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3 **Clustering and interfacial segregation of radiogenic Pb in a mineral host – inclusion system**

4 **Tracing two-stage Pb and trace element mobility in monazite inclusions in rutile**

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17

Abstract

18 Accessory minerals like zircon, rutile and monazite are routinely studied to inform about the timing and
19 nature of geological processes. These studies are underpinned by our understanding of the transfer
20 processes of trace elements and the assumption that the isotopic systems remain undisturbed. However,
21 the presence of microstructures or Pb-bearing phases in minerals can lead to the alteration of the Pb
22 isotopic composition. To gain insight into the relationship between Pb isotopic alterations from inclusions
23 and microstructures, this study focused on inclusions from an ultra-high temperature metamorphic rutile.
24 The studied inclusions are submicron monazites, a common mineral rich in Pb but normally not present in
25 rutile. The sample is sourced from Mt. Hardy, Napier Complex, East Antarctica, an ultra-high
26 temperature (UHT) metamorphic terrane. By applying correlative analytical techniques including electron
27 backscatter diffraction mapping, transmission electron microscopy (TEM), and atom probe tomography, it
28 is shown monazite inclusions are often in contact with low-angle boundaries and yield no preferred
29 orientation. TEM shows the monazite core has a mottled texture due to the presence of radiation damage
30 and nanoclusters associated with the radiation damage defects that are rich in U, Pb, and Ca. Some
31 monazites exhibit a core-rim structure. The rim yields clusters composed of Ca- and Li-phosphate that
32 enclose Pb nanoclusters that are only present in small amounts compared to the core, with Pb likely
33 diffused into the rutile-monazite interface. These textures are the result of two-stages of Pb mobility.
34 Initial Pb segregation was driven by volume diffusion during UHT metamorphism (2500 Ma). The second
35 stage is a stress-induced recrystallization during exhumation, leading to recrystallization of the monazite
36 rim and trace element transport. The isotopic signature of Pb trapped within the rutile-monazite interface
37 constrains the timing of Pb mobility to c. 550 Ma.

38

39 **Keywords:** Rutile, Monazite, Pb mobility, Atom Probe Tomography

40

Introduction

41 Rutile (TiO₂) is a common accessory mineral that is stable at mid- to lower-crustal metamorphic
42 conditions (Clark et al., 2019; Dacheille et al., 1968; Harley, 2016). Zr-in-rutile geothermometry and rutile
43 U-Pb geochronology are increasingly used to provide temperature and time constraints on the evolution of
44 metamorphic terranes (Clark et al., 2018; Kooijman et al., 2010; Meinhold, 2010; Smye and Stockli,
45 2014; Watson et al., 2006; Zack et al., 2004a). Mineral inclusions in rutile are geologically significant
46 with inclusions of corundum in rutile providing constraints on growth mechanisms (Daneu et al., 2014),
47 whilst the molar fraction of Zr-bearing mineral inclusions can be used to re-integrate Zr concentrations
48 for the estimation of peak metamorphic temperatures (Mitchell and Harley, 2017; Pape et al., 2016). The
49 ability of rutile to trap and isolate inclusions from retrograde metamorphic overprints and fluid alteration
50 events may allow inclusion assemblages in rutile to be used to characterize metamorphic conditions in
51 ultra-high pressure (UHP) and ultra-high temperature (UHT) metamorphic terranes (Hart et al., 2016;
52 Hart et al., 2018).

