



# NON STATIONARY POT MODEL FOR EXTREME TEMPERATURE AND PRECIPITATION ANALYSIS IN BURKINA FASO

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## Extreme Climate Events, Climate Change

### Objectives

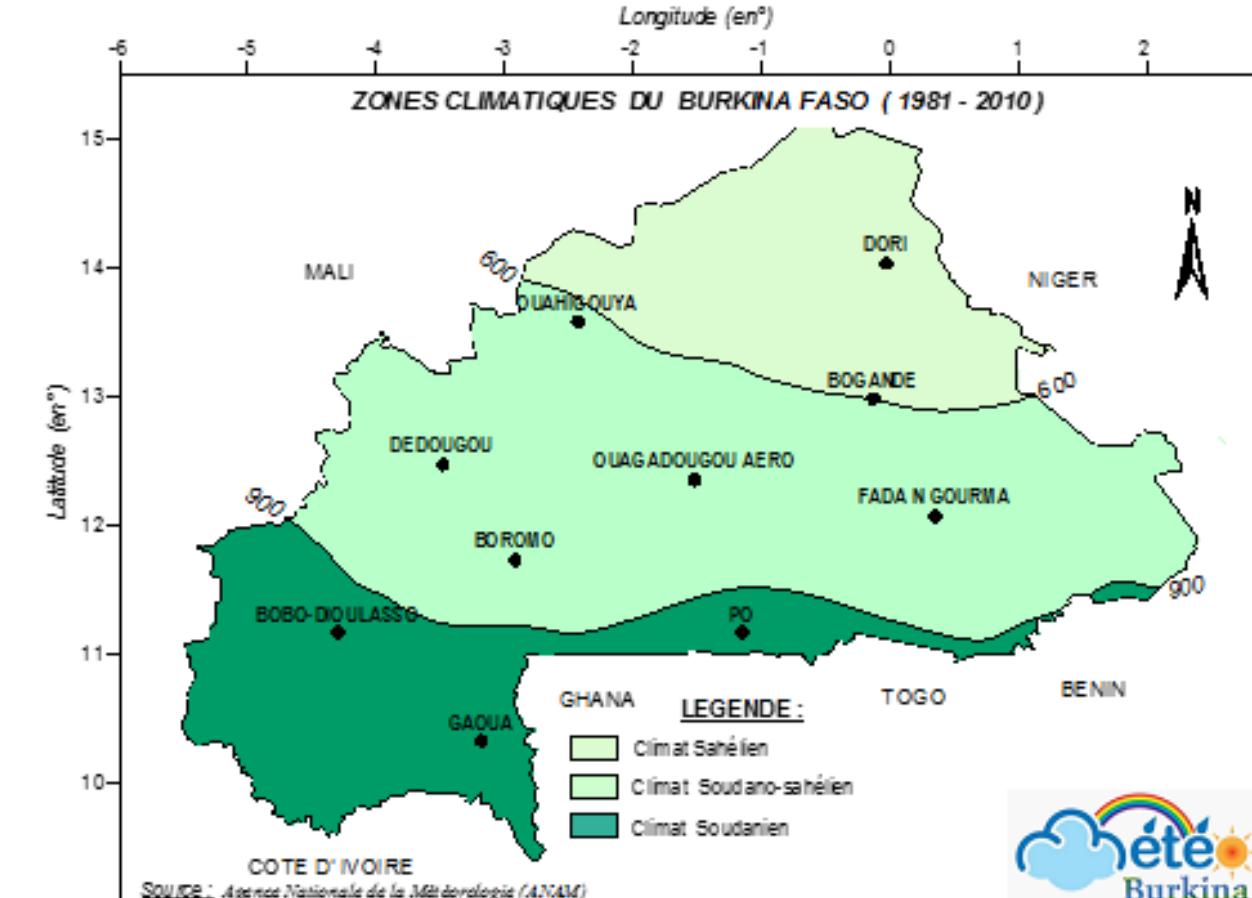
### Ouaga, 2009

- detect trends of extreme weather, in a non-stationary context;
- calculate the extreme quantiles, from the extrapolation of identified trends;
- study the spatial dependence structures of extremes of temperature and precipitation.



