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1 Trajectory, fate, and magnitude of continental microplastic loads to the
2 inner shelf: A case study of the world's largest coastal shallow lagoon.

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15 Key Words: Coastal Plume; South Atlantic; Numerical Modelling; Plastic Fiber
16 Transport; Accumulation Hotspots.

18 Abstract

19 The Patos Lagoon estuary is a highly significant ecosystem where freshwater from a
20 vast and densely populated area continuously flows into the Atlantic Ocean by coastal
21 plumes, exporting not only freshwater but also sediment, nutrients, plastics, and other
22 contaminants. In this work, numerical modeling tools together with field data were used
23 to assess for the first time the capacity of the coastal plume to export microplastics (MPs)
24 to the inner shelf under different hydrodynamic conditions. Two field surveys were
25 conducted during plume events to quantify MP concentrations and validate the model
26 approach. A bottom-up approach was employed to estimate the potential MP export from
27 the estuary's domain to the Atlantic Ocean. MP concentration in surface plume waters
28 ranged from 0.20 items m⁻³ to 1.37 items m⁻³, confirmed by FTIR as synthetic polymers
29 in a 90%, being Polypropylene (PP) and Polyethylene (PE) the most abundant in a 73%.
30 The accumulation pattern was observed on the plume's frontal system, consistent with
31 simulation results. The estimated average MP potential export rate attained 9.0 million
32 items day⁻¹ during moderate plume events and 47.5 million items day⁻¹ during high

33 discharge plume events. Strong discharge events, coupled with intense northeast winds,
34 facilitated rapid southwestward export of MPs. Conversely, moderate to weak discharge
35 events retained MPs closer to the estuary's mouth, enabling either longer trajectories or
36 earlier deposition. Significant MP accumulation hotspots were identified in the gyre
37 between the jetties and Cassino beach, as well as in the saline front within the plume
38 boundaries. These accumulation zones may function as reservoirs for MP particles,
39 potentially posing threats to local ecosystems. Understanding these dynamics is crucial
40 for ongoing monitoring efforts to assess potential harmful interactions over time.

41

42 **1. Introduction**

43 Understanding the dispersion of plastic waste from the continent to the ocean,
44 considering the physicochemical environmental variables, is one of the foremost current
45 scientific challenges (Alfonso et al., 2021). Estimating MP fluxes has been one of the
46 most pursued but still unattained objectives in this research field (Jambeck et al., 2015).
47 It is estimated that inadequate management of plastic waste in continental areas
48 contributes from 4 to 12×10^6 metric tons yr^{-1} (MT) of plastics to the ocean, varying as a
49 function of the development index of each country (Alfonso et al., 2021). Brazil, ranking
50 fourth globally in plastic waste generation, produces 11.3×10^6 MT yr^{-1} (equivalent to 1
51 kg per person per week), with only 1.28% (145×10^3 MT) effectively recycled and
52 reintroduced into the production chain (Zamora et al., 2020). Mismanaged plastic waste
53 ends up in landfills and open dumps, from where it can be primarily transported to larger
54 water bodies through freshwater discharge and wind (Lozoya et al., 2015; Alencar et al.,
55 2023), with rivers contributing to 80% of the total marine waste (Meijer et al., 2021). This
56 implies that the emission probability of a country depends on its coastal extension and
57 precipitation regime, with Brazil being responsible for around $37,799$ MT yr^{-1} of plastic
58 waste leakage into the ocean (Meijer et al., 2021).

59 Patos Lagoon, the world's largest coastal lagoon located in the Southwestern Atlantic
60 Ocean (SWAO), is considered the sixth hotspot of plastic leakage to the ocean in South
61 America, only after La Plata River, Guanabara Bay, Amazon, São Francisco, and
62 Tocantins Rivers (Alencar et al., 2023). As part of the Southeast South America (SESA)
63 region, Patos Lagoon is subject to the impact of El Niño-Southern Oscillation (ENSO)
64 (Bjerknes, 1969), increasing precipitation and affecting sea level, wind, and wave patterns

65 (Barreiro, 2010; Schossler et al., 2018; Bitencourt et al., 2020b; Rehbein & Ambrizzi,
66 2022). Previous studies have demonstrated a positive impact of ENSO on continental
67 discharge in the Northern part of the lagoon associated with an increase in pressure
68 gradient between the estuary and the ocean, thus favoring plume formation (Távora et al.,
69 2019). However, in the current challenging scenario of global marine plastic pollution,
70 the impacts of different intensities of ENSO on plastics export and distribution still remain
71 surprisingly unexplored (Morishige et al., 2007; Doyle and Gilfillan, 2016). Earlier
72 investigations have demonstrated that extreme events and rainfall seasons over river
73 basins are related to stronger MP transport and mobilization (Bertoldi et al., 2021;
74 Emmerik et al., 2023a, b; Lima et al., 2023). In this context, continental discharges into
75 the SWAO during such extreme events account for 7% of the Patos Lagoon's total
76 volume, resulting in plumes with areas of up to 1,500 km² extending up to 50 km off the
77 estuary mouth (Fernandes et al., 2002) under a microtidal regime and dominance of NE-
78 SW winds (22% and 12%, respectively), where river discharge displays not only seasonal
79 and interannual variability (Moller et al., 2001, Távora et al., 2019; Bitencourt et al.,
80 2020a) but also interdecadal variability (Bortolin et al., 2022). Therefore, a combination
81 of local and non-local factors governs the dynamics of the coastal plume (Marques et al.,
82 2010b; Jung et al., 2020).

83 While the study of plastic contamination in aquatic environments has received
84 numerous global scientific contributions (e.g., Abalansa et al., 2020; Alfonso et al., 2021,
85 and references therein), comprehensive and multidimensional studies in the SWAO are
86 limited (Rodríguez et al., 2020; Pazos et al., 2021; Pinheiro et al., 2021; Lacerda et al.,
87 2022; Schicchi et al., 2024). According to dos Santos (2023), Patos Lagoon received
88 plastic waste inputs ranging from 21,670 MT to 107,190 MT yr⁻¹ (equivalent to 6.54 to
89 32.35 g day⁻¹ per person) between 2010 and 2017, composed of polyethylene (PE),
90 polypropylene (PP), and polyvinyl chloride (PVC) as the main polymers. Alves et al.
91 (2022) identified MPs (mainly fibers) as chronological markers in sediment cores within
92 the Patos Lagoon estuary and Mirim Lagoon (**Fig. 1**), with the first appearance dating
93 back to 1970 CE. This finding highlights the impact of the Anthropocene and industrial
94 activities on the local sedimentary history. Recent studies have also identified MPs in
95 oceanic waters in South Brazil, where the highest concentrations of MPs in surface water
96 at two Rio Grande do Sul stations, near the coastal plume influence area (Lacerda et al.,
97 2022). In this context, the Patos Lagoon is potentially crucial as both a source and sink of

98 anthropogenic waste, and ulterior export to the coastal region, make the estuary a
99 "hotspot" of marine pollution (Pazos et al., 2021).

100 Continental plastic loads, together with direct plastic inputs from fisheries, offshore
101 industries, and other maritime activities (Boucher and Friot, 2017; Boucher et al., 2020),
102 undergo various physical, chemical, and biological degradation processes driven by
103 external environmental factors, such as mechanical action from wind and waves, solar
104 radiation, oxygen availability and interaction with organisms. These processes lead to the
105 formation of smaller size categories and dynamic changes in particle properties over time
106 (Verma et al., 2016; Chamas et al., 2020). Consequently, estuaries are key compartments
107 for both dispersion and fragmentation of macroplastics (>25 mm) into mesoplastics (5
108 mm – 25 mm) and MPs (1 μm - 5 mm), experiencing changes in dimensions and densities
109 and altering their buoyancy properties (Jalón-Rojas et al., 2019). These characteristics
110 distinguish them from fine sediment particles, thus justifying the need for specific
111 analytical tools to understand their behavior and transport trends in different marine
112 compartments (Jalón-Rojas et al., 2019).

113 Numerical modeling has proven to be an essential tool for studying the behavior and
114 dispersion of MPs in marine environments, because of the need for specific approaches
115 due to the unique characteristics of MPs compared to suspended sediment particles
116 (Critchell and Lambrechts, 2016; Jalón-Rojas et al., 2019). Current numerical modeling
117 tools, complemented with field validation, offer a low-cost methodology with high spatial
118 and temporal resolution and integration of both experimental and field data to achieve
119 high model accuracy (Hardesty et al., 2017). Hybrid models combining hydrodynamics
120 with Lagrangian process-based approaches and statistical data have been widely used to
121 predict areas of high MP accumulation (Uzun et al., 2022). In this regard, the TrackMPD
122 model, developed by Jalón-Rojas et al. (2019) as an extended Lagrangian-based open-
123 source tool, integrates hydrodynamic models and real-world process equations for
124 tracking plastic particles in natural environments (Jalón-Rojas et al., 2019). TrackMPD
125 applications have been worldwide demonstrated in recent years in various coastal and
126 estuarine environments (Cheng et al., 2020; Baudena et al., 2022; 2023; Liao et al., 2023;
127 Kim et al., 2023; Schicchi et al., 2023).

128 In this sense, this approach was performed with the objective of understanding the
129 role of coastal plumes in the dispersion dynamics of MPs in inner shelf ecosystems, to
130 test the hypothesis that different discharge intensities would determine different patterns

131 in the magnitude of export, distribution and fate of MPs to the SWAO. The transport of
132 MPs was assessed for weak (2014-2015) and very strong (2015-2016) El Niño intensities
133 to identify different "hotspots" of MP accumulation in the export region as a function of
134 plume intensity. Two field sampling surveys validated the model and provided the basis
135 for estimating MP export flux, thus elucidating the coastal plume role in the dispersion of
136 continental MP export to the SWAO.

137 2. Methods

138 2.1. Study area

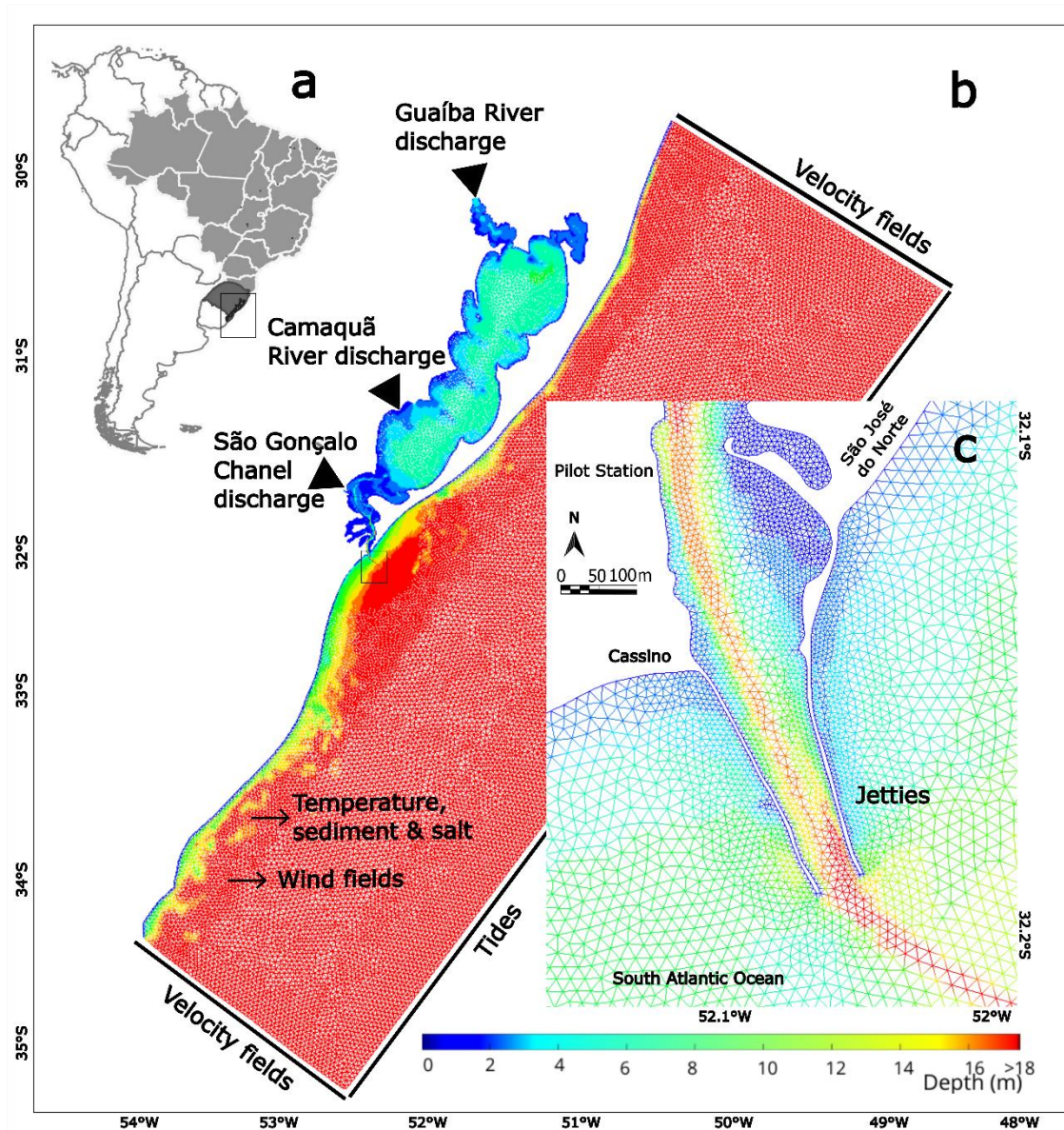
139 Patos Lagoon (30-32°S; 50-52°W, **Fig. 1**) is the world's largest choked coastal lagoon
140 (Kjerfve, 1986), with an area of 10,360 km², a maximum length of 240 km, and a mean
141 depth of 5 m. Draining a watershed of approximately 200,000 km², it is connected to the
142 SWAO through a long and narrow channel (700 m wide at the mouth) delimited by a pair
143 of 4 km long jetties (**Fig.1c**, Franzen et al., 2022). The channel's dredged portion,
144 designed for navigation purposes, maintains depths varying from 16 to 20 m. The lagoon
145 receives continental inputs from very turbid tributaries, i.e., Guaíba and Camaquã River,
146 and São Gonçalo Channel (connection with Mirim Lagoon, **Fig. 1**). These tributaries
147 together yield an average annual discharge of 2,400 m³ s⁻¹, carrying a significant
148 contribution of suspended sediments of approximately 5.1x10⁶ t yr⁻¹ (Jung et al., 2020).
149 Guaíba River is the main tributary, accounting for approximately half of the fluvial
150 discharge into the lagoon (average flow of 1,253 m³ s⁻¹) and carrying mean annual
151 discharges of 2.99 × 10⁶ ton yr⁻¹ of suspended sediments from a sub-drainage basin of
152 about 84,763 km². Camaquã River contributes 0.24 × 10⁶ ton yr⁻¹, and drains a basin of
153 about 21,657 km², while São Gonçalo Channel contributes 1.88 × 10⁶ ton yr⁻¹ (Jung et
154 al., 2020) from a basin area of 28.499 km². Extreme discharge events from the tributaries
155 over the north region of the system are related to the inter-annual variability driven by
156 ENSO oscillations (Távora et al., 2020, Bortolin et al, 2022).