53 Despite the uses of inclusion analyses, the presence of small (<1 µm diameter) inclusions in rutile has the
54 potential to compromise compositional and isotopic analysis of the host mineral by techniques, such as
55 Laser ablation inductively coupled plasma mass spectrometry (LA-ICP-MS), which analyses significantly
56 larger volumes than the inclusion volume. Little is currently known about the geochemistry of sub-
57 micrometer mineral inclusions or how their presence and composition may affect the analysis of the host.
58 In part, this reflects the difficulty in measuring the trace element and isotopic compositions of such small
59 inclusions. Furthermore, there is little information regarding how such inclusions may evolve as the
60 pressure-temperature conditions of the host rock change over time, or if such inclusions are truly isolated
61 from the matrix of the rock in which the host mineral resides. In this contribution, the structural setting
62 and compositional character of micrometer-scale (200-1100 nm) monazite inclusions within ultra-high
63 temperature metamorphic rutile are studied to provide some constraints on host-inclusion relationships
64 and determine the mechanisms by which compositional heterogeneity in the inclusion may be developed.

65 The mobility of Pb in monazite has received significant attention in nanoscale studies using transmitted
66 electron microscopy (TEM) and atom probe tomography (APT) (Bingen and Van Breemen, 1998;
67 Fougereuse et al., 2021a; Fougereuse et al., 2018; Grand'Homme et al., 2016; Seydoux-Guillaume et al.,
68 2003; 2012; 2019; Turuani et al., 2022, 2023). However, such studies have typically analyzed the internal
69 parts of large monazite crystals, ignoring the potential complications associated with Pb mobility adjacent
70 to mineral-mineral interfaces that characterize small inclusions.

71 The complementary use of APT, TEM and electron backscattered diffraction (EBSD) provides unique
72 insight into Pb migration within the inclusions (Reddy et al., 2020). The results obtained show evidence
73 for multiple mechanisms of Pb and trace element mobility, provide important insights into the segregation
74 of Pb to mineralogical grain boundaries and yield important new observations that influence the
75 interpretation of isotopic data from sub-micrometer monazite grains in a mineral – host system.

76

77 **Geological setting and sample description**

78 The sample used for this study was collected from Mt Hardy in the Napier Complex of East Antarctica
79 (Fig. 1a) and is from an Archean Mg-Al rich metapelitic layer. The Napier Complex consists of a suite of
80 Archean (2990 – 2800 Ma, Black, Williams, & Compston,
81 1986 Harley & Black, 1987. ortho- and paragneisses that underwent a complex deformational and
82 metamorphic history including two ultrahigh temperature (UHT) metamorphic events at c. 2850 and c.
83 2580 Ma with their timing constraint by zircon dating. (Clark et al., 2018; Harley, 2016; James and Black,
84 1981; Sheraton et al., 1987). The c. 2500 Ma UHT event is widely believed to have resulted in the
85 formation of the highest regional-scale temperatures for exposed crustal rocks on Earth, with temperature
86 estimates, coming from the region where Mt Hardy is located, of between 1050–1120 °C (Harley and
87 Motoyoshi, 2000; Hokada, 2001). These high temperatures are thought to have persisted for in excess of
88 100 Myr (Clark et al., 2018; Harley, 2016). Two younger events, mostly in the southern region of the

89 Napier Complex have been described by Black et al. (1983); Black et al. (1984). Zircon and monazite U-
90 Pb dating constrain an amphibolite-facies metamorphic event to take place between 1073-1094 Ma. Rb-Sr
91 ages record new magmatic and hydrothermal activity to take place around 522 Ma (Black et al., 1983). A
92 study on galena inclusions in monazite (Turuani et al., 2022) shows the latter event (522 Ma) was more
93 localized while the event at 1073 Ma was more pervasive within the Napier Complex.

94 The sample is comprised of dry residual layers of porphyroblastic garnet–cordierite–spinel–sapphirine–
95 osumilite (now replaced by symplectites of K-feldspar–cordierite–dendritic orthopyroxene) and quartz–
96 mesoperthite leucosomes (Fig. 1b). Rutile occurs as inclusions in garnet and within both the residual
97 layers and leucosomes, rutile is inferred to form part of the peak UHT assemblage. In this sample, a
98 population of sector zoned “soccer-ball” type zircons interpreted to be metamorphic in origin yielded a
99 concordia age of 2492 ± 7 Ma and a Ti-in-zircon temperature of $915 \pm 30^\circ\text{C}$ (Clark et al., 2018). The
100 monazite inclusions reside in two rutile grains. The monazites studied were extracted from rutile in direct
101 contact with garnet, mesoperthite, and traces of biotite. Submicron monazite inclusions are also present in
102 the biotite present in the direct vicinity of the rutile grain of interest. No large monazite grains are present
103 within this sample. However, the earliest monazite in the Napier complex likely crystallized around 3070
104 Ma (Black et al., 1984).

