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- 1 Effects of monsoons and storms on the structuring and diversity of picoeukaryotic 2 microbial communities in a tropical coastal environment
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- 10
- 11 Abstract

Picoeukaryotes are key components in marine ecosystems that play crucial roles in food webs 12 and biogeochemical cycles. Despite their significance, many aspects of their community 13 ecology and diversity remain understudied. Here, we investigated the taxonomic and functional 14 15 diversity of picoeukaryotic communities in response to monsoonal patterns and weather 16 disturbances brought about by storms, characterizing tropical coastal regions. To do this, water samples were collected almost weekly or bi-weekly at a single location in a tropical coastal 17 environment covering the late northeast (NE) and southwest (SW) monsoons. We then 18 performed high-throughput amplicon sequencing of the V4 region of the 18S rRNA gene to 19 20 generate taxonomic profiles of the communities across time. Clustering based on environmental parameters grouped our samples into months associated with NE monsoon, SW 21 monsoon, and stormy SW monsoon, demonstrating seasonality influenced by monsoons and 22 storms, typically observed in tropical coastal waters. In comparison, clustering based on 23 abundance only grouped the samples into NE and SW monsoon, with most communities during 24 storm period joining the NE monsoon samples. These samples exhibited greater diversity, with 25 smaller taxa such as Syndiniales, Prymnesiophyceae, Picozoa, Cercozoa, Stramenopiles, and 26 Chlorophytes being the most abundant groups present. In contrast, SW monsoon samples have 27 28 lower diversity but have become generally dominated by large-celled taxa, mostly diatoms. 29 Multivariate and correlation analyses both revealed nitrate as the strongest environmental driver of the picoeukaryotic community structuring. Meanwhile, network analysis grouped the 30 31 taxa into three modules, more consistent with the clustering based on environmental 32 parameters, implying that although storms may not significantly change the community composition, they may however influence the dominating taxa. Each module was composed of 33 a unique set of co-occurring taxa, highlighting high turnover of picoeukaryotic communities 34 between each season. In addition, our results showed that SW monsoon-associated module had 35 higher interconnectivity than other modules, suggesting that the interactions during this period 36 37 may be less species-specific, thus, more adaptable than during NE monsoon. However, we observed that extreme fluctuations caused by storms could have possibly allowed for selection 38 of dominant taxa. Shotgun metagenomic sequencing of representative samples from each 39 40 monsoon period also revealed that differently abundant functional genes, particularly genes 41 associated to nitrogen metabolism, might have also helped in adaptation to the changing nutrient conditions. Our observations provide new insights on the potential trajectory of
 microbial communities under environmental stresses, which are important in understanding the
 implications of emerging threats such as coastal eutrophication and climate change.

- Keywords: picoeukaryotic communities, seasonal monsoons, tropical coastal ecosystems, 18S
 rRNA gene, shotgun metagenomics, co-occurrence network
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- 48

49 1. INTRODUCTION

50 Marine picoeukaryotes are unicellular eukaryotes belonging to the size fraction of less than 3.0 µm (Worden & Not, 2008). They are ubiquitous in marine ecosystems and play crucial roles 51 52 in the biogeochemical cycles (Massana et al., 2004; Vaulot et al., 2008) with their functional 53 role in the ecosystem being defined by their trophic modes (Worden & Not, 2008). 54 Traditionally, picoeukaryotes can be classified as photoautotrophs, which are considered part 55 of the phytoplankton that can perform photosynthesis, or as heterotrophs, which play a key role in remineralizing organic matter in the ocean (Massana et al., 2004; Not et al., 2009; Xie et al., 56 57 2021). Some picoeukaryotes may also have both phototrophic and heterotrophic capabilities and are considered as mixotrophs (Sanders & Gast, 2012). However, studying picoeukaryotes 58 have been challenging owing to their very small size and lack of morphological landmarks 59 (Massana et al., 2004), resulting in the limited understanding of their ecology and diversity, 60 despite contributing significantly to overall biomass and carbon standing stock (Worden et al., 61 2004). The advent of molecular methods (e.g., cloning, high throughput sequencing) have 62 consistently confirmed the high diversity of picoeukaryotes, with majority of the recovered 63 operational taxonomic units (OTUs) belonging to novel lineages, highlighting the need for 64 more in-depth studies on their taxonomic and even functional diversity (Massana et al., 2004; 65 Romari & Vaulot, 2004). 66

Previous works demonstrated that picoeukaryotes also exhibit community patterns and 67 structuring akin to their larger counterparts, where changes in the environmental conditions can 68 significantly influence their community structure (Lambert et al., 2019; Logares et al., 2020). 69 Specifically, a few studies demonstrated that environmental conditions linked to seasonal 70 71 variability could potentially be more significant driver than differences caused by spatial 72 variability. This is particularly true for regions with pronounced seasonal shifts such as in the 73 temperate and polar waters. In a temperate upwelling region, for example, multivariate analysis 74 grouped the samples according to season, but not by location (Rocke et al., 2020). Similarly,

in a temperate coastal basin, the community structure of pico- and nanoeukaryotes significantly 75 differed between summer and winter, but not between sampling sites (Wang et al., 2020). 76 However, most of these studies were done in temperate or polar waters or in freshwater bodies, 77 which cannot represent the same conditions as in marine tropical waters that experience less 78 pronounced seasonality. Initial work on marine picoeukaryotes in the tropics revealed that 79 80 temperature may potentially have the most significant influence on their community structuring (dela Peña et al., 2021; Liu et al., 2021; Logares et al., 2020). For example, a recent study in 81 the Western Pacific Ocean, encompassing the regions influenced by the North Equatorial 82 83 Current, Kuroshio Current, and Mindanao Current, showed that environmental variation, particularly changes in temperature, dissolved oxygen (DO), phosphate, and silicate 84 concentrations, was the main driver of picoeukaryotic community structuring in this area (Liu 85 86 et al., 2021). However, in a study using the data from both *Malaspina-2010* and *Tara Oceans* expeditions covering large or 'planetary' geographic scale, it was reported that dispersal 87 88 limitation was more influential in shaping picoeukaryotic communities in tropical and subtropical surface waters than heterogeneous (environmental) selection (Logares et al., 2020). 89 90 In addition, other than environmental factors, co-occurrence analyses derived from the *Tara* Oceans data consistently reported that biotic interactions and positive associations dominate 91 92 over environmental factors in driving the structuring of planktonic community (Lima-Mendez 93 et al., 2015; Sunagawa et al., 2020; Vincent & Bowler, 2020). Open oceans however exhibit less heterogeneity over large spatial and time scales compared to coastal environments, which 94 95 are more vulnerable to short-term changes (Lambert et al., 2019).

In the tropics, the high seasonal variability along the coast is primarily brought about by the 96 97 periodic reversal of the winds during southwest (SW) and northeast (NE) monsoon (Rivera, 1997; Wyrtki, 1961). The NE monsoon, also known as the 'dry season', is typically observed 98 99 from November to March, and characterized by strong northeasterly winds and relatively less precipitation (Matsumoto et al., 2020). Meanwhile, the SW monsoon or 'wet season', from 100 101 June to September, is characterized by heavier precipitation and strong, southwesterly winds (Matsumoto et al., 2020). During this season, tropical cyclones are more frequently 102 103 experienced, which may cause disturbances in water conditions (Cayanan et al., 2011; Rivera, 1997). In particular, external pollution loads including silt, nutrients, and organic matter 104 brought about by runoffs during the typhoons from terrestrial sources can have significant 105 impacts on environmental conditions, and thus, also influence the microbial communities 106 (Lambert et al., 2019; Rivera, 1997). Previous studies investigating the phytoplankton 107

dynamics in such coastal setting have provided insights on how the community changes on a 108 seasonal basis and suggested possible environmental drivers during each season (Azanza & 109 Benico, 2013; Canini et al., 2013; dela Peña et al., 2021; Escobar et al., 2013; Yap et al., 2004). 110 These studies reported that phytoplankton community composition varied with changing 111 monsoons and inferred that the changes were associated with conditions associated with 112 monsoon periods. However, the same studies focused only on the microphytoplankton 113 communities and were limited using traditional methods such as microscopy, except for De La 114 Peña et al. (2021), where they observed the previously unexplored vast diversity of 115 116 picoeukaryotes in Philippine waters using high throughput amplicon sequencing. Recent 117 studies have also consistently reported the presence of picoeukaryotic groups in tropical coasts (Amin et al., 2021; dela Peña et al., 2021; Mitbavkar & D'souza, 2021; Qiu et al., 2019; 118 119 Rajaneesh et al., 2018). In particular, Amin et al. (2021) reported that the distribution of picoeukaryotes progressively decreased from the coast of Peninsular Malaysia towards the 120 121 open ocean during SW monsoon. These reports highlighted the contribution of picoeukaryotes to the productivity and biogeochemical cycles in a tropical coastal habitat. Understanding how 122 123 picoeukaryotic communities change with monsoonal shifts will provide a more accurate picture of the overall productivity and functioning of the tropical coastal system. Further, as the base 124 125 of the food web, changes in picoeukaryotic communities may also cascade to higher trophic 126 levels. Thus, there is a need to understand how picoeukaryotes are influenced by conditions brought about by monsoons and storms, especially in tropical coasts that are more vulnerable 127 to effects of intensifying storms and other impacts of climate change (Kurtay et al., 2021). 128

Here, our goal was to provide a more in-depth understanding of how picoeukaryotic 129 130 communities respond to environmental changes brought about by monsoonal shifts in coastal environments, including disturbances caused by tropical storms, by looking at the changes in 131 132 their taxonomic and functional diversity. We hypothesize that shifts in monsoons will result in significant changes in the physical and chemical regimes of the water column in coastal 133 134 habitats, also greatly affecting the proliferating microbial communities. To do this, we collected samples at a single location on an almost weekly or bi-weekly basis from January to October 135 136 (16 samples in total), capturing the communities from late NE to SW monsoon. We then performed high-throughput amplicon sequencing of the V4 region of the 18S rRNA gene of 137 the entire microbial eukaryotic community to obtain their taxonomic profiles across time. The 138 change in environmental conditions along with the monsoon is expected to be accompanied by 139 the dominance of taxa with specific adaptive mechanisms to the new conditions. This would 140

also be reflected in the abundant functional genes that may be differentially expressed between 141 monsoonal periods. For this, we performed shotgun metagenomic sequencing on select 142 samples representing each monsoonal period and investigated genes that may indicate 143 adaptation to prevailing environmental conditions. Finally, we also performed co-occurrence 144 network analysis, which extends community analysis beyond the diversity indices by 145 visualizing interactions between taxa including those previously uncharacterized (Vincent & 146 Bowler, 2020). This allowed gaining more insights on the underlying processes that structure 147 the community and predict possible changes of food web systems in response to monsoonal 148 149 shifts and storms.

150

151 2. MATERIALS AND METHODS

152 2.1. Sample collection and processing

To monitor the changes in the community across monsoon periods, a single sampling site 153 (16°22'56.7"N 119°54'46.4"E) was maintained at a tropical marine coastal area in front of the 154 Bolinao Marine Laboratory located in Bolinao, Pangasinan at the northwestern side of the 155 Philippines (Supplementary Figure S1). It is a relatively well-studied site where majority of 156 157 phytoplankton work has been done, including spatial and temporal investigations, evolving into becoming a model site for the Philippines (Azanza et al., 2004; Yap et al., 2004; Azanza & 158 Benico, 2013; Escobar et al., 2013; Albelda et al., 2019; dela Peña et al., 2021). Surface 159 seawater samples (ca. 1 m below) were collected at an approximately weekly to biweekly basis 160 from January to October 2019 to investigate temporal variability of the entire picoeukaryotic 161 microbial community. 162

Water samples were collected using a Niskin-type bottle sampler (General Oceanics, Inc.). A 163 164 total of 2 L of water was collected each time and pre-filtered immediately through a 120 µm sieve to remove zooplankton and other large debris. Samples were collected in sterilized 165 Nalgene bottles (acid and sterile water washed) and immersed in ambient seawater in an icebox 166 during transport. Separate water samples for Chlorophyll a (Chl a) analysis were also collected 167 from the same Niskin bottle and kept in an icebox until they reached the laboratory. 168 Environmental parameters including temperature, salinity, and dissolved oxygen (DO) were 169 170 measured every sampling using a multiparameter (Hanna HI9829 and YSI Pro 2030).

In the laboratory, the pre-filtered seawater samples for DNA extraction were serially filtered 171 through a 42 mm polycarbonate (PC) 3 µm-pore size filter (Whatman), which represents the 172 large micro-nanoeukaryotes, and through a 0.2 µm Sterivex filter (Millipore) or a 42 mm 0.2 173 µm-pore size filter (Whatman and Membrane Solutions) corresponding to the picoeukaryotes. 174 A total of 5x10 mL of the serially filtered seawater was also collected and stored in separate 175 conical tubes used for nutrient analyses including nitrite, nitrate, phosphate, and silicate using 176 spectrophotometry following the methods based on Parsons et al. (1984). The filters were 177 transferred to 2 mL sterile microcentrifuge tubes and DNA was preserved by adding at least 178 179 1.6 mL RNALater (Ambio). Samples were stored at -20°C until extraction. For Chl a analysis, 180 500 mL of seawater samples were filtered through Whatman GF/F filters and extracted with 90% acetone and subsequently analyzed using a Trilogy laboratory fluorometer following the 181 182 method of Parsons et al. (1984).

183 2.2. DNA extraction and high throughput sequencing

Total genomic DNA was only extracted from the 0.2 µm-pore size filters to focus on the 184 185 picoeukaryotic communities using the DNeasy PowerSoil Kit (Qiagen, Germany) following the manufacturer's protocol with some modifications. Briefly, the filters were cut into smaller 186 187 pieces before being transferred to PowerBead tubes provided by the kit. An additional 0.5 mL of Buffer AL lysis buffer (Qiagen, Germany) was pipetted inside the tube to ensure complete 188 immersion of the filters to enhance cell lysis. Subsequently, 60 µL of Solution C1 was added, 189 190 then vortexed at maximum speed for 10 minutes. This was followed by the procedures 191 prescribed by the kit's manufacturer. DNA quality and yield were measured using spectrophotometry (Nanodrop, Thermo Fisher Scientific). Integrity of the extracted DNA was 192 193 also checked by agarose gel electrophoresis. In addition, PCR using universal eukaryotic primer pair 4616F and 4618R (Logares et al., 2007) targeting near full-length 18S rRNA gene 194 was performed before sending the samples for sequencing. 195

Here, two sequencing approaches were used, namely 18S rRNA gene amplicon sequencing to
probe on community shifts with monsoonal change and their co-occurrences (January to
October), and shotgun metagenome sequencing to further look at potential changes in dominant
functional genes with monsoons. For the targeted sequencing, all DNA extracts (*n*=16) that
passed the quality check were sent to the Integrated Microbiome Resource (IMR) housed in
Dalhousie University, Canada for Illumina multiplex sequencing using E572F/E1009R
forward and reverse primers (Comeau et al., 2011) targeting the V4 region of the 18S rRNA

203 gene. Raw amplicon sequences are available in the Sequence Read Archive (SRA) database of
204 NCBI under Accession No. PRJNA656691.

For shotgun metagenome sequencing, samples representing the intermonsoon (April) where community diversity is expected to be greatest (dela Peña et al., 2021) and towards the end of the southwest monsoon (September), were sent to Macrogen Genome Center (Macrogen NGS Division, Seoul, Korea). Libraries were prepared using the Nextera XT DNA Library Preparation Kit. Sequencing was performed on the Illumina NovaSeq 6000 platform with 150 bp paired-end setting. Raw shotgun sequencing data are also available in the SRA database of NCBI with Accession No. PRJNA788625.

212 2.3. Post-sequencing processing and analysis

Bioinformatics processing and quality filtering of the resulting 18S rRNA reads were done in 213 Quantitative Insights in Microbial Ecology v.2 (QIIME 2; Bolyen et al., 2019). Raw reads were 214 first imported as a QIIME 2 artifact using the gime tools 'import' method. After quality-215 checking, the reads were assembled with a minimum overlap of 12 bp using the q2-vsearch 216 plugin with the 'join-pairs' method (Rognes et al., 2016). Low-quality sequences were 217 removed, and the resulting data set was subsequently processed for Operational Taxonomic 218 219 Unit (OTU) clustering using the q2-vsearch plugin with the 'cluster-features-open-reference' method. Under this method, OTUs were clustered against an inhouse curated reference 220 database (Philippine Reference 18S rRNA Gene Database, unpublished), which uses a 221 combination of the SILVA v138.1, Nordicana Reference database v2.0 (Lovejoy et al., 2016), 222 223 and additional sequences amplified from Philippine isolates curated from available GenBankdeposited sequences at 98% similarity, and any sequences that had no match were then 224 clustered de novo (Rideout et al., 2014). Singletons and chimeras were removed using the q2-225 vsearch plugin with the 'uchime-denovo' method (Rognes et al., 2016). Assignment of 226 taxonomic identity was performed using the sci-kit learn plugin against the inhouse curated 227 reference database. Fungi, and metazoa-related reads were excluded from the analysis. The 228 final OTU dataset used for community analysis was then rarefied at 1,559 reads per sample 229 using single rarefaction based on the sample (Jun15BML) with the lowest number of reads to 230 make the diversity indices comparable. 231

For the shotgun sequencing data, both samples were first subjected to Kaiju (web server version) taxonomic classification (Menzel et al., 2016) using the maximum exact matches

