

# The effect of stress barriers on unconventional-singularity-driven frictional rupture

Barnaby Fryer, Mathias Lebihain, Corentin Noël, Federica Paglialunga,

François Passelègue

## ► To cite this version:

Barnaby Fryer, Mathias Lebihain, Corentin Noël, Federica Paglialunga, François Passelègue. The effect of stress barriers on unconventional-singularity-driven frictional rupture. Journal of the Mechanics and Physics of Solids, 2024, 193, pp.105876. 10.1016/j.jmps.2024.105876 . hal-04711620

## HAL Id: hal-04711620 https://hal.science/hal-04711620v1

Submitted on 27 Sep 2024

**HAL** is a multi-disciplinary open access archive for the deposit and dissemination of scientific research documents, whether they are published or not. The documents may come from teaching and research institutions in France or abroad, or from public or private research centers. L'archive ouverte pluridisciplinaire **HAL**, est destinée au dépôt et à la diffusion de documents scientifiques de niveau recherche, publiés ou non, émanant des établissements d'enseignement et de recherche français ou étrangers, des laboratoires publics ou privés.

## Highlights

#### <sup>2</sup> The effect of stress barriers on unconventional-singularity-driven frictional rupture

- <sup>3</sup> Barnaby Fryer, Mathias Lebihain, Corentin Noël, Federica Paglialunga, François Passelègue
- Stress barriers promote rupture arrest by decreasing energy flux to the crack tip
- Long-tailed frictional weakening must be included in LEFM to predict crack arrest
- This long-tailed weakening depends on the in-situ stress conditions
- Fault preconditioning is experimentally demonstrated to contain dynamic events

# The effect of stress barriers on unconventional-singularity-driven frictional rupture

Barnaby Fryer<sup>a</sup>, Mathias Lebihain<sup>b</sup>, Corentin Noël<sup>a</sup>, Federica Paglialunga<sup>a</sup>, François Passelègue<sup>a</sup>

<sup>a</sup>Université Côte d'Azur, CNRS, Observatoire de la Côte d'Azur, IRD, Géoazur, Valbonne, France

<sup>b</sup>Laboratoire Navier, CNRS (UMR 8205), École des Ponts ParisTech, Université Gustave Eiffel, Champs-sur-Marne,
 France

#### 15 Abstract

10

11

12

Whether or not energy dissipation is localized in the vicinity of the rupture tip, and whether any distal energy dissipation far from the crack tip has a significant influence on rupture dynamics are key questions in the description of frictional ruptures, in particular regarding the application of Linear Elastic Fracture Mechanics (LEFM) to earthquakes. These questions are investigated experimentally using a 40-cm-long experimental frictional interface. Three independent pistons apply a normal load with a fourth piston applying a shear load, enabling the application of a heterogeneous stress state and stress barriers. After loading the frictional interface to a nearcritical state, subsequent unloading of one normal-load piston leads to dynamic ruptures which propagate into the heterogeneous stress fields. The ruptures in these experiments are found to be driven by unconventional singularities, characterised by an ever-increasing breakdown work with slip, and as a result do not conform to the assumptions of LEFM. As these experimental stress barriers inhibit slip, they therefore also reduce the breakdown work occurring outside of the cohesive zone. It is shown that this distal weakening, far from the crack tip, must be considered for the accurate prediction of rupture arrest length. These experiments are performed in the context of a proposed stimulation technique for Enhanced Geothermal Systems (EGSs). It has previously been suggested, through theoretical arguments, that stress barriers could be induced through the manipulation of pore pressure such that there is reduced seismic hazard during the shear stimulation of EGSs. This stimulation technique, known as preconditioning, is demonstrated here to reduce the mechanical energy flux to the crack tip, G, while also increasing the fracture energy,  $G_c$ . Preconditioning is shown to be capable of arresting seismic rupture and reducing co-seismic slip, slip velocity, and seismic moment at preconditioning stresses which are reasonably achievable in the field. Due to the fully-coupled nature of seismic rupture and fault slip, preconditioning also reduces distal weakening and its contribution to the propagation of induced seismic ruptures. In a similar vein, heterogeneous pore pressure fields associated with some seismic swarms can be used to explain changes in stress drop within the swarm without recourse to material or total-stress heterogeneity.

*Keywords:* Frictional ruptures, Stress barriers, Rupture dynamics, Unconventional singularity,
 Preconditioning

September 27, 2024

<sup>\*</sup>Corresponding author: barnaby.fryer@geoazur.unice.fr Preprint submitted to Journal of the Mechanics and Physics of Solids

#### 18 1. Introduction

The injection of fluid into the subsurface has been associated with induced seismicity on a 19 significant number of occasions. These injection operations have had a variety of purposes, such 20 as fluid disposal/storage, improved hydrocarbon recovery (e.g., Raleigh et al. (1976)), and reser-21 voir stimulation (e.g., Bao and Eaton (2016)). In the case of reservoir stimulation for Enhanced 22 Geothermal Systems (EGSs), the mechanical stimulation of a fault or fracture zone is, in the-23 ory, achieved through the induced shear dilation of the targeted discontinuity (e.g., Lee and Cho 24 (2002)). The shearing itself is induced through the increase in pore pressure associated with fluid 25 injection, which acts to reduce the fault's effective normal stress and initiate shear failure. Ideally, 26 this process occurs aseismically, such that seismic waves are not radiated. However, as injection 27 continues, the rupture front bounding the stimulated region grows, potentially surpassing the area 28 of increased pore pressure if the fault is critically stressed (Garagash and Germanovich, 2012; 29 Bhattacharya and Viesca, 2019; Cebry and McLaskey, 2021). Eventually, this rupture front may 30 reach a critical length, known as the nucleation length (Ida, 1972; Okubo and Dieterich, 1984; 31 Campillo and Ionescu, 1997; Uenishi and Rice, 2003), and begin accelerating, resulting in the 32 33 dynamic rupture of the fault and the radiation of potentially-damaging seismic waves. The understanding of how to achieve shear dilation in an aseismic manner while avoiding the dynamic 34 rupture of the fault is an obstacle for the global development of EGSs, and a number of field-scale 35 EGS projects have been impeded or halted due to the seismicity they have induced (e.g., Basel, 36 Switzerland (Häring et al., 2008); Pohang, South Korea (Kim et al., 2018)). 37

The success or failure of an EGS operation is strongly dependent on the state of stress in the 38 subsurface. This is because the state of stress not only influences the susceptibility of a given fault 39 to shear dilation (e.g., Lee and Cho (2002)), but also affects the nucleation length (Okubo and 40 Dieterich, 1984), rupture velocity (Ben-David et al., 2010), slip velocity (Okubo and Dieterich, 41 1984), rupture halting (Husseini et al., 1975), and radiated seismic energy (Scholz, 1968) of 42 earthquakes. In particular, recent experimental and theoretical studies have highlighted that the 43 rupture length, i.e., the size of the earthquake that will eventually be induced, is contingent 44 upon the stress distribution along the fault (Bayart et al., 2016; Galis et al., 2017). Indeed, 45 rupture arrest is described by the flux of total potential energy per unit extension of the crack 46 tip (commonly termed G) and the dissipated energy related to the unit extension of the crack 47 tip (commonly termed  $G_c$  or fracture energy); both of these terms depend on the stress state. It 48 is generally assumed that weakening is localized near the crack tip, such that, while slip may 49 continue, there is no further weakening once slip has reached a critical slip distance, allowing 50 the stress and strain fields around the crack tip to be described by a conventional singularity. 51 However, recently, instances of continued power-law weakening beyond this critical slip distance 52 have been observed experimentally, resulting in unconventional singular stress and strain fields 53 (Paglialunga et al., 2024). This raises the question as to whether or not the energy release rate 54 and, therefore rupture arrest, are truly governed exclusively by near-tip weakening (Garagash, 55 2021; Brener and Bouchbinder, 2021b; Kammer et al., 2024). It is further unclear what effect, if 56 any, stress barriers and the stress state in general have on distal weakening and if a reduction in 57 distal weakening promotes rupture arrest. 58

Engineering operations are capable of influencing the state of stress in the Earth's crust. For example, fluid production was first connected to land subsidence nearly a century ago (Pratt and Johnson, 1926), and fluid injection was linked to seismicity in the 1960's and 1970's (Raleigh et al., 1976). Concurrently, in the oil and gas industry, operators began actively manipulating the *in-situ* stress with hydraulic fracturing operations, first employed in the 1940's (Clark, 1949), which rely on the reduction of the minimum principal effective stress through the increase of
 fluid pressure (Hubbert and Willis, 1957). Later, operators began recognizing that the *in-situ* stress could be manipulated prior to initiating a hydraulic fracture for that fracture's eventual
 benefit (Shuck, 1977).

Similarly, the mining industry has been employing a de-stressing technique referred to as 68 preconditioning since the 1950's (Roux et al., 1957). In this operation, a rock face is artificially 69 fractured in order to relieve stress and reduce the occurrence and severity of rockbursts. Still in 70 its relative infancy, the EGS industry has not sufficiently developed methods to reduce the hazard 71 72 of induced seismic events related to stimulation operations. However, the potential to direct and inhibit stimulation treatments by manipulating pore pressure has been recognized for two 73 decades already (Baria et al., 2004), and methodologies to inhibit the occurrence of large seismic 74 events such as cyclic stimulation/fatigue hydraulic fracturing (Zimmermann et al., 2010; Zang 75 et al., 2013; Hofmann et al., 2019; Noël et al., 2019), control theory (Stefanou and Tzortzopoulos, 76 2022), fracture caging (Frash et al., 2021), and preconditioning (Fryer et al., 2020, 2023; Jalali 77 et al., 2023) have been suggested, developed, and, in the case of cyclic stimulation, even deployed 78 at the field scale. 79

Here, using the effective stress principle, the preconditioning of an EGS through the de-80 velopment of stress barriers (Fryer et al., 2023) is demonstrated experimentally using a biaxial 81 apparatus. In practice these stress barriers are proposed to be achieved through an extended (but 82 short compared to the background stressing rate) period of production, reducing pore pressure 83 and increasing effective stress, prior to a comparatively-shorter period of injection, stimulating 84 the fault by inducing slip, Figure 1(a-c). Here, this procedure will be simulated in dry conditions 85 with changes in total normal stress, Figure 1(d,e). It is shown that preconditioning has the ca-86 pability to halt nucleated dynamic ruptures, reducing the hazard associated with fluid injection 87 into a fault. Preconditioning achieves this moderation in hazard through the abatement of energy 88 available to the propagating rupture via a diminished stress drop and the increase in the energy 89 required to continue dynamic propagation via an increased fracture energy; in accordance with 90 previous findings and predictions (Husseini et al., 1975; Freund, 1990; Kammer et al., 2015; 91 Bayart et al., 2016; Galis et al., 2017; Bayart et al., 2018; Gvirtzman and Fineberg, 2021; Cebry 92 et al., 2022; Fryer et al., 2023; Barras et al., 2023). Further, it will be shown that distal weaken-93 ing, far from the crack tip, exhibits a first-order control on rupture arrest and is also reduced by 94 preconditioning and stress barriers in general. 95

#### 96 2. Methods

#### 97 2.1. Experimental Setup

The biaxial apparatus, CrackDyn, located at Géoazur in Valbonne, France, was used to per-98 form experiments on two polymethylmethacrylate (PMMA) rectangular-prism blocks ( $40 \times 10$ 99  $\times$  1 cm and 45  $\times$  10  $\times$  1.8 cm), yielding a fault length, L, of 40 cm and a fault width of 1 cm, 100 Figure 1(f). The position of any point along the sample is given by its horizontal, x, and vertical, 101 y, coordinates. The apparatus engages with the sample blocks via three independent vertical pis-102 tons, responsible for applying normal forces, and one horizontal piston, responsible for applying 103 a shearing force. The pistons are driven with two ENERPAC P141 hydraulic pumps and transmit 104 load to the sample blocks via steel sample holders. The sample holders transmitting the normal 105 stress to the fault system are discontinuous, allowing for a heterogeneous normal load to be ap-106 plied along the fault length, Figure 1f. The force applied by each piston is recorded at 500 Hz by 107

a Scaime K13 load cell located in between it and the sample holder. Thirteen 350-Ohm strain-108 gauge rosettes (39 total strain gauges), recording at 2 MHz, are glued at 3 mm from the simulated 109 fault interface. The strain gauges' signals are amplified by a factor of 100 by Elsys SGA-2 MK2 110 amplifiers. Twenty Brüel & Kjær type 8309 accelerometers are glued either horizontally or verti-111 cally at approximately 1 cm from the simulated fault interface and record at 2 MHz with a cutoff 112 frequency of 54 kHz. Finally, three high-intensity light sources emit light which traverses the 113 sample to arrive at a Phantom TMX 6410 high-speed camera. The light is cross-polarized with 114 two linear polarizing filters. During the experiments, the camera was triggered using a piezo-115 electric sensor and recorded frames at 500 kHz (1.7  $\mu$ sec exposure time with a 1280 $\times$ 32 pixel 116 resolution across the fault length, i.e., 350 µm per pixel). The high-speed camera and the use of 117 polarized light with a birefringent material allow for the use of photoelasticity, which enables the 118 tracking any propagating dynamic ruptures (Rosakis et al., 1999; Nielsen et al., 2010; Schubnel 119 et al., 2011; Latour et al., 2013; Gounon et al., 2022; Paglialunga et al., 2023). The piezoelectric 120 sensor signal used for the camera triggering was split between the recording computers and used 121 to synchronize the different recording systems. The data treatment is addressed in the appendix 122 A. 123

#### 124 2.2. Experimental Approach

The sample is loaded in steps, Figure 1(d,e), by increasing the normal stress and shear stress 125 applied at the pistons in increments of 30 bar. After initial loading, the normal stresses applied 126 by all three vertical pistons are equivalent and set to either 60, 90, or 120 bar nominal normal 127 stress,  $\sigma^0$ , i.e., as read on the pump's analogue gauge in the hydraulic pressure lines. The shear 128 stress is set to a value just below the shear stress required to initiate a dynamic event, in this 129 case corresponding to 100, 145, or 190 bar nominal shear stress. The volume at the shear pump 130 was then maintained constant for the rest of the experiment. At this stage the sample fault is 131 considered to be loaded and representative of a tectonically-loaded natural fault, Figure 1a. If the 132 sample is to be preconditioned, the normal stress applied by all three vertical pistons is increased 133 by a nominal normal stress of  $\Delta \sigma^{\rm p}$ , such that  $\frac{\Delta \sigma^{\rm p}}{\sigma^{\rm 0}}$  is equivalent to approximately 0.08, 0.16, or 0.24, Figure 1(b,d). In practice, if the normal total stress on a fault is given by a lithostatic gradient of 23  $\frac{MPa}{km}$  and the pore pressure by a hydrostatic gradient of 10  $\frac{MPa}{km}$ , these values of  $\frac{\Delta \sigma^{\rm p}}{\sigma^{\rm 0}}$  correspond to producing fluid from the fault such that the effective hydrostatic gradient on the fault is reduced to approximately 9, 8, or 7  $\frac{MPa}{km}$ . At each normal stress a base case is also tested twice whereby the stress is not present and  $\Delta \sigma^{\rm p} = 0$ , corriging as reference. 134 135 136 137 138 twice whereby the stress is not preconditioned and  $\frac{\Delta\sigma^p}{\sigma^0} = 0$ , serving as reference. Injection is then simulated by reducing the stress applied by the right-hand-side vertical piston (closest to 139 140 the horizontal piston), leaving the other two vertical pistons applying the preconditioned stress, 141 Figure 1(d-f). The unloading generally results in an audible acoustic event, after which the shear 142 and then normal stresses are unloaded and the sample reset. 143

#### 144 **3. Experimental Results**

The local stress recorded by the strain gauge rosettes show that the stress profiles are highly reproducible between experiments, Figure 2. The strain gauges show that preconditioning increases the normal stress,  $\sigma_{yy}$ , with smaller changes to shear stress,  $\sigma_{xy}$ . Strain gauge measurements were further able to provide stress profiles just before and just after the nucleation of the principal dynamic event associated with each experiment, Figure 2.

The stress-induced birefringence of the PMMA samples allows for photoelasticity measurements, which provide a second independent method to track the propagating ruptures. When



Figure 1: An overview of preconditioning and the experimental procedure. (a-c) illustrates preconditioning as it could be applied at the field scale, with the process simulated in the laboratory here under dry conditions. (a) Tectonic loading results in a critically-stressed fault at depth. (b) A well is drilled into the fault. Fluid is produced from the fault, reducing the pore pressure and increasing the effective normal stress. (c) Fluid is injected into the fault on a time scale significantly shorter than the production phase in (b). The increased pore pressure locally reduces the effective stress and initiates slip. This slipping or rupture front may outpace the fluid pressure front, but will encounter the low-pore-pressure region, which acts as a fracture-energy and reduced-shear-stress-drop barrier (Fryer et al., 2023). In the context of this figure, t is time,  $t_0$  is the start of operations,  $t_p$  is the time when fluid production stops, and  $t_s$  is the time of shut-in. (d,e) The stress profiles of two example experiments, illustrating the left-most vertical load (green), central vertical load (purple), right-most vertical load (red), and shear load (blue). Initially, the normal and shear loads are increased, simulating tectonic loading (a). If preconditioning is applied (b,d), the vertical loads are increased. In the case preconditioning is not applied, this step is skipped (e). The right-most vertical load is then reduced, simulating fluid injection. This results in the nucleation of a dynamic rupture. Note the macroscopic stress drop present in the non-preconditioned case (e). This macroscopic stress drop is not present in the preconditioned case (d) and is the first piece of evidence suggesting that the preconditioned case halted rupture propagation, containing the event, whereas the non-preconditioned case ruptured the entire experimental fault, reaching the sample edges and resulting in reduced stiffness and extra slip (e.g., Kilgore et al. (2017)). (f) The experimental setup with the inset showing the camera setup. (A) high-speed camera, (B) biaxial setup, (C) linear polarizer, (D) light source, (E) stopper, (F) load cell, (G) vertical piston, (H) sample holder, (I) strain gauge rosette, (J) acoustic sensor (trigger), (K) PMMA sample, (L) sample-sample interface, (M) displacement sensor (only in Supplementary Material), (N) accelerometer, and (O) horizontal piston.



Figure 2: Stress profiles deduced from strain gauge measurements for two experiments performed at a nominal normal stress of 120 bar. (a,c,e) No preconditioning is performed. (b,d,f) A preconditioning of 30 bar is applied. (a,b) The stress profile in time found from one strain gauge located at a position near the sample edge (x = 40 cm), in the area where the sample is unloaded. The initial normal (green) and shear (red) stresses, representative of the *in-situ* stress resulting from tectonic loading have a yellow background. The unloading phase has a red background. In (b) the preconditioning phase has a green background. Note both experiments have a foreshock during the unloading phase which is contained on one side of the sample. The normal (c,d) and shear (e,f) stress profiles across the entire surface of the sample, using all working strain gauges at specific moments: after "tectonic" loading (circle), after preconditioning (triangle), during unloading but prior to the dynamic event (square), and after the dynamic event (star). Markers represent values found from the strain gauges. Lines are linear interpolations. Note the similarities in the initial stress profiles (circles) between the two cases. Additionally, note the stress increase in (d,f) related to preconditioning (triangles). Finally, note the reduced shear stress in (e) after the event across the entirety of the sample (stars). The shear stress is increased in the preconditioned case (f). The case without preconditioning resulted in the rupturing of the entire sample; the preconditioned case had a contained event.



Figure 3: Three experiments are shown, all performed at a nominal normal stress of 120 bar. (a,d) No preconditioning is performed. (b,c,e,f) Preconditioning is performed up to a nominal normal stress of (b,e) 130 and (c,f) 140 bar. (a-c) The videogram profiles illustrate the progression of the rupture fronts and are overlain by the local displacement recorded by the horizontally-oriented accelerometers (black lines). The displacements are zeroed to the location of the sensor and deviation from this position indicates local displacement. The scale for the displacement is found on the right. The rupture front is traced with a dotted line. Note that certain spatial bands where the grey-scale remains uniform on the videograms correspond to areas where light was not able to pass through the sample at the height of the fault. This is due to it being blocked by a sensor. (d-f) The slip profiles as calculated from the accelerometers positioned along the fault interface. An initial time, here denoted by the color bar as 0.00 microseconds, is chosen just before slip initiates. Then, four additional evenly-spaced times are taken starting from this moment at 0.08, 0.16, 0.24, and 0.32 microseconds. The "final" slip profile after 1.2 microseconds is shown in black. The positions of the accelerometers are denoted by white circles, with linear interpolation between these points.

these measurements are taken in combination with the slip profiles, computed from the ac-152 celerometers, the rupture fronts of the dynamic events can be traced, Figure 3(a-c). Measure-153 ments of slip were used to build slip profiles across the fault interface, Figure 3(d-f). The slip 154 profiles show that the larger the magnitude of the preconditioning, the smaller the total cumu-155 lative slip and slip velocities are. Note finally that the characteristic rise times (slip durations) 156 computed by the accelerometers are similar to the characteristic source durations (the time re-157 quired for the rupture to traverse the sample) observed on the videograms, meaning that these 158 ruptures exhibit crack-like, as opposed to pulse-like, behaviour. 159

Preconditioning results in halted and/or slowed rupture and reduced slip, slip velocity, stress 160 drop, and moment magnitude, Figures 3 and 4. The videogram measurements were manually 161 traced, allowing for the creation of rupture profiles for the dynamic events associated with each 162 experiment, Figure 4(a-c). In all cases where preconditioning was not applied, the rupture was 163 able to traverse the entire interface, highlighting that, while the stress state along the interface is 164 heterogeneous (all ruptures experienced deceleration at approximately x = 15 cm, presumably 165 associated with a higher normal stress at this position, Figure 2(c,d)), on the whole the tested 166 conditions can be considered representative of a critically-stressed fault. However, with the exception of one experiment ( $\sigma^0 = 120$  bar;  $\frac{\Delta \sigma^P}{\sigma^0} = 0.08$ ), preconditioning was able to halt the 167 168 dynamically-propagating rupture before it reached the sample edges at all tested nominal normal 169

stresses ( $\sigma^0 = 60, 90, \text{ and } 120 \text{ bar}$ ) and all tested levels of preconditioning ( $\frac{\Delta \sigma^P}{\sigma^0} = 0.08, 0.16$ , 170 (0.24). This result is further confirmed by the slip profiles, Figure 4(d-f). Generally, although 171 the results are influenced by whether or not the rupture reaches the free boundary at the sample 172 edges (Kilgore et al., 2017), preconditioning results in reduced total slip behind the crack tip 173 and reduced slip velocity, Figure 4(d-i). In fact, the accelerometers provide a near-continuous 174 measure of slip along the interface that can be used to estimate the seismic moment,  $M_0$ , of each 175 event as  $M_0 = \mu A_s D$  (Aki, 1966), where  $\mu$  is the dynamic shear modulus, D is the slip, and  $A_s$ 176 the area of the ruptured region along the interface. Without preconditioning, experiments with 177 nominal normal stresses of 60, 90, and 120 bar yield average seismic moments of 416, 940, and 178 1608 N·m, respectively. With a preconditioning of  $\frac{\Delta \sigma^{P}}{\sigma^{0}} = 0.24$ , experiments with nominal normal stresses of 60, 90, and 120 bar yield seismic moments of 86, 474, and 495 N·m, respectively, 179 180 corresponding to a 50 to 79% reduction in seismic moment due to preconditioning. Finally, 181 strain gauge measurements from before and after the passing of the rupture front allow for the 182 calculation of a continuous stress drop profile, Figure 4(j-l). Preconditioning results in smaller 183 stress drops behind the rupture front, as well as negative stress drops (increases in shear stress) 184 in the cases that it is able to halt the propagating rupture. The stress drop profile,  $\Delta \sigma_{xy}(x)$ , will 185 later be seen to be of importance for the energy flux reaching the crack tip of the propagating 186 rupture. 187

#### 188 4. Discussion

#### 189 4.1. Fault preconditioning for reduced hazard during EGS stimulation

<sup>190</sup> During crack propagation, the mechanical energy release rate, or the flux of total potential <sup>191</sup> energy per unit extension of a crack's tip, G, is equivalent to the dissipated energy related to that <sup>192</sup> same unit extension, which is known as the fracture energy,  $G_c$ , such that (Griffith, 1921; Freund, <sup>193</sup> 1990),

$$G = G_{\rm c}.$$
 (1)

<sup>194</sup> The arrest of a dynamically-propagating shear crack can occur due to an increase in  $G_c$  or a <sup>195</sup> reduction in G, such that G falls below  $G_c$  (Husseini et al. (1975); Rice (1980), page 594; Freund <sup>196</sup> (1990), Eqn. 7.4.27; Kammer et al. (2015); Bayart et al. (2016)).

In this context, fault preconditioning results in a reduction in the energy available to a propa-197 gating rupture and in a fracture energy barrier (Fryer et al., 2023). These predictions are based on 198 and in agreement with Linear Elastic Fracture Mechanics (LEFM) theory (Husseini et al., 1975; 199 Rice, 1980; Freund, 1990; Kammer et al., 2015; Bayart et al., 2016; Galis et al., 2017; Bayart 200 et al., 2018; Gvirtzman and Fineberg, 2021; Cebry et al., 2022; Paglialunga et al., 2022). Here, 201 the possible effects of preconditioning have been explored experimentally, demonstrating gen-202 eral agreement with previously-made predictions. In particular, preconditioning has been shown 203 to be capable of slowing and/or halting a dynamically propagating rupture as well as reducing 204 the slip and slip velocity behind the dynamically-propagating crack tip, resulting in a smaller 205 seismic moment for nucleated events. This effect of preconditioning on rupture propagation can 206 be principally the result of (i) the increase in the energy required to continue crack propagation 207 (the fracture energy,  $G_c$ ), (ii) the reduction of the mechanical energy flux at the crack tip (G), or 208 (iii) a combination of both. 209



Figure 4: (a-c) The crack tip progression for each experiment, manually picked from videograms. Dots represent manually-picked points; linear interpolation is used between these points. (d-f) The total (maximum) slip calculated from the accelerometers, calculated up to 1.5 msec after the trigger. (g-i) The maximum slip velocity up to 1.5 msec after the trigger. For (d-i), white dots represent the locations of accelerometers; linear interpolation is used between these locations. (j-l) The stress drop resulting from the dynamic event calculated using the strain gauges. White dots represent the locations of strain gauge rosettes; linear interpolation is used between these locations. Stress drops are only plotted for positions behind the crack tip; however note that the strain gauge at 0.5 cm was not working for the experiments at 120 bar. Experiments are performed at (a,d,g,j) 60, (b,e,h,k) 90, and (c,f,i,l) 120 bar nominal normal stress. For all plots, the color bar indicates the amount of preconditioning used in each experiment.

#### $_{210}$ 4.1.1. Influence of fault preconditioning on $G_{c}$

The dependence of  $G_c$  on normal stress was examined through two calibration experiments, both exhibiting complete ruptures. The determination of  $G_c$  began with the calculation of the breakdown work,  $w_b$ , (Tinti et al., 2005; Brener and Bouchbinder, 2021b,a),

$$w_{\rm b}(D) = \int_0^D \left( \sigma_{\rm xy}(\delta) - \sigma_{\rm xy}(D) \right) \mathrm{d}\delta, \tag{2}$$

where  $d\delta$  is an increment of slip. Note that the definition of breakdown work here is slightly modified from that introduced by Tinti et al. (2005) to follow Brener and Bouchbinder (2021b,a); Paglialunga et al. (2022, 2024), and is precisely the definition of fracture energy in Abercrombie and Rice (2005), stemming from Palmer and Rice (1973). The breakdown work was determined for multiple events at various strain gauge rosettes.

Initially, once the rupture tip first passes a specific location along the fault and slip initiates, 219 breakdown work scales as  $w_{\rm b} \propto D^2$  (Tinti et al., 2005; Brener and Bouchbinder, 2021b); however, 220 this phase is generally not seen here to due the frequency response of the accelerometers. At 221 larger values of slip, the relationship between  $w_b$  and D undergoes a change, such that  $w_b \propto D^m$ , 222 223 where *m* is a fitting parameter. The slip which corresponds to this change in dependence is known as the critical slip distance,  $D_c$ , and corresponds to a cross-over slip scale. Within the framework 224 of LEFM, weakening is localized in an infinitesimal region near the crack tip, so that no further 225 weakening occurs when  $D \ge D_c$ . This results in  $w_b \propto D^0$  and corresponds to a singularity 226 order,  $\xi$ , of -0.50 (e.g., Svetlizky et al. (2020); Shlomai et al. (2021); Brener and Bouchbinder 227 (2021b)). The breakdown work achieved at  $D = D_c$  is equivalent to the fracture energy, such that 228  $G_{\rm c} = w_{\rm b} \left( D = D_{\rm c} \right).$ 229

In these experiments, breakdown work is generally a continuously-increasing function of slip 230 even when  $D \ge D_c$ , and can be fit as  $w_b (D \ge D_c) = AD^m$ , Figure 5b and appendix B, where A is 231 a fitting parameter. There is therefore a cross-over behaviour occurring at  $D = D_c$  from a small-232 slip behaviour where  $w_b \propto D^2$  to a large-slip behaviour where  $w_b \propto D^m$ , as described by previous 233 authors (Viesca and Garagash, 2015; Brantut and Viesca, 2017; Brener and Bouchbinder, 2021b; 234 Paglialunga et al., 2024). The latter results in unconventional singularity orders in stress and 235 strain,  $\xi = \frac{m-1}{2-m}$ , ranging between -0.52 and -0.18, with mean and median values of -0.28 236 and -0.24, respectively, appendix B. Such values depart from the "conventional" -0.5 value of 237 LEFM and are close to the value -0.25 expected for flash heating (Brantut and Viesca, 2017; 238 Brener and Bouchbinder, 2021b; Paglialunga et al., 2024) and thermal pressurization (Viesca 239 and Garagash, 2015). In order to further confirm the presence of these unconventional singular-240 ity orders, the stress perturbations associated with passing rupture tips were fit to the ruptures' 241 singular fields using both the unconventional singularity orders found from the trend between  $w_b$ 242 and D and the conventional singularity order of  $\xi = -0.50$  (Irwin (1957); Freund (1990), Eqn 243 4.3.23; Paglialunga et al. (2024), Eqns 7-9), appendix B. The use of unconventional singular-244 ity orders generally resulted in superior fits, Figure 5a, thereby providing further evidence for 245 the unconventional nature of these singularities. These experimental results are a demonstration 246 of the numerical prediction by Lambert and Lapusta (2020) that breakdown energy is "neither a 247 constant material property nor uniquely defined by the amount of slip attained during the rupture" 248 and builds upon previous discussions by, for example, Abercrombie and Rice (2005); Kammer 249 250 et al. (2024) who considered the scale dependence of earthquake rupture.

The fitting of the singular field with the conventional singularity order of  $\xi = -0.50$  further allows for the calculation of  $G_c^{\text{LEFM}}$  (e.g., Svetlizky et al. (2020)). This value of  $G_c^{\text{LEFM}}$  and the



Figure 5: a) An example of the fitting (dark and dashed lines) of  $K_{II}^{(\xi)}$  to the stress change seen by one strain gauge rosette (light lines, green for  $\sigma_{xx}$  and red for  $\sigma_{xy}$ ) as the rupture front passes it. The dashed lines are fit as a conventional singularity with  $\xi = -0.5$ ; the dark lines are fit as an unconventional singularity with  $\xi = -0.23$ . These data are from a calibration experiment and are filtered at 100 kHz. b) From the same experiment, the breakdown work plotted as a function of slip. After the trend between  $w_b$  and D turns over, becoming approximately linear in log-log space, it is fit as  $w_b = AD^m$ . The beginning of this trend corresponds to  $D = D_c$  and  $w_b$  ( $D = D_c$ ) =  $G_c$  and is marked with a blue point. Further, the singularity order can then be calculated as  $\xi = \frac{m-1}{2-m}$ . The cross over in scaling at  $D = D_c$  has been described by Viesca and Garagash (2015); Brantut and Viesca (2017); Brener and Bouchbinder (2021b); Paglialunga et al. (2024). c) The fracture energy,  $G_c$ , versus normal stress,  $\sigma_{yy}$ , based on calibration experiments performed for this study using either LEFM fits (triangles) or the change in trend between breakdown work and slip (circles). Data taken from Paglialunga et al. (2022) are also plotted (squares).

value of  $G_c$  found from the trend between  $w_b$  and D allow for the calibration of  $G_c$ 's dependence 253 on normal stress, Figure 5c and appendix C.  $G_c$ 's dependence on normal stress has been pre-254 viously demonstrated (Okubo and Dieterich, 1981, 1984; Bayart et al., 2016, 2018; Paglialunga 255 et al., 2022), and the results here are in agreement with previous findings, Figure 5(c). Therefore, 256 the increase in normal stress associated with pore pressure preconditioning is predicted to cause 257 in an increase in fracture energy, resulting in a fracture energy barrier. These fracture energy 258 barriers represent *local* impediments to propagation and have been previously shown to be capa-259 ble of halting dynamically-propagating ruptures (Husseini et al., 1975; Bayart et al., 2016, 2018; 260 Gvirtzman and Fineberg, 2021). It should be noted, however, that, for the normal stresses inves-261 tigated here,  $G_c$  varies only from approximately 5 to 10  $\frac{J}{m^2}$ , depending on whether the LEFM or 262 breakdown-work-derived values are used. As will be demonstrated later, these results imply that 263 fracture energy barriers do not exert the primary control on rupture arrest in these scenarios. 264

#### $_{265}$ 4.1.2. Influence of fault preconditioning on the energy flux, G

Complementary to this effect on the fracture energy, fault preconditioning is also expected to reduce the energy flux to the crack tip. In the case of rupture arrest during mode-II propagation in plane stress conditions (in the framework of LEFM - these equations are undefined for unconventional singularities),  $\lim_{\nu \to 0} G = G^{\text{stat}}(a) = \frac{(K_{\text{II}}^{\text{stat}}(a))^2}{E}$  (Freund (1990), Eqn. 5.3.10). The quantity  $K_{\text{II}}^{\text{stat}}(a)$ , the static mode-II stress intensity factor, can then be calculated explicitly as (Tada et al. (2000), Eqn. 8.3; Kammer et al. (2015); Bayart et al. (2016)),

$$K_{\rm II}^{\rm stat}(a) = \frac{2}{\sqrt{\pi a}} \int_0^a \frac{\Delta \sigma_{\rm xy}(s) F\left(\frac{s}{a}\right)}{\sqrt{1 - \left(\frac{s}{a}\right)^2}} {\rm d}s,\tag{3}$$

where  $F\left(\frac{s}{a}\right) = 1 + 0.297 \left(1 - \left(\frac{s}{a}\right)^{\frac{3}{4}}\right)$ , *a* is the crack length, and d*s* is an increment of crack length. In the frameworks of LEFM and cohesive zone models, the residual friction is often treated as a

local property of the interface (e.g., Bayart et al. (2016); Ke et al. (2018)), such that the residual 274 shear stress can be predicted based on the known normal stress prior to the rupture. This enables 275 the calculation of first  $K_{II}^{\text{stat}}$  and then  $G^{\text{stat}}$  from the initial loading conditions. In the case of 276 conventional singularities, this approach provides values of  $G^{\text{stat}}$  which show excellent agreement 277 with observed rupture arrest when compared to  $G_c$  (e.g., Bayart et al. (2016, 2018); Ke et al. 278 (2018)). However, in the case of the unconventional singularities observed here, the calculation 279 of  $\Delta \sigma_{xy}$  is less clear. Behind the observed rupture arrest length, in an analogous fashion to LEFM, 280 the measured stress drops (which, unlike for conventional-singularity-driven ruptures, include 281 282 long-tailed weakening) might be taken after the shear stress achieves a relatively constant value. In front of the observed rupture arrest length, and again as in LEFM (e.g., Bayart et al. (2016); Ke 283 et al. (2018)), an attempt might be made to predict the residual shear stress were the crack to pass 284 a given location. The challenge arises from the fact that shear cracks driven by unconventional 285 singularities do not present locally-constant values of residual friction. Consequently, variations 286 in stress drop, slip, and rupture length are all interdependent, as observed in elastodynamics 287 (e.g., Madariaga (2015)). Nevertheless, to adhere to a procedure typical of LEFM, an attempt 288 was made to characterize the "residual" friction of the fault ahead of the observed rupture arrest 289 length, and thereby predict the residual shear stress and stress drop that would result after the 290 passing of a rupture front. This involved determining the "residual" friction of all complete 291 ruptures, taking the median of this ensemble, and using this value to predict the stress drop 292 that would have occurred for the contained ruptures were they to outgrow the location of their 293 actual arrest. This residual friction profile therefore varies spatially and is assumed to be a local 294 property of the interface. 295

This approach enables the estimation of  $G^{\text{stat}}$ , representative of the mechanical energy flux to 296 the crack tip. G<sup>stat</sup> is reduced during the preconditioning experiments, Figure 6(a-c), reflecting 297 the reduction in energy flux when preconditioning is applied and occurring due to reduction in 298 stress drop available to propagate the crack, Figure 4(j-1). This stress drop reduction, in turn, is 299 due to the increase in the "residual" shear stress, appendix D. Due to the preconditioning phase, 300 the normal stress distribution prior to rupture nucleation is increased from its *in-situ* value, while 301 the shear stress distribution remains approximately unchanged, Figure 2(d,f). As the "residual" 302 shear stress depends directly on the normal stress, this implies that the drop in shear stress avail-303 able has been reduced by preconditioning. Indeed, preconditioning and stress barriers in general 304 (i.e., local decreases in shear stress and/or local increases in normal stress) can even result in a 305 negative stress drop, whereby shear stress is actually increased by the passing rupture, removing 306 energy from the propagating crack tip, Figures 4(j-1). In fact, whether or not a stress drop is pre-307 dicted to be negative or positive has been suggested numerous times as a criterion for dynamic 308 rupture propagation (Garagash and Germanovich, 2012; Cebry et al., 2022), and is inherent in 309 the LEFM framework (Freund (1990), Eqn 7.4.27; Bayart et al. (2016, 2018); Ke et al. (2018)) 310 as a negative stress drop leads to a sharp decrease in energy release rate, see Figure 6(a-c). These 311 stress drop barriers represent *global* impediments to rupture propagation as the stress drop across 312 the entire interface of the slipping fault must be taken in to consideration. These experimental 313 results highlight that pore pressure changes realistically achievable with pore pressure precondi-314 tioning (i.e., tens of percent of the normal stress) lead to relatively large changes in G<sup>stat</sup> when 315 compared to the potential changes in fracture energy, in conditions relevant for laboratory exper-316 iments (up to two orders of magnitude for  $G^{\text{stat}}$  compared to approximately a factor of four for 317  $G_{\rm c}$ ). Indeed, at x = 15 cm, a common position for rupture arrest,  $G_{\rm c}$  is predicted to be increased 318 by a factor of approximately two due to preconditioning, see appendix C. Conversely, at this 319 position  $G^{\text{stat}}$  is increased by a factor of 10 and 100 for experiments at a nominal normal stress 320

of 90 and 120 bar, respectively. There is limited apparent effect for experiments performed at a nominal normal stress of 60 bar. Hence, the primary mechanism leading to the arrest of a rupture related to fault or reservoir preconditioning is the reduction of energy flux to the crack tip.

#### 324 4.2. Influence of barrier size on rupture arrest

Our results demonstrated that fault preconditioning can halt potential seismic rupture prop-325 agation due to the reduction of the energy flux to the crack tip. However, several studies have 326 shed light on ruptures skipping over spatially-limited barriers (Das and Aki, 1977; Cebry et al., 327 2023). The focus here has been principally placed on testing different magnitudes of pore pres-328 sure preconditioning, with  $\frac{\Delta \sigma^{P}}{\sigma^{0}} = 0.08, 0.16, 0.24$ . However, as previously highlighted (Fryer 329 et al., 2023), the extent of this barrier is also of significance. If the period of production leading 330 to preconditioning is too short, the stress barrier may be large in magnitude, but will be small in 331 extent. It is therefore possible that dynamically-propagating ruptures will be able to overcome 332 these barriers and continue propagating. To illustrate this, preconditioning was only applied to 333 one of the pistons (both central- and periphery-only preconditioning were tested, affecting ap-334 proximately one third of the fault length), with the result compared to the base case where the 335 entire sample interface is preconditioned, Figure 7. As can be seen, the ruptures in these cases 336 are liable to skip past the barrier for low values of  $\frac{\Delta \sigma^{P}}{\sigma^{0}}$ , highlighting the significance of achiev-337 ing pore-pressure decreases which are not just large in magnitude, but also in extent. Stress 338 drops behind the crack tip recorded 1 msec after the trigger were also used to reconstruct the 339  $\frac{G^{\text{stat}}}{G_c}$  profiles of these events. Generally, good but imperfect agreement is achieved with theory as 340 the case with preconditioning applied to only the left-most piston never achieves values of  $G^{\text{stat}}$ 341 significantly lower than G<sub>c</sub> and is the only case presented in which the rupture is not arrested, 342 Figure 7a. Note further that the rupture renucleates beyond the preconditioning stress in the case 343 that preconditioning is only applied to the middle piston, Figure 7b. 344

#### 345 4.3. Can the effect of fault preconditioning on rupture length be predicted?

Following Barras et al. (2023), their equation 31, the arrest length,  $\bar{L}_{arr}$ , of a crack-like rupture encountering a stress barrier can be predicted in the framework of LEFM for the case of homogeneous loading and constant residual friction by,

$$\bar{L}_{\rm arr} = -\frac{\bar{x}_{\rm b}\bar{\tau}_0}{\bar{\tau}_{\rm b}},\tag{4}$$

where  $\bar{\mathbf{x}}_{b} = \frac{\mathbf{x}_{b}}{H}$  is the position of the barrier normalized by the damage zone, *H*, and  $\bar{\tau}_{0}$  and  $\bar{\tau}_{b}$  are the dimensionless stress parameters outside and within the barrier,

$$\bar{\tau}_{0} = \frac{\frac{\sigma_{xy}^{0}}{\sigma_{yy}^{0}} - f_{r}}{f_{p} - f_{r}}, \quad \bar{\tau}_{b} = \frac{\frac{\sigma_{xy}^{b}}{\sigma_{yy}^{b}} - f_{r}}{f_{p} - f_{r}},$$
(5)

resulting in, for the dimensional form of arrest length,

$$L_{\rm arr} = x_{\rm b} \frac{\sigma_{\rm xy}^0 - f_{\rm r} \sigma_{\rm yy}^0}{\sigma_{\rm xy}^b - f_{\rm r} \sigma_{\rm yy}^b}.$$
 (6)

Here,  $\sigma_{xy}^0$  and  $\sigma_{yy}^0$  are the shear and normal stresses in the zone outside the barrier, and  $\sigma_{xy}^b$  and  $\sigma_{yy}^b$  are the shear and normal stresses within the barrier, respectively.  $f_p$  and  $f_r$  are the peak and



residual friction coefficients, respectively. This equation represents a simplified approach with 354 its use of unrealistic homogeneous stress values, see Figure 2. However, it may still provide 355 a first-order estimate of the arrest length of the rupture. Considering that two out of the three 356 pistons maintain their preconditioned load during the unloading of the third piston, the barrier 357 position can be taken as one-third of the sample length, or approximately 13 cm. In the case 358 that the nominal normal load is 120 bar, the shear stress can be taken as approximately 1.2 MPa, 359 Figure 2(e,f). The normal load is more heterogeneous. At the moment of nucleation, in the case 360 that  $\sigma^0 = 120$  bar, the normal stress is approximately 0.8 MPa in the unloaded zone and 1.6 MPa 361 in the barrier region. Finally, the measured residual friction is heterogeneous across the sample 362 interface, but to a first order might be taken as 0.5. Crudely, this yields an arrest length estimate 363 of 26 cm which is in agreement with the measured rupture lengths for the preconditioned cases, 364 Figure 4(a-c). It should be emphasized, however, that there is a large degree of arbitrariness in 365 this calculation, which inherently does not consider the large degree of heterogeneity along the 366 sample interface, despite the tested material's synthetic nature and the highly-controlled labo-367 ratory environment. Indeed, for the experiments performed at a nominal normal stress of 120 368 bar without preconditioning, the normal stress in the barrier region is only marginally lower, 369 Figure 2, yet this small discrepancy in normal stress is sufficient to enable complete rupturing 370 in all cases which are not preconditioned. It is difficult to see how such fine margins could be 371 accurately captured by an approach which does not incorporate heterogeneity. More success 372 may be had applying this approach to either larger-scale or more extreme cases, where the stress 373 conditions are not on the border between halting or encouraging rupture propagation. 374

Conversely, the approach of Section 4.1.2 integrates the heterogeneity present along the fault 375 into its prediction of rupture halting and has been successful in this regard (e.g., Bayart et al. 376

Figure 6 (preceding page): While Equation 3 can be used in the prediction of rupture arrest length using the full stress drop, the stress drop at  $D = D_c$  does not provide satisfactory predictions. Therefore, in order to describe shear crack growth when driven by unconventional singularities for which energy dissipation is not exclusively located at the crack tip, continued weakening after  $D = D_c$  must be incorporated; the consideration of energy dissipation localized in the vicinity of the rupture tip without the inclusion of tail processes is not sufficient. (a-c) The profiles of  $G^{\text{stat}}$  for each experiment. G<sup>stat</sup> is based on a combination of the measured stress drop behind the ultimate crack tip arrest length and the predicted stress drop ahead of the ultimate observed crack tip arrest length. The prediction ahead of the ultimate crack tip arrest length is based on the median of the ensemble of residual friction values calculated from experiments without preconditioning, as these experiments exhibited complete ruptures, appendix D. The uncertainty bounds are based on the first and third quartile of the same ensemble of residual friction values. (d-f) The evolution of  $\frac{G^{\text{stat}}}{G_c}$  for each experiment.  $G_c$  profiles are calculated based on the measured normal stress and the trend shown in Figure 5c. (g-i) The evolution of  $\frac{G_{D_c}^{\text{stat}}}{G_c}$  based on the profiles of  $G_{D_c}^{\text{stat}}$  built using the stress drop at  $D = D_c$ . Note that for certain strain gauges it was not

possible to accurately pick  $D_c$ . At these strain gauges the full stress drop was used, such that  $\frac{G_{D_c}^{\text{stat}}}{G_c}$  is overestimated. The friction profile used to predict stress drops ahead of the rupture tip was also built using the stress drop at  $D = D_c$ . (d-i) The position at which the rupture was arrested, based on the videograms, is shown with a square. Where possible the square is placed on the relevant  $\frac{G^{\text{stat}}}{G_c}$  curve; otherwise it is placed at  $\frac{G^{\text{stat}}}{G_c} = 1$ . j) The predicted rupture length versus the measured rupture length for the cases that the energy flux is calculated using the full stress drop, G<sup>stat</sup>, and the stress drop at  $D = D_c$ ,  $G_{D_c}^{\text{stat}}$ , k) The breakdown work development with slip for all experiments based on a strain gauge rosette located outside the approximate point of nucleation (rosette located at x = 29.9 cm). Emphasis is placed on complete ruptures, with preconditioned experiments in a faded color. The ruptures from calibration experiments at the same rosette are shown in faded grey. The breakdown work for all calibration experiments can be found in the appendix B. I) The difference in residual shear stress at  $D = D_c$  compared to when  $D = D_{end}$  versus the slip occurring after  $D = D_c$ .  $D_{end}$ is the final value of slip. The color bar shows the inverted value of  $\xi$ , the singularity order of the rupture at the location of the strain gauge. Experiments performed at (a,d,g) 60, (b,e,h) 90, and (c,f,i) 120 bar nominal normal stress. (a-k) The color bar indicates the amount of preconditioning used in each experiment.



Figure 7: Three separate experiments, all performed at a nominal normal stress of 90 bar and a preconditioning stress of 7 bar. Unloading is performed only by the right-hand side piston. (a-c) The videogram illustrates the progression of the rupture front and is overlain by the local displacement recorded by the horizontally-oriented accelerometers (black lines). The scale for the displacement found from the accelerometers is found on the right. The rupture front is traced with a dotted teal line. Note that certain spatial bands where the grey-scale remains uniform on the videograms correspond to areas where light was not able to pass through the sample at the height of the fault. This is due to it being blocked by a sensor. (d-f) The ratio of  $G^{\text{stat}}$  to  $G_c$  calculated 1 msec after the trigger. As the ratio falls below one the rupture is predicted to halt. White dots mark the locations of strain gauge rosettes. (a,d) Preconditioning is only applied by the leftmost piston. The rupture is initially slowed before accelerating again after approximately 2.5 msec. (b,e) Preconditioning is only applied to middle piston. The rupture is halted but is then able to jump past the barrier and reinitiate on the lefthand side of the sample. (c,f) Preconditioning is applied to all three pistons. The rupture is halted and does not reinitiate. Note that rupture halting is predicted in (f) despite the large values of  $\frac{G^{\text{stat}}{G_c}}{F_c}$  predicted on the left-hand side of the sample.

(2016, 2018); Ke et al. (2018)). However, unlike the predictions by Barras et al. (2023) and 377 others (e.g., Kammer et al. (2015)), these approaches require the characterization of the residual 378 friction (and therefore residual shear stress) along the entirety of the fault, and generally assume 379 that the residual friction is a local interface property. This poses a problem for the case of un-380 conventional singularities, where a steady-state residual friction is not reached and the amount 381 of weakening and residual shear stress depend on the amount of slip, Figure 6(k,l). Interestingly, 382 by considering the full stress drop in Equation 3, the calculation of  $G^{\text{stat}}$  yields satisfactory pre-383 dictions of rupture arrest length when compared to  $G_c$ , Figure 6(d-f,j). Unfortunately, due to the 384 continued weakening, the *a priori* predictions of these full stress drops is not obvious, with slip, 385 rupture length, and stress drop being fully coupled. 386

One alternative possibility is to assume that the energy dissipation relevant for driving the 387 rupture is exclusively located at the crack tip, in the process zone. If this were to be the case, 388 it might be possible to characterize the interface's residual friction at  $D = D_c$  such that a priori 389 predictions could be performed. To test this assumption, the stress drop up to  $D = D_c$  was used 390 in Equation 3 to ultimately calculate the tip-localized static energy flux,  $G_{D_c}^{\text{stat}}$ .  $G_{D_c}^{\text{stat}}$  is always less 391 than or equal to  $G^{\text{stat}}$  due to continued long-tail weakening which occurs when  $D > D_c$  (Lambert 392 and Lapusta, 2020; Brener and Bouchbinder, 2021b,a; Paglialunga et al., 2022, 2024). Values of 393  $G_{D_c}^{\text{stat}}$  were compared to  $G_c$  to predict crack arrest, Figure 6(g-i). This methodology under predicts 394 ultimate crack length, Figure 6(j), implying that  $G_{D_c}^{\text{stat}}$  is an under-prediction of the energy driving 395 rupture propagation. The continued weakening occurring when  $D > D_c$ , Figure 6(k,l), which 396 is not localized at the crack tip, therefore plays a significant role in the propagation of ruptures 397 driven by unconventional singularities. Further, as this continued weakening must be considered 398 when calculating the energy flux driving the crack tip, the characterization of a residual friction 399 to be used in the *a priori* prediction of stress drop is difficult or even impossible for these ruptures 400 as a true residual friction is not achieved even after  $D \gg D_c$ . Several weakening mechanisms are 401 expected to induce long-tailed weakening (i.e. unconventional singularity orders), such as ther-402 mal pressurization (Viesca and Garagash, 2015), flash heating (Brantut and Viesca, 2017), and 403 melt lubrication of the full fault interface (Di Toro et al., 2011). Therefore, long-tailed weaken-404 ing may be a common phenomenon for natural earthquakes. Nevertheless, it cannot be excluded 405 that the residual friction might reach a steady-state value after a large amount of slip such that an 406 equilibrium between the heat production and the heat dissipation within the fault zone has been 407 achieved (although this seems unlikely for thermal pressurization considering the relatively long 408 time scale of fluid diffusion). If this steady state residual friction is attained, the singularity is 409 expected to become conventional, allowing for predictions using LEFM. However, steady state 410 will require large amounts of slip (>10 meters (Di Toro et al., 2011), likely corresponding to a 411 large moment-magnitude,  $M_{\rm w} > 7$ , earthquake), resulting in cohesive zones on the order of kilo-412 meters (considering slip velocities on the order of meters per second and rupture velocities of 413 many hundreds of meters per second). These large sections the fault which do not achieve suffi-414 cient slip will not reach a steady state friction. This steady state is further unlikely to be achieved 415 in the context of induced seismicity and smaller magnitude events as thermal slip equilibrium 416 distances depend inversely on normal stress (Di Toro et al., 2011). 417

In order for fracture mechanics to one day provide *a priori* predictions of rupture arrest for unconventional singularities, it must account for this continued weakening (Brener and Bouchbinder, 2021b,a). Ideally, an equation of motion describing the propagation of the ruptures should be developed to replace the edge-localized Equation 1 for unconventional-singularity-driven ruptures. It may be that, in the case of rupture arrest, this equation of motion will lend itself to cases

where  $G_{\rm c}$  can be neglected, potentially leading rupture propagation domains, such as a stress-423 drop-dominated domain, analogous to the zero-toughness asymptotic solutions for fluid-driven 424 mode-I fractures (Savitski and Detournay, 2002). At any rate, it has been shown here that stress 425 barriers not only affect the LEFM-defined parameters  $G_c$  and G, they also further reduce the 426 breakdown work by reducing the slip and weakening (stress drop) occurring outside of the co-427 hesive zone, thereby further reducing the energy flux propagating to the crack tip beyond what 428 can be predicted by LEFM. Indeed, the events with the smallest amount of additional weaken-429 ing when  $D > D_c$  are those with the largest stress barriers. This additional, distal weakening 430 has been shown to be significant for the accurate calculation of the energy flux to the crack tip. 431 These results are even more relevant for induced seismicity, since shear ruptures propagating in 432 the presence of fluids are inherently expected to be driven by unconventional singularities (i.e., 433 even without the activation of thermal weakening, as for instance in the case of mode-I frac-434 tures), due to the diffusion of the fluid pressure along the fault and in the surrounding medium 435 (Garagash et al., 2011). 436

#### 437 4.4. Natural stress barriers and earthquake swarms

Heterogeneous pore pressure fields associated with some seismic swarms can induce stress 438 barriers which can be used to explain changes in stress drop within the swarm without recourse 439 to material or total-stress heterogeneity. Earthquake swarms are common in areas of volcanism 440 and geothermal activity and occur such that the number and magnitude of earthquakes fluctuate 441 in time whilst not presenting a distinct, larger-magnitude main shock (Mogi, 1963). Swarms 442 have been previously suggested to be related to heterogeneous material properties and stress 443 distributions (Mogi, 1963), with evidence from laboratory (Scholz, 1968) and earthquake moni-444 445 toring (Ross et al., 2020) studies. For example, the 2016-2019 earthquake swarm near Cahuilla, California is thought to have been due to fluid influx and to be located along a fault with per-446 meability barriers (Ross et al., 2020). In fact, before the occurrence of a moment magnitude, 447  $M_{\rm w}$ , 4.4 earthquake, the swarm was seen to migrate in space and time, exhibiting progressively 448 lower stress drops (Ross et al., 2020). While the influence of mechanical properties is likely in 449 this case (Ross et al., 2020), for a constant shear stress (on the time scale of the fluid diffusion), 450 a lower fluid pressure will result in an increased effective normal stress and, therefore, an in-451 creased "residual" shear stress, resulting in an ultimately lower stress drop, as has been shown 452 in this work. Since the fluid source in this  $M_{\rm w}$  4.4 earthquake's case was thought to be coming 453 from a point source (a broken seal connected to a deeper reservoir) (Ross et al., 2020), the earth-454 quakes occurring farther from the point source were likely occurring in and/or propagating into 455 zones of lower fluid pressure and therefore higher residual shear stress, offering an additional 456 explanation for trends in stress drop seen for this swarm. It should be further noted that, as with 457 injection-induced seismicity, fluid-driven earthquake swarms can be expected to be characterised 458 by unconventional-singularity-driven ruptures due to their reliance on fluid pressure, making the 459 results presented here particularly applicable. 460

#### 461 5. Conclusion and Outlook

<sup>462</sup> Preconditioning has been demonstrated experimentally, with stress changes that correspond to pore pressure changes which are reasonably achievable in the field. Preconditioning consistently results in rupture arrest at stress changes which correspond to pore pressure reductions of  $2 \frac{MPa}{km}$ . All tested magnitudes of preconditioning (1, 2, and  $3 \frac{MPa}{km}$ ) result in reduced co-seismic slip, slip velocity, stress drop, and seismic moment. However, these experiments were performed
 on analogue material at laboratory scales and stresses. Meso-scale testing could be pursued to
 test the influence of these limitations.

The experimental ruptures investigated here were driven by unconventional singularities, 469 such that breakdown work continuously increased with slip. LEFM requires scale separation 470 between edge-localized dissipation and linear elastic driving energy. The breakdown work in-471 creasing with slip breaks this scale separation as the energy dissipation is not exclusively local-472 ized near the crack tip (Brener and Bouchbinder, 2021b,a; Paglialunga et al., 2022, 2024). This 473 makes the *a priori* use of a constant residual frictional coefficient to predict rupture arrest unten-474 able. In order to precisely predict rupture arrest the full evolution of shear stress with slip must 475 be considered, and an equation of motion for unconventional-singularity-driven ruptures should 476 be developed. Here, ruptures which propagated into stress barriers were characterised by less 477 slip and less weakening (stress drop) behind the crack tip. This means that stress barriers arrest 478 ruptures by not just increasing the "residual" shear stress reached in the cohesive zone of the 479 passing crack tip, they also reduce the long-tailed weakening occurring farther behind the crack 480 tip by inhibiting slip and rupture advance. In summary, stress barriers increase fracture energy, 481 reduce energy flux to the crack tip resulting from cohesive-zone weakening, and reduce energy 482 flux to the crack tip resulting from distal, long-tailed weakening. All three of these effects aid 483 in the arrest of the rupture and explain not only why preconditioning has the ability to reduce 484 seismic hazard but also why natural fluid-driven earthquake swarms might exhibit lower stress 485 drops away from the fluid source. 486

#### 487 Acknowledgements

B.F., C.N., and both F.P.'s acknowledge funding from the European Union (ERC Starting
Grant HOPE num. 101041966). B.F. acknowledges Hannes Hofmann for a useful discussion.
The authors thank Julien Ambre for his help preparing sensors. Dmitry Garagash is thanked for
helpful comments and suggestions.