105

106

Methods

107 **Electron backscatter diffraction & electron dispersive X-ray spectroscopy**

108 The thin section has previously been analyzed by LA-ICP-MS. Following LA-ICP-MS, the thin section
109 was re-polished and carbon-coated and analyzed by SEM. SEM characterization was undertaken on a
110 TESCAN Mira3 field emission scanning electron microscope housed at the John de Laeter Centre (JdLC),
111 Curtin University, Perth, Australia and included electron backscatter diffraction (EBSD) and energy-
112 dispersive x-ray spectroscopy (EDS) mapping. Analyses were undertaken with an accelerating voltage of
113 20 kV and an instrument specific ‘beam intensity’ of 17 using a working distance of 20 mm and 70°

114 specimen tilt. For mapping of rutile, a 500 nm step size was used and the match units for rutile were
115 obtained from the American Mineralogist crystallography database (801 448-45x) (Swope et al., 1995).
116 Match units for monazite were derived from the crystallographic data of (Ni et al., 1995), following
117 (Erickson et al., 2015). EBSD mapping of the monazite inclusions was done using a 50 nm step size.
118 Post-processing of EBSD data was performed in Matlab®, version R2022b, with the free toolbox MTEX
119 Version 5.8.2 (Bachmann et al., 2010). Images are constructed using local misorientation and the mean
120 misorientation of the rutile grain. The maps were plotted using the ‘roma’ scientific color map (Cramer,
121 2018). Post-processing involved a noise-reduction by removal of groups of <5 adjacent pixels (‘Wild’
122 spikes/shards) with crystal orientations within 10° of those of their neighbors. The procedure was
123 followed by applying a 5 × 5 pixel median filter. The filter smooths the orientations in the EBSD data
124 while preserving subgrain boundaries. Slip system characterization was undertaken following established
125 methodologies (Lloyd et al., 1997; Prior et al., 2002; Reddy et al., 2007).

126

127 **Focused-ion-beam SEM and TEM**

128 Analysis by transmission electron microscope (TEM) and atom probe tomography (APT) requires
129 samples to be prepared in the form of electron-transparent thin foils (TEM) or needle-shaped specimens
130 (APT). These samples were prepared at two different institutes.

131 **Curtin University.** TEM and APT specimens analyzed at Curtin University were prepared on a
132 Tescan Lyra3 Ga Focused-Ion Beam Scanning Electron Microscope (FIB-SEM), housed in the John de
133 Laeter Centre, Curtin University. The Tescan Lyra3 comes equipped with an electronically operated
134 nanomanipulator and Pt gas injection system. The FIB was operated with an accelerating voltage of 30
135 kV for the monoisotopic (mass 69 Da) Ga⁺-source. For backscatter electron (BSE) imaging in SEM-
136 mode, an accelerating voltage of the electron beam was varied between 5, 10 & 20 kV to provide a depth
137 control on the size of the monazite inclusions below the surface. All, TEM and APT, specimens were
138 extracted from one, out of the two, grains that contained monazite inclusions.

139 The TEM foil was prepared by cutting a vertical slice using the FIB and lifted out using the
140 nanomanipulator. The TEM foil was mounted on a copper half-grid and thinned until electron transparent
141 (< 100 nm) followed by a 2 kV clean-up routine to remove damage and Ga implantation caused by the 30
142 kV ion beam. After initial imaging by TEM, the sample was further thinned to 30-50 nm to reduce the
143 curvature effect of the inclusion during Scanning TEM (S/TEