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# **Characterize the magnetic signal generated in the**  CIDITS **magnetosphere from geomagnetic observations from 1999 to 2024**

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Study the magnetospheric field up to spherical harmonic (SH) degree 6 (L=6) over 26 years from 1999 to 2024 covered 2 solar cycles through normal distributions (  $\kappa(\mu=m,\,\sigma^2=\mathcal{C}_m)$  ) with a time resolution of an hour, based on a Kalman filter and correlation-based technique.

mean prior model  $m_0$ , prior covariance matrix  $C_0$  as theoretical statistical prior information

main magnetic field **EXAPP CONTRIBUTION IS removed from the** hourly-mean INTERMAGNET observatory data over 26 years from 1999 to 2024 by subtracting the MCM model ( Ropp et Lesur [2020])

geomagnetically quiet time. We then derive a time series of model coefficients describing the external field and its induced response in the Earth's interior and compare their time variations with the magnetospheric ring current index (RC), a magnetic proxy determined by other methodologies (Leonie et al. [2017]). The correlation between our axial dipole spherical harmonic coefficient (q10) and the external RC index is 0.9. Semi-annual, monthly, and daily variations, as well as tidal variations, are highlighted in the temporal spectra domain of magnetospheric gauss coefficients. A key element of our approach is ensuring that all model parameters are accompanied by robust error estimates.

### **Objective**:

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# **Introduction**

In the near-Earth geomagnetic field region (Langel et al. [1996]), the main contribution to the geomagnetic field is from the core, but fields generated from the magnetosphere during magnetic storms may have geomagnetic field contributions up to hundreds of nanoteslas (nT). A challenging task is to separate the magnetospheric fields from other contributing sources of internal and external origin. Magnetic field measurements are available at different altitudes, from the surface to the orbits of the Low Earth Orbiting satellites. Magnetic survey satellite data provide an homogeneous spatial resolution but their spatiotemporal resolution is coarse thus limiting our ability to study rapidly varying external magnetic fields (CC. Finaly et al. [2017]). In contrast, ground geomagnetic observatories provide datasets with high temporal resolution of an hour over a large range of latitudes and longitudes that are particularly well suited to characterize magnetospheric signals. We describe here an approach to modeling the magnetospheric signal by replying on a Kalman filter approach (Kalman [1960]) and correlation-based technique (Holshneider et al. [2016]). We use hourly averaged magnetic field triplet measurements at the location of 218 magnetic observatories (see Figure 1) during

> **Figure 1**: Distribution of the 218 geomagnetic observatory positions over 26 years

• **Prior information:** 

• **Analysis step:**

 adjustment of models at the *k*th hour by fitting INTERMAGNET observatory data, based on the Least Squares method

$$
d = Am + e
$$
  

$$
n_k^* = m_k + (A^t C_e^{-1} A + C_k^{-1})^{-1} A^t C_e^{-1} (d - Am_k)
$$
  

$$
C_k^* = (A^t C_e^{-1} A + C_k^{-1})^{-1}
$$

### • **Prediction step:**

# **I Approach -- Kalman Filter**

 prediction of the next hour's model (at the *(k+1)*th hour) from the adjusted model at the step *k* and based on a prediction operator P

$$
m_{k+1}=Pm_k^*\\
$$
  

$$
C_{k+1}=PC_kP^t+C_w
$$

**Figure 2:** Annual correlation coefficient between the  $q_1^0$  and the Dst index (top); between  $\boldsymbol{q_1^0}$  and RC index (bottom) from 1999 to 2024. The orange curve presents the correlation for geomagnetically quiet time and the blue one for all times.

**Table 1:** Root-Mean-Square Errors (RMSE) in nT and Correlation Coefficient (CC) between the RC index, Dst index and minus the modeling SH Gauss coefficient (top:  $q_1^0$ , bottom:  $i_1^0$ ) from 1 January 1999 to 1 August 2024.

Quiet time

 $-27.5$ 

#### **RMSE**  $cc$ Average  $RC<sup>e</sup>$  $RC^e$  $-q_1^0$ Est Est  $RC^e$ Est **Full period**  $-11.68$  $-10.63$ 12.62 13.16  $0.7$ 0.53  $-9.19$ Quiet period (-30<Dst<30 nT)  $-6.68$  $-8.55$  $0.7$ 0.92  $-5.3$ 4.26 8.47 Active period (|Dst|≥30 nT) 29.7 31.92  $-0.17 -0.17$  $-42.1$  $-23.25$  $-36.4$  $RC<sup>i</sup>$  $RC^i$  $-i_1^0$  $RC<sup>i</sup>$ **Ist Ist** Ist **Full period** 0.48  $-0.03 -0.77$ 4.63 4.81 0.46  $-2.94$ Quiet period (-30<Dst<30 nT)  $-0.5$ 2.51 2.17  $0.77$  0.74 1.3  $-5.3$ Active period ( $|Dist| \ge 30$  nT) 10.64 12.35 -0.06 -0.19  $-8.17 -2.4 -11.65$



Full time

### • **Smoothing step:**

 a posteriori smoothing of the calculated series, based on Gaussian hypothesis for the correlation time

```
m_k^s = m_k^* + G_k(m_{k+1}^s - m_{k+1})C_k^s = C_k^* - G_k(C_{k+1} - C_{k+1}^s)G_k^t
```
# **INTERMAGNET observatory data**



2, Observatory data are selected during geomagnetically quiet time (Dst between -30 nT and 30 nT) and cleaned for outliers.

The geomagnetic vector components after the selection are modeled in terms of spherical harmonics with the sum of the magnetospheric fields, induced fields, crustal offsets and ionospheric statistical contribution at each observatory location:

$$
r, \phi) = \sum_{\ell,m} q_{\ell}^{m} \left(\frac{r}{a}\right)^{\ell-1} Y_{\ell}^{m}(\theta, \phi)
$$
  
+ 
$$
\sum_{\ell,m} i_{\ell}^{m} \left(\frac{a}{r}\right)^{\ell+2} Y_{\ell}^{m}(\theta, \phi)
$$
  
+ 
$$
\sum_{i \in N_{obs}} \delta(r - r_{i}, \theta - \theta_{i}, \phi - \phi_{i}) O_{i}
$$
  
+ 
$$
\sum_{i \in N_{obs}} \delta(r - r_{i}, \theta - \theta_{i}, \phi - \phi_{i}) I O_{i}
$$

 $B($ 



## **Parameters**

# **Description of the hourly models**

### **Future works**

Figure 7: asymmetric structure. Left: The 26-year average of the ratio between external GC  $q_2^0$  and  $q_1^0$  (blue bar chart) and of the ratio between  $q_3^0$  and  $q_1^0$  (orange bar chart) in function of the months. Right: The 26-year average of the  $q_1^0$ on function of the months.

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### **References**

# **Assessment of the hourly models**

- Applying the methodology during the geomagnetically active periods
- Applying better statistics from theoretical or semi-empirical models in order to separate the different contributions in the magnetosphere and develop a better description of magnetospheric time variations over solar cycles



**Figure 4:** The geomagnetic observatory positions are shown with 26y residual σ (diagonal elements of Fig.3: right) greater than 10nT during geomagnetically quiet time. The colorbar indicates the average σ over 3 components (x,y,z)

**Figure 3:** The covariance matrix of residuals between observatory data and modelling signal for the 218 observatories over 26 years during all geomagnetic times (Left) and quiet times (Right). The shape of the covariance matrix is (3×218,3×218).

### • **Comparison with RC and Dst index**



• **Data Fitting**

**Figure 5:** Average Power spectrum of hourly Spherical Harmonic (SH) Gauss coefficient as a function of SH degree over 26 years for all geomagnetic time (Left) and for geomagnetically quiet time only (Right). The solid lines are a priori power spectrum (black for the magnetospheric field, grey for the induced field). The colored lines are posterior power spectrums (blue for the magnetospheric field, red for the induced field).



**Figure 6:** The temporal spectrogram analysis of magnetospheric modeling Gauss Coefficient. Top: The median temporal spectra of each Gauss coefficient (in each row) up to degree 6. Bottom: The zoom of the black box of top figure, the median temporal spectrum of the external Gauss coefficient for degree L=1. The bleu spectrum is for q10 signal, the orange for q11 and the green for q1-1. The median temporal spectrum covered periods from 2 hours to 8760 hours (one year).



### • **Solar cycles**

• **Asymetric structure, Russel-Mcpherron effect**







### • **Temporal power spectra**