157 The residence time of particles within the system was calculated at approximately 1.5
158 yr (for an average discharge of 1,000 m³ s⁻¹, Moller et al., 1996), and was estimated at 68
159 days for maximum discharges (10,000 m³ s⁻¹, Fernandes et al., 2002), showing a decrease
160 inversely related to river discharge. The concentration of suspended solids decreases from
161 the Northern end of the lagoon (38 g m⁻³) to the Southern end in the estuarine area (10 g
162 m⁻³), which highlights the retention and filtration capacity of the lagoon. Mud depocenters

163 form in the central silty zones where current velocities are smaller than 0.1 m s^{-1} and depth
164 is higher than 5 m (Bortolin et al., 2022). Littoral sandy zones are dominated by sediment
165 resuspension when current velocities are higher than 0.2 m s^{-1} and depth is lower than 5
166 m.

167 The wind regime is mainly determined by NE wind (22%) interrupted by the passage
168 of frontal systems from SW (12%) in time intervals of 3 – 16 days (Moller et al., 2001).
169 Under conditions of high discharge (i.e., $> 3,000 \text{ m}^3 \text{ s}^{-1}$) combined with NE winds,
170 hypopycnic plumes holding strong salinity stratification off the lagoon's mouth are
171 observed as less dense freshwater from the river discharge flows over the denser saline
172 water (Calliari et al., 2009). The formation of the continental discharge plume over the
173 SWAO carries a suspended sediment concentration of up to 12 g m^{-3} (Jung et al., 2020).
174 The Patos Lagoon coastal plume is a small-scale plume that exhibits surface variability
175 resulting from two principal modes. The first one, associated with the southwest and
176 southeast plume displacement, forced by northeasterly winds, accounts for 70% of the
177 plume variability, the upward velocities in the near field, the stratification processes, and
178 potential energy anomalies (Marques et al., 2009, 2010a). The second one, related to the
179 north and east plume migration due to south quadrant wind during the passage of frontal
180 systems, accounts for the remaining 19% of coastal variability, promoting downward
181 velocities and mixing conditions in the near field (Marques et al., 2010a).

182 Tides display a mean range of 0.3 m and characterize the lagoon as microtidal, mainly
183 of diurnal regime (Moller et al., 2001). Tidal effects and Earth rotation play a secondary
184 role in the plume behavior, mainly affecting its shape, mixing process, and offshore
185 penetration (Monteiro et al., 2011). Additionally, bed friction impacts mainly the
186 alongshore direction in the coastal region (Marques et al., 2010b). In this sense, the
187 advection and deformation of the dominant density fields control the mixing and
188 stratification of the Patos Lagoon coastal plume, considering time scales associated with
189 the influences of winds (5- and 15-day processes) and river discharge (30-day processes)
190 (Marques et al., 2009).



191

192 *Fig. 1. Patos Lagoon (a) bathymetry for the numerical domain and numerical mesh (b) for the simulations,*
 193 *where the main tributaries, i.e., Guaíba River, Camaquã River, and São Gonçalo Channel are indicated.*
 194 *Boundary conditions used to force the model are depicted with black solid lines. The location of the jetties*
 195 *(c) represents the connection with the SWAO and the area of strongest influence of the coastal plume of*
 196 *suspended sediments and organic/inorganic debris.*

197 2.2. Hydrodynamic modeling

198 The TELEMAC-3D model (www.opentelemac.org) has been widely employed to
 199 study the dynamics of Patos Lagoon itself (Fernandes et al., 2002), its estuary (Fernandes
 200 et al., 2004) and the resulting coastal plume (Marques et al. 2009; 2010a; Monteiro et al.
 201 2011, Fernandes et al, 2021). TELEMAC-3D solves the 3D Reynolds-Averaged Navier-

202 Stokes Equations, considering the Boussinesq and Hydrostatic approximations, based on
203 the Characteristics Method and Finite Element Method (Hervouet, 2007; Villaret et al.,
204 2013). The TELEMAC-3D model also solves the mass conservation equation, which
205 simulates the temporal and spatial variation of active tracers such as salinity, temperature,
206 and suspended sediments, incorporating the flocculation process. The model domain was
207 discretized with an unstructured finite element mesh according to Fernandes et al. (2021)
208 (**Fig. 1b**), based on digitalized nautical charts of bathymetric data obtained from the
209 Brazilian Navy and complementary data provided by the Rio Grande Port Authority
210 (Portos RS). The domain covers from 29° and 36°S to 48° and 54°W (**Fig. 1b**) and reaches
211 the 3700 m depth isobath. The numerical mesh consists of approximately 33,000
212 elements, 50,000 nodes, and seven sigma levels in the vertical, with results output every
213 6 h.

214 For the oceanic boundary conditions (**Fig. 1b**), sea level data and regional tidal
215 velocity fields were obtained from the OSU Tidal Inversion System (OTIS – Egbert and
216 Erofeeva, 2002), an internal coupling to TELEMAC (TPXO). Temperature and salinity
217 fields were derived from the HYCOM + NCODA Global project (Hybrid Coordinate
218 Ocean Model, <https://hycom.org/>), with a temporal and spatial resolution of 3 h and 0.08°,
219 respectively. At the surface boundary, data from ECMWF ERA-Interim and ERA5
220 (European Centre for Medium-Range Weather Forecast, <http://www.ecmwf.int/>) were
221 applied with a temporal and spatial resolution of 6 h and 0.75°, respectively. All these
222 data were interpolated in time and space for each numerical grid point. Daily river
223 discharge data from the main tributaries (Guaíba and Camaquã) provided by the National
224 Water Agency (ANA, www.hidroweb.ana.gov.br) were used for continental boundaries.
225 Discharge data for the São Gonçalo Channel were obtained as water level data from the
226 Mirim Lagoon Agency (ALM, <https://wp.ufpel.edu.br/alm/>), and transformed into flow
227 data using the Rating Curve Method (Oliveira et al., 2015).

228 The hydrodynamic output of TELEMAC was calibrated and validated for the study
229 region by several authors (Fernandes et al., 2001, 2002, 2005, 2007; Marques et al.,
230 2010a; Bitencourt et al, 2020a; Lisboa et al., 2022), always yielding from good to
231 excellent values of both RMSE and RMAE. The detailed calibration and validation of the
232 numerical mesh used in this study have already been published in Fernandes et al. (2021).

233 Time series of current velocities, salinity, wind intensity, and direction were extracted
234 for the jetties area at 6 h resolution to feed discussion of the microplastic transport model

235 results. In this sense, the mean position of the salinity field was analyzed and used to
236 calculate each plume duration, together with the cross-sectional vertical profile of the
237 mean dispersion of the plumes. Pearson correlation tests were performed to evaluate the
238 relationship between these variables in the plume's behavior, followed by linear
239 regression to infer the type of such relationship.

240 *2.3. Microplastic transport modeling*

241 The TrackMPD model (Jalón-Rojas et al., 2019) simulates MP transport. TrackMPD
242 consists of a 3D, extended Lagrangian numerical model developed in Matlab as an ad hoc
243 open-access source to analyze the transport of plastic waste in marine and coastal systems.
244 The model interpolates the hydrodynamic unstructured mesh from TELEMAC-3D to
245 produce a structured mesh. To achieve a good resolution at the area of interest (i.e., jetties,
246 **Fig. 1c**) without compromising computational time, the domain size was reduced,
247 covering the sub-space between 52.38°-51.63°W and 32.89°-31.81°S, with a subsequent
248 differential refinement for a better definition on the jetty's region.

249 TrackMPD differs from traditional Lagrangian particle tracking models by modifying
250 particle characteristics over time and attributing specific behaviors. It consists of
251 interactive and independent modules, which make it possible to incorporate all physical
252 processes (such as advection, vertical and horizontal dispersion, sinking, resuspension,
253 stranding, washing-off, and wind transport) together with the physical properties of the
254 materials (size, density, and shape) and their behaviors (e.g., degradation and biofouling,
255 modifying the above characteristics in the indicated time frame). For further details on
256 TrackMPD calculations, readers can refer to Jalón-Rojas et al., (2019), Jalón-Rojas &
257 Marieu (2023), and the Supplementary Data (Annex I).

258 *2.4. Selection of simulation parameters for TrackMPD*

259 The simulated periods were chosen based on the Ocean Niño Index (ONI) values
260 (<https://ggweather.com/enso/oni.htm>). NOAA developed ONI to identify positive or
261 negative El Niño (La Niña) events in the Tropical Pacific using the three-month average
262 surface temperature (SST) anomaly for the Niño 3.4 region (5°N-5°S, 120°-170°W). The
263 ONI classification allows identifying the intensity of events for at least three consecutive
264 overlapping three-month periods, between "Weak" (with an anomaly of 0.5 to 0.9 SST),
265 "Moderate" (1.0 to 1.4), "Strong" (1.5 to 1.9) and "Very Strong" (≥ 2.0) events. In this

266 sense, we selected neutral (Mar/2014-Oct/2014), weak (Oct/2014-May/2015), and very
267 strong (Aug/2015-Mar/2016) El Niño phases as simulation periods. For each period, three
268 well-represented plumes were chosen from the salinity and current velocity field,
269 considering a time scale from hours to a few days for each simulation, where the processes
270 of coastal plume formation are more important. The selected periods of simulation and
271 further statistics are provided in results section 3.3.

272 Sensitivity tests were conducted to identify numerical configurations that most
273 accurately represented the simulation processes. For each parameter, activation within the
274 control scenario (involving passive particles in the advective field provided by the
275 TELEMAC velocity field) was carried out, with the aim of selecting values that closely
276 resembled those of the control simulations. In this sense, vertical and horizontal
277 dispersion coefficients were tested, together with the computational time for internal
278 calculation and output liberation, the number of particles, and the release point.
279 Horizontal (K_h) and vertical (K_v) dispersion coefficients were set up to typical values
280 from coastal systems: 0.25 for K_h and 10^{-5} for K_v (Bogucki et al., 2005; Diez et al. 2008;
281 Jalón-Rojas et al., 2019), as the small variations in these parameters didn't show
282 significant changes. The computational time was also tested, varying from 2 sec, 30 sec,
283 1 min, 5 min, 10 min, and 15 min. The 10 minutes timestep was selected since it showed
284 the best resolution in the particle's trajectories, without being as computationally
285 expensive, as was observed for the shorter times. The particle amount was optimized by
286 testing with 240 particles to ensure a comprehensive representation and facilitate even
287 distribution across the available logical processors for efficient simulations. Finally, the
288 release point was selected by initially considering the sampling sites and subsequently
289 determining the location where the export behavior of the coastal plume through the
290 jetties could be most effectively observed.

291 Because the time scale was used to focus on plume behavior, the effects of
292 degradation and biofouling were disregarded for the simulations, and only the deposition
293 and resuspension modules were activated. Then, for each plume event simulated, we use
294 two configurations of TrackMPD: (1) passive scenario only with advection and dispersion
295 (incorporating the values of the horizontal and vertical dispersion constants K_h and K_v ,
296 respectively) and (2) sinking scenario with advection, dispersion together with the
297 deposition resuspension modules activated. **Table 1** resumes the main parameters chosen
298 for each scenario.

