



HAL
open science

From dome to duplex: Convergent gravitational collapse explains coeval intracratonic doming and nappe tectonics, central Australia

Youseph Ibrahim, Patrice Rey, Donna Whitney, Christian Teyssier, Françoise Roger, Valérie Bosse, Bénédicte Cenko

► To cite this version:

Youseph Ibrahim, Patrice Rey, Donna Whitney, Christian Teyssier, Françoise Roger, et al.. From dome to duplex: Convergent gravitational collapse explains coeval intracratonic doming and nappe tectonics, central Australia. *Geology*, In press, 10.1130/g51721.1 . hal-04401661

HAL Id: hal-04401661

<https://hal.science/hal-04401661>

Submitted on 18 Jan 2024

HAL is a multi-disciplinary open access archive for the deposit and dissemination of scientific research documents, whether they are published or not. The documents may come from teaching and research institutions in France or abroad, or from public or private research centers.

L'archive ouverte pluridisciplinaire **HAL**, est destinée au dépôt et à la diffusion de documents scientifiques de niveau recherche, publiés ou non, émanant des établissements d'enseignement et de recherche français ou étrangers, des laboratoires publics ou privés.

1 From dome to duplex: convergent gravitational collapse
2 explains coeval intracratonic doming and nappe tectonics,
3 central Australia

4 **Youseph Ibrahim¹, Patrice F. Rey¹, Donna L. Whitney², Christian Teyssier², Françoise**
5 **Roger³, Valérie Bosse⁴ and Bénédicte Cenki³**

6 ¹School of Geosciences, University of Sydney, NSW 2006, Australia

7 ²Department of Earth Sciences, University of Minnesota, Minnesota 55455, USA

8 ³Géosciences Montpellier, Université de Montpellier, CNRS, 34095 Montpellier cedex, France

9 ⁴Laboratoire Magmas et Volcans, Campus Universitaire des Cézeaux, 63178 Aubière Cedex,
10 France

11 **ABSTRACT**

12 In central Australia, an apparently coeval gneiss dome (Entia Dome) developed adjacent
13 to a thrust belt (Arltunga Nappe Complex) within an intracratonic setting. Here we employ a
14 combination of fieldwork, geochronology, and numerical modeling to investigate the structure
15 and tectonic evolution of these features. We present a structural model linking an extensional
16 domain comprising the Entia Dome, across a transitional zone containing the Bruna décollement
17 zone and the Illogwa shear zone, into a contractional zone comprising thrusts and duplexes of the
18 Arltunga Nappe Complex. Supported by numerical modeling, we propose a tectonic model in
19 which the dome and nappe complex formed synchronously because of the convergent
20 gravitational collapse of the 30-40 km deep Paleozoic Harts Range rift.

21 **INTRODUCTION**

22 Defying traditional plate tectonics, the Alice Springs Orogeny (ASO) developed between
23 450 to 300 Ma deep within the interior of the Australian continent (Fig. 1), at significant
24 distances from any plate boundaries. It spans ~600 km, following a WNW-ESE striking corridor
25 amidst the weaker regions between the North, South, and West Australian cratons (Nixon et al.,
26 2022). Before the ASO, a large portion of the Australian Paleoproterozoic interior remained
27 buried beneath the Neoproterozoic to Devonian Centralian Superbasin, remnants of which

28 include the Amadeus Basin (Fig. 1). During much of the Palaeozoic Era, Australia existed as a
29 Gondwana promontory bordered by active margins to the NW (Metcalf, 2021), N (Jost et al.,
30 2018) and Pacific margins (Rosenbaum, 2018). Consequently, the intracratonic ASO has been
31 attributed to a combination of mantle dynamics (Houseman and Molnar, 2001; Roberts and
32 Houseman, 2001) and far-field stresses (Klootwijk, 2013; Silva et al., 2018).

33 The region NE of the Amadeus basin presents an intriguing geological association of the
34 Entia gneiss dome, in which Paleoproterozoic upper amphibolite facies gneisses were exhumed
35 alongside a series of nappes (Fig. 1). The nappe complex comprises the amphibolite to granulite
36 facies rocks (10 kbar, 850°C; Hand et al., 1999; Mawby et al., 1999) of the Harts Range
37 Metamorphic Complex (HRMC), the Paradise Nappes, and their lower-grade Neoproterozoic
38 covers (Forman, 1971). These isoclinal fold nappes were thrust onto the greenschist facies Ruby
39 Gap duplex and White Range duplex, which consist of Amadeus basin sequences, overlapping
40 the Amadeus basin (Dunlap and Teyssier, 1995). Collectively, these nappes constitute the
41 Arltunga Nappe Complex.

42 Understanding of the regional geology took a transformative turn when detrital zircon
43 data revealed that the HRMC, into which the Entia dome was emplaced, did not belong to the
44 Paleoproterozoic basement. Instead, these gneisses are the metamorphic equivalent of
45 Neoproterozoic to Palaeozoic marine siliciclastic and limestone sediments of the Amadeus Basin
46 with interlayered metavolcanics (Maidment et al., 2013; Tucker et al., 2015). Amphibolite to
47 granulite facies metamorphism between 470 to 450 Ma is associated with bedding-parallel
48 fabrics, suggesting syn-rift high-T metamorphism within the Harts Range Rift, a 30 to 40 km
49 deep sub-basin adjacent to the Amadeus Basin (Hand et al., 1999; Mawby et al., 1999; Maidment
50 et al., 2005; Tucker et al., 2015). Granulite facies conditions seem to have been confined to
51 deeper parts of the Harts Range Rift, with little effect on the adjacent Paleoproterozoic basement
52 (Mawby et al., 1999).

53 Episodic pegmatite emplacement between 450 and 300 Ma within the Entia Dome and
54 the HRMC suggests the maintenance of high-T conditions in the deeper crust during most of the
55 Paleozoic (Buick et al., 2008). Hence, the Entia Dome is a post-460 Ma structure that ascended
56 through the HRMC during the ASO. The driver for protracted partial melting following an
57 episode of high-grade metamorphism in this intracontinental setting remains enigmatic (Buick et
58 al., 2008; Asimus et al., 2023).

59 This paper focuses on the structure and tectonic history of the region between the Entia
60 Dome and the Ruby Gap Duplex (RGD) across the intervening Illogwa Shear Zone (ISZ; Fig. 1).
61 We document a structural continuity from the Entia Dome to the RGD and argue that the dome
62 and duplex, active between 340 and 310 Ma, represent an extensional and contractional domain,
63 respectively, linked by a translational domain that consists of the Bruna Décollement Zone
64 (BDZ) and the ISZ. We propose that the deep Harts Range rift basin led to a significant
65 gravitational potential anomaly that relaxed via inward convergent gravitational collapse.

66 **A TRANSECT FROM DOME TO DUPLEX**

67 **Entia Dome**

68 The Entia Dome comprises migmatitic quartzofeldspathic gneisses and layered
69 amphibolites and calc-silicates. A composite layer-parallel compositional banding fabric defines
70 an internal architecture involving two subdomes separated by a planar high-strain zone (Fig. 1,
71 2A). The subdome cores host the Huckitta and Inkamulla granodiorites dated (U-Pb zircon) at
72 1762 Ma and 1773 Ma, respectively (Maidment et al., 2005). Both granodiorites display a
73 magmatic to solid-state fabric concordant with overlying migmatitic gneisses. On average,
74 lineations within the dome trend to the NE (Fig. 1) and plunge 20-30° to the NE. Polyphase
75 structures in the migmatites show evidence of at least three fabric-forming events (Fig 2D-E),
76 with a clear regional partitioning. In the dome, the metatexite compositional layering and layer
77 parallel foliation is folded into tight to isoclinal recumbent folds, ranging from centimeter to 10's
78 of meters in amplitude (Fig. S1A-B). The sub-horizontal migmatitic layering displays locally
79 sub-horizontal asymmetric boudinage and low-angle extensional shear bands with melt
80 segregation that points to a strong extensional tectonic regime above the solidus (Fig. S1C). The
81 magmatic fabric in the granodiorites preserves remnants of older structures (Fig. S2). At the
82 margin of the dome, the recumbent folds are refolded by cascading recumbent folds (Fig. S3).
83 Granite-bearing extensional fractures, some axial planar to these second-generation folds, show
84 that these folds also developed at temperatures above the solidus (Fig. 2D-E). In between the
85 subdomes, the high-strain zone records a strong horizontal NW-SE shortening, with tight to
86 isoclinal upright folds (Fig. S4). The strain distribution is consistent with that expected during
87 the exhumation of a migmatitic-cored dome, with vertical shortening and horizontal extension on

88 top of the dome, gravity folding along its margin (Cruden, 1990), and horizontal shortening
89 between the subdomes (Rey et al., 2017).

90 U-Pb ages of metamorphic monazite from Entia gneisses between 365 and 308 Ma
91 (Wade et al., 2008 and references therein; Varga et al., 2021), as well as U-Pb ages of
92 metamorphic zircon growth in the Huckitta Granodiorite at $\sim 332 \pm 3$ Ma (Maidment et al., 2005)
93 and $\sim 330 \pm 6$ Ma (Hand et al., 1999) point to a protracted metamorphism from 360 to 310 Ma,
94 peaking at 7-9 kbar and 680-720°C (Varga et al., 2021 and references therein) and melt present
95 conditions (Asimus et al., 2023). To better constrain the age of deformation, we conducted U-Th-
96 Pb LA-ICPMS geochronology (Fig. S8) on monazite from a folded leucocratic vein within a
97 metatexite between the subdomes (Figs. 1, 2D, S4). Two distinct date populations identified at
98 339 ± 4 Ma and 314 ± 4 Ma, as well as a concordant date at 365 ± 8 Ma, are interpreted as two
99 pulses of doming and melt extraction during a ~ 365 Ma to 310 Ma period of protracted
100 metamorphism, melting and deformation.

101 **South Margin of the Entia Dome**

102 The Bruna décollement zone deforms the contact between the Paleoproterozoic Entia
103 gneisses and the Neoproterozoic to Paleozoic Irindina Gneiss at the base of the HRMC (Fig. 1)
104 (James et al., 1989). Separating the basement and cover, the Bruna Granitic Gneiss (BGG) (Fig.
105 S5) has been interpreted as a ~ 1745 Ma (U/Pb zircon) laccolith (Ding and James, 1985;
106 Mortimer et al., 1987; Cooper et al., 1988). The BGG records heterogeneous deformation (James
107 et al., 1989). Along the S margin of the Entia Dome, the BDZ corresponds to a prominent S-
108 dipping normal shear zone. A strong L-S fabric carries a mineral and stretching lineation
109 plunging to the SSW (Fig. 2F). Sm/Nd whole rock-garnet-hornblende isochron dating of BGG
110 from the sheared NW margin of the Entia Dome yielded an age of 449 ± 10 Ma, interpreted to
111 reflect a phase of Ordovician deformation (Mawby et al., 1999). Similar ages have been reported
112 in the Harts Range (e.g., Buick et al., 2008). To constrain the age of shearing along the S margin
113 of the dome, we have performed *in situ* LA-ICPMS U-Pb dating of titanite from an isoclinally
114 folded aplite vein (Fig. 2C) hosted in sheared BGG (Fig. 1). We obtained a date of 315 ± 4 Ma
115 (Fig. S8), interpreted as the minimum emplacement age for the vein and maximum age for
116 shearing and folding, closely agreeing with the 314 ± 4 Ma age obtained in the high-strain zone
117 between the subdomes.

118 South of the Entia Dome, the BDZ flattens into a broad synform (Fig. 2A). The L-S
119 fabric is flat-lying, with the prominent stretching lineation gently plunging to the SSW or NNE.
120 In some localities, strain shadows around K-feldspar and S-C fabrics in the BGG support a top-
121 to-the-SW sense of shear (Fig. S6A-B). However, deformation was heterogeneous, varying in
122 intensity and involving both simple and pure shear (Fig. S6C). Underlying the BGG, the
123 Strangways gneisses are exposed in the core of a few discrete antiforms (Fig. 1). The BGG thins
124 out above the roof thrust of the duplex (Fig. 2).

125 **Ruby Gap duplex and the Illogwa shear zone**

126 South of the BGG, the ISZ consists of quartzofeldspathic gneisses and schists, with mafic
127 and calc-silicate interlayering, reminiscent of the rocks in the Entia Dome. The foliation dips 30-
128 50° to the N and carries a down-dip stretching and muscovite lineation. We have observed S-C
129 fabrics, δ -clasts, and σ -clasts recording clear top-to-the-SW kinematics (Fig. 2E, Fig. S7). This
130 shear zone is in structural conformity with the overlying upper amphibolite facies BDZ to the N
131 and the underlying greenschist facies RGD to the S. Although the ISZ is often presented as a
132 retrograde shear zone (Foden et al., 1995), it is part of an inverted metamorphic gradient linking
133 the upper-amphibolite facies at the Entia Dome's S margin to the greenschist facies fabric of the
134 RGD (Collins and Teyssier, 1989; Dunlap and Teyssier, 1995). We speculate that the ISZ fabric
135 is the axial planar fabric of an anticlinal fold nappe made mainly of Paleoproterozoic gneisses,
136 which could be structurally underneath the Paradise Nappes (Figs. 2, 3). $^{207}\text{Pb}/^{206}\text{Pb}$ zircon ages
137 of 1794 ± 6 Ma and 1769 ± 3 Ma are interpreted as crystallization ages and later alteration,
138 respectively. In contrast, the 327 ± 2 Ma $^{40}\text{Ar}-^{39}\text{Ar}$ age obtained on muscovite (Reno and Fraser,
139 2021) defining the lineation (Fig. S7) can be interpreted as the age of the ASO fabric.

140 Underneath and in structural conformity with the ISZ, the RGD is a S-directed duplex
141 comprising Amadeus basin sequences and the underlying Paleoproterozoic basement. To the
142 West, it occupies the lowest structural level of the Arltunga Nappe Complex, underlying the
143 Paradise and HRMC nappes and coeval with the structurally equivalent White Range duplex
144 (Dunlap and Teyssier, 1995) (Fig. 1). To the first order, what is preserved is the duplicated limb
145 of an anticlinal fold nappe (Fig. 2A). The basement and the competent Heavitree Quartzite are
146 attached, and the overlying incompetent Bitter Springs Fm (Fig. 2B) acts as a décollement.
147 Thrusting occurred between 336 and 311 Ma ($^{40}\text{Ar}-^{39}\text{Ar}$ on white mica) at temperatures

148 straddling the ductile-brittle transition (250-300 °C; Dunlap, 1997). Under these conditions,
149 thrust sheets internally deform by stretching in the transport direction.

150 **Summary**

151 The structure across our transect consists of the Entia Dome with a double-dome
152 architecture, adjacent to the Arltunga Nappe Complex to the S, comprising a stack of imbricated
153 metamorphic fold nappes including the ISZ (Figs. 2A, 3). The structural geology is conformal,
154 with a progressive decrease in metamorphic conditions from upper amphibolite facies in the
155 Entia Dome and the BGG, to lower amphibolite facies across the ISZ and greenschist facies in
156 the RGD (Figs. 2A, 3). Structures within the Entia Dome and the BDZ developed under
157 extension, while the Arltunga Nappe Complex developed under contraction. Exhumation of the
158 Entia Dome, shearing along the BDZ and ISZ, and thrusting of the Arltunga Nappe Complex are
159 age-bracketed between ~345 and ~310 Ma. The structural continuity and geochronology indicate
160 that these features are synchronous and interrelated.

161 **TECTONIC MODEL**

162 The Entia Dome was emplaced into the 30-40 km deep Harts Range Rift basin (i.e.,
163 Irindina Province), the base of which reached granulite facies metamorphism and melting at
164 ~470 Ma (Maidment et al., 2013; Tucker et al., 2015). The density contrast between the
165 Neoproterozoic to Paleozoic sedimentary basin infill and the metamorphic Paleoproterozoic
166 basement creates a strong horizontal pressure gradient, leading to significant gravitational
167 stresses acting laterally towards the basin (Rey et al., 2001). As the geotherm in the basin
168 reached melting conditions, these gravitational forces may have overcome viscous strength to
169 drive centripetal flow, forcing the exhumation of the Entia Dome and the development of gravity
170 nappes.

171 To test this hypothesis, we ran a set of thermal-mechanical numerical experiments using
172 the *Underworld* framework (Moresi et al. 2007). The model setup comprises a 34 km deep rift
173 basin, embedded into a metamorphic basement overlying a viscous mantle (Fig. 4A). The
174 temperature field reflects conditions at ~470 Ma and delivers partial melting in the deeper level
175 of the basin as well as the basement. Rheology is temperature, stress, strain, strain rate and melt
176 dependent. Detailed model information is provided in the data repository.

177 We have tested static, mild, and faster convergent and divergent velocity conditions. All
178 sets of experiments lead to the exhumation of a basement-cored double-dome, pushing the basin
179 infill upward and laterally into gravity nappes (Fig. 4B-D). This experiment (Fig. 4) suggests that
180 the contractional collapse of the Harts Range Rift basin exhumed the Entia Dome and triggered
181 S-SW gliding of gravity-driven fold nappes, accommodated by isoclinal folding and foliation-
182 parallel slip. At cooler conditions, limbs of the isoclinal fold nappes are imbricated forming the
183 RGD and White Range duplex. A transitional domain links the extensional exhumation of the
184 dome to the contractional deformation at the front of the nappe complex (Fig. 4D). The delay
185 between granulite facies conditions at 470-450 Ma and doming at ~340-310 Ma may be due to
186 rifting's divergent velocity balancing the convergent gravitational velocity until ~355 Ma and
187 slow conductive heating delaying basement weakening and convergent flow.

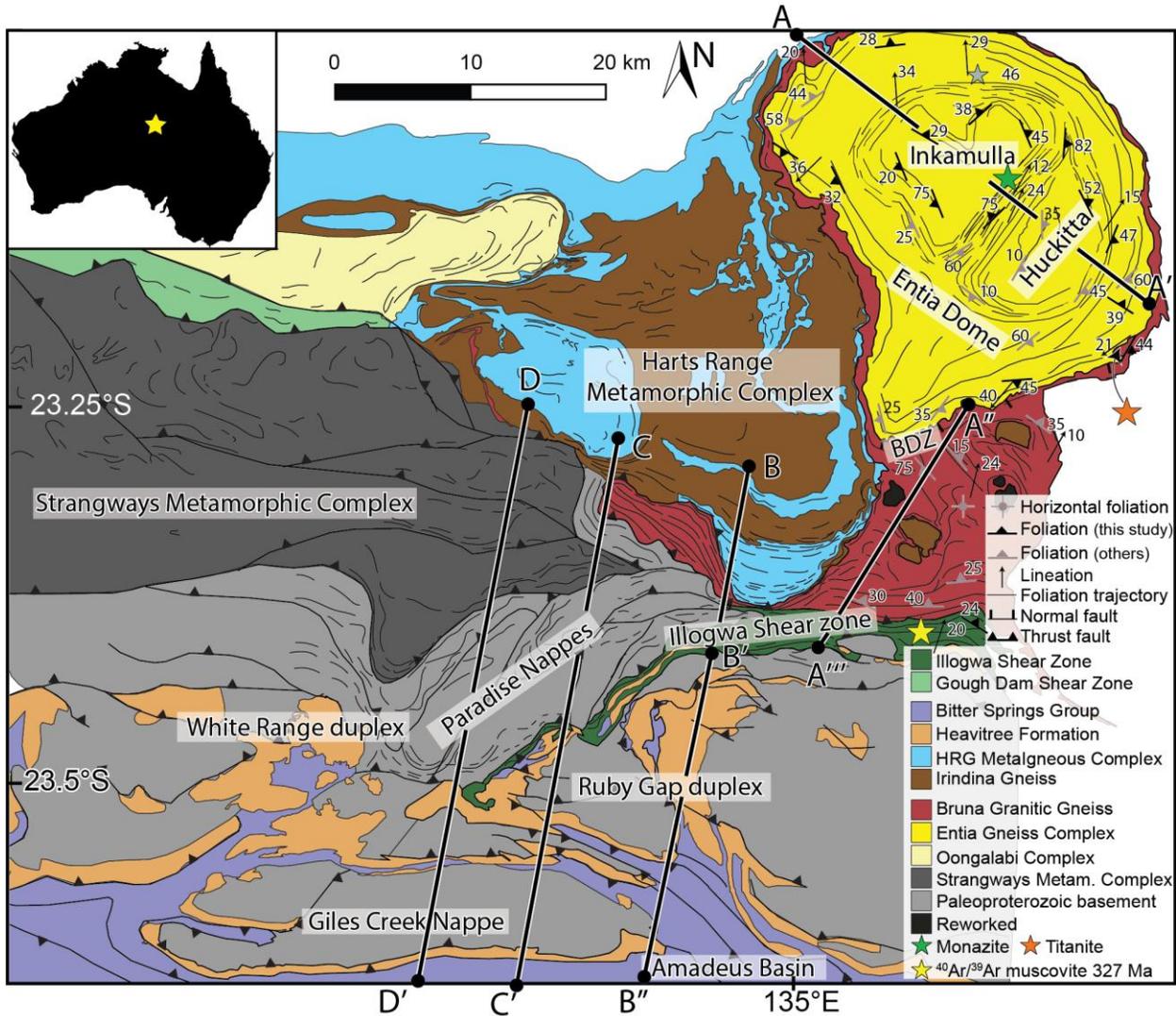
188 **CONCLUSIONS**

189 We propose a tectonic model linking the Entia Dome and the RGD based on i/ the
190 structural and metamorphic continuity between the dome and the duplex, ii/ geochronological
191 constraints showing synchronous dome exhumation and nappe emplacement from ~340 to ~310
192 Ma, and iii/ coupled thermo-mechanical numerical experiments illustrating the geodynamics.
193 Our model invokes the contractional gravitational collapse of the Neoproterozoic to Devonian
194 Harts Range Rift basin as the driver for exhumation of the Entia dome, associated with
195 gravitational spreading of the basin infill. This gravity-driven deformation was superimposed on
196 far-field stresses driven by plate tectonics and may have contributed to deformation along the N
197 margin of the Amadeus basin. While the processes described here explain some Carboniferous
198 features observed in the east Arunta region, it cannot explain the totality of the ASO strain field.