(MEM) mode with a minimum match length of 11. Additionally, the nr+euk was used as the 234 reference database and the SEG low-complexity region filtering was enabled. Afterwards, 235 reads that were unclassified and were inferred as belonging to the Eukaryota superkingdom 236 (NCBI taxonomy ID: 2759) were extracted from both samples. The extracted reads were then 237 quality checked using FastQC v0.11.9 (Andrews, 2010) and trimmed and filtered using 238 Trimmomatic v0.39 (Bolger et al., 2014) (Parameters: SLIDINGWINDOW:5:18, 239 MINLEN:75, LEADING:3, TRAILING:3, AVGQUAL:20). Subsequently, the remaining read 240 pairs from each of the samples were merged using PEAR v0.9.11 (Zhang et al., 2014) and the 241 242 merged and unmerged reads were concatenated together into one file. To functionally annotate the shotgun data, the SUPER-FOCUS pipeline v0.35 (Silva et al., 2016) was utilized with 243 DIAMOND (Buchfink et al., 2021) as the aligner and DB_100 as the reference database 244 (Parameters: e-value=1e⁻⁵, minimum percent identity=60%, minimum alignment length=15 245 amino acids, normalization=0, FOCUS=0). Finally, the generated functional abundance table 246 247 of SUPER-FOCUS was imported into Statistical Analysis of Metagenomic Profiles (STAMP v. 2.1.3) (Parks et al., 2014) for differential abundance analysis. Statistical significance was 248 249 calculated using Fisher's exact test with Benjamini-Hochberg false discovery rate method of multiple test correction. Additionally, the Newcombe-Wilson method was applied to calculate 250 251 for the confidence intervals and features with a q-value of less than 0.05 were considered significant. Lastly, a combination of basic local alignment search tool (BLAST) and 252 phylogenetic analysis of the four most abundant photosynthetic genes was performed to 253 validate if they belong to the dominant species identified by the amplicon sequencing. Briefly, 254 255 quality filtered reads from the two samples were assembled individually using MEGAHIT (Li et al., 2015). Afterwards, all PsaA, PsbA, PsbB, and PsbC amino acid sequences were retrieved 256 from the manually curated subset of UniProt (UniProt Consortium, 2019), the SWISS-PROT 257 database (Boeckmann et al., 2003). The assembled metagenomes were then queried against the 258 retrieved sequences using MetaEuk v5.34c21f2 (Levy et al., 2020) to predict the protein 259 sequences in the contigs respective to the reference sequences used. Next, the predicted 260 261 sequences and the reference sequences were concatenated together and aligned using MAFFT 262 v7.475 (Katoh et al., 2013) applying the local pair method with a maximum iteration of 1000. Finally, the phylogenetic tree was constructed from the aligned sequences using FastTree 263 v2.1.10 (Price et al., 2010) with default parameters. 264

265 2.4. Statistical and ecological analyses

Alpha diversity indices such as Observed OTUs, Chao1, and Shannon were estimated from the 266 rarefied OTU table using the 'alpha' function in the package 'microbiome' in R (Lahti et al., 267 2019). Phylogenetic diversity (Faith's PD) was also calculated using the 'pd' function in the R 268 package 'picante' (Kembel et al., 2010). Boxplots showing the alpha diversity indices were 269 produced using the 'ggboxplot' function in the R package 'ggpubr' (Kassambara, 2020). Beta 270 271 diversity was also assessed by calculating for the weighted UniFrac distances using the 'GUniFrac' R package (Chen et al., 2022). Non-metric multidimensional scaling (NMDS) 272 ordination plot was then performed based on the weighted UniFrac distances using the R 273 274 package 'vegan' (Oksanen et al., 2020). Correlation between environmental variables and 275 community composition was explored using canonical correspondence analysis (CCA) also using 'vegan'. In addition to CCA, possible drivers of clustering were also investigated using 276 277 linear regression, with principal coordinates as x-axis and measured environmental parameters 278 as y-axis. Most of the statistical tests in this study were performed in the R environment v.3.6.2 279 (R Core Team, 2019).

280 2.5. Weighted co-occurrence network analysis

To further investigate the possible co-occurrence vis-à-vis interactions between taxa, network 281 282 analysis was performed using the R package 'Weighted Gene Co-expression Network Analysis (WGCNA)' (Langfelder & Horvath, 2008). An OTU matrix without the low abundance OTUs 283 (<10 counts) across all time points or 0 values for at least 11 time points to yield more robust 284 285 seasonal patterns (Bryant et al., 2016) was used for WGCNA. The resulting matrix was checked for any missing values using the 'GoodSamplesGenes' function included in the package. 286 Subsequently, the OTU matrix was used to identify the modules and construct the network 287 manually following the method of Zhang & Horvath (2005). Briefly, we first chose to set the 288 soft-thresholding power to 6 by using the 'pickSoftThreshold' function. This soft-thresholding 289 power was used to encode the connection strength (weight) between each pair of OTU, 290 resulting to an adjacency matrix. To reduce noise and spurious associations, the adjacency 291 matrix was transformed into the Topological Overlap Matrix (TOM), which reflects the relative 292 interconnectedness between OTUs. Then, average linkage hierarchical clustering using the 293 294 TOM was generated as data input to obtain a clustering dendrogram of OTUs. The clusters in the resulting dendrogram correspond to the modules, which are subgroups of highly correlated 295 OTUs with similar distribution patterns across the samples. Module identification based on the 296 dendrogram was done using the 'cutreeDynamic' function. Lastly, the modules were used to 297

calculate for eigengene values to correlate with the environmental parameters and elucidate potential drivers of the module clustering. The modules were exported to Cytoscape v.3.8.2 (Csardi & Nepusz, 2006) for visualization but only included edges with weight of ≥ 0.5 to filter out less significant associations. Key topological features of the network, such as the degree (number of edges), shortest path distribution, and betweenness centrality, were also calculated for each node using the plugin 'Network Analyzer' in Cytoscape to assess significant associations between taxa.

305

306 3. RESULTS



307 3.1. Environmental context

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Fig. 1. A) UPGMA clustering of samples based on log values of measured environmental

310 parameters. The samples were grouped into NE, SW and storm clusters; B) Barplots showing

- the relative abundance of binned OTUs at the class level present in the picoplankton fraction.
- An in-house curated reference database including entries from both SILVA v.138.1 and
- 313 Northern Reference Database v.1.0 was used as the reference for the barplot. The OTU
- dataset used was rarefied at 1,559 reads; C) Shifts in the most abundant taxa (represented by

the OTU with the highest read count) per module driven by changes in local environmentalconditions influenced by monsoons and by storms.

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Results showed that the surface waters were marked by strong seasonal variation, where 318 measured physico-chemical parameters clustered the samples into three main groups (Figure 319 1A). Groupings include January to mid-June, which were months associated with the late 320 321 northeast (NE) monsoon to intermonsoon, and late June-October associated with southwest (SW) monsoon. Notably, some samples collected in August (Aug27BML) and September 322 (Sep17BML and Sep24BML) also separated into another group outside of the clusters 323 associated with the monsoon periods. These samples coincided with episodic disturbance when 324 tropical storms were reported near the study site (PAGASA, 2019) during the collection period, 325 herein referred to as the 'storm' cluster. Significant differences on the dissimilarities based on 326 environmental conditions (p<0.05) were found using ANOSIM (Supplementary Figure S3A). 327

The shift in monsoon as well as the disturbance caused by storms were accompanied by drastic 328 changes in environmental conditions. Generally, months during NE season had lower 329 temperature, lower Chl a concentrations, and higher salinity and nitrate values compared to 330 331 months during SW season (Table 1, Supplementary Figure S2). Temperature generally increased in the surface waters from NE (27.8°C - 30.6°C) to SW monsoon (30.2°C - 32.8°C; 332 t-test, p < 0.05). Meanwhile, salinity values steadily declined from an average of 30.8 + 0.7333 during the NE monsoon to 29.2 + 0.7 during the SW monsoon (t-test, p<0.05) in the surface 334 waters. A similar trend was observed for nitrate concentrations, which decreased from an 335 average of $1.17 + 0.44 \,\mu\text{M}$ to $0.09 + 0.03 \,\mu\text{M}$ (t-test, p<0.05) during the SW monsoon. Chl a 336 concentrations fluctuated from a minima of 0.31 mg m⁻³ during the NE monsoon to a maxima 337 of 21.46 mg m⁻³ in the SW monsoon (t-test, p < 0.05). In comparison, samples associated with 338 the storm were observed to have the lowest salinity values, averaging only 19.7 + 3.97, lower 339 by 10 units when compared to both monsoon periods. The observed decrease in salinity could 340 be attributed to the higher amount of rainfall and runoff during those sampling dates (Table 2). 341 In addition, the lowest phosphate concentrations were recorded, from approximately 0.8 µM 342 for both monsoon periods to $0.36 + 0.13 \,\mu\text{M}$ during stormy conditions. However, the highest 343 dissolved silicate concentrations $(25.03 + 10.07 \mu M)$ were recorded for those same samples, 344 compared to the average silicate values during NE and non-stormy SW monsoon samples 345

346 (Table 1), respectively. Meanwhile, nitrite concentrations ($<1.0 \mu$ M) were observed to be 347 relatively low for the entire duration of sampling.

