (see **Supplementary Table 1**). To our  
97 knowledge, no previous work has utilized global climate reanalysis products—datasets  
98 generated by assimilating historical weather station data into global circulation models—for  
99 training  $ET_0$  estimation models. Such reanalysis datasets provide near-global coverage of past  
100 weather conditions at various resolutions (often down to 0.5 degrees), sub-daily time steps,  
101 and spanning multiple decades.

102           In this paper, we explicitly test the hypothesis that ML models, trained from reanalysis  
103 data using only a few simple on-site measurements (e.g., daily minimum and maximum  
104 temperatures and rainfall), can achieve  $ET_0$  estimation accuracies on par with widely used  
105 empirical methods. By tapping into the historical depth and spatial consistency of the NASA  
106 Power reanalysis dataset, we aim to develop locally adapted models capable of capturing the  
107 long-term statistical relationships underlying  $ET_0$ -PM. Confirming this hypothesis would greatly  
108 simplify the  $ET_0$  estimation process, providing a practical and precise tool for routine  
109 agricultural management and crop modelling in regions with sparse meteorological  
110 infrastructure.

111           Unlike previous approaches limited by geographically sparse or short-term training  
112 datasets, our method exploits the rich spatiotemporal coverage of reanalysis data to train and  
113 evaluate locally adapted ML models systematically. Furthermore, we examine the sensitivity  
114 of model performance to different subsets of input variables—a dimension rarely explored in  
115 the literature. To evaluate our methodology, we train local extreme gradient boosting (XGB)  
116 models (chosen for their robustness in regression tasks, e.g., as shown in hydrological studies  
117 such as (Niazkar et al., 2024)) using NASA Power data. We then compare the  $ET_0$  estimates  
118 generated by these models—under various combinations of meteorological inputs—with  $ET_0$   
119 values derived from the FAO-56-PM approach using ground measurements from twenty  
120 weather stations spread across West Africa. We assess model performance both statistically  
121 and qualitatively and discuss the implications of our findings for local calibration. Finally, we  
122 present a Python software module called RF- $ET_0$ , allowing researchers and practitioners to  
123 easily train and apply these XGB models anywhere in the world.

## 124 2. Materials and methods

### 125 2.1. General workflow

126 The workflow for training, evaluating, and testing XGB models to estimate reference  
127 evapotranspiration (ET<sub>0</sub>) involved several key steps. First, ground truth ET<sub>0</sub> values (ET<sub>0</sub>-PM-  
128 GT) were computed using the FAO-56 Penman-Monteith equation and daily meteorological  
129 data collected from weather stations. These included air temperature (TMIN, TMAX), wind  
130 speed, rainfall, relative humidity (HR), and solar radiation (IRRAD).

131 To complement these observations, daily time series data for each site were retrieved  
132 from the NASA Power reanalysis dataset, spanning 1979–2020. Using the same FAO-56  
133 Penman-Monteith equation, ET<sub>0</sub> values (ET<sub>0</sub>-PM-RM) were computed from the reanalysis  
134 data to serve as the target for model training. The reanalysis dataset was split into a training  
135 subset (historical data up to one year before station observations) and a validation subset (the  
136 period immediately preceding the observations).

137 XGBoost (XGB) models were trained locally for each site using combinations of  
138 meteorological variables derived from reanalysis data as input features (e.g., TMIN, TMAX,  
139 WIND, IRRAD), forming distinct model classes. Each model class represented the variable  
140 configurations measurable by different types of weather stations. Model training used default  
141 hyperparameters with 1000 boosting rounds for robust regression.

142 Model performance was evaluated on the validation dataset by comparing predicted  
143 ET<sub>0</sub> values against ET<sub>0</sub>-PM-RM, using root mean square error (RMSE) and the coefficient of  
144 determination ( $R^2$ ) as metrics. Finally, the trained models were tested against ground truth  
145 ET<sub>0</sub> values (ET<sub>0</sub>-PM-GT) derived from station observations. Performance was assessed  
146 across agro-ecological zones to analyze variability and identify the most reliable variable  
147 combinations for different climatic conditions.

## 148 2.2. In situ meteorological datasets

149 For this study, we used the data from the archive of the AfricaRice weather station  
150 network (<https://dataverse.harvard.edu/dataverse/Weather>), providing a representative set of  
151 meteorological observations for West Africa, as the data provided from their  
152 agrometeorological stations is openly available, homogeneous, and quality controlled. We  
153 selected sites for which original metadata (coordinates and elevation) were provided, and  
154 containing all the necessary variables to compute  $ET_0$  using the FAO-56 standardized  
155 Penman-Monteith equation: air temperature, rainfall, wind speed, vapour pressure, and global  
156 radiation. A total of twenty sites were retained for the study, (**Figure 1, Table 1**). spanning  
157 from 17°N, 18°W to 4°N, 18°E covering seven distinct agro-ecological zones as classified by  
158 the FAO Global Agro-ecological Zones V4, simplified (FAO and IIASA, 2021). Of these,  
159 seventeen sites fall within the "lowland tropics", further categorised into nine "humid", five  
160 "sub-humid", and three "semi-arid" zones; the remaining sites are distributed among  
161 "desert/arid climate", "land with severe soil/terrain constraints", and "land with ample irrigated  
162 soils" (**Figure 2**), ensuring a representative sample of the agro-ecological variability within  
163 West Africa, albeit with some underrepresentation of arid zones.

164 The duration of these records varies depending on the station, and spans from March  
165 19, 2012, to December 31, 2020. **Table 2** describes dataset durations and periods considered  
166 for the analysis on each station. To establish a ground truth for reference evapotranspiration  
167 ( $ET_0$ -PM-GT), we utilised the FAO-56 standardized Penman-Monteith equation, implemented  
168 through the `reference_ET` function in the Python Crop Simulation Environment (PCSE)  
169 package v.5.4.0 (de Wit, 2023).

## 170 2.3. Climate reanalysis datasets

171 To complement our in-situ data, we sourced daily weather data from the NASA Power  
172 API, which offers a comprehensive depiction of long-term meteorological conditions at a



173 regional scale. The data for our measurement sites was retrieved using the  
174 NASA Power Weather Data Provider function of PCSE v.5.4.0. NASA Power provides a detailed  
175 dataset including solar parameters from the GEWEX/SRB 3.0 and the CERES FlashFlux  
176 versions 2 and 3 models, as well as meteorological parameters from the GMAO MERRA-2  
177 and GEOS 5.12.4 FP-IT models on a granular  $0.5^\circ \times 0.5^\circ$  latitude/longitude grid. such as the  
178 NASA Prediction of Worldwide Energy Resources (NASA Power) (Sparks, 2018; Zhang et al.,  
179 2008).

180 For our study, the dataset spanned from January 1st, 1984, to June 1st, 2022. It  
181 included a range of daily meteorological variables crucial for our analysis: minimal temperature  
182 (TMIN), maximal temperature (TMAX), average temperature (TEMP), rainfall (RAIN), average  
183 wind speed (WIND), average vapour pressure (VAP), and global radiation (IRRAD). We  
184 computed daily reference evapotranspiration values from reanalysis model ( $ET_0$ -PM-RM) from  
185 these reanalysis model variables. This computation was done using the reference\_ET function  
186 from PCSE v.5.4.0, integrating station-specific elevation data from the AfricaRice dataset  
187 metadata, instead of the broader  $0.5^\circ \times 0.5^\circ$  elevation averages typically provided by NASA  
188 Power.

## 189 2.4. XGB model training and validation

190 In this study, XGBoost (XGB) models were trained and validated using daily weather  
191 data from NASA Power for each of the twenty sites. We chose to work only with this family of  
192 models as there are recognized as robust and adapted to a variety of tasks. Furthermore, we  
193 wanted to focus on the comparison in the performance of models of the same nature, but  
194 including an increasing number of predictor variables, rather than comparing different model  
195 formalisms. To evaluate the feasibility and performance of this approach, in this work we will  
196 train twenty local XGB models using data from NASA Power, as its API and companion tools  
197 allow for very efficient requests. For each site, a data partitioning was done to split the data  
198 into training and validation subsets. Specifically, the training dataset included NASA Power

199 records until a year prior to the first on-site observations, while the days leading up to these  
200 first observations formed the validation dataset. The in-situ observations themselves served  
201 as the test set. The periods used for each site are detailed in **Table 2**.

202 Our approach utilised eleven distinct combinations of daily variables from the NASA  
203 Power reanalysis model, forming the basis for 11 XGB model classes (**Table 3**). The target  
204 output for these models was the  $ET_0$ -PM-RM value. Model training was conducted using  
205 XGBoost v1.6.2 (Chen and Guestrin, 2016) within a Python v3.9.6 environment. To optimise  
206 performance, we increased the number of boost rounds to 1000 while keeping other  
207 hyperparameters at their default settings.

208 The performance of each XGB model was assessed through two key metrics: root  
209 mean squared error (RMSE) and coefficient of determination ( $R^2$ ). This evaluation involved  
210 comparing the XGB-estimated  $ET_0$  values with the  $ET_0$ -PM-RM values from the validation set,  
211 allowing for a robust analysis of model precision and accuracy over the 20 locations.

## 212 2.5. Comparison of the locally-trained neural network estimation 213 models with other $ET_0$ estimation methods

214 In this study, we benchmarked the performance of our locally-trained XGB models for  
215 estimating  $ET_0$  against traditional empirical methods described in **Table 4**. This comparison  
216 was based on  $ET_0$  calculations on test sets from in situ datasets (**Table 2**). We employed two  
217 primary performance metrics for this analysis: the correlation coefficient and root mean square  
218 error (RMSE). These metrics were applied to evaluate both our neural network approach and  
219 the selected empirical formulas, with a focus on their accuracy in estimating  $ET_0$ -PM-GT  
220 across the twenty study sites.

221 To facilitate a comprehensive understanding of the results, we visualised the  
222 performance comparisons using Taylor diagrams. This graphical representation allowed us to  
223 simultaneously assess the correlation, standard deviation, and RMSE of the different methods.

224 Additionally, we conducted a qualitative analysis to provide deeper insights into the strengths  
225 and limitations of each approach in various agro-climatic conditions.

## 226 3. Results

### 227 3.1. Training and validation of the XGB models

228 In our study, an average of 10,000 records from NASA Power were utilised for each  
229 site to train the different XGB model classes presented in **Table 3**. Analysing the performance  
230 metrics against the validation sets (left-out training data from NASA Power), as detailed in  
231 **Table 5**, revealed distinct patterns in model efficacy. Notably, model class M11 emerged as  
232 the top performer, achieving a mean  $R^2$  of 0.99 and a mean RMSE of just 0.09 mm. This  
233 superior performance aligns with our expectations, considering that M11's set of training  
234 variables - or features - closely mimics the variables used by the FAO56-PM-  $ET_0$  formula.  
235 This suggests a high capability of XGB models to approximate the FAO56-PM-  $ET_0$  method  
236 accurately.

237 Model classes M9 and M10 also demonstrated strong results, with  $R^2$  values around  
238 0.95, but their RMSEs were higher, approximately 0.15 mm. Model class M7, while showing  
239 promising results with a mean  $R^2$  of 0.87, exhibited variability across different sites. In contrast,  
240 the least effective models (M1-M4) recorded  $R^2$  values around 0.62 and an RMSE of 0.68 mm,  
241 indicating a lower level of performance.

242 These findings imply that the performance of  $ET_0$ -PM-RM estimation might vary with  
243 agro-ecological zones (AEZs) as the main driving factors of evapotranspiration are not  
244 identical in all environments. A notable observation was the underperformance of model class  
245 M7 in north-western sites, suggesting that adding more features could potentially enhance  
246 accuracy. Specifically, including the irradiance (IRRAD) feature in the models significantly  
247 improved  $ET_0$  estimations by incorporating the radiative component, as evidenced by the  
248 improved validation metrics.

## 249 3.2. Comparison of the trained XGB performance against 250 empirical $ET_0$ estimation formulas from ground data

251 Then, from the in-situ values of the test set, we calculated  $ET_0$  using established  
252 empirical formulas (presented in **Table 4**), estimated  $ET_0$  using our different XGB model  
253 classes, and compared these calculations with the  $ET_0$  value computed using in situ weather  
254 measurements ( $ET_0$ -PM-GT) across the twenty sites (**Table 6**). The performance of each  
255 method was quantified using the root mean square error (RMSE) and correlation coefficient,  
256 with the results further illustrated in Taylor diagrams (representing correlation coefficient,  
257 RMSE and reference value in one plot) for a visual comparison (**Figure S1**).

258 A qualitative analysis was performed from these Taylor diagrams, in which we consider  
259 as a suitable model a model in which the RMSE was below 1 mm, and that is inside the highest  
260 available and the least high available correlation coefficient quadrant. The models were then  
261 scored and ranked based on their adherence to these criteria at each site (**Table 7**). According  
262 to the data, XGB model class 11 was distinguished as the most accurate, followed by the  
263 Irmak-a model. Subsequent ranks were shared by Hargreaves, Irmak-c, Valiantzas-a, and  
264 XGB-M8 models, which displayed comparable performance.

265 Notably, the Irmak-a model showed significant variability in RMSE across the sites,  
266 with a mean variability of 0.75, as opposed to 0.71 for M8 and 0.48 for M9. This variability is  
267 crucial as it affects the reliability and consistency of the model's performance in diverse  
268 climatic contexts, as represented by the various sites in West Africa.

269

## 270 4. Discussion

271 Our study contributes to the active research on machine learning (ML) methodologies  
272 for estimating  $ET_0$  from a limited dataset of meteorological variables. In a field inclined towards  
273 the application of various ML approaches, we've selected XGB, and specifically extreme

274 gradient boosting (XGBoost) due to its proven robustness and established track record  
275 (Nielsen, 2016). Central to our discourse is the challenge of training data acquisition—a facet  
276 that, in our view, has not received due attention. The prevalent practice in the literature utilises  
277 weather data from a small number of locations over short time series to train models (see  
278 **Supplementary Table 1** for a curated selection of studies on the topic). These models are  
279 then expected to extrapolate  $ET_0$  estimations to larger, often regional scales, which risks  
280 significant extrapolation errors during application.

281         We propose a methodological pivot, employing a transversal ML approach that utilises  
282 local meteorological data from assimilation databases. This strategy has the distinct  
283 advantage of being deployable in areas devoid of ground weather data, offering uniform  
284 datasets that boast extensive historical records and spatial representativity—crucial for  
285 models that need to operate across varied geographic expanses. We harnessed the breadth  
286 of NASA Power's data, spanning a grid resolution of  $0.5^\circ \times 0.5^\circ$ , effectively training a model  
287 for every  $\sim 3000\text{km}^2$  under West Africa's latitudes. However, we acknowledge that such a  
288 resolution may not fully encapsulate the microclimatic diversity within these regions, especially  
289 in areas with complex terrain. One solution to this limitation could involve debiasing reanalysis  
290 data using historical ground measurements, potentially through methods such as linear  
291 regression, prior to training the ML models (Chandler, William S., 2011).

292         Moreover, our approach stands to be further refined by incorporating data from sources  
293 with finer resolutions, such as the fifth generation of European ReAnalysis climate products  
294 (Copernicus ERA5) (Hersbach et al., 2020; Muñoz-Sabater et al., 2021). This would offer a  
295 more detailed  $0.1^\circ \times 0.1^\circ$  resolution and an extensive historical reach back to 1950—allowing  
296 the inclusion of rarer climatic events in our training set, however at a price of lengthier API  
297 calls thus overall computation time. Such granularity is anticipated to enhance the  
298 representativeness of training data and, consequently, the performance of ML models,  
299 although this hypothesis warrants additional research to quantify the influence of training data  
300 volume on model performance.

301 For example, the Hargreaves-Samani model (George H. Hargreaves and Zohrab A.  
302 Samani, 1985), the Schendel model (Schendel, U., 1967), and the Priestley-Taylor model  
303 (Priestley and Taylor, 1972) are based only on solar radiation and air temperature values ; the  
304 Turc radiation-based model (Turc, L, 1961) uses temperature and relative humidity ; the  
305 Makkink-based models (Allen, Richard, 2003; Hansen, 1984; Makkink, G.F., 1957) use  
306 radiation-related parameters. A thorough review of these methods can be found in (McMahon  
307 et al., 2013).

308 Aiming to serve practitioners with scant meteorological infrastructure, our study sought  
309 to validate the accuracy and precision of XGB-based  $ET_0$  estimation models against the FAO-  
310 56 standardized PM equation. Our experiments scrutinised the impact of a limited set of  
311 input variables on  $ET_0$  output quality. We also compared the ML-derived  $ET_0$  estimates with  
312 those calculated from ground truth data across diverse West African climates, ensuring our  
313 models' efficacy relative to conventional methods.

314 The practicality of our approach is underscored by its minimal data requirements: a  
315 farm manager equipped with only a min/max thermometer and a rain funnel shall be able to  
316 use a locally trained model to estimate  $ET_0$  with a level of precision and accuracy usually  
317 attainable by also having anemometer or hygrometer readings. Existing literature suggests  
318 different combinations of variables for efficient  $ET_0$  estimation by models : (Granata, 2019)  
319 suggests mean temperature, mean relative humidity, and net solar radiation; wind and sunny  
320 hours as the strict minimum information for most locations according to (Shabani et al., 2020),  
321 and most empirical formulas need at least information about wind speed plus deficit pressure  
322 or solar radiation with air temperature. Meanwhile, our approach showed that it can yield  
323 models that perform relatively well, starting with only information about minimum, maximum,  
324 and average daily air temperature. With a dependency on a high number of devices to  
325 measure each variable needed, there is a greater risk in data gaps leading to impossibility to  
326 calculate  $ET_0$ ; this issue of missing or incomplete meteorological data has been discussed in  
327 (Koudahe et al., 2018). Our finding suggests that practitioners may only need a quality

328 thermometer for reliable  $ET_0$  estimation, simplifying an on-farm weather data acquisition  
329 process that is already challenging in the region (Dinku, 2019). Stacking up variables will make  
330 the estimations more reliable. Compared to other models that overestimate  $ET_0$ , our models  
331 demonstrated better stability and less bias across various climates, although this observation  
332 necessitates further validation with a broader dataset. The use of twenty weather stations  
333 provides a solid foundation for our performance analysis, but a more expansive dataset would  
334 enable a more comprehensive understanding of the models' applicability and robustness.

335 Most of these empiric models have been initially defined in a given climatic  
336 environment and are less reliable or at least need recalibration when used in different climates.  
337 In the context of climate change, with anticipated shifts in climatic zones (Cui et al., 2021), our  
338 universally applicable method, supported by global assimilation data, could significantly aid in  
339  $ET_0$  estimation efforts. Our XGB-  $ET_0$  models, though in need of further validation, show  
340 potential for at least equivalent performance compared to traditional empirical models, with  
341 the added benefit of being more robust across different AEZs. This robustness is particularly  
342 important in reducing bias risks, a common issue with empirical formulas lacking quality  
343 assessment avenues.

344 To enable wider application and continued research, we have made our XGB-  $ET_0$   
345 code publicly available under the GPLv3 licence (doi: [10.5281/zenodo.10463369](https://doi.org/10.5281/zenodo.10463369), repository  
346 URL: <https://github.com/SARRA-cropmodels/RF-ET0>). This initiative allows researchers and  
347 practitioners to conduct their own localised model training and  $ET_0$  estimations, utilising the  
348 model classes we've explored in this work.

## 349 Conclusion

350 This research represents a significant stride in the application of machine learning for  
351 agricultural water management, particularly within the challenging context of West Africa. The  
352 study has successfully demonstrated the efficacy of XGB models in estimating  
353 evapotranspiration (ET) using limited meteorological data. Our approach, which utilizes

354 extensive historical and spatial data from NASA Power's climate reanalysis datasets, marks a  
355 departure from traditional methods reliant on short-term, location-specific weather data. The  
356 ability of these models to accurately approximate ET estimations, as validated against ground  
357 measurements from diverse West African agro-ecological zones, underscores their potential  
358 in areas with limited meteorological infrastructure. This methodology not only simplifies the  
359 data requirements for ET estimation but also proves to be robust across various climatic  
360 conditions, enhancing its applicability for sustainable water management practices. The  
361 outcomes of this research contribute to the broader body of knowledge by showcasing the  
362 practicality of machine learning in agricultural contexts and underscore the importance of  
363 adopting innovative technologies in the face of climate change and resource constraints. The  
364 findings of this study hold significant implications for precision irrigation and water resource  
365 management, providing a sustainable and accessible solution to a region facing acute water-  
366 related challenges.

## 367 Data Statement

368 All data, models, or code generated or used during the study are available in a  
369 repository or online in accordance with funder data retention policies. Data used for validation  
370 is available at the DOIs indicated in **Table 2**. Additionally, the climate reanalysis datasets  
371 utilized in this study, provided by NASA Power, are publicly accessible and can be retrieved  
372 for extended time series and various spatial resolutions through their API and data services.  
373 The XGB-  $ET_0$  code developed for this study is publicly available under the GPLv3 licence,  
374 ensuring open access and facilitating further research and practical applications. The code  
375 can be accessed at the following repository URL: [https://github.com/SARRA-cropmodels/RF-](https://github.com/SARRA-cropmodels/RF-ET0)  
376 [ET0](https://github.com/SARRA-cropmodels/RF-ET0) (doi:10.5281/zenodo.10463369).



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- 385           • Jérémy LAVARENNE: Conceptualization, Methodology, Software, Formal Analysis,  
386           Investigation, Writing - Original draft, Visualization, Supervision
- 387           • Audrey BROUILLET: Conceptualization, Methodology, Writing - Review and editing

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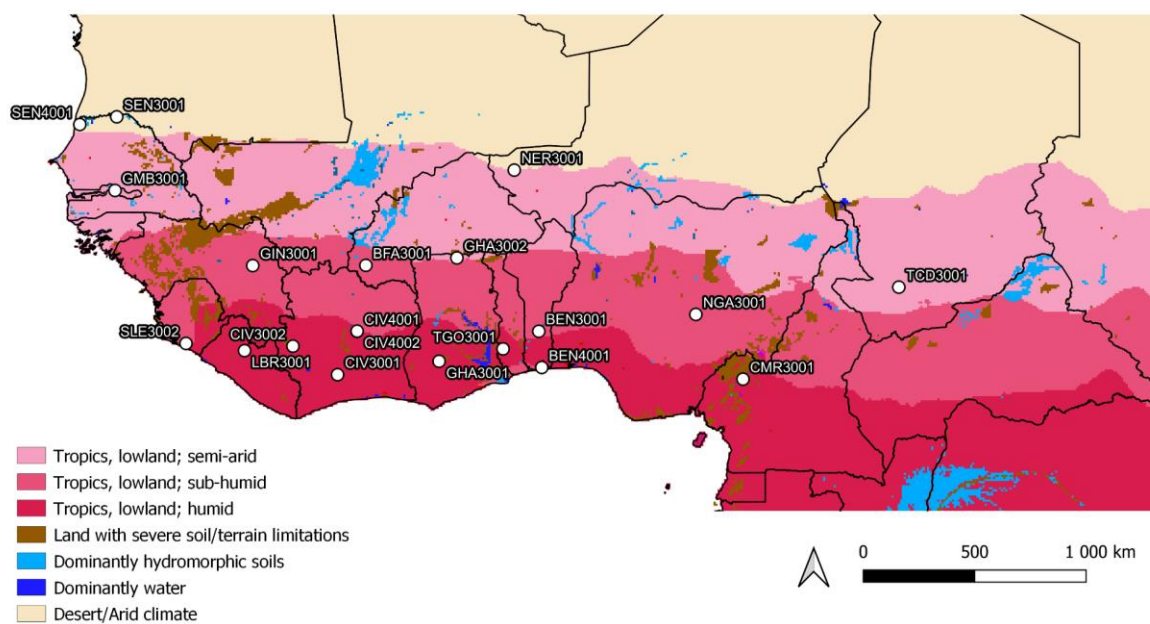
554

555 **Figures**

556 **Figure 1**

557 **Map of Weather Measurement Sites and Agro-Ecological Zones**

558 This color map of West Africa displays the geographic locations of the twenty weather  
559 measurement sites used in the study, overlaid on a representation of global agro-ecological  
560 zones as defined by the FAO and IIASA. Each site is indicated by a dot symbol. The color  
561 legend for the agro-ecological zones (AEZ) is provided to differentiate between the zones.



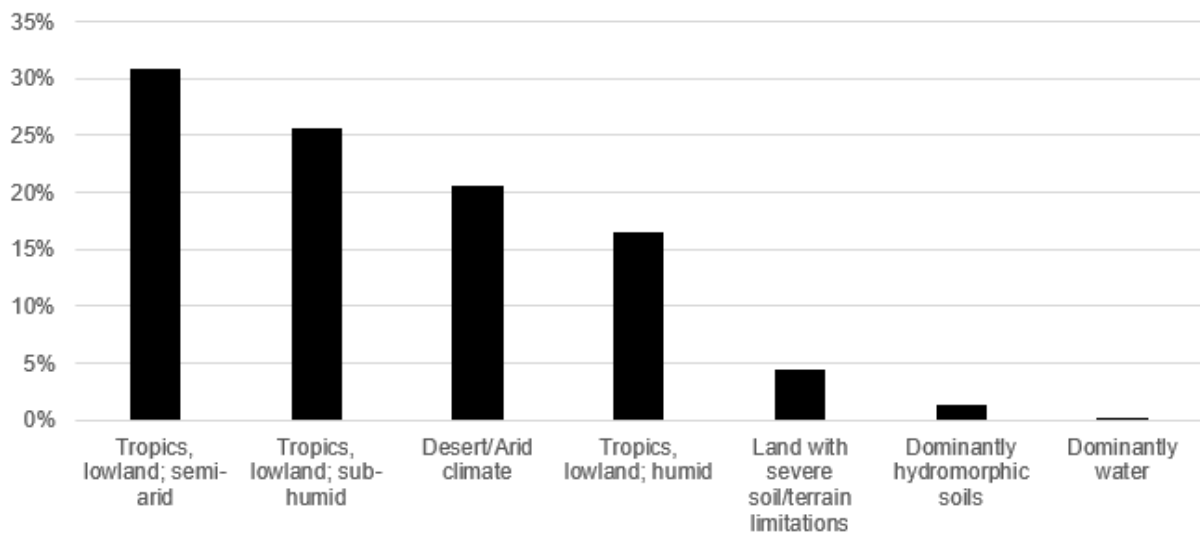
562

563 **Figure 2**

564 **Surface Occupancy Histogram of AEZ**

565 This histogram illustrates the distribution of surface occupancy percentages for the various  
566 Agro-Ecological Zones (AEZ) within the study's geographic extent. Each bar represents the  
567 proportion of an AEZ relative to the total area considered.





568

569

570 **Tables**

571 **Table 1**

572 **In Situ Daily Data Collection Sites**

573 This table enumerates the weather stations selected for the collection of daily in situ  
 574 meteorological data. Detailed within are the geographical coordinates (latitude and longitude),  
 575 altitude, country of location, data collection period, DOI for dataset reference, and the agro-  
 576 ecological zones represented by each station.

Station name	Station code	Latitude	Longitude	Altitude (m)	Country	Measurement start	Measurement end	Dataset DOI	Agro-ecological zone
Glazoue	BEN3001	7.8800	2.2100	199	Benin	2012-07-05	2016-09-22	doi:10.7910/DVN/GVQYG	Tropics, lowland; sub-humid
AfricaRice_Calavi	BEN4001	6.4200	2.3300	41	Benin	2012-03-19	2018-01-08	doi:10.7910/DVN/WTIMYZ	Tropics, lowland; humid
Farakoba	BFA3001	10.5400	-4.7600	319	Burkina Faso	2013-06-21	2018-03-08	doi:10.7910/DVN/TQX3L9	Tropics, lowland; sub-humid
Gagnoa_HUB	CIV3001	6.1300	-5.9000	248	Ivory Coast	2015-07-25	2018-06-13	doi:10.7910/DVN/A6SGP3	Tropics, lowland; humid
Man	CIV3002	7.2800	-7.7000	376	Ivory Coast	2012-07-21	2019-05-31	doi:10.7910/DVN/DQXCJC	Tropics, lowland; humid
Mbe_Low	CIV4001	7.8744	-5.1138	273	Ivory	2014-10-	2020-12-	doi:10.79	Tropics,

land					Coast	03	31	10/DVN/ HVKBPL	lowland; humid
Mbe_Upl and	CIV4002	7.8856	-5.1123	303	Ivory Coast	2014-10- 10	2018-08- 05	doi:10.79 10/DVN/2 C7XKY	Tropics, lowland; humid
Ndop	CMR300 1	5.9500	10.4400	1185	Cameroo n	2012-08- 03	2015-06- 24	doi:10.79 10/DVN/Z SP7NQ	Land with severe soil/terrai n limitations
Gimbi	COD300 1	-5.2200	13.2700	364	Democrat ic Rep. Congo	2013-08- 03	2014-02- 01	doi:10.79 10/DVN/I XPDM3	Tropics, highland; sub- humid
Kumasi	GHA3001	6.6700	-1.8100	236	Ghana	2012-06- 23	2016-04- 20	doi:10.79 10/DVN/ EUPQPZ	Tropics, lowland; humid
Navrongo	GHA3002	10.8400	-1.0900	173	Ghana	2013-05- 14	2014-05- 17	doi:10.79 10/DVN/ H5MX3J	Tropics, lowland; semi-arid
Kankan	GIN3001	10.5300	-9.3200	432	Guinea	2013-11- 09	2016-03- 16	doi:10.79 10/DVN/6 PHUPC	Tropics, lowland; sub- humid
Central_ River	GMB300 1	13.5600	-14.8700	41	Gambia	2013-03- 07	2020-03- 02	10.7910/ DVN/ZLH PFP	Tropics, lowland; semi-arid
Cari	LBR3001	7.1000	-9.6500	273	Liberia	2015-10- 08	2018-02- 21	doi:10.79 10/DVN/ MJQVZO	Tropics, lowland; humid

Tillabery_ HUB	NER3001	14.3800	1.2200	234	Niger	2012-07- 29	2015-07- 03	doi:10.79 10/DVN/ Q6ZDBV	Tropics, lowland; semi-arid
Lafia	NGA3001	8.5600	8.5400	150	Nigeria	2012-08- 09	2018-02- 25	doi:10.79 10/DVN/8 ERRWL	Tropics, lowland; sub- humid
Fanaye	SEN3001	16.5300	-14.8000	11	Senegal	2013-02- 06	2020-12- 31	doi:10.79 10/DVN/9 HVVWP	Desert/Ar id climate
Ndiaye	SEN4001	16.2200	-16.2900	13	Senegal	2013-02- 01	2019-12- 14	doi:10.79 10/DVN/ HTLG35	Land with ample irrigated soils
Tormabu m	SLE3002	7.3900	-12.0100	9	Sierra Leone	2014-05- 28	2016-05- 01	doi:10.79 10/DVN/ VWGJHE	Tropics, lowland; sub- humid
Tandjile_ East	TCD3001	9.6600	16.7200	345	Chad	2013-06- 07	2013-07- 03	doi:10.79 10/DVN/ WKP6HR	Tropics, lowland; humid
Adeta	TGO3001	7.1700	0.7700	295	Togo	2013-04- 04	2015-12- 05	doi:10.79 10/DVN/0 SQFGQ	Tropics, lowland; humid

578 Table 2

579 **Dataset Durations for Model Development**

580 Summary of dataset lengths for ground weather observations, detailing the periods utilised for  
 581 the training, validation, and testing of XGB models.

station code	on-site observations interval		subset duration (days)		
	start	end	train set	validation set	test set
BEN3001	2012-07-05	2016-09-22	10048	365	1540
BEN4001	2012-03-19	2018-01-08	9940	365	2121
BFA3001	2013-06-21	2018-03-08	10399	365	1721
CIV3001	2015-06-25	2018-06-13	11133	365	1084
CIV3002	2012-07-21	2019-05-31	10064	365	2505
CIV4001	2014-10-03	2020-12-31	10868	365	2281
CIV4002	2014-10-10	2018-10-05	10875	365	1456
CMR3001	2012-08-03	2015-06-24	10077	365	1055
COD3001	2013-08-03	2014-02-01	10442	365	182
GHA3001	2012-06-23	2016-04-20	10036	365	1397
GHA3002	2013-05-14	2014-05-17	10361	365	368
GIN3001	2013-11-09	2016-03-16	10540	365	858
GMB3001	2013-03-07	2020-03-07	10293	365	2557
LBR3001	2015-10-08	2018-02-21	11238	365	867
NER3001	2012-01-01	2015-07-03	9862	365	1279
NGA3001	2012-01-01	2018-02-25	9862	365	2247
SEN3001	2013-01-01	2020-12-31	10228	365	2921
SEN4001	2013-02-01	2019-12-31	10259	365	2524
SLE3002	2014-01-01	2016-06-22	10593	365	903
TCD3001	2013-01-12	2013-07-03	10239	365	172
TGO 3001	2013-01-01	2015-12-05	10228	365	1068
		average	10361		1481

582

583 **Table 3**

584 **Training Features for XGB Model Classes**

585 This table specifies the combinations of meteorological variables used to train the eleven  
 586 distinct XGB model classes, with each variable corresponding to data typically collected by  
 587 standard meteorological instruments. X indicates the correspondence between the set of  
 588 variables and the minimum set of sensors needed to measure these variables on site

model alias	Features used for training	maximum– minimum thermometer	temperature data logger	rain gauge	anemome ter	hygrom eter	pyrano meter
M1	TMIN, TMAX	x					
M2	TEMP, TMIN, TMAX		x				
M3	TMIN, TMAX, RAIN	x		x			
M4	TEMP, TMIN, TMAX, RAIN		x	x			
M5	TEMP, TMIN, TMAX, RAIN, WIND		x	x	x		
M6	TEMP, TMIN, TMAX, RAIN, VAP		x	x		x	
M7	TEMP, TMIN, TMAX, RAIN, IRRAD		x	x			x
M8	TEMP, TMIN, TMAX, RAIN, WIND, VAP		x	x	x	x	
M9	TEMP, TMIN, TMAX, RAIN, WIND, IRRAD		x	x	x		x
M10	TEMP, TMIN, TMAX, RAIN, VAP, IRRAD		x	x		x	x
M11	TEMP, TMIN, TMAX, RAIN, WIND, VAP, IRRAD		x	x	x	x	x

589

590 **Table 4**

591 **Empirical methods used for  $ET_0$  estimations.**

592  $ET_0$  reference evapotranspiration (mm);  $u_2$  represents the wind speed measured at 2 m from  
 593 the ground ( $ms^{-1}$ );  $(es - ea)$  vapour saturation deficit (kPa);  $T_a$  is the average air temperature  
 594 ( $^{\circ}C$ );  $T_{max}$ —maximum temperature ( $^{\circ}C$ );  $T_{min}$ —minimum temperature ( $^{\circ}C$ );  $R_s$  is the short  
 595 wavelength solar radiation ( $MJ\ m^{-2}d^{-1}$ );  $\phi$  represents the latitude of the station in radian  
 596 degree, and  $\lambda$  is the latent heat of vaporisation ( $MJ.m^{-2}d^{-1}$ ).

Category	Reference	Formula
Aerodynamic	Dalton (Dalton, J, 1802)	$ET_0 = (0.3648 + 0.07223u_2)(es - ea)$
	Trabert (Trabert, W, 1896)	$ET_0 = 0.3075\sqrt{u_2}(es - ea)$
	Penman (Penman, H. L., 1963)	$ET_0 = 0.35(1 + 0.24u_2)(es - ea)$
	Rohwer (Rohwer, Carl, 1931)	$ET_0 = 0.44(1 + 0.27u_2)(es - ea)$
	Mahringer (Mahringer, 1970)	$ET_0 = 0.15072\sqrt{3.6}u_2(es - ea)$
Temperature	Hargreaves (George H. Hargreaves and Zohrab A. Samani, 1985)	$ET_0 = 0.0135 \times 0.408R_s(T_a + 17.8)$
Radiation	Jensen & Haise (Jensen and Haise, 1963)	$ET_0 = 0.025(T_a - 3)R_s$
	Abtew (Abtew, 1996)	$ET_0 = 0.53 \frac{R_s}{\lambda}$
	Oudin (Oudin, Ludovic, 2004)	$ET_0 = R_s \times \frac{T_a + 5}{100}$
	Irmak (a) (Irmak et al., 2003a, 2003b)	$ET_0 = 0.149R_s + 0.079T_a - 0.611$
	Irmak (b) (Irmak et al., 2003a, 2003b)	$ET_0 = 0.174R_s + 0.353T_a - 0.642$
	Irmak (c) (Irmak et al., 2003a, 2003b)	$ET_0 = 0.156R_s - 0.0112T_{max} + 0.0733T_{min} - 0.478$