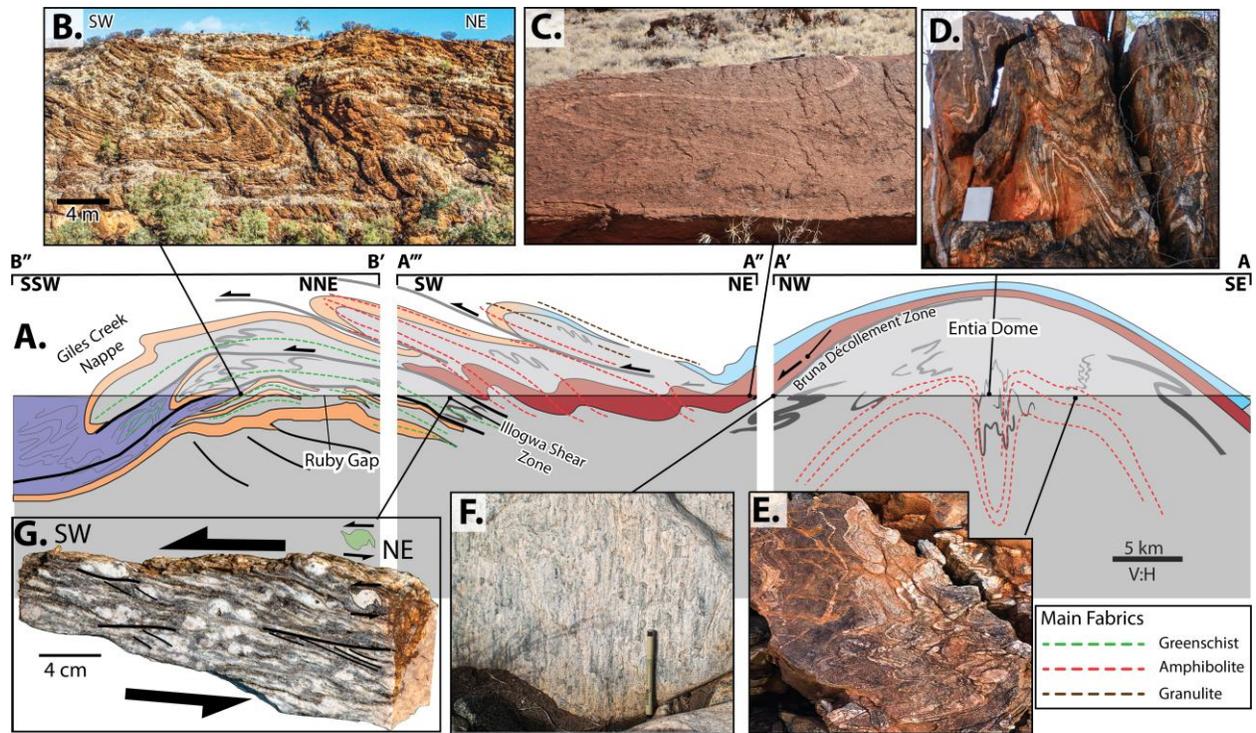
199 **ACKNOWLEDGMENTS**

200 We acknowledge funding by Australian Research Council grants ARC-LP190100146 and ARC-
201 DP220100709, and support by NCI's National Computational Merit Allocation Scheme, Pawsey
202 Supercomputing Centre, and Underworld provided by AuScope. B.C. received funding from
203 European Horizon 2020 grant no. 793978. We thank our reviewers for their comments and
204 insights.

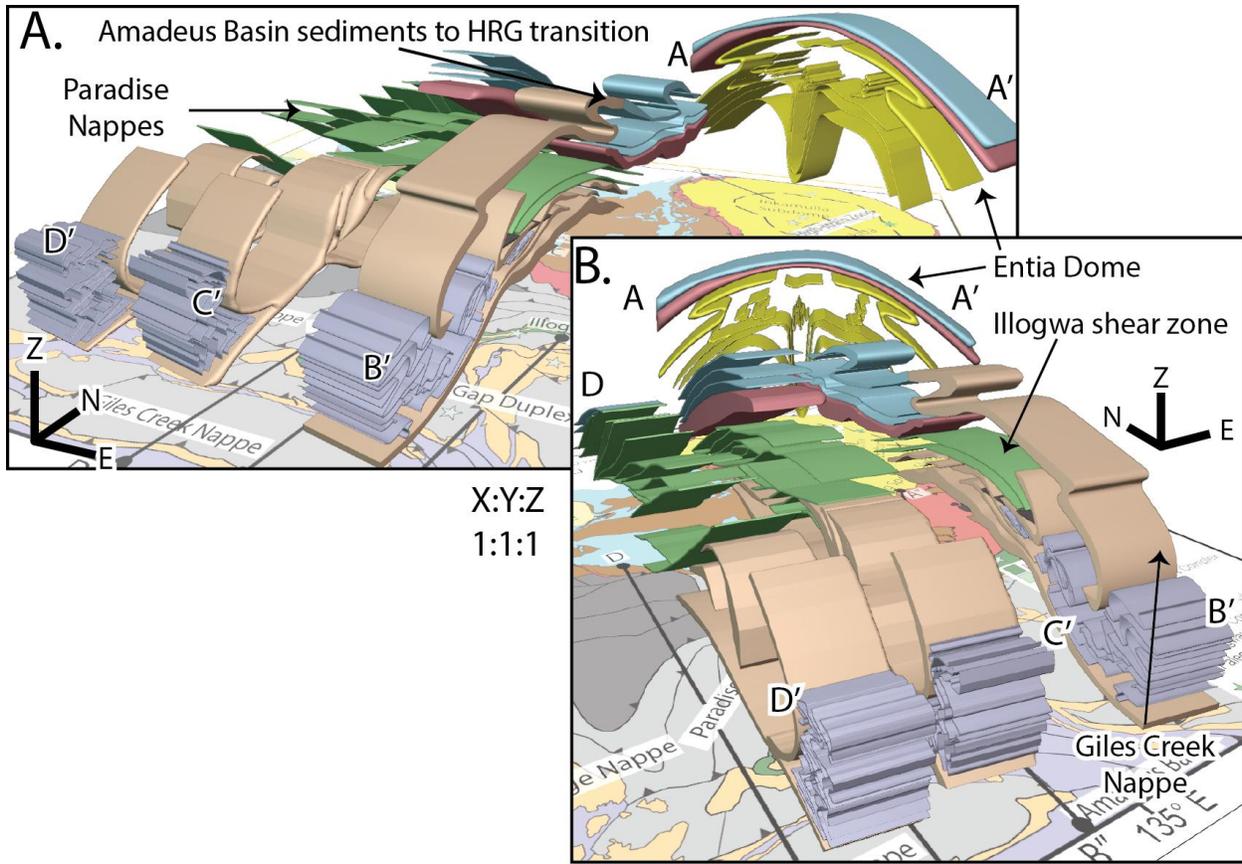
205



207
 208 Figure 1. Geological map of the study area, central Australia, including location of
 209 geochronology data ($^{40}\text{Ar}/^{39}\text{Ar}$ on muscovite from Reno and Fraser, 2021), using a combination
 210 of our own data and data from Forman (1971), Shaw et al. (1984), Shaw and Freeman (1990),
 211 and Dunlap (1992). HRG—Harts Range Group.

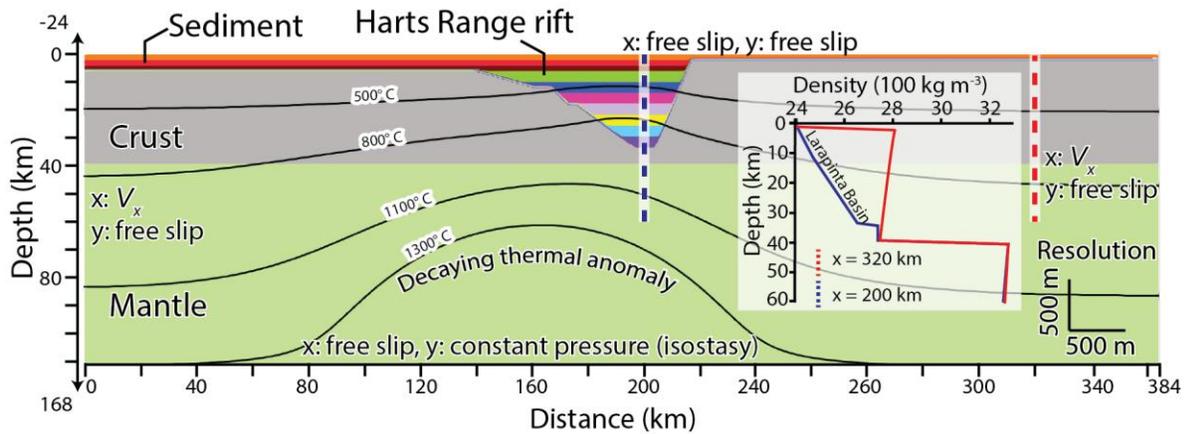


212
 213 Figure 2. (A) Composite cross section from Ruby Gap Duplex to Entia Dome. Major thrust faults
 214 are shown as thick black lines. Lighter shading indicates interpretation above the current
 215 topography. The Harts Range Metamorphic Complex, including the Irindina Gneiss, is shown in
 216 light blue. Section lines and legend are shown in Figure 1. (B) Folds in the Bitter Springs Group.
 217 (C) Folded aplite vein with axial planar foliation in the Bruna décollement zone (pen for scale).
 218 (D, E) Upright folds (notebook for scale) (D) refolded into cascading folds (pen for scale) (E) in
 219 the high-strain dome envelope. (F) L-S tectonite within the Bruna décollement zone. (G)
 220 Kinematic indicators in the Illogwa shear zone.

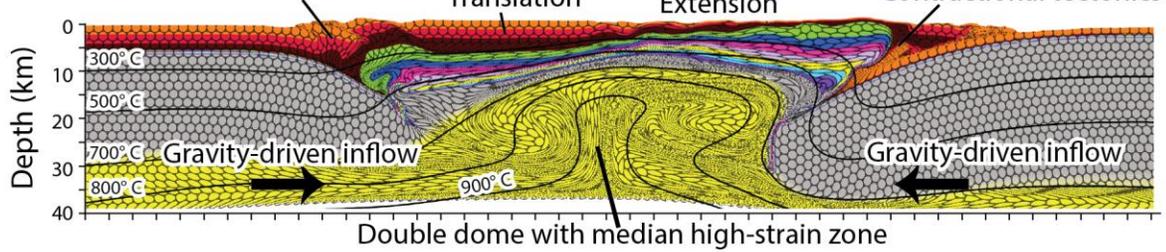


221
 222 Figure 3. Two schematic views of the structural architecture of Entia Dome and Arltunga Nappe
 223 Complex. Paradise Nappes are colored green. Legend and section lines are shown in Figure 1.
 224 HRG—Harts Range Group.

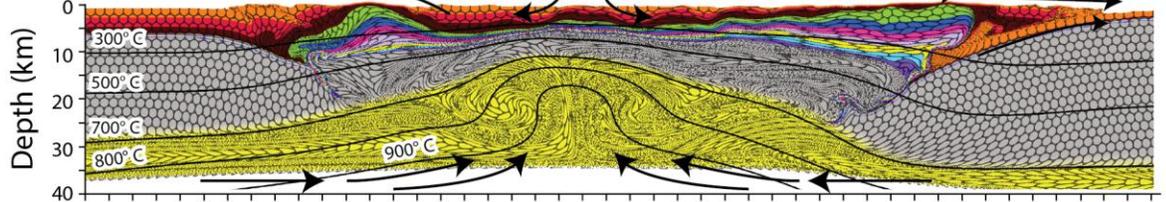
A. 0 m.y. Pre-doming lithosphere architecture



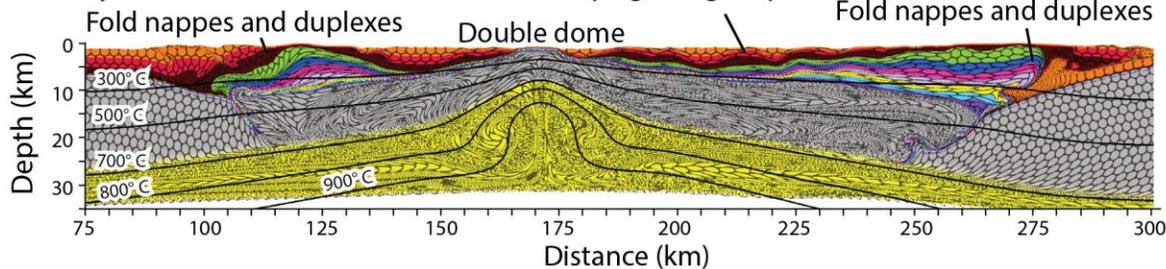
B. 3.4 m.y. Contractional tectonics



C. 6.7 m.y.



D. 10 m.y.



225
 226 Figure 4. Numerical experiment showing contractional gravitational collapse of a deep and hot
 227 rift basin triggering exhumation of a gneiss dome and associated gravity nappes. (A) Initial
 228 model setup showing materials distribution, density profiles, boundary conditions, and isotherms.
 229 Velocity boundary condition V_x is applied to both vertical walls. V_x is compressional at 0.05
 230 mm yr^{-1} for 4 m.y. and extensional at 0.5 cm yr^{-1} for the following 6 m.y. Decaying heat
 231 anomaly approximates temperature conditions associated with rifting. Blue and red dashed lines

232 show locations of density profile. (B–D) Model evolution. Finite strain ellipses and isotherms are
233 overlaid on the deforming material. Thin black arrows show flow direction.

234 REFERENCES

235 Asimus, J.L., Daczko, N.R., and Ezad, I.S., 2023, Melt-present deformation at the Entia Dome,
236 Central Australia: A metamorphic core complex formed during lower crustal tectonic extrusion:
237 *Lithos*, v. 448–449, <https://doi.org/10.1016/j.lithos.2023.107170>.

238 Buick, I.S., Storkey, A., and Williams, I.S., 2008, Timing relationships between pegmatite
239 emplacement, metamorphism and deformation during the intra-plate Alice Springs Orogeny,
240 central Australia: *Journal of Metamorphic Geology*, v. 26, p. 915–936,
241 <https://doi.org/10.1111/j.1525-1314.2008.00794.x>.

242 Collins, W.J., and Teyssier, C., 1989, Crustal scale ductile fault systems in the Arunta Inlier,
243 central Australia: *Tectonophysics*, v. 158, p. 49–66, [https://doi.org/10.1016/0040-](https://doi.org/10.1016/0040-1951(89)90314-4)
244 [1951\(89\)90314-4](https://doi.org/10.1016/0040-1951(89)90314-4).

245 Cooper, J.A., Mortimer, G.E., and James, P.R., 1988, Rate of Arunta Inlier evolution at the
246 eastern margin of the Entia Dome, central Australia: *Precambrian Research*, v. 40–41, p. 217–
247 231, [https://doi.org/10.1016/0301-9268\(88\)90069-1](https://doi.org/10.1016/0301-9268(88)90069-1).

248 Cruden, A.R., 1990, Flow and fabric development during the diapiric rise of magma: *The Journal*
249 *of Geology*, v. 98, p. 681–698, <https://doi.org/10.1086/629433>.

250 Ding, P., and James, P.R., 1985, Structural evolution of the Harts Range Area and its
251 implication for the development of the Arunta Block, central Australia: *Precambrian Research*,
252 v. 27, p. 251–276, [https://doi.org/10.1016/0301-9268\(85\)90015-4](https://doi.org/10.1016/0301-9268(85)90015-4).

253 Dunlap, W.J., 1992, Structure, kinematics, and cooling history of the Arltunga Nappe Complex,
254 central Australia [Ph.D. thesis]: University of Minnesota, 276 p.

255 Dunlap, W.J., 1997, Neocrystallization or cooling? ⁴⁰Ar/³⁹Ar ages of white micas from low-
256 grade mylonites: *Chemical Geology*, v. 143, p. 181–203, [https://doi.org/10.1016/S0009-](https://doi.org/10.1016/S0009-2541(97)00113-7)
257 [2541\(97\)00113-7](https://doi.org/10.1016/S0009-2541(97)00113-7).

258 Dunlap, W.J., and Teyssier, C., 1995, Thermal and structural evolution of the intracratonic
259 Arltunga Nappe Complex, central Australia: *Tectonics*, v. 14, p. 1182–1204,
260 <https://doi.org/10.1029/95TC00335>.

261 Foden, J., Mawby, J., Kelley, S., Turner, S., and Bruce, D., 1995, Metamorphic events in the
262 eastern Arunta Inlier, Part 2. Nd-Sr-Ar isotopic constraints: *Precambrian Research*, v. 71, p. 207–
263 227, [https://doi.org/10.1016/0301-9268\(94\)00062-V](https://doi.org/10.1016/0301-9268(94)00062-V).

264 Forman, D.J., 1971, The Arltunga Nappe Complex, Macdonnell Ranges, Northern Territory,
265 Australia: *Journal of the Geological Society of Australia*, v. 18, p. 173–182,
266 <https://doi.org/10.1080/00167617108728756>.

267 Hand, M., Mawby, J.O., Kinny, P., and Foden, J., 1999, U–Pb ages from the Harts Range,
268 central Australia: Evidence for early Ordovician extension and constraints on Carboniferous
269 metamorphism: *Journal of the Geological Society*, v. 156, p. 715–730,
270 <https://doi.org/10.1144/gsjgs.156.4.0715>.
271
272 Houseman, G., and Molnar, P., 2001, Mechanisms of lithospheric rejuvenation associated with
273 continental orogeny, in Miller, J.A., et al., eds., *Continental Reactivation and Reworking*:
274 Geological Society, London, Special Publication 184, p. 13–38,
275 <https://doi.org/10.1144/GSL.SP.2001.184.01.02>.
276
277 James, P.R., Macdonald, P., and Parker, M., 1989, Strain and displacement in the Harts Range
278 detachment zone: A structural study of the Bruna Gneiss from the western margin of the Entia
279 Dome, central Australia: *Tectonophysics*, v. 158, p. 23–48, [https://doi.org/10.1016/0040-](https://doi.org/10.1016/0040-1951(89)90313-2)
280 [1951\(89\)90313-2](https://doi.org/10.1016/0040-1951(89)90313-2).
281
282 Jost, B.M., Webb, M., and White, L.T., 2018, The Mesozoic and Palaeozoic granitoids of north-
283 western New Guinea: *Lithos*, v. 312–313, p. 223–243,
284 <https://doi.org/10.1016/j.lithos.2018.04.027>.
285
286 Klootwijk, C., 2013, Middle–Late Paleozoic Australia-Asia convergence and tectonic extrusion
287 of Australia: *Gondwana Research*, v. 24, p. 5–54, <https://doi.org/10.1016/j.gr.2012.10.007>.
288
289 Maidment, D.W., Hand, M., and Williams, I.S., 2005, Tectonic cycles in the Strangways
290 Metamorphic Complex, Arunta Inlier, central Australia: Geochronological evidence for
291 exhumation and basin formation between two high-grade metamorphic events: *Australian*
292 *Journal of Earth Sciences*, v. 52, p. 205–215, <https://doi.org/10.1080/08120090500139414>.
293
294 Maidment, D.W., Hand, M., and Williams, I.S., 2013, High grade metamorphism of sedimentary
295 rocks during Palaeozoic rift basin formation in central Australia: *Gondwana Research*, v. 24,
296 p. 865– 885, <https://doi.org/10.1016/j.gr.2012.12.020>.
297
298 Mawby, J., Hand, M., and Foden, J., 1999, Sm-Nd evidence for high-grade Ordovician
299 metamorphism in the Arunta Block, central Australia: *Journal of Metamorphic Geology*, v. 17,
300 p. 653–668, <https://doi.org/10.1046/j.1525-1314.1999.00224.x>.
301
302 Metcalfe, I., 2021, Multiple Tethyan ocean basins and orogenic belts in Asia: *Gondwana*
303 *Research*, v. 100, p. 87–130, <https://doi.org/10.1016/j.gr.2021.01.012>.
304
305 Moresi, L., Quenette, S., Lemiale, V., Mériaux, C., Appelbe, B., and Mühlhaus, H.-B., 2007,
306 Computational approaches to studying non-linear dynamics of the crust and mantle: *Physics of*
307 *the Earth and Planetary Interiors*, v. 163, p. 69–82, <https://doi.org/10.1016/j.pepi.2007.06.009>.
308
309 Mortimer, G.E., Cooper, J.A., and James, P.R., 1987, U-Pb and Rb-Sr geochronology and
310 geological evolution of the Harts Range ruby mine area of the Arunta Inlier, central Australia:
311 *Lithos*, v. 20, p. 445–467, [https://doi.org/10.1016/0024-4937\(87\)90029-6](https://doi.org/10.1016/0024-4937(87)90029-6).

312 Nixon, A.L., Glorie, S., Fernie, N., Hand, M., De Vries Van Leeuwen, A.T., Collins, A.S.,
313 Hasterok, D., and Fraser, G., 2022, Intracontinental fault reactivation in high heat production
314 areas of central Australia: Insights from apatite fission track thermochronology: *Geochemistry,*
315 *Geophysics, Geosystems*, v. 23, <https://doi.org/10.1029/2022GC010559>.
316

317 Reno, B.L., and Fraser, G.L., 2021, Summary of results—Joint NTGS-GA geochronology
318 project: Constraining cooling and deformation in the eastern Aileron Province through
319 $^{40}\text{Ar}/^{39}\text{Ar}$ stepheating of hornblende, muscovite, and biotite: *Northern Territory Geological*
320 *Survey Record 2021-001*, <https://geoscience.nt.gov.au/gemis/ntgsjspui/handle/1/91246>.
321

322 Rey, P., Vanderhaeghe, O., and Teyssier, C., 2001, Gravitational collapse of the continental
323 crust: Definition, regimes and modes: *Tectonophysics*, v. 342, p. 435–449,
324 [https://doi.org/10.1016/S0040-1951\(01\)00174-3](https://doi.org/10.1016/S0040-1951(01)00174-3).
325

326 Rey, P.F., Mondy, L., Duclaux, G., Teyssier, C., Whitney, D.L., Bocher, M., and Prigent, C.,
327 2017, The origin of contractional structures in extensional gneiss domes: *Geology*, v. 45, p. 263–
328 266, <https://doi.org/10.1130/G38595.1>.
329

330 Roberts, E.A., and Houseman, G.A., 2001, Geodynamics of central Australia during the
331 intraplate Alice Springs Orogeny: Thin viscous sheet models, in Miller, J.A., et al., eds.,
332 *Continental Reactivation and Reworking: Geological Society, London, Special Publication 184*,
333 p. 139–164, <https://doi.org/10.1144/GSL.SP.2001.184.01.08>.
334

335 Rosenbaum, G., 2018, The Tasmanides: Phanerozoic tectonic evolution of eastern Australia:
336 *Annual Review of Earth and Planetary Sciences*, v. 46, p. 291–325,
337 <https://doi.org/10.1146/annurev-earth-082517-010146>.
338

339 Shaw, R.D., and Freeman, M.J., 1990, Quartz, Northern Territory, sheet 5951: Australia Bureau
340 of Mineral Resources, Geology and Geophysics, scale 1:100,000, <https://geoscience.nt.gov.au/gemis/ntgsjspui/handle/1/81611>.
341

342 Shaw, R.D., Stewart, A.J., and Rickard, M.J., 1984, Arltunga–Harts Range region, Northern
343 Territory: Australia Bureau of Mineral Resources, Geology and Geophysics, scale 1:100,000,
344 <https://geoscience.nt.gov.au/gemis/ntgsjspui/handle/1/81909>.
345

346 Silva, D., Piazzolo, S., Daczko, N.R., Houseman, G., Raimondo, T., and Evans, L., 2018,
347 Intracontinental orogeny enhanced by far-field extension and local weak crust: *Tectonics*, v. 37,
348 p. 4421–4443, <https://doi.org/10.1029/2018TC005106>.
349

350 Tucker, N.M., Hand, M., and Payne, J.L., 2015, A rift-related origin for regional medium-
351 pressure, high-temperature metamorphism: *Earth and Planetary Science Letters*, v. 421, p. 75–
352 88, <https://doi.org/10.1016/j.epsl.2015.04.003>.
353

354 Varga, J., Raimondo, T., Morrissey, L., Kelsey, D.E., and Hand, M., 2021, Pressure–
355 temperature–time constraints on gneiss dome formation in an intracontinental orogen: *Journal of*
356 *Metamorphic Geology*, v. 40, p. 457–488, <https://doi.org/10.1111/jmg.12635>.
357

358 Varga, J., Raimondo, T., Hand, M., Curtis, S., and Daczko, N., 2022, Hydration, melt production
359 and rheological weakening within an intracontinental gneiss dome: *Lithos*, v. 432–433, [https://](https://doi.org/10.1016/j.lithos.2022.106872)
360 doi.org/10.1016/j.lithos.2022.106872.

361
362 Wade, B.P., Hand, M., Maidment, D.W., Close, D.F., and Scrimgeour, I.R., 2008, Origin of
363 metasedimentary and igneous rocks from the Entia Dome, eastern Arunta region, central
364 Australia: A U-Pb LA-ICPMS, SHRIMP and Sm-Nd isotope study: *Australian Journal of Earth*
365 *Sciences*, v. 55, p. 703–719, <https://doi.org/10.1080/08120090801982868>.