348 3.2. Community diversity and composition

We observed a clear shift in community composition both at the binned reads (relative 349 abundance; Figure 1B) and OTU levels (Figure 1C). In the samples associated with the NE 350 monsoon, the most abundant picoplanktonic taxa present (>1% of the total reads per sample) 351 were Chlorophyta, Syndiniales, Cryptophyta, Prymnesiophyceae, Picozoa, and Cercozoa. 352 Larger taxa such as ciliates and dinoflagellates were also detected at more than 1% of the total 353 reads of the picoplankton fractions but could be related to cell breakage during collection or 354 filtration (Onda et al., 2017). In contrast, larger phytoplankton taxa such as diatoms dominated 355 the communities during SW monsoon, making up 44% (July) to as much as 97% (September) 356 of the entire community, consistent with previous reports on microphytoplankton composition 357 in Bolinao (Escobar et al., 2013; Yap et al., 2004). However, the dominance of diatoms was 358 less pronounced during the dates when storms occurred. Instead, the most abundant 359 360 picoplanktonic taxa present in the storm samples were Chlorophyta and Prymnesiophyceae.



Fig. 2. A) Hierarchical clustering of OTUs based on weighted UniFrac distances.

362 Communities were only separated into two distinct clusters; B) NMDS ordination plot also

based on weighted UniFrac distances of all OTUs (stress=0.078) showing temporal shifts
influenced by nitrate (contour lines).

365

In contrast to the clustering based on environmental data (Figure 1A), clustering of beta diversity based on weighted UniFrac distances (Figure 2A) showed only two significant community clusters based on their phylogenetic composition (PERMANOVA, p<0.01). The communities during the NE and SW monsoon periods still separated into two significantly distinct clusters (r = 0.70, p<0.001, Supplementary Figure S3B), but no observed separate or distinct cluster for the storm samples. Instead, Aug27BML and Sep17BML samples grouped with the NE monsoon cluster, while Sep24BML sample grouped with the SW monsoon cluster.

373



Fig. 3. Non-phylogenetic alpha diversity indices (Observed OTUs, Chao and Shannon) and
phylogenetic diversity (Faith's PD) of the total eukaryotic communities present in the

376 picoplankton fraction. The ** symbol denotes a significant p-value<0.01, * means a

significant *p*-value<0.05, and NS means non-significant *p*-value>0.05.

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A decrease in alpha diversity of the total community was further observed from NE to SW monsoon (Figure 3). The different alpha diversity indices we used consistently demonstrated that the mean diversity of the whole community was higher during NE monsoon (mean observed OTUs = 252 ± 59 ; Chao1 = 379.95 ± 57.34 ; Shannon = 4.22 ± 0.74 ; PD = $27.06 \pm$

4.39) than during SW monsoon (mean observed OTUs = 124 + 53; Chao1 = 209.94 + 89.65; 383 Shannon = 2.66 ± 0.66 ; PD = 15.44 ± 5.91). The drastic decrease in diversity during SW 384 monsoon was due to the dominance of diatoms observed during that period. SIMPER analysis 385 (Table 3) further showed that Chlorophyta and diatoms primarily contributed to the significant 386 differences observed between monsoons. Chlorophyta dominated during NE monsoon, 387 particularly during the colder months of January and February, and steadily declined in 388 abundance until it was almost undetectable during SW monsoon when diatoms started to 389 dominate. Furthermore, we also observed a decrease in alpha diversity during the stormy days 390 391 (mean observed OTUs = 137 ± 79 ; Chao1 = 194.77 ± 109.33 ; Shannon = 3.05 ± 1.22 ; PD = 14.21 ± 7.79) compared to the NE monsoon. However, we only observed significant 392 differences between the alpha diversities of communities during the NE monsoon and storm 393 period for Chao1 (p < 0.05) and Faith's PD (p < 0.05). Meanwhile, no significant difference was 394 observed for all alpha diversity measures between communities during the SW monsoon and 395 the storm period (p < 1.0). 396

397 3.3. Potential drivers of community structuring



Fig. 4. Canonical correlation analysis (CCA) plot based on total picoplanktonic communities and measured environmental parameters where the environmental factors explained 21% (sum of CCA1 and CCA2 axes) of the variance. Forward selection with function 'ordistep' revealed that the best explanatory factor was nitrate. Sample names are color-coded according to monsoon season, where red is for samples during NE monsoon, blue is for samples during non-stormy SW monsoon, and green is for samples during stormy SW monsoon.

407 Canonical correspondence analysis (CCA) and multiple linear regression analysis further 408 provided insights on the potential drivers influencing the temporal variability of the 409 picoplanktonic communities. At the total community level, both analyses identified nitrate as 410 the strongest environmental factor driving the seasonal community patterns. However, only a 411 total of 21% can be accounted for the variance observed in the total picoplankton communities by the measured environmental parameters based on CCA Axes 1 and 2 (Figure 4). Forward selection in CCA only identified nitrate as the best explanatory variable. This was further explored using NMDS (Figure 2B) overlaid with the contour plots of nitrate, which revealed a relatively more scattered clustering of NE samples collected when nitrate was from 0.4 to 1.2 μ M. Meanwhile, the plot showed a tighter clustering of SW samples collected when nitrate concentrations were lower from <0.1 to 0.6 μ M.

In addition, while nitrate was identified to be the strongest predictor, CCA triplot also showed 418 that temperature, dissolved silicate, and salinity could also be significantly influencing the 419 variability of the communities (Figure 4). Interestingly, samples belonging to the storm cluster, 420 421 particularly Sep17BML, correlated with high dissolved silicate levels while Sep24BML sample with low salinity. We also observed that NE monsoon samples were associated with 422 high nitrate levels while samples collected from SW monsoon with high temperature based on 423 the triplot. To further support our CCA results, we then performed multiple linear regression 424 425 analysis with PC1 (49.6% of the variance observed) which showed temperature (p=0.010) significantly positively correlating with it, while nitrate (p=0.09) and DIN:DIP ratio (p=0.022)426 427 to be significantly negatively correlated. In contrast, silicate (p=0.081) and dissolved oxygen levels (p=0.062) positively correlated with PC2 (16.7% of the variance observed), but only 428 slightly significant. PC1 and PC2 were based on the PCoA of weighted UniFrac distances. 429

430 3.4. Shifts in community interactions



Fig. 5. Networks of the microbial eukaryotic communities during the Northeast (NE)

433 monsoon (Module I), Southwest (SW) monsoon (Module 2) and during storms (Module 3).

434 The colors of the nodes represent the phylum of the OTUs while the sizes represent

435 betweenness centrality.

436

Shifts in the abundant taxa driven by changes in the environment also affected the possible 437 438 trophic structure. In our resulting network (Figure 5), the 238 most abundant OTUs grouped only into three modules, namely those OTUs abundant during the NE monsoon (Module 1), 439 440 SW monsoon (Module 2), and during storms (Module 3), similar to the clustering based on environmental parameters. Module 1 was dominated by a chlorophyte taxa, which was 441 observed to peak in January (Figure 1C), and had a total of 2,754 reads, accounting for around 442 35% of the total OTU reads belonging to Module 1. We validated the phylogenetic placement 443 of the chlorophyte sequences, which revealed that the most abundant sequence in our dataset 444 with barcode "3f5ff" was more closely related to Ostreococcus Clade E (Supplementary Figure 445 446 S5). Meanwhile, Module 2 was dominated by the diatom Skeletenoma, with a total read count of 1,561, which was approximately 21% of the total OTU reads. It was observed to have highest 447 abundance during late June at the onset of SW monsoon but remained present the entire season 448

(Figure 1C). Lastly, the most abundant OTU in Module 3 was the chlorophyte Picochlorum 449 with 740 reads, accounting for approximately 32% of the total OTU reads. This taxon was 450 observed to be the most abundant during dates associated with the storm cluster (Figure 1C). 451 Spearman's rank correlation of each module (denoted as ρ) with the measured environmental 452 variables further highlighted the seasonality of the modules (Supplementary Figure S4). For 453 example, Module 1 correlated with lower temperatures ($\rho = 0.34$, p>0.10), low Chl a 454 concentrations ($\rho = 0.58$, p<0.05), higher nitrate concentrations ($\rho = 0.46$, p>0.05), and higher 455 DIN: DIP values ($\rho = 0.74$, p < 0.01), which marked the NE monsoon, although only the 456 457 correlations for Chl a and DIN: DIP were significant. Low temperature, low Chl a 458 concentrations, and higher nitrate values were environmental conditions observed during the 459 NE monsoon period.

The turnover in seasonal conditions coincided with the changes in the co-occurring taxa and 460 dominant OTUs. Based on Figure 5, Modules 1 and 3, which correspond to NE and storm 461 clusters, respectively, had lower number of edges (42 and 28) and nodes (14 and 10) compared 462 to the separate Module 2 (SW cluster) that had 43 nodes and 918 edges (Figure 5). In Module 463 1, the chlorophytes Ostreococcus and Trebouxiophyceae dominated, which also coincided with 464 their peak abundances during the late NE monsoon. Ostreococcus specifically was the most 465 connected (7 edges), mostly with dinoflagellates. Interestingly, at the start of the intermonsoon, 466 the chlorophytes decreased and were replaced by picozoa as the most abundant taxa. Picozoa 467 OTUs accounted for 56% of the total read counts during March and formed a sub-network 468 together with prymnesiophytes. In contrast, most of the abundant OTUs in Module 2 (SW 469 monsoon) were diatoms, accounting for 51% of the total reads and with 16% of all the total 470 471 edges in this module. The diatoms were connected to almost all other taxa present in Module 2. Parasitic Syndiniales were the second most abundant in this module accounting for 9.5% of 472 473 the total reads, and significantly co-occurring with their potential hosts, such as the dinoflagellates (12% of module total) and ciliates (8.8% of the module total). OTUs that were 474 475 unclassified were also abundant particularly in June, peaking at 9.2% of the module total. Notably, the unclassified OTUs only formed connections with each other, and formed a 476 477 separate sub-network that was not significantly affiliated with the diatom OTUs. Finally, Module 3 had the lowest number of OTUs with chlorophyte OTUs being the most abundant 478 taxa, making up around 46% of the module total. Specifically, the chlorophyte *Picochlorum* 479 (32% of the module total) was the most abundant taxon, which was significantly associated 480 only with other chlorophytes and a Cercozoa. Similar to Module 2, parasitic Syndiniales were 481

- 482 also abundant in this module (14% of the module total), forming significant connections only
- 483 with ciliates (16% of the module total), which could potentially be their hosts.



484 3.5. Changes in functional community profiles

Fig. 6. Relative abundance of the "Level 1" functional categories based on eukaryotic and
unclassified sequences retrieved from shotgun metagenomic analysis of the representative
samples for NE and SW monsoon. "Others" category is composed of 12 other low abundance
"Level 1" functional categories.

489

For the shotgun data, our two representative samples, namely April (Apr19BML) for late NE 490 491 to early intermonsoon and September (Sept30BML) for late SW to early intermonsoon, each had a total of 42,948,874 and 42,677,963 sequence reads, respectively. We first extracted 492 493 sequences inferred as Eukaryota or unclassified from both datasets, and only these were used for further processing to focus on picoeukaryotes. Using these filters, around 11% of the total 494 495 sequences were extracted from the April sample while ca. 24% from September. We also observed significant differences between the resulting profiles in each sample based on the 496 497 relative abundance of 'Level 1' functional categories (Figure 6). These differences were particularly visible in genes related to photosynthesis, being 5 times lower in NE monsoon 498 499 (1.28%) than in SW monsoon (6.38%). Genes related to protein metabolism along with phage-500 and prophage-related functions were also significantly lower during NE monsoon than in SW

- monsoon. In contrast, genes for the metabolism of amino acids and derivatives were slightly 501
- higher during NE monsoon (10.05% vs. 8.11%). 502
- 503

Α

95% confidence intervals Photosystem_II_CP43_protein_(PsbC)(1) ю < 1e-15 photosystem II protein D1 (PsbA)(2) < 1e-15 нон photosystem_I_P700_chlorophyll_a_apoprotein_subuni... юн < 1e-15 photosystem_I_P700_chlorophyll_a_apoprotein_subuni... нон < 1e-15 Photosystem_II_CP47_protein_(PsbB)(1) ю < 1e-15 photosystem_II_protein_D2_(PsbD)(3) 💾 ю < 1e-15 Putative_chaperon-like_protein_Ycf39_for_quinone_b... 0 < 1e-15 photosystem_l_subunit_ll_(PsaD)(1) 0 | < 1e-15 "photosystem_II_P680_chlorophyII_A_apoprotein,_put... | 0 < 1e-15 0 "Photosystem_II_protein_PsbV,_cytochrome_c550(1)" < 1e-15 0.0 13 -12 -1.0 -0.8 -0.6 -0.4 -0.2 0.0 0.2 Proportion (%) Difference between proportions (%) 🔲 30-Sep 🔲 19-Apr В 95% confidence intervals Nitrite-sensitive transcriptional repressor NsrR(1) < 1e-15 Nitrous oxide reductase maturation protein NosR(1) 4.44e-15 Ferredoxin-dependent glutamate synthase (EC 1.4.7... 4.12e-8 Cyanoalanine nitrilase (EC 3.5.5.4) # plant-induce... 8.74e-6 _ "Nitrate ABC transporter, permease protein(1)" 2 27e-5 Glutamate_synthase_[NADPH]_large_chain_(EC_1.4.1.1... 6.85e-5 Assimilatory_nitrate_reductase_large_subunit_(EC:1... 2.78e-4 Nitrogen_regulation_protein_NtrC(1) -----2.78e-4 "Putative_thiol:disulfide_oxidoreductase,_nitrite_... P юн 6.59e-4 Asparagine_synthetase_[glutamine-hydrolyzing]_(EC_.. 6.91e-4 -----Putative_glutamine_amidotransferase_YafJ(1) 占 7.01e-3 юн Glutamate_synthase_[NADPH]_large_chain_(EC_1.4.1.1... 🏪 0.011 Ammonium_transporter(1) 0.015 "Urea_ABC_transporter,_permease_protein_UrtB(2)" 🖡 0.018 Ureidoglycolate dehydrogenase (EC 1.1.1.154)(1) i HO 0.018 Nitrite_reductase_[NAD(P)H]_small_subunit_(EC_1.7.... P 0.021 I Ю-Uroporphyrinogen-III_methyltransferase_(EC_2.1.1.1... 占 ю-0.046 0.0 0.1 -0.06 -0.04 -0.02 0.00 0.02 0.04 0.06 0.08

Difference between proportions (%)

🔲 19-Apr 🛛 🔲 30-Sep

504 Fig. 7. Metagenomic profile comparisons of genes associated with A) photosynthesis and B) nitrogen metabolism created using STAMP. Positive differences between proportions imply 505 506 that Apr19BML (blue) dataset have greater abundance for the given genes, whereas negative differences show that Sept30BML (orange) dataset have higher abundance. Only the most 507 significant genes with corrected *p*-values (q-values) $< 1.0e^{-15}$ were included in the 508 photosynthesis gene profile while significant genes with corrected *p*-values (*q*-values)<0.05 509 510 were included in the nitrogen metabolism gene profile.

Proportion (%)

We then compared the abundance profiles of photosynthetic-related genes at a finer level and 511 noted that almost all genes related to both photosystems I and II were higher during SW 512 monsoon when diatoms dominated (Figure 7A). Most of the significant photosynthetic-related 513 genes detected (corrected p < 0.05) were related to the proteins that form the photosystem II 514 core complex, including CP43 (PsbC) and CP47 (PsbB) core antenna proteins as well as the 515 D1 (PsbA) and D2 (PsbD) reaction center core proteins. These genes were all detected at an 516 abundance of almost 10 times lower in NE monsoon than in SW monsoon. In particular, the 517 proportions of D1 (PsbA) and D2 (PsbD) during NE monsoon were only around 0.11% and 518 519 0.07% respectively, lower than 1.2% and 0.31% proportions in SW monsoon. Similarly, the proportions of CP43 (PsbC) and CP47 (PsbB) increased from 0.11% and 0.14% during NE to 520 0.87% and 0.78% respectively in SW samples. Furthermore, genes important to the electron 521 transport activity in photosystem I, specifically P700 (PsaA) and P700 (PsaB), were also lower 522 in abundance during NE monsoon at 0.11% and 0.31%, respectively, and increased to 1.10% 523 524 and 1.27% during SW monsoon. We further identified the taxonomic identity of the reads that were associated with the most differentially abundant photosynthetic genes (PsbC, PsbA, PsaA, 525 526 *PsbB*), and observed that the April sample was composed of around 51% chlorophytes (Figure 8A) while the September sample was composed of approximately 90% diatoms (Figure 8B), 527 528 consistent with the amplicon datasets. The phylogenetic trees using the predicted protein sequences further agree with our previous results (Supplementary Figure S6). In general, genes 529 from the April sample were associated with chlorophytes particularly Ostreococcus, while 530 531 genes from the September sample with diatoms such as *Thalassiosira*.

Another significant change observed was those related to nitrogen metabolism, consistent with 532 533 the correlations showing that nitrate was the most significant driver of community structuring (Figure 7B). Results suggest that genes associated with ammonia assimilation were the most 534 535 dominant for both seasons, although lower for NE monsoon than in SW monsoon (0.16% vs. 536 (0.20%). Majority were ammonium transporter or regulator during NE monsoon (0.09%) and were slightly more abundant than during SW monsoon (0.06%). Meanwhile, genes associated 537 with ammonia assimilation mostly were glutamate synthase during the SW monsoon (0.14%) 538 539 and were higher than NE monsoon (0.06%).



- 541 **Fig. 8.** Potential taxonomic identity based on BLAST against the nt database of the reads
- 542 from A) Apr19BML and B) Sep30BML samples, that were associated with the top four most
- 543 differentially abundant photosynthetic genes (*PsbC*, *PsbA*, *PsaA*, *PsbB*).

