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1 Monitoring of natural oil seepage in the Lower Congo Basin using

2 SAR observations

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11 **Abstract**

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Synthetic Aperture Radar (SAR) is a spaceborne tool allowing near real-time imagery over large ground footprints of areas hundreds of kilometres wide. SAR also allows persistent observations of the process of oil discharges, benefiting from (1) day and night observations, (2) independence of cloud cover and (3) high revisiting frequencies. The interpretation of 82 SAR scenes over the Lower Congo Basin for the 1994-2012 period allowed the recognition of 1400 natural seepage slicks associated with 102 individual seep sites. We acquired an additional SAR data set with a short revisit time over a selected prolific area. The data set consisted of 22 SAR scenes acquired over a 10-day period, meaning a maximum revisiting period of 12 hours between subsequent SAR acquisitions. The short-term approach shows that seepage slicks were detected with wind speeds between 1.5 and 6.5 m/s (2.91 to 12.63 knots). Both long-term and short-term monitoring evidences that seepage events are intermittent, with the occurrence rate of oil release ranging from 5 to 80%. Short-term monitoring shows that the seepage pace is independent between seep sites, suggesting contrasted controlling factors. It also shows that the residence time of seepage slicks is systematically shorter than 12 hours, which is far below reported values. Understanding the seepage slick residence time is essential to provide an accurate estimation of the fluid flow and to compute the volumetric outflow. The integration of in-situ mooring points measuring the current velocity 10 m below the sea surface, in addition to local wind field and slick length, allowed us to more accurately estimate the residence time of oil slicks at the sea surface before vanishing. The results show that the sea surface residence time of oil slicks is limited to a few hours, with a median value of 3h15min. These new residence time estimations allowed us to propose a quantification of the regional oil output in the Lower Congo Basin, estimated at 4380 m³/year. This area may therefore be considered as the world's third biggest oil-supplying province from natural leakages.

1. Introduction

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36 Natural discharges of hydrocarbon fluids occur in most of the world's petroleum basins 37 (Kvenvolden and Cooper, 2003; National Research Council, 2003; Kvenvolden and Rogers, 2005; 38 Etiope, 2015). Because of its sensitivity to centimetre wavelengths (e.g., McCandless and 39 Jackson, 2003), Synthetic Aperture Radar (SAR) is a remotely operated tool widely used in ocean 40 studies for the recognition of sea surface patterns. Satellites with an on-board SAR system are 41 used to detect natural liquid hydrocarbon seepage at the sea surface (Johannessen et al., 2000; Espedal and Johannessen, 2000; MacDonald et al., 2002; Brekke and Solberg 2005; Garcia-42 43 Pineda et al., 2009, Fingas and Brown 2014) in complement to spaceborne optical data 44 (MacDonald et al., 1993; Macdonald et al., 2002; Hu et al., 2009; Chen and Hu, 2014). The 45 presence of oil flattens the sea surface and significantly reduces the amount of energy 46 backscattered to the satellite (Fresnel specular reflection) compared to the surrounding free water. 47 Sea surface covered by oil seepage appears as black streaks on SAR images (Johannessen et al., 48 2000; Espedal and Johannessen, 2000; Williams and Lawrence, 2002; Ivanov et al., 2007; 49 Zatyagalova et al., 2007; Garcia-Pineda et al., 2009; Körber et al., 2014). Up to now, research on 50 sea surface manifestations of thermogenic seeps has focused on a few case studies, including the 51 Caspian Sea (Ivanov et al., 2007; Zatyagalova et al., 2007), the Black Sea (Körber et al., 2014) and 52 the Gulf of Mexico (Garcia-Pineda et al., 2014, Macdonald et al., 2015). The SAR method also 53 serves as a primary tool for the evaluation of released volumes of heavy hydrocarbon discharges 54 (MacDonald et al., 1993; Mitchell et al., 2000; Garcia-Pineda et al., 2014), even if quantifications 55 may have been potentially underestimated with this technique (Hu et al., 2009; Garcia-Pineda et 56 al., 2010; Smith et al., 2014). Oil seep systems are usually accompanied by gas escapes (Solomon 57 et al., 2009; Hu et al., 2012) originating from the thermal cracking and the second methanogenesis from anaerobic biodegradation (Head et al., 2003; Peters et al., 2007). For environmental 58 59 purposes, quantifying natural hydrocarbon escapes is therefore compulsory to understand their 60 implications on the global carbon budget. Understanding the amount of expelled hydrocarbon at the regional scale along continental margins also presents great implications for petroleum systems evaluation. The sea surface residence time of oil slicks is an important input parameter for the quantification of oil output but still remains poorly documented due to low revisiting frequencies of SAR data. The Lower Congo Basin is considered as an active petroleum province where geophysical investigations of the seafloor have allowed the recognition of high-density fluid flow features (Gay et al., 2003; Gay et al., 2006a, b, c; Gay et al., 2007; Andresen et al., 2010; Andresen and Huuse, 2011; Anka et al., 2013) of probable thermogenic origin (Gay et al., 2002; Gay et al., 2006b; Andresen 2012; Anka et al., 2013). Natural oil seepage slicks have been recognised at the sea surface at the regional scale in the Lower Congo Basin from observation using SAR images (Williams and Lawrence 2002; Dhont et al., 2013; Deffontaines et al., 2014) confirming that this area is a major active thermogenic fluid releasing domain and an excellent natural work site for studying regional fluid release dynamics and associated expelled hydrocarbon volumes. The aim of this paper is to combine short-term and long-term SAR monitoring to compare the pace of thermogenic seeps distributed along the Lower Congo Basin and to quantify the volumes of expelled hydrocarbons. The study comprises an extensive archive comprising 82 SAR scenes acquired between 1994 and 2012 used to detect and map seepage slicks and to determine the location of recurrent oil seep sites. From these long-term observations, we assessed the activity of individual active seeps in order to quantify expelled oil volumes. To address the question of the seep lifespan, we set up a short-term monitoring with the acquisition of 22 SAR scenes every 12 hours during 10 days over a specific area recognised as a major active seeping province. The tasking of overlapping SAR scenes over an active seeping area with a half day revisiting period constitutes a first order approach to study the recurrence and synchronicity of fluid discharges at a regional scale. For the first time, this study allows us to compare simultaneously the activity of several active oil seeping sites, benefiting from a minimum of two SAR scenes each day. The comparison between the local wind fields derived from SAR data and the occurrence of oil slicks

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also helps to improve our knowledge of the relationships between slick detection and weather conditions. In addition, the assessment of short-term seepage pace is decisive to categorise the active, dormant and passive seep sites for research teams and exploratory cruises.

2. The Lower Congo Basin

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2.1. Geological context

The Lower Congo Basin is located in offshore Angola, in the southern part of the Congo Delta (Fig. 1A). The study area is located on a passive margin resulting from the opening of the South Atlantic domain initiated during the Early Cretaceous (Duval et al., 1991; Cainelli and Mohriak 1999; Browfield and Charpentier, 2006). During the Neocomian to Barremian, the development of syn-rift lakes gave rise to the prolific mudstone source rocks of the Bucomazi formation (Burwood 1999). During the late Aptian, the post-rift shallow environments allowed the development of thick hypersaline deposits overlaying a sag basin under conditions of regional thermal subsidence (Marton et al., 2004; Brun and Fort, 2004; Fort et al., 2004; Beglinger et al., 2012). During the Albian, sedimentation was characterised by carbonate ramp development (Da Costa et al., 2000; Anka et al., 2009; Beglinger et al., 2012). As thermal subsidence and progressive basin deepening continued, post-rift salt tectonics triggered by high sedimentation rates caused strong deformation of the overburden. Subsequent evolution of the margin was controlled by salt tectonics. The study area is sub-divided into three main provinces (Fig. 1B): an internal extensional area where the salt layer has been extremely thinned, a mid-slope transitional province where post-salt sedimentary series have been translated towards the basin, and a compressive domain consisting of a wide area dominated by severe deformation associated with salt diapirism, followed by a basinward compressive toe with extensive salt napes (Brice et al., 1982; Duval et al., 1991; Schoellkopf and Patterson, 2000; Marton et al., 2004; Brun and Fort, 2004; Anka et al., 2013). The seaward part of the basin deformed by salt tectonics is delimited from the abyssal plain by a distal compressive salt front (Anka et al., 2009). Intra-saliferous basins are infilled with turbidites and display an alternation of mudstone, claystone, shale and sandstone.

2.2. Seafloor manifestation of fluid vents

Geophysical investigations in the Lower Congo Basin report fluid escape features occurring as focused depressions on the seafloor and sub-seafloor fluid migration chimneys (Gay et al., 2003; Gay et al., 2007; Sahling et al., 2008; Anka et al., 2013). A large variety of pockmarks, defined as seafloor expressions of focused fluid flows (King and McLean., 1970), have been described with diameters varying from 30 m (Gay et al., 2007) up to 3150 m (Charlou et al., 2004; Ondreas et al., 2005; Gay et al., 2006c). The Aptian salt withdrawal and associated gravitational deformation greatly influenced the location of thermogenic seeps by creating efficient migration pathways along salt flanks as demonstrated by several publications (e.g. Gay et al., 2007; Andresen et al., 2010; Andresen and Huuse, 2011; Andresen, 2012). Wenau et al. (2014a, b) described active escape features on the seafloor expelling gas bubble trains imaged from side scan sonar at the salt-induced deformation front. From their fauna inventory, Jones et al. (2014) showed abundant seafloor topographic reliefs, consisting of biodegraded oil which forms asphalt mounds with diameters ranging from 0.5 to 50 m in the compressive salt province.

2.3. Oceanographic context

Understanding water mass movements is necessary for seepage studies because the strength of underwater currents steers the trajectory of oil droplets in the water column, whilst sea surface currents and wind determine the shape and lifespan of the oil slicks. The Fig. 1A displays a summary of oceanographic transport along the Angolan coast. Surface water circulation is prominently dominated by the southward Angolan Current (AC) (Moroshkin et al., 1970, Hardman-Mountfort et al., 2003), which results from the combination of different surface and subsurface currents, including the South Equatorial Under-Current (SEUC), Equatorial Under Current (EUC), the steady state surface Angolan Dome (AD) (Doi et al., 2007), and the

subsurface Angolan Gyre (AG), which are respectively surface and subsurface permanent gyres whose depth influence reach 500 m (Peterson and Stramma, 1991; Schneider et al., 1996; Holmes et al., 1997; Stramma and England, 1999). The AC is bordered to the south by a frontal convergence area with the Benguela current (Shannon, 2001), namely the Angolan-Benguela Front Zone (ABFZ) (Lass et al., 2000; Boyer et al., 2000), roughly located along the Angolan-Namibian frontier. However, the coastal and surfacing component of the Benguela Coastal Current (BCC) bypasses the ABFZ, resulting in a current flowing towards the north along the Angolan coasts (Hopkins et al., 2013). The Congo River outflow also inflects water circulation in the northern area. Seasonal variations are characterised by two warm periods from February to April and from September to November (Hardman-Mountford et al., 2003), alternating with two cold periods characterised by strong upwellings. Intermediate water circulations are controlled by the extension of the SICC to the south (Southern Intermediate Countercurrent) for water depths comprised between 500 and 1200 m (Stramma and England, 1999; Stramma and Scott, 1999). The main orientation of deepwater currents along the Angolan coast is controlled by the North Atlantic Deep Water known to occur at 2000 m depth (NADW; Stramma and England, 1999; Arhan et al., 2003) and flowing to the south at 2 cm/s (Lynn, 1971).

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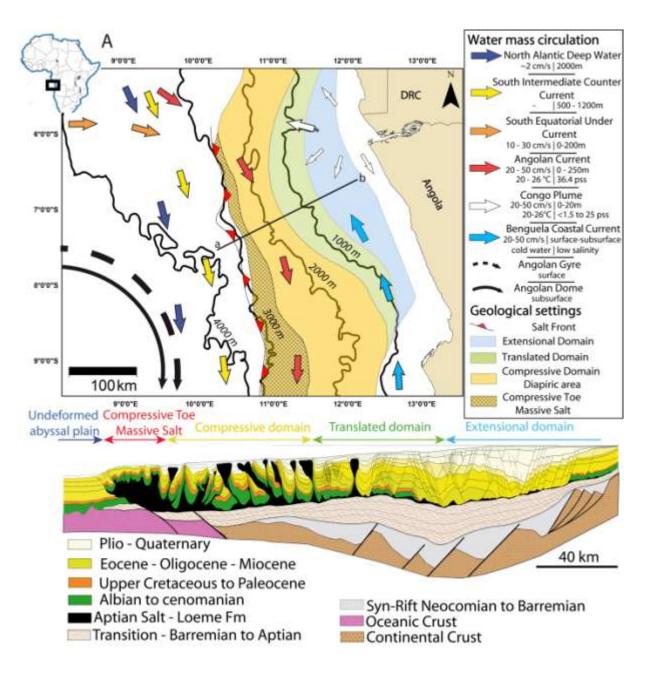


Fig. 1. A: Location map of the study area (see the African map within the black frame in the top left corner) superimposed with an overview of main oceanic currents and geological settings. The orientation of coloured arrows represents the directions of the hydrodynamic components of the basin. The characteristics of each current, including velocity, depth influence, temperature and salinity, appear in the legend when available. **B:** Geological cross-section of the Lower Congo Basin (a-b on Fig. 1A) including pre-rift, syn-rift, transitional and post-rift series (modified from Guilbot et al., 2012). See location of the section from the black line in Fig. 1A.

3. Dataset and methods

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3.1. SAR acquisition parameters for oil slick detection

Synthetic Aperture Radar (SAR) is sensitive to centimetre-scale capillary waves induced by the local wind at the sea surface (Franceschetti et al., 2002; McCandless and Jackson, 2003). Oil surfactants smooth the sea surface and significantly reduce capillary waves, which triggers a contrast in roughness between oil-covered areas and free water. A minimum wind speed is required to trigger capillary waves and therefore to identify seepage slicks. The lower wind speed limit for seepage slick detection varies between case studies and authors, from 1.5 m/s (Skrunes, 2014, Fingas and Brown, 2014), 2-3 m/s (Espedal and Wahl, 1999; Johannessen et al., 2000; Girard-Arduin et al., 2005, Brekke and Sölberg, 2005) to 3.5 m/s (Garcia-Pineda et al., 2009). In case of strong winds, natural oil slicks are also affected by capillary waves, which prevents the detection of oil at the sea surface. Seepage slicks from natural discharges are considered to be detectable up to 6-7 m/s (Bern et al., 1992; Espedal and Wahl, 1999). Some authors consider that thicker slicks can be identified for higher wind speeds (Alpers and Espedal, 2004; Brekke and Sölberg, 2005; Körber et al., 2014). In addition to the sea surface currents, it is assumed that 3% of the wind speed accounts for the oil slick drift velocity (Bern, 1992; Espedal and Wahl, 1999; Kim et al., 2014). Previous studies demonstrated that short wavelengths of the X- (2.5 to 3.75 cm), C- (3.75 to 7.5 cm) and Ku-bands (1.6 to 2.5 cm) are more efficient in detecting oil slicks at the sea surface (Gade et al., 1998; Espedal and Wahl, 1999; Espedal and Johannessen, 2000; Johannessen et al., 2000; Alpers and Espedal, 2004; Brekke and Sölberg, 2005; Girard-Arduin et al., 2005; Skrunes, 2014). According to Bragg's reflection theory, the backscattered signal decreases with increasing incidence angles between 20° to 70° (Alpers and Espedal, 2004). Ideal incidence angles for slick detection range from 20° to 45° (Girard-Arduin et al., 2005; Brekke and Sölberg, 2005; Garcia-Pineda et al., 2010). Vertical SAR polarisation (VV) is the most efficient mode to detect oil slicks at the sea surface (Brekke and Solberg, 2005).

3.2. SAR dataset

3.2.1. Long-term acquisitions

Satellite SAR sensors are deployed on polar orbiting platforms, which enables worldwide observations with a rapid repeat rate. The SAR operational mode allows day and night imagery, regardless of cloud cover, which is one of the main advantages compared to passive optical imagery. In order to find oil seep sites, we gathered a collection of 82 SAR images from different satellite sensors (Fig. 2A). Archive data acquired from 1994 to 2012 were used to detect recurrent oil slicks and afterward to identify the sea surface location of natural oil seepage.

The dataset provided by the European Space Agency (ESA) is composed of ASAR data (Advanced Synthetic Aperture Radar) acquired in Image Mode Precision (IMP) by the European Remote Sensing satellites (ERS 1 and 2) between 1994 and 2001 and Wide Swath Mode (WSM) Envisat scenes acquired between 2002 and 2012, both operating in C-Band (Table 1). We complemented this dataset with RADARSAT, TerraSAR-X and Cosmo-SkyMed data. The map of SAR coverage density varies between 40 to 70 scenes (Fig. 2A). SAR scenes were converted from N1 raw format to 8-bit geoTiff format, displayed in greyscale for visual interpretation with GIS software (ArcInfo, ESRI). Scenes were kept speckled to preserve whole details. We applied linear stretching to counterbalance signal loss with increasing incidence angle from the orbital path.

Table 1
 SAR data used for oil slick mapping during both long-term and short-term monitoring.

			Amount of SAR data		
Satellite/ Sensor	Band/ Wavelength	Pixel Spacing (m)	Long-term monitoring (1994 - 2012)	Short-term monitoring (12 - 22/10/2012)	
Envisat /WSM	C (5.8 cm)	75	46	0	
ERS/IMP	C (5.8 cm)	12.5	17	0	
Cosmo-SkyMed	X (3.1 cm)	3-15	13	15	
TerraSAR – X	X (3.1 cm)	20	2	4	

RADARSAT-2	C (5.3 cm)	25-50	4	3

3.2.2. **Short-term monitoring**

In order to investigate the seepage recurrence, synchronicity and lifespan on sea surface, we performed a short-term monitoring based on the tasking of SAR images during a restricted period of 10 days. We focused this monitoring on the area recognised for the largest amount of expelled oil, located 135 km from the Angolan coast to the southwest of the mouth of the Congo River (Fig. 2B). Tasking of 22 satellite images during 10 days provided at least one image every 12 hours (Table 2). The dataset consists of a combination of X-band (TerraSAR and Cosmo-SkyMed) and C-band (RADARSAT) data. The radar density coverage constructed from the stack of multi-temporal images varies spatially, including a central area of maximum coverage density imaged by a series of at least 20 images out of the 22 tasked.

Table 2
 Listing and characteristics of SAR scenes acquired during short-term monitoring

Scene	Date	Time	Sensor	Scene	Date	Time	Sensor
number	(dd/mm/yyy	(hh:mm)		number	(dd/mm/yyyy)	(hh:mm)	
	y)						
1	12/10/2012	05:40	Cosmo-SkyMed	12	18/10/2012	05:22	Cosmo-SkyMed
2	13/10/2012	04:58	RADARSAT - 2	13	18/10/2012	16:47	Cosmo-SkyMed
3	13/10/2012	16:53	Cosmo-SkyMed	14	18/10/2012	17:37	TerraSAR - X
4	13/10/2012	17:28	TerraSAR - X	15	19/10/2012	05:40	Cosmo-SkyMed
5	15/10/2012	05:06	TerraSAR - X	16	19/10/2012	16:47	Cosmo-SkyMed
6	15/10/2012	05:40	Cosmo-SkyMed	17	20/10/2012	04:54	RADARSAT - 2
7	15/10/2012	16:41	Cosmo-SkyMed	18	20/10/2012	16:35	Cosmo-SkyMed
8	16/10/2012	05:34	Cosmo-SkyMed	19	21/10/2012	04:58	TerraSAR - X
9	16/10/2012	17:42	RADARSAT - 2	20	21/10/2012	05:28	Cosmo-SkyMed
10	17/10/2012	05:34	Cosmo-SkyMed	21	21/10/2012	16:53	Cosmo-SkyMed
11	17/10/2012	06:53	Cosmo-SkyMed	22	22/10/2012	06:53	Cosmo-SkyMed

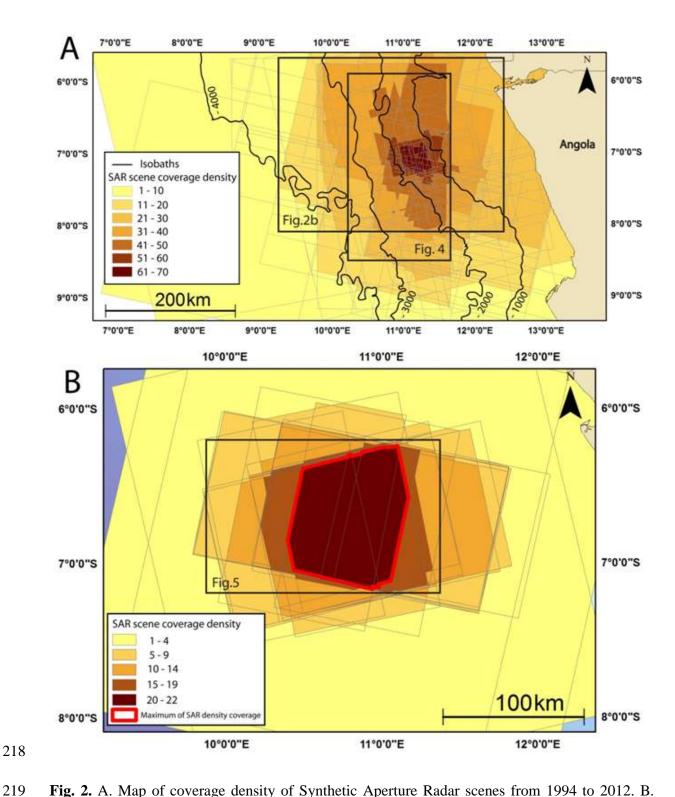


Fig. 2. A. Map of coverage density of Synthetic Aperture Radar scenes from 1994 to 2012. B. Coverage density of Synthetic Aperture Radar tasked during the short-term monitoring. The maximum coverage density of SAR scenes tasked during short-term monitoring occurs within the red polygon outline and covers 6736 km² with a water depth ranging from 1000 to 2750 metres.

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3.2.3. Wind and current assessments

In order to observe the wind effect on slick detection, we compared the regional wind field with the number of slicks detected and their size with the local wind field extracted at the location of the oil slicks. Real-time wind field measurements are not available over the study area, hence we used the regional wind field provided by the European Centre for Medium-Range Weather Forecasts (ECMWF). The wind field is calculated every 6 hours with a spatial resolution of the grid of 1/8° (~14 km). The forecasted wind field has been averaged over 24 hours for better comparison with the number of slicks detected in each image. In order to observe small-scale variations of the wind field, we assessed the local wind from SAR scene derivation with the scatterometer technique. Spatial resolution is 1150 m with a maximum speed uncertainty of 2 m/s and 15° in direction (for details see Johannessen et al., 2000 and Zecchetto, 2015). The wind dataset is composed of discrete data including forecasts and SAR-derived wind fields which represents a certain limit for the understanding of the wind effect on slick dissipation between two SAR acquisitions. The current velocity is measured from ADCP (Acoustic Doppler Current Velocity) tools that use particle displacement to quantify current speed (Rowe and Yound 1979; Gordon and RDI, 1996). The data were collected during 6 months of mooring between January 2009 and July 2009 at a depth of 10 m below the sea surface. The current velocity was not available during short-term monitoring, but the integration of predating mooring data represents a good proxy to approach the

3.3. Identification of natural leakages at the sea surface

- Sea surface smoothing may be induced by a wide collection of look-alike seep features. Table 3 sums up look-alikes found during SAR scene interpretation.
- 246 **Table 3:**

current velocity distribution.

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Overview of look-alike features detected on SAR scenes and dismissal features

Look-alike type	Origin	SAR features	Dismissal method	References
Low Wind Areas	Insufficient wind speed to trigger capillary waves	Extensive dark patches	Masking of the affected region	Espedal and Wahl (1999)
Bioturbations	Surface covered area smoothes the sea surface	Threadlike extensive and scattered films	Masking of the affected region or SAR scene dismissal	Gade et al. (1998) Espedal and Johannessen (2000)
Boat spill	Ship hold cleaning	Straight slick caused by moving source, shining point at boat location	Slick mapping with appropriate "spill" rank attribution	Espedal and Wahl (1999)
Platform/ Pipeline spills	Accidental/operational production pollution	Petroleum infrastructure data integration (IHS) and shining platform	Slick mapping with appropriate "spill" rank attribution	Trivero and Biamino (2010)
Rain cells	Local opposite winds at the edge of rain cells cause roughness wipeouts	High roughness values in the internal cell area compared to free water and rounded features	Masking of the affected region	Espedal and Johannessen (2000)
Solitons or Internal waves	Decrease backscattered energy associated with wave-related slopes	Multiple parallel trends interspersed by greater backscattered values	Visual dismissal	Jackson et al. (2013)

Petroleum infrastructure database integration (IHS) helps to discard leakage originating from pipes and/or petroleum platforms. In addition, boats or platforms appear as shiny diffracting points on SAR data, which make a distinctive feature compared to natural slicks. Spills from anthropogenic activities represent about 15% of the total surfacing oil in this study. Rain cells and solitons were occasional, in contrast to bioturbations which forced us to dismiss about 20% of SAR scene collection. Regions of SAR scenes containing low wind areas, rain cells or bioturbations were removed from the data collection to preserve the SAR coverage density which is used to quantify the seepage activity. The shape of natural oil slicks is induced by the strength and orientation of the wind and surface currents (Espedal and Wahl, 1999). Wind and surface currents induce short-term variations of sea surface motion. The sinuosity of black patterns is a first-order criterion to detect oil content slicks (Fig. 3 A - D). Sharp radiometric contrast between oil-covered areas and the surrounding free water correspond to the sideward detectable edge of the

slicks and allowed us to identify the extent of individual seepage slicks (see examples in Fig. 3 A and B). Computer-based automatic picking (Brekke and Solberg, 2005; Garcia-Pineda et al., 2009) certainly provides a quick way to delineate dark features but is still not able to distinguish natural seepage from anthropogenic pollutions or look-alikes (Suresh et al., 2013). Garcia-Pineda et al. (2009) showed that the classification using neuronal networks between dark areas and free water is equivalent to manual picking. In our case, visual interpretation is more powerful, as the input of the photo-interpreter allows a confidence rank attribution which prevents the overinterpretation of dark features. Therefore, we based our interpretation on manual picking instead of computer-based automatic detection. We compiled on a map the outlines of each highconfidence slick identified on each SAR image (Fig. 3E). Once mapped, diverging black streaks evidence active seeping areas of thermogenic fluids (Kornacki et al., 1994; MacDonald et al., 1996; de Beukelear et al., 2003; Zatyagalova et al., 2007; Garcia-Pineda et al., 2010; Garcia-Pineda et al., 2014; Körber et al., 2014). The centre of each diverging structure corresponds to the probable impact area of the fresher oil droplets at the sea surface. The Oil Slick Origin (OSO, Garcia-Pineda et al., 2010; Körber et al, 2014) is visually drawn for each seepage slick from the location of the proximal detectable edge of the slick to the seeping area (see white dots in Fig. 3E). Diverging slicks allowed us to sort seepage slicks and associated OSOs between recurrent seep sites. We associated nearby and recurrent double-pattern slicks as being associated with two different seep sites (Fig. 3). From the location of each OSO, we computed the kernel density with the ArcGIS toolbox in order to better display areas of active seepage. We also calculated the geometric mean centres (Garcia-Pineda et al., 2010) for each diverging structure, corresponding to mean values of each OSO location. The spatial dispersion of the OSOs generated by a single seep and detected in multiple SAR images is related to the horizontal deflection of the oil plume through the water column. We drew the minimum size circles encompassing each single OSO expelled from individual seep sites to evaluate the offset range values generated by underwater deflection (Garcia-Pineda et al., 2010)

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by using the minimum bounding geometry tool of the ArcGIS toolbox (Fig. 3E). The radius serves as a proxy to quantify the horizontal deflection of oil droplets in the water column when rising towards the sea surface. We also computed the ratio between the number of oil slicks detected and the SAR scene coverage density, namely the occurrence rate, which gives an indication of the frequency and release activity of each seep. Compared to SAR data coverage, seep sites with high occurrence percentage frequently expel oil slicks, conversely to the low occurrence percentage that corresponds to weaker seeping sites with few slicks detected during the monitoring interval. We assessed the average oil volume at the sea surface from the method developed by MacDonald et al. (2015), who gridded their area and considered the total surfacing volume divided by the SAR coverage density for each cell.

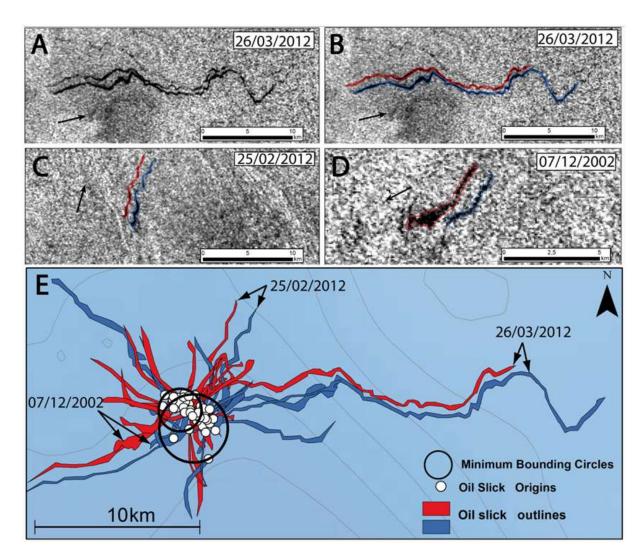


Fig. 3. A. Extract from SAR WSM scene dated 26/03/2012. Grey speckled background represents backscattered energy from sea surface free water. Black streaks depict the location of seepage slicks. B. Same SAR image extract complemented with manual picking. C and D. ASAR WSM SAR sub-scenes respectively dated 25/02/2012 and 07/12/2002 superimposed with manual picking of oil slicks. The three different SAR extracts overlap at a common spatial location. Black arrows show the flow direction of surface currents during SAR scene acquisition. E. Compilation map of oil slick outlines mapped from the entire dataset, including slicks displayed in A., B., C. and D. at two major seep sites. All scenes are provided by the European Spatial Agency (ESA).

4. Oil seep analysis and results

4.1. Long-term observations

We mapped a total of 1600 oil slicks interpreted from the analysis of the 82 SAR data over 18 years. The integration of the IHS international exploration & production database and the recognition of diffracting points associated with the boat or platform location allowed us to remove about 200 slicks associated with anthropogenic activities. The remaining 1400 seepage slicks are displayed in Fig 4A. Among them, 1240, i.e. 90%, of the slicks are recurrent, forming diverging patterns on the compilation map. The length of the slicks varies from a few hundred metres to 45 km. We recognised 102 active seep sites whose geometric mean centres are displayed in coloured dots in Fig. 4B. The water depth of the seeping province ranges roughly from 1100 to 2600 m at a distance of 130 to 200 km from the Angolan coastline. Active seep sites are distributed all along the southern part of the Congo Basin over an area of 21,000 km². The seeping province is well delimited from the basinward edge of the translated province and mainly included in the compressive province associated with salt diapirs (Figs. 1 and 4). Occurrence rates range between 5% for the weakest seeps and 80% for the most active (Fig. 4B), suggesting that the flux of fluid differs between seep sites.

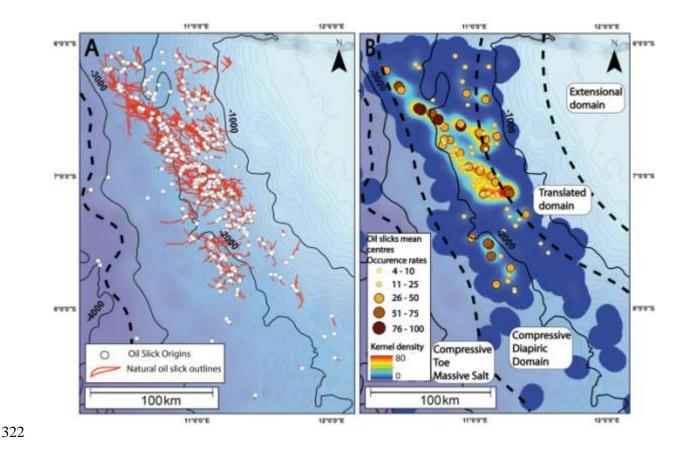


Fig. 4. A. Compilation map of oil slick outlines in red polygons superimposed with Oil Slick Origins in white dots (see details in text). B. Kernel density computed from the density of seepage slicks. Coloured dots show the location of geometric mean centres of OSOs. The dashed black lines indicate the boundaries of each province related to the salt tectonics presented in Fig. 1B.

4.2.Short-term monitoring

4.2.1. General observations

We identified a total of 231 high-confidence oil seepage slicks from the interpretation of the 22 tasked scenes. The locations of recurrent seep sites are consistent between the two monitorings. Recurrent slicks during short-term monitoring are associated with 53 seep sites identified from long-term monitoring (coloured dots in Fig. 5 - see location in Fig. 2). A total of 102 oil seep sites were mapped from long-term monitoring, suggesting that about 50% of seeping sites were passive during short-term monitoring (black stars in Fig. 5).

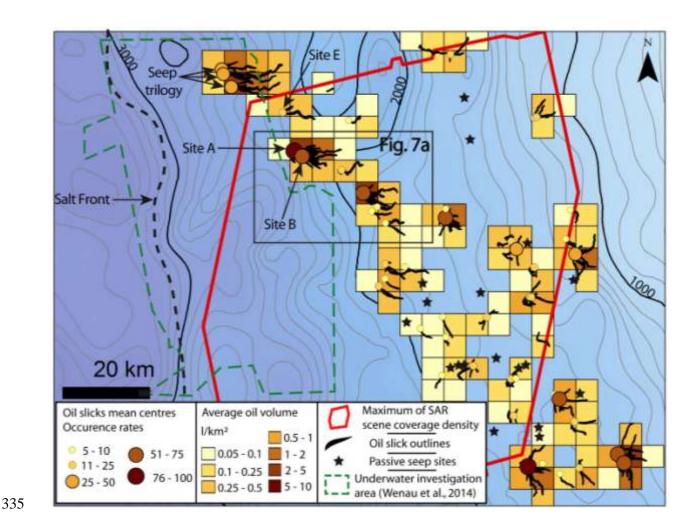


Fig. 5. Compilation map of oil slick outlines (black polylines) mapped from SAR scenes tasked during short-term monitoring and focused on the maximum coverage density area (see location in Fig. 2A). Coloured dots represent geometric mean centres of OSOs expelled from individual seep sites. Size and colours correspond to the occurrence percentages. Inactive seepage site locations during short-term monitoring are displayed as black stars. The salt tectonic boundary is displayed as a black dashed line. Background continuous lines are 100 m isobaths.

The area of highest radar coverage, defined as the area imaged by a minimum of 20 images out of the 22 tasked scenes, encompasses the 51 seeping sites observed during long-term monitoring (inside the red frame in Fig. 5).

In this area, 130 slicks with lengths varying from 1000 to 16,000 metres are associated with 33 geometric mean centres or active sites (Fig. 5). The spatial spreading of OSOs retrieved from the computation of minimum bounding circles for each seep sites provides the estimation of the deflection offset in the water column. The spatial dispersion of surfacing slick areas varies between 80 and 1020 m, which remains low compared to the water depth of the seeping province (from 1100 to 2600 m). As shown by the Fig. 1, the regional hydrodynamics involves multiple oceanic components with various speed and depth influences. The horizontal deflection of oil droplets in the water column is directly influenced by this hydrodynamic complexity. The linear law between the water depth and the horizontal deflection established in the Gulf of Mexico (Garcia-Pineda et al., 2010) does not apply for the Lower Congo Basin. As a matter of fact, the greatest OSO dispersion values account for shallower areas in this study where the AC meets the BCC's influence, probably resulting in greater current variability, whereas the deepest seep areas are associated with restricted OSO spatial spreading.

4.2.2. Correlation with the wind field

The study of seepage recurrence requires good knowledge of the wind speed to ensure the consistency of the detection. As a first approach, we used the wind speed field forecasted at the regional scale (ECMWF) to confirm that weather conditions were suitable all along the 10-day acquisition period (Fig.6). The reported values of suitable wind speed for the detection of oil slicks (Espedal and Wahl, 1999; Alpers and Espedal, 2004; Girard-Arduin et al., 2005; Brekke and Sölberg, 2005; Garcia-Pineda et al., 2009; Fingas and Brown, 2014) confirm that the weather conditions were suitable during short-term monitoring (from 3.7 to 6.5 m/s (Fig. 6) for the recognition of oil slicks on SAR data.

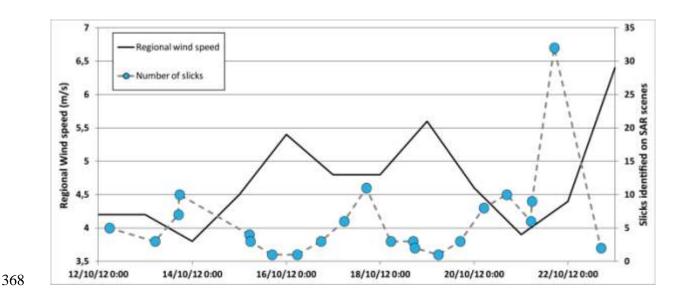


Fig. 6. Superimposed graphs of modelled regional wind speed (ECMWF) displayed in black line and the number of slicks in the high-density SAR coverage area.

The two curves displaying the wind speed and the number of slicks show an anti-correlation, where higher wind speeds correspond to fewer detected slicks and vice versa. The curve displaying the number of detected slicks through time reveals three peaks characterised by an increase of the number of detected slicks (scenes 3 to 4, scenes 9 to 11, scenes 17 to 21 - see Table 2). These periods are interspersed by stages of low detection of slicks on SAR scenes especially at the following scenes: 5 to 8, and 12 to 16. These stages correspond to periods of regional wind speed intensification.

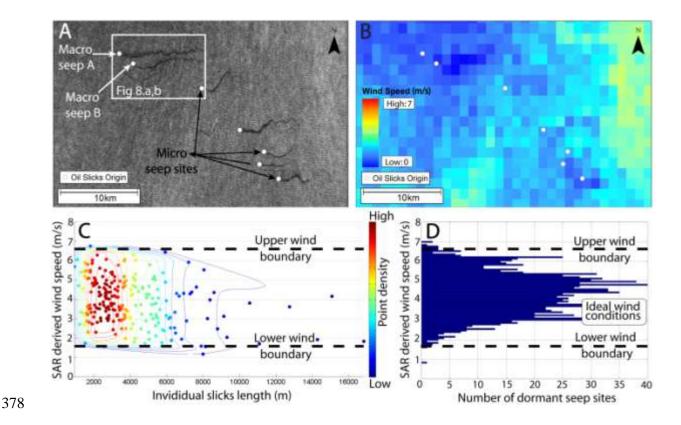


Fig. 7. A. Extract from the SAR scene dated 21 October 2012, 16:53 (scene 21). Black streaks show the location of oil slicks at the sea surface. B. Local wind values are retrieved from SAR scene derivation at the locations of OSOs displayed as white dots. C. Plot of the length of individual slicks and the local wind speed value at the location of OSOs. The colour of the dots ranging from red to blue respectively represents high to low point density. D. Histogram of the distribution of local wind values at the location of recurrent oil seeping areas revealed from long-term monitoring and measured during dormant stages in short-term monitoring. The vertical black dashed lines correspond to the probable ideal boundaries for seepage slick detection.

Even though the modelled wind provides a regional assessment of weather conditions, the applications related to local wind quantification remain questionable. Therefore, in order to quantify more accurately the ideal weather conditions for slicks in the study area, we integrated the local wind measurement extracted from SAR scenes from the scatterometer technique at the location of seepage slicks (Fig. 7A, B). Local wind values vary between 1.5 and 6.5 m/s at the

slick locations (Fig. 7C). This observed detectability agrees with lower and upper values reported in the literature.

The information related to the population density of slick lengths shows that most of the slicks (137 out of 231) were detected with lengths ranging from 1000 to 4000 metres. The population density progressively decreases with the increase of the slick length. Slicks longer than 8000 m were occasional and exclusively associated with local winds between 2 to 4 m/s, meaning that no elongated slicks were detected with wind speeds above 4 m/s during the 10 days of monitoring (Fig. 7C).

We then plotted the number of dormant seep sites on the SAR-derived winds (Fig. 7D). Considering that ideal wind values range from 1.5 to 6.5 m/s in the study area, dormant seep phases clearly occur during periods of ideal wind conditions. Only a restricted portion of dormant seep sites (~5%) occur outside the ideal detection boundary conditions (Fig. 7D). This shows that the disappearance of slicks at recurrent seep sites is associated with short-term intermittence.

4.2.3. *Macro-seepage sites*

We defined macro-seepage sites as prolific leaking spots with occurrence rates above 50%. Within the area of SAR maximum coverage density, the acquisition of SAR data simultaneously at the nearby major seep sites allowed us to study the seepage pace for major seep sites during short-term monitoring. Seep sites A and B are distinct seepage sites, as the identification of recurrent double-pattern slicks allowed us to distinguish slicks expelled from one site to the other (Fig. 7A). The two sites are only 2250 m apart; this proximity constitutes a first-order benefit to compare the pace of two different seeps, independently of weather conditions (Fig. 8 A, B). Most of the time, sites A and B were both active, as seepage slicks were identified at both sites for 13 scenes out of 20 (Fig. 8 C).

We also identified asynchronous active phases from the recognition of isolated seepage slicks on 4 images at macro-seep A, while macro-seep B was dormant (pinpointed with black arrows in Fig.

8). Based on previous works showing that oil slicks are best detected for wind speeds less than 6-7 m/s (Bern et al., 1992; Espedal and Wahl, 1999), the maximum wind value of 6.5 m/s at these specific sites during short-term monitoring cannot be considered as the factor controlling the absence of seepage (Fig. 8D). In addition, the fact that oil slicks are observed for higher wind speeds during the monitoring period implies that weather conditions were suitable for slick detection. To summarise, all the conditions were satisfied during the three common and asynchronous cessation events for an optimal detection of oil slicks.

Furthermore, based on the fact that the disappearance of the seepage slicks is observed between two consecutive SAR acquisitions implies that oil dispersion is effective in less than 12 hours. The reappearance of oil slicks on the following image also indicates that cessation of oil release can occur under restricted time scales.

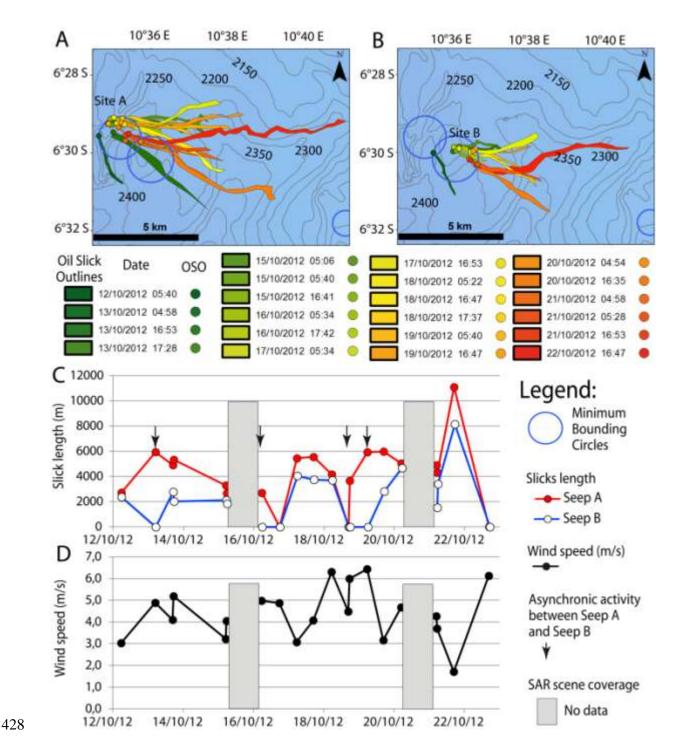


Fig. 8. A, B. Compilation of oil slick footprints and associated OSOs revealed from short-term monitoring for macro-seep sites A and B. Colour scale refers to subsequent acquisition dates and is applied for both seepage slicks and OSOs. Blue circles are the deflection offset computed from individual OSO locations. C. Red and blue graphs display the length of individual seepage slicks for seep site A and B respectively. D. Graph of the local wind values derived from SAR scenes.

Non-imaged areas were masked with grey squares. Black arrows pinpoint SAR scenes where seepage slicks were only visible from one of the two sites A and B.

4.2.4. Transience and persistence at micro-seepage sites

A micro-seepage is defined here as a seeping spot with an occurrence rate below 50%. A large number of micro-seepage sites (~1/3) are only associated with a single active event that occurred during the acquisition of SAR scene No. 21 (see yellow dots in Fig. 5). Isolated slick detection occurred during a period of regional increase in the amount of detected oil, while roughly 60% of seeping sites were active. Wind values varied between 1.2 and 6.1 m/s at the OSO location, which is slightly below observed average weather conditions (Fig. 7). Apart from lower occurrence rates, oil slicks expelled from micro seepage sites are revealed by shorter and narrower slicks, showing that the amount of oil released is weaker (Fig. 7A). Micro-seep E identified from 3 recurrent and diverging slicks appears as a new seepage feature on the compilation map that we confidently ranked as natural slicks (Fig. 5). Prior to short-term monitoring, the area was passive, as the SAR data coverage reaches 40 overlapping scenes between 1994 and 2012 (Fig. 2A), but none revealed the presence of seepage slicks at this location. The slicks are characterised by typical curvilinear shapes that have similar trends to the neighbouring slicks on the images, which comforts our presumptions on the natural origin of expelled slicks. This example corresponds to the minimum period of seepage activity observed during the short-term monitoring, revealing that oil release can be ephemeral and temporally constrained to a few days.

5. Discussion

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5.1. Wind requirements for seepage slick detection

The weather requirements for slick detection have been reported by different authors to be 3 to 10 m/s (Espedal and Wahl, 1999; Alpers and Espedal, 2004; Girard-Arduin et al., 2005; Brekke and

Sölberg, 2005), 3.5 to 7 m/s (Garcia-Pineda et al., 2009) and 1.5 to 10 m/s (Fingas and Brown, 2014). During the short-term monitoring, regional wind values ranged among the most restrictive conditions as reported by Garcia-Pineda et al. (2009). The results from seepage studies in the Lower Congo Basin show that weather requirements are constrained between 1.5 and 6.5 m/s (Fig. 7C). Such differences are presumably associated with seepage slick characteristics such as the slick thickness or the oil type (API gravity) that largely varies between petroleum provinces. The wind is also known to affect seepage slicks by separating heavy and light components of oil slicks and spreading oil at the sea surface (Alpers and Espedal, 2004; Marmorino et al., 2008). SAR scene No. 21 records the greatest number of seepages, where two thirds of the imaged sites were active (Figs. 6 and 8). The SAR acquisition was predated by low weather conditions, with regional wind speed around 4 m/s during the 24 preceding hours (Fig. 6). This suggests that ideal weather requirements for seepage slick detection might call for regional low wind conditions not only during, but also predating the SAR acquisition.

5.2. Seepage slick residence time

Seepage slick residence time was defined from the estimated age of the oil droplet at the distal edge of the slicks (De Beukelear et al., 2003). The restricted length of oil slicks, in addition to brief variations of slick orientation prevent seepage shapes from being fitted between subsequent scenes. This suggests that oil slick dissipation is effective between the acquisitions of two subsequent SAR scenes, typically separated by 12 h.

The residence time is approached from the individual slick length divided by the sea surface drift velocity (Macdonald et al., 2015). The velocity of the surface slick displacement is inferred to be induced by 100% of the surface currents in addition to 3% of the wind speed (Kim et al., 2014). The current measurements were not available during short-term monitoring and large scale oceanographic models are unsuitable to characterise local dynamics of the sea surface. We

integrated mooring data to characterise the distribution of the surface/subsurface current velocity (Fig. 9A).

The mooring data were recorded from January 2009 to July 2009, i.e. three years before short-term monitoring. The integration of ADCP data allowed us to compute a statistical study of the residence time of oil slicks at the sea surface during short-term monitoring using the near surface velocity of the current (10 m below the sea surface) and the local wind from derived SAR scenes. Surface velocity ranged between 0 and 81 cm/s during the 6 months of mooring. Among this range, 85% of measured velocity lies between 5 and 30 cm/s with a median velocity value of 17 cm/s. We made three estimations of residence time considering a lower bound for a conservative current velocity of 5 cm/s, an upper bound for an overstating current velocity of 30 cm/s and a best estimate for the median current velocity of 17 cm/s (Fig. 9B).

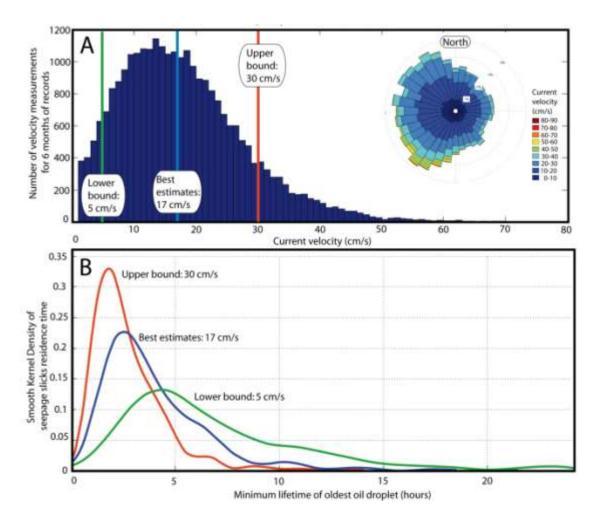


Fig. 9. A: Histogram of the distribution of current speed recorded from 6 months of mooring measurements 10 m below the sea surface. We considered an upper and a lower bound from the overall distribution of the current velocity. The wind rose to the right shows the distribution of current orientation and speed. B: Smooth kernel density of the distribution of the minimum residence time of the oldest oil droplet of individual slicks mapped during short-term monitoring. The residence time computation considers the local wind value at the location of OSOs, surface current velocity and slick length. The estimation of theoretical residence times shows that their distribution gathers around the median value for the upper, median and lower bound. The distribution progressively decreases from this median value. The median residence time value for the upper bound at 30 cm/s is 2h15min, 3h15min for the median value of 17 cm/s and 5 h 30min for the lower bound at 5 cm/s. The slick residence time for median estimates is mostly distributed between 1 and 7 hours. The natural seepage residence time was inferred to be comprised between 8 and 24 h by Mitchell et al. (2000) and MacDonald et al. (2015) for the Gulf of Mexico. Seepage residence time in this study is lower than previously suggested, as best estimates suggest that median residence time is 3h15min. The residence times from this study are rather in agreement with Marmorino et al. (2008), who suggested that they concentrate below 5 hours. Restricted residence time has great implication on the hydrocarbon flux and therefore on quantifications of the regional oil output. The short lifetime of seepage slicks can be explained from the relative expelled amount that is probably lower in the Lower Congo Basin than in the Gulf of Mexico, but also by severe weathering conditions along warm Angolan coasts (Hardman-Mountford et al., 2003) that largely contribute to spill dispersion (Leifer et al., 2012). However, these observations refer to short-term monitoring, where the length of the oil slicks

varies between 1000 and 17,000 m (Fig. 7), yet seepage slicks were mapped with greater lengths

during long-term monitoring (Fig. 3A and B). Observations of slicks greater than 30 km long

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remained unusual (less than 1% of slick collection) but suggest that slick residence time can occasionally be greater than values computed from short-term monitoring. Such exceptional observations of lengthened slicks could also be explained by a substantial increase in surface displacement velocity or predating by consistent low wind conditions before SAR acquisition.

5.3. Evidence of leakage interruption

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The non-detection of seepage slicks on SAR scenes can either be explained by unsuitable wind conditions or as an evidence of a consistent decrease of the amount of expelled oil that prevents the detection of slicks by the SAR system. The comparison between the distributions of local wind values at recurrent seep sites between active and dormant stages remained concordant (Fig. 7). The wind distribution at the location of recurrent sites during dormant periods were imaged during the ideal weather conditions (Fig. 7D) which demonstrates that the non-detection of seepage slicks is associated with a short-term intermittence. It was established in the Black Sea (Greinert et al., 2006) and Gulf of Mexico (Garcia-Pineda et al., 2010) that gas seepage rates from cold seeps may be highly variable over time. Leifer et al., (2004) and Garcia-Pineda et al. (2016) also demonstrated this variability for oil seepage in the Santa Barbara basin and the Gulf of Mexico, respectively. Concordant observations performed from long-term and short-term monitoring suggesting that oil seepage is also discontinuous in the Lower Congo Basin are listed below: - The great variability of occurrence rates values ranging from 5 to 80% shows that the amount of released oil significantly differs between seeping sites, suggesting recurrent interruption of oil release with varying frequencies between seep sites. - Recurrent twin slicks were identified at the location of the two macro-seep sites A and B, which are separated by a distance of 2250 m. The identification of concordant shapes for neighbouring slicks on SAR data (Figs. 3 and 8) proves that they are expelled simultaneously from distinct

sources (MacDonald et al., 1993; Garcia-Pineda et al., 2010; Körber et al., 2014). Therefore, the identification of a single seep suggests the leakage interruption of one of the two sites instead of unsuitable meteorological conditions (Fig. 8).

- Seep site E (see location in Fig. 6) was characterised by a restricted active period of a few hours.

The observations of recurrent slicks over a restricted period can be either explained as associated with ephemeral seep sites or with a consistent increase in the amount of expelled oil.

- The seep triplets in the northernmost area (Fig. 6) lie in the underwater investigation area studied by Wenau et al. (2014). The detection of acoustic anomalies is commonly used to identify seepage bubble trains in the water column (De Beukelear et al., 2003; Körber et al., 2014). The area was considered as inactive, as the water column records did not reveal any typical acoustic anomaly at this location. This demonstrates that the three oil seeps were dormant during the few days of underwater investigations and further supports the hypothesis of intermittent seepage activity.

The integration of SAR data characterised by a swath several hundred of kilometres wide provides an interesting means to monitor several oil seeps simultaneously. Mesoscale events, such as seismic events (Zatyagalova et al., 2007) or hydrostatic variation of pressure (Boles, 2001; Glasby et al., 2001; Chanton et al., 2003; Rollet et al., 2006) were demonstrated to be associated to a variation of the amount of expelled fluids. Conversely, our results regarding the regional seepage pace indicate that individual seep sites can have their own periodicity (see examples of seep sites A and B) with miscellaneous active and inactive stages.

5.4. Insights on seepage pace from short- and long-term monitoring

The short-term pace of seepage activity appears highly variable. The maximum duration of seepage activity was observed for Seep A, where slicks were detected on six consecutive scenes showing that they remained active during at least 3 days. In addition, inactive periods of macroseep sites are restricted to daily periods (Fig. 9), revealing that the seeping mechanism may be transient at a daily scale. The recurrent lack of seepage slick detection above macro-seepage sites

suggests at least a great decrease of both the volume of expelled oil and the slick thickness preventing the slick detection from SAR scenes detectability threshold.

Among the 102 seep sites (Fig. 4), 48 minor seep sites (occurrence rate lower than 30%) were passive during the short-term monitoring acquisitions (black stars in Fig. 5), revealing that inactive stages can be longer than 10 days. Occurrence rates of macro-seep sites are consistent between short-term monitoring and long-term monitoring. This suggests that the relative macro-seep activity observed over the short-term period remain constant over decades.

In addition, non-recurrent slicks were observed from 140 seepage slicks (10%) during long-term monitoring (isolated white dots in Fig. 4A). Assuming that the acquisition frequency of SAR scene acquisitions was low between 1994 and 2012 (4 to 5 scenes each year in average), non-recurrent slicks detected during the long-term monitoring are presumably associated with ephemeral expulsions sites, as demonstrated by the short-term monitoring (seep site E). The high variability of the amount of expelled oil suggests that the evaluation of the seepage activity during scientific exploration campaigns should be based on a consistent period of underwater records.

5.5. Oil output quantification

From oceanographic investigations and SAR observations, the minimum detectable thickness for oil film can be estimated at 0.1 µm (Allen, 1970; Macdonald et al., 1993; Mitchell, 2000; Macdonald et al., 2002; Garcia-Pineda et al., 2009). From the Bonn agreements (2009), which established a relationship between the visual aspect of oil slicks and their thickness in the offshore domain, Garcia-Pineda et al. (2013) considered that the thickness of an oil slick is 1 µm for 5% of the fresher slick surface and 0.1 µm for the remaining 95% of the surface. Based on this assumption, a slick of 1 km² may therefore contain 145 l of oil. Considering this thickness, the sum of oil-covered surface within each grid cells gives a total average amount of 1625 litres at the sea surface (Fig. 5). Considering a median residence time at the sea surface computed with the subsurface current (Fig. 9) and the local wind velocity (Fig. 7), we performed a volumetric

computation of regional oil output for the 102 natural seeps along the Angolan coasts (Table 4). The median estimate, considered as the best estimate and computed using an average surface current velocity (Fig 9), indicates that 500 litres are expelled each hour, which represents an annual amount of 4380 m³ of oil naturally expelled in the Lower Congo Basin.

Table 4

Estimation of the regional output of oil from the 102 natural seeps along the Lower Congo Basin, assessed from average surface oil and SAR data coverage.

Average surface oil volume (l)	Lower Bound	Median estimate	Upper bound
Considered subsurface current velocity (cm/s)	5	17	30
Median residence time	5h30min	3h15min	2h15min
Average output (l/h)	295.45	500	722.22
Annual amount (l/yr)	$2.58*10^6$	4.38*10 ⁶	6.32*10 ⁶
Annual amount (bbl/yr)	1.62*10 ⁴	2.75*10 ⁴	3.97*10 ⁴

To give a comparison, Macdonald et al. (2015) assessed the annual output across the entire Gulf of Mexico as between 25,300 m³ and 94,800 m³ per year. Hornafius et al. (1999) estimated oil seepage along the Santa Barbara basins around 5840 m³. For the southwestern Caspian Sea, Zatyagalova et al. (2007) evaluated a minimum natural seepage rate of 1700 m³. Estimations performed by Körber et al. (2014) suggest that about 360 m³ are expelled for the eastern Black Sea. The most accurate estimation presented in this paper therefore rank the Lower Congo Basin as the third biggest oil-supplying area in the world. The computations of expelled volumes based on the gridded method developed by Macdonald et al. (2015) are detached from the dimensions of the study areas. The orders of magnitude of the surface oil volume are comparable between the Gulf of Mexico (<15 l/km²) and the Lower Congo Basin (<10 l/km²), where oil seep systems are related to salt tectonics in both cases. The seepage intensity per unit area should be compared for miscellaneous geological contexts to understand bypass system efficiency. Thermogenic seep

sites are also associated with a consistent methane flux (Solomon et al., 2009; Hu et al., 2012), benefiting from both thermal cracking and biodegradation (Head et al., 2003; Peters and Moldowan, 2007). Even if the gas quantification associated with oil escapes remains an open question, this amount is presumably considerable. The assessment of contemporary seepage volumes also provides calibration tips for the evaluation of fossil seep systems. Ongoing work integrating the SAR dataset together with subsurface geophysical imagery will help to associate sea surface evidences of oil leakage with seafloor venting features.

Conclusions

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This study integrates a combination of long-term (82 images over 18 years) and short-term (22 images over 10 days) monitoring of SAR scenes to understand the seepage situation in the Lower Congo Basin. The compilation map of recurrent seepage slicks forms diverging oil slick patterns from the central origin area, which allowed the recognition of 102 recurrent oil seeping sites. The seeping area is constrained to the compressive and translated salt-related structural domains, where salt diapirs create efficient oil migration pathways to the sea surface. It was demonstrated that the occurrence rates strongly vary between seeping sites, ranging from 5 to 80%. The correlation of SAR observations with wind speeds showed that slick detection is possible for wind values ranging from 1.5 to 6.5 m/s. Local winds at dormant sites remain within ideal weather conditions demonstrating the intermittent activity and the high variability of seepage flow rates over short periods. This allowed the identification of macro-seep sites of high occurrence rates and micro-seepage sites releasing hydrocarbons sparsely or ephemerally towards the sea surface. Merely half of the seeping sites revealed from 18-year long-term monitoring were active during 10-day short-term monitoring, suggesting that the remaining half remained dormant. Short-term monitoring shows that the slick residence time is low, as the seepage signal is lost between two consecutive SAR scenes separated by 12 h. Sea surface speed displacement estimated from the integration of local winds and subsurface current measurements, coupled with slick lengths allowed us to provide for the first time a statistical study and a quantitative approach to assess the surface slick residence time. This distribution concentrates around a median value of 3h15min. The best estimates of the volumetric outflow give an annual volume of 4380 m³ of oil naturally expelled in the Lower Congo Basin.

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