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# The impacts of elevated CO<sub>2</sub> on forest growth, mortality, and recovery in the Amazon rainforest

Yitong Yao<sup>1,2</sup>, Philippe Ciais<sup>1</sup>, Emilie Joetzjer<sup>3</sup>, Wei Li<sup>4</sup>, Lei Zhu<sup>1,4</sup>, Yujie Wang<sup>2</sup>,  
Christian Frankenberg<sup>2,5</sup>, and Nicolas Viovy<sup>1</sup>

<sup>1</sup>Laboratoire des Sciences du Climat et de l'Environnement, LSCE/IPSL, CEA-CNRS-UVSQ,  
Université Paris-Saclay, 91191 Gif-sur-Yvette, France

<sup>2</sup>Division of Geological and Planetary Sciences, California Institute of Technology, Pasadena, CA 91125, USA

<sup>3</sup>Université de Lorraine, AgroParisTech, INRAE, UMR Silva, 54000 Nancy, France

<sup>4</sup>Department of Earth System Science, Ministry of Education Key Laboratory for Earth System Modeling,  
Institute for Global Change Studies, Tsinghua University, 100084 Beijing, China

<sup>5</sup>Jet Propulsion Laboratory, California Institute of Technology, Pasadena, CA 91109, USA

**Correspondence:** Yitong Yao (yyao2@caltech.edu)

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**Abstract.** The Amazon rainforest plays a crucial role in global carbon storage, but a minor destabilization of these forests could result in considerable carbon loss. Among the external factors affecting vegetation, elevated CO<sub>2</sub> (eCO<sub>2</sub>) levels have long been anticipated to have positive impacts on vegetation, including the direct enhancement of both photosynthesis and productivity and increasing water use efficiency. However, the overall impact of eCO<sub>2</sub> on the net carbon balance, especially concerning tree-mortality-induced carbon loss and recovery following extreme drought events, has remained elusive. Here, we use a process-based model that couples physiological CO<sub>2</sub> effects with demography and both drought mortality and resistance processes. The model was previously calibrated to reproduce observed drought responses of Amazon forest sites. The model results, based on factorial simulations with and without eCO<sub>2</sub>, reveal that eCO<sub>2</sub> enhances forest growth and promotes competition between trees, leading to more natural self-thinning of forest stands. This occurs following a growth–mortality trade-off response, although the growth outweighs the tree loss. Additionally, eCO<sub>2</sub> provides water-saving benefits, reducing the risk of tree mortality during drought episodes. However, extra carbon losses could still occur due to an eCO<sub>2</sub>-induced increase in background biomass density, leading to “more carbon available to lose” when severe droughts happen. Furthermore, we found that eCO<sub>2</sub> accelerates drought recovery and enhances drought resistance and resilience. By delving into the less-explored aspect of tree mortality response to eCO<sub>2</sub>, the model improvements advance our understanding of how carbon balance responds to eCO<sub>2</sub>, particularly regarding mechanisms of continuous competition-induced carbon loss vs. pulses of drought-induced carbon loss. These findings provide valuable insights into the intricate ways in which rising CO<sub>2</sub> influences forest carbon dynamics and vulnerability, offering a critical understanding of the Amazon rainforest’s evolution amidst more frequent and intense extreme climate events.

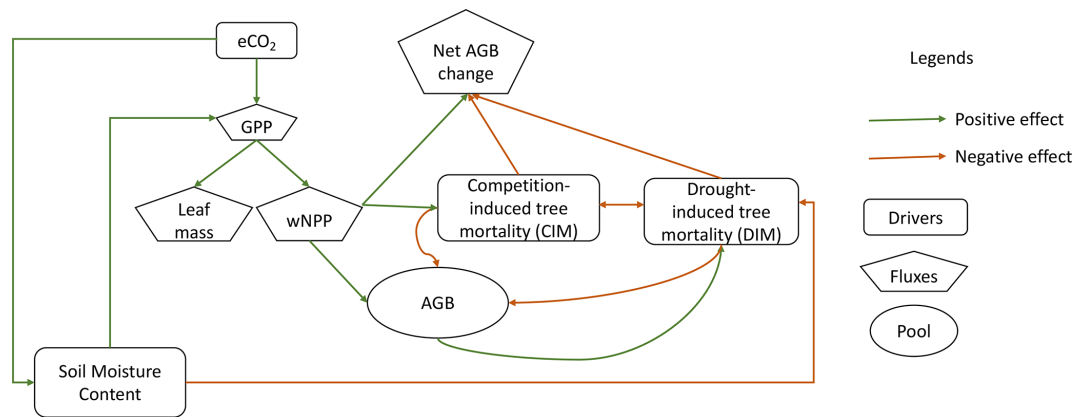
## 1 Introduction

The intact Amazon rainforest influences present and future global carbon dynamics, accounting for a carbon sink of 0.42–0.65 PgC yr<sup>-1</sup> between 1990–2007 (Pan et al., 2011) and containing 40 % of the world's tropical-forest above-ground biomass (Liu et al., 2015), a large carbon stock that is projected to be vulnerable to climate change (Boulton et al., 2022). Preserving this carbon stock is essential for regulating global CO<sub>2</sub> levels and stabilizing the Earth's climate. As climate change progresses and CO<sub>2</sub> levels rise, tropical rainforests may both increase carbon sequestration and become destabilized by climate risks. The impact of elevated CO<sub>2</sub> (eCO<sub>2</sub>) on carbon sequestration can be separated into two components: direct effects related to higher leaf carboxylation rates, which may translate into a higher leaf area index, higher tree productivity, and increased biomass (known as CO<sub>2</sub> fertilization), and indirect effects, i.e., partial stomatal closure and subsequently increased water use efficiency (WUE; CO<sub>2</sub> physiological forcing) (Smith et al., 2020). In turn, higher leaf area increases transpiration and can increase water stress. However, the potential translation of increased individual-tree growth rates into biomass accumulation at the ecosystem level remains uncertain, given that eCO<sub>2</sub> not only enhances carbon inputs at the ecosystem level but also amplifies carbon loss through growth–mortality trade-offs with higher growth, possibly leading to more competition between trees and higher mortality rates. Such “high gain/high loss” patterns reflecting a coupling between growth and mortality have been identified across spatial gradients (Needham et al., 2020; Stephenson and van Mantgem, 2005; Walker et al., 2021) and also seem to emerge in terms of temporal trajectories, with a parallel increase in growth and mortality, observed, for example, in the Amazon (Lewis et al., 2004) and other rainforests from long-term inventories (Hubau et al., 2020). For example, censuses of forest inventory plots within intact tropical forests in Amazonia have repeatedly revealed a faster increase in carbon losses from tree mortality that surpasses the increase in carbon gains attributed to both tree growth and new-tree recruitment, resulting in a decline in the net forest carbon sink (Hubau et al., 2020). Although a positive effect of eCO<sub>2</sub> on increased tree loss has been hypothesized by McDowell et al. (2022), establishing a significant correlation between carbon loss and eCO<sub>2</sub> has proven to be elusive (Hubau et al., 2020).

