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Wave propagation in random 2D turbulence: a multi-scale approach

- ³ Valentin Resseguier^{1,2}[†], Erwan Hascoët³ and Bertrand Chapron⁴
- 4 ¹INRAE, UR OPAALE, 17 avenue de Cucillé, F-35044 Rennes, France
- 5 ²LAB, SCALIAN DS, 2 Rue Antoine Becquerel, 35700 Rennes, France
- 6 ³Oceandatalab, 870 Rte de Deolen, 29280 Locmaria-Plouzané, France
- 7 ⁴LOPS, Ifremer, 1625 Rte de Sainte-Anne, 29280 Plouzané

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To study two-dimensional dispersive waves propagating through turbulent flows, a new and 9 less restrictive fast waves approximation is proposed using a multiscale setting. In this ansatz, 10 large scale and small scale of the turbulence are treated differently. Correlation lengths of 11 the random small scale turbulence components can be considered negligible in the wave 12 packet propagating frame. Still, the large-scale flow can be relatively strong, to significantly 13 impact wavenumbers along the propagating rays. New theoretical results, numerical tools and 14 proxies are derived to describe ray and wave action distributions. All model parameters can 15 robustly be calibrated from the large-scale flow component only. We illustrate our purpose 16 with ocean surface gravity waves propagating in different types of surface currents. The 17 multiscale solution is demonstrated to efficiently document wave trapping effects by intense 18 jets. 19

20 Key words:

21 **MSC Codes** 76B15,76M35, 82B31,82C31, 37H10, 37L55

22 1. Introduction

- 23 This paper aims to revisit the ray-path concept for fast waves propagating over heterogeneous
- 24 turbulent flows. Considering ocean surface wave propagation, many authors have already
- 25 discussed the random changes of rays subject to a random current (Voronovich 1991; White
- ²⁶ & Fornberg 1998; Smit & Janssen 2019), and consequences on wave action distributions.
- 27 Closures have been derived in the Eulerian setting (Bal & Chou 2002; Klyatskin & Koshel
- 28 2015; Borcea et al. 2019; Kafiabad et al. 2019; Bôas & Young 2020; Garnier et al.
- 29 2020). Some of these approaches can be traced back to wave-wave interactions models, e.g.
- 30 McComas & Bretherton (1977) (see also Kafiabad *et al.* 2019, and reference therein). In most

† Email address for correspondence: valentin.resseguier@inrae.fr

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31 cases, the central assumption is either time-delta-correlated turbulent velocity (Voronovich 1991; Klyatskin 2005; Klyatskin & Koshel 2015) and/or fast waves in comparison to fluid 32 flow velocities (White & Fornberg 1998; Dysthe 2001; Bal & Chou 2002; Borcea et al. 2019; 33 Kafiabad et al. 2019; Smit & Janssen 2019; Bôas & Young 2020; Garnier et al. 2020; Boury 34 35 et al. 2023; Wang et al. 2023). Medium variations may be slow and delta-correlations are 36 hardly justifiable in a fixed frame. Though, attached to a fast propagating wave group, the 37 medium may seem to vary rapidly, and the delta-correlation assumption makes more sense. Another common assumption is frozen turbulence. In such a case, weak currents also implies 38 conservation along ray of intrinsic frequency, wavenumber, and group velocity magnitude 39 in two dimension (Boury et al. 2023). Subsequently, most of wave dynamics models neglect 40 variations and diffusion of frequency or wavenumber. 41

The diffusion of the wave action at large distance with a multiscale decomposition of the current has already been reported (Bal & Chou 2002). However, an explicit formulation for the diffusivity has solely been derived for a zero large-scale current. More generally, fast wave models mostly rely either on zero or constant current components at larger scales. West (1978), for instance, discussed acoustic waves in a two-component random media, but no velocity was involved.

Hereafter, the proposed two-scale velocity decomposition falls into the family of stochastic 48 transport models (Kunita 1997; Mikulevicius & Rozovskii 2004; Resseguier et al. 2020a; 49 Zhen et al. 2023), including dynamics under Location Uncertainty (LU) (Mémin 2014; 50 Resseguier et al. 2017a) and Stochastic Advection by Lie Transport (SALT) (Holm 2015). 51 52 Under this framework, the small-scale velocity component is delta-correlated in time (Cotter et al. 2017). Up to usual source terms, fluid dynamics quantities (temperature, momentum, 53 etc) are transported by both the large-scale revolved component and that random unresolved 54 55 turbulence component. The stochastic closures obtained are conservative. Nonlinear wave Hamiltonian dynamics and wave influence on currents (e.g. stokes drift) have then been 56 derived by (e.g. Crisan & Holm 2018; Bauer et al. 2020; Holm 2021; Holm & Luesink 57 2021; Dinvay & Mémin 2022; Holm et al. 2023). Considering a single-wavevector current, 58 solutions for a monochromatic shallow water wave were developed by Mémin et al. (2022). 59 In the present study, our objective is restricted to the influence of turbulent flows on linear 60 61 waves.

After first recalling the principles of the ray tracing method, we present the multiscale framework for fast wave dynamics, its physical grounds and a calibration method for the closure. Simplified stochastic equations are then derived for the ray dynamics and the wave action spectrum, in both Lagrangian and Eulerian settings. For illustrative examples, numerical tools, analytic models and proxies are applied to ocean surface gravity waves propagating through two types of 2D turbulent flows: a typical slow homogeneous turbulence and a jet case.

69 2. Characteristics of wave-packet rays

70 Isolating a single progressive group of quasi-regular wave train, it follows a form

71 $h(\mathbf{x},t)e^{i\phi(\mathbf{x},t)} + c.c.$, for most properties. Typically, h would be the local wave height with

meters units. If a packet is to be followed, the phase, $\phi(\mathbf{x}, t)$, must smoothly vary along the

- 73 propagation, i.e. $\phi(\mathbf{x}, t)$ is differentiable. The relative frequency is then $\omega = -\partial_t \phi(\mathbf{x}, t)$, and
- the wave number vector $\mathbf{k} = \nabla \phi(\mathbf{x}, t)$, with wavenumber $k = \|\mathbf{k}\|$ and direction given by
- 75 the normalized wave-vector, $\tilde{k} = k/k = \begin{pmatrix} \cos \theta_k \\ \sin \theta_k \end{pmatrix}$. To first order, such a train of waves is

76 dispersive and the intrinsic frequency reads

77

80

$$\omega - \mathbf{k} \cdot \mathbf{v} = \omega_0 = \begin{cases} \operatorname{cst.} \frac{1}{\alpha} k^{\alpha}, & \alpha \neq 0\\ \operatorname{cst.} \log(k), & \alpha = 0 \end{cases}$$
(2.1)

and propagates with its group velocity $v_g = \nabla_k \omega$, constantly modified by the local velocity of the currents v,

$$\frac{\mathrm{d}\boldsymbol{x}_r}{\mathrm{d}t} = \boldsymbol{v}_g = \boldsymbol{v}_g^0 + \boldsymbol{v},\tag{2.2}$$

81 where \mathbf{x}_r is the centroid of a wave group, $\mathbf{v}_g^0 = \frac{\partial \omega_0(k)}{\partial k} \tilde{\mathbf{k}}$ is the group velocity without 82 currents, i.e. solely depending on the wave vector. For $\alpha = 1$, the medium is non-dispersive 83 (e.g. acoustic waves). $\alpha = 1/2$ corresponds to gravity waves over deep ocean ($\omega_0 = \sqrt{gk}$). 84 The dominant wave-vector \mathbf{k} within the group evolves according to

85
$$\frac{\mathrm{d}\boldsymbol{k}}{\mathrm{d}t} = -\boldsymbol{\nabla}\boldsymbol{v}^{\mathsf{T}}\boldsymbol{k}.$$
 (2.3)

Equations (2.2)-(2.3) are the Hamilton's eikonal equations. Along the propagating ray, velocity gradients induce linear variations. Decelerating currents will, for instance, shorten waves, and reduce the group velocity. Traveling over fields of random velocities v, the wavevector k will also become randomly distributed. Scattering of ocean surface wave packets by random currents can generally be assumed to be weak, with ||v|| of order 0.5 m.s⁻¹, much smaller than $v_g^0 = ||v_g^0||$ of order 10 m.s⁻¹. Yet, cumulative effects of these random surface currents can lead to strong convergence or divergence between initially nearby ray trajectories.

⁹⁴ To complete the wave field description, $E(\mathbf{x},t) = \frac{1}{2}\rho gh^2(\mathbf{x},t)$ and $A(\mathbf{x},t) =$ ⁹⁵ $E(\mathbf{x},t)/\omega_0(k(\mathbf{x},t))$ denote energy and action by unit of surface. *E* is expressed in ⁹⁶ J/m^2 and *A* in $J.s^{-1}/m^2$. To avoid spurious notations, we set the multiplicative constant ⁹⁷ $\frac{1}{2}\rho g$ to unity. The wave action is considered to be an adiabatic invariant in absence of source ⁹⁸ terms. Wave action is then crucial to anticipate wave transformations by currents (White ⁹⁹ 1999). Unlike wave energy, wave action is conserved, in the absence of wave generation or ¹⁰⁰ dissipation. This action is the integral over wave-vectors of the action spectrum, *N*, also ¹⁰¹ related to the wave energy spectrum, *E*:

102
$$A(\mathbf{x},t) = \int d\mathbf{k} N(\mathbf{x},\mathbf{k},t) = \int d\mathbf{k} \frac{E(\mathbf{x},\mathbf{k},t)}{\omega_0(\mathbf{k},t)}.$$
 (2.4)

Action and energy spectrum quantify action and energy by unit of surface (unit of x) and by unit of wave-vector surface (unit of k). Consider the (x, k) variable change between different times t_i and t_f integrating the characteristic eikonal equations (2.2)-(2.3)

106
$$\begin{pmatrix} \boldsymbol{x}_r(t_i) \\ \boldsymbol{k}(t_i) \end{pmatrix} \mapsto \begin{pmatrix} \boldsymbol{x}_r(t_f) \\ \boldsymbol{k}(t_f) \end{pmatrix}.$$
(2.5)

According to the Liouville theorem for Hamiltonian mechanics (Landau & Lifshits 1960, §46), the state-space of the "packet-by-packet" approach (the (x, k) space) does not contract nor dilates along time. Readers not familiar with Hamiltonian dynamics may see the divergence free of the 4-dimensional flow (2.5) – i.e. $\nabla_x \cdot \frac{dx_r}{dt} + \nabla_k \cdot \frac{dk}{dt} = 0$ – as the divergence free of incompressible flow velocities, leading naturally to volume-preserving dynamics. Therefore, if wave dissipation is neglected, the wave action spectrum *N* is conserved (Lavrenov (2013)), i.e.