### Datasets

#### Stations

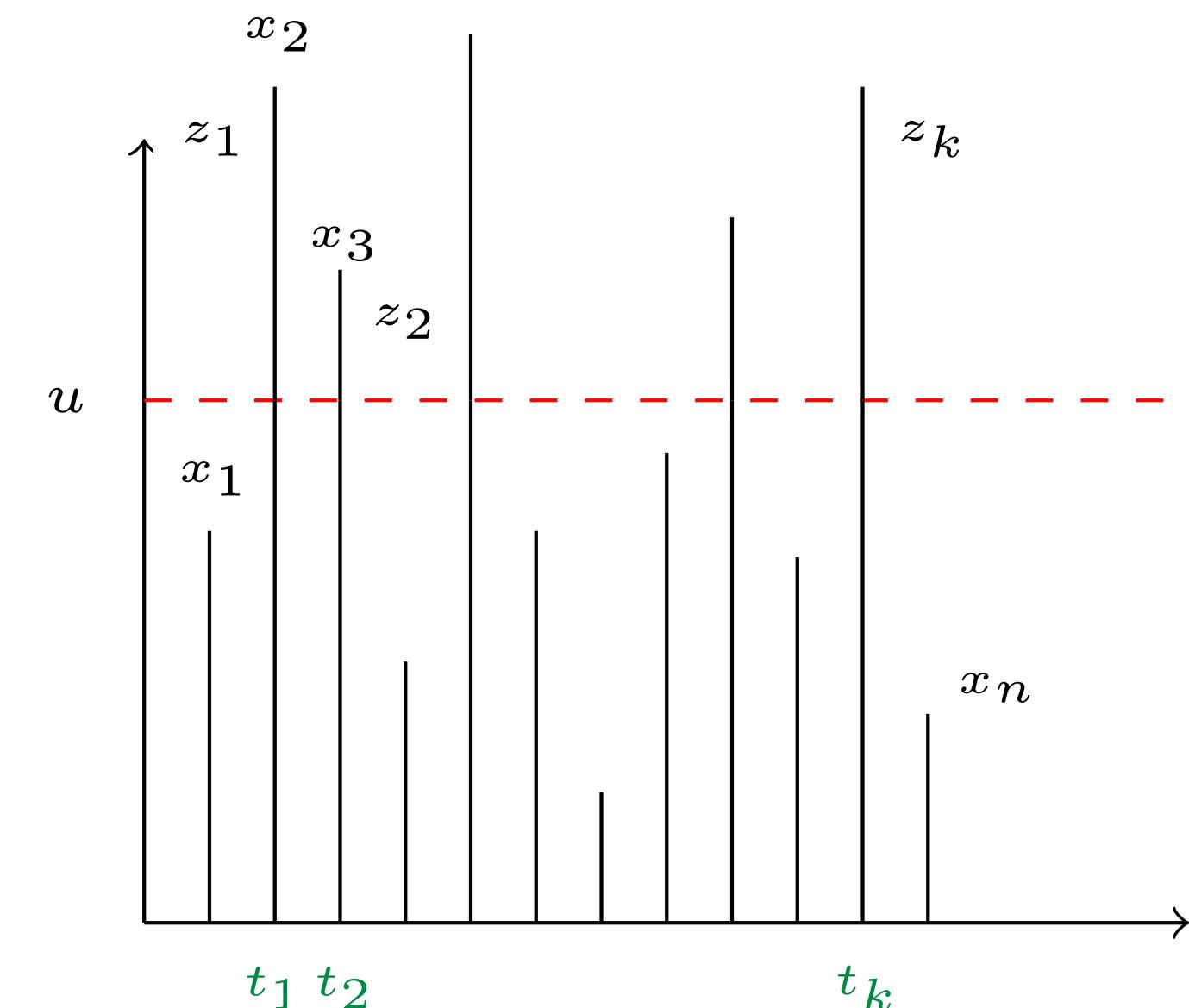


#### Series

- Spatial uniformity, representativeness of the different climatic regimes;
- the quality of the data;
- Hot days: from March 15th to June 15th (temperatures);
- Rainy days: from June to October (precipitation).