Compared to the research concentrating on vegetation productivity in response to eCO<sub>2</sub>, less attention has been directed toward the response of carbon loss, although minor disruptions of mortality rates in high-biomass systems, such as the intact Amazon rainforests, could trigger substantial carbon loss. An increase in tree mortality can reduce plant carbon residence time and consequently counteract the enhanced productivity (Friend et al., 2014). A comprehensive understanding of the response of tree mortality to eCO<sub>2</sub> is thus crucial for unraveling forest biomass carbon dynam-

ics. Carbon loss can arise from internal ecosystem processes, such as competition-induced self-thinning, the death of demoted trees, and the death of individual large trees (forming gaps), and external drivers, such as extreme climate events, insects, and pathogens (Fig. 1). McDowell et al. (2018) outlined two potential mechanisms underlying the connection between eCO<sub>2</sub> and an increased tree mortality rate. First, enhanced individual-tree growth rates could accelerate self-thinning due to increased competition. Second, eCO<sub>2</sub> makes trees larger and more vulnerable to the external environmental conditions of wind damage, drought, and heat (Gora and Esquivel-Muelbert, 2021; Maia et al., 2020). These two mechanisms correspond to competition-induced carbon loss and drought-induced carbon loss, and they pose threats to smaller trees and larger trees separately. However, eCO<sub>2</sub> also has the potential to promote tree survival by improving water use efficiency during drought (Brienen et al., 2017a, b; Van der Sleen et al., 2015). Liu et al. (2017) demonstrated, using simulations with a detailed soil–plant hydraulic model, that eCO<sub>2</sub> mitigates drought risks by decreasing the fraction of days when the daily minimum xylem water potential is below a critical threshold. Findings from a global model by Yao et al. (2023) simulating hydraulics and demography also indicated that drought exposure could be alleviated under eCO<sub>2</sub> in the Amazon. Besides, eCO<sub>2</sub> effects are also regulated by hydroclimatic conditions. Fatichi et al. (2016) revealed that indirect effects on productivity from eCO<sub>2</sub> tend to be more pronounced in water-limited ecosystems, although severe water stress can offset the expected CO<sub>2</sub> fertilization effects (Kolby Smith et al., 2016). The magnitude of the water-saving effect is also modulated by the intensity and duration of water stress events (Birami et al., 2020). Therefore, given the interplay of enhanced photosynthesis, heightened competition, vulnerability due to larger size, and mitigating effects from water-saving benefits, the impact of eCO<sub>2</sub> on carbon balance is not a straightforward monotonic relationship. The relative rates at which gross carbon fluxes change with eCO<sub>2</sub> play a crucial role in determining the net changes in aboveground biomass (AGB) (Fig. 1).

Given the unavailability of free-air CO<sub>2</sub> enrichment (FACE) experiments and the scarcity of in situ measurements within the complex Amazon rainforest, employing process-based modeling emerges as a promising approach for investigating how eCO<sub>2</sub> influences the accumulation and loss of biomass. Koch et al. (2021) demonstrated that models participating in Coupled Model Intercomparison Project 5 (CMIP5) and CMIP6 could reproduce the response of carbon gains attributed to tree growth in response to environmental drivers but were rather unable to reproduce the carbon losses observed in inventory data. This model shortcoming primarily results from the fact that CMIP5- and CMIP6-participating models do not include processes related to tree competition and that most of them ignore climate-induced mortality processes, although a new generation of global models are under development to address these re-



**Figure 1.** A schematic plot illustrating the influence of  $e\text{CO}_2$  on forest biomass carbon balance. A positive effect means that the impacted variable is expected to increase under  $e\text{CO}_2$  compared to pre-industrial  $\text{CO}_2$ . A positive effect on net AGB change means that  $e\text{CO}_2$  leads to a higher net carbon gain for AGB. When  $\text{CO}_2$  concentration rises, water use efficiency increases due to partial closure of stomata. Consequently, soil water consumption decreases, leading to an increase in soil moisture content (positive effect on soil moisture content (SMC)). This increase in SMC offers a degree of mitigation against drought exposure, known as water-saving benefits (negative effect on drought-induced mortality (DIM)). Simultaneously,  $e\text{CO}_2$  enhances photosynthesis, resulting in greater carbon gains (positive effect on canopy leaf mass and woody net primary production (wNPP)). This enhanced tree growth also intensifies competition, leading to natural self-thinning (positive effect on competition-induced mortality (CIM)). While a reduction in drought exposure through higher SMC offers a reduction in DIM, the presence of “more carbon available to lose” under  $e\text{CO}_2$  contributes to increased drought-induced biomass carbon loss due to drought mortality (negative effect of DIM on AGB). The combined effects of enhanced carbon gain, reduced drought exposure, and higher biomass density can influence net biomass carbon change either positively or negatively. GPP: gross primary productivity.

search gaps (Eller et al., 2020; Koven et al., 2020; Yao et al., 2022). Yu et al. (2022) underscored this issue and showed that biomass loss due to the tree mortality was overestimated in dynamic global vegetation models (DGVMs) when compared to historical forest inventory measurements and that incorporating observation-based constraints into these models would lead to a reduction in carbon sink predictions by the end of the century. Therefore, conducting modeling studies with realistic representations of tree mortality and incorporating observation-based constraints are crucial steps toward achieving a more reliable projection of the evolution of carbon stocks in intact Amazon forests.

As the representation of mortality in most process-based models is based on prescribing a fixed-loss fraction of standing biomass stocks (Adams et al., 2013), there is a clear need for a more realistic simulation of tree-mortality-induced carbon loss. In Yao et al. (2023), an empirical hydraulic-failure module and a light-competition tree mortality module in the ORCHIDEE (Organising Carbon and Hydrology In Dynamic Ecosystems) land surface model were tested over the Amazon rainforest. This model has been calibrated to reproduce tree-size-dependent mortality rates at a site in Caxiuana (a long-term throughfall exclusion experiment) and has proven effective in reproducing the increasing carbon loss due to tree mortality rate and the resulting basin-scale deceleration in the net carbon sink, observed in inventories from Hubau et al. (2020).

In this study, we explore the impact of  $e\text{CO}_2$  on forest growth, tree mortality, and drought recovery in the Ama-

zon rainforest. Our analysis leverages the newly upgraded process-based model ORCHIDEE-CAN-NHA r7236 with competition- and drought-induced mortalities, following the methodology outlined in Yao et al. (2023). We conducted two factorial simulations – one with rising  $\text{CO}_2$  levels since 1901 and one without. The three key specific questions we address here are as follows. (i) Does  $e\text{CO}_2$  lead to a greater increase in tree mortality compared to productivity? (ii) Does  $e\text{CO}_2$  promote carbon loss more during wet years compared to dry years? (iii) Does  $e\text{CO}_2$  alleviate the impact of drought on net carbon balance of AGB and benefit drought recovery? Our hypothesis is that  $e\text{CO}_2$  leads to a smaller increase in tree mortality compared to carbon gain and that this net benefit for AGB changes is greater during dry years, contributing to accelerated drought recovery. The process-based model ORCHIDEE-CAN-NHA has been well calibrated. In brief, this model incorporates a mechanistic plant hydraulic architecture simulating water potentials at 30 min intervals within the soil–root–stem–leaf continuum. It also includes a drought-exposure-related tree mortality scheme and accounts for size-dependent tree mortality rates under exposure conditions. For a detailed model description, calibration information, and validation against observed datasets, please refer to Yao et al. (2022) and Sect. 2.

## 2 Materials and methods

### 2.1 Model description

ORCHIDEE-CAN-NHA incorporates a plant hydraulic architecture that enables the modeling of water potential and hydraulic conductance along the vertical profile of plants. This module considers both vertical water flow driven by water potential gradients and the movement of water into or out of water storage pools, regulated by water capacitance. By simulating the plant hydrodynamics, we derive a critical indicator of plant water stress: the percentage loss of conductance (PLC). PLC has been demonstrated to correlate with tree mortality (Choat et al., 2012), e.g.,  $\psi_{50}$ , which represents the water potential at which 50% of conductance is lost. Building on the simulation of  $\psi_{50}$ , we have integrated an empirical tree mortality module that is based on drought exposure, which can help reproduce the size-dependent tree mortality pattern of higher tree mortality rates in cohorts with a larger circumference class. Within this framework, two crucial empirical parameters have been introduced: the drought exposure threshold and the fraction of tree mortality once this threshold is reached. These two parameters were calibrated using observed water potentials, sap flux transpiration, and stem mortality rates from a long-term throughfall exclusion experiment conducted at the Caxiuanã site located in the northeastern Amazon (Yao et al., 2022). The calibrated model has proven accurate in capturing the sensitivity of carbon fluxes to drought and the long-term trends in net carbon sink dynamics. This was demonstrated by comparing the simulated sensitivity of biomass loss rates and growth rates to water deficit against plot observations for the droughts of 2015 and 2010 (Yao et al., 2023). In addition to drought-induced tree mortality, ORCHIDEE also parametrizes the light-competition-induced self-thinning process (Joetzjer et al., 2022), which accounts for competition-induced tree mortality. The self-thinning process in ORCHIDEE is regulated by the quadratic mean diameter, where smaller trees are killed in order of priority.

As ORCHIDEE is a cohort-based model, we obtain woody carbon gain, woody carbon loss, and biomass carbon pools for 20 cohorts, associated with increasing circumference/diameter classes from small trees to large trees. Carbon gain in our model refers to wNPP, specifically that for cohorts with a diameter above 10 cm, aligning with inventory protocols. Carbon loss represents the amount of live biomass (with a diameter > 10 cm) that is transferred to the woody litter pool due to tree mortality, resulting from events of continuous competition-induced mortality (killing small trees) and events regarding pulses of drought-induced mortality (killing large trees). Then we aggregate the grid-level carbon gain and carbon loss to the basin level, following the approach used by Brienen et al. (2015). Observational time series of carbon gains, losses, and the net carbon balance for Amazonian forests are obtained from Brienen et al. (2015). To

gain a deeper insight into how eCO<sub>2</sub> impacts carbon loss, we examined changes in both competition-induced tree mortality (self-thinning (CIM)) and drought-induced tree mortality (DIM) as distinct components. For drought mortality, we compared drought exposure under constant CO<sub>2</sub> to assess how eCO<sub>2</sub> alleviates the risk of tree mortality from hydraulic failure.

Following a TRENDY-type protocol (Seiler et al., 2022), we have implemented two distinct scenarios in our study. The first scenario maintains a constant CO<sub>2</sub> concentration at the 1901 level while varying climate forcing (A1), whereas the second scenario permits variations in both CO<sub>2</sub> concentration and climate forcing (A2). Here, A2 is similar to S2 from the TRENDY protocol despite the fact that we did not consider land cover change.

### 2.2 Climate-forcing data

The gridded climate-forcing dataset employed is CRUJRA v2.1 (Harris et al., 2020), used in the TRENDY simulations. CRUJRA v2.1 was created by re-gridding data taken from Japanese Reanalysis (JRA) data, a product of the Japanese Meteorological Agency. This re-gridded dataset was adjusted to align with monthly observation-based Climatic Research Unit (CRU) TS4.04 data (Harris et al., 2020). CRUJRA v2.1 provides 6 h meteorological variables from January 1901 to December 2019 at a spatial resolution of 0.5°.

### 2.3 Drought characteristics

Following Papastefanou et al. (2022), for the evaluation of drought area and severity, the maximum cumulative water deficit (MCWD) was used to compare droughts, as seen in Eqs. (1) and (2), with a fixed value for evapotranspiration (ET) of  $\sim 100$  mm per month being used (Phillips et al., 2009). When monthly rainfall ( $P_m$ ) is below 100 mm, the forest undergoes a water deficit. This water deficit accumulates over the hydrological year from October in the previous year to September in the current one. The MCWD is the most negative value of the water deficit across all 12 months.

$$\text{CWD}_m = \text{CWD}_{m-1} + (P_m - 100) \text{ if } P_m < 100, \\ \text{else } \text{CWD}_m = 0, \quad (1)$$

where  $m$  is the month from October to September.

$$\text{MCWD} = \min(\text{CWD}_m), m = 1, \dots, 12. \quad (2)$$

Then the decadal mean of the MCWD over the entire period ( $\mu_{\text{MCWD}}$ ) was subtracted from the MCWD of a year with drought ( $\text{MCWD}_i$ ), giving an MCWD anomaly (Eq. 3).

$$\text{MCWD anomaly} = \text{MCWD}_i - \mu_{\text{MCWD}} = 1980\text{--}2019 \quad (3)$$

We derived  $Z$  scores of MCWD time series at the annual scale by following Eq. (4) from Feldpausch et al. (2016), cal-



culated according to

$$Z_{\text{MCWD}} = \frac{\text{MCWD}_i - \mu_{\text{MCWD}}}{\sigma_{\text{MCWD}}} \quad i = 1980\text{--}2019, \quad (4)$$

where  $\sigma_{\text{MCWD}}$  is the interannual standard deviation of the MCWD.

## 2.4 Drought resistance and resilience

For each drought event, drought *resistance* is defined as the change in the net biomass carbon sink during drought disturbance relative to its pre-drought state. A positive value indicates that drought conditions lead to an increase in the net carbon sink relative to non-stressed conditions, while negative values indicate a decrease in the net biomass carbon sink. A more negative value indicates higher vulnerability. Drought *resilience* refers to the ability of the net carbon sink to recover to the pre-drought state. It is computed as the difference in the net carbon sink between the post-drought period and the pre-drought state relative to the pre-drought period. Positive values indicate full recovery, where the net carbon sink after drought stress surpasses the pre-drought state, while negative values indicate incomplete recovery. A more negative ratio represents a more limited capacity for recovery. The calculations for drought resistance and resilience of net biomass carbon change followed Tao et al. (2022). We also used the net biomass carbon balance from 2 years before and 2 years after the drought event to represent pre- and post-drought forest conditions, respectively (Tao et al., 2022). Resistance and resilience were calculated for each pixel and for all the drought events during the past 4 decades, following Eqs. (1) and (2), and were reported at the basin scale by taking the median value across drought-affected pixels ( $Z_{\text{MCWD}}$  below  $-1$ ).

$$\text{Drought resistance} = \frac{Y_e - Y_{\text{pre}}}{Y_{\text{pre}}} \quad (5)$$

$$\text{Drought resilience} = \frac{Y_{\text{post}} - Y_{\text{pre}}}{Y_{\text{pre}}} \quad (6)$$

Here,  $Y_{\text{pre}}$  is the pre-drought value of net biomass carbon change (carbon gain (CG) minus carbon loss (CL)),  $Y_{\text{post}}$  is the post-drought value of net biomass carbon change, and  $Y_e$  is the signal during the drought event for net biomass carbon change.

## 3 Results

### 3.1 The mean carbon gains and carbon losses over the Amazon rainforest

Our model simulations over the past 4 decades, driven by varying CO<sub>2</sub> and climate forcing, reveal that CG slightly surpasses CL at the basin scale, resulting in a positive net carbon balance (“Net” in Fig. 2a). When we separate years

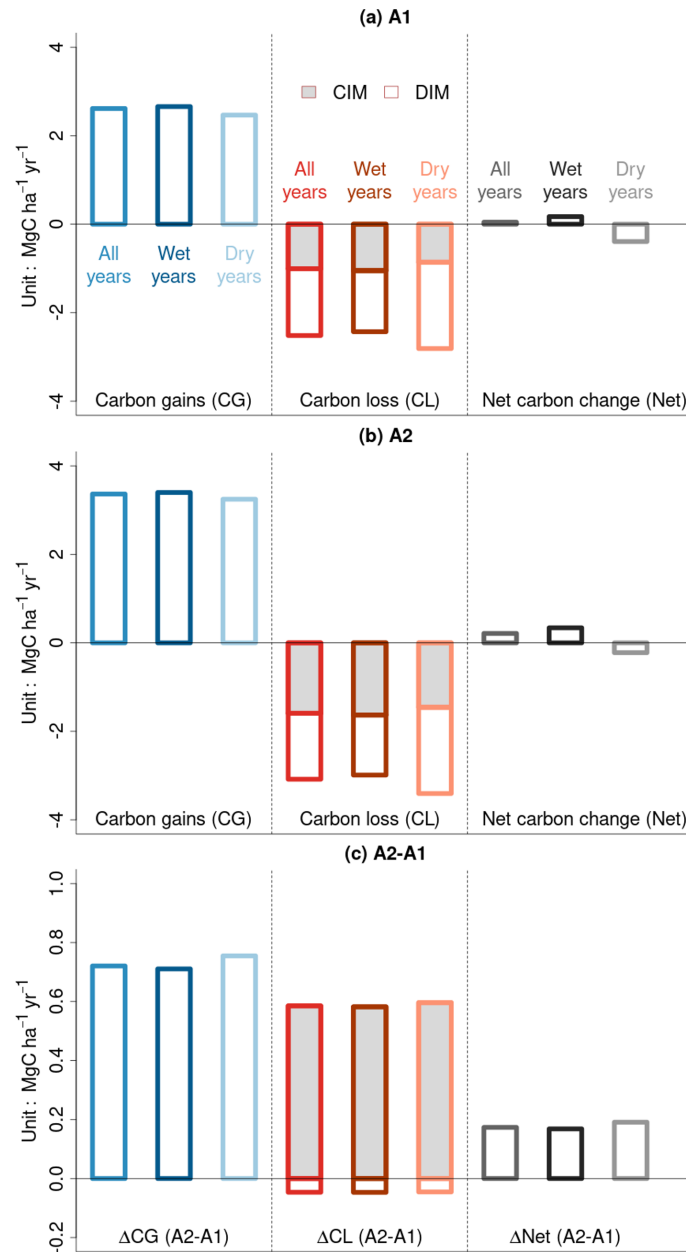
into wet and dry categories based on the basin-scale median of Z-transformed MCWD ( $Z_{\text{MCWD}} > 0$  for wet years and  $Z_{\text{MCWD}} < 0$  for dry years), we find a net carbon sink during wet years and a net carbon source during dry years. This pattern arises because CG is lower and CL is higher during dry years compared to during wet years and vice versa. During dry years, competition-induced carbon loss (CIM) is a bit lower compared to during wet years, and increased drought-induced tree mortality (DIM) leads to higher CL. Under constant CO<sub>2</sub> concentration conditions (A1 scenario), both CG and CL decrease compared to the A2 scenario, with the reduction in CG being more pronounced, resulting in a much smaller net carbon sink than under eCO<sub>2</sub> conditions (Fig. 2b). Carbon loss from CIM also decreases during dry years compared to during wet years, and its fraction in the total carbon loss becomes lower as increases in drought-induced tree dieback can suppress self-thinning. When comparing the model simulations with and without eCO<sub>2</sub>, higher  $\Delta\text{CG}$  during dry years compared to during wet years suggests the CO<sub>2</sub> fertilization effect is more efficient during dry years (Fig. 2c).  $\Delta\text{CL}$  is primarily affected by CIM, even though  $\Delta\text{CIM}$  and  $\Delta\text{DIM}$  have opposing effects on it (Fig. 2c). Our model simulation thus implies that the CO<sub>2</sub> fertilization effect plays a dominant role in augmenting aboveground forest productivity (carbon gains) and (to a lesser extent) biomass loss rates from mortality. Our estimate falls within the upper range of the trend distribution, which is consistent with existing studies on the effects of eCO<sub>2</sub>, including those employing process-based models, analytical solutions, and ecological optimality theory (Table S1 in the Supplement).

### 3.2 Effect of eCO<sub>2</sub> on the trends of the net biomass carbon sink, carbon gains, and carbon losses

Our A2 simulation, which accounts for varying climate and CO<sub>2</sub> concentration, shows a decline in the net biomass carbon sink since 1980, decreasing at a rate of  $0.006 \text{ MgC ha}^{-1} \text{ yr}^{-2}$  ( $6 \text{ kgC ha}^{-1} \text{ yr}^{-2}$ ). This decelerating trend can be predominantly attributed to the increase in carbon loss resulting from tree mortality, which amounts to  $0.014 \text{ MgC ha}^{-1} \text{ yr}^{-2}$ , surpassing the enhanced carbon gain trend of  $0.008 \text{ MgC ha}^{-1} \text{ yr}^{-2}$ . The trend of the biomass sink in the A2 scenario has the same sign as that in forest inventory records, but it is 60 % smaller in magnitude (see Fig. 3). When CO<sub>2</sub> concentration is held constant, the A1 scenario indicates a larger decline in the net carbon sink. This more negative trend is primarily driven by reduced carbon gains, while carbon loss increases less.

### 3.3 Effect of eCO<sub>2</sub> on competition-induced and drought-induced carbon losses

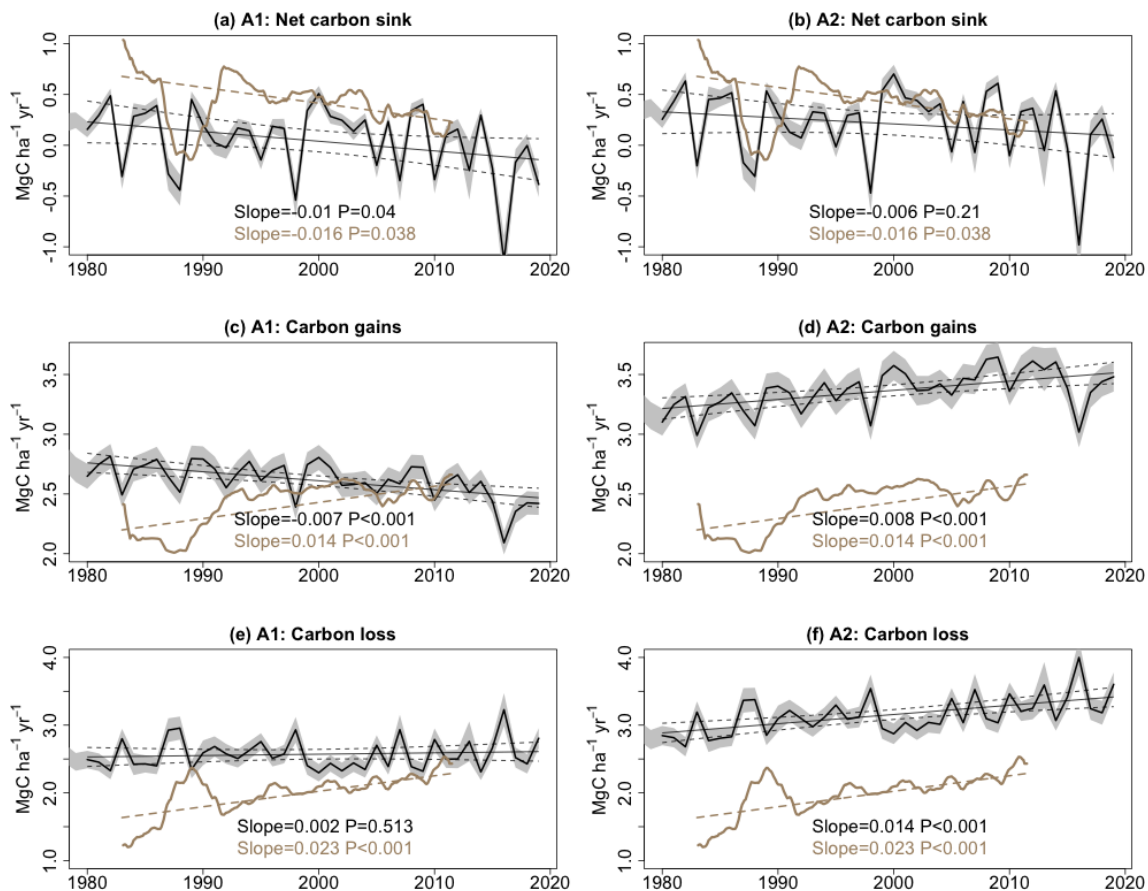
As described in Sect. 3.1, carbon loss resulting from tree mortality can be categorized into two distinct processes:



**Figure 2.** Basin-scale averages of forest biomass carbon gains (CG), forest biomass carbon losses (CL), and the net balance of gains minus losses (Net) over the last 4 decades. **(a)** eCO<sub>2</sub>. **(b)** Constant CO<sub>2</sub> at pre-industrial levels. **(c)** Difference between eCO<sub>2</sub> and constant CO<sub>2</sub> ( $\Delta$ ). By convention, gains are positive and losses are negative in **(a)** and **(b)**. In panel **(c)**,  $\Delta\text{CL}$  is calculated as the difference between the absolute values of CL in panels **(a)** and **(b)**. Dry years are defined as those in which the median of the Z-transformed cumulative water deficit ( $Z_{\text{MCWD}}$ ) at the basin scale falls below 0. For biomass carbon loss, CIM represents competition-induced self-thinning processes, and DIM represents drought-induced tree mortality processes.

competition-induced mortality (CIM) and drought-induced mortality (DIM). Figure 4 illustrates the trends in carbon loss attributed to these two processes. In the A2 scenario, simulated CIM displays no significant trend (slope =  $0.001 \text{ MgC ha}^{-1} \text{ yr}^{-2}$ ;  $P = 0.738$ ). However, when CO<sub>2</sub> remains constant in the A1 scenario, this term exhibits a notable decrease (slope =  $-0.01 \text{ MgC ha}^{-1} \text{ yr}^{-2}$ ;  $P < 0.001$ ).

In contrast, both the A1 and A2 scenarios exhibit significant increasing trends in DIM (slope =  $0.013 \text{ MgC ha}^{-1} \text{ yr}^{-2}$ ;  $P < 0.05$ ). Consequently, the lack of a significant overall trend in total carbon loss in the A1 scenario can be attributed to the opposing effects of CIM and DIM. Our model simulations reaffirm that the increasing carbon loss in the A2 scenario is primarily a result of higher drought-induced tree



**Figure 3.** Trends in (a, b) the net carbon sink for aboveground live biomass, (c, d) carbon gains, and (e, f) carbon losses from tree mortality; these were obtained from ORCHIDEE model simulations (black lines) and forest inventory data (brown lines). Shading represents the 95 % confidence interval. The slopes and associated *P* values are from linear regression models. It should be noted that the number of measurements for each year in the inventory varies and a “linear mixed-effects model” was used to account for the weight associated with different sampling plot areas and their monitoring time lengths. Therefore, in the inventory-pooling results, the trend of the net carbon sink is not equal to the difference between trends in carbon gain and carbon loss.

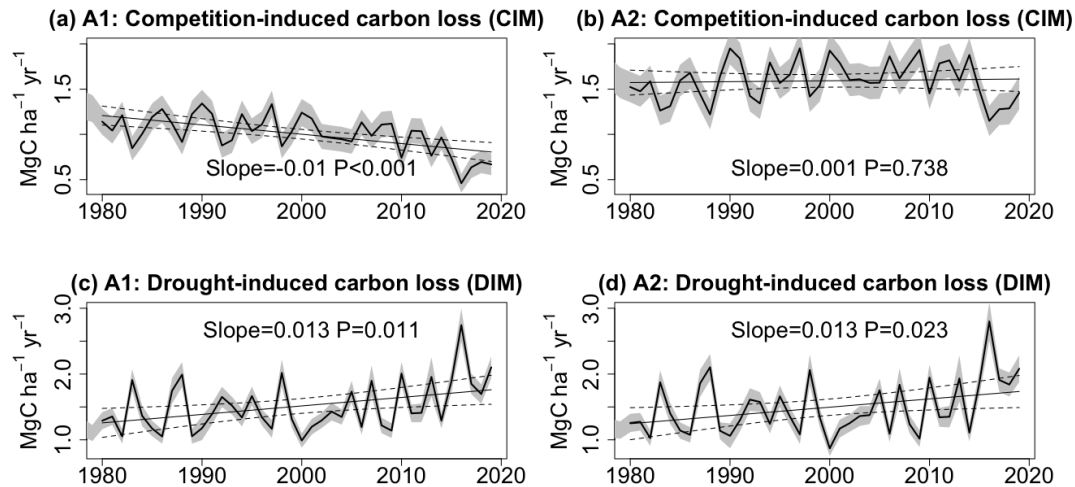
mortality. Without the sustained increase in CO<sub>2</sub>, the net carbon sink would have diminished at a faster pace (slope = -0.01 MgC ha<sup>-1</sup> yr<sup>-2</sup> in the A1 scenario, and slope = -0.006 MgC ha<sup>-1</sup> yr<sup>-2</sup> in the A2 scenario).

### 3.4 Water conditions mediate biomass carbon flux responses to eCO<sub>2</sub>

The impact of eCO<sub>2</sub> on biomass carbon sinks is influenced by prevailing water conditions. To explore how hydroclimatic conditions regulate the impact of eCO<sub>2</sub> on carbon gain and carbon loss, we focused on three recent megadrought events (2005, 2010, and 2015/2016). In line with the methodology applied by Pan et al. (2022) to eCO<sub>2</sub> vs. control experiments, we used the ratio between the enhanced CO<sub>2</sub> scenario (A2) and the constant CO<sub>2</sub> scenario (A1) to assess the relative response (*R*) of the ecosystem level to carbon gain (*R*<sub>CG</sub>) and carbon loss (*R*<sub>CL</sub>). During these three drought events, we found that forests in drier climate zones (a more nega-

tive MCWD) exhibited greater *R*<sub>CG</sub> compared to their wetter counterparts (Fig. 5), and this model response prevails across all cohorts, with larger-sized cohorts showing lower *R*<sub>CG</sub> and less negative sensitivity of *R*<sub>CG</sub> to the MCWD due to more carbon allocation to smaller cohorts (Fig. S1 in the Supplement). Interestingly, *R*<sub>CL</sub> does not show a monotonic change from small to large cohorts, even though the average over the drought epicenter indicates higher *R*<sub>CL</sub> in smaller cohorts (Fig. S2). Self-thinning may not always occur due to suppression by DIM. Therefore, even though higher *R*<sub>CG</sub> is found under eCO<sub>2</sub> along the water stress gradient, self-thinning does not always change coordinately. Here, *R*<sub>CL</sub> is mainly contributed by drought-induced carbon loss (*R*<sub>CL\_DIM</sub>), where self-thinning is suppressed. However, *R*<sub>CL\_DIM</sub> does not exhibit a significant correlation with the MCWD (Fig. 5). It is worth noting that DIM-induced carbon loss is influenced by two key factors: background biomass density and tree mortality rate. The former one is boosted by eCO<sub>2</sub>, indicating “more carbon available to lose” (Fig. 5),





**Figure 4.** Trends in carbon losses due to (a, b) competition-induced tree mortality (CIM) and (c, d) drought-induced tree mortality (DIM) with and without eCO<sub>2</sub>.

while the response of the latter is the opposite because eCO<sub>2</sub> leads to a reduction in stomatal conductance and transpiration, alleviating water stress, which is indicated by fewer drought exposure days (Fig. S3) and a lower fraction of trees killed due to DIM in the A2 scenario than in the A1 scenario across most regions within the epicenter of the drought events ( $Z_{MCWD} < -1$ ; Fig. 5). This suggests that eCO<sub>2</sub> has a positive impact on mitigating the effects of drought on biomass loss driven by DIM.

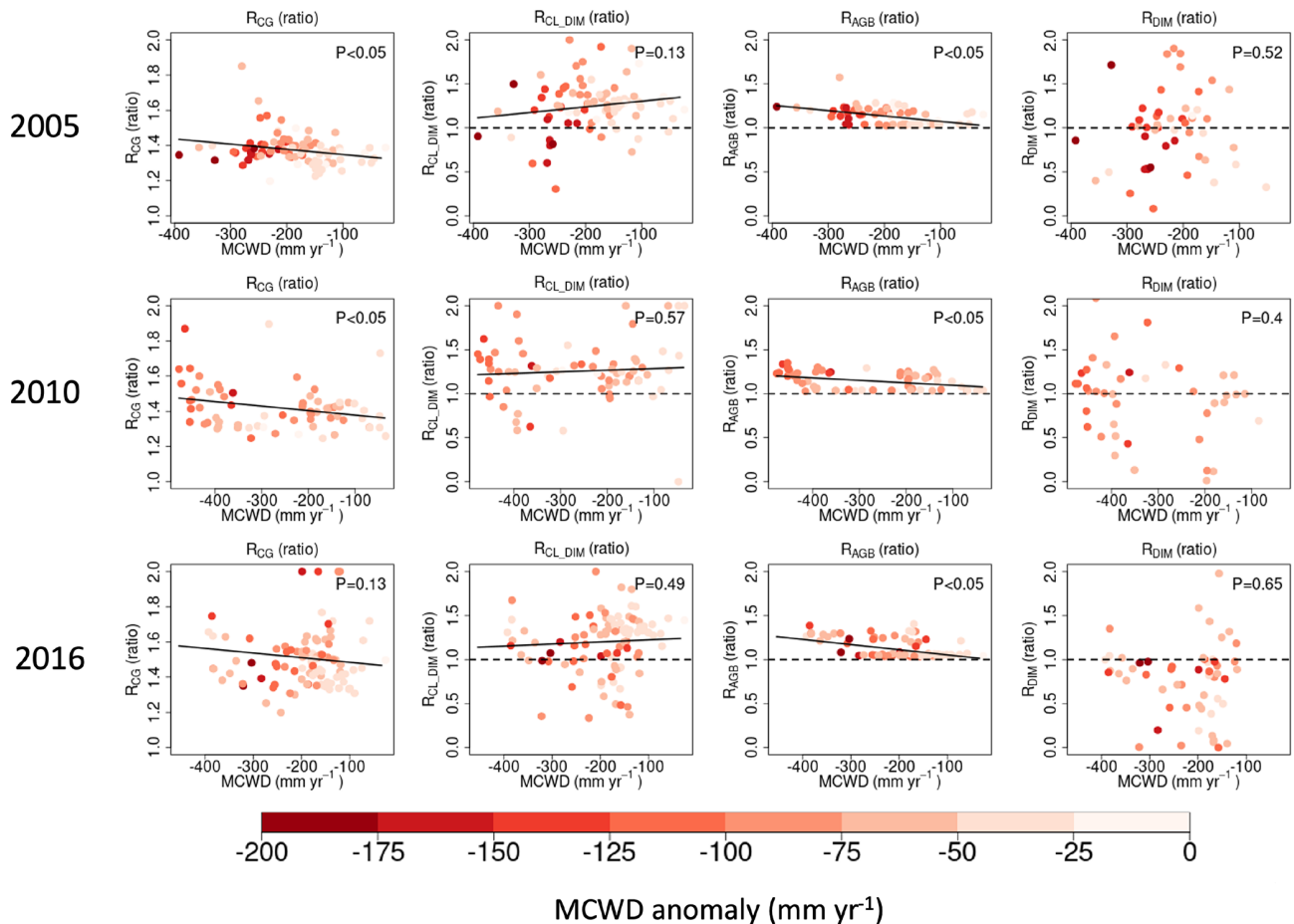
### 3.5 Effect of eCO<sub>2</sub> on drought recovery