Combinatory	Valiantzas (a) (Valiantzas, 2013)	$ET_0 = (0.0393Rs \times \sqrt{Ta + 9.5}) - 0.19Rs^{0.6}\varphi^{0.15} + 0.048(Ta + 20)\left(1 - \frac{Hr}{100}\right)u^{2.07}$
	Valiantzas (b) (Valiantzas, 2013)	$ET_0 = (0.0393Rs \times \sqrt{Ta + 9.5}) - 0.19Rs^{0.6}\varphi^{0.15} + 0.078(Ta + 20)\left(1 - \frac{Hr}{100}\right)$
	Valiantzas (c) (Valiantzas, 2013)	$ET_0 = (0.0393Rs \times \sqrt{Ta + 9.5}) - 0.19Rs^{0.6}\varphi^{0.15} + 0.0061(Ta + 20)(1.12Ta - T_{min} - 2)^{0.7}$

597

598



599 **Table 5**

600 **Validation Metrics for XGB Models**

601 This table provides the root mean square error (RMSE) and coefficient of determination (R<sup>2</sup>)  
602 obtained from the validation sets used in the training of the eleven XGB model classes.

	station	model class											
		M1	M2	M3	M4	M5	M6	M7	M8	M9	M10	M11	
RMSE (mm)	BEN3001	0,43	0,43	0,42	0,42	0,42	0,41	0,20	0,41	0,15	0,16	0,06	
	BEN4001	0,50	0,48	0,47	0,47	0,47	0,46	0,14	0,47	0,12	0,12	0,06	
	BFA3001	0,77	0,78	0,74	0,76	0,49	0,68	0,61	0,48	0,29	0,51	0,10	
	CV3001	0,41	0,41	0,40	0,41	0,40	0,42	0,13	0,41	0,10	0,10	0,06	
	CV3002	0,38	0,38	0,38	0,38	0,37	0,38	0,15	0,37	0,12	0,12	0,07	
	CV4001	0,55	0,56	0,55	0,56	0,50	0,53	0,35	0,48	0,25	0,32	0,09	
	CV4002	0,55	0,56	0,55	0,56	0,50	0,53	0,35	0,48	0,25	0,32	0,09	
	CMR3001	0,41	0,42	0,39	0,39	0,40	0,37	0,12	0,38	0,12	0,09	0,05	
	COD3001	0,69	0,67	0,70	0,68	0,67	0,66	0,06	0,66	0,06	0,06	0,04	
	GHA3001	0,43	0,44	0,43	0,44	0,43	0,42	0,24	0,42	0,19	0,20	0,07	
	GHA3002	0,81	0,82	0,81	0,83	0,53	0,72	0,67	0,51	0,34	0,54	0,12	
	GIN3001	0,71	0,72	0,71	0,73	0,50	0,69	0,58	0,50	0,33	0,53	0,11	
	GMB3001	0,96	0,94	0,95	0,96	0,57	0,81	0,90	0,50	0,46	0,74	0,13	
	LBR3001	0,40	0,40	0,39	0,40	0,38	0,39	0,11	0,38	0,10	0,10	0,06	
	NER3001	1,14	1,14	1,13	1,14	0,57	0,98	1,09	0,50	0,48	0,95	0,16	
	NGA3001	0,75	0,75	0,76	0,76	0,53	0,63	0,66	0,50	0,34	0,43	0,12	
	SEN3001	1,44	1,45	1,43	1,45	0,61	1,22	1,33	0,48	0,53	1,09	0,19	
	SEN4001	1,09	1,08	1,07	1,06	0,63	0,90	1,03	0,45	0,62	0,75	0,19	
	SLE3002	0,55	0,55	0,54	0,55	0,53	0,55	0,10	0,54	0,09	0,09	0,06	
	TCD3001	0,97	0,95	0,98	0,97	0,53	0,79	0,83	0,53	0,31	0,63	0,11	
	TGO3001	0,43	0,43	0,43	0,43	0,43	0,42	0,18	0,42	0,14	0,17	0,07	
		<i>mean</i>	0,68	0,68	0,68	0,68	0,50	0,62	0,47	0,47	0,26	0,38	0,09
		<i>median</i>	0,55	0,56	0,55	0,56	0,50	0,55	0,35	0,48	0,25	0,32	0,09
R2	BEN3001	0,58	0,57	0,60	0,59	0,60	0,62	0,91	0,62	0,95	0,94	0,99	
	BEN4001	0,43	0,46	0,49	0,49	0,50	0,51	0,96	0,50	0,97	0,97	0,99	
	BFA3001	0,76	0,76	0,78	0,77	0,90	0,82	0,85	0,91	0,97	0,90	1,00	
	CV3001	0,33	0,32	0,34	0,32	0,35	0,28	0,93	0,31	0,96	0,96	0,99	
	CV3002	0,61	0,62	0,61	0,62	0,62	0,62	0,94	0,62	0,96	0,96	0,99	
	CV4001	0,75	0,74	0,75	0,74	0,80	0,77	0,90	0,81	0,95	0,91	0,99	
	CV4002	0,75	0,74	0,74	0,73	0,79	0,77	0,90	0,81	0,95	0,91	0,99	
	CMR3001	0,68	0,67	0,72	0,71	0,70	0,74	0,97	0,73	0,97	0,98	1,00	
	COD3001	0,11	0,15	0,09	0,14	0,16	0,18	0,99	0,19	0,99	0,99	1,00	
	GHA3001	0,72	0,71	0,72	0,71	0,72	0,73	0,91	0,73	0,95	0,94	0,99	
	GHA3002	0,78	0,77	0,78	0,77	0,91	0,83	0,85	0,91	0,96	0,90	1,00	
	GIN3001	0,82	0,82	0,82	0,81	0,91	0,83	0,88	0,91	0,96	0,90	1,00	
	GMB3001	0,81	0,82	0,81	0,81	0,93	0,86	0,83	0,95	0,96	0,89	1,00	
	LBR3001	0,42	0,42	0,44	0,42	0,47	0,44	0,96	0,48	0,97	0,96	0,99	
	NER3001	0,57	0,57	0,58	0,57	0,89	0,68	0,61	0,92	0,92	0,70	0,99	
	NGA3001	0,82	0,82	0,82	0,82	0,91	0,88	0,87	0,92	0,96	0,94	1,00	
	SEN3001	0,55	0,54	0,55	0,54	0,92	0,67	0,61	0,95	0,94	0,74	0,99	
	SEN4001	0,59	0,59	0,60	0,61	0,86	0,72	0,64	0,93	0,86	0,80	0,99	
	SLE3002	0,54	0,54	0,55	0,54	0,57	0,54	0,99	0,56	0,99	0,99	0,99	
	TCD3001	0,74	0,75	0,73	0,74	0,92	0,83	0,81	0,92	0,97	0,89	1,00	
	TGO3001	0,65	0,65	0,65	0,64	0,65	0,66	0,94	0,66	0,96	0,95	0,99	
		<i>mean</i>	0,62	0,62	0,63	0,62	0,72	0,66	0,87	0,73	0,96	0,91	0,99
		<i>median</i>	0,65	0,65	0,65	0,64	0,79	0,72	0,90	0,81	0,96	0,94	0,99

603

604

605 **Table 6**

606 **Performance Comparison of XGB and Empirical ET0 Formulas**

607 This table presents the average metrics from the comparison between XGB model estimations  
608 and those calculated using empirical ET<sub>0</sub> formulas against the standard ET<sub>0</sub>-PM-GT values.  
609 Metrics include RMSE and R<sup>2</sup> values for each model and formula.



611 **Table 7**

612 **Taylor Diagram Analysis**

613 Summary of the qualitative analysis based on Taylor diagrams for model performance across  
614 the 20 study sites.

615

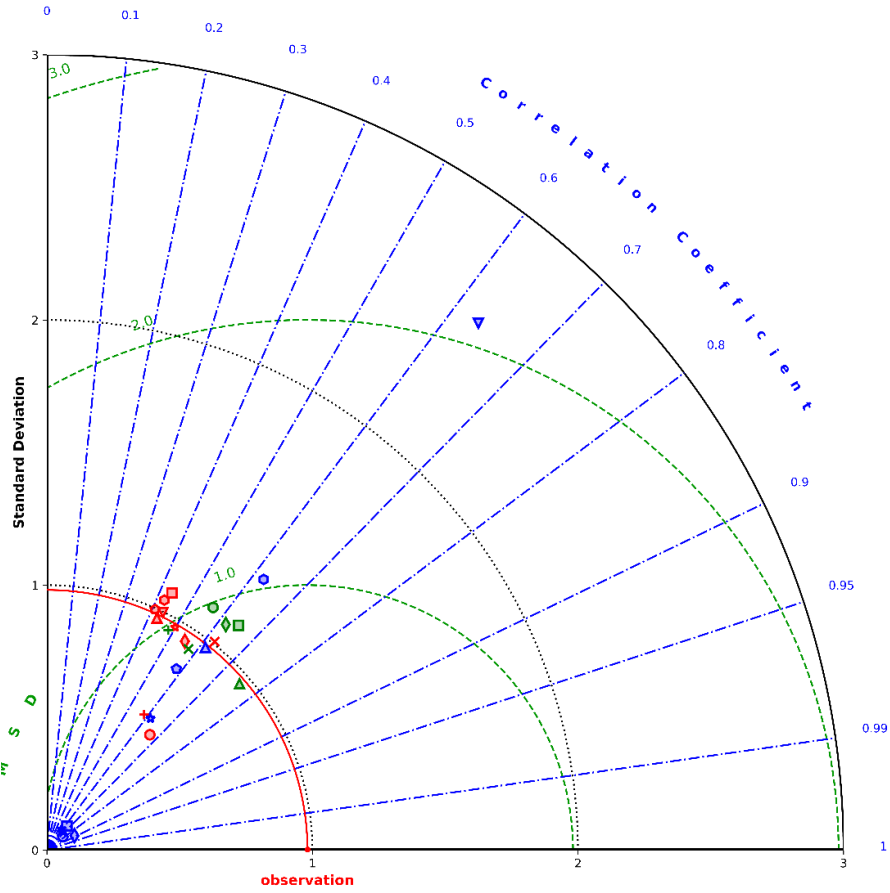
	Tropics, lowland, sub-humid		Tropics, lowland, humid			Tropics, lowland, semi-arid		Desert/Arid climate	Land with severe soil/terrain limitations		Land with ample irrigated soils		SUM	RANK
	BEN3001	BFA3001	BEN4001	CV3001	CV3002	CV4001	CV4002	GHA3001	GHA3002	GMB3001	NER3001	SEN3001		
Dakon														
Tesart														
Pompan														
Rohrufer														
Haininger														
Heigraves	X	X								X				
Jansensraas														
Abbew														
Oudin														
Imak-a	X	X												
Imak-b														
Imak-c	X	X												
Valantzas-a	X	X												
Valantzas-b	X	X												
Valantzas-c														
M1														
M2														
M3														
M4														
M5														
M6														
M7														
M8														
M9	X	X												
M10	X	X												
M11	X	X	X	X	X	X	X	X	X	X	X	X	X	X

# Supplementary figures

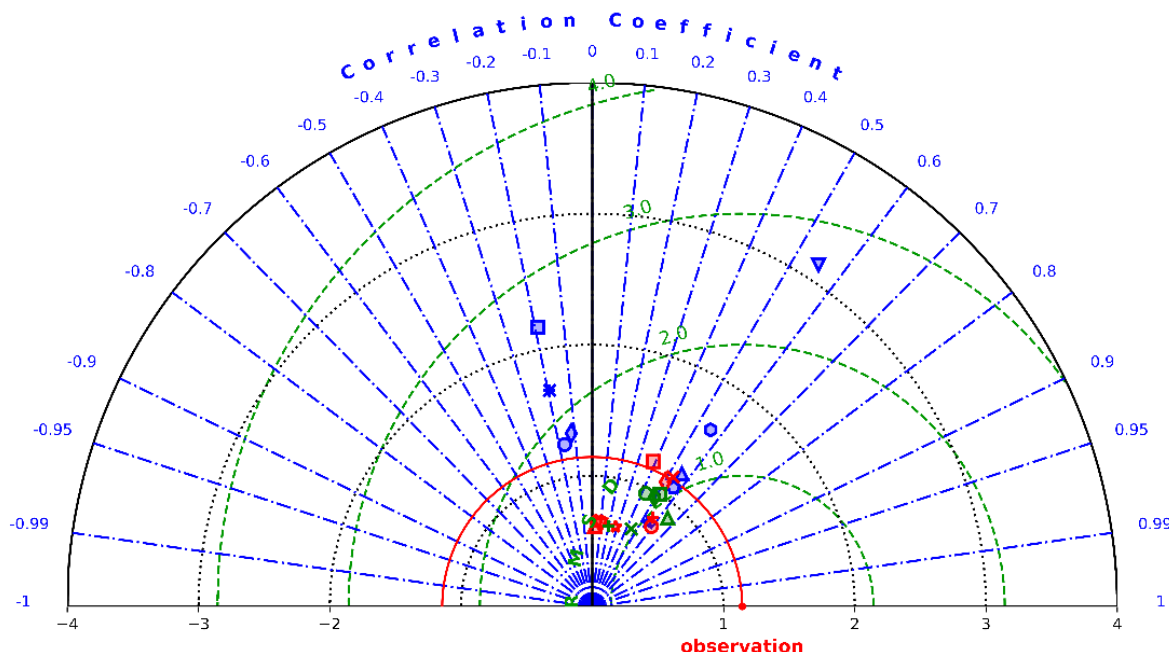
## Figure S1

Taylor diagrams for comparison of our approach versus empirical methods

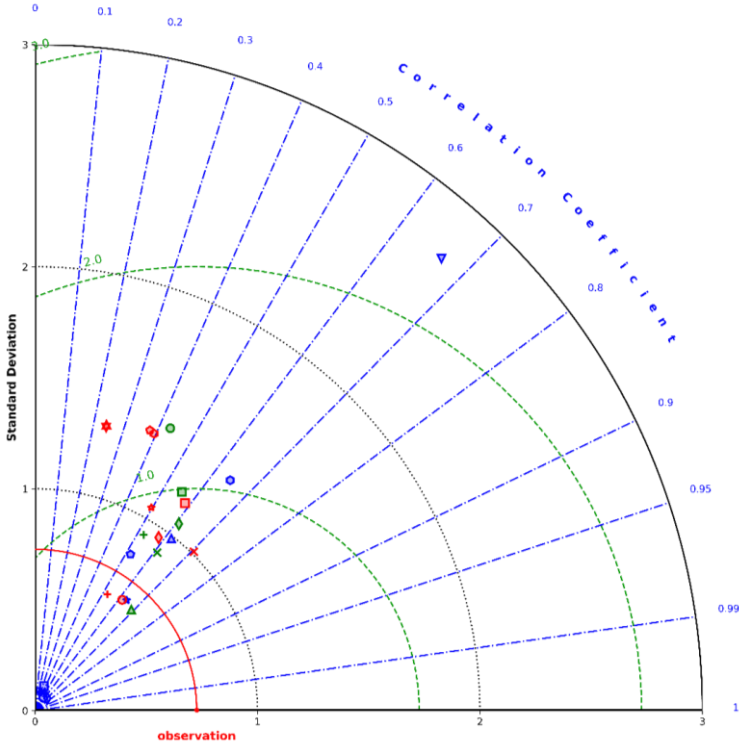
S1(a) Taylor diagram for station BEN3001



S1(b) Taylor diagram for station BEN4001

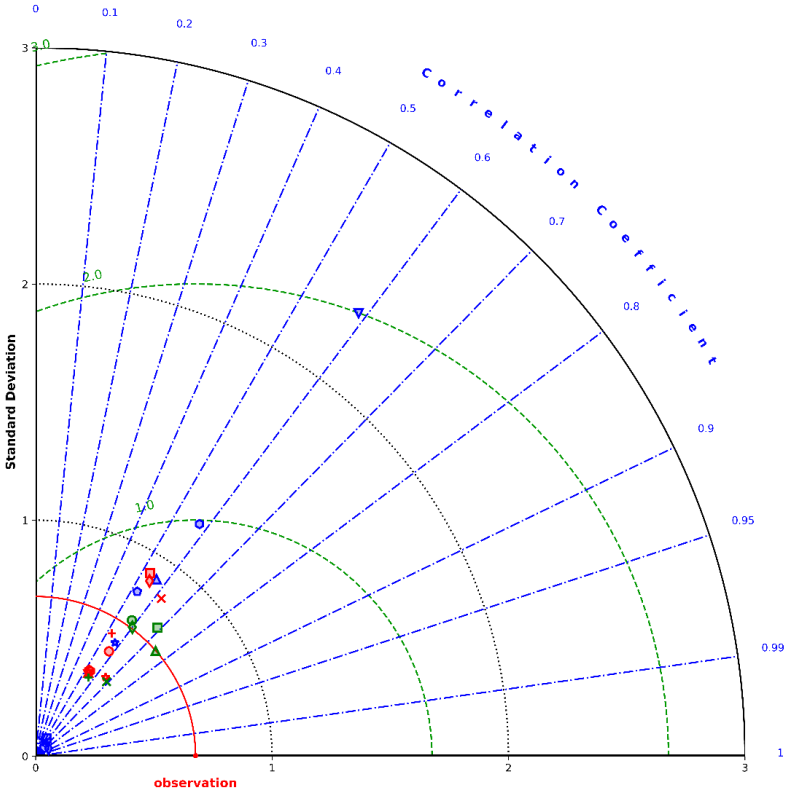


S1(c) Taylor diagram for station BFA3001

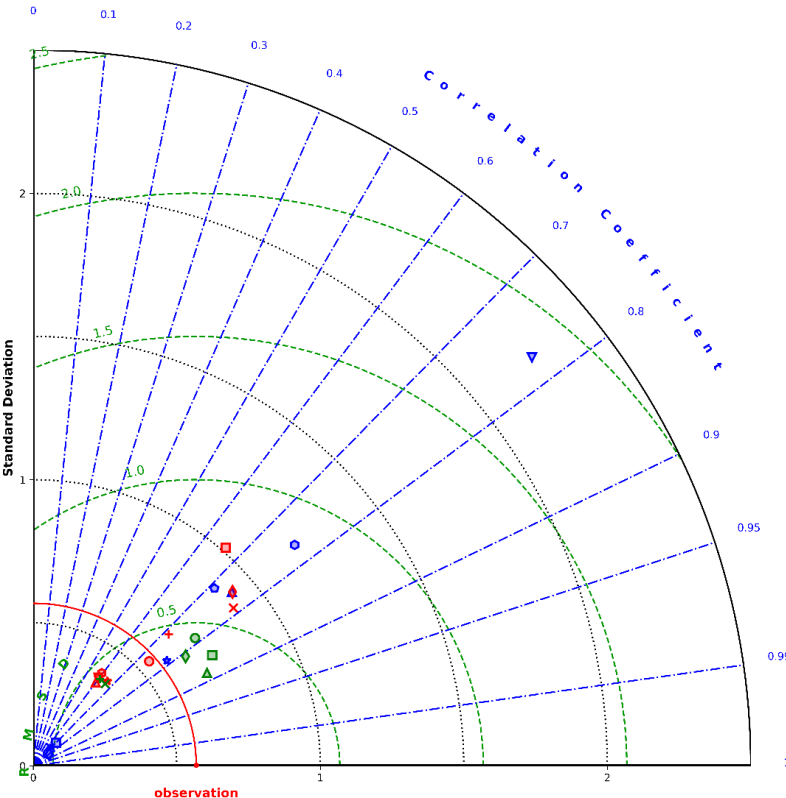




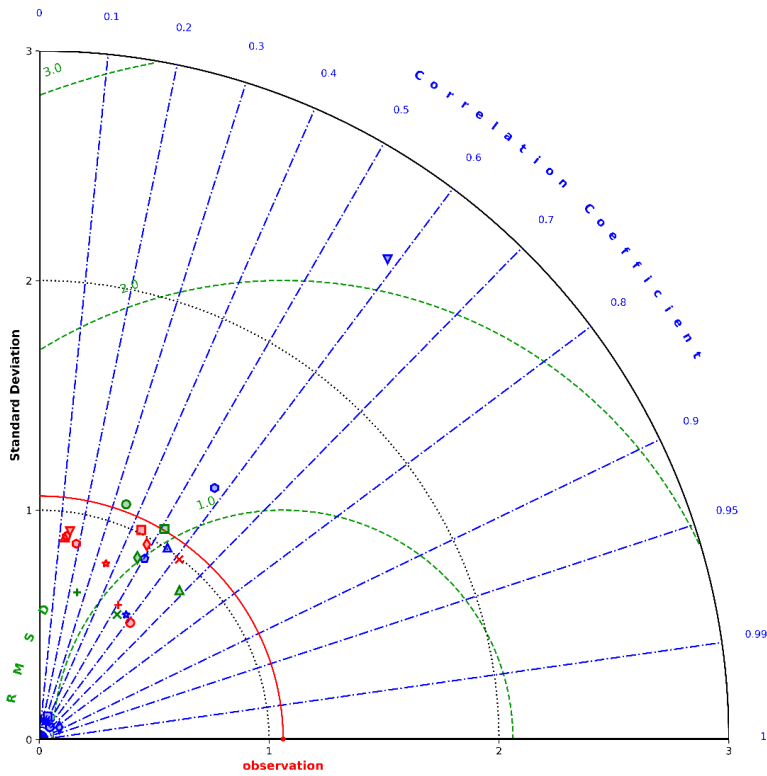
S1(d) Taylor diagram for station CIV3001



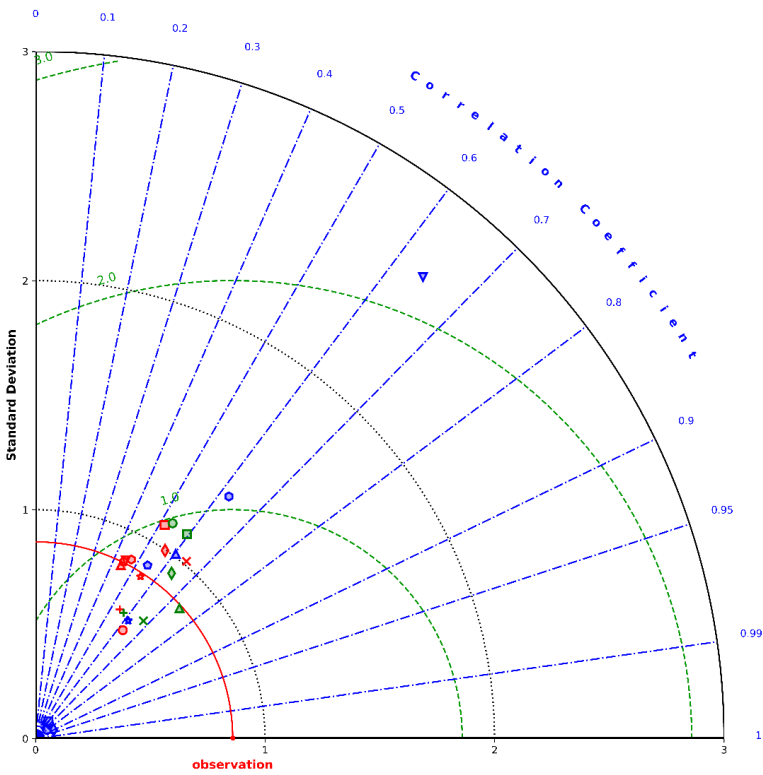
S1(e) Taylor diagram for station CIV3002



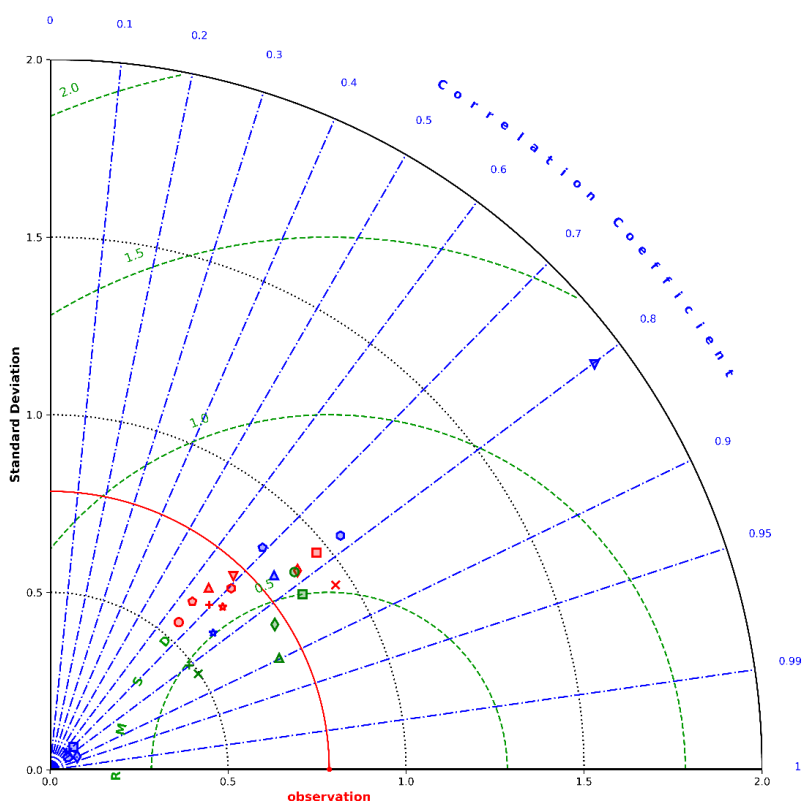
S1(f) Taylor diagram for station CIV4001