545 **4. DISCUSSION**

546 4.1. Shifts in environmental conditions

Clustering of the samples based on measured environmental parameters resulted in three major 547 groups, namely from January to early June, from late June to October, and some samples from 548 August to September. The first two groups are the months historically associated with NE and 549 SW monsoons, respectively (Wyrtki, 1961), while the last group, although being within the 550 SW monsoon, were the days associated with the occurrence of storms or heavy rainfall events. 551 552 Such observations are typical of tropical coastal settings, such as in Bolinao, where changes in conditions with season are mainly brought about by the reversal of the monsoon winds (Wyrtki, 553 1961; Rivera, 1997). The ensuing changes in the conditions due to monsoon-driven shifts as 554 555 well as episodic disturbances (i.e., storms) have significant influences on the local 556 environmental conditions and thus, on the productivity of the different marine ecosystems including the coastal habitats (Canini et al., 2013; Mitbavkar et al., 2015). 557

Changes in temperature were observed between the identified monsoon periods but were very 558 small when compared to temperate areas. The highest temperature was observed around July, 559 560 consistent with previous observations that water temperature slightly lags with the change in air temperature, with maximum values observed around June and July (Escobar et al., 2013; 561 Rivera, 1997). These warm conditions may result in thermal stratification, resulting in 562 shallower mixed layer depth reducing nutrient transport from deeper waters to the surface 563 (Gittings et al., 2018). Salinity was also observed to be generally higher during NE period 564 possibly due to the higher evaporation rates and minimum freshwater input (Rivera, 1997), 565 except during La Nina years (Ramos et al., 2017), although inputs from submarine groundwater 566 discharge have also been reported in the area (Cardenas et al., 2010). In contrast, heavy 567 precipitation during SW monsoon with strong freshwater input from runoff and riverine 568 569 discharge significantly dilutes the surface waters lowering the salinity (Canini et al., 2013; Ravelo et al., 2022). Such discharges from land and rivers however also bring in the needed 570 571 nutrients, causing a net influx of resources for phytoplankton growth and primary productivity (Ferrera et al., 2016; Yap et al., 2004). Here, a significant increase in Chl a was observed with 572 increased phytoplankton biomass during the SW monsoon, similar to previous reports in the 573 574 area (Ferrera et al., 2016; Yap et al., 2004).

575 Occurrence of storms characterized by strong winds (> 62 kph) and heavy rainfall (> 9 mm) 576 may also reduce the stratification and cause turnover of the water column, also replenishing 577 the nutrients in the surface waters (Gittings et al., 2018; Rivera, 1997). Interestingly, some of 578 the samples collected in August and September formed a separate group, which were

incidentally collected when gale or tropical storm warnings were up. Specifically, Aug27BML 579 sample was collected when a tropical cyclone wind signal was raised over the area due to 580 tropical storm 'Jenny' (PAGASA, 2019). Meanwhile, although a storm signal was not raised 581 during the collection day for Sep17BML sample, rains and heavy winds were still present 582 brought by the trough or extension of tropical depression 'Nimfa' (PAGASA, 2019), which 583 passed by the area a few days before. The lowest salinity values within the SW monsoon period 584 were indeed recorded during these stormy days. We also observed low nutrient concentrations 585 (nitrate, nitrite, and phosphate) but high Chl a concentrations, possibly indicating active uptake 586 587 by phytoplankton (Ferrera et al., 2016). Notably, unlike the other periods within the SW 588 monsoon, the storm cluster was observed to have higher dissolved silicate values (Figure 1A; Figure 3). Silicate generally during SW monsoon should be high due to freshwater forcing; 589 590 however, diatoms were likely to have utilized the silicate, thus, possibly explaining the low 591 silicate measurements (Escobar et al., 2013). During storms however, diatoms were the most 592 affected as characterized by a significant decline in their abundance, resulting also in lower uptake and thus, buildup of silicate in the sea surface. These fluctuating conditions caused by 593 594 monsoonal shifts, accompanied by increased sediment transport and riverine input during SW monsoon make establishing predictable patterns of microbial communities difficult (Lambert 595 596 et al., 2019). These changes in local conditions caused by the monsoon-driven shift as well as 597 the episodic disturbances could select organisms that could adapt, and thus, would have significant effects in the picoeukaryotic communities. 598

599 4.2. Effects of monsoons and storms on community patterns

In Bolinao, our results revealed two distinct picoeukaryotic communities similar to our 600 601 monsoon-based grouping, indicating the significant effect of monsoonal shifts on the structuring of microbial communities (dela Peña et al., 2021). Community and phylogenetic 602 diversities among communities during NE monsoon were higher compared to the ones during 603 SW monsoon, when diatoms started to dominate. No distinct storm cluster was observed, 604 instead most communities during stormy period grouped together with NE communities, likely 605 because disturbance caused by the storm events disrupted the dominance of diatoms, causing 606 an increase in community diversity. For example, during Aug19BML, a week before a typhoon 607 was recorded in the area, diatoms accounted for approximately 80.2% of the total community 608 reads. A week after, on Aug27BML, during typhoon 'Podul' (local name 'Jenny' in the 609 Philippines), the diatoms only accounted for around 1.2% of the total community reads. Instead 610

of diatoms, we observed that there was an increase in the abundance of chlorophytes 611 specifically, *Picochlorum*, which are organisms known to thrive in freshwater brought in by 612 the storms (Foflonker et al., 2016). Further, abundance of mixotrophic or wholly heterotrophic 613 species such as Cercozoa and Cryptophyta increased, perhaps due to the decrease of light 614 intensity related to the increased turbulence and terrigenous inputs that accompanied the 615 storms. The increase in community diversity during short-term disturbance is consistent with 616 the intermediate disturbance hypothesis (Fox, 1979), which postulates that increase in 617 disturbance results to an increase in diversity up to a certain level of strength until diversity 618 619 decreases or stabilizes. In addition, we also observed an increase in the phylogenetic diversity 620 from pre-storm (Aug19BML: PD = 19.74) to during a typhoon (Aug27BML: PD = 23.19), rising to above the average PD observed during the SW monsoon. This may indicate that 621 622 additional niches were created, and that more phylogenetically distinct species may co-occur during storms than during the SW monsoon (Galand et al., 2016). 623

Multivariate approach further revealed that nutrients, specifically nitrate, which was observed 624 to have decreased from NE to SW monsoon, was the strongest predictor of shifts in 625 picoeukaryotic community structuring. This is consistent with previous microphytoplankton 626 studies, and was likely because our site was at the opening of the Guiguiwanen channel, where 627 most of the mariculture activities take place making the area N-limited (Ferrera et al., 2016). 628 629 This shift in the community was also accompanied by the shift in the general taxonomic composition especially in the most abundant phytoplankton groups. This is consistent with 630 631 previous reports, which observed opposite patterns between small and large-celled taxa in coastal environments (Collado-Fabbri et al., 2011). This shift in the abundant taxa to the larger 632 633 diatoms during SW monsoon also resulted in lower diversity (Shannon), rareness (Chao1), and richness (no. of observed OTUs) compared to NE monsoon. The detection of larger taxa such 634 635 as diatoms even though our study was limited to pico-fraction could be due to cell breakage during filtration (dela Peña et al., 2021; Onda et al., 2017) or some pennate diatoms maybe 636 637 passing through the filter membranes such as Pseudo-nitzschia which was relatively abundant in our study (Vaulot et al., 2008). Nevertheless, such abundance of reads could still indicate a 638 639 sudden and significant shift in the cells being filtered when the samples were taken. Generally, we observed that communities associated with NE monsoon experienced relatively colder and 640 more saline environmental conditions relative to those associated with SW monsoon. 641 Productivity (using Chl a as the proxy) was also lower during the NE monsoon, despite 642 recording higher mean nitrate concentration compared to SW monsoon. In contrast, during SW, 643

the lower nitrate concentration despite potential inputs from runoff and riverine/groundwater 644 discharge (Ferrera et al., 2016) may be due to active uptake by opportunistic chain-forming 645 diatoms such as *Skeletonema* and *Thalassiosira*, which overwhelmingly dominated during this 646 period, accompanied by a drastic increase in Chl a. Aside from nitrate, turbulence, here 647 represented by strong winds and heavy rainfall events during tropical storms, was not measured 648 directly but may also be a significant driver of community structuring since previous studies 649 have reported that turbulence could select for particular species and regulate phytoplankton 650 community structure in aquatic environments (Trigueros & Orive, 2001; Zhou et al., 2016). 651 652 Margalef's Mandala hypothesis predicted that in high nutrient and high turbulent conditions, 653 r-strategists such as diatoms will be more favored over K-strategists such as dinoflagellates 654 (Margalef, 1978). In the Philippines, based on rainfall and wind speed data recorded in 2019 from a weather station approximately 90 km east of our study site (Table 2), rainfall was higher 655 during SW monsoon. Thus, it can be deduced that disturbance was more frequent during this 656 657 period. Moreover, there were 21 typhoons in 2019, with 13 of them making landfall and passed through our study site during the same period (PAGASA, 2019). The higher frequency of 658 659 typhoons and greater disturbance would have caused more frequent high turbulent mixing conditions, which have been reported to favor diatoms possibly due to their high adaptability 660 661 to fluctuating light conditions and to reduce grazing pressure during turbulent conditions 662 (Margalef, 1978; Zhou et al., 2016). However, we were not able to uncover how much force was optimal for diatoms to continue to thrive in this area as studies have also noted that 663 excessive turbulent force might have negative effects such as mechanical damage to 664 phytoplankton (Fraisse et al., 2015; Thomas & Gibson, 1990). In addition, as previously noted, 665 turbulence associated with storm events may likely disrupt proliferation of diatoms due to 666 decrease of light availability. 667

668 Smaller phytoplankton taxa have been suggested to be more adapted to low-nutrients conditions due to their higher volume-to-surface area ratio (Li et al., 2009). Aside from this 669 670 however, genes related to nutrient assimilation, particularly nitrogen may also help. Indeed, changes in taxonomic profiles were accompanied by changes in the abundant functional genes 671 672 especially those related to nutrient assimilation and photosynthesis. Our shotgun sequencing data for example revealed that ammonium transporter (AMT) genes were abundant for both 673 seasons, but slightly higher during NE monsoon. It would appear that uptake of ammonium is 674 prioritized over other forms of inorganic nitrogen. This is likely because ammonium is less 675 costly to assimilate since it can be directly used for amino acid synthesis via the glutamate 676

synthetase/glutamine oxoglutarate aminotransferase (GS/GOGAT) cycle in photosynthetic 677 eukaryotes (Hopes & Mock, 2015; McDonald et al., 2010; Sanz-Luque et al., 2015). AMT 678 genes have also been reported for chlorophytes, prymnesiophytes, and diatoms (McDonald et 679 al., 2010; Sanz-Luque et al., 2015), which were the major taxa identified to comprise our 680 representative NE monsoon sample (Figure 8A). In contrast, genes for both NADPH-681 682 dependent and ferredoxin-dependent glutamate synthase increased during SW monsoon. The presence of both functional genes gives a competitive advantage in assimilating ammonia since 683 the former allows synthesis even during conditions of low irradiance while the latter performs 684 685 during conditions of high or continuous irradiance (Suzuki & Rothstein, 1997). The increased 686 presence of these genes highlighted the importance of increased capacity for ammonia 687 assimilation during this period. We also observed that genes related to transport of nitrate and 688 genes related to reduction of nitrite decreased in SW monsoon. The lower abundance of genes for the transport and reduction of nitrogen coupled with increase for ammonia assimilation 689 690 concurs with previous culture studies on the response of the diatoms Thalassiosira pseudonana and Skeletonema costatum to nitrogen starvation (Hockin et al., 2012; Takabayashi et al., 691 692 2005). The increase in biomass of photosynthetic taxa, especially diatoms, could also explain the increase of genes responsible for photosynthesis from NE to SW monsoon. This is 693 694 consistent with previous studies where increase in photosynthetic taxa likely resulted in higher 695 relative abundance of photosynthesis-related genes (Warden et al., 2016; Yergeau et al., 2017).

CCA however only accounted for 21% of the variance observed for axes 1 and 2 (Figure 4), 696 697 indicating that other factors not measured may be contributing to the observed community structures. These include current speed and tides, total suspended solids (TSS), and light 698 699 availability (Bryant et al., 2016; Canini et al., 2013; Yap et al., 2004), noting that the site was 700 also near a mariculture area (Albelda et al., 2019; dela Peña et al., 2021). Yap et al. (2001) for example reported that TSS was higher during NE monsoon due to the lower current speed 701 during this period, resulting in longer residence time of suspended solids in the water column. 702 703 The high TSS might have caused lower light availability during the NE monsoon (Canini et 704 al., 2013; Yap et al., 2004). During these low-light conditions, picoeukaryotes with smaller 705 biomass compete better than diatoms which might explain the abundance of Chlorophyta 706 during NE monsoon. These environmental trends were similar to those reported in previous studies in this site (Escobar et al., 2013; Ferrera et al., 2016). Overall, our results demonstrated 707 that community composition and structure were significantly distinct for each monsoon season 708 709 indicating the high turnover of communities caused by the accompanying conditions in each monsoon period. Our results further showed that disturbances such as storm events may disrupt
dominance of a species, thus, increasing community diversity.

712 4.3. Co-occurrence and implications

Microbial network analysis is an explorative technique used in microbiome studies to predict 713 significant biotic interactions as well as to explore the contribution of specific taxa to global 714 biogeochemical processes (Faust, 2021; Jones et al., 2018; Lima-Mendez et al., 2015). Using 715 716 this approach, our resulting network (Figure 5) had a low average clustering coefficient (CI) of 0.824, which implies that our system is highly interconnected (Jones et al., 2018). In addition, 717 its average shortest path length (AL) was 1.26, further indicating that the response of the 718 microbial picoeukaryotes in this ecosystem to disturbance is fast (Faust & Raes, 2012). These 719 network properties are typical for microbial communities and are similar to those reported in 720 freshwater lakes (Debroas et al., 2017; Jones et al., 2018) and in marine open ocean systems 721 722 (Berdjeb et al., 2018; Lima-Mendez et al., 2015), where seasonal turnover seems to be 723 significant or environmental heterogeneity is high.

724 We filtered our network to only show strong co-occurrence patterns (correlation threshold = >0.5) between taxa, which could be interpreted as a possible direct relationship (*e.g.*, grazing) 725 726 or as taxa potentially preferring similar environmental conditions (Faust & Raes, 2012; Lima-727 Mendez et al., 2015). Our results were also consistent with our monsoon-based clustering, implying that although community composition did not shift drastically during storms, 728 dominance of certain taxa appeared to have been significantly altered. Further, similar patterns 729 730 were also reported in other tropical coastal sites, where stratified conditions due to increasing temperature and lower nutrients during the intermonsoon likely favored the growth of the 731 mixotrophic prymnesiophytes and heterotrophic picozoa (dela Peña et al., 2021). We observed 732 733 that the NE monsoon-associated module involved two dominant chlorophytes, identified as 734 Ostreococcus belonging to Clade E, and Trebouxiophyceae. Interestingly, Ostreococcus was more connected than Trebouxiophyceae, co-occurring with mixotrophic dinoflagellates and 735 heterotrophic eukaryotes, despite both taxa accounting for similar percentage of module total. 736 Although reports of grazing activity on Ostreococcus have been very limited, larger 737 microplankton taxa such as dinoflagellates have been known to feed on chlorophytes (Jeong et 738 al., 2010). It also had a higher betweenness centrality (0.468), which suggests that 739 Ostreococcus had higher control over the interactions of other nodes in the module than 740 741 Trebouxiophyceae. This observation may hint of a food preference for Ostreococcus in both

microplankton and nanoplankton predators. This potential predator-prey relationship however
needs further exploration to determine its validity and explore the value of *Ostreococcus* as the
base of the tropical trophic food web.

745 During SW monsoon, the dominant taxa in terms of abundance shifted from chlorophytes to diatoms. However, despite their abundance, we observed that diatom OTUs were not the most 746 747 connected. This is consistent with observations for co-occurrence networks using the global data of Tara Oceans where the most abundant OTUs did not have the most connections 748 749 (Vincent & Bowler, 2020). We also observed that in contrast to the NE-associated module, the SW-associated module was highly interconnected, suggesting that the interactions were not 750 751 species-specific and possibly that the predators have no problem shifting from one prey to another (Käse et al., 2021; Vincent & Bowler, 2020). The flexibility of the possible predators 752 753 and parasites might indicate higher stability and adaptability of the trophic food web system 754 during SW monsoon than during NE monsoon (Käse et al., 2021). Interestingly, we also 755 observed that, instead of diatoms, the parasitic Syndiniales, which are mostly obligate marine parasites (Guillou et al., 2008), had the highest number of connections for this module similar 756 757 to oceanic interactome (Lima-Mendez et al., 2015), demonstrating the ubiquity of these parasitic groups. The abundance of Syndiniales associations in our network further highlights 758 the possible significant role of parasitism in this coastal ecosystem. Previous studies have 759 760 already suggested the significant ecological impacts of parasites like Syndiniales on the population dynamics of the phytoplankton by promoting mortality of their hosts (Chambouvet 761 762 et al., 2008). However, host-parasite associations are difficult to determine just based on the network alone since the interactions during this period within and between genera can also 763 764 become less species-specific (Vincent & Bowler, 2020). Most of the Syndiniales in this module were identified as Syndiniales Group II, with two OTUs further assigned to the genus 765 Amoebophrya. It is a genus classified as a "species complex" composed of strains with varying 766 degrees of host specificity, ranging from extremely species specific to nonspecific 767 768 (Chambouvet et al., 2008; Chen et al., 2018; Coats & Park, 2002; Kim, 2006). Recently, a polyphasic approach combining genetic and phenotypic characters delineated eight individual 769 cryptic Amoebophrya species, with different host ranges (Cai et al., 2020). It is possible that 770 771 the Syndiniales present are generalists and can infect various hosts, but it is also highly likely that the Syndiniales were in the free-living dinospore stage and acted as food for other protists 772 773 taxa such as ciliates. Further studies should be done to investigate and establish the role of 774 parasitism in shaping phytoplankton community dynamics in this ecosystem.

In contrast to the non-stormy SW period, the dominating taxa during stormy events shifted 775 back to the smaller chlorophyta from diatoms, particularly being dominated by Picochlorum. 776 It is a fast-growing organism that can adapt to extreme environmental fluctuations particularly 777 salinity (Foflonker et al., 2016; Goswami et al., 2022), which was very evident during the 778 stormy days in Bolinao. Notably, *Picochlorum* did not form any connections with other taxa 779 780 and instead formed a sub-network consisting almost exclusively of OTUs belonging to the same taxa. This demonstrated that these chlorophyte OTUs occurred sporadically in high 781 abundance, possibly brought about by the large volume of freshwater discharge during the 782 783 storm (Käse et al., 2021). The main source of freshwater discharge in our study site are rivers and submarine groundwater around the area (Rivera, 1997). The other sub-network in this 784 module consisted mostly of ciliates co-occurring with parasitic Syndiniales classified as Group 785 786 I. One of the Syndiniales OTU was further assigned to the genus Euduboscquella (Coats et al., 2012), known as a parasite infecting mainly tintinnid ciliates (Horiguchi, 2015), thus 787 highlighting a potentially significant parasitic interaction during this period. 788

Overall, our results revealed how specific taxa respond to changes in conditions brought about by changes in monsoons. Our networks demonstrated a potentially more stable food web system during the SW monsoon than during the NE monsoon due to the presence of more adaptable predators and parasites, however, fluctuations caused by storms appeared to select for dominant taxa. The changes we have observed here to the base of the food web system may have impacts on higher trophic levels, and thus, should be considered in future studies.

795

796 **5. CONCLUSIONS**

Our combination of community structuring and network analysis indicates a high turnover of 797 798 picoeukaryotic communities associated with monsoon changes. Network analysis further revealed that communities during the SW monsoon might have higher stability and adaptability 799 800 to natural environmental fluctuations than during the NE monsoon, however, drastic environmental changes due to storm events could select for dominant taxa. This resulted in 801 higher diversity of communities when storms occurred compared to non-stormy SW monsoon 802 conditions. Stability of the SW communities might be due to the presence of less selective 803 804 predators and parasites as well as higher variability of prey options. In addition, functional genes, particularly genes related to nitrogen assimilation, might have also enabled them to 805

adapt to changing nutrient conditions. This study provided insights on how picoeukaryotic
communities and their key taxa respond to changes caused by monsoonal shifts and by storms
in a tropical coastal ecosystem. The implications in our study should help in further
understanding picoeukaryotic community adaptation and resilience in a tropical coastal
ecosystem.

811

812 TABLES

Table 1. Sampling dates in 2019 (Date) and measured environmental parameters of the
samples used in this study. *NS* denotes no sample, *nr* denotes no replicate, and *nd* denotes
that the actual concentration is below the detected limit of the instrument. The mean values <u>+</u>
standard deviation of two replicates were reported for nitrate, nitrite, phosphate and silicate.
DIN: DIP ratio was calculated using the sum of the average nitrate and nitrite values divided
by the average phosphate concentration.

Date	T (°C)	Salinity	Chl <i>a</i> (mg m ⁻³⁾	NO3 ⁻ (μΜ)	NO2 ⁻ (μΜ)	PO4 ³⁻ (μM)	DIN: DIP	SiO4 ⁴⁻ (μΜ)	DO (mg L ⁻¹)
31-Jan-19	27.8	32.4	NS	2.67 <u>+</u> 0.03	0.29 <u>+</u> 0.01	1.55 <u>+</u> 0.05	1.91	6.36 <u>+</u> 0.12	6.0
17-Feb-19	27.9	33.3	7.24	2.59 <u>+</u> 0.22	0.11 <u>+</u> 0.00	1.63 <u>+</u> 0.06	1.66	6.69 <u>+</u> 0.02	4.7
9-Mar-19	30.3	29.3	3.16	0.66 <u>+</u> 0.05	nd	0.59 <u>+</u> 0.09	1.06	2.73 <u>+</u> 0.04	5.7
5-Apr-19	29.7	31.7	1.51	nd	0.02 <u>+</u> 0.14	0.47 <u>+</u> 0.12	0.05	5.57 <u>+</u> 0.52	5.6

19-Apr-19	30.5	29.1	1.41	1.79 <u>+</u> 0.41	nd	0.33 <u>+</u> 0.01	5.48	8.72 <u>+</u> 2.03	6.2
15-Jun-19	30.6	31.2	0.31	0.25 <u>+</u> 0.27	0.07 <u>+</u> 0.06	0.45 <u>+</u> 0.27	0.71	5.45 <u>+</u> 0.07	3.6
28-Jun-19	31.3	30.9	8.41	0.14 <u>+</u> 0.14	nd	1.11 <u>+</u> 0.04	0.12	4.81 <u>+</u> 0.22	3.7
5-Jul-19	31.2	30.8	20.5	0.19 <u>+</u> 0.19	nd	0.16 <u>+</u> 0.02	1.18	0.59 <u>+</u> 0.00	6.5
19-Jul-19	34.3	28.3	1.60	0.24 <u>+</u> 0.13	0.13 <u>+</u> 0.08	0.24 <u>+</u> 0.14	1.54	6.15 <u>+</u> 0.28	6.2
19-Aug-19	30.2	26.6	7.32	0.04 <u>+</u> 0.01	nd	0.38 <u>+</u> 0.02	0.10	3.09 <u>+</u> 0.08	5.1
27-Aug-19	29.8	27.2	NS	0.08 ± 0.08	nd	0.11 <u>+</u> 0.02	0.75	15.74 <u>+</u> 0.02	5.3
17-Sep-19	30.3	18.2	10.67	0.11 <u>+</u> 0.00	nd	0.49 <u>+</u> 0.01	0.22	45.16 <u>+</u> 5.34	6.0
24-Sep-19	30.4	13.7	32.71	0.15 <u>+</u> 0.08	0.05 <u>+</u> 0.20	0.49 <u>+</u> 0.10	0.40	14.20 <u>+</u> 0.95	6.6
30-Sep-19	30.6	27.5	17.18	0.11 <u>+</u> 0.05	nd	0.23 <u>+</u> 0.02	0.47	4.84 <u>+</u> 0.95	6.4

7-Oct-19	32.8	29.0	NS	0.05 (nr)	0.02 (nr)	0.83 <u>+</u> 0.02	0.08	3.06 <u>+</u> 0.34	5.4
28-Oct-19	32.3	30.1	21.5	0.03 <u>+</u> 0.02	nd	1.92 <u>+</u> 0.07	0.02	1.93 <u>+</u> 0.10	2.1

Table 2. Rainfall (mm) and wind speed (m/s) data of each sampling date were recorded from

a station ~90 km east of Bolinao (Dagupan monitoring site of the Philippine Atmospheric,

823 Geophysical and Astronomical Services Administration-PAGASA).

SAMPLING DATE	RAINFALL, mm	WIND SPEED, m/s
January 31, 2019	nil	2
February 17, 2019	nil	3
March 9, 2019	nil	3
April 5, 2019	nil	3
April 19, 2019	0.1	3
June 15, 2019	4	3
June 28, 2019	54.4	2
July 5, 2019	nil	2
July 19, 2019	33	2
August 19, 2019	nil	2
August 27, 2019	1.6	3
September 17, 2019	3.4	3

September 24, 2019	nil	3
September 30, 2019	nil	3
October 7, 2019	nil	3
October 28, 2019	nil	2

- **Table 3.** SIMPER analysis showing the most influential OTUs and their respective percent
- 827 contribution to the pairwise dissimilarity between NE, SW and Storm samples.

Taxon	Most resolved taxonomy	Percent contribution to pairwise dissimilarity			Average	abundance	e (reads)
		NE/ Storm (27.46%)	NE/ SW (40.09%)	SW/ Storm (53.81%)	NE	SW	Storm
Diatomea	Chaetoceros	0.55	10.55	17.32	53	1817	114
Diatomea	Skeletonema	nil	7.72	13.14	7	1298	6
Diatomea	Skeletonema	2.69	8.33	11.15	1	1395	299
Diatomea	Thalassiosira	9.05	4.20	3.05	8	711	1011
Chlorophyta	Ostreococcus	5.10	3.27	nil	588	41	23
Chlorophyta	Picochlorum	2.70	nil	3.34	37	8	336
Chlorophyta	Micromonas	3.22	2.61	0.81	439	2	82
Diatomea	Rhizosolenia	nil	1.44	2.55	14	255	4
Dinoflagellata	Dinophyceae	2.47	1.64	nil	274	0	0
Diatomea	Leptocylindrus	1.68	0.33	2.45	186	241	0



830 Supplementary Fig. S1. Map of Bolinao, Pangasinan, Philippines. Red dot represents the

831 sampling station. Arrows represent wind direction during each monsoon season.









their trends over time.

Α





В







Module-trait relationships

Supplementary Fig. S4. Correlation of each module eigengene to every environmental
variable. The y-axis corresponds to the modules while the x-axis corresponds to the
environmental variables. The values are the Spearman's rank correlation coefficient (ρ) while
the values in the parentheses indicate the respective *p*-values. The colors in the matrix
indicate direction and magnitude of each coefficient ρ as given by the scale.





Supplementary Fig. S5. Maximum likelihood analysis of chlorophyte sequences retrieved 849 from our study (shown in boldface) as well as reference sequences from Viprey et al., 2008 850 and Monier et al., 2016 using MEGA v10.2.6. The most abundant Ostreococcus sequence in 851 852 a similar tropical coastal site reported by dela Peña et al., 2021 (shown as denovo1) was also included in the analysis. 853







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Supplementary Fig. S6. Phylogenetic trees constructed from the predicted protein sequences
[(A) *psbC*, (B) *psbA*, (C) *psaA*, (D) *psaB*] as well as reference sequences.

862

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