299 The characterization of the particles was defined from one of the most represented
300 items in field sampling (Section 3.1, **Fig. 2**) in terms of MPs type and average size, the
301 same as referred to in other coastal and ocean environments from the literature (Zhao et
302 al., 2019). Polyester fibers, highly produced in the industry (Geyer et al. 2017), were
303 reported as a land source important contribution from industrial and urban washing
304 machines, that are not being retained by effluent treatment plants (Napper and Thompson,
305 2016). Thus, polyester fiber was selected for the TrackMPD simulations and the reference
306 value of density according to the type of particle was used as 1380 kg m³ (Nguyen et al.
307 2022). The sinking behavior was set up by adding the constant settling velocity value
308 (Ws) of 0.1 mm s⁻¹ proposed as the lowest value experimentally found by Nguyen et al.
309 (2022) for polyester fibers and referenced by Jalón-Rojas et al. (2021) as a relatively low
310 value characteristic of small fibers.

Table 1. Summary of the particle parameters such as size, polymer type, density, and settling velocity, and the configurations selected for the simulations, such as timestep for the calculations and for the outputs, dispersion constants, and particle behaviors modules selection for activation.

GENERAL PARAMETERS									
Cores	TimeStep Calc	TimeStep Out	# days	Particle Shape	Particle Polymer	Particle Density	Particle Size	Particle Dequi	
	10	10	Plume duration	Fiber	Polyester	1380*	0.0015	0.0012	
TRACKMPD SIMULATION CONDITIONS									
		Kh	Kv	Initial Ws Option	Beaching	Refloating	Deposition	Resuspension	Biofouling
1. PASSIVE: (Advection + Dispersion)		0.25	10 ⁻⁵	0	No	No	No	No	No
2. SINKING: (Advection + Dispersion + Deposition)		0.25	10 ⁻⁵	0.1 mm s ⁻¹ *	No	No	Yes	Yes	No

*Nguyen et al, 2022.

311 2.5. Model results postprocessing

312 For each simulation, total particle displacements were calculated as the resulting
313 distance from the liberation point to the final position, and total trajectories were
314 determined as the sum of distances covered by particles at each timestep. Statistical
315 analysis included a one-way ANOVA with posthoc Tukey HSD analysis to assess the
316 statistical differences in total displacement and trajectory. In addition to the ANOVA, a

317 linear regression, and Principal Component Analysis (PCA) were applied to reduce
318 dimensionality and characterize the relationship between particle distances, water
319 discharge, plume duration, and wind intensity.

320 The spatial distribution of MPs was represented by calculating the density probability
321 maps for surface and vertical dispersion. For this purpose, the longitude, latitude, and
322 depth of each particle at each time step was computed, the number of particles was
323 quantified in the grid, and then normalized with respect to the total number of data points
324 in the specified arrays to obtain a probability distribution in the study domain. Vertical
325 density figures (Supplementary Data, Fig. S8, S12, among others) represent results as
326 depth versus time probability, showing the time steps until final fate.

327 *2.6. Model validation with field data*

328 The transport trends of MPs were validated by undertaking two field surveys with a
329 total of seven surface water collections within the study area in the context of the
330 PLUMPLAS Project (Zavialov et al., 2020). MP density accumulation results were
331 compared to TrackMPD simulations results (similar configuration to those described
332 above) performed for the same data of field collections.

333 Sampling was carried out during plume formation events, corresponding to June 3rd
334 and July 21st, 2022, using a 300 μm "Manta" type surface trawl, 30 x 60 cm mouth and
335 2 m long (Zheng et al., 2021). Surface layer suspended material was collected for
336 approximately 15 min at each transect (**Fig. 3**). The in-situ measurements were
337 synchronic with Acoustic Doppler Current Profiler (ADCP) and conductivity,
338 temperature, and depth (CTD) readings, acquired at 0.1 m intervals. These measurements
339 provided information on current velocity and water flow filtered by the net, along with
340 the physical characterization of the water mass (**Table 2**) as recently recommended by
341 Defontaine and Jalón-Rojas (2023).

342 Samples were stored in glass jars and taken for post-processing in the laboratory,
343 following the methodology suggested by Zheng et al. (2021) and adapted for our case
344 study. All steps were performed with special care to avoid environmental and systemic
345 cross-contamination, using cotton lab coats, manipulation bell, glass materials such as
346 Petri dishes, filtration system and containers, and metal tools such as spatulas and
347 tweezers, all cleaned with distilled-filtered water beforehand (Bogdanowicz et al. 2021).

348 Samples were processed with 30 mL of 35% H₂O₂ previously filtered and digested in an
349 oven at 40°C for at least 48 h until organic matter was removed (Pfeiffer & Fisher 2020).
350 Subsequently, samples were filtered in a vacuum system, inside a closed hood to avoid
351 environmental contamination, on 95 mm nitrocellulose filters with 300 µm porosity, and
352 stored in glass petri dishes. The dry filters were analyzed under a tri-ocular stereoscope
353 (COLEMAN, model NSZ 606T) coupled to a camera and the Opticam Microscopy
354 OPTHD software (version 3.7.11443.20180326) inside a closed protective hood, and
355 items identified as potential plastic were characterized at 10x magnification according to
356 size, shape, and color. These debris particles were chemically characterized using
357 Fourier-transform infrared (FTIR) spectroscopy analysis (Bruker's Alpha P) in ATR
358 mode using a zinc selenide (ZnSe) crystal in the region of 4000-600 cm⁻¹ with a resolution
359 of 4 cm⁻¹ and 128 scans. The particle spectra were processed and identified using the
360 Open Specy library (<https://openanalysis.org/openspecy/>) (Cowger et al., 2021) and then
361 the total MP concentration was corrected incorporating the percentage of synthetic
362 particles observed by the FTIR analysis.

363 Blanks were used to assess contamination of both the lab procedure itself and the
364 environment (Bogdanowicz et al., 2021). In this sense, each sample was accompanied by
365 a blank Petri dish with a clean filter throughout the procedure, which was treated the same
366 as the samples were. To evaluate the contamination of the whole filtration system, an
367 extra blank was performed using the same procedure as the samples, but without a sample.
368 When plastics were detected in the blanks, the same items were finally removed from the
369 sample and those recorded in the system blank were also removed for all samples.

370 The concentration results of MPs were also utilized to extrapolate the MP export rate
371 of the upper layer of the estuary mouth, by employing a bottom-up approach (Boucher et
372 al., 2019). This involved considering the observed surface concentration of MPs at the
373 sampling point near the mouth of the jetties. This place corresponds to where the water
374 flow was estimated for each survey, and where the lowest concentration of MPs was
375 recorded. The water discharge on the upper 1 m layer was calculated by extracting the
376 water current velocity fields from the modeling results and by multiplying them by the
377 surface area, estimated at 700 m². This allowed us to calculate the total export of MPs by
378 the surface layer of the plumes under differential conditions of water discharge, in a
379 conservative way.

380 3. Results

381 3.1. Field data for model validation

382 3.1.1. Observed MPs

383 MPs were detected in all seven surface water samples (**Table 2**). A total of 2564
384 potential plastic particles were counted and identified, ranging from 93 to 725 items per
385 transect (**Fig. 2**), with an average of 0.85 items m⁻³ (SE = 0.48 items m⁻³). Length ranged
386 from 0.04 to 98.97 mm, categorizing most of the particles as MPs (i.e., 1 µm - 5 mm)
387 where the highest size frequencies were detected (92%). Mesoplastics (5 mm - 25 mm)
388 were detected with lower abundance (8%), and finally, just a few macroplastics were
389 observed (> 25 mm) (**Fig. 2**). Fibers represented 47% of the total MPs observed, being
390 the rest 53% fragments (including films and spheres).

391 The relative percentage of polymeric composition per sampling station in each field
392 survey is presented, as analyzed by FTIR spectroscopy is shown in Fig. 2. A total of 209
393 items was processed, constituting 8% of the total particles, with 189 items (90%)
394 identified as synthetic polymers. Eight different polymeric types were identified, where
395 Polypropylene (PP) and polyethylene (PE) accounted for 76% of confirmed plastic
396 particles, being the most representative polymers from recovered items. In the analysis of
397 June-MPs, polypropylene (PP) was the dominant polymer, i.e., 36%, followed by
398 polyethylene (PE) 30%, alkyd varnish from polyester resin paints 20%, polyester (PES)
399 11%, and polystyrene (PS) at 2%. In July-MPs, a higher polymeric diversity was
400 observed, especially in sampling station 1. Polyethylene was the most representative type
401 with 55%, followed by polypropylene 24%, polystyrene 9.66%, polyester 6%, epoxy
402 resin 2%, rubber 2%, alkyd varnish 1%, and ethylene-vinyl acetate (EVA) 1%.

Table 2. Field sampling characteristics and MP concentration for modelling validation.

	Station	Transect distance (m)	Filtered Volume (m ³)	Water Discharge* (m ³ s ⁻¹)	Plastic Fragments	Plastic Fibers	Total Plastics	Plastic concentration (items m ⁻³)
June -03-2022	1	2037	367	5900	59	114	174	0.47
	2	2408	433		102	138	239	0.55
	3	2315	417		50	33	83	0.20

July- 21- 2022	1	2100	378	10576	177	121	298	0.79
	2	1716	309	11472	124	107	231	0.75
	3	2554	460	11087	427	204	632	1.37
	4	2890	520	7418	283	368	652	1.25

*Water discharge through the lagoon outlet, estimated using fixed ADCP located at Pilot Station (**Fig. 1c**), multiplied by the current speed by the section area for the high level (10,000 m²) by the flow period (estimated in 12 hours).

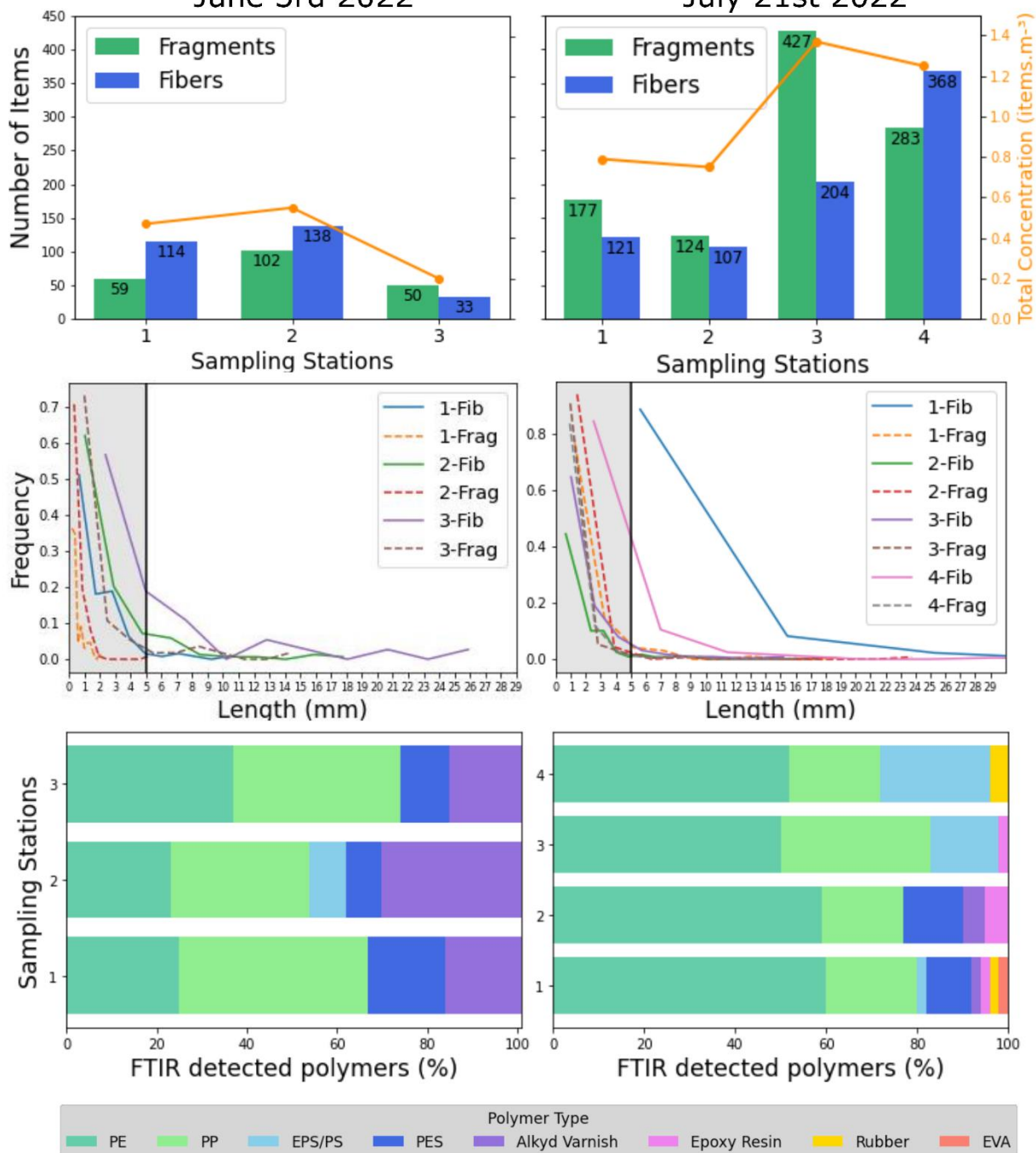
MP concentration was corrected according to the results of the FTIR analysis, i.e., 90% of the observed potential MPs.

403 **Fig. 2** shows the characterization of analyzed particles for both field samplings (June
404 3rd to the left, July 21st to the right of the plot), where the number of items per sample
405 (top) and the size distribution frequency (middle) are shown, together with the FTIR
406 results (bottom). During the June 3rd field survey (**Fig. 2** - left), the mean concentration
407 of all transects was 0.40 items m⁻³ (SE = 0.19 items m⁻³). In the first transect, outside the
408 jetties, a total of 174 plastic items were identified (0.47 items m⁻³), including 114 plastic
409 fibers (0.31 items m⁻³) with an average length of 1.82 mm and 59 plastic fragments (0.16
410 items m⁻³) with an average length of 0.52 mm. The second transect, set perpendicular to
411 the first, yielded 239 plastic items (0.55 items m⁻³), comprising 138 plastic fibers (0.31
412 items m⁻³) with an average length of 0.19 mm, and 102 plastic fragments (0.23 items m⁻³)
413 with an average length of 0.55 mm. The third transect, conducted next to the jetties'
414 mouth, showed a total of 83 potential plastic items (0.20 items m⁻³), which consisted of
415 33 plastic fibers (0.08 items m⁻³) with an average length of 5.29 mm, and 50 plastic
416 fragments (0.12 items m⁻³) with an average length of 2.08 mm.

417 During the July 21st field survey (**Fig. 2** - left), the samples attained a mean
418 concentration of 1.04 items m⁻³ (SE = 0.31 items m⁻³). The first transect was located next
419 to the jetties' mouth, where a total of 298 plastic items (0.79 items m⁻³) was observed,
420 including 121 plastic fibers (0.32 items m⁻³) with an average length of 5.39 mm, and 177
421 plastic fragments (0.47 items m⁻³) with an average length of 1.95 mm. The second transect
422 was conducted within the jetty's channel, resulting in a total abundance of 231 plastic
423 particles (0.75 items m⁻³), consisting of 107 plastic fibers (0.34 items m⁻³) with an average
424 length of 1.21 mm, and 124 plastic fragments (0.40 items m⁻³) with an average length of
425 1.22 mm. The third transect, located outside the jetties where the plume front was
426 observed, exhibited a total abundance of 632 items (1.37 items m⁻³), including 204 plastic

427 fibers ($0.44 \text{ items m}^{-3}$) with an average length of 2.16 mm and 427 plastic fragments (0.93
428 items m^{-3}) with an average length of 0.94 mm. The fourth transect, was conducted in the
429 Pilot Station area transversely set to the channel and recorded a total abundance of 652
430 plastic items ($1.25 \text{ items m}^{-3}$), consisting of 368 plastic fibers ($0.71 \text{ items m}^{-3}$) with an
431 average length of 3.13 mm and 283 plastic fragments (0.54 item m^{-3}) with an average
432 length of 1.28 mm.

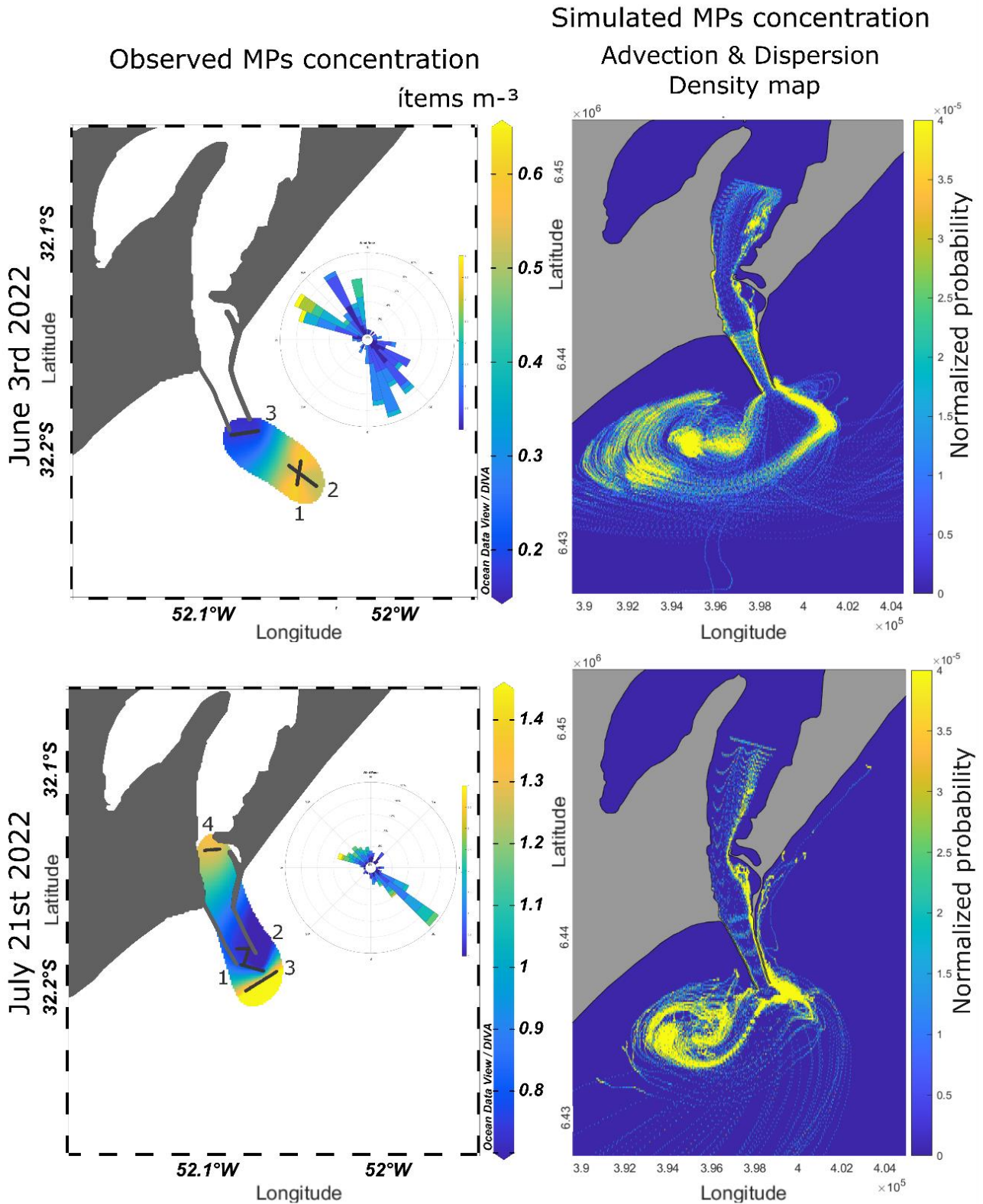
Observed MPs concentration in surface water
June 3rd 2022 July 21st 2022



433 **Fig. 2.** Field observed results and FTIR analysis from the manta net surface samplings for both June 3rd
 434 (left) and July 21st, 2022 (right). The uppermost plots show number of items corrected according to the
 435 results of the FTIR analysis, following the classification as fragments (green) and fibers (blue) for each
 436 sampling station, and the second Y-axis (on the right) expresses in orange line the total concentration
 437 values for each station in items.m⁻³. Middle-plots show the length frequency distributions, where each MP
 438 category (1 μm - 5 mm) is pointed out within the vertical a gray box, where the solid lines correspond to
 439 fibers and dotted lines correspond to fragments. The lowermost plots show the relative polymeric
 440 composition of identified MPs by FTIR spectroscopy, according to the color references described below.

441 3.1.2. Model validation

442 For model validation purposes, **Fig. 3** shows the concentration (items m⁻³) of field-
443 sampled MPs on the left side of the plot together with the TrackMPD simulation results
444 displayed as a probability density map for each field survey (June 3 at the top, July 21 at
445 the bottom) to the right of the panel. TrackMPD results revealed a good agreement
446 between observed and simulated data. While a direct quantitative comparison of these
447 variables is not feasible, a qualitative assessment reveals a consistent concentration
448 pattern in both observed and simulated scenarios. Notably, both situations exhibit a
449 similar trend of higher MPs concentration near the saline front adjacent to the jetties'
450 mouth. This consistent replication of observed patterns strongly indicates the model's
451 reliability for reproducing the transport features of MPs.



452 *Fig. 3. Field experiment composition for model validation. Comparison between observed versus simulated*
 453 *distribution of MPs in the coastal plume. The left section of the plot shows the spatial abundance*
 454 *distribution together with the wind rose characterization, where black lines refer to the location of each*
 455 *sampling station. Right plots show the TrackMPD results, simulated for the same day, expressed as a*
 456 *density map. Note the similar accumulation pattern on the convergence areas of the coastal plume for both*
 457 *observed and simulated data.*

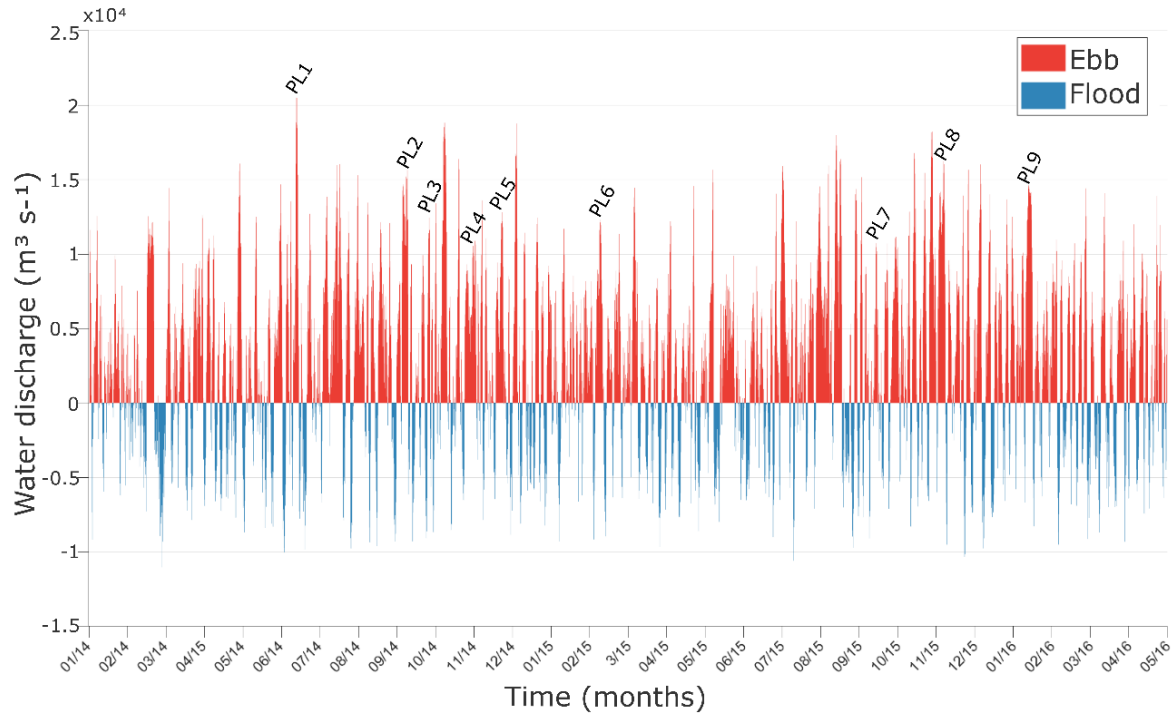
458 3.2. Estimation of MP export trough observed coastal plume events – a bottom-up
459 approach

460 During the June 2022 plume event, the mean water discharge value was lower than
461 the second event of July 2022, which exhibited a strong mean discharge. The mean water
462 discharge calculated at the surface layer (1 m) represented, on average, 7% of the mean
463 water discharge out of the entire water column between the jetties on the analyzed plumes.
464 For the June 2022 survey, by extrapolating the MP concentration observed in transect 3
465 ($0.20 \text{ items m}^{-3}$; **Fig. 2**), the water discharge restricted to the surface layer ($413 \text{ m}^3 \text{ s}^{-1}$),
466 resulted in 83 particles exported per second. This means a total export of approximately
467 7.1×10^6 items per day. Similarly, for the July 2022 survey, considering the MP
468 concentration in transect 2 ($0.75 \text{ items m}^{-3}$; **Fig. 2**) and a water discharge value of 803 m^3
469 s^{-1} , the number of MPs being exported per second attained 602 items. Thus, a total export
470 of 52×10^6 items per day was calculated.

471 3.3. Coastal plume hydrodynamics

472 The time series of modeled water discharge between from the lagoon to the ocean,
473 from January 2014 to May 2016, is reported in **Fig. 4**, where positive (negative) values
474 indicate ebb (flood) flows. The mean discharge at the mouth was $2,566 \text{ m}^3 \text{ s}^{-1}$, within
475 well-represented plume events selected based on the ONI index. Table 3 shows the
476 descriptive statistics for each of the nine simulated plumes, where the ONI condition,
477 duration, direction of deflection, and mean discharge were described. The first three
478 plumes (i.e., PL1, PL2, PL3) formed under an ENSO neutral period, the fourth, fifth and
479 sixth plume (PL4, PL5, PL6) developed under weak ENSO events, while the last three
480 plumes (PL7, PL8, PL9) formed during very strong ENSO events.

481 Results suggest that there was no significant relationship in terms of water discharge
482 at the mouth and the ONI index, but there was a significant positive relationship between
483 the water discharge and the wind intensity ($r = 0.74$; $p < 0.05$). The linear regression for
484 these variables showed a y-axis intercept value of 620.3 and the coefficient for wind
485 intensity was 1,112.7, with an R^2 of 0.54 (Fig. S4).



486 **Fig. 4.** Modeled water exchange at the cross-sectional area of the jetties. Positive values refer to the
 487 discharge outflow (ebb - red), and negative values refer to the inflow (flood - blue). Each plume selected
 488 for further investigation is indicated above the discharge peaks (i.e., PL1 through PL9).

489 Plumes PL1, PL3, PL8, and PL9 were classified as high-discharge events at the
 490 mouth, with mean discharge values of $10,497 \text{ m}^3 \text{ s}^{-1}$, $12,555 \text{ m}^3 \text{ s}^{-1}$, $12,139 \text{ m}^3 \text{ s}^{-1}$, $11,167$
 491 $\text{m}^3 \text{ s}^{-1}$ respectively (**Table 3**). Similarly, PL2, PL4, PL5, PL6, and PL7 were classified as
 492 either low or intermediate discharge events with mean water discharge of $6,121 \text{ m}^3 \text{ s}^{-1}$,
 493 $6,956 \text{ m}^3 \text{ s}^{-1}$, $6,601 \text{ m}^3 \text{ s}^{-1}$, $7,439 \text{ m}^3 \text{ s}^{-1}$, $6,308 \text{ m}^3 \text{ s}^{-1}$, respectively. All events showed the
 494 predominance of northerly winds with different combinations of east and west wind
 495 components. Fig. 5 shows three out of the nine plumes (PL1, PL5, and PL8) to depict the
 496 common hydrodynamic trends of the plumes. The supplementary material reports the
 497 remaining plumes (Fig. S1, S2, S3).

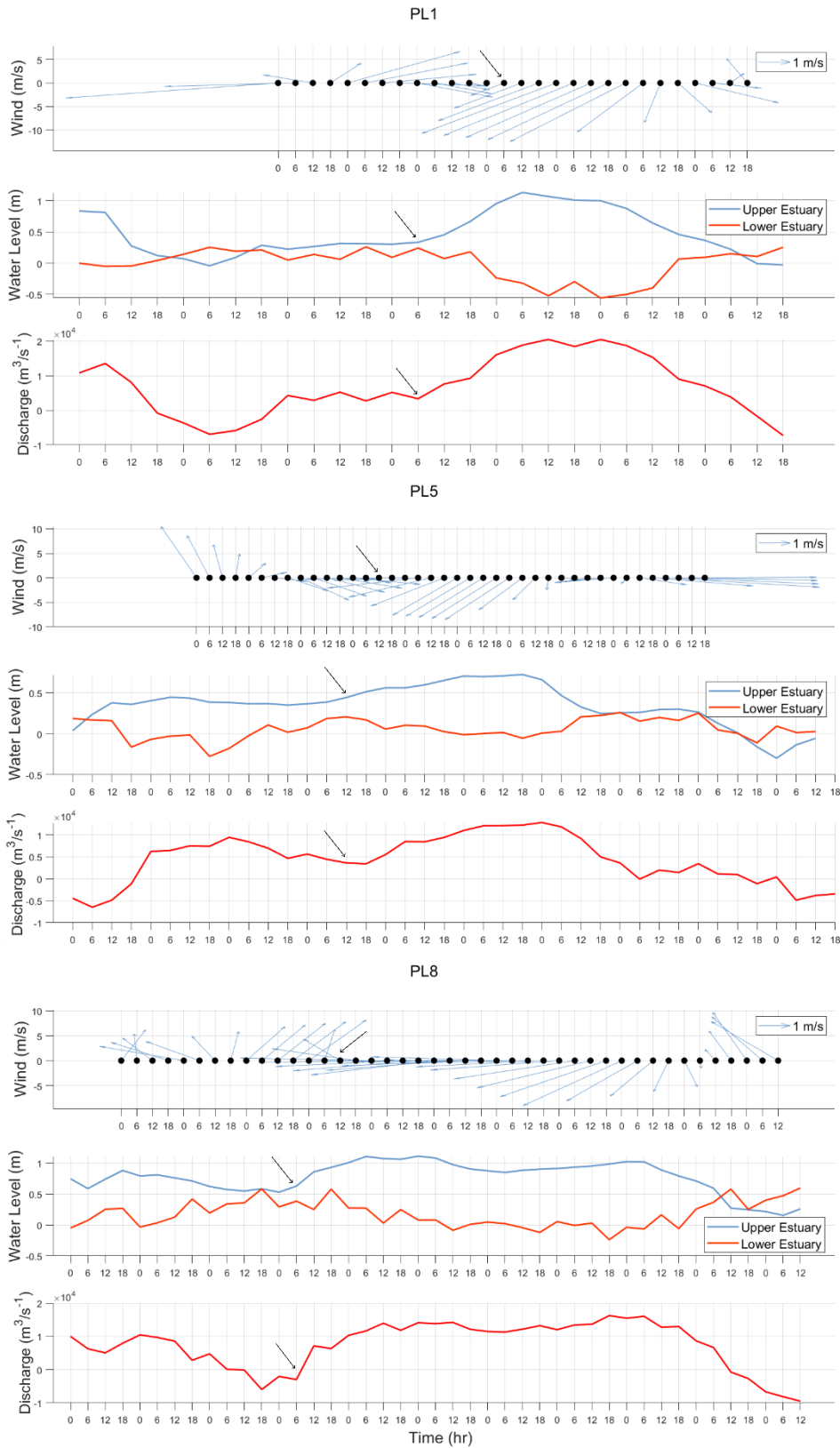
Table 3. Mean discharge, wind intensity, plume duration, and ONI condition for each simulated plume. High discharge events of PL1, PL3, PL8 and PL9 are depicted with bold.

REF	El Niño Condition (ONI)	PERIOD			Mean Discharge ($\text{m}^3 \cdot \text{s}^{-1}$)	Mean Wind Intensity ($\text{m} \cdot \text{s}^{-1}$)
		Initial date	Final date	Duration (days)		
PL1	Neutral	2014/06/10	2014/06/13	3	10,497	8.2
PL2	Neutral	2014/07/29	2014/08/03	6	6,121	4.7

PL3	Neutral	2014/09/02	2014/09/06	5	12,555	10.5
PL4	Weak	2014/10/22	2014/10/29	8	6,956	5.3
PL5	Weak	2014/11/16	2014/11/22	6	6,601	6.1
PL6	Weak	2015/02/01	2015/02/06	5	7,439	8.1
PL7	Very Strong	2015/09/19	2015/09/24	5	6,308	7.9
PL8	Very Strong	2015/10/24	2015/10/31	7	12,139	7.7
PL9	Very Strong	2016/01/20	2016/01/25	5	11,167	8.3

498

499 As shown in **Fig. 5**, the onset of each plume event was observed under the sustained
500 incidence of northeasterly winds. This wind condition modulated the water level
501 difference between the upper and lower reaches of the estuary, thus generating the
502 necessary barotropic pressure gradient for the freshwater export as a coastal plume to the
503 ocean. The highest peak of discharge from the jetties' mouth (i.e., PL1), attained a
504 maximum discharge value of $20,509 \text{ m}^3 \text{ s}^{-1}$ and a mean discharge of $10,497 \text{ m}^3 \text{ s}^{-1}$ during
505 the three-day plume event. During PL5, the maximum discharge value was $12,807 \text{ m}^3 \text{ s}^{-1}$
506 **(Fig. 4)**, while the mean discharge during the six-day plume event was $6,601 \text{ m}^3 \text{ s}^{-1}$. The
507 wind incidence was predominantly from the northeast, but as the wind direction shifted
508 to the south, the water flux was quickly reversed **(Fig. 4)**, and the plume broke up and
509 ceased to disperse freshwater to the ocean. For PL8, the maximum discharge value was
510 $16,246 \text{ m}^3 \text{ s}^{-1}$ and mean discharge of $11,167 \text{ m}^3 \text{ s}^{-1}$ for the seven-day plume event.



511

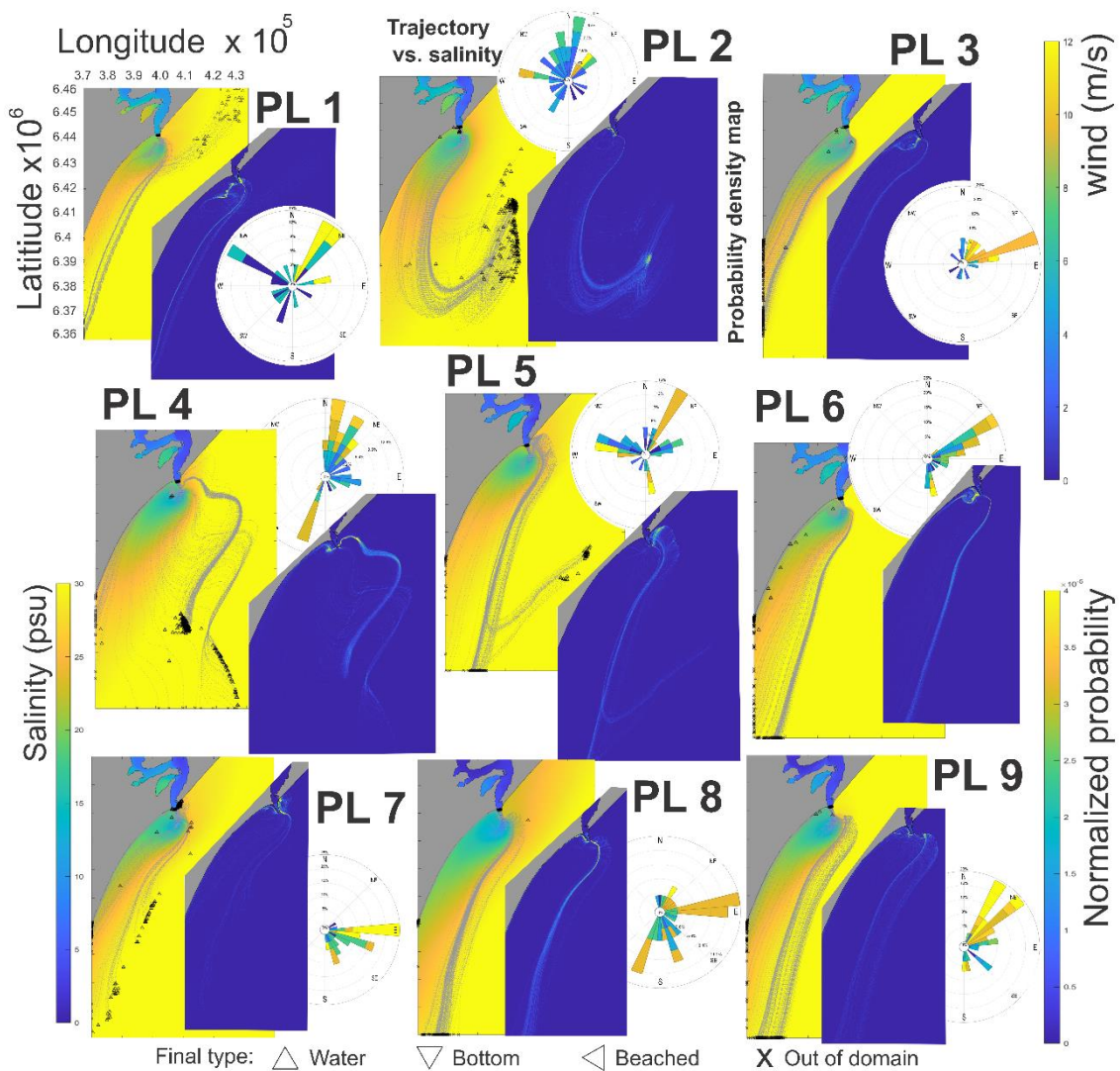
512 *Fig. 5. Wind intensity and direction, water level (upper and lower estuary) and water discharge at the*
 513 *jetties mouth for the high discharge plumes PL1, PL5 and PL8. The time series extraction for this analysis*
 514 *started three days before each plume formation. The black arrows indicate the moment of plume formation*
 515 *and the time of particle release for the TRACKMPD simulations.*

516 3.4. MP dynamics simulations

517 3.4.1. Passive scenario: Advection – Dispersion simulations

518

519 The results of the first configuration are presented in **Fig. 6** as surface trajectories over
 520 the mean salinity and density maps for surface probability for all nine plumes. Such
 521 figures also show the density probability map and the wind rose characterization for a
 522 complete understanding of both MP and plume behavior. For this first set of simulations
 523 (**Fig. 6**), the particles were not only affected by advection, but also for the turbulent
 524 dispersive process. Under this configuration, the MP dynamics in the coastal plume area
 525 of influence was mainly controlled by the water discharge and wind direction.



526

527 **Fig. 6.** Simulated MP trajectories for advection and dispersion conditions for all nine plumes. For each
 528 plume (subplots 1-9), yellow figures are the surface 2D trajectories plotted in relation to the mean salinity,
 529 where the symbols indicate the final fate of particles, as referenced in the legend (Water ▲, Bottom ▼,

530 *Beached* (◀) and *Out of Domain* (✕). Blue figures show the density maps for the surface dispersion. Wind
531 roses are also plotted for each plume to complement the understanding of the MP behavior.

532 As a general trend, all simulated particles of all nine plumes remained on the surface,
533 with a south-west displacement trend, where 69% of them were transported out of the
534 domain (Table S2), and vertically most of them reached the 5 m isobath. By the end of
535 the simulations, particles were transported out of the domain through the south-west
536 ocean boundary, about 49 and 86 km from the liberation point. The particle trajectories
537 to the southwesterly direction parallel to the coastline represent the north-easterly wind
538 response, characteristic during such plume events. When the surface trajectory was
539 analyzed together with the mean salinity field (**Fig. 6**), particle trajectories followed the
540 freshwater dispersion and were more likely to occur on the plume boundary, over the 25
541 isohalines approximately, where the saline front was located. Following this pattern, the
542 particles accumulated rather on the small gyre generated between the west side of the
543 jetties and Cassino beach, and on the north-easterly region of the jetties (**Fig. 6**).

544 Plumes events PL1, PL3, PL6, PL7, PL8 and PL9 exhibited the same accumulation
545 pattern mainly in the gyre region and in the southwest pathway contouring the saline front
546 about 9 km offshore. On PL1, approximately 20% of the particles showed a final fate
547 inside the domain, and the remaining 80% was transported to the offshore region, out of
548 domain after the 4 days of plume duration, with a tendency of accumulation in two
549 regions. Firstly, in the gyre southwest of the jetties' mouth, and secondly along the east
550 side of the jetties, showing first an accumulation spot in the northeast region, but then
551 being transported to the southwest following the parallel pathway to the coastline,
552 contouring the freshwater plume (**Fig. 6** PL1). The maximum depth attained by the
553 particles during this event was 5 m (Fig. S8), lasting for 3 days after the liberation, but
554 most of the particles were accumulated between the surface and 2 m depth (Fig. S9).

555 PL3 (**Fig. 6** PL3) had only 3% of the MP particles inside the domain at the end of the
556 simulation, and the remaining 97% were recorded out of domain. During this 5-day event,
557 particles showed a vertical dispersion between the 2 m depth, reaching a maximum depth
558 of 4 m in the gyre region south of the jetties and Cassino beach (Fig. S12) after the fourth
559 day (Fig. S13). PL6 lasted 5 days and a total of 3% of the particles remained inside the
560 domain with the remaining 97% out of domain. Vertically, they reached a maximum
561 depth of 6 m in the gyre accumulation spot, at the end of the simulation period (Fig. S16,
562 S17). PL7 also lasted 5 days. The final fate of the particles showed that 60% stayed inside

563 the domain and 40% were transported out of domain by the end of the simulation. Same
564 as the behavior observed for PL1, the particles showed a dispersion firstly to the north,
565 where they first accumulated and then were transported to the southwest quadrant.
566 Vertically, the densest depth was observed between surface and 1 m depth, but the
567 maximum depth was 4 m during the second day of simulation (Fig. S20). PL8 event
568 lasted 7 days, with almost all particles transported out of the domain also in the south-
569 west boundary. The vertical distribution was on average between the surface and the 3 m
570 depth during the fourth day, with only a few particles reaching a maximum depth of 5 m
571 (Fig. S24, Fig. S25). PL9 lasted 5 days and at the end of the simulation, only 1% of
572 particles remained in the domain (and 99% out). Such a 1% consisted only of 3 particles
573 located between the jetties and Cassino beach. Most of the particles displayed a vertical
574 dispersion between the surface and 2 m depth. After the second day of simulation a depth
575 of 4 m was attained (Fig. S29).

576 PL2, PL4 and PL5 also showed MP dispersion patterns to an offshore direction. PL2
577 lasted 6 days and the expelled particles by the coastal plume were transported firstly to
578 the south-west, and then shifted to a north-east trajectory. In this case, the dispersion of
579 all the particles was completely maintained inside the domain (except for only one MP
580 particle) during the plume simulation. The densest areas were mainly located along the
581 westerly jetty outside the channel, in the saline frontline and in an offshore area to the
582 south direction. The vertical dispersion for this event was concentrated between the
583 surface and the 2 m depth, although the maximum depth in this area was 7 m at the fourth
584 day of simulation (Fig. S33). PL4 lasted 8 days and 75% of the particles remained in the
585 domain by the end of the simulation (and 25% out). The surface dispersion pattern for
586 this event showed northeasterly trajectories, where the densest regions were observed,
587 that ended up in an offshore area on the south quadrant of the domain, reaching a
588 maximum depth of 6 m (Fig. S36), after the fifth day of simulation (Fig. S37). The second
589 densest region for this event was also located in the gyre between the westerly jetty and
590 Cassino beach. PL5 consisted of a 6-day event where 17% of the MP particles stayed in
591 the domain, following an eastward trajectory, and the remaining 83% was transported
592 outside of domain by the sixth day of simulation. The densest areas were identified close
593 to the saline front to the northeast of the jetties. On the vertical axis they reached a
594 maximum depth of 4.5 m (Fig. S40), with a higher accumulation until 2 m depth, during
595 the three-day simulation (Fig. S41).

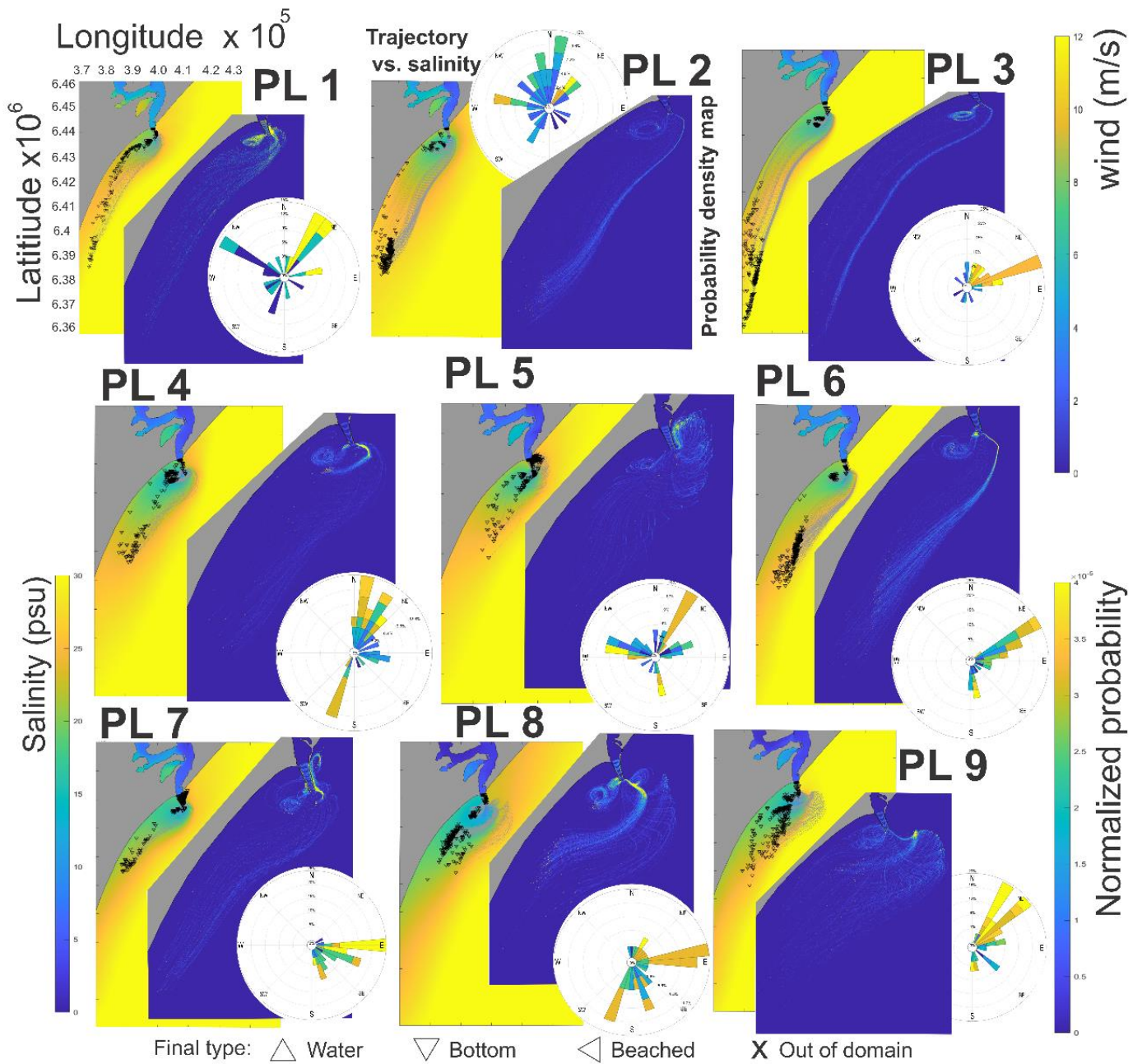
596 3.4.2. *Sinking scenario: Advection – Dispersion – Deposition simulations*

597 **Fig. 7** shows the results for the second set of simulations, where the deposition process
598 was activated as an input to the analysis of MP trajectories to represent the sinking
599 behavior in synthetic MP fiber. For all nine plumes, the main area of deposition followed
600 the freshwater dispersion with a southwest direction, parallel to the coast. Thus, 100% of
601 the particles were deposited at the bottom within the domain during the 48-hour plume
602 simulation for all the events.

603 In PL1, PL5 and PL7, two main areas of accumulation and deposition were detected,
604 the first one on the backside of the eastern jetty and the second one in the gyre region
605 between the western jetty and Cassino beach. During this dispersion pattern, particles
606 being expelled from the middle of the jetties mouth, followed first a trajectory to the east
607 and then shifted to the west during the deposition. Particles being expelled from the west
608 region, entered the gyre and deposited, whereas the remaining fraction was transported
609 southwest parallel to the coast. Particularly, plastic fibers in PL1 reached a maximum
610 depth of 13 m (Fig. S44), after 36 hours of simulation (Fig. S45). In PL5 they deposited
611 at a maximum depth of 16 m (Fig. S48) by the 48 hours of simulation (Fig. S49). PL7
612 also showed a deposition area both inside and outside the jetties' mouth, reaching 8 to 12
613 m depth, but in the southwest coastal region, the 12 m depth isobath was similarly attained
614 (Fig. S52).

615 For PL2, PL3 and PL6 there were two main areas of fiber deposition, one between
616 the jetties and Cassino beach, and the other in the southwest portion of the domain, just
617 off the freshwater boundary. The highest probability of occurrence of these MP fibers
618 under such event conditions was observed in the saline front and in the gyre. Thus, the
619 vertical trajectory plots for these events indicate that particle dynamics during the
620 deposition process, formed a loop until reaching a depth of 10 to 12 m in the gyre next to
621 the jetties in less than 24 hours. However, for PL2, the deepest deposition point was
622 recorded in the southwest portion of the domain, i.e., at 14 m depth after 30 hours of
623 simulation (Fig. S56, S57), where a maximum horizontal displacement of 61.1 km was
624 measured (Section 3.4.3, **Table 4**). Similarly, for PL3 and PL6, fibers reached a maximum

625 depth of 15 m in the southwest deposition area within 42 and 48 hours of simulation (Fig.
 626 S60 and Fig. S64 respectively).



627 **Fig. 7.** Simulated MP transport and accumulation for PL1 through PL9 considering advection, dispersion,
 628 and deposition processes. For each plume, yellow figures show the surface 2D trajectories plotted over the
 629 mean salinity, where the symbols indicate the final fate of the particles, as referenced in the legend
 630 (Water \blacktriangle , Bottom \blacktriangledown , Beached \blacktriangleleft and Out of Domain \times). Blue figures show the density maps for the surface
 631 dispersion. Wind rose plots are also reported for each case.

632 Finally, for plumes PL4, PL8 and PL9, most of the deposition occurred in the gyre
 633 between Cassino beach and the jetties, and in the southwest coastal region, following the

634 freshwater dispersion. PL8 also showed deposition areas inside the jetties. Accordingly,
635 for these simulations, the densest areas were observed within the gyre and in the eastern
636 side of the saline front just next to the jetties' mouth. In PL4 fibers reached a maximum
637 distance of 38.5 km away from the jetties' mouth in the southwest direction parallel to
638 the coast. Particles attained a depth of 10 to 12 m (Fig. S68) within 24 hours of simulation,
639 but a deepest localization of nearly 15 m was detected in the southwest coastal region
640 within less than 40 hours of simulation (Fig. S69). In PL8, the fibers reached the 11 m
641 depth isobath in the gyre region (Fig. S72), by the first 30 hours of simulation, whereas
642 the maximum depth of 16 m was reached at the 48 hours of simulation (Fig. S73). In PL9,
643 the maximum displacement distance of the fibers attained 39.6 km, and they deposited at
644 16 m depth (Fig. S76) within 48 hours of simulation (Fig. S77).