### Methodology ([1], [2], [3])

#### Stationary POT



- the law of excesses is given by:

$$S - GP(z | \sigma_u, \xi_u) = 1 - \left(1 + \xi_u \frac{z}{\sigma_u}\right)^{-\frac{1}{\xi_u}}.$$

- The law of occurrences is given by :

$$P(N[t_2 - t_1] = k) = \frac{[\lambda(t_2 - t_1)]^k}{k!} e^{-\lambda(t_2 - t_1)}.$$

#### Non-stationary POT

- In this new statistical framework  $Z \sim NS - GP(\sigma_u(s, t), \xi_u(s, t))$  whose distribution given by:

$$\text{NS-GP}(z) = \begin{cases} 1 - \left(1 + \xi_u(s, t) \frac{z}{\sigma_u(s, t)}\right)^{-\frac{1}{\xi_u(s, t)}}, & \xi_u(s, t) \neq 0 \\ 1 - \exp\left(-\frac{z}{\sigma_u(s, t)}\right), & \xi_u(s, t) = 0 \end{cases}$$

- In this case, for  $B = [t_1, t_2]$ , the law of occurrences of the excesses

$$P(N[B] = k) = \exp(-\Lambda(B)) \frac{\Lambda(B)^k}{k!}, \quad \Lambda(B) = \int_{t_1}^{t_2} \lambda(s, t) ds,$$

- the non-stationary Return Level (RL)  $z_m$ , is solution of the equation:

$$\sum_{t=t_0}^{t_0+mn_z} \left(1 + \frac{\xi}{\sigma_u(s, t)} (z_m(s) - u)\right)^{-\frac{1}{\xi}} \lambda(s, t) = 1.$$

#### Likelihood, Ratio Test

$$\begin{cases} \sigma_u(s, t) = \exp(\sigma_0 + \sigma_1 t + \sigma_2 t^2) \\ \lambda(s, t) = \alpha_0 + \alpha_1 t + \alpha_2 t^2 \\ \xi_u(t) = \xi_0 \end{cases}, s \in \mathcal{S} \subset \mathbb{R}^2.$$

- Estimation of the parameters of  $\lambda(s, t)$ :

$$\ell(\mathbf{t} | \Theta) = -An_z \left( \alpha_0 + \frac{1}{2} \alpha_1 An_z + \frac{1}{3} \alpha_2 A^2 n_z^2 \right) + \sum_{i=1}^k \log(\alpha_0 + \alpha_1 t_i + \alpha_2 t_i^2)$$

- Estimation of the parameters of  $\sigma_u(s, t)$ :

$$\ell(\mathbf{z} | \Theta) = \begin{cases} \sum_{i=1}^k (\sigma_0 + \sigma_1 t_i) - \left(1 + \frac{1}{\xi}\right) \sum_{i=1}^k \log(1 + \xi_0 z_i \exp(-( \sigma_0 + \sigma_1 t_i))), \xi_u \neq 0. \\ -\sum_{i=1}^k (\sigma_0 + \sigma_1 t_i) \\ -\sum_{i=1}^k z_i \exp(-( \sigma_0 + \sigma_1 t_i)), \xi = 0 \end{cases}$$

- Likelihood ratio test

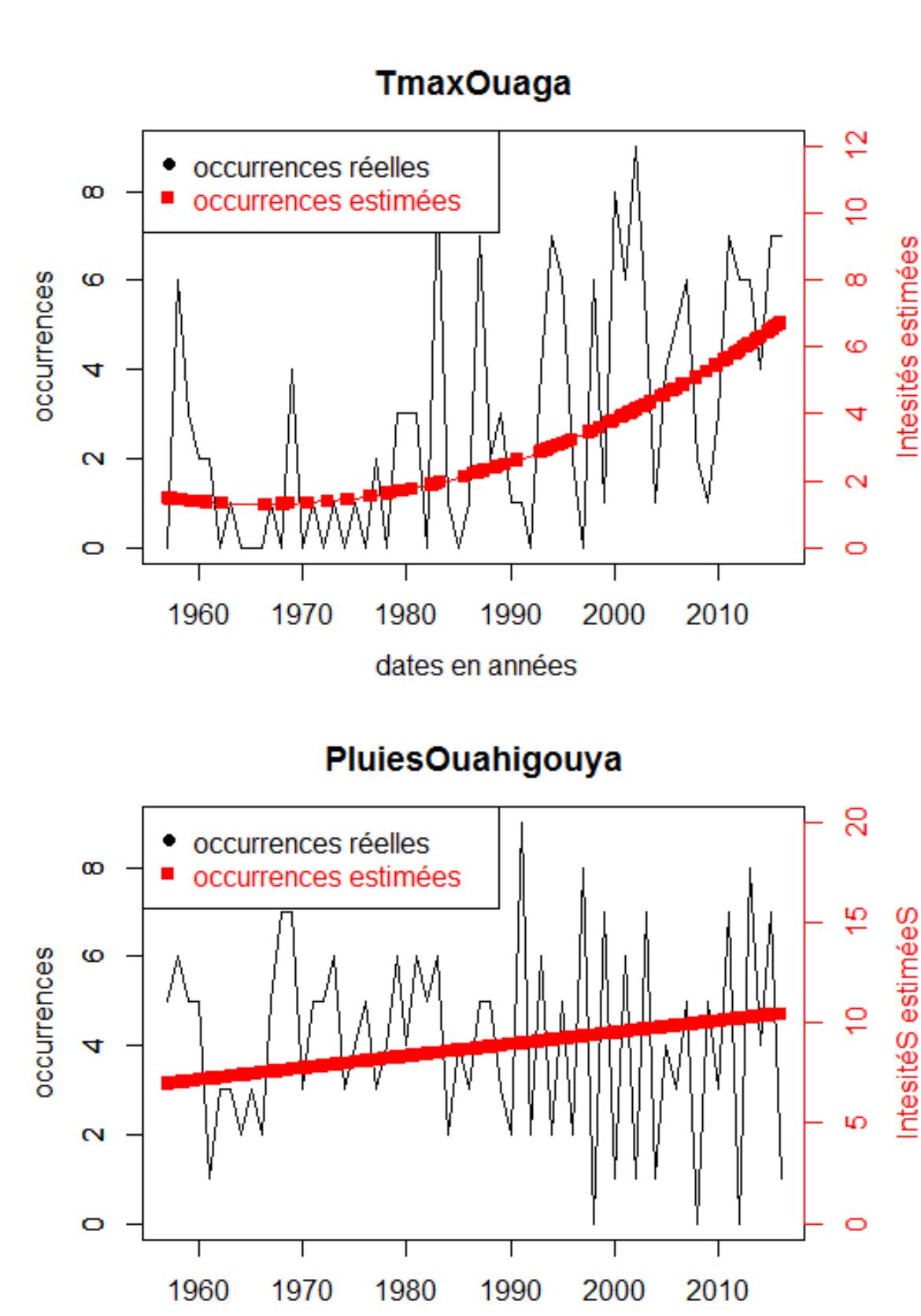
### Results

#### Trends of Extremes

Likelihood	Ratio Oua	Test Dor	Test Ohg	Test Déd	Results Bog
Tmin	$\sigma$ 1	0	0	0	0
	$\lambda$ 2	2	1	1	2
Tmax	$\sigma$ 1	1	0	1	0
	$\lambda$ 2	1	0	2	0
Pluies	$\sigma$ 0	0	0	0	0
	$\lambda$ 0	0	0	1	0
	Fad	Bob	Bor	Pô	Gaou
Tmin	$\sigma$ 1	0	1	0	1
	$\lambda$ 2	1	1	0	2
Tmax	$\sigma$ 0	1	0	0	0
	$\lambda$ 2	2	2	0	1
Pluies	$\sigma$ 0	0	0	0	0
	$\lambda$ 0	0	0	0	1

- An upward trend in extreme temperatures (Tmin and Tmax);
- The scale parameter remains constant in more than half of the stations;
- we have not detect a trend for extreme precipitation.