After a drought, eCO<sub>2</sub> supports faster or more complete recovery due to growth enhancement (De Kauwe et al., 2021). eCO<sub>2</sub> can also enable deeper root depth and thus better access to deeper soil moisture (Iversen, 2010). In addition to investigating the impact of eCO<sub>2</sub> on carbon loss during droughts, our study delved into whether eCO<sub>2</sub> confers benefits to drought recovery, considering the likely promoted growth and two different tree mortality regimes of CIM and DIM. To identify past drought events, we calculated the MCWD for each year and each pixel. Droughts were defined as pixels in a year with an MCWD Z score falling below  $-1$  since 1980. This threshold, while not extremely severe, was chosen to ensure enough pixels for analysis. Following the methodology outlined by Tao et al. (2022), we aggregated the pixel-level results to the basin scale by calculating the median value across all pixels. Considering the findings of Tao et al. (2022), which indicated increased drought resistance across a gradient of drought severity using C-band radar signals, we examined whether the process-based model could reproduce this drought response. Figure 6 illustrates the relationship between drought resistance and resilience of net biomass carbon change in relation to drought severity, expressed as  $Z_{MCWD}$ . In both the A1 and A2 scenarios,

we observed a well-expected declining trend in resistance as drought severity increased. When drought severity was defined using the absolute value of the MCWD anomaly, a similar declining trend of resistance for net carbon change was found (Fig. S4). Notably, the rainforests did not fully recover to their pre-drought conditions following drought events of a higher severity – in the A1 scenario,  $62 \pm 12\%$  of the area located in the drought epicenter showed resilience below 0, whereas in the A2 scenario, this was  $59 \pm 16\%$ . When examining the differences in drought resistance and resilience between the A1 and A2 scenarios, our findings indicate that eCO<sub>2</sub> enhances both drought resistance and resilience. For drought-sensitive areas where the tree mortality regime shifts from CIM to DIM, their drought resistance and resilience are lower than in insensitive areas due to higher carbon loss. This suggests that elevated CO<sub>2</sub> levels contribute to improved forest resilience and enhance the ability of forests to withstand and recover from drought events.

## 4 Discussion

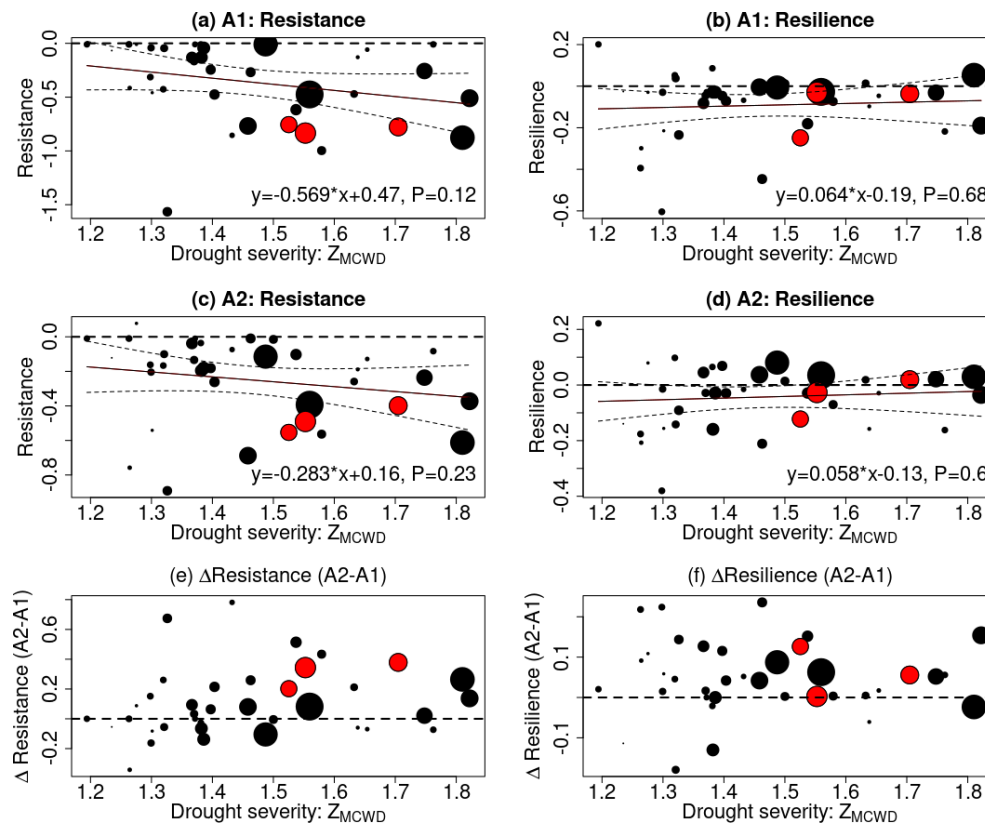
The CO<sub>2</sub> fertilization effect has gained widespread recognition as a primary driver of vegetation greening, observed as an increase in leaf area index (LAI) in most regions around the world (Zhu et al., 2016). Our model, forced by eCO<sub>2</sub>, also captures LAI greening in most areas over the basin, while a constant CO<sub>2</sub> setup results in browning (Fig. S5). However, it is important to clarify that this greening phenomenon does not necessarily translate into increased biomass accumulation because the carbon allocation shift and processes that control biomass loss can be decoupled from the increasing trend of LAI and productivity (Fan et al., 2023). The dynamics of the net carbon sink are fundamentally shaped by the balance between productivity and carbon loss. The concept of a growth–mortality trade-off, or high gain/high loss, has



**Figure 5.** The effect of eCO<sub>2</sub> on carbon gain ( $R_{CG}$ ), carbon loss due to drought-induced tree mortality ( $R_{CL\_DIM}$ ), AGB ( $R_{AGB}$ ), and the proportion of trees affected by DIM ( $R_{DIM}$ ) in relation to water conditions and drought intensity in the years 2005, 2010, and 2016. Water conditions are characterized by the MCWD during the drought year (on the horizontal axis), where negative values indicate a drier climate. Dots are color-coded to reflect the drought intensity characterized by the MCWD anomaly, with darker colors indicating more severe water deficits. The dots shown in the panels correspond to pixels located in the epicenter of the drought, featuring  $Z_{MCWD}$  values below  $-1$ . This threshold is set to ensure an adequate number of pixels for our analysis.

been observed across spatial gradients (Wright et al., 2010), but it remains more difficult to prove in terms of expressed temporal changes as “faster gains imply higher losses”. This is because one can observe an acceleration in both gains and losses, as noted by Hubau et al. (2020), but a causal relation between the two cannot be empirically proven. A demography model including more detailed stand dynamics could help test this hypothesis. eCO<sub>2</sub> enhancing tree competition in wet tropical forests has the potential to couple faster gains with higher losses and thus may not universally lead to a boost in the net carbon sink due to an accelerated carbon turnover rate (Walker et al., 2021). Future CO<sub>2</sub> fertilization could potentially increase not only recruitment and growth but also tree mortality (McDowell et al., 2020). In our study, we conducted model experiments to disentangle the effect of eCO<sub>2</sub> on forest biomass carbon fluxes, including carbon gain, carbon loss, and net carbon balance. Our model simula-

tions revealed a compelling insight. When we deactivated the CO<sub>2</sub> fertilization effect, the model simulated an LAI browning trend (Fig. S5) and a declining trend in carbon gain compared to forest inventory observations (Fig. 2). These findings align with a similar compensatory effect on LAI from rising CO<sub>2</sub> in TRENDY models for tropical forests, as noted by Winkler et al. (2021). Our results also underscored that eCO<sub>2</sub> not only drives enhanced biomass carbon gains but also plays a pivotal role in shaping carbon losses (Figs. 3 and 4). Specifically, turning off the eCO<sub>2</sub> effect leads to a dampened increase in carbon loss. This attenuation is primarily attributable to the muted response of natural self-thinning-related tree mortality, which is determined by less growth (i.e., the deceleration of carbon gain). Meanwhile, the more negative net carbon sink trend observed in the A1 scenario is primarily driven by the deceleration of carbon gain. The basin-scale average shows that CL\_DIM (carbon loss driven



**Figure 6.** The effect of eCO<sub>2</sub> on drought recovery in terms of resistance and resilience of net biomass carbon change. For each drought event, the median resistance, resilience, and drought severity of the drought-affected pixels ( $Z_{MCWD}$  below  $-1$ ) were calculated (shown as dots in each panel). Drought severity was defined by the  $Z_{MCWD}$ . The size of the dots indicates the area within the drought epicenter. The dots for 2005, 2010, and 2016 are color-coded in red. A trend fit was applied to panels (a)–(d), with its equation labeled and 95 % confidence interval marked by dashed lines. Resistance was computed, where lower values mean a more pronounced reduction in the net biomass carbon sink during a drought. Resilience was also computed, where positive values indicate complete recovery of the forest to pre-drought conditions, while negative values signify incomplete recovery. In panels (a)–(d), a horizontal dashed line indicates resistance and resilience values of zero. In panels (e)–(f), a horizontal dashed line was drawn to denote no change in resistance or resilience under eCO<sub>2</sub> conditions.

by drought-induced tree mortality) is lower under eCO<sub>2</sub> but that CL\_ST (carbon loss driven by self-thinning) is highly promoted under eCO<sub>2</sub>. Thus,  $\Delta CL_{DIM}$  and  $\Delta CL_{ST}$  oppose each other, although  $\Delta CL_{ST}$  dominates the magnitude of  $\Delta CL$  (Fig. 1c). Such net forest biomass carbon loss in the absence of eCO<sub>2</sub> was similar to the findings of de Almeida Castanho et al. (2016), in which multiple model simulations were used. They showed a decline in biomass over the past several decades when not considering eCO<sub>2</sub> effects, although tree mortality had not been incorporated yet. Our work advances this understanding by separating the contributions of competition-induced and drought-induced tree carbon loss and providing evidence that eCO<sub>2</sub> benefits the net forest biomass carbon sink over the Amazon rainforest, although such benefit may still not be sufficient to offset the carbon loss caused by prolonged external climate stressors, such as long-term temperature-induced carbon loss (Sullivan et al., 2020).

The response of biomass carbon flux dynamics to eCO<sub>2</sub> is intricately linked to water availability. Taking the years 2005, 2010, and 2016 as examples, the enhancement ratio of carbon gain ( $R_{CG}$ ) exhibits a negative correlation with the water deficit metric of the MCWD, indicating that the eCO<sub>2</sub> effect is more pronounced in drier regions. Pan et al. (2022) demonstrated that extra-tropical woody ecosystems characterized by drier baseline climates tend to exhibit a higher average enhancement in aboveground carbon gain in response to eCO<sub>2</sub>. Our model simulations suggest that a similar pattern may persist in wet tropical forests within the Amazon as well. The negative interactions found between eCO<sub>2</sub> and water deficit suggest that vegetation in drier climates can benefit more from the combination of enhanced photosynthesis, reduced photorespiration, and increased WUE. Alongside potential growth advantages, the tree-ring-width data analysis by Zuidema et al. (2020) reveals that eCO<sub>2</sub>-induced partial stomata closure and reduced transpiration may attenuate the cooling effect on leaf surfaces, potentially pushing

leaf surface temperatures beyond the optimal range for photosynthesis. Our modeling results, showing enhanced carbon gain over the Amazon rainforest, suggest that temperature may not have exceeded the optimal range in this case, given the mean annual land surface temperature under eCO<sub>2</sub> increases by around 0.26 °C at most during the 2015–2016 El Niño (Fig. S6). Nonetheless, the potential shift in sensitivity within the carbon flux response of tropical forests to eCO<sub>2</sub>, depending on the balance between benefits and potential temperature-related challenges, emphasizes the necessity for optimizing and refining the process-based representation of climate–growth relationships. For instance, employing various sensitivity scenarios involving temperature intervals and CO<sub>2</sub> concentrations can offer valuable insights for identifying the critical threshold beyond which the benefits of eCO<sub>2</sub> do not outweigh the adverse high-temperature effects.

While self-thinning-induced carbon loss is heightened in the eCO<sub>2</sub> scenario because of increased competition, drought-induced carbon loss might not always be exacerbated under eCO<sub>2</sub>, although there is more biomass built up over time that is available for carbon loss, revealing the influence of water-saving conditions on carbon loss. An important indicator we employed in our model, cumulative drought exposure, reveals a reduction over most areas when accounting for the water-saving effects driven by eCO<sub>2</sub> (Fig. S3). Regarding tree mortality triggered by drought events, our modeling work confirms alleviated drought exposure, which is consistent with previous modeling findings that eCO<sub>2</sub> brings about water saving due to increased WUE, thus enhancing vegetation productivity and decreasing the probability of forest dieback in the eastern Amazon basin being threatened by drier and warmer climate scenarios (Huntingford et al., 2013; Lapola et al., 2009; Zhang et al., 2015); however, these studies mainly considered enhanced productivity with respect to sustaining the biomass rather than carbon loss changes. The strength of this mitigation effect depends on the intensity and duration of stress (Lapola et al., 2009) and could postpone the point at which a forest shifts from being a carbon sink to a carbon source (Feng et al., 2018). However, the decreased frequency of tree mortality risk, combined with an increase in background biomass stock facilitated by eCO<sub>2</sub>, contributes to uncertainty regarding the fate of carbon loss.

In our modeling study, we found an increase in both WUE and tree growth (Figs. 1 and S7). While van der Sleen et al. (2015) reported no growth stimulation of tropical trees by means of CO<sub>2</sub> fertilization, they did find an increase in WUE in their tree-ring-width analysis, which focused on a fixed tree size class. However, it has been argued by Brienen et al. (2017b) that this approach, aimed at removing the effect of tree size, might lead to a biased interpretation of growth trends, particularly when there is a clustered age distribution with fast-growing and slow-growing trees coexisting. Concerning the less controversial increase in WUE, Brienen et al. (2017a) suggested that the observed trends could be related to developmental effects rather than solely being the

result of climate and CO<sub>2</sub> effects on WUE. It is crucial to investigate and distinguish these factors as WUE varies with tree development stages, especially in broadleaf forests. Partial stomatal closure, driven by increasing constraints in water transport with tree height and increasing photosynthesis due to greater light availability with tree height, can lead to changes in intrinsic WUE (McDowell et al., 2011). To better isolate the effects of external factors, such as eCO<sub>2</sub>, through size-stratified sampling and account for the varying tree gas regulation strategies throughout a tree's lifespan, it is essential to incorporate a stratified simulation of stomatal conductance as well as corresponding photosynthesis. While our ORCHIDEE model features a stratified LAI pattern, it does not yet include stratified simulation of a stomatal conductance. Implementing a more detailed water budget per canopy layer would provide a more comprehensive understanding of tree-height-related shifts in WUE.

Regarding drought resistance and resilience, C-band radar data have demonstrated a decrease in resistance to drought over the Amazon rainforest during the past 3 decades, while forest resilience has not declined significantly (Tao et al., 2022). Our model simulation also detected the well-expected phenomenon of decreased resistance with increasing drought severity. The resistance and resilience of net biomass carbon balance were found to benefit from eCO<sub>2</sub>, which is broadly consistent with the enhanced resistance to drought due to restricted stomatal conductance and improved WUE observed in Feng et al. (2018). This suggests that eCO<sub>2</sub> can enhance the recovery of ecosystem carbon uptake after short-term drought events.

Several uncertainties warrant in-depth investigations. Whether eCO<sub>2</sub> would lead to biomass growth also depends on the carbon allocation strategies and, subsequently, the carbon turnover rate (Friend et al., 2014; Hofhansl et al., 2016), which has been found to be governed by hydraulic constraints, such as the hydraulic adjustment of the ratio between leaf area and sapwood area (Trugman et al., 2019). Given that both tree productivity and mortality responses during drought are sensitive to hydraulic traits (Anderegg et al., 2016, 2018), incorporating varying hydraulic traits that are adaptive to the environment will be highly important (Madani et al., 2018). For example, tree mortality risk is intricately linked to plant water use strategies, with isohydric tree species exhibiting a lower xylem embolism risk due to their tendency to close stomata earlier to conserve water. In contrast, anisohydric tree species, characterized by less conservative water use strategies, may derive more significant benefits from eCO<sub>2</sub>-induced partial stomatal closure. Additionally, the analysis of water deficit affiliation has indicated that genera affiliated with wetter climate regimes exhibit a higher risk of drought-induced tree mortality (Esquivel-Muelbert et al., 2017). Exploring the interactions between eCO<sub>2</sub> and varying hydraulic vulnerabilities would be a potential avenue for further examining the effects of eCO<sub>2</sub> on biomass carbon dynamics. Besides hydraulic-failure-induced tree mortality,



other possible sources, including carbon starvation, should be included as more carbon gain enhanced by eCO<sub>2</sub> would delay the depletion of carbohydrate reserves. Furthermore, it is essential to consider the legacy effects of drought, a dimension that has not been addressed in process-based modeling. Yang et al. (2023) used a first-order kinetics model to account for the gradual decomposition of coarse woody debris, yielding a better correspondence between net biomass carbon change and variability in the atmospheric CO<sub>2</sub> growth rate. The legacy effects from tree mortality should be carefully revisited, given the evidence suggesting that external drivers can lead to increased mortality for at least 2 years after a climatic event (Aleixo et al., 2019). Regarding the strength and persistence of eCO<sub>2</sub>, previous studies have suggested that such fertilization effects could slow down (Peñuelas et al., 2017), and the eCO<sub>2</sub> effect has declined in recent years, possibly due to nutrient limitation (Winkler et al., 2021). Wieder et al. (2015) demonstrated that accounting for nitrogen and nitrogen–phosphorus limitation lowers the projected productivity and could even turn terrestrial ecosystems into carbon sources. Fleischer et al. (2019) highlighted the important role of phosphorus acquisition and use strategies in regulating the forest response to eCO<sub>2</sub>, effectively reducing the expected stimulation by 50 % over the Amazon rainforest. The lack of downregulation on fertilization in the model could lead to an overestimation of eCO<sub>2</sub> effects. In addition to the absence of downregulation due to nutrient availability, uncertainties in carbon allocation could also contribute to differences in baseline values compared to inventory data. In the ORCHIDEE model, carbon allocation among biomass components adheres to the “pipe model” theory, which dictates the relationship between leaf area, sapwood area, and fine-root area (Sitch et al., 2003). However, the carbon allocation process remains relatively unconstrained and requires further observation data for benchmarking purposes. Given that nutrient availability influences productivity and adjustments in carbon allocation, a nutrient-enabled version of the model would help elucidate ecosystem responses to eCO<sub>2</sub>. Therefore, estimating the strength and persistence of the CO<sub>2</sub> fertilization effect under future climate scenarios remains challenging (Nölte et al., 2023). Additional observations are imperative, and the Amazon free-air CO<sub>2</sub> enrichment (AmazonFACE) project will be a strong observational constraint on our knowledge of the rainforest response to eCO<sub>2</sub> (Lapola and Norby, 2014). We have also provided estimates of carbon gain and carbon loss in response to the planned CO<sub>2</sub> increase (i.e., 200 ppm above ambient levels) for this forest for the period from 2010 to 2020. Our simulations indicate an enhancement in gross primary productivity (GPP) of ~ 34 % and an enhancement in wNPP of ~ 55 % (diameter at breast height (DBH) > 10 cm) throughout the simulation period. These values are higher compared to those from simulations conducted with nutrient-cycle-enabled models, as reported by Fleischer et al. (2019). Obtaining more experimental data to illustrate the interac-

tions between water and nutrient availability and their impacts on the CO<sub>2</sub> fertilization effect would aid in constraining model responses, thus enabling more accurate predictions of the Amazon rainforest’s response to future climate change.

In our study, we conducted offline simulations and found that eCO<sub>2</sub> leads to an increase in WUE (Fig. S7), which could partially mitigate drought risk through soil–atmosphere feedback mechanisms. However, it is noteworthy that CO<sub>2</sub>-induced physiological effects reduce ET (evapotranspiration) and, subsequently, precipitation in a land–atmosphere coupled mode. A recent study using coupled climate model simulations has highlighted that the reduction in ET under eCO<sub>2</sub> and its impact on precipitation contribute to potential water stress (Li et al., 2023; Skinner et al., 2018). Tree dieback indeed leads to a reduction in plant transpiration and also decreases the soil moisture consumption. We found that eCO<sub>2</sub> leads to an increase in land surface temperature of 0–0.26 °C, based on the simulation during the 2015–2016 El Niño (Fig. S6). Therefore, given the contribution of moisture recycling to precipitation over the Amazon rainforest, a comprehensive investigation into the effects of eCO<sub>2</sub> on biomass carbon dynamics, such as whether eCO<sub>2</sub> can mitigate the negative effects of water stress due to changes in precipitation, should be conducted in a coupled mode, including a tree mortality module to capture the intricate interactions among these components.

## 5 Conclusions

In summary, this work offers a comprehensive quantitative assessment at the basin scale of how eCO<sub>2</sub> influences above-ground biomass carbon gain and carbon loss in a warming and increasingly water-stressed climate. We systematically disentangle the effect of eCO<sub>2</sub> in this complex ecosystem. Our findings not only underscore the role of eCO<sub>2</sub> in shaping the high gain/high loss pattern but also highlight its water-saving benefits. Additionally, we identify an enhancement in drought resistance and resilience attributed to eCO<sub>2</sub> as it accelerates drought recovery. Our improved model, which separates tree mortality schemes into competition-driven and drought-driven mechanisms, offers a more comprehensive understanding of carbon fluxes in response to eCO<sub>2</sub>, a perspective that cannot be solely attained through field experiments. With the likelihood of more frequent and intense drought events in the near future, these findings serve as a compelling impetus for further modeling and observational efforts aimed at gaining deeper insights into the role of eCO<sub>2</sub> in predicting the forest biomass carbon budget and ecosystem vulnerability within the Amazon rainforest.

**Code availability.** The ORCHIDEE-CAN-NHA r7236 model code used in this study is deposited at [https://forge.ipsl.jussieu.fr/orchidee/browser/branches/publications/ORCHIDEE\\_CAN\\_NHA](https://forge.ipsl.jussieu.fr/orchidee/browser/branches/publications/ORCHIDEE_CAN_NHA) (last access: 17 June 2021) and



archived at <https://doi.org/10.14768/8C2D06FB-0020-4BC5-A831-C876F5FBBFE9> (Yao, 2021).

**Data availability.** CRUJRA v2.1 climate-forcing data can be found at <https://catalogue.ceda.ac.uk/uuid/10d2c73e5a7d46f4ada08b0a26302ef7> (Harris, 2020). The forest inventory dataset can be found at [https://doi.org/10.5521/ForestPlots.net/2014\\_4](https://doi.org/10.5521/ForestPlots.net/2014_4) (Brienen et al., 2014).

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**Author contributions.** PC and YY designed the study. YY ran the simulation, analyzed the outputs, and drafted the paper. All authors contributed to the final paper.

**Competing interests.** The contact author has declared that none of the authors has any competing interests.

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