645 *3.4.3. MP particles displacement and trajectories*

646

647 **Table 4** shows the mean, maximum and minimum horizontal displacement and
648 trajectories traveled by the simulated MPs in the proposed scenarios, combining the nine
649 plume events and two model configurations: (1) Passive: advection + dispersion + sinking
650 and (2) Sinking: advection + dispersion + sinking + deposition). We observed differences
651 in the distances attained by the particles in each plume depending on the scenario
652 simulated and the environmental conditions forcing the MP behavior. In this sense, to
653 understand the significance of these physical drivers, they were statistically analyzed
654 together with the resulting distances. Under the passive configuration, the particle
655 displacements showed a stronger positive relationship with the water discharge ($r = 0.34$;
656 $p < 0.001$) than that of the wind intensity ($r = 0.18$; $p < 0.001$). However, when the
657 trajectory was added to the analysis, the relationship was inverted, showing negative
658 correlations with the wind intensity ($r = -0.48$; $p < 0.001$) and the water discharge ($r = -$
659 0.16 ; $p < 0.001$). An ANOVA test showed differences in the total distance between
660 plumes ($F = 207.1$; $p < 0.001$, Fig. S5). Also, the PCA diagram (Fig. 8), showed that the
661 MP mean displacement was rather related to the mean discharge whereas the trajectory
662 was related to the event duration.

663 Under the second configuration of sinking behavior, the displacements attained by the
664 particles characterized as fibers, from the liberation point until bottom deposit, showed a
665 weak positive relationship with water discharge ($r = 0.13$; $p < 0.001$), and wind intensity
666 ($r = 0.11$; $p < 0.001$). Similarly, the particle trajectories showed a low positive correlation

667 coefficient with water discharge ($r = 0.19$; $p < 0.001$), and with wind intensity ($r = 0.12$;
 668 $p < 0.001$), differing from the previous scenario where this relationship was negative.

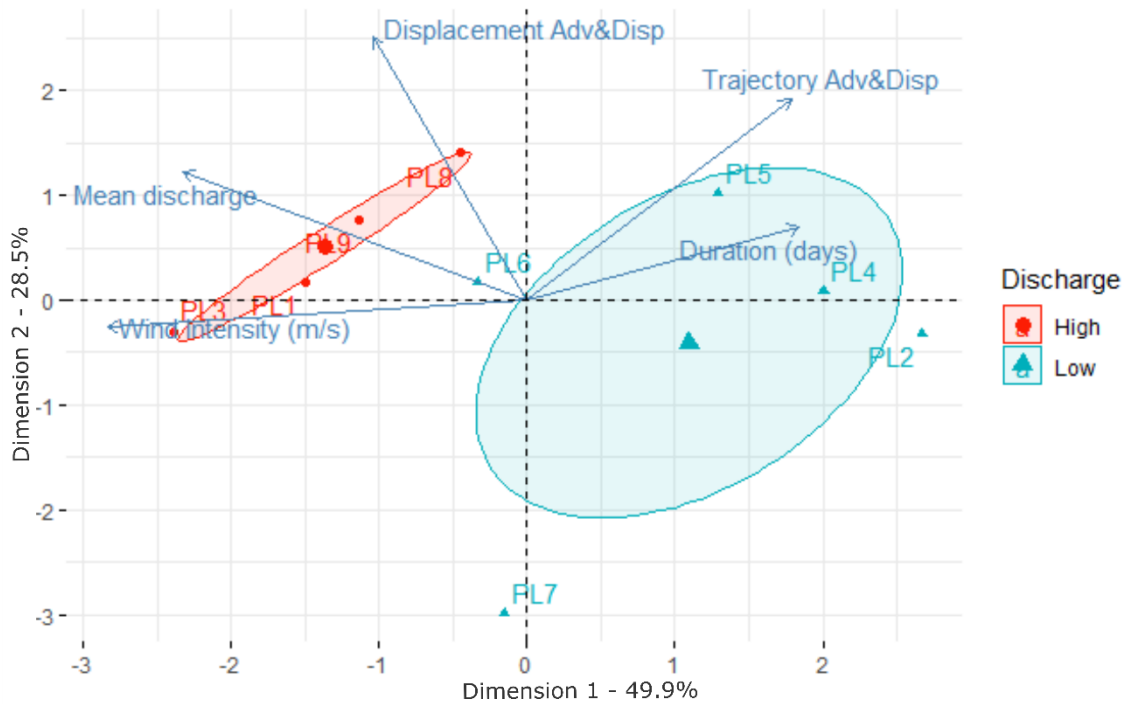
Table. 4. Summary of the simulated surface displacement and trajectory attained by MP particles in each TrackMPD simulation.

Plume Event	Scenario	Distance	Mean (m)	Max (m)	Min(m)	Standard Deviation
PL1	Passive	displacement	68,365	85,264	13,381	22,062
		trajectory	81,012	89,574	52,289	10,342
	Sinking	displacement	17,551	59,236	991	16,822
		trajectory	24,151	60,861	1003	15,541
PL2	Passive	displacement	38,534	60,753	316	14,080
		trajectory	109,857	131,260	21,251	24,753
	Sinking	displacement	43,562	61,098	2022	18,664
		trajectory	49,319	63,274	3260	13,371
PL3	Passive	displacement	59,372	78,629	856	13,605
		trajectory	64,995	82,272	24,645	10,743
	Sinking	displacement	43,294	83,602	1530	27,914
		trajectory	48,956	87,257	1532	25,314
PL4	Passive	displacement	56,834	83,590	5211	10,620
		trajectory	83,569	114,378	41,643	12,610
	Sinking	displacement	11,772	38,538	892	10,494
		trajectory	18,424	40,561	893	9336
PL5	Passive	displacement	73,754	85,394	20,624	16,168
		trajectory	99,724	145,827	46,594	19,678
	Sinking	displacement	5995	34,691	509	6882
		trajectory	12,905	37,666	897	6947
PL6	Passive	displacement	72,398	85,385	4115	17,272
		trajectory	79,016	87,534	37,389	11,800
	Sinking	displacement	26,021	46,697	999	14,197

		trajectory	30,997	48,905	1002	13,130
	Passive	displacement	30,867	83,381	423	25,837
PL7		trajectory	46,551	102,087	14,766	23,063
	Sinking	displacement	7001	32,193	187	9884
		trajectory	10,981	34,982	730	10,169
	Passive	displacement	74,531	85,229	10,587	15,809
PL8		trajectory	84,021	101,666	56,250	15,290
	Sinking	displacement	14,094	34,730	148	7684
		trajectory	23,806	42,026	667	10,796
	Passive	displacement	73,018	85,353	3150	16943
PL9		trajectory	80,772	98,294	17,932	16,152
	Sinking	displacement	14,295	39,622	1015	8171
		trajectory	24,158	42,542	1016	7825

669

670 A Principal Component Analysis (PCA, **Fig. 8**) explained the variability of the system
671 in two principal dimensions. Dimension 1 accounted for nearly 50% of the variance and
672 was interpreted as wind intensity and water discharge gradient, together with the
673 displacement of the particles, on the left side of the axis, opposite to the plume duration
674 and trajectory. Dimension 2 accounted for 28.5% of the total variance and was interpreted
675 mainly as the displacement gradient. The plume events were separately grouped within
676 the ordination diagram, where PL1, PL3, PL8 and PL9 (red dots in **Fig. 8**), were
677 associated with higher discharge at the mouth and wind intensity. On the other hand, PL2,
678 PL4, PL5, PL6 and PL7 (light blue triangles in **Fig. 8**) were plotted to the left of the
679 diagram and associated to low water discharge plume event duration.



680
 681 *Fig. 8. PCA diagram for the advection and dispersion scenario for all nine simulated plumes. Note that*
 682 *water discharge and wind intensity are negatively associated to the plume duration.*
 683

684 **4. Discussion**

685 *4.1. MP export of the Patos Lagoon coastal plume*

686 *4.1.1. Observed data during the model validation*

687

688 This study introduces for the first-time data on the concentration of MPs exported
 689 from the plume of the world’s largest choked coastal lagoon to the Southwestern Atlantic
 690 Ocean (SWAO). The magnitude of export can be considered as large values, since the
 691 lagoon hold large industrial cities and harbors. The observed accumulation patterns in the
 692 field corroborated the pattern inferred from TrackMPD simulations (**Fig. 2**). Results
 693 illustrated the tendency of the plume’s frontal system to accumulate MPs during the
 694 outward displacement from the estuary. As observed for the higher MP concentration,
 695 this pattern was observed not only in transects 1 and 2 during the June 3rd survey, but also
 696 in transect 3 during the July 21st. The higher concentration of MPs in the Pilot Station of
 697 Rio Grande, observed for transect 4 during the July 21st survey, could be related to the
 698 high traffic of daily vessels, which is also supported by the presence of paint fragments
 699 in the net samples. Fortunately, both sets validate the use of TrackMPD model for MP
 700 surface transport.

701 Previous investigations on the presence/abundance of MPs in this region were
702 conducted across different natural compartments, including the salt marshes of Patos
703 Lagoon (Pinheiro et al., 2021), Cassino Beach (Wetzel et al., 2004; de Ramos et al., 2021),
704 and in offshore areas of the SWAO (Lacerda et al., 2022). However, because of the
705 absence of standardization of procedures due to variations in environmental matrices,
706 sampling methods, and reported result units, comparisons must be interpreted with
707 caution (Pasquier et al., 2022).

708 According to our results, the concentration of MPs observed in surface waters of Patos
709 Lagoon coastal plume ranged from 0.20 items m⁻³ (~59,611 items km⁻²) to 1.37 items m⁻³
710 (~412,294 items km⁻²) for all seven transects. When comparing our findings to those of
711 Lacerda et al. (2022), where they detected 19,267 items km⁻² in the oceanic offshore
712 station, close to our study area, our results were approximately one order of magnitude
713 higher because of the inherent dilution processes of the ocean. Compared to similar
714 studies in the Guaíba River, the main lagoon tributary, Bertoldi et al. (2021) reported
715 average concentrations ranging from 11.9 ± 0.6 to 61.2 ± 6.1 items m⁻³, our results
716 indicate concentrations approximately one order of magnitude lower, as Guaíba River is
717 located very close to a large urban contamination source. This gradient in MP
718 concentrations, from the primary tributary to the open ocean via the estuary, underscores
719 the important role of Patos Lagoon as a buffering system prior to MP export from land to
720 sea (Lopez et al., 2021, Meijer et al., 2021) but also de dilution oceanic process. A recent
721 work from Queiroz et al., (2022) showed comparisons of MP abundance in surface waters
722 of the Western Atlantic Ocean, to assess the MP fluxes from freshwater systems to open
723 ocean in South America (Queiroz et al., 2022). Using a 64-µm mesh, they reported for
724 the Amazon basin a mean abundance of MPs ranging from 2672 ± 1167 items m⁻³ to 4772
725 ± 2761 items m⁻³ during the dry and rainy seasons, respectively. Such a high abundance
726 may be attributable not only to the basin dimension, but also to the methodology used,
727 since almost 50% of the plastics corresponded to fractions smaller than 300 µm (the mesh
728 size used in the present work) (Queiroz et al., 2022). Also from Brazilian waters, in the
729 Goiana River estuary, a mean concentration of 0.26 items m⁻³ (300-µm mesh) was
730 reported (Lima et al., 2014). In the Tamandaré Bay, Lima et al., (2023) reported a range
731 of MP concentration of 7.8 ± 3.5 to 3.7 ± 0.6 items m⁻³ (using a 64-µm mesh plankton
732 net), while in the Tamandaré coastal plume, the authors reported a range of 4.0 ± 0.2 to
733 44.6 ± 55.5 items m⁻³ during low – high rainfall respectively (Lima et al., 2023). The

734 FTIR polymeric characterization also showed consistency with similar regional studies
 735 where PP, and PE are the most abundant type of MPs (Rodriguez et al., 2020, Bertoldi et
 736 al., 2021). These are the most manufactured and lightest plastics, thus they are more easily
 737 found in surface water samples worldwide (Geyer et al., 2017, Zamora et al., 2020).

738

739 *4.1.2. Estimation on export rate.*

740

741 **Table 5** shows the detailed MP export estimations calculated for all nine simulated
 742 plumes, based on this distinction according to the water discharge as identified by the
 743 PCA (**Fig. 8**). July-2022 presented a mean discharge in the same order of magnitude as
 744 the simulated plumes PL1, PL3, PL8 and PL9, grouped as high discharge events by the
 745 PCA (**Fig. 8**). By considering the estimated MP flow for both the lower and higher water
 746 discharge conditions during the June and July 2022 surveys, it would be possible to
 747 extrapolate the number of MPs being exported in each of the nine simulated plume events.
 748 As an approximate estimation, the MP export during the low-moderate water discharge
 749 simulated events ranged from 8.3×10^6 items day⁻¹ to 10.5×10^6 items day⁻¹ (mean: $9.0 \times$
 750 10^6 items day⁻¹), while the export during the high-water discharge simulated events
 751 ranged from 42.4×10^6 items day⁻¹ to 56.6×10^6 items day⁻¹ (mean: 47.5×10^6 items day⁻
 752 ¹). In this sense, the potential export magnitude of simulated plumes encompasses a wide
 753 range between 42 and 297 million of microplastic items per event.

***Table 5.** Estimation of total MP export per day and per plume event based on the observed surface data obtained from the field survey and extrapolated to the mean water discharge of the surface layer and the duration of each simulated/observed plume.*

	Plume event	Duration (days)	Surface water discharge (m ³ s) [*]	MP flux (items day ⁻¹)	Total MP flux
Low discharge events	June – 2022	6	413	7.1 x 10⁶	42.8 x 10⁶
	PL2	6	519	8.9 x 10 ⁶	53.8 x 10 ⁶
	PL4	8	609	10.5 x 10 ⁶	84.2 x 10 ⁶
	PL5	6	516	8.9 x 10 ⁶	53.5 x 10 ⁶

	PL6	5	482	8.3×10^6	41.7×10^6
	PL7	5	491	8.5×10^6	42.5×10^6
High discharge events	July - 2022	5	803	52.0×10^6	260×10^6
	PL1	3	684	44.3×10^6	133×10^6
	PL3	5	717	46.4×10^6	232×10^6
	PL8	7	655	42.4×10^6	297×10^6
	PL9	5	874	56.6×10^6	283×10^6

* Calculated for the surface layer of 700 m²

754

755 The MP export values estimated in the present study for all nine plume events were
756 calculated as a function of the mean water discharge restricted to the surface layer. Due
757 to the inherent water column heterogeneity, the MP concentration estimated here cannot
758 be unfortunately extrapolated to the entire water column. The calculations were
759 consequently limited to the surface layer, confined to of the uppermost 1 m layer. These
760 values ranged from millions to tens of millions of plastics items being exported
761 throughout the coastal plume. However, these values serve as an approximation to the
762 potential MP export of the system only under these conditions. In this context, we
763 assumed that during such discharge events, the surface layer exports MPs at comparable
764 rate during simulations. As also pointed out by Zhao et al., (2019), the extrapolations for
765 the MP fluxes through the estuary mouth must be therefore approached cautiously, by
766 only recognizing them as potential exportation rate.

767 The estimated potential export rate of MPs revealed significant findings. On average,
768 during low to moderate simulated plume events, the export rate was approximately 9.0
769 million items day⁻¹, while during high water discharge simulated plume events, it
770 increased to 47.5 million items day⁻¹. Similarly, Zhao et al., (2019), estimated for the
771 Changjiang Estuary (China) an export rate of 16 - 20 x 10⁸ items yr⁻¹ (utilizing a pump
772 with a 60 µm size methodology). The potentially comparable values between different
773 papers, however, are only comparable in terms of abundance instead of export rates, and

774 abundances were measured using different mesh size (Queiroz et al., 2022). Other recent
775 studies on plastic riverine transport during high river flood conditions focused on the
776 visual sorting of macroplastics for the export calculations, making it difficult to compare
777 to our results (Emmerik et al., 2023, Lima et al., 2023). Our findings underscore the key
778 role of coastal plumes as transport vehicles of a substantial volume of anthropogenic
779 materials from expansive, urbanized basins to coastal and oceanic environments.

780

781 *4.2. Hydrodynamic conditions*

782

783 Despite the temporal co-variation between discharge tributary anomalies of the Patos
784 Lagoon, and the ENSO phases as indicated by the ONI index (Távora et al., 2019,
785 Bitencourt et al., 2020b), the absence of a statistical relationship between the water
786 discharge and the ONI index can be attributed to the inherent complexity of the coastal
787 system. Notably, during the ENSO neutral phase, we observed the most extreme
788 discharge events throughout the jetties' mouth, which surpassed the peaks observed
789 during both weak and very strong El Niño events. In this regard, previous studies on
790 suspended sediment contributions from the Patos Lagoon plume demonstrated that the
791 ENSO cycles did not significantly modify the spatial variability of inner shelf suspended
792 sediments (Lisboa et al., 2022). This could be related to the influence of other mechanisms
793 of ocean-atmosphere interaction exerting hierarchically more important roles than the
794 precipitation anomalies, such as the Southern Annular Mode (SAM), associated with the
795 ENSO phases (Schossler et al., 2018). Moreover, various authors have demonstrated that
796 the Patos Lagoon coastal plume is primarily modulated by the wind influence at synoptic
797 time scale, changing its direction in less than 24 h (Marques et al., 2009) in the alongshore
798 direction (Marques et al., 2010a). In this regard, water discharge over the north portion
799 of the lagoon is related to time scales of around 30 days (Marques et al., 2009) and plays
800 a secondary role in the plume hydrodynamics.

801 The aforementioned evidence not only supports the statistical relationships between
802 water discharge at the mouth and the wind intensity in the estuary mouth, but also
803 corroborates that the wind influence is the main hydrodynamic modulator at this area.
804 Similarly, PCA showed a close relationship between wind and water discharge, grouping
805 together specific plume events under such characteristics (**Fig. 8**). We detected three
806 groups of plumes. PL6 and PL7 were grouped in the middle section of the horizontal axis
807 and represented events of low water discharge but high wind intensity. PL1, PL3, PL8

808 and PL9 represented events of high-water discharge and high wind intensity. Finally,
809 PL2, PL4 and PL5 were characterized by low water discharge and low wind intensity. In
810 addition to synoptic data, climate modes of variability also play an important role as
811 northeasterly winds reported as typical for El Niño events (Marques et al., 2010a, 2011,
812 Bitencourt et al., 2020b), which intensify the plume displacement to the southwest
813 quadrant as observed in this study for PL1 through PL9.

814

815 *4.3. Simulated MP dispersion patterns under coastal plume dynamics*

816

817 *4.3.1. Considerations for advection-dispersion scenarios*

818

819 In response to the influence of northeasterly winds, the primary direction of particle
820 dispersion of all plume events was towards the southwest of the domain. However, when
821 comparing the particles displacements and trajectories between plume events for
822 advection and dispersion scenarios, we observed differences in terms of main physical
823 variables.

824 The particle displacements (trajectories) exhibited a positive (negative) correlation
825 with water discharge and wind intensity. In plume events of high-water discharge, such
826 as PL1, PL3, PL8, and PL9, the particles reached longer displacements quickly, traveling
827 shorter but straight trajectories. These events showed a high percentage of particles out
828 of the domain (80%, 98%, 100%, 99% respectively). The dispersion pattern in this group
829 of events followed a direct southwest path, being the particles trapped in an alongshore
830 circulation. Areas of highest particle density were observed along the coast, thus
831 delineating the front of the saline plume, and within the gyre between the west jetties and
832 Cassino beach. During such transport process, the particles formed a loop in this gyre.
833 These areas of accumulation located to the south of the jetties mouth were previously
834 reported by Marques et al. (2009; 2010b) and Fernandes et al. (2021) as recirculation
835 zones where vertical velocities are intensified, and sediment deposition is promoted.
836 Hence, it appears to resemble the same pattern as observed in the vertical particle
837 trajectories where they converge downward while performing this loop, supporting the
838 occurrence of a potential depocenter of these anthropogenic particles.

839 In contrast, the other group of plume events exhibited lower discharge values. Despite
840 this fact, PL6 and PL7 also experienced strong wind intensities, resulting in particle

841 transport patterns comparable to those of high discharge events. Such events displayed
842 high-density areas in the plume front and within the gyre. In the case of PL7, there was
843 also a high-density region to the northeast behind the jetties.

844 On the other hand, PL2, PL4, and PL5 shared lower values of water discharge and
845 wind intensities and displayed lower percentages of particles being transported out of
846 domain (0%, 25%, and 83% respectively). This is evident by the longer particle
847 trajectories during the simulation periods, performing more offshore transport patterns.
848 In this sense, PL2 showed a hotspot of accumulation at the farthest point of the inner
849 shelf, being susceptible to a wider redistribution by oceanic currents. Both PL5 and PL7,
850 exhibited the accumulation pattern behind the jetties front of São José do Norte.

851 4.3.2. Study limitations: considerations for sinking scenarios

852 As the simulated vertical behavior of MPs within the plume was based in only one
853 W_s value, there is a limitation for inferring all the MP potential scenarios in the study
854 area considering the variability of this parameter over different water conditions and
855 biofilm growth (Mendrik et al., 2023). Also, this validation process based on surface
856 sampling conducted and reported also as a validation method elsewhere (Atwood et al.,
857 2019, Gorman et al., 2020, Lopez et al., 2021), does not allow to properly validate the
858 vertical transport of MPs. *Ad hoc* field surveys with this specific objective should be
859 conducted to provide a comprehensive understanding of the vertical behavior of MPs on
860 coastal plumes.

861 Another limitation arises from the use of a constant value of K_v , knowing that this
862 can also change over time and space. TrackMPD can read spatio-temporal values of K_v
863 from TELEMAC. In that case if K_v is not well estimated in TELEMAC, it can be a
864 limitation for TrackMPD.

865 To enhance our understanding of this complex issue, future studies should build upon
866 this work by exploring different periods, considering a range of climatic conditions, and
867 examining materials with varying characteristics, including differences in density and
868 shape. For more accurate calibration of numerical models, additional field experiments
869 are essential to evaluate in-situ characteristics and parameters such as specific coefficients
870 of diffusion, settling velocities and the evolution of plastic density in real estuarine
871 conditions. Moreover, more field surveys should be conducted, specifically targeting
872 different gradients and accumulation hotspots, and determining the MP concentration in

873 the whole water column. This comprehensive approach will contribute significantly to
874 the scientific understanding of the multifaceted nature of this problem.

875 *4.4. Ecological consequences for observed and simulated findings*

876 The inferred patches where MPs accumulated represent areas of elevated ecological
877 risk. These areas are characterized by the coexistence of diverse groups of organisms with
878 both natural and human significance (Lemos et al., 2022), and they might potentially
879 interact with the accumulated MPs.

880 The increased accumulation observed in the plume fronts, located between the
881 freshwater-marine interface, is a recurring phenomenon observed for most simulations.
882 These areas are prone to high levels of instability, driven by shearing stress (Pritchard,
883 2000), which triggers mixing processes contributing to the degradation and fragmentation
884 of MPs (Wang et al., 2022). Moreover, these frontal regions are impacted by the
885 accumulation of MPs, which can disrupt nutrient capture and biogeochemical cycles
886 associated with frontal systems (Galgani et al., 2021; Wang et al., 2022). This suggests
887 that numerous processes are likely increasing the bioavailability of MPs to a wide range
888 of organisms in the food webs (Wang et al., 2022), specially during high discharge events
889 (Lima et al., 2023). Lima et al., (2023) reported MP intense inputs from river to coastal
890 region, during high discharge plume periods, where almost 70% of them ranged in the
891 same size of zooplankton communities, thus demonstrating the elevated risk for
892 zooplanktivorous organisms to ingest MPs for food (Lima et al., 2023). Supporting this,
893 previous research in the study area has reported eight marine fish species, highly valued
894 for commercial purposes, having ingested MPs (Neto et al., 2020). Similar findings have
895 been documented for five species of sea turtles (Rizzi et al., 2019; Bugoni et al., 2001),
896 procellariiform birds (Colabuono et al., 2009), fur seals (Denuncio et al., 2017; Lima et
897 al., 2023), and sperm whales (Brentano et al., 2020). These relevant reports, together with
898 our findings, denote the importance of this underestimated problem for the environmental
899 quality and health of the marine ecosystem influenced by Patos Lagoon coastal plume.

900 **5. Conclusion**

901 The present study represents the first attempt to understand the presence, abundance,
902 and distribution of MPs in the surface waters off Patos Lagoon estuary, which is set in a
903 large urban conglomerate holding one of the largest harbors of Brazil. This is also the
904 first report on coastal plume potential export rate of MPs and the dispersal mechanisms

905 influenced by hydrodynamics. The potential utility of the TrackMPD model as a tool for
906 assessing and mitigating this anthropogenic pressure on the local environment under
907 varying ambient conditions is explored. The combination of numerical modeling tools
908 with field observations, was shown to be invaluable and cost-effective for simulating a
909 wide range of natural conditions and identifying microplastic hotspots in the area.

910 The coastal plume showed surface MP concentrations ranging from 0.20 to 1.37 items
911 m^{-3} , thus identifying this region of South America as a hot spot of transport and export of
912 MP to the inner shelf, potentially exporting values of 9.0 - 47.5 million items day^{-1} . This
913 conclusion is supported by the condition that the mouth of the estuary is as permanently-
914 open system. There is a strong relationship between the discharge flow and the wind
915 intensity as main hydrodynamic modulator for dispersion of the plume.

916

917 **6. Credit authorship contribution statement**

918 **Carolina Rodriguez Perez:** Conceptualization, Methodology, Formal Analysis,
919 Investigation, Data Curation, Writing – original draft, Writing – review and editing,
920 Visualization. **Pablo Silva:** Data Curation, Methodology, Visualization. **Laura Moreira:**
921 Methodology. **Larissa Zacher:** Methodology. **Andreia Fernandes:** Resources, review,
922 and editing. **Remi Bouyssou:** Methodology, Visualization. **Isabel Jalón-Rojas:**
923 Methodology, Visualization, Resources, review, and editing. **Osmar Moller:**
924 Methodology, Funding acquisition, Resources. **Felipe Garcia-Rodriguez:**
925 Conceptualization, Formal analysis, Writing – review and editing, Supervision. **Grasiela**
926 **Lopes Leães Pinho:** Conceptualization, Formal analysis, Resources, Writing – review
927 and editing, Supervision. **Elisa Fernandes:** Conceptualization, Formal analysis,
928 Resources, Writing – review and editing, Supervision

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934 **Declaration of competing interest**

935 The authors declare that they have no known competing financial interests or personal
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947 **Supplementary data**

948 Supplementary data to this article can be found online at ..

949 **8. References**

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