114
$$N(\mathbf{x}_{r}(t_{i}), \mathbf{k}(t_{i}), t_{i}) = N(\mathbf{x}_{r}(t_{f}), \mathbf{k}(t_{f}), t_{f}).$$
(2.6)

- 115 This result is extremely useful because it only involves quantities of the characteristics, i.e.
- 116 each Fourier mode can be modified independently of the others. The wave energy spectrum
- 117 can be computed from the characteristics

118
$$E\left(\mathbf{x}_{r}(t_{f}), \mathbf{k}(t_{f}), t_{f}\right) = \frac{\omega_{0}(\mathbf{k}(t_{f}))}{\omega_{0}(\mathbf{k}(t_{i}))} E\left(\mathbf{x}_{r}(t_{i}), \mathbf{k}(t_{i}), t_{i}\right).$$
(2.7)

starting with an initial incoming wave spectrum $E(\mathbf{x}_r(t_i), \mathbf{k}(t_i), t_i)$ for every wave-vectors $k(t_i)$, starting from a small set of spatial points $\mathbf{x}_r(t_i)$.

121 **3.** A new fast wave assumption

Eikonal equations (2.2)-(2.3) are driven by currents and their gradients. Commonly, the Eulerian current v is decomposed into a low-frequency large-scale component \overline{v} and a transient small-scale unresolved component v':

$$v = \overline{v} + v'. \tag{3.1}$$

126 Current gradients naturally follow the same scale separation. From now on, we shall consider 127 divergence-free two-dimensional currents only.

- 128 3.1. The ray Lagrangian correlation time
- To better characterize the wave dynamics in such a random environment, the covariance of the fluid velocity can be evaluated in the wave group frame. To take into account the small-scale unresolved component v', its Eulerian spatio-temporal covariance is considered, assuming statistical homogeneity and stationarity for the Eulerian velocity $v'_E(t, x) = v'(t, x)$

133
$$C_{ij}^{v'_E}(\delta t, \delta \mathbf{x}) = \mathbb{E}\left(v'_i(t, \mathbf{x})v'_j(t + \delta t, \mathbf{x} + \delta \mathbf{x})\right) = \mathbb{E}\left(v'_i(t, \mathbf{x}_r(t))v'_j(t + \delta t, \mathbf{x}_r(t) + \delta \mathbf{x})\right) (3.2)$$

134 where x_r is solution of (2.2) with an arbitrary initial position x_r^0 . Then, we define, $v'_R(t) =$

135 $v'(t, \mathbf{x}_r(t))$, the Lagrangian velocity along the ray $\mathbf{x}_r(t)$. The temporal covariance of the 136 small-scale component v' – in the wave group frame – is the covariance of that Lagrangian velocity:

137
$$C_{ij}^{\nu'_R}(\delta t) = \mathbb{E}\left(\nu'_i(t, \mathbf{x}_r(t))\nu'_j(t+\delta t, \mathbf{x}_r(t+\delta t))\right) = C_{ij}^{\nu'_E}(\delta t, \mathbf{x}_r(t+\delta t) - \mathbf{x}_r(t)),$$
 (3.3)

138 Assume for example a typical isotropic form for the Eulerian covariance:

139
$$C^{\nu'_{E}}(\delta t, \delta \mathbf{x}) = C\left(\frac{|\delta t|}{\tau_{\nu'}} + \frac{||\delta \mathbf{x}||}{l_{\nu'}}\right), \tag{3.4}$$

the covariance can be evaluated in the wave group frame for small time increment δt :

141
$$C^{\nu'_{R}}(\delta t) = C\left(\frac{|\delta t|}{\tau_{\nu'}} + \frac{\|\boldsymbol{x}_{r}(t'+t) - \boldsymbol{x}_{r}(t')\|}{l_{\nu'}}\right) = C\left(\left(\frac{1}{\tau_{\nu'}} + \frac{\|\boldsymbol{v}_{g}\|}{l_{\nu'}}\right)|\delta t| + O(\delta t^{2})\right), \quad (3.5)$$

since $\mathbf{x}_r(t'+t) - \mathbf{x}_r(t') = \mathbf{v}_g \delta t + O(\delta t^2)$. Therefore, $\left(\frac{1}{\tau_{v'}} + \frac{\|\mathbf{v}_g\|}{|\mathbf{v}_{v'}|}\right)^{-1}$ is the correlation time of $\mathbf{v}'(t, \mathbf{x}_r(t))$. For fast waves, the along-ray correlation time of the small-scale velocity can be approximated by $l_{v'}/v_g^0$. Note that eikonal equations (2.2)-(2.3) involve both velocity and velocity gradients. The above derivation is also valid for the small-scale velocity gradients ($\nabla \mathbf{v}^{\mathsf{T}}$)' $(t, \mathbf{x}_r(t))$. The ratio ϵ , between that along-ray correlation time and the characteristic time of the wave group properties evolution, will then control the time decorrelation

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148 assumption of v':

$$\epsilon = \frac{l_{\nu'}}{\nu_g^0} \|\boldsymbol{\nabla} \boldsymbol{\nu}^{\mathsf{T}}\| \sim \frac{l_{\nu'}}{l_{\nu}} \frac{\|\boldsymbol{\nu}\|}{\nu_g^0}.$$
(3.6)

149

This time scale estimation can be obtained from spatio-temporal covariances more general than (3.4) (not shown) even though the derivation is more technical. Note the Eulerian smallscale velocity ν' is not necessarily time uncorrelated, as assumed in Voronovich (1991); Klyatskin & Koshel (2015). Yet, for small enough ϵ , the Lagrangian small-scale velocity along the ray can be considered time uncorrelated. From the expression of ϵ , such a condition depends upon:

156 • v_g^0 , the fast wave group velocity

157 • $\|v\|$, often slow but not always negligible compared to the intrinsic wave group, v_g^0 . 158 • $l_{v'}/l_v$, related to the separation between large scales \overline{v} and small scales v', e.g. the

• $l_{\nu'}/l_{\nu}$, related to the separation between large scales $\overline{\nu}$ and small scales ν' , e.g. the spatial filtering cutoff of the large-scale velocity $\overline{\nu}$, but also related to its kinetic energy (KE) distribution over spatial scales, typically the spectrum slope.

This along-ray partial time-decorrelation assumption is less restrictive than the usual fast 161 162 wave approximation (White & Fornberg 1998; Dysthe 2001; Bal & Chou 2002; Borcea et al. 2019; Kafiabad et al. 2019; Smit & Janssen 2019; Bôas & Young 2020; Garnier et al. 163 2020; Boury *et al.* 2023; Wang *et al.* 2023) – say $\frac{\|\mathbf{v}\|}{v_g^0} \ll 1$ – and than the SALT-LU time-decorrelation used for turbulence dynamics (Mémin 2014; Holm 2015; Cotter *et al.* 2017; 164 165 Resseguier *et al.* 2020*a*) – say $\frac{l_{v'}}{l_v} \ll 1$. Similarly, this last validity criterion can be obtained 166 replacing in (3.2)-(3.6) x_r by the fluid particle Lagrangian path x (solution of $\frac{dx}{dt} = v$) and thus v_g^0 by v. These asymptotic models often rely on averaging or homogenization techniques (Papanicolaou & Kohler 1974; White & Fornberg 1998) to derive Markovian 167 168 169 dynamics involving various types of diffusivity. 170

171

3.2. Ray absolute diffusivity and turbulence statistics: calibration

Diffusivity is a natural tool to specify statistics of uncorrelated random media. For waves in random media, we shall specify multi-point statistics, and the Fourier space is convenient for this purpose. We will first present scalar diffusivity and then distribute it over spatial scales to fully calibrate the random velocity ν' , i.e. choose some parameter values to set the statistics of that velocity field. As such, we will obtain a closed model to derive analytic results and generate samples for simulations.

The absolute diffusivity (or Kubo-type formula) usually corresponds, in the so-called diffusive regime, to the variance per unit of time of a fluid particle Lagrangian path $\frac{d\mathbf{x}(t)}{dt} =$ $\mathbf{v}_L(t) = \mathbf{v}(t, \mathbf{x}(t))$. It is approximately equal to the velocity variance times its correlation time. The Eulerian velocity covariance (3.4) will thus induce an absolute diffusivity (Piterbarg & Ostrovskii 1997; Klyatskin 2005)

183
$$\frac{1}{2}a^{L} = \int_{0}^{+\infty} d\delta t \ C^{\nu'_{L}}(\delta t) = \int_{0}^{+\infty} d\delta t \ C^{\nu'_{E}}(\delta t, \mathbf{x}(t+\delta t) - \mathbf{x}(t)) \approx \frac{1}{2}\tau_{\nu'} \ C(0).$$
(3.7)

This diffusivity well describes effects of fast-varying eddies but is not appropriate in our case. Indeed, along a propagating wave group, $\frac{d\mathbf{x}_r(t)}{dt} = \mathbf{v}_g^0(t) + \mathbf{v}_R(t)$, a ray absolute diffusivity occurs and slightly differs from the usual absolute diffusivity to become

187
$$\frac{1}{2}a^{R} = \int_{0}^{+\infty} \mathrm{d}\delta t \ C^{\nu_{R}'}(\delta t) \approx \frac{1}{2} \left(\frac{1}{\tau_{\nu'}} + \frac{\|\boldsymbol{v}_{g}\|}{l_{\nu'}}\right)^{-1} C(0) \approx \frac{1}{2} \frac{l_{\nu'}}{v_{g}^{0}} C(0). \tag{3.8}$$

188 The absolute diffusivity sets the amplitude of the small scale velocity v'. Indeed, since the

209

189 kinetic energy of a time-continuous white noise is infinite, it has no physical meaning. It is

- 190 more relevant to deal with absolute diffusivity rather than kinetic energy in order to describe
- 191 the statistics of the time-uncorrelated velocity. To calibrate its spatial correlations, we may

192 focus on its Fourier transform, $\widehat{v'}(\kappa, t)$, denoting by $\kappa = \kappa \begin{pmatrix} \cos \theta_{\kappa} \\ \sin \theta_{\kappa} \end{pmatrix}$, the surface current

wave-vector. By analogy with the current kinetic energy spectra $E_{\kappa} = \frac{1}{2} \oint_{0}^{2\pi} d\theta_{\kappa} \kappa \frac{\|\widehat{v}(\kappa,t)\|^{2}}{(2\pi)^{2}}$,

194 Resseguier *et al.* (2017*b*, 2020*b*) decompose the absolute diffusivity scale by scale:

195
$$a^R = \int_0^{+\infty} A^R_{\nu'}(\kappa) \mathrm{d}\kappa.$$
(3.9)

Referring it to Absolute Diffusivity Spectral Density (ADSD), it is defined as the kinetic energy spectra multiplied by the correlation time at each scale, $\tau(\kappa)$. Unlike Resseguier *et al.* (2017*b*, 2020*b*), that correlation time is here imposed by the wave dynamics. Therefore, by analogy with (3.8) we choose a correlation time $\tau^R(\kappa) = \frac{1/\kappa}{\nu_{\mu}^0(k)}$ and then

200
$$\frac{1}{2}A^{R}(\kappa) = \frac{1}{2}\tau^{R}(\kappa) E_{\kappa}(\kappa) = \frac{1}{2}\frac{1/\kappa}{v_{g}^{0}(k)} E_{\kappa}(\kappa), \qquad (3.10)$$

201 where k denotes the wave wavenumber and κ the current wavenumber.

To calibrate an equivalent noise, we model \mathbf{v}' by $\sigma dB_t/dt$, where dB_t/dt is a spatiotemporal white noise and σ denotes a spatial filtering operator which encodes spatial correlations through its ADSD, $A_{v'}^R$ and the horizontal incompressibility condition ($\nabla \cdot \sigma = 0$). For incompressibility, we work with the curl of a streamfunction. To generate a homogeneous and isotropic streamfunction, we can filter a one-dimensional white noise \dot{B} with a filter $\check{\psi}_{\sigma}$ (Resseguier *et al.* 2017*b*), that is $\check{\psi}_{\sigma} \star \dot{B}$ where \star denotes a spatial convolution. The velocity field is hence :

$$\mathbf{v}' = \sigma \mathrm{d}B_t / \mathrm{d}t = \mathbf{\nabla}^{\perp} \breve{\psi}_{\sigma} \star \mathrm{d}B_t / \mathrm{d}t,$$

with ∇^{\perp} the two-dimensional curl. That formula is easily written and implementable in Fourier space (see equation (A 2)). To define the streamfunction filter, we note that $\frac{\pi\kappa^3}{(2\pi)^2} |\tilde{\psi}_{\sigma}(\kappa)|^2 = \frac{1}{2} \oint_0^{2\pi} d\theta_{\kappa} \kappa \frac{\|\widehat{\sigma dB_t}(\kappa)\|^2}{(2\pi)^2 dt} = A_{\nu'}^R(\kappa)$, i.e. the filter can be fully defined by the small-scale ADSD, $A_{\nu'}^R$. To close our model, we assume an ADSD power law:

(3.11)

214
$$A^R(\kappa) \approx A_0^R \kappa^{-\mu}.$$
 (3.12)

It enables automatic closure calibration: $A_{\nu'}^R(\kappa) = A_0^R \kappa^{-\mu} - A_{\overline{\nu}}^R(\kappa)$, from instantaneous large-scale current statistics $A_{\overline{\nu}}^R$ only (Resseguier *et al.* 2020*b*) as illustrated in figure 1.

217 4. Statistical wave dynamics

In a stochastic framework, the Stratonovich or Itō notations can both be used (Kunita 1997; Oksendal 1998). Under Stratonovich calculus rules, expressions become similar to deterministic ones. Specifically, stochastic versions of linearized dynamical equations are obtained by replacing v by $\overline{v} + \sigma \circ dB_t/dt$. Then, the stochastic transport of phase, $\frac{d}{dt}\phi = \omega_0(||\nabla \phi||)$, i.e. – up to that velocity replacement – the Stratonovich dispersion relation is exactly (2.1). The method of characteristics also applies. Note, one can switch from Stratonovich to Itō notations, where v' corresponds to $\sigma dB_t/dt$. The characteristics



Figure 1: KE spectrum $(m^2.s^{-2}/(\text{rad}.m^{-1}))$ (left) and ADSD $(m^2.s^{-1}/(\text{rad}.m^{-1}))$ (right) of the resolved high-resolution velocity, A^R , in red, low-resolution velocity, $A^R_{\overline{y}}$, in blue, and modeled stochastic velocity, $A_{\nu'}^R(\kappa) = A_0^R \kappa^{-\mu} - A_{\overline{\nu}}^R(\kappa)$, in green. For the ADSD power law, $A^R(\kappa) \approx A_0^R \kappa^{-\mu}$, we impose the theoretical KE spectrum slope $-\frac{5}{3}$ (black solid line), coherently with homogeneous SQG dynamics (see "Numerical results") section). The residual ADSD (green line) is set to extrapolate that power law at small scales.

equations (2.2)-(2.3) also remain unchanged for homogeneous and isotropic ν' : 225

226
$$\begin{cases} \mathrm{d}\boldsymbol{x}_r = (\boldsymbol{v}_g^0 + \overline{\boldsymbol{v}})\mathrm{d}t + \sigma \mathrm{d}\boldsymbol{B}_t, \\ \mathrm{d}\boldsymbol{k} = -\boldsymbol{\nabla}(\overline{\boldsymbol{v}}\mathrm{d}t + \sigma \mathrm{d}\boldsymbol{B}_t)^{\mathsf{T}}\boldsymbol{k}. \end{cases}$$
(4.1)

4.1. Single-ray stochastic differential equations

When studying a single ray in an homogeneous and isotropic turbulence (3.11), the wave-228 vector dynamics simplifies. In the local crest-oriented frame, the influence of small scale 229 currents can be solely represented by four one-dimensional white noise forcings. 230

Notably, dynamics of wave-vectors (2.3) are similar to tracer gradient dynamics (Bühler 231 2009; Plougonven & Zhang 2014). Only the coupled ray path dynamics (2.2) differs. 232 233 Accordingly, we follow the notations and derivations of the mixing analysis from Lapeyre et al. (1999) and references therein. Without loss of generality, the large-scale velocity can 234 be parameterized as 235

227

236
$$\overline{\boldsymbol{v}} = \overline{\boldsymbol{v}} \begin{pmatrix} \cos \overline{\theta} \\ \sin \overline{\theta} \end{pmatrix}$$
 and $\nabla \overline{\boldsymbol{v}}^{\mathsf{T}} = \frac{1}{2} \begin{bmatrix} \overline{\sigma} \sin 2\overline{\phi} & \overline{\omega} + \overline{\sigma} \cos 2\overline{\phi} \\ -\overline{\omega} + \overline{\sigma} \cos 2\overline{\phi} & -\overline{\sigma} \sin 2\overline{\phi} \end{bmatrix}$. (4.2)

Figure 2 provides a synthetic view of angles involved. The dynamics wave group centroid x_r is 237 directly driven by the large current wave group velocity, $v_g^0 + \overline{v}$. The influence of the large-scale currents gradients on the wavevector dynamics (4.1), expressed in the local crest-oriented 238 239 frame $(\tilde{k}, \tilde{k}^{\perp})$, is straightforward (Lapeyre *et al.* 1999). The small-scale currents force the ray 240 dynamics through a stochastic noise. For a single ray $(\mathbf{x}_r, \mathbf{k}) = (x_r, y_r, k \cos \theta_k, k \sin \theta_k)$, 241 this noise can be rigorously described by four independent one-dimensional white noises 242 only (see Appendix A), $\dot{B}_t^{(1)}$, $\dot{B}_t^{(2)}$, $\dot{B}_t^{(3)}$, and $\dot{B}_t^{(4)}$, and: 243

244
$$\frac{\mathrm{d}}{\mathrm{d}t}x_r = v_g^0 \cos\theta_k + \overline{v}\cos\overline{\theta} + \sqrt{a_0}\dot{B}_t^{(1)}, \qquad (4.3)$$

245
$$\frac{\mathrm{d}}{\mathrm{d}t}y_r = v_g^0 \sin\theta_k + \overline{v}\sin\overline{\theta} + \sqrt{a_0}\dot{B}_t^{(2)}, \qquad (4.4)$$

246
$$\frac{\mathrm{d}}{\mathrm{d}t}\log k = -\overline{\sigma}\sin(\zeta) + \gamma_0 + \sqrt{\gamma_0}\dot{B}_t^{(3)}, \qquad (4.5)$$

247
$$\frac{\mathrm{d}}{\mathrm{d}t}\theta_k = \frac{1}{2}(\overline{\omega} - \overline{\sigma}\cos(\zeta)) + \sqrt{3\gamma_0}\dot{B}_t^{(4)}, \qquad (4.6)$$



Figure 2: Schematic view of vectors and angles involved in single-ray dynamics. \overline{S}_{-} and \overline{S}_{+} are respectively compression and dilatation axes associated with the large-scale velocity gradient $\nabla \overline{v}^{T}$.

248 where $\zeta = 2(\theta_k + \overline{\phi})$ and

249

250

$$a_0 = \frac{1}{2\mathrm{d}t} \mathbb{E} \|\sigma \mathrm{d}B_t\|^2 = \int_0^{+\infty} A_{\nu'}^R(\kappa) \mathrm{d}\kappa, \qquad (4.7)$$

$$\gamma_0 = \frac{1}{8\mathrm{d}t} \mathbb{E} \|\boldsymbol{\nabla}_{\boldsymbol{x}} (\sigma \mathrm{d}B_t)^\mathsf{T}\|^2 = \frac{1}{4} \int_0^{+\infty} k^2 A_{\nu'}^R(\kappa) \mathrm{d}\kappa.$$
(4.8)

Diffusivity constants depend through (3.10) on both the correlation length and the spectrum 251 slope of the small-scale velocity. In contrast to the classical fast wave approximation, the 252 wavenumber does vary. This is due to (i) the finite large-scale strain rate, $\overline{\sigma}$, and (ii) the 253 254 small-scale isotropic velocity model (3.11). This isotropy assumption and its implication are discussed in Appendix C. Note that neither large-scale nor small-scale component is assumed 255 to be steady, even though that Eulerian velocity unsteadiness is only a secondary process in 256 the wave dynamics. The fast temporal variations seen by the wave are mainly driven by the 257 large wave speed and not by the Eulerian velocity unsteadiness. The current unsteadiness can 258 also lead to wavenumber variations (Dong et al. 2020; Cox et al. 2023; Boury et al. 2023). 259 Given a known wavevector angle, it leads to a wavenumber evolution 260

261
$$k(t) = k(0) \exp\left(-\int_0^t \overline{\sigma} \sin(2(\theta_k + \overline{\phi}))dt'\right) \exp\left(\gamma_0 t + \sqrt{\gamma_0}B_t^{(3)}\right), \tag{4.9}$$

and hence to the complete wavevector distribution, i.e. the wave spectrum. The second exponential factor in (4.9) is a geometric Brownian motion. Its mean diverges in time exponentially rapidly. Physically, shear and strain of v' tends to shorten the wavelength ((Voronovich 1991; Boury *et al.* 2023)) leading to this exponential divergence. This factor has a log-normal distribution, suggesting possible extreme transient wavenumber events. This generalizes previous results Voronovich (1991); Klyatskin & Koshel (2015), obtained with neglecting the time-correlated current component, \overline{v} .

For completeness, the action distribution over space and wave vector can be derived. Some approaches consider finite-size wave trains either through additional equations (Jonsson 1990; White & Fornberg 1998) or re-meshing (Hell *et al.* in preparation). Otherwise, each ray transports its action spectrum (2.6) and we need to numerically combine many rays (Lavrenov 2013), or rely on analytic approximations. Typically, we solve (4.3)-(4.5) exhibiting, $p(\mathbf{x}, \mathbf{k} | \mathbf{x}_r^0, \mathbf{k}_r^0, t)$, the distribution of the ray (\mathbf{x}, \mathbf{k}) at time *t* given initial conditions

 $(\mathbf{x}_r^0, \mathbf{k}^0)$. Then, by analogy with tracers in incompressible turbulence (Piterbarg & Ostrovskii 275 276 1997, equation (1.31), see also Appendix D) we can evaluate the wave action spectrum mean - or any point-wise statistics - as follows 277

278
$$\mathbb{E}N(\boldsymbol{x}, \boldsymbol{k}, t) = \iint \mathrm{d}\boldsymbol{x}_r^0 \mathrm{d}\boldsymbol{k}^0 \ N^0(\boldsymbol{x}_r^0, \boldsymbol{k}_r^0) p(\boldsymbol{x}, \boldsymbol{k} | \boldsymbol{x}_r^0, \boldsymbol{k}^0, t), \qquad (4.10)$$

where N^0 is the initial wave action spectrum. Integrating this expression over wavevectors, 279 we note that the distribution inside the integrals changes 280

281
$$\mathbb{E}A(\boldsymbol{x},t) = \iint \mathrm{d}\boldsymbol{x}_r^0 \mathrm{d}\boldsymbol{k}^0 \ N^0(\boldsymbol{x}_r^0,\boldsymbol{k}_r^0) p(\boldsymbol{x}|\boldsymbol{x}_r^0,\boldsymbol{k}^0,t).$$
(4.11)

The wave action mean solely depends of group positions distribution. Multi-point action 282 statistics – e.g. focusing $\mathbb{E} \| \nabla_x A \|^2$ – rely on multi-ray correlations, encoded in the stochastic 283 characteristic equations (4.1), but not the simplified model (4.3)-(4.6). Alternatively, Eulerian 284 descriptions of wave action dynamics directly provide action distribution over space and wave 285 vector. 286

4.2. Eulerian dynamics and action diffusion 287

Wave action spectrum is transported along a 4-dimensional volume-preserving stochastic 288 flow (4.1). Again by analogy with incompressible turbulence (Resseguier et al. 2017a), the 289 stochastic transport of wave action spectrum in Itō notations reads 290

291
$$\partial_t N + (\mathbf{v}_g^0 + \overline{\mathbf{v}} + \sigma \frac{\mathrm{d}B_t}{\mathrm{d}t}) \cdot \nabla_{\mathbf{x}} N + \left(-\nabla_{\mathbf{x}} (\overline{\mathbf{v}} + \sigma \frac{\mathrm{d}B_t}{\mathrm{d}t})^{\mathsf{T}} \mathbf{k}\right) \cdot \nabla_{\mathbf{k}} N$$

292
$$= \begin{bmatrix} \nabla_{\mathbf{x}} \\ \nabla_{\mathbf{k}} \end{bmatrix} \cdot \left(\mathbf{D} \begin{bmatrix} \nabla_{\mathbf{x}} \\ \nabla_{\mathbf{k}} \end{bmatrix} N \right) = \frac{1}{2} a_0 \Delta_{\mathbf{x}} N + \frac{1}{2} \gamma_0 \frac{1}{k} \partial_k \left(k^3 \partial_k N \right) + \frac{3}{2} \gamma_0 \partial_{\theta_k}^2 N.$$
(4.12)

The RHS is reminiscent to Eq. (3.16) in Bôas & Young (2020) and Eq. (36) in Smit & 293 Janssen (2019), and more generally to rapid wave models. Nevertheless, equation (4.12) is 294 not averaged and explicitly involves large-scale currents and noise terms (terms with factor 295 $\frac{dB_t}{dt}$). Differences with Smit & Janssen (2019); Bôas & Young (2020) for the diffusivity 296 estimates and the detailed computation of the 4×4 diffusion matrix **D** can be found in 297 Appendix A. Itō notations of (4.12) explicitly separate mean terms (e.g. diffusion terms) and 298 zero-mean noise terms. Here, the Eulerian Itō notations reveal that coefficients $\frac{1}{2}a_0$, $\frac{1}{2}\gamma_0$, and 299 $\frac{3}{2}\gamma_0$ act to diffuse wave action in space, wavenumber and wave-vector angle, respectively. 300

5. Numerical experiments 301

302 To illustrate these developments, we consider ocean surface gravity waves propagating over a dynamical flow region. Ray tracing through synthetic surface currents will provide a 303 benchmark. It will be shown that a broad range of the current scales can be replaced by 304 the stochastic parametrization (3.11) without affecting ray scattering and action distribution. 305 Theoretical results (4.3)-(4.12) will suggest approximate analytic solutions. 306

Simplified upper ocean dynamics are considered to follow: 308

309
$$(\partial_t + \boldsymbol{v} \cdot \boldsymbol{\nabla}) \Theta = 0 \text{ with } \boldsymbol{v} = -\boldsymbol{\nabla}^{\perp} (-\Delta)^{-\xi} \Theta,$$
 (5.1)

310

where Θ stands for the buoyancy, ∇^{\perp} the curl and Δ the Laplacian. Two extreme cases: the Surface Quasi-Geostrophic dynamics ($\xi = \frac{1}{2}$) (Held *et al.* 1995; Lapeyre 2017), abbreviated 311



Figure 3: Current velocity norm of the SQG homogeneous turbulence (left) and of the 2D Euler jet current at high-resolution (512×512) .

312 SQG, and the two-dimensional Euler dynamics ($\xi = 1$), abbreviated 2D Euler. SQG is a dynamics with an extreme locality (KE spectrum slope -5/3) whereas 2D Euler has an 313 extreme non-locality (KE spectrum slope -3). The objective is to test how the proposed 314 closures apply to both dynamics to be equally useful for any more realistic upper ocean 315 316 dynamics. Additionally, test cases are developed to assess the multiscale stochastic closure in both homogeneous and heterogeneous propagating medium. Moreover, we would like to 317 challenge our closure beyond the validity of rapid wave models. In our first test case, surface 318 fast waves travel in a homogeneous and isotropic SQG turbulence. Then, we simulate waves 319 propagating in a spatially heterogeneous 2D Euler turbulence, mimicking an oceanic jet. For 320 both SQG and 2D Euler dynamics, a reference simulation is obtained at a resolution 512×512 321 322 for a 1000-km squared domain, with the help of a pseudo-spectral code (Resseguier et al. 2017b, 2020b). Once initialized, the current velocity v is about 0.1 m.s⁻¹ for the homogeneous 323 turbulence and 1 m.s^{-1} for the jet (see figure 3). 324

325

5.2. Rays scattering in homogeneous SQG turbulence

A wave system enters the bottom boundary, propagating to the top. The carrier incident wave 326 has an intrinsic wave group velocity of 10 m/s, i.e. wavelength $\lambda = 250$ m. Its envelope 327 is Gaussian with an isotropic spatial extension of 30λ . The left panels of figures 4 and 328 5 illustrates the resulting dynamics, spreading the wavevectors (figure 5) of the incoming 329 waves. From bottom to top, spectral diffusion occurs (figure 5), in the direction orthogonal 330 (here k_x) to the propagation (here k_y), in line with the additive noise appearing in equation 331 (4.6). This scattering accelerates – along the propagation – the wave position spread (figure 332 4). This acceleration is explained by the ray equation (4.3) dominated by the intrinsic wave 333 334 group velocity.

To mimic a badly resolved \overline{v} field, v is smoothed at a resolution 32×32 . Using this 335 coarse-scale current, middle panels of figures 4 and 5, the scattering - described in the 336 previous paragraph – is strongly depleting in comparison to ray tracing in fully-resolved 337 turbulence. The spectral diffusion induced by small-scale turbulence is missing. Thus, the 338 spatial spreading also is narrower compared to high-resolution simulations. A stochastic 339 current v' is then added for ray tracing (4.1). This stochastic component is divergence-free 340 341 and has a self-similar distribution of energy across spatial scales (3.11) (see figure 1). The resulting spatial and spectral spreads are now comparable to simulations with high-resolution 342

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Figure 4: Swell (wavelength $\lambda = 250$ m) interacting with a high-resolution (512×512) deterministic SQG current (left), a low-resolution (32×32) deterministic SQG current (middle) and a low-resolution (32×32) deterministic SQG current plus (one realization of) the time-uncorrelated stochastic model (right) – colored by the corresponding wave amplitude, $h(t) = \sqrt{\omega_0(k(t))N(t=0)}$ (right-hand side colorbar) – computed by forward advection and superimposed on the current vorticity $\omega = \nabla^{\perp} \cdot \nu$. The red cross indicate where the bidirectional wave spectra of figure 5 are computed.

currents. For this setting, the stochastic closure provides satisfying results for a sufficiently well-resolved large-scale current. The key decorrelation ratio $\epsilon = \frac{l_{v'}}{l_v} \frac{\|v\|}{v_g^0}$ indeed depends on the resolution through $l_{v'}$. The large-scale current \overline{v} is resolved on a 32×32 grid, i.e. with a resolution $l_{v'} = \frac{\|\nabla v^{\top}\|}{\|\nabla v^{\top}\|} l_v = 0.33 l_v$. As such $\epsilon = 4.1 \times 10^{-3}$, computed with $v_g^0 \approx 10 \text{ m.s}^{-1}$ and $\|v\| \approx 0.12 \text{ m.s}^{-1}$, so $\frac{\|v\|}{v_g^0} \approx 1.2 \times 10^{-2}$, which is sufficiently small to make the proposed model applicable.

From the ADSD estimate (equation (3.10) illustrated by figure 1) and (4.7)-(4.8), eval-349 uations of the diffusivity coefficients a_0 and γ_0 are straightforward. Previously discussed 350 Smit & Janssen (2019), the spatial diffusivity is extremely weak: $a_0 = 6.4 \times 10^{-1} \text{ m}^2 \text{ s}^{-1}$ 351 (spatial variations in ray equations (4.3)-(4.4) of about $\sqrt{a_0t} = 230$ m during 1 day). In contrast, the spectral angle diffusivity is large: $3\gamma_0 = 3.0 \times 10^{-8}$ rad².s⁻¹. Along our 1-day 352 353 simulation, neglecting large-scale velocity influence, (4.6) leads to a Brownian wave vector 354 angle variations $\delta\theta_k = \theta_k - \theta_k(0) = \sqrt{3\gamma_0}B_t^{(4)}$ with a standard deviation $\sigma_{\delta\theta_k} = \sqrt{3\gamma_0 t} = 5.2 \times 10^{-2}$ rad $\approx 3.0^\circ$, eventually increasing the wave group spectral maximal extension from 355 356 $\pm 2\sigma_{k_x} = \pm 2\frac{2\pi}{30\lambda} = \pm 1.7 \times 10^{-3} \text{ rad.m}^{-1} \text{ to } \pm 2\sigma_{k_x} \approx \pm 2\sqrt{\left(\frac{2\pi}{30\lambda}\right)^2 + \left(k\sigma_{\delta\theta_k}\right)^2} = \pm 3.1 \times 10^{-3}$ 357 rad.m⁻¹, confirmed by figure 5. This figure also illustrates the wave action diffusion induced 358 by diffusivity γ_0 , well predicted by the Eulerian wave action model (4.12). In this scattering 359

by diffusivity γ_0 , well predicted by the Eulerian wave action model (4.12). In this scattering regime, the increased angle variability leads, by advection, to a spatial spread. The simplified ray equation (4.3) gives $\delta x \approx \int_0^t v_g^0 \cos \theta_k dt' \approx v_g^0 \int_0^t \delta \theta_k dt' \approx v_g^0 \sqrt{3\gamma_0} \int_0^t B_{t'}^{(4)} dt'$ with maximal extension $\pm 2\sigma_x \approx \pm 2v_g^0 \sqrt{\gamma_0 t^3} \approx \pm 52$ km, in agreement with figure 4. Finally, we estimate a first-order delay along the propagation $\delta t = t - (y - y(0))/v_g^0 \approx \int_0^t (1 - \sin \theta_k) dt' \approx$ $\frac{1}{2} \int_0^t \delta \theta_k^2 dt' \approx \frac{3}{2}\gamma_0 \int_0^t (B_{t'}^{(4)})^2 dt'$, with mean value $\mathbb{E}\delta t = \frac{3}{4}\gamma_0 t^2$.



Figure 5: Bidirectional wave spectra, computed by backward advection, at 8 locations along a vertical axis (the mean wave propagation direction) resulting from a swell interacting with a high-resolution (512 × 512) deterministic SQG current (panel (a)), a low-resolution (32 × 32) deterministic SQG current (panel (b)) and a low-resolution (32 × 32) deterministic SQG current plus (one realization of) the stochastic model (3.11) (panel (c)). The spatial locations where the spectra are calculated are highlighted on figure 4 by the red crosses.

5.3. Wave groups trapped in a 2D Euler turbulent jet

Tests are now performed for rays traveling in fast and strongly heterogeneous 2D Euler flows. Classical fast wave models – assuming flows of weak amplitude and often uniform statistics – are expected to fail here. Jets exhibit strong current gradients (e.g. Kudryavtsev *et al.* 2017) creating strong ray focusing and possibly rogue events. Passing through localized spatial structures, caustics can appear but solely from unrealistically collimated wave trains (White & Fornberg 1998; Heller *et al.* 2008; Wang *et al.* 2023). Occurrences strongly reduce for

- 372 finite directional spread (Slunyaev & Shrira 2023). Here, wave groups are trapped in a jet, but nonlinear wave interactions are neglected. The high-resolution numerical simulations (see 373 figure 6) reveal that even linear wave trains are well trapped in adversarial currents. Freund 374 & Fleischman (2002) observed a similar behavior for acoustic waves in a 3D turbulent jet. 375 Note that during our simulation, rays cross the domain several times (because of the doubly 376 periodic boundary conditions, see Appendix E for technical details). Top (resp. bottom) of 377 378 the jet, the vorticity and thus - at first order - rays curvatures (Dysthe 2001) are negative (resp. positve). Therefore, rays oscillate around the jet. A toy model can explain this behavior. 379 Following the multiscale stochastic approach (4.3)-(4.6), wave scattering is also taken into 380 account. 381
- For very-coarsed-grained (4×4) current $\overline{\nu}$, oscillation remains, but most of the scattering 382 383 vanishes, as illustrated by figure 7. Moreover, the curvature of the jet creates artificial wave focusing at t = 8 and 10 days. Introducing a time-uncorrelated model (3.11) corrects the 384 resolution issue on figure 8. Figure 9 plots the current ADSD. The current is strong ($||v|| \approx 1.4$ 385 m.s⁻¹), and the usual fast wave approximation cannot be applied $\left(\frac{\|v\|}{v_g^0} \approx 1.2 \times 10^{-1}\right)$. However, 386 the proposed modified fast wave model is valid, even at the very coarse 4×4 resolution. [VR:(à discuter)] Indeed, 2D Euler spectra are steeper than for SQG dynamics, and the length scale 387 388 ratio is already significant at this resolution, $\frac{l_{v'}}{l_v} = 0.14$, and the derived time-decorrelation 389 ratio is small: $\epsilon = \frac{l_{\nu'}}{l_{\nu}} \frac{\|\boldsymbol{\nu}\|}{v_g^0} = 1.6 \times 10^{-2}.$ 390
- Furthermore, by approximating the under-resolved current \overline{v} , an analytic stochastic solution can be obtained for a ray traveling against the current. The large-scale pattern of the jet takes a quadratic form
- 394

$$\overline{u} \approx \overline{U}_0 - \frac{1}{2}\overline{\beta}\left(y - \frac{L_y}{2}\right)^2$$
 and $\overline{v} \approx 0$, with $\overline{U}_0, \overline{\beta} < 0.$ (5.2)

Note, the toy model (5.2) simply considers a straight jet, neglecting its curvature. For weak subgrid currents and a ray, $(x_r, y'_r + \frac{L_y}{2}, k, \theta_k)$, propagating mainly to the right, θ_k is small and the simplified ray equation (4.4) determines the group position with respect to the jet y'_r

- 398 $\frac{\mathrm{d}}{\mathrm{d}t}y'_r \approx v_g^0 \sin(\theta_k) = v_g^0 \theta_k + O(\theta_k^2). \tag{5.3}$
- For frozen turbulence, the wave-number and hence v_g^0 will not significantly vary. The other ray equation (4.3) localizes the group along the jet, $x_r \approx x_r(0) + (v_g^0 - \overline{u})t$, dropping the $O(\theta_k^2)$ from now on. Moreover, $\tilde{k}^{\perp} \cdot \nabla \overline{\nu}^{\mathsf{T}} \tilde{k} \approx -\partial_y \overline{u}$ and the dynamics of wave vector angle (4.6) simplifies to a stochastic oscillator equation:

403
$$\frac{d^2}{dt^2}y'_r = v_g^0 \frac{d}{dt} \theta_k = -\partial_y (v_g^0 \overline{u}) + v_g^0 \sqrt{3\gamma_0} \dot{B}_t^{(4)} = -\overline{\omega}_r^2 y'_r + v_g^0 \sqrt{3\gamma_0} \dot{B}_t^{(4)}.$$
(5.4)

with $\overline{\omega}_r = \sqrt{|v_g^0 \overline{\beta}|}$. Here $v_g^0 \overline{u}$ plays the role of a potential, trapping the rays in the jet vicinity, whereas the noise accounts for wave scattering. Solution of this linear equation is known (e.g. Resseguier *et al.* 2017*a*, Eq.(51)-(55)):

407
$$y_r(t) = \underbrace{\frac{L_y}{2} + y_r'(0)\cos(\overline{\omega}_r t) + \frac{v_g^0}{\overline{\omega}_r}\theta_k(0)\sin(\overline{\omega}_r t)}_{=\mathbb{E}(y_r(t))} + \underbrace{Y_{\gamma_0}\sqrt{\overline{\omega}_r}\int_0^t \sin(\overline{\omega}_r(t-r))dB_r^{(4)}}_{=y_r''(t)}, (5.5)$$

408 with $Y_{\gamma_0} = v_g^0 \sqrt{3\gamma_0/\overline{\omega}_r^3}$. The wavevector angle solution is similar. The solution ensemble



Figure 6: Rays facing a high-resolution (512 × 512) deterministic 2D Euler jet current – colored by the corresponding wave amplitude, $h(t) = \sqrt{\omega_0(k(t))N(t=0)}$ (right-hand side colorbar) – computed by forward advection and superimposed on the current vorticity $\omega = \nabla^{\perp} \cdot v$ (top colorbar).



Figure 7: Rays facing a low-resolution (4×4) deterministic 2D Euler jet current – colored by the corresponding wave amplitude, $h(t) = \sqrt{\omega_0(k(t))N(t=0)}$ (right-hand side colorbar) – computed by forward advection and superimposed on the current vorticity $\omega = \nabla^{\perp} \cdot v$ (top colorbar).



Figure 8: Rays facing a low-resolution (4 × 4) deterministic 2D Euler jet current plus (one realization of) the time-uncorrelated stochastic model – colored by the corresponding wave amplitude, $h(t) = \sqrt{\omega_0(k(t))N(t=0)}$ (right-hand side colorbar) – computed by forward advection and superimposed on the low-resolution current vorticity $\overline{\omega} = \nabla^{\perp} \cdot \overline{\nu}$ (top colorbar).



Figure 9: ADSD $(m^2 . s^{-1} / (rad. m^{-1}))$ of the resolved high-resolution jet velocity in red, low-resolution jet velocity in blue, and modeled stochastic velocity, in green. The theoretical spectrum slope -3 (black solid line) is imposed, consistent with homogeneous 2D Euler dynamics. The residual ADSD (green line) is set to extrapolate that power law at small scales.



Figure 10: Vorticity shear $\partial_{y}^{2}u$ of the deterministic 2D Euler jet current at high-resolution (512×512) (left), at low-resolution (4×4) (middle), and the corresponding swell system period $2\pi/\overline{\omega}_r$. Far from the jet (±200 km away), the vorticity shear becomes zero or even positive, so period larger than 10 days are cropped.

mean, $\mathbb{E}y_r$, is a simple coherent deterministic oscillator. This mean solution describes well 409 the interaction between the group and the under-resolved current from figure 7. From the 410 coarse-scale vorticity shear plotted in figure 10 in the vicinity of the jet, we can estimate 411 $\overline{\beta} = -2.7 \times 10^{-11} \text{m}^{-1} \text{.s}^{-1}$. It yields an oscillation frequency $\overline{\omega}_r = 1.3 \times 10^{-5} \text{rad.s}^{-1}$ i.e. a 412 period of $2\pi/\overline{\omega}_r = 5.7$ days, in agreement with the ray tracing simulations. Note that the 413 high-resolution vorticity shear, left panel of figure 10, does not suggest any relevant values 414 to explain the ray oscillations. Only the proposed multiscale current decomposition provides 415 a quantitative explanation for these oscillations, and by extension for trapping rays inside the 416 jet. Added to the mean solution, the random parts, $y''_r(t)$, are continuous summations of zero-417 mean incoherent wave fluctuations. At each time r, the additive random forcing introduces 418 an oscillation. But, the influence of the past excitations is weighed by sine wave due to the 419 phase change. The group position and wavevector angle are Gaussian random variables (as 420 linear combinations of independent Gaussian variables). Therefore, their finite dimensional 421 422 law (*i.e.* the multi-time probability density function) are entirely defined by their mean and covariance functions. Specifically, 423

$$424\mathbb{E}\left(y_r''(t)y_r''(t+\tau)\right) = \frac{1}{4}Y_{\gamma_0}^2\left(\cos(\overline{\omega}_r\tau)\left(2\overline{\omega}_rt - \sin(2\overline{\omega}_rt)\right) + \sin(\overline{\omega}_r\tau)\left(1 - \cos(2\overline{\omega}_rt)\right)\right).$$
(5.6)
425 In particular, the variance of the vertical positions reads $\sigma_y^2(t) = \frac{1}{4}Y_{\gamma_0}^2\left(2\overline{\omega}_rt - \sin(2\overline{\omega}_rt)\right).$ At

426 $t = 2\pi/\overline{\omega}_r$, the group has oscillated once around the jet and the maximal position extension 427 reaches $\pm 2\sigma_y = \pm 2\sqrt{\pi} Y_{\gamma_0} = \pm 42$ km, well confirmed by ray simulations. In contrast, usual 428 fast wave models (e.g. Smit & Janssen 2019) do not consider the interplay between smooth 429 and rough currents, and hence solely predict a classical scattering with a much faster vertical 430 location spreading: $\pm 2\sigma_y = \pm 2\sqrt{(2\pi)^3/3} Y_{\gamma_0} = \pm 217$ km. For large time, our multiscale 431 approach predicts a scaling in *t*, much slower than the usual scattering t^3 scaling.

From the group vertical location and wavevector angle, we can also solve (4.5) analytically to estimate the group wavenumber variations. For small wavevector angles, $-\int_0^t \overline{\sigma} \sin(\zeta) dt' \approx 2 \int_0^t \overline{\omega} \theta_k dt' = 2\overline{\beta} \int_0^t y'_r \theta_k dt'$ and (4.9) together with the analytic solutions for y'_r and θ_k give a closed stochastic expression for the group wavenumber. Thus, the wavenumber factor $\exp(2\overline{\beta} \int_0^t y'_r \theta_k dt')$ oscillates at frequency $2\overline{\omega}_r$ and the oscillations modulate the wave amplitude: $h = \sqrt{E} = \sqrt{\omega_0 N} = \operatorname{cst.} k^{\frac{1}{4}}$. The modulations are associated with wave-current energy exchanges Boury *et al.* (2023), visible in the colored rays, figures 6, 7 and 8, when the groups enter and exit the jet.

Finally, the conditional ray distribution, $p(\mathbf{x}, \mathbf{k} | \mathbf{x}_r^0, \mathbf{k}^0, t)$, the action spectrum mean from (4.10) and the action mean from (4.11) can all be derived. For a system initially localized in $(0, \frac{L_y}{2})$ with action A^0 , wavenumber k^0 and a $\sigma^0_{\delta\theta_k}$ -width Gaussian angular spreading, propagating to the right, the action mean reads

444
$$\mathbb{E}A(x, y, t) = A^0 \delta\left(x - (v_g^0(k^0) - \overline{u}(y))t\right) \mathcal{N}\left(y - \frac{L_y}{2} \big| \tilde{\sigma}_y^2(t) \right), \tag{5.7}$$

445 with $\mathcal{N}\left(\bullet | \tilde{\sigma}_{y}^{2}(t)\right)$, a Gaussian function with variance $\tilde{\sigma}_{y}^{2}(t) = \sigma_{y}^{2}(t) + \left(\frac{v_{g}^{0}}{\overline{\omega}_{r}}\sin(\overline{\omega}_{r}t)\sigma_{\delta\theta_{k}}^{0}\right)^{2}$. 446 The action is advected in the horizontal direction, and slowly diffuses along the vertical 447 direction.

448 6. Conclusion

Developed to generalize the ray-path concept for waves propagating over an heterogeneous 449 turbulence, a practical stochastic framework is derived. For fast waves, the smallest scales 450 of a turbulent flow decorrelate along the wave propagation. Flows with steeper spectra 451 452 decorrelate faster, leading to a broader validity range of fast wave approximations. The proposed framework encodes both large-scale refraction and random scattering effects on 453 wave statistical properties. The mean wave-action statistics are directly linked to resolved 454 strain-rate and vorticity, but also to unresolved KE spectral properties. Both Eulerian and 455 Lagrangian views are presented. A convenient calibration method is also proposed for the 456 457 subgrid parametrization.

As anticipated, random horizontal currents delay wave arrival and augment the initial 458 radiative transport equation with a directional diffusive term. These phenomena are illustrated 459 with numerical simulations, analytical solutions, and quantitative proxies describing weak 460 homogeneous turbulence. Using these proxies, measured delays in ray arrivals, estimated 461 wave energy spectral characteristics and decays, and/or varying directional spread shall then 462 be used to more quantitatively interpreted the upper the turbulent underlying flow properties. 463 The generalized fast wave approximation does takes into account wavenumber variation 464 and handles strong heterogeneous flows, like localized jets with strong current gradients. As 465 compared to numerical simulations, numerical and theoretical results explain and quantify 466 ray trapping effects by jets, unlike usual fast wave approaches. 467

Among fast wave literature, isotropic diffusion and, hence, wavenumber diffusion may (e.g. Voronovich 1991) or may not (e.g. Bôas & Young 2020) comes into play (see Appendix 470 C for details). Future works could adapt our convenient stochastic calculus framework to the second models family. Besides, further analytical developments could consider finite-471 size wave groups, their dynamics (Jonsson 1990; White & Fornberg 1998) and statistical 472 distributions, or alternatively the Eulerian action dynamics (4.12) with all its multi-point 473 stochastic structure. When achieved, this next theoretical development could provide new 474 means to analyze wave dynamics with subsequent fast simulations of ensembles. Beside 475 comprehension and analysis, our stochastic simulation tools aim at eventually facilitate 476 future ensemble-based data assimilation algorithms (Smit et al. 2021). 477

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- 486 Author contributions. VR developed the theory. VR and EH wrote the code and performed numerical
- 487 experiments. VR and BC wrote the paper.

488 Appendix A. Stochastic forcing covariance

In this appendix, we will compute the conditionnal covariance of the stochastic forcing of our eikonal characteristic equations (4.1), that is:

491
$$2\boldsymbol{D} \triangleq \frac{1}{\mathrm{d}t} \mathbb{E}_t \left\{ \begin{pmatrix} \sigma \mathrm{d}B_t \\ \mathrm{d}\boldsymbol{\eta}_t \end{pmatrix} \begin{pmatrix} \sigma \mathrm{d}B_t \\ \mathrm{d}\boldsymbol{\eta}_t \end{pmatrix}^\mathsf{T} \right\} = \begin{bmatrix} \boldsymbol{a} & \boldsymbol{\Sigma}_{\boldsymbol{\eta},\sigma} \\ \boldsymbol{\Sigma}_{\boldsymbol{\eta},\sigma}^\mathsf{T} & \boldsymbol{\Sigma}_{\boldsymbol{\eta}} \end{bmatrix}$$
(A1)

492 . where, $d\eta_t = -\nabla (\sigma dB_t)^T k$, denotes the wave-vector stochastic forcing and, $\Sigma_{\eta} dt$, its 493 covariance, and $\mathbb{E}_t \{\bullet\} \stackrel{\triangle}{=} \mathbb{E} \{\bullet | \mathbf{x}_r(t), \mathbf{k}(t)\}$ stands for the conditional expectation evaluated 494 with given characteristics $(\mathbf{x}_r(t), \mathbf{k}(t))$ at the current time *t*. Note that in the appendix we 495 use Itō notations only.

The subgrid velocity, $v' = \sigma dB_t/dt$ is constructed in Fourier space with a divergence-free isotropic spatial filter $\nabla^{\perp} \check{\psi}_{\sigma}$ (see (3.11)).

498
$$\widehat{\boldsymbol{v}'}(\boldsymbol{\kappa},t) = \int d\boldsymbol{x} \ \boldsymbol{v}'(\boldsymbol{x},t) e^{-i\boldsymbol{\kappa}\cdot\boldsymbol{x}} = \widehat{\sigma dB_t}/dt(\boldsymbol{\kappa},t) = i\boldsymbol{\kappa}^{\perp} \widetilde{\psi}_{\sigma}(\boldsymbol{\kappa}) \widehat{dB_t}(\boldsymbol{\kappa})/dt, \qquad (A\,2)$$

499 where κ^{\perp} is the vector directly orthogonal to κ . The computation of the variance tensor **a** 500 is classical and straightforward from the definition of the inverse Fourier transform and the 501 identity $\mathbb{E}\left\{\widehat{dB_t}(\kappa_1)\widehat{dB_t}^*(\kappa_2)\right\} = (2\pi)^2 \delta(\kappa_1 - \kappa_2) dt$, where * denotes complex conjugate. We 502 simply need to split the integral of the stochastic forcing spectrum over the current wavevector

503 $\kappa = \kappa(\cos \theta_{\kappa}, \sin \theta_{\kappa})$:

504
$$\boldsymbol{a} = \frac{1}{(2\pi)^4 \mathrm{d}t} \iint \mathrm{d}\boldsymbol{\kappa}_1 \mathrm{d}\boldsymbol{\kappa}_2 \mathbb{E}_t \left\{ (\widehat{\sigma \mathrm{d}B_t})(\boldsymbol{\kappa}_1, \boldsymbol{k}) (\widehat{\sigma \mathrm{d}B_t}^{\mathsf{T}})^*(\boldsymbol{\kappa}_2, \boldsymbol{k}) \right\} e^{\mathrm{i}(\boldsymbol{\kappa}_1 - \boldsymbol{\kappa}_2) \cdot \boldsymbol{x}}, \quad (A3)$$

505
$$= \frac{1}{(2\pi)^2} \int d\kappa \, \kappa^2 |\tilde{\psi}_{\sigma}(\kappa)|^2 \begin{pmatrix} -\sin\theta_{\kappa} \\ \cos\theta_{\kappa} \end{pmatrix} \begin{pmatrix} -\sin\theta_{\kappa} \\ \cos\theta_{\kappa} \end{pmatrix}^{\mathsf{T}}, \tag{A4}$$

$$506 \qquad \qquad = \frac{1}{(2\pi)^2} \int_0^{+\infty} \oint_0^{2\pi} d\kappa d\theta_\kappa \kappa^3 |\tilde{\psi}_{\sigma}(\kappa)|^2 \begin{pmatrix} \sin^2 \theta_\kappa & -\sin \theta_\kappa \cos \theta_\kappa \\ -\sin \theta_\kappa \cos \theta_\kappa & \cos^2 \theta_\kappa \end{pmatrix}, \quad (A5)$$

507
$$= \frac{2}{2\pi} a_0 \oint_0^{2\pi} d\theta_\kappa \begin{pmatrix} \sin^2 \theta_\kappa & -\sin \theta_\kappa \cos \theta_\kappa \\ -\sin \theta_\kappa \cos \theta_\kappa & \cos^2 \theta_\kappa \end{pmatrix},$$
(A 6)

(A7)

508

509 where a_0 is defined by (4.7).

 $= a_0 \mathbb{I}_d$

510 Now, the Fourier transform of the wave-vector stochastic forcing is

511
$$d\widehat{\boldsymbol{\eta}}_t = -\overline{\boldsymbol{\nabla}(\sigma dB_t)}^{\mathsf{T}} \boldsymbol{k} = -i\boldsymbol{\kappa}(i\boldsymbol{\kappa}^{\perp}\widetilde{\boldsymbol{\psi}}_{\sigma}\widehat{dB_t}) \cdot \boldsymbol{k} = \boldsymbol{\kappa}(\boldsymbol{\kappa}^{\perp} \cdot \boldsymbol{k})\widetilde{\boldsymbol{\psi}}_{\sigma}\widehat{dB_t} = -\boldsymbol{\kappa}(\boldsymbol{k}^{\perp} \cdot \boldsymbol{\kappa})\widetilde{\boldsymbol{\psi}}_{\sigma}\widehat{dB_t}.$$
 (A 8)

512 Then, applying the crest-oriented rotation matrix, $M_k = \begin{bmatrix} \tilde{k} & \tilde{k}^{\perp} \end{bmatrix}$, leads to

513
$$d\widehat{\mathbf{Z}}_{t} = \mathbf{M}_{k}^{\mathsf{T}} d\widehat{\boldsymbol{\eta}}_{t} = -\begin{pmatrix} \widetilde{\mathbf{k}} \cdot \boldsymbol{\kappa} \\ \widetilde{\mathbf{k}}^{\perp} \cdot \boldsymbol{\kappa} \end{pmatrix} (\mathbf{k}^{\perp} \cdot \boldsymbol{\kappa}) \widetilde{\psi}_{\sigma} \widehat{\mathbf{d}B_{t}} = -\begin{pmatrix} \cos \delta\theta \sin \delta\theta \\ \sin^{2} \delta\theta \end{pmatrix} \boldsymbol{\kappa}^{2} k \, \widetilde{\psi}_{\sigma} \widehat{\mathbf{d}B_{t}}, \quad (A 9)$$

514 with $\delta \theta = \theta_{\kappa} - \theta_{k}$. From there, we can evaluate the conditional covariance matrix $\Sigma_{Z} = \frac{1}{dt} \mathbb{E}_{t} \{ d\mathbf{Z}_{t} d\mathbf{Z}_{t}^{\mathsf{T}} \}$ of $d\mathbf{Z}_{t}$ as before:

516
$$\Sigma_{Z} = \frac{1}{(2\pi)^{4} dt} \iint d\kappa_{1} d\kappa_{2} \mathbb{E}_{t} \left\{ (d\widehat{\boldsymbol{Z}}_{t})(\boldsymbol{\kappa}_{1}, \boldsymbol{k}) (d\widehat{\boldsymbol{Z}}_{t}^{\mathsf{T}})^{*}(\boldsymbol{\kappa}_{2}, \boldsymbol{k}) \right\} e^{\mathbf{i}(\boldsymbol{\kappa}_{1} - \boldsymbol{\kappa}_{2}) \cdot \boldsymbol{x}}, \tag{A 10}$$

517
$$= \frac{1}{(2\pi)^2} \int_0^{+\infty} \oint_0^{2\pi} d\kappa d\delta\theta \ \kappa^5 k^2 |\tilde{\psi}_{\sigma}(\kappa)|^2 \begin{pmatrix} \cos^2 \delta\theta \sin^2 \delta\theta & \cos \delta\theta \sin^3 \delta\theta \\ \cos \delta\theta \sin^3 \delta\theta & \sin^4 \delta\theta \end{pmatrix} (A \ 11)$$

518
$$= \gamma_0 k^2 \begin{bmatrix} 1 & 0 \\ 0 & 3 \end{bmatrix}.$$
 (A 12)

519 Finally, we come back to the canonical frame to get Σ_{η}

520
$$\Sigma_{\eta} = \mathbb{E}_{t} \left\{ \mathrm{d}\boldsymbol{\eta}_{t} \mathrm{d}\boldsymbol{\eta}_{t}^{\mathsf{T}} \right\} = \boldsymbol{M}_{k} \Sigma_{Z} \boldsymbol{M}_{k}^{\mathsf{T}} = \gamma_{0} k^{2} \left[\tilde{\boldsymbol{k}} \tilde{\boldsymbol{k}}^{\mathsf{T}} + 3 \tilde{\boldsymbol{k}}^{\perp} \left(\tilde{\boldsymbol{k}}^{\perp} \right)^{\mathsf{T}} \right].$$
(A 13)

521 For noises cross-correlations, by isotropy, it is also straightforward to show that

522
$$\Sigma_{\eta,\sigma} = 0. \tag{A14}$$

523 The stochastic forcings of x_r and k are hence (conditionally) independent from one another.

524 Appendix B. Single ray dynamics

The Itō noise $\begin{pmatrix} \sigma dB_t \\ d\eta_t \end{pmatrix}$ is white in time and conditionally Gaussian. Its conditional singlepoint distribution is fully determined by its zero mean and its local covariance matrix (given by equations (A 1), (A 7), (A 13) and (A 14)). In particular, we can replace this noise by another zero-mean Gaussian vector with the same covariance without changing the singleray dynamics – typically replacing σdB_t by $\sqrt{a_0} \begin{pmatrix} dB_t^{(1)} \\ dB_t^{(2)} \end{pmatrix}$ and $d\mathbf{Z}_t$ by $-\sqrt{\gamma_0}k \begin{pmatrix} dB_t^{(3)} \\ \sqrt{3}dB_t^{(4)} \end{pmatrix}$. It yields the simplified ray equations (4.3)-(4.4).

531 Then note that from Itō lemma (Oksendal 1998) $d\tilde{k} = d \begin{pmatrix} \cos \theta_k \\ \sin \theta_k \end{pmatrix} = \tilde{k}^{\perp} d\theta_k - \frac{1}{2} \tilde{k} d < 532 \quad \theta_k, \theta_k >_t$ where $\langle \bullet, \bullet \rangle_t$ denotes the quadratic covariation. Thus,

533
$$\mathbf{d}\mathbf{k} = \mathbf{d}k\,\tilde{\mathbf{k}} + k\mathbf{d}\tilde{\mathbf{k}} + \mathbf{d} < k, \,\tilde{\mathbf{k}} >= (\mathbf{d}k - \frac{1}{2}k\mathbf{d} < \theta_k, \theta_k >_t)\tilde{\mathbf{k}} + (k\mathbf{d}\theta_k + \mathbf{d} < k, \theta_k >_t)\tilde{\mathbf{k}}^{\perp} (\mathbf{B} 1)$$

534 Projecting this equation and $d\mathbf{k} = -\nabla \overline{\mathbf{v}}^{\mathsf{T}} \mathbf{k} dt + d\eta_t$ on $\tilde{\mathbf{k}}$ and $\tilde{\mathbf{k}}^{\perp}$, we have

$$\begin{cases} dk = -\tilde{k} \cdot \nabla \overline{v}^{\mathsf{T}} k dt + (dZ_t)_1 + \frac{1}{2} k d < \theta_k, \theta_k >_t \\ k d\theta_k = -\tilde{k}^{\perp} \cdot \nabla \overline{v}^{\mathsf{T}} k dt + (dZ_t)_2 - d < k, \theta_k >_t \end{cases},$$
(B 2)

535

 $\begin{cases} \mathrm{d}k = -\tilde{k} \cdot \nabla \overline{\nu}^{\mathsf{T}} k \, \mathrm{d}t + (\mathrm{d}Z_t)_1 + \frac{1}{2} k^{-1} \mathrm{d} < Z_2, Z_2 >_t \\ \mathrm{d}\theta_k = -\tilde{k}^{\perp} \cdot \nabla \overline{\nu}^{\mathsf{T}} \tilde{k} \, \mathrm{d}t + k^{-1} (\mathrm{d}Z_t)_2 + \frac{1}{2} k^{-2} \mathrm{d} < Z_1, Z_2 >_t \end{cases}$ (B 3)

The treatment of the large-scale terms $\tilde{k} \cdot \nabla \overline{v}^{\mathsf{T}} \tilde{k}$ and $\tilde{k}^{\perp} \cdot \nabla \overline{v}^{\mathsf{T}} \tilde{k}$ is classical. Interested readers can refer to Lapeyre *et al.* (1999) for details. From Ito lemma again, d log $k = dk/k - \frac{1}{2}d < k, k > t/k^2$ leading to the simplified wave-vector dynamics (4.5)-(4.6).

540 Appendix C. Subgrid flow anisotropy and comparison with other works

Throughout this paper, we have considered an isotropic model for the stochastic subgrid 541 velocity (3.11). The isotropic diffusivity matrix $\mathbf{a} = a_0 \mathbb{I}_d$ is a good illustration of this. In 542 contrast, many authors (e.g. White & Fornberg 1998; Bôas & Young 2020; Smit & Janssen 543 2019) assume isotropic and homogeneous turbulence and obtain anisotropic stochastic 544 subgrid models for $\frac{\|\tilde{\boldsymbol{\nu}}\|}{v_{\perp}^{0}} \to 0$. In these approaches, the integral over $\delta\theta$ in diffusivity matrix 545 computations (A 4) and (A 12) involve singular integrations over the direction $v_g^0 = v_g^0 \tilde{k}$. It makes appear a Dirac delta function, $2\pi\delta(\kappa \cdot v_g^0) = \frac{2\pi}{\kappa v_g^0} (\delta(\theta_\kappa - \theta_k - \frac{\pi}{2}) + \delta(\theta_\kappa - \theta_k + \frac{\pi}{2}))$ 546 547 (see Appendix in Bôas & Young 2020). This precision imposes a statistical anisotropy for σdB_t (oriented along k) and $d\eta_t$ (oriented along k^{\perp}), eventually leading to a covariance 548 549 $\Sigma_Z = \gamma_0 k^2 \begin{vmatrix} 0 & 0 \\ 0 & 16 \end{vmatrix}$ (Eq. (3.17) in Bôas & Young (2020) and Eq. (24) in Smit & Janssen 550 (2019)), no noise dZ_1 , and no Brownian motion $B_t^{(3)}$. Moreover, because of scaling assumption, Bôas & Young (2020) neglect the spatial diffusivity matrix, **a**, while Smit 551 552 & Janssen (2019) find $\mathbf{a} = 4a_0 \left(\mathbb{I}_d + \frac{5}{4} \tilde{\mathbf{k}} \tilde{\mathbf{k}}^{\mathsf{T}} \right)$ (Eq. (22)-(23)). In this anisotropic framework, 553 the Stratonovich wavevector equation (2.3), $dk = -\nabla (\overline{\nu} dt + \sigma \circ dB_t)^T k$, would involve an 554 additional drift term in Ito notations. 555

Further developing this anisotropic stochastic closure is an interesting avenue. A multiscale anisotropic stochastic closure would involve wavenumber variations but no wavenumber diffusion. Nevertheless, in the present study, we adopt the isotropic model for σdB_t , which is much more convenient for multi-ray numerical simulations.

560 Appendix D. Action spectra and ray distribution

561 Here we highlight the link between mean action spectral density and the ray distribution.

562 We denote by N^0 the initial wave action spectrum. We first use the definition of the Dirac

measure then we perform a variable change corresponding to the characteristic (2.5) from $t = t_i$ to $t = t_f$:

565
$$\mathbb{E}N(\boldsymbol{x},\boldsymbol{k},t) = \mathbb{E}\iint \mathrm{d}\boldsymbol{x}_r \mathrm{d}\boldsymbol{k}_r \ N(\boldsymbol{x}_r,\boldsymbol{k}_r,t)\delta((\boldsymbol{x}_r,\boldsymbol{k}_r)-(\boldsymbol{x},\boldsymbol{k})), \tag{D1}$$

566
$$= \mathbb{E} \iint \mathbf{d} \mathbf{x}_r^0 \mathbf{d} \mathbf{k}_r^0 N(\mathbf{x}_r^0, \mathbf{k}_r^0, 0) \delta((\mathbf{x}_r, \mathbf{k}_r)(\mathbf{x}_r^0, \mathbf{k}_r^0, t) - (\mathbf{x}, \mathbf{k})), \quad (D 2)$$

567
$$= \iint \mathbf{d} \mathbf{x}_r^0 \mathbf{d} \mathbf{k}^0 \ N^0(\mathbf{x}_r^0, \mathbf{k}_r^0) p(\mathbf{x}, \mathbf{k} | \mathbf{x}_r^0, \mathbf{k}^0, t), \tag{D3}$$

where the standard relation between the Dirac measure and the probability distribution function has been used.

570 Appendix E. Jet simulation

Again, currents are simulated at a resolution 512×512 on a 1000-km-width squared domain [0, L_x] × [0, L_y] through the same code. A backward velocity v_{Bk} forces a leftward jet structure.

574
$$\partial_t \omega + \mathbf{v} \cdot \nabla \omega = S_\omega \text{ with } \mathbf{v} = \nabla^\perp \Delta^{-1}(\omega + \omega_{\text{Bk}}).$$
 (E1)

575 S_{ω} encompasses the linear drag and the hyperviscosity with coefficient $1/\tau_F = 3.22 \times 10^{-8} s^{-1}$ 576 and $\nu_{\rm HV}/dx^8 = 3.33 \times 10^{-9} s^{-1}$ respectively. The background vorticity, $\omega_{\rm Bk}$, is a smooth step 577 function with a wavy interface at $y = Y_{\rm Bk}(x)$:

578
$$\omega_{Bk}(x,y) = \Omega_{Bk}\left(\frac{1}{2} - \operatorname{erf}\left(\frac{y - Y_{Bk}(x)}{L_y^{\omega}}\right)\right) \text{ with } Y_{Bk}(x) = L_y\left(\frac{1}{2} + \frac{1}{30}\cos\left(\frac{2\pi}{L_x}x\right)\right).$$
 (E 2)

To better highlight the interplay between ray oscillations and scattering, we consider very-collimated swells, with a spatial extension of $100\lambda = 25$ km.

Besides, the curvature of the simulated jet can force an additional faster oscillation around the jet for small enough wavevector angle. Indeed, a wave group traveling exactly rightward would cross an alternation of positive and negative vorticity regions with a period $L_x/(v_g^0 - \overline{U}_0) \approx 1$ day $< 2\pi/\overline{\omega}_r$. Here, we set an initial wavevector angle large enough to prevent the additional harmonics.

REFERENCES

- BAL, GUILLAUME & CHOU, TOM 2002 Capillary–gravity wave transport over spatially random drift. *Wave Motion* 35 (2), 107–124.
- BAUER, WERNER, CHANDRAMOULI, PRANAV, CHAPRON, BERTRAND, LI, LONG & MÉMIN, ETIENNE 2020
 Deciphering the role of small-scale inhomogeneity on geophysical flow structuration: a stochastic
 approach. *Journal of Physical Oceanography* 50 (4), 983–1003.
- Bôas, ANA B VILLAS & YOUNG, WILLIAM R 2020 Directional diffusion of surface gravity wave action by
 ocean macroturbulence. *Journal of Fluid Mechanics* 890.

BORCEA, LILIANA, GARNIER, JOSSELIN & SOLNA, KNUT 2019 Wave propagation and imaging in moving
 random media. *Multiscale Modeling & Simulation* 17 (1), 31–67.

BOURY, SAMUEL, BÜHLER, OLIVER & SHATAH, JALAL 2023 Fast-slow wave transitions induced by a random
 mean flow. *Physical Review E* 108 (5), 055101.

597 BÜHLER, OLIVER 2009 Waves and mean flows. Cambridge University Press.

COTTER, COLIN J, GOTTWALD, GEORG A & HOLM, DARRYL D 2017 Stochastic partial differential fluid
 equations as a diffusive limit of deterministic lagrangian multi-time dynamics. *Proceedings of the Royal Society A: Mathematical, Physical and Engineering Sciences* 473 (2205), 20170388.

- 601Cox, MICHAEL R, KAFIABAD, HOSSEIN A & VANNESTE, JACQUES 2023 Inertia-gravity-wave diffusion by602geostrophic turbulence: the impact of flow time dependence. Journal of Fluid Mechanics 958, A21.
- CRISAN, DAN & HOLM, DARRYL D 2018 Wave breaking for the stochastic camassa-holm equation. *Physica D: Nonlinear Phenomena* 376, 138–143.
- DINVAY, EVGUENI & MÉMIN, ETIENNE 2022 Hamiltonian formulation of the stochastic surface wave problem.
 Proceedings of the Royal Society A 478 (2265), 20220050.
- DONG, WENJING, BÜHLER, OLIVER & SMITH, K SHAFER 2020 Frequency diffusion of waves by unsteady
 flows. *Journal of Fluid Mechanics* 905, R3.
- DYSTHE, KRISTIAN B 2001 Refraction of gravity waves by weak current gradients. *Journal of Fluid Mechanics* 442, 157–159.
- FREUND, JB & FLEISCHMAN, TG 2002 Ray traces through unsteady jet turbulence. *International Journal of Aeroacoustics* 1 (1), 83–96.
- GARNIER, JOSSELIN, GAY, ETIENNE & SAVIN, ERIC 2020 Multiscale analysis of spectral broadening of acoustic
 waves by a turbulent shear layer. *Multiscale Modeling & Simulation* 18 (2), 798–823.
- Held, I., PIERREHUMBERT, R., GARNER, S. & SWANSON, K. 1995 Surface quasi-geostrophic dynamics.
 Journal of Fluid Mechanics 282, 1–20.
- HELL, MOMME C, FOX-KEMPER, BAYLOR & CHAPRON, BERTRAND in preparation An efficient wave model for
 surface wave growth and propagation in coupled climate models. *Journal of Advances in Modeling Earth Systems*.
- HELLER, EJ, KAPLAN, L & DAHLEN, A 2008 Refraction of a gaussian seaway. *Journal of Geophysical Research: Oceans* 113 (C9).
- HOLM, D. 2015 Variational principles for stochastic fluid dynamics. *Proceedings of the Royal Society of London A: Mathematical, Physical and Engineering Sciences* 471 (2176).
- HOLM, DARRYL D 2021 Stochastic variational formulations of fluid wave–current interaction. *Journal of nonlinear science* 31 (1), 4.
- HOLM, DARRYL D, HU, RUIAO & STREET, OLIVER D 2023 On the interactions between mean flows and
 inertial gravity waves. *arXiv preprint arXiv:2302.04838*.
- HOLM, DARRYL D & LUESINK, ERWIN 2021 Stochastic wave–current interaction in thermal shallow water
 dynamics. *Journal of Nonlinear Science* 31, 1–56.
- 630 JONSSON, IVAR G 1990 Wave-current interactions. The sea 9, 65–120.
- KAFIABAD, HOSSEIN A, SAVVA, MILES AC & VANNESTE, JACQUES 2019 Diffusion of inertia-gravity waves
 by geostrophic turbulence. *Journal of Fluid Mechanics* 869, R7.
- KLYATSKIN, V. 2005 Stochastic equations through the eye of the physicist: Basic concepts, exact results and
 asymptotic approximations. Elsevier.
- KLYATSKIN, VI & KOSHEL, KV 2015 Anomalous sea surface structures as an object of statistical topography.
 Physical Review E 91 (6), 063003.
- KUDRYAVTSEV, VLADIMIR, YUROVSKAYA, MARIA, CHAPRON, BERTRAND, COLLARD, FABRICE & DONLON,
 CRAIG 2017 Sun glitter imagery of ocean surface waves. part 1: Directional spectrum retrieval and
 validation. Journal of Geophysical Research: Oceans 122 (2), 1369–1383.
- 640 KUNITA, H. 1997 Stochastic flows and stochastic differential equations, , vol. 24. Cambridge university press.
- 641 LANDAU, LEV DAVIDOVICH & LIFSHITS, EVGENIĬ MIKHAĬLOVICH 1960 Mechanics, , vol. 1. CUP Archive.
- 642 LAPEYRE, GUILLAUME 2017 Surface quasi-geostrophy. Fluids 2 (1), 7.
- LAPEYRE, G., KLEIN, P. & HUA, B. 1999 Does the tracer gradient vector align with the strain eigenvectors in
 2D turbulence? *Physics of Fluids* 11 (12), 3729–3737.
- LAVRENOV, IGOR 2013 Wind-waves in oceans: dynamics and numerical simulations. Springer Science &
 Business Media.
- MCCOMAS, C HENRY & BRETHERTON, FRANCIS P 1977 Resonant interaction of oceanic internal waves.
 Journal of Geophysical Research 82 (9), 1397–1412.
- MÉMIN, E. 2014 Fluid flow dynamics under location uncertainty. *Geophysical & Astrophysical Fluid Dynamics* 108 (2), 119–146.
- MÉMIN, ETIENNE, LI, LONG, LAHAYE, NOÉ, TISSOT, GILLES & CHAPRON, BERTRAND 2022 Linear wave
 solutions of a stochastic shallow water model. In *Stochastic Transport in Upper Ocean Dynamics Annual Workshop*, pp. 223–245. Springer Nature Switzerland Cham.
- MIKULEVICIUS, R. & ROZOVSKII, B. 2004 Stochastic Navier–Stokes equations for turbulent flows. SIAM
 Journal on Mathematical Analysis 35 (5), 1250–1310.
- 656 OKSENDAL, B. 1998 Stochastic differential equations. Spinger-Verlag.

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- PAPANICOLAOU, G. & KOHLER, W. 1974 Asymptotic theory of mixing stochastic ordinary differential
 equations. *Communications on Pure and Applied Mathematics* 27 (5), 641–668.
- PITERBARG, LEONID & OSTROVSKII, A 1997 Advection and diffusion in random media: implications for sea
 surface temperature anomalies. Kluwer Academic.
- PLOUGONVEN, RIWAL & ZHANG, FUQING 2014 Internal gravity waves from atmospheric jets and fronts.
 Reviews of Geophysics 52 (1), 33–76.
- Resseguier, Valentin, Li, Long, Jouan, Gabriel, Dérian, Pierre, Mémin, Etienne & Bertrand,
 CHAPRON 2020a New trends in ensemble forecast strategy: uncertainty quantification for coarse-grid
 computational fluid dynamics. Archives of Computational Methods in Engineering pp. 1–82.
- Resseguier, VALENTIN, MÉMIN, ETIENNE & CHAPRON, BERTRAND 2017a Geophysical flows under location
 uncertainty, part I random transport and general models. *Geophysical & Astrophysical Fluid Dynamics* 111 (3), 149–176.
- Resseguier, Valentin, Mémin, Etienne & Chapron, Bertrand 2017b Geophysical flows under location
 uncertainty, part II quasi-geostrophy and efficient ensemble spreading. *Geophysical & Astrophysical Fluid Dynamics* 111 (3), 177–208.
- Resseguier, Valentin, Pan, Wei & Fox-Kemper, Baylor 2020b Data-driven versus self-similar
 parameterizations for stochastic advection by lie transport and location uncertainty. *Nonlinear Processes in Geophysics* 27 (2), 209–234.
- SLUNYAEV, AV & SHRIRA, VI 2023 Extreme dynamics of wave groups on jet currents. *Physics of Fluids* 35 (12).
- SMIT, PB, HOUGHTON, IA, JORDANOVA, K, PORTWOOD, T, SHAPIRO, E, CLARK, D, SOSA, M & JANSSEN, TT
 2021 Assimilation of significant wave height from distributed ocean wave sensors. *Ocean Modelling* 159, 101738.
- SMIT, PIETER B & JANSSEN, TIM T 2019 Swell propagation through submesoscale turbulence. *Journal of Physical Oceanography* 49 (10), 2615–2630.
- VORONOVICH, A. 1991 The effect of shortening of waves on random currents. In *Proceedings of nonlinear water waves*. Bristol.
- WANG, HAN, BÔAS, ANA B VILLAS, YOUNG, WILLIAM R & VANNESTE, JACQUES 2023 Scattering of swell by
 currents. arXiv preprint arXiv:2305.12163.
- 686 WEST, BRUCE J 1978 Ray paths in a fluctuating environment. Physical Review A 18 (4), 1646.
- 687 WHITE, BENJAMIN S 1999 Wave action on currents with vorticity. Journal of Fluid Mechanics 386, 329–344.
- WHITE, BENJAMIN S & FORNBERG, BENGT 1998 On the chance of freak waves at sea. Journal of fluid
 mechanics 355, 113–138.
- ZHEN, YICUN, RESSEGUIER, VALENTIN & CHAPRON, BERTRAND 2023 Physically constrained covariance
 inflation from location uncertainty. *EGUsphere* 2023, 1.