A CC-BY public copyright license has been applied by the authors to the present document
 and will be applied to all subsequent versions up to the Author Accepted Manuscript arising from
 this submission, in accordance with the grant's open access conditions.



Distributed under a Creative Commons Attribution — 4.0 Interna-

```
497 tional licence
```

495

496

#### 498 Data availability

The data collected for this work and used in the analyses is publicly available (Fryer et al., 2024).

#### 501 Appendix A. Tracking the rupture

Strain gauge data are filtered with a 6-th order low-pass Butterworth filter at 28 kHz, except for the fitting of the singular field where the filter is set to 100 kHz. The strain,  $\epsilon_i$ , measured on <sup>504</sup> each strain gauge, i, is calculated as,

$$\epsilon_i = \frac{-4V_i}{V_{\text{ex}} \left( G_{\text{f}} G_{\text{amp}} \left( 1 + \frac{2V_i}{V_{\text{ex}} G_{\text{amp}}} \right) \right)},\tag{A.1}$$

where  $V_i$  is the voltage measured by the *i*-th strain gauge,  $V_{ex}$  the excitation voltage,  $G_f$  the gauge factor, and  $G_{amp}$  the amplification gain. Depending on the orientation of the strain gauge from the vertical direction in degrees, these strains are referred to as  $\epsilon_{315}$ ,  $\epsilon_0$ , and  $\epsilon_{45}$ . These strains are then used to calculate the strain tensor at the location of the strain gauge rosette as,

$$\epsilon_{xx} = \epsilon_{315} + \epsilon_{45} - \epsilon_0, \quad \epsilon_{yy} = \epsilon_0, \quad \epsilon_{xy} = \frac{\epsilon_{45} - \epsilon_{315}}{2}.$$
 (A.2)

The stress tensor was then calculated using Hooke's Law in plane stress, considering the strain rate dependence of the Young's Modulus of PMMA. A static Young's Modulus of 3.3 GPa and a dynamic Young's Modulus of 5.7 GPa were employed. Poisson's ratio was taken as 0.33, appropriate for PMMA. The static Young's Modulus was used to find all loading-phase stresses. The dynamic Young's Modulus was used to calculate the dynamic stress drop, following Bayart et al. (2016, 2018).

For each frame, the grayscale image recorded by the camera is taken as a 1280 x 32 matrix where the value ranges from zero (black) to a maximum value (white). The average greyscale map of the first 20 frames serve as a reference. All the future frames are compared to this reference in order to track stress changes along the interface and, therefore, track the rupture front position (e.g., Nielsen et al. (2010); Schubnel et al. (2011)).

The data from the accelerometers is converted from mV into  $\frac{m}{sec^2}$  using the calibration provided by Brüel & Kjær. These accelerations are integrated twice in time to find both the velocity and displacement (e.g., Schubnel et al. (2011)). These data are unfiltered and multiplied by two, considering they represent movement on only one of the two similarly-sized blocks.

#### 524 Appendix B. Characterizing dynamic rupture

To further confirm the presence of the unconventional singularity orders, the stress perturbation associated with a passing rupture can be described as (Irwin (1957); Freund (1990), Eqn 4.3.23; Paglialunga et al. (2024), Eqns 7-9),

$$\Delta \sigma_{\rm xx} \left( r, \theta \right) = \frac{4 \left( \xi + 1 \right) K_{\rm II}^{\left( \xi \right)}}{\hat{D} \sqrt{2\pi}} \left[ \alpha_{\rm s} \left( 1 + 2\alpha_{\rm d}^2 - \alpha_{\rm s}^2 \right) r_{\rm d}^{\xi} \sin \left( \xi \theta_{\rm d} \right) - \alpha_{\rm s} \left( 1 + \alpha_{\rm s}^2 \right) r_{\rm s}^{\xi} \sin \left( \xi \theta_{\rm s} \right) \right],$$
  
$$\sigma_{\rm xy} \left( r, \theta \right) = \sigma_{\rm xy}^{\rm min} + \frac{2 \left( \xi + 1 \right) K_{\rm II}^{\left( \xi \right)}}{\hat{D} \sqrt{2\pi}} \left[ 4\alpha_{\rm s} \alpha_{\rm d} r_{\rm d}^{\xi} \cos \left( \xi \theta_{\rm d} \right) - \left( 1 + \alpha_{\rm s}^2 \right)^2 r_{\rm d}^{\xi} \cos \left( \xi \theta_{\rm s} \right) \right],$$
  
$$\Delta \sigma_{\rm yy} \left( r, \theta \right) = \frac{4 \left( \xi + 1 \right) \alpha_{\rm s} \left( 1 + \alpha_{\rm s}^2 \right) K_{\rm II}^{\left( \xi \right)}}{\hat{D} \sqrt{2\pi}} \left[ r_{\rm d}^{\xi} \sin \left( \xi \theta_{\rm d} \right) - r_{\rm s}^{\xi} \sin \left( \xi \theta_{\rm s} \right) \right], \quad (B.1)$$

where *r* and  $\theta$  are polar coordinates centered on a steadily-moving crack tip,  $K_{II}^{(\xi)}$  is the instantaneous mode-II stress intensity factor,  $\hat{D} = 4\alpha_d\alpha_s - (1 + \alpha_s^2)^2$  is the Rayleigh function (Freund (1990), Eqn 4.3.8), and  $\alpha_d = \sqrt{1 - (v/C_d)^2}$  and  $\alpha_s = \sqrt{1 - (v/C_s)^2}$  are velocity factors (Freund (1990), Eqn 4.3.12), where *v* is the rupture velocity and  $C_d$  and  $C_s$  are the P- and S-wave

Table B.1: Results of the fitting procedure used to find  $G_c$  described in the text. The ID contains the nominal normal stress of the calibration experiment, the the event number, and the strain gauge rosette number.  $\sigma_{yy}$  is the local normal stress measured by the strain gauges, v the local rupture velocity found from the videograms, l is the cohesive zone size and is found by choosing  $D_c$  on the plot of  $w_b$  vs. D (see main text for plot and Paglialunga et al. (2024), for example) and finding the time required to achieve this  $D_c$  since the passing of the rupture front. This time is multiplied by v (assumed constant) to yield l. The error in l is found based on the error in  $D_c$  (which affects the time to achieve  $D_c$ ) and an assumed error of 10  $\frac{m}{sec}$  in v. Note that the uncertainty in l can be considered to be larger than the error presented here due to the sensitivity of l to other parameters.  $G_c$  ( $D = D_c$ ),  $D_c$ , and  $\xi$  are found from the plot of breakdown work versus slip.

ID	$\sigma_{\rm yy}$ [MPa]	$v\left[\frac{m}{\text{sec}}\right]$	$G_{\rm c}\left(\xi = -\frac{1}{2}\right)\left[\frac{\rm J}{\rm m^2}\right]$	<i>l</i> [cm]	<i>D</i> <sub>c</sub> [μm]	$G_{\rm c} \left( D = D_{\rm c} \right) \left[ \frac{\rm J}{\rm m^2} \right]$	ξ[-]
200/1/11	1.95	214	9.77	$1.2 \pm 0.2$	$16 \pm 3$	$7.47 \pm 0.38$	$-0.24 \pm 0.03$
200/2/10	1.76	281	4.94	$0.6 \pm 0.2$	$9 \pm 5$	$2.55 \pm 0.18$	$-0.52\pm0.04$
200/2/11	2.07	178	8.10	$1.3 \pm 0.1$	$12 \pm 1$	$4.51 \pm 0.30$	$-0.24 \pm 0.04$
200/3/7	2.61	145	4.69	$4.2 \pm 0.3$	$19 \pm 0$	$5.06 \pm 0.22$	$-0.18 \pm 0.25$
200/3/11	2.19	159	8.12	$1.6 \pm 0.4$	$15 \pm 5$	$6.01 \pm 0.45$	$-0.27\pm0.06$
300/4/5	2.39	26	13.68	$2.3 \pm 1.0$	$19 \pm 1$	$6.43 \pm 0.40$	$-0.23\pm0.06$

velocities, respectively.  $\sigma_{yy}$  is positive in compression. The polar coordinates are corrected 532 for distortion as  $\theta_d = \arctan(\alpha_d \tan(\theta))$ ,  $\theta_s = \arctan(\alpha_s \tan(\theta))$ ,  $r_d = r\sqrt{1 - (v\sin(\theta)/C_d)^2}$ , and 533  $r_{\rm s} = r \sqrt{1 - (v \sin(\theta) / C_{\rm s})^2}$  (Freund (1990), Eqn 4.3.12).  $\sigma_{\rm xy}^{\rm min}$  is the minimum value of shear 534 stress achieved after the passing of the rupture. By fitting all three stress components to Equa-535 tion B.1,  $K_{II}^{(\xi)}$  can be found for each location along the fault with reliable strain gauge measure-536 ments. The singularity order,  $\xi$ , is -0.5 for conventional-singularity-driven shear cracks, which 537 reduces Equation B.1 to the classical equations (Freund (1990), Eqn 4.3.23). Here,  $K_{II}^{(\xi)}$  was fit 538 with both  $\xi = -0.5$  and the value of  $\xi$  found from the fitting of the trend between breakdown 539 work to slip. The use of unconventional singularity orders generally resulted in superior fits, 540 thereby providing further evidence for the unconventional nature of these singularities. Note that 541 this inversion requires a constant rupture velocity. The variation in rupture velocities during these 542 experiments may lead to error in the inversion. 543

The fits performed assuming a conventional singularity order (i.e.,  $\xi = -0.5$ ) were used to provide a point of comparison for the values of  $G_c$  found considering the trend between  $w_b$  and *D*. Following LEFM, *G*, which depends on the velocity of propagation, can be related to the the dynamic mode-II stress intensity factor,  $K_{II}^{(\xi=-0.5)}$ , as (Freund (1990), Eqn 5.3.10; Bayart et al. (2016)),

$$G(v) = \frac{\alpha \left(1 - v^2\right)}{E} f_{\rm II}(v) \left[K_{\rm II}^{(\xi = -0.5)}(a, v)\right]^2, \tag{B.2}$$

where  $\alpha = 1$  in plane stress conditions, *a* is the crack length, and  $f_{\text{II}}(\nu)$  is a decreasing function of velocity equivalent to (Freund (1990), Eqn 5.3.11),

$$f_{\rm II}(\nu) = \frac{\alpha_{\rm s} \nu^2}{\hat{D}(1-\nu) C_{\rm s}^2},\tag{B.3}$$

having the property  $\lim_{\nu \to 0} f_{\text{II}}(\nu) = 1$  (Freund (1990), page 234). Considering then that during dynamic rupture propagation  $G = G_{\text{c}}$ , the fit performed using Equation B.1 assuming  $\xi = -0.5$ , allows for the calculation of  $G_{\text{c}}$ . This value of  $G_{\text{c}}$  and the value of  $G_{\text{c}}$  found from the trend between  $w_{\text{b}}$  and D allow for the calibration of  $G_{\text{c}}$ 's dependence on normal stress, Table B.1.