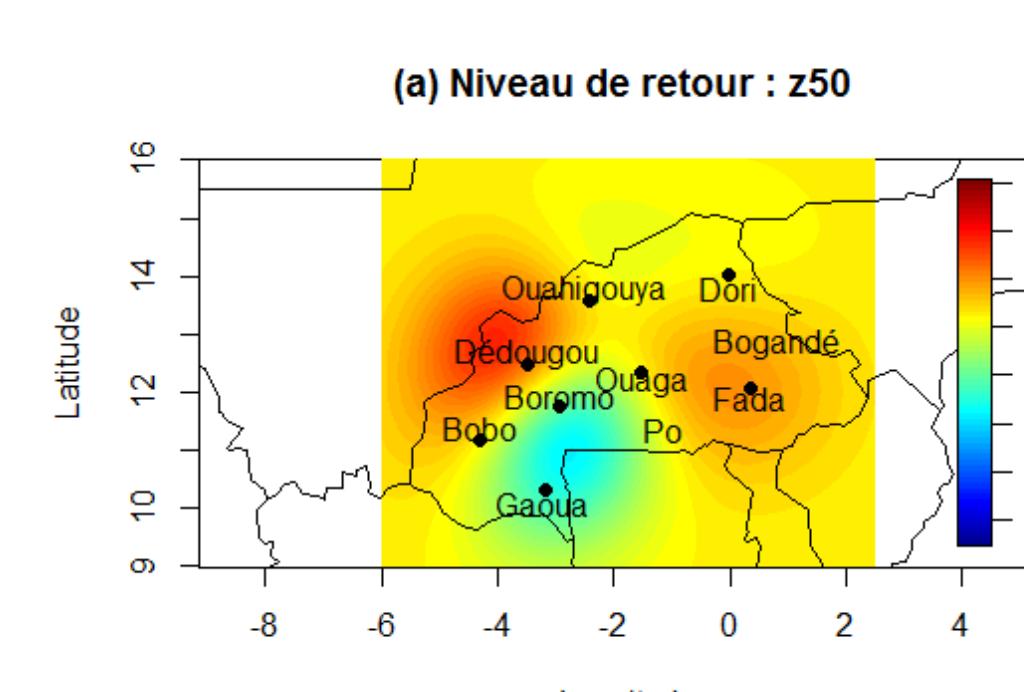
#### Some graphs



#### 50RL[95% CI].

Stations	Tmin	Tmax	Precipitation
Oua	31.8 [30.8, 33.8]	44.4 [43.1, 48.2]	275.2 [179.3, 477.6]
Dor	35.1 [34.1, 36.9]	46.2 [45.6, 46.7]	158.1 [135.1, 283.5]
Ohg	35.3 [33.5, 37.7]	45.8 [45.0, 46.8]	219.4 [149.6, 326.3]
Déd	35.5 [35.2, 38.0]	44.4 [43.5, 45.7]	172.2 [139, 328.6]
Bog	33.7 [32.4, 36.5]	39.9 [39.4, 40.6]	220.2 [154.7, 377.3]
Fad	31.0 [30.1, 31.9]	44.7 [44.0, 46.0]	200.9 [151.3, 437.2]
Bob	30.1 [29.4, 30.5]	41.1 [40.6, 41.4]	196.5 [149.4, 500]
Bor	33.3 [31.9, 35.4]	44.4 [43.8, 45.8]	224.5 [152.4, 377.2]
Pô	30.8 [30.1, 31.8]	44.4 [43.0, 47.9]	231.9 [156.7, 384.5]
Gao	29.8 [28.7, 31.3]	42.1 [41.7, 42.5]	249.6 [164.3, 418.7]

#### Spatial structure of RL50



- A slight spatial structure for extreme temperatures.

- for precipitation, no spatial structure could be exhibited on these estimates.

### Conclusion and perspectives

- This study has described recent changes in temperature and extreme precipitation in Burkina Faso in a non-stationary setting;
- Over the period 1957-2016, the intensity of occurrences of extreme temperatures shows a slight (polynomial) upward trend with a slight spatial structure;
- No trend has been detected for extreme precipitation nor spatial structure;
- In perspective, we plan to better integrate the spatial aspect using the  $\ell$ -Pareto processes (Ferreira and de Haan (2014) Domby and Ribatet (2015)).

### References

- [1] Parey S., Malek F., Laurent C., Dacunha-Castelle D., (2007), *Trends and climate evolution: Statistical approach for very high temperatures in France*, Climatic Change 81:331–352
- [2] Parey S., Hoang T.T.H. , Dacunha-Castelle D., (2010b), *Different ways to compute temperature return levels in the climate change context*, Environmetrics 2010; 21: 698–718.
- [3] Acero F. J. , Garcia J. A., Gallego M. C., Parey S., Dacunha-Castelle D. , (2014), *Trends in summer extreme temperatures over the Iberian Peninsula using nonurban station data*, Journal of Geophysical Research: Atmospheres, Vol. 119, 39–53