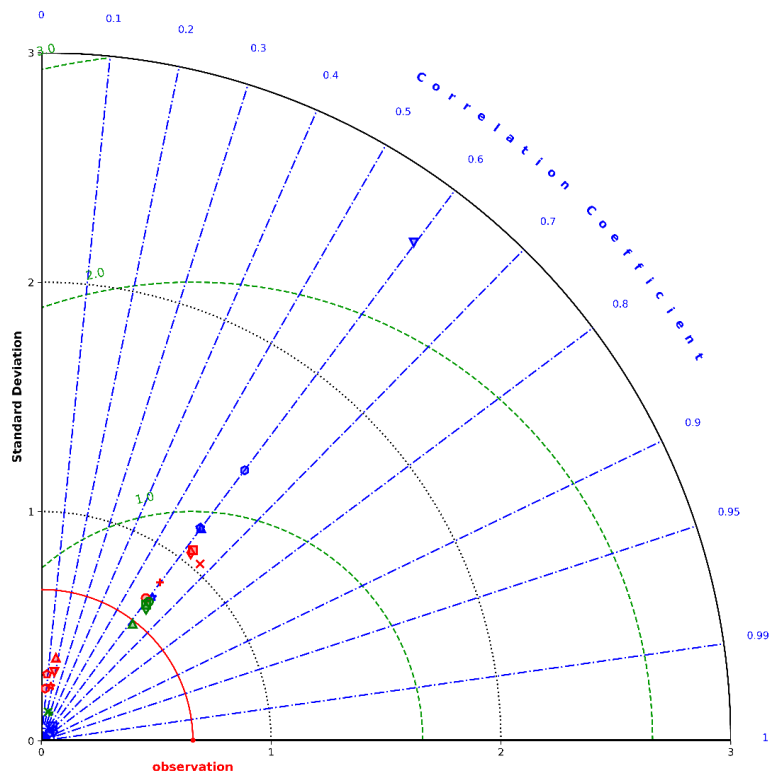
S1(g) Taylor diagram for station CIV4002



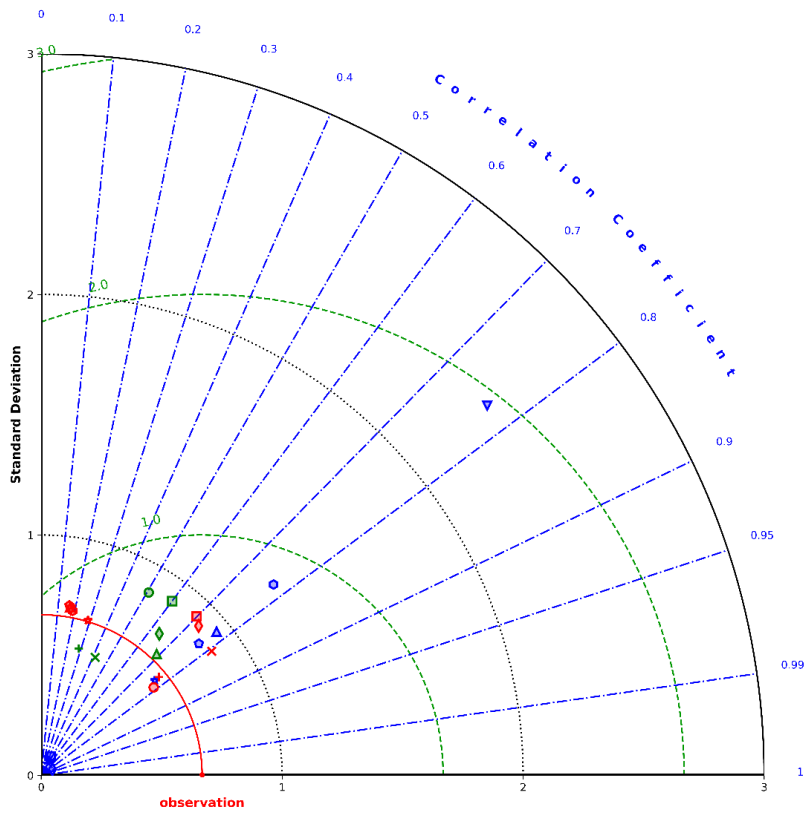
S1(h) Taylor diagram for station CMR3001



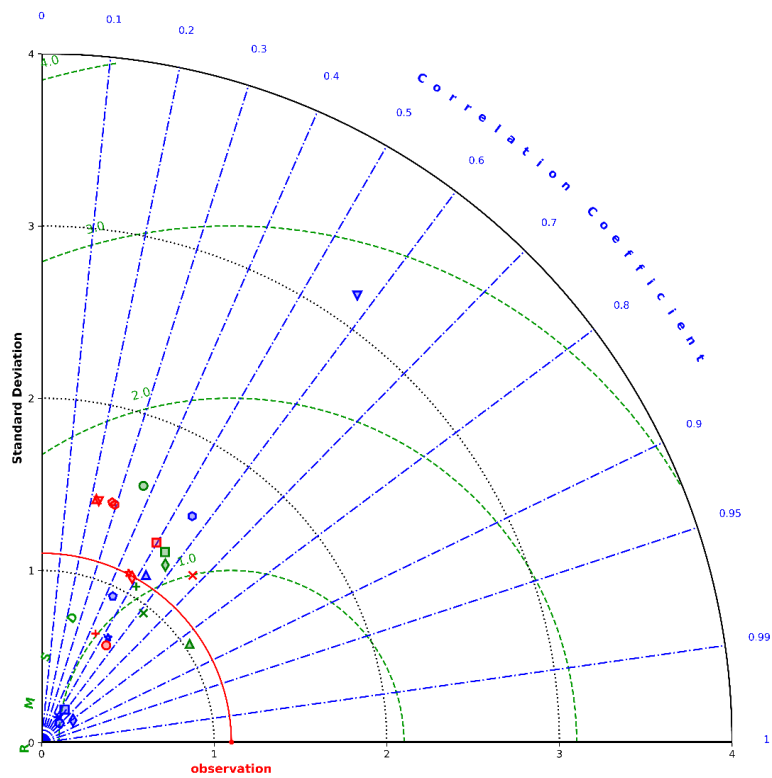
S1(i) Taylor diagram for station COD3001



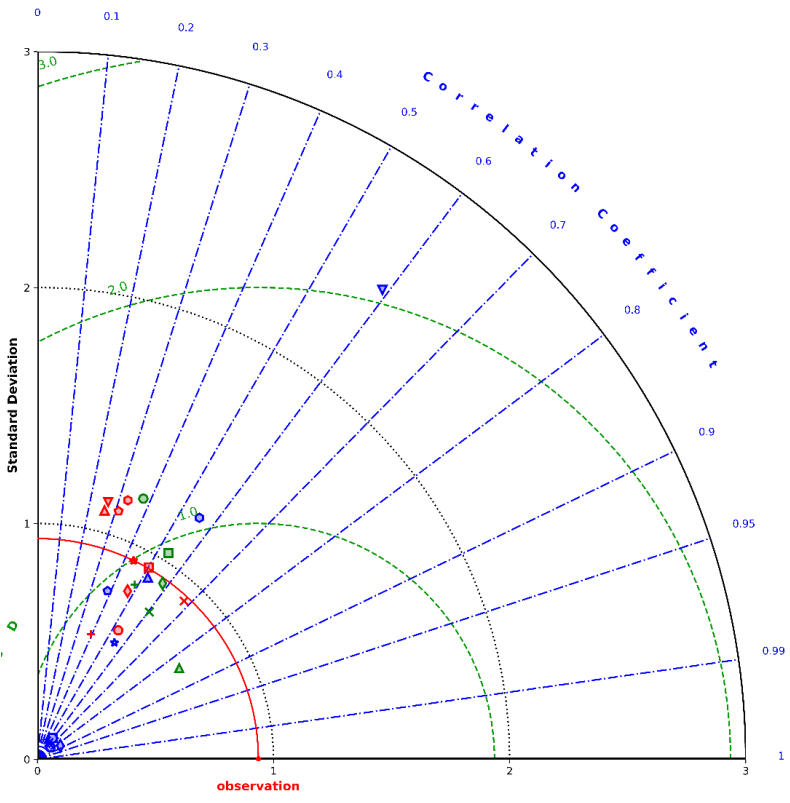
S1(j) Taylor diagram for station GHA3001



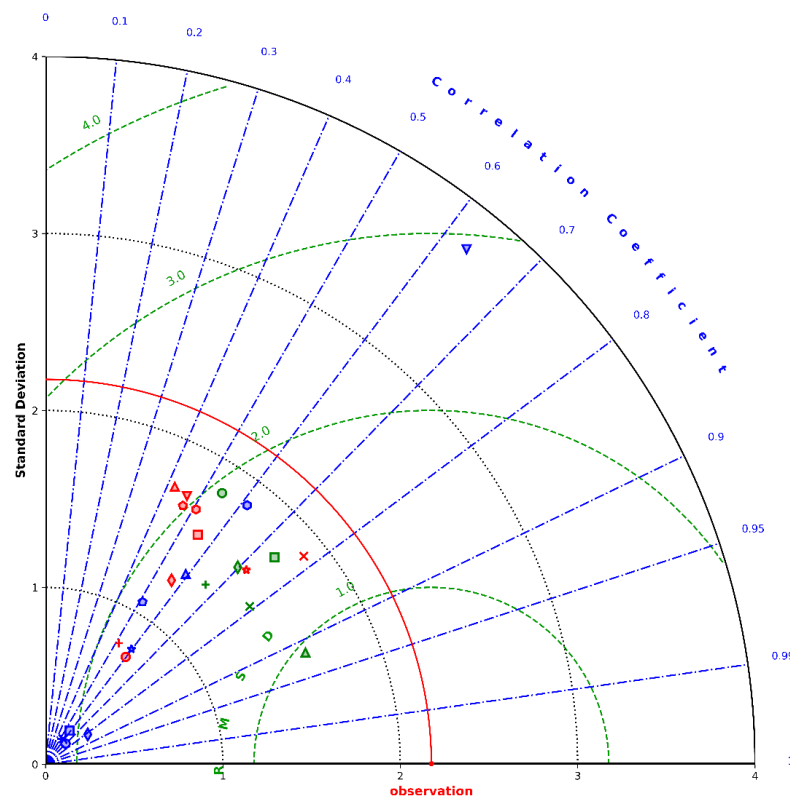
S1(k) Taylor diagram for station GHA3002



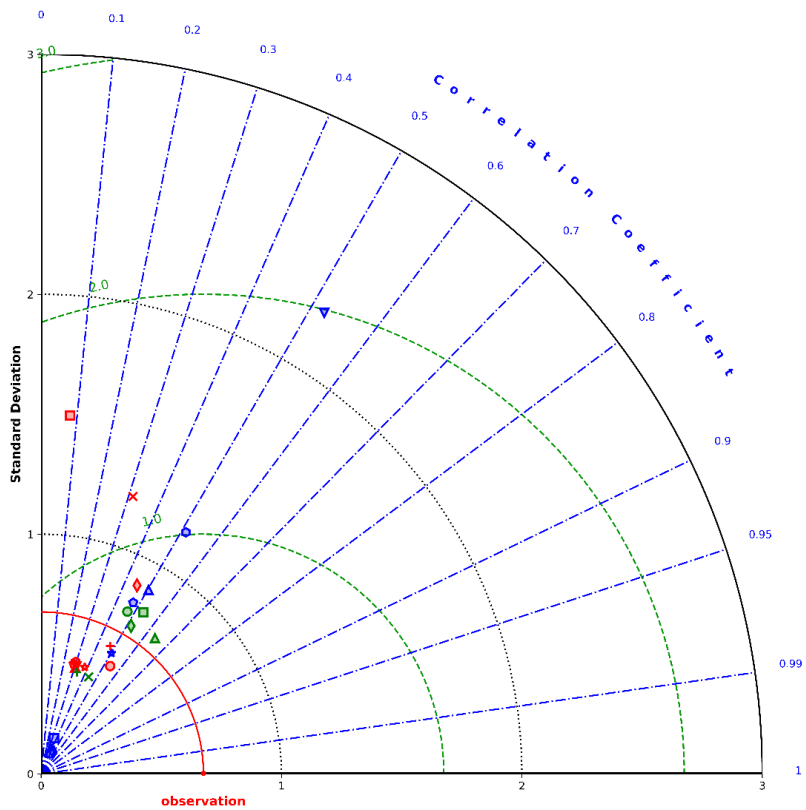
S1(l) Taylor diagram for station GIN3001



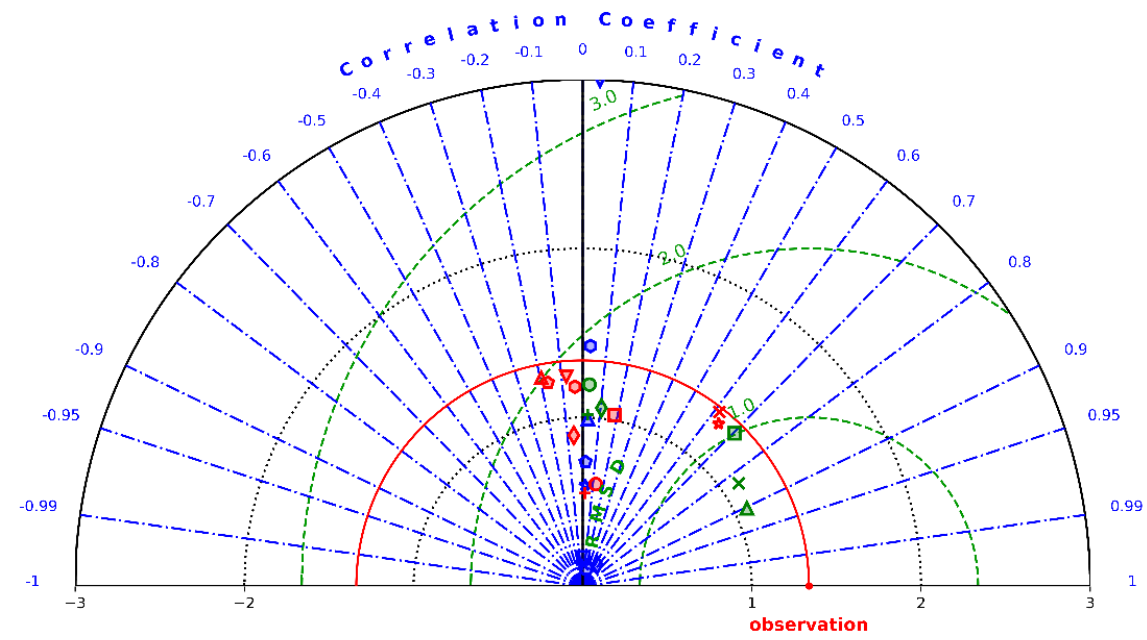
S1(m) Taylor diagram for station GMB3001



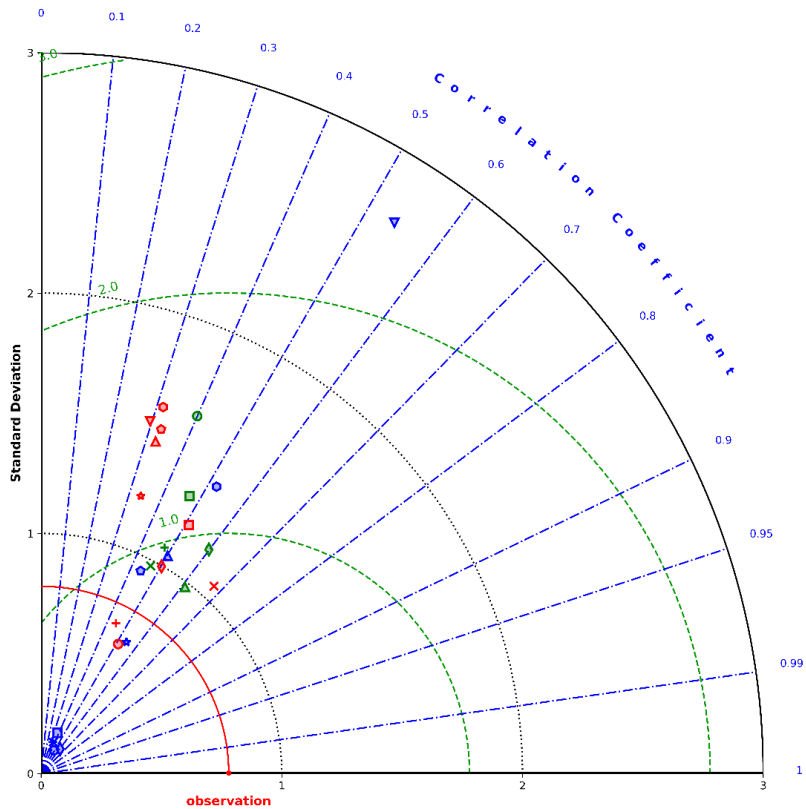
S1(n) Taylor diagram for station LBR3001



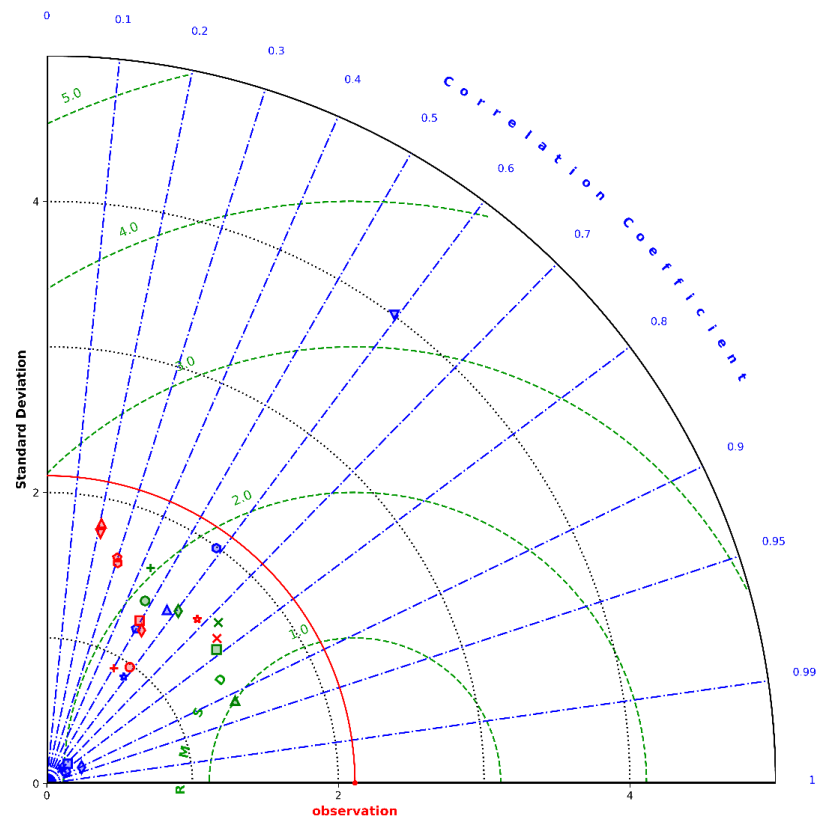
S1(o) Taylor diagram for station NER3001



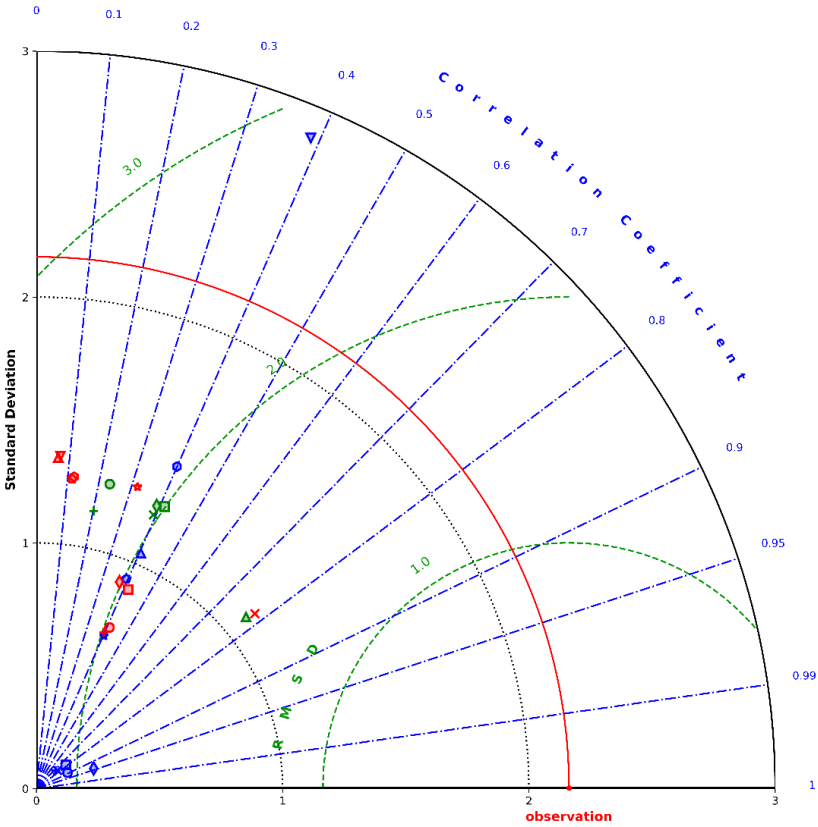
S1(p) Taylor diagram for station NGA3001



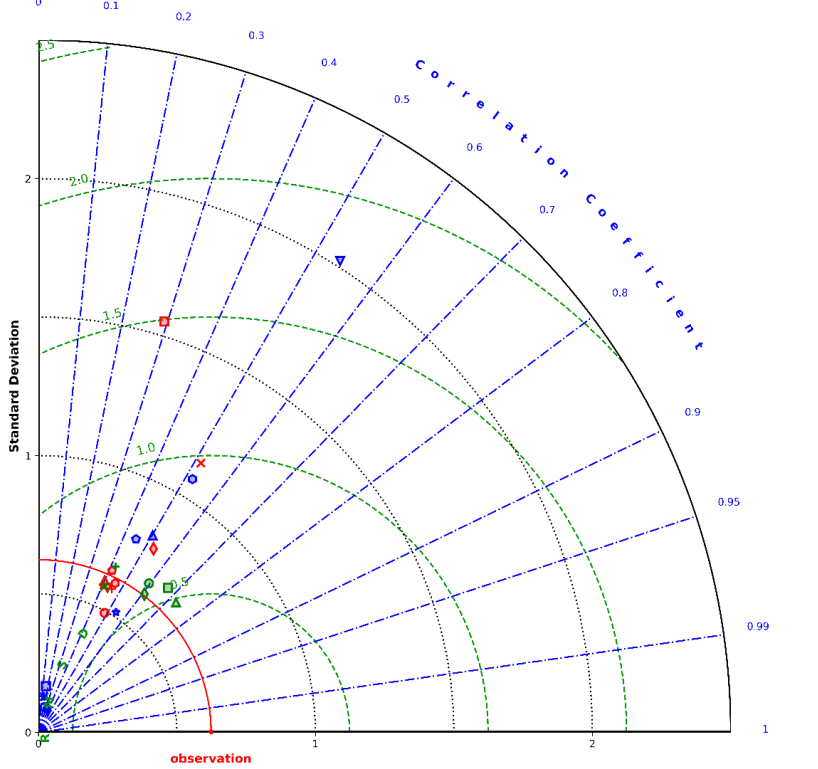
S1(q) Taylor diagram for station SEN3001



S1(r) Taylor diagram for station SEN4001

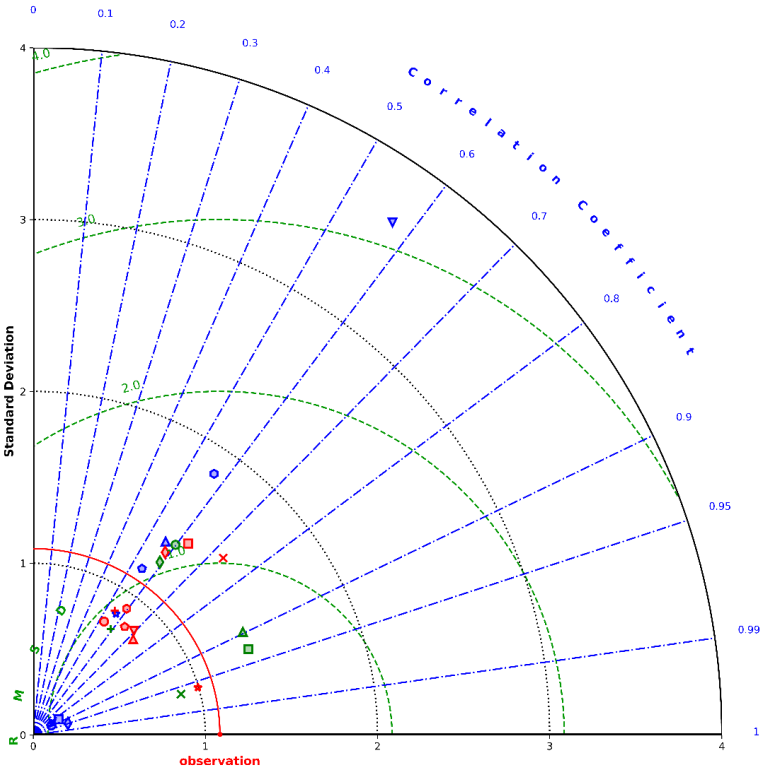


S1(s) Taylor diagram for station SLE3002

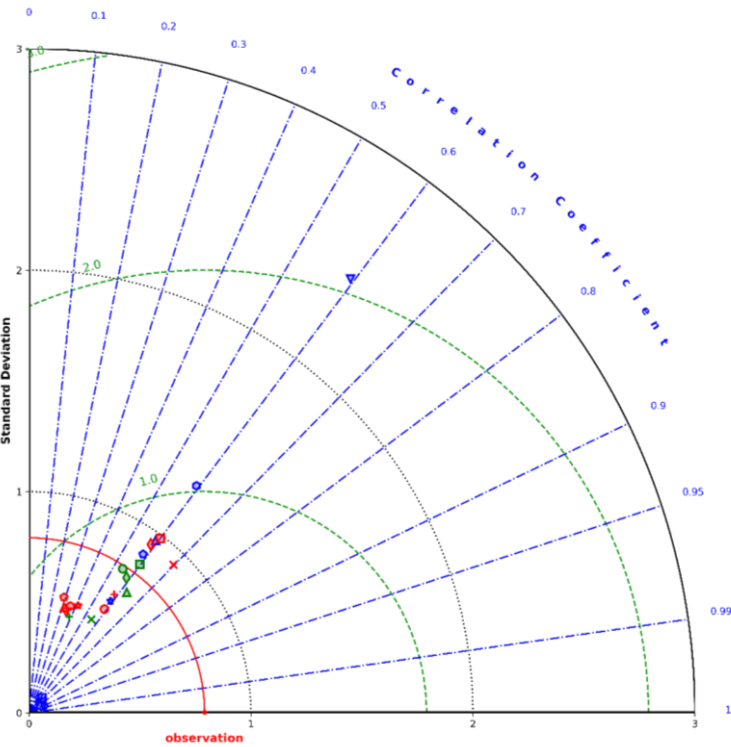




S1(t) Taylor diagram for station TCD3001



S1(u) Taylor diagram for station TGO3001



### S1(v) Symbol legend for plotted models

+	ET0-dalton	□	ET0-valiantzas-b
○	ET0-trabert	◇	ET0-valiantzas-c
×	ET0-penman	△	ET0-RF-M1
□	ET0-rohwer	▽	ET0-RF-M2
◇	ET0-mahringer	☆	ET0-RF-M3
△	ET0-hargreaves	⊕	ET0-RF-M4
▽	ET0-jensen-haise	+	ET0-RF-M5
☆	ET0-abtew	+	ET0-RF-M6
⊕	ET0-oudin	○	ET0-RF-M7
☆	ET0-irmak-a	×	ET0-RF-M8
+	ET0-irmak-b	□	ET0-RF-M9
⊕	ET0-irmak-c	◇	ET0-RF-M10
×	ET0-valiantzas-a	△	ET0-RF-M11

# Supplementary Tables

## Supplementary Table 1

### Summary of Selected Literature on Machine Learning Applications for ET<sub>0</sub> Estimation

This table presents a curated selection of recent studies that compare and utilise machine learning methods to estimate reference evapotranspiration (ET<sub>0</sub>). The literature compilation was achieved through a search for "machine learning ET<sub>0</sub> estimation" in Google Scholar for the years 2016-2021. The table collates the most pertinent studies identified in the initial pages of search results, with additional relevant papers supplemented by the authors' further reading.

Reference	Compared machine learning models	Origin and type of data	Climate	Location
(Feng et al., 2016)	extreme learning machine (ELM), neural networks optimized by genetic algorithm (GANN), wavelet neural networks (WNN)	13 weather stations, 1994-2013 period, daily data	warm humid	Central Sichuan, China
(Gocic et al., 2016)	ELM	2 weather stations, 1980-2010 period, daily data	ND	Serbia
(Kisi, 2016)	least square support vector regression (LSSVR), multivariate adaptive regression splines (MARS), M5Tree	2 weather stations, 1982-2006 period, monthly data	mediterranean	Western Turkey
(Patil and Deka, 2016)	ELM, artificial neural networks (ANN), least square support vector machine (LSSVM)	2 weather stations, 1970-2010 period, weekly data	arid	Rajasthan, India

(Yassin et al., 2016)	ANN, gene expression programming (GEP)	19 weather stations, 1980-2010 period, daily data	arid	Saudi Arabia
(Yin et al., 2017)	Genetic algorithm support vector machine (GA-SVM), ANN	1 weather station, 2009-2010 period, daily data	temperate	Qilian Mountains, NW China
(Antonopoulos and Antonopoulos, 2017)	ANN	1 weather station, 2009-2013 period, daily data	mediterranean	Greece
(Feng et al., 2017a)	random forest (RF), generalized regression neural network (GRNN)	2 weather stations, 2009-2014 period, daily data	warm humid	Sichuan, SW China
(Feng et al., 2017b)	ELM, GRNN	6 weather stations, 1961-2014 period, daily data	warm humid	Sichuan, SW China
(Karimi et al., 2017)	GEP, support vector machine (SVM)	8 weather stations, 1997-2006 period, daily data	humid	South Korea
(Mehdizadeh et al., 2017)	GEP, polynomial SVM, radial basis function SVM, MARS	44 weather stations, 1951-2010 period, monthly data	arid	Iran
(Pandey et al., 2017)	evolutionary regression (ER), ANN, multi-nonlinear regression (MNL), SVM	2 weather stations, 2008-2012 period, daily data	ND	Arunachal Pradesh, India
(Fan et al., 2018)	RF, M5Tree, gradient boosting decision tree (GBDT), extreme gradient boosting/XGBoost	8 weather stations, 1961-2010 period, daily data	various	China

	(XGB), SVM, ELM			
(Keshtegar et al., 2018)	adaptive neuro-fuzzy inference system (ANFIS), M5Tree, ANN	3 weather stations, 20 year data (period unknown), daily data	cold arid	Central Anatolia, Turkey
(Kisi and Alizamir, 2018)	wavelet extreme learning machine (WELM), WNN, ANN, online sequential ELM (OS-ELM)	2 weather stations, 1985-2005 period, daily data	semi arid	Central Anatolia, Iran
(Mehdizadeh, 2018)	MARS, GEP with lagged ET0	6 weather stations, 2000-2014 period, daily data	semi arid, arid, hyper arid	Iran
(Zhang et al., 2018)	SVM, ANN, ANFIS	9 weather stations, MODIS remote sensing data, 2010-2013 period, daily data	ND	Hexi corridor, NW China
(Ehteram et al., 2019)	SVM with cuckoo algorithm (SVM-CA), gaussian process regression (GPR), M5Tree, ANFIS	1 weather station, 1990-2016 period, daily data	humid	Himalaya, India
(Ferreira et al., 2019)	ANN, SVM	203 weather stations, 2001-2015 period, daily data	various	Brazil
(Granata, 2019)	M5Tree, bagging, RF, support vector regression (SVR)	1 weather station, 2000-2004 period, daily data	subtropical humid	Florida, USA
(Huang et al., 2019)	gradient boosting with categorical features support/CatBoost (CAT),	5 weather stations, 2001-2015 period,	humid	southern China

	RF, SVM	daily data		
(Mosavi and Edalatifar, 2019)	ANFIS, particle swarm optimization (PSO), principal component analysis (PCA)	2 weather stations, 1981-2011 period, monthly data	arid	Iran
(Nourani et al., 2019)	feed forward ANN, ANFIS, SVR	14 weather stations, 1987-2018 period, monthly data	various	Turkey, Cyprus, Iran, Iraq, Lybia
(Reis et al., 2019)	ANN, ELM, multiple linear regression (MLR)	5 weather stations, 1996-2016 period, daily data	warm tropical, dry winter	Verde Grande Basin, Brazil
(Wu and Fan 2019)	multi-layer perceptron (MLP), GRNN, ANFIS, SVM, kernel-based nonlinear extension of Arps decline (KNEA), M5Tree, XGB, MARS	14 weather stations, 2001-2015 period, daily data	various	China
(Wu et al., 2019)	ANN, RF, GBDT, XGB, MARS, SVM, ELM, KNEA	11 weather stations, 2011-2015 period, daily data	subtropical humid	Jiangxi province, China
(Zhang et al. 2019)	SVM	5 weather stations, 1990-2014 period, daily data	ND	Shaangxi, China
(Ashrafzadeh et al., 2020)	seasonal autoregressive integrated moving average (SARIMA), SVM	4 weather stations, 1993-2014 period, daily data	ND	Northern Iran
(Chia et al., 2020)	SVM	4 weather stations 2014-2018 period, daily data	tropical	west coast Malaysia
(Dos Santos	cubist regression, ANN, SVM,	23 weather stations,	various	MATOPIBA,

Farias et al., 2020)	MLR	2000-2016 period, daily data		Brasil
(Hossein Kazemi et al., 2020)	GEP	7 weather stations, 2008-2017 period, daily data	arid	Iran
(Kim et al., 2020)	RF	9 weather stations + MODIS remote sensing data + numerical weather prediction data, 2013-2019 period, daily data	ND	South Korea
(Mohammadi and Mehdizadeh, 2020)	SVR with whale optimization algorithm (WOA)	3 weather stations, 2000-2014 period, daily data	semi arid, arid, hyper arid	Iran
(Wu et al., 2020)	Adaptive boosting (ADA), GBDT, XGB, light gradient boosting decision machine (LGB), CAT	10 weather stations, 1997-2016 period, daily data	temperate, subtropical and tropical	eastern monsoon zone, China
(Yu et al., 2020)	ANN, SVR, ELM	1 weather station, 2005-2017 period, daily data	semiarid, arid	Xinjiang Uygur, China
(Zhang et al., 2020)	CAT, GRNN, RF	15 weather stations, 1996-2015 period, daily data	arid, semi- arid	northern China
(Adnan et al., 2020)	ANN, MARS, M5Tree	2 weather stations, 1968-2015 period, daily data	mediterranea n	southern Turkey

(Ayaz et al., 2021)	LSTM, GBDT, RF, SVR	2 weather stations, 1965-2015 period, daily data	hot semi arid, west coast marine climate	Hyderabad, India, Waipara, New Zealand
(Bellido-Jiménez et al., 2021)	MLP, ELM, GRNN, SVM, RF, XGB	5 weather stations, 2000-2018 period, daily data	semi arid	Andalusia, Spain
(Chia et al., 2021a)	MLP, SVM, ANFIS, bootstrap aggregating, Bayesian model averaging (BMA), ELM	4 weather stations, undetermined period, daily data	ND	peninsular Malaysia
(Chia et al., 2021b)	ELM-PSO hybrid, ELM-moth-flame optimization (MFO) hybrid, ELM-WOA hybrid	3 weather stations, 2014-2018 period, daily data	tropical	Eastern Malaysia
(Rashid Niaghi et al., 2021)	GEP, SVM, MLR, RF	6 weather stations, 2003-2016 period, daily data	continental	North Dakota and Minnesota, USA
(Sattari et al., 2021)	GPR, SVR, Broyden Fletcher Goldfarb Shanno ANN, LSTM	1993-2018 period, monthly regional measurements	arid, semi-arid	Turkey
(Wu et al., 2021)	ELM coupled with K-means clustering and firefly algorithms, ANFIS, RF, M5Tree	26 weather stations, 1966-2015 period, monthly data	ND	Poyang Lake basin, Southern China
(Yan et al., 2021)	XGB-WOA	8 weather stations, 1966-2015 period, daily data	arid, humid	China



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