#### <sup>555</sup> Appendix C. Dependence of the fracture energy $G_c$ on normal stress

 $G_{\rm c}$  depends on both the stress drop and the critical slip distance (Rice (1980), Chapter 5; 556 Ohnaka (2003)) and therefore approximately quadratically with normal stress. Locally, both 557 peak and residual shear stress vary linearly with normal stress, due to changes in real contact 558 area (Bayart et al., 2016), with real contact area depending (sub-)linearly on normal stress (Bow-559 den and Tabor, 1938; Archard, 1957).  $D_c$  can be estimated to vary linearly with normal stress 560 assuming a constant rupture velocity and purely slip weakening behaviour (Ida, 1972; Palmer 561 and Rice, 1973; Rice, 1980), with D<sub>c</sub>'s dependence on normal stress shown experimentally for 562 PMMA (Paglialunga et al., 2022) and granite (Passelègue et al., 2016). Considering that fracture energy is estimated as  $G_c = D_c(\sigma_{yy})\sigma_{yy}\frac{f_p-f_r}{2}$  (e.g., Ida (1972); Palmer and Rice (1973); Rice (1980), Chapter 6; Okubo and Dieterich (1981, 1984)) in the linear slip weakening case, where 563 564 565  $f_{\rm p}$  and  $f_{\rm r}$  are the peak and residual friction coefficients, respectively, the result is that fracture 566 energy is predicted to vary approximately quadratically with normal stress. This dependence 567 could be even stronger as  $f_p - f_r$  is also considered to scale with normal stress (Passelègue et al., 568 2016), for example through the activation of dynamic weakening mechanisms, such as flash 569 heating (e.g., Brantut and Viesca (2017)), where residual friction is greatly reduced. It should 570 be noted that  $K_{II}^{\text{stat}}$  also depends linearly on stress drop, implying a quadratic dependence of  $G^{\text{stat}}$ 571 on normal stress in an LEFM framework. The activation of thermal weakening mechanisms at 572 higher normal stresses, leading to long-tailed weakening, can be expected to further strengthen 573 this dependence and may lead to a scale dependence in the relative importance of  $G^{\text{stat}}$  and  $G_{\text{c}}$  at 574 higher normal stresses and longer rupture lengths. The measured normal stress and empirically-575 estimated fracture energy profiles are displayed in Figure C.8. 576

#### 577 Appendix D. Residual shear stress

In these experiments, the residual shear stress is not a constant, nor a material property, Figure D.9. As described in the main text, normal stress barriers result in an increase in residual shear stress due to their ability to impede frictional rupture, thereby reducing distal weakening. Further, in a frictional setting the residual shear stress is generally thought as dependent on normal stress. As normal stress barriers represent zones of increased normal stress, the residual shear stress can be expected to be larger in these cases.

In order to predict the stress drop ahead of the ultimate crack length (i.e., in zones where the rupture did not pass and stress drop could therefore not be measured), the residual friction 585 of the samples was characterized. This was done by taking the residual friction of all six non-586 preconditioned experiments as these experiments exhibited complete ruptures. The median and 587 first and third quartile values of this ensemble, Figure D.10, were then used in combination with 588 the measured normal and shear stresses to predict the stress drops of unruptured zones in the 589 calculation of Equation 3. Note the edges of the sample exhibit large errors in residual friction 590 due to the low values of normal stress present in these locations. Generally these areas are 591 insignificant for rupture arrest prediction, as ruptures arrest typically between 10 and 20 cm. 592

#### 593 References

Abercrombie, R., Rice, J., 2005. Can observations of earthquake scaling constrain slip weakening? Geophysical Journal
 International 162, 406–424. Https://doi.org/10.1111/j.1365-246X.2005.02579.x.



Figure C.8: (a-c) The normal stress measured by the strain gauges just before nucleation, found using the static Young's Modulus and limited to a minimum value of zero. (d-f) The predicted value of  $G_c = \sigma_{yy}^2 + 0.001$  (with  $\sigma_{yy}$  in MPa) just before nucleation across the entire laboratory fault considering the measured normal stress. Experiments performed at (a,d) 60, (b,e) 90, and (c,f) 120 bar nominal normal stress. The color bar indicates the amount of preconditioning used in each experiment.



Figure D.9: The measured residual shear stress of each experiment, taken only for locations behind the arrested crack tip. Results for (a) 60, (b) 90, and (c) 120 bar nominal normal stress. The color bar indicates the amount of preconditioning used in each experiment.



Figure D.10: The residual friction calculated based on all 6 non-preconditioned experiments, with estimated error shown as grey shading. The error and value of the residual friction are based on the first, second, and third quartile values of residual friction for the ensemble of experiments.

- Aki, K., 1966. Generation and propagation of G waves from the Niigata Earthquake of June 16, 1964. Part 2. Estimation
   of earthquake moment, released energy, and stress-strain drop from the G wave spectrum. Bulletin of the Earthquake
   Research Institute 44, 73–88.
- Archard, J., 1957. Elastic deformation and the laws of friction. Proceedings of the Royal Society of London Series A
   243, 190–205. Https://dx.doi.org/10.1098/rspa.1957.0214.
- Bao, X., Eaton, D., 2016. Fault activation by hydraulic fracturing in western Canada. Science 354, 1406–1409. Https://doi.org/10.1126/science.aag258.
- Baria, R., Michelet, S., Baumgaertner, J., Dyer, B., Gerard, A., Nicholls, J., Hettkamp, T., Teza, D., Soma, N., Asanuma,
   H., Garnish, J., Megel, T., 2004. Microseismic monitoring of the World's largest potential HDR reservoir, in: Twenty Ninth Workshop on Geothermal Reservoir Engineering Stanford University, Stanford, California, January 26-28,
   2004.
- Barras, F., Thøgersen, K., Aharonov, E., Renard, F., 2023. How do earthquakes stop? Insights from a minimal model of
   frictional rupture. Journal of Geophysical Research: Solid Earth 128. https://doi.org/10.1029/2022JB026070.
- Bayart, E., Svetlizky, I., Fineberg, J., 2016. Fracture mechanics determine the lengths of interface ruptures that mediate
   frictional motion. Nature Physics 12, 166–170. Https://doi.org/10.1038/nphys3539.
- Bayart, E., Svetlizky, I., Fineberg, J., 2018. Rupture dynamics of heterogeneous frictional interfaces. Journal of Geo physical Research: Solid Earth 123, 3828–3848. Https://doi.org/10.1002/2018JB015509.
- Ben-David, O., Cohen, G., Fineberg, J., 2010. The dynamics of the onset of frictional slip. Science 330, 211–214.
   Https://doi.org/10.1126/science.1194777.
- Bhattacharya, P., Viesca, R., 2019. Fluid-induced aseismic fault slip outpaces pore-fluid migration. Science 364, 464–
   468. Https://doi.org/10.1126/science.aaw7354.
- Bowden, F., Tabor, D., 1938. The area of contact between stationary and between moving surfaces. Proceedings of the
   Royal Society of London. Series A 169, 391–413. https://doi.org/10.1098/rspa.1939.0005.
- Brantut, N., Viesca, R., 2017. The fracture energy of ruptures driven by flash heating. Geophysical Research Letters 44,
   6718–6725. Https://doi.org/10.1002/2017GL074110.
- Brener, E., Bouchbinder, E., 2021a. Theory of unconventional singularities of frictional shear cracks. Journal of the
   Mechanics and Physics of Solids 153. Https://doi.org/10.1016/j.jmps.2021.104466.
- Brener, E., Bouchbinder, E., 2021b. Unconventional singularities and energy balance in frictional rupture. Nature Communications 12. Https://doi.org/10.1038/s41467-021-22806-9.
- Campillo, M., Ionescu, I., 1997. Initiation of antiplane shear instability under slip dependent friction. Journal of Geo physical Research 102, 20363–20371. Https://doi.org/10.1029/97JB01508.
- Cebry, S., Ke, C., McLaskey, G., 2022. The role of background stress state in fluid-induced aseismic
   slip and dynamic rupture on a 3-m laboratory fault. Journal of Geophysical Research: Solid Earth 127.
   Https://doi.org/10.1029/2022JB024371.
- Cebry, S., McLaskey, G., 2021. Seismic swarms produced by rapid fluid injection into a low permeability laboratory
   fault. Earth and Planetary Science Letters 557. Https://doi.org/10.1016/j.epsl.2020.116726.
- 632 Cebry, S., Sorhaindo, K., McLaskey, G., 2023. Laboratory earthquake rupture interactions with a high normal stress

- bump. Journal of Geophysical Research: Solid Earth 128. Https://doi.org/10.1029/2023JB027297. 633
- Clark, J., 1949. A hydraulic process for increasing the productivity of wells. Journal of Petroleum Technology 1. 634 635 Https://doi.org/10.2118/949001-G.
- Das, S., Aki, K., 1977. Fault plane with barriers: A versatile earthquake model. Journal of Geophysical Research 82, 636 5658-5670. Https://doi.org/10.1029/JB082i036p05658. 637
- Di Toro, G., Han, R., Hirose, T., De Paola, N., Nielsen, S., Mizoguchi, K., Ferri, F., Cocco, M., Shimamoto, T., 2011. 638 Fault lubrication during earthquakes. Nature 471, 494–498. Https://doi.org/. 639
- Frash, L., Fu, P., Morris, J., Gutierrez, M., Neupane, G., Hampton, J., Welch, N., Carey, J., Kneafsey, T., 2021. Fracture 640 caging to limit induced seismicity. Geophysical Research Letters 48. Https://doi.org/10.1029/2020GL090648. 641
- Freund, L., 1990. Dynamic Fracture Mechanics. Cambridge University Press. 642
- Fryer, B., Lebihain, M., Noël, C., Paglialunga, F., Passelègue, F., 2024. Data set for: The effect of stress barriers on 643 unconventional-singularity-driven seismic rupture: An experimental demonstration of fault preconditioning". doi:10. 644 5281/zenodo.13772819 645
- Fryer, B., Lebihain, M., Violay, M., 2023. Single-well pore pressure preconditioning for Enhanced Geothermal System 646 stimulation. Journal of Geophysical Research: Solid Earth 128. Https://doi.org/10.1029/2022JB025443. 647
- Fryer, B., Siddiqi, G., Laloui, L., 2020. Injection-induced seismicity: strategies for reducing risk using high stress 648 path reservoirs and temperature-induced stress preconditioning. Geophysical Journal International 220, 1436-1446. 649 Https://doi.org/10.1093/gji/ggz490. 650
- Galis, M., Ampuero, J., Mai, P., Cappa, F., 2017. Induced seismicity provides insight into why earthquake ruptures stop. 651 Science Advances 3. Https://doi.org/10.1126/sciadv.aap7528. 652
- Garagash, D., 2021. Fracture mechanics of rate-and-state faults and fluid injection induced slip. Philosophical Transac-653 654 tions of the Royal Society A 379. doi:https://doi.org/10.1098/rsta.2020.0129.
- Garagash, D., Detournay, E., Adachi, J., 2011. Multiscale tip asymptotics in hydraulic fracture with leak-off. Journal of 655 Fluid Mechanics 669, 260-297. doi:10.1017/S002211201000501X. 656
- Garagash, D., Germanovich, L., 2012. Nucleation and arrest of dynamic slip on a pressurized fault. Journal of Geophys-657 ical Research 117. doi:https://doi.org/10.1029/2012JB009209. 658
- Gounon, A., Latour, S., Letort, J., El Arem, S., 2022. Rupture nucleation on a periodically heterogeneous interface. 659 Geophysical Research Letters 49. Https://doi.org/10.1029/2021GL096816. 660
- Griffith, A., 1921. Vi. the phenomena of rupture and flow in solids. Philosophical Transactions of the Royal 661 Society of London. Series A, Containing Papers of a Mathematical or Physical Character 221, 163-198. 662 Https://doi.org/10.1098/rsta.1921.0006. 663
- Gvirtzman, S., Fineberg, J., 2021. Nucleation fronts ignite the interface rupture that initiates frictional motion. Nature 664 Physics 17, 1037-1042. Https://doi.org/10.1038/s41567-021-01299-9. 665
- Häring, M., Schanz, U., Ladner, F., Dyer, B., 2008. Characterisation of the Basel 1 enhanced geothermal system. 666 Geothermics 37, 469-495. Https://doi.org/10.1016/j.geothermics.2008.06.002. 667
- Hofmann, H., Zimmermann, G., Farkas, M., Huenges, E., Zang, A., Leonhardt, M., Kwiatek, G., Martinez-Garzon, P., 668
- Bohnhoff, M., Min, K., Fokker, P., Westaway, R., Bethmann, F., Meier, P., Yoon, K., Choi, J., Lee, T., Kim, K., 669 2019. First field application of cyclic soft stimulation at the Pohang Enhanced Geothermal System stie in Korea. 670 Geophysical Journal International 217, 926–949. Https://doi.org/10.1093/gji/ggz058. 671
- Hubbert, M., Willis, D., 1957. Mechanics of hydraulic fracturing. Petroleum Transactions, AIME 210, 153-168. 672
- Husseini, M., Jovanovich, D., Randall, M., Freund, L., 1975. The fracture energy of earthquakes. Geophysical Journal 673 International 43, 367-385. Https://doi.org/10.1111/j.1365-246X.1975.tb00640.x. 674
- Ida, Y., 1972. Cohesive force across the tip of a longitudinal-shear crack and Griffith's specific surface energy. Journal 675 of Geophysical Research 77. Https://doi.org/10.1029/JB077i020p03796. 676
- Irwin, G., 1957. Analysis of stresses and strains near the end of a crack traversing a plate. Journal of Applied Mechanics 677 24, 361-364. Https://doi.org/10.1115/1.4011547. 678
- Jalali, M., Selvadurai, P., Dal Zilio, L., Durand, V., Obermann, A., Meier, M., Amann, F., Zappone, A., Shakas, A., Ri-679 naldi, A., Maurer, H., Bröker, K., Hertrich, M., Doonechaly, N., Achtziger-Zupancic, P., Wenning, Q., Ma, X., Cocco, 680 M., Wiemer, S., Giardini, D., 2023. An attempt to control induced seismicity and on-fault stress pre-conditioning in 681 the Bedretto Underground Laboratory, Switzerland, in: Proceedings: EGU General Assembly 2023, Vienna, Austria, 682
- 24-28 Apr 2023. 683 Kammer, D., McLaskey, G., Abercrombie, R., Ampuero, J., Cattania, C., Cocco, M., Dal Zilio, L., Dresen, G., Gabriel, 684
- A., Ke, C., Marone, C., Selvadurai, P., Tinti, E., 2024. Earthquake energy dissipation in a fracture mechanics frame-685 work. Nature Communications 15. Https://doi.org/10.1038/s41467-024-47970-6. 686
- Kammer, D., Radiguet, M., Ampuero, J., Molinari, J., 2015. Linear elastic fracture mechanics predicts the propagation 687 distance of frictional slip. Tribology Letters 57. Https://doi.org/. 688
- Ke, C., McLaskey, G., Kammer, D., 2018. Rupture termination in laboratory-generated earthquakes. Geophysical 689 Research Letters 45, 12784-12792. Https://doi.org/10.1029/2018GL080492. 690
- Kilgore, B., McGarr, A., Beeler, N., Lockner, D., 2017. Earthquake source properties from instrumented laboratory 691 25

- stick-slip, in: Thomas, M., Mitchell, T., Bhat, H. (Eds.), Fault Zone Dynamic Processes: Evolution of Fault Properties
   During Seismic Rupture, Geophysical Monograph. John Wiley & Sons, Inc., chapter 8, pp. 151–169.
- Kim, K., Ree, J., Kim, Y., Kim, S., Kang, S., Seo, W., 2018. Assessing whether the 2017  $M_w$  5.4 Pohang earthquake in South Korea was an induced event. Science 360, 1007–1009. Https://doi.org/10.1126/science.aat6081.
- Lambert, V., Lapusta, N., 2020. Rupture-dependent breakdown energy in fault models with thermo-hydro-mechanical
   processes. Solid Earth 11, 2283–2302. Https://doi.org/10.5194/se-11-2283-2020.
- Latour, S., Schubnel, A., Nielsen, S., Madariaga, R., Vinciguerra, S., 2013. Characterization of nucleation during
   laboratory earthquakes. Geophysical Research Letters 40, 5064–5069. https://doi.org/10.1002/grl.50974.
- Lee, H., Cho, T., 2002. Hydraulic characteristics of rough fractures in linear flow under normal and shear load. Rock
   Mechanics and Rock Engineering 35, 299–318. Https://doi.org/10.1007/s00603-002-0028-y.
- Madariaga, R., 2015. Seismic Source Theory, in: Treatise on Geophysics, Second Edition. Elsevier B.V.. chapter 4, pp. 51–71.
- Mogi, K., 1963. Some discussions on aftershocks, foreshocks and earthquake swarms the fracture of a semi-infinite
   body caused by an inner stress origin and its relation to the earthquake phenomena (3rd paper). Bulletin of the
   Earthquake Research Institute 41, 615–658.
- Nielsen, S., Taddeucci, J., Vinciguerra, S., 2010. Experimental observation of stick-slip instability fronts. Geophyscial Journal International 180, 697–702. Https://doi.org/10.1111/j.1365-246X.2009.04444.x.
- Noël, C., Passelègue, F., Giorgetti, C., Violay, M., 2019. Fault reactivation during fluid pressure oscillations: Tran sition from stable to unstable slip. Journal of Geophysical Research: Solid Earth 124. Https://doi.org/10.1029/
   2019JB018517.
- Ohnaka, M., 2003. A constitutive scaling law and a unified comprehension for frictional slip failure, shear fracture of
   intact rock, and earthquake rupture. Journal of Geophysical Research 108. Https://doi.org/10.1029/2000JB000123.
- Okubo, P., Dieterich, J., 1981. Fracture energy of stick-slip events in a large scale biaxial apparatus. Geophysical
   Research Letters 8, 887–890. Https://doi.org/10.1029/GL008i008p00887.
- Okubo, P., Dieterich, J., 1984. Effects of physical fault properties on frictional instabilities produced on simulated faults.
   Journal of Geophysical Research 89, 5817–5827.
- Paglialunga, F., Passelègue, F., Brantut, N., Barras, F., Lebihain, M., Violay, M., 2022. On the scale dependence in the dynamics of frictional rupture: Constant fracture energy versus size-dependent breakdown work. Earth and Planetary Science Letters 584. Https://doi.org/10.1016/j.epsl.2022.117442.
- Paglialunga, F., Passelègue, F., Latour, S., Gounon, A., Violay, M., 2023. Influence of viscous lubricant
   on nucleation and propagation of frictional ruptures. Journal of Geophysical Research: Solid Earth 128.
   Https://doi.org/10.1029/2022JB026090.
- Paglialunga, F., Passelègue, F., Lebihain, M., Violay, M., 2024. Frictional weakening leads to uncon ventional singularities during dynamic rupture propagation. Earth and Planetary Science Letters 626.
   Https://doi.org/10.1016/j.epsl.2023.118550.
- Palmer, A., Rice, J., 1973. The growth of slip surfaces in the progressive failure of over-consolidated clay. Proc. R. Soc.
   Lond. A 332, 527–548. Https://doi.org/10.1098/rspa.1973.0040.
- Passelègue, F., Schubnel, A., Nielsen, S., Bhat, H., Deldicque, D., Madariaga, R., 2016. Dynamic rupture processes inferred from laboratory microearthquakes. Journal of Geophysical Research: Solid Earth 121, 4343–4365.
   Https://doi.org/10.1002/2015JB012694.
- 732 Pratt, W.E., Johnson, D.W., 1926. Local subsidence of the Goose Creek oil field. The Journal of Geology 34, 577–590.
- Raleigh, C., Healy, J., Bredehoeft, J., 1976. An experiment in earthquake control at Rangely, Colorado. Science
   Https://doi.org/10.1126/science.191.4233.1230.
- Rice, J., 1980. The mechanics of earthquake rupture, in: Dziewonski, A., Boschi, E. (Eds.), Physics of the Earth's Interior. Società Italiana di Fisica, Bologna, Italy, pp. 555–649.
- Rosakis, A., Samudrala, O., Coker, D., 1999. Cracks faster than the shear wave speed. Science 284, 1337–1340.
   Https://doi.org/10.1126/science.284.5418.1337.
- Ross, Z., Cochran, E., Trugman, D., Smith, J., 2020. 3D fault architecture controls the dynamism of earthquake swarms.
   Science 368, 1357–1361. Https://doi.org/10.1126/science.abb0779.
- Roux, A., Leeman, E., Denkhaus, H., 1957. De-stressing: a means of ameliorating rockburst conditions. Part I—The
   concept of de-stressing and the results obtained from its application. The Journal of The South African Institute of
   Mining and Metallurgy , 101–119.
- Savitski, A., Detournay, E., 2002. Propagation of a penny-shaped fluid-driven fracture in an impermeable rock:
   asymptotic solutions. International Journal of Solids and Structures 39, 6311–6337. Https://doi.org/10.1016/S0020-7683(02)00492-4.
- Scholz, C., 1968. The frequency-magnitude relation of microfracturing in rock and its relation to earthquakes. Bulletin
   of the Seismological Society of America 58, 399–415.
- Schubnel, A., Nielsen, S., Taddeucci, J., Vinciguerra, S., Rao, S., 2011. Photo-acoustic study of sub shear and supershear ruptures in the laboratory. Earth and Planetary Science Letters 308, 424–432.

- 751 Https://dx.doi.org/10.1016/j.epsl.2011.06.013.
- Shlomai, H., Kammer, D., Adda-Bedia, M., Arias, R., Fineberg, J., 2021. Unstable cracks trigger
   asymptotic rupture models in bimaterial friction. Journal of the Mechanics and Physics of Solids 149.
   Https://doi.org/10.1016/j.jmps.2021.104330.
- Shuck, L., 1977. Method for selectively orienting induced fractures in subterranean Earth formations. U.S. Patent No.
   4005750.
- Stefanou, I., Tzortzopoulos, G., 2022. Preventing instabilities and inducing controlled, slow-slip in frictionally unstable
   systems. Journal of Geophysical Research: Solid Earth 127. Https://doi.org/10.1029/2021JB023410.
- Svetlizky, I., Albertini, G., Cohen, G., Kammer, D., Fineberg, J., 2020. Dynamic fields at the tip of sub Rayleigh and supershear frictional rupture fronts. Journal of the Mechanics and Physics of Solids 137.
   Https://doi.org/10.1016/j.jmps.2019.103826.
- Tada, H., Paris, P., Irwin, G., 2000. The Stress Analysis of Cracks Handbook, Third Edition. ASME Press.
- Tinti, E., Spudich, P., Cocco, M., 2005. Earthquake fracture energy inferred from kinematic rupture models on extended
   faults. Journal of Geophysical Research 110. Https://doi.org/10.1029/2005JB003644.
- Uenishi, K., Rice, J., 2003. Universal nucleation length for slip-weakening rupture instability under nonuniform fault
   loading. Journal of Geophysical Research 108. Https://dx.doi.org/10.1029/2001JB001681.
- Viesca, R., Garagash, D., 2015. Ubiquitous weakening of faults due to thermal pressurization. Nature Geoscience 8,
   875–879. Https://doi.org/10.1038/NGEO2554.
- Zang, A., Yoon, J., Stephansson, O., Heidbach, O., 2013. Fatigue hydraulic fracturing by cyclic reservoir treatment enhances permeability and reduces induced seismicity. Geophysical Journal International 195, 1282–1287.
   Https://doi.org/10.1093/gji/ggt301.
- 772 Zimmermann, G., Moeck, I., Blöcher, G., 2010. Cyclic waterfrac stimulation to develop an En-
- hanced Geothermal System EGS-Conceptual design and experimental results. Geothermics 39, 59–69.
- 774 Https://doi.org/10.1016/j.geothermics.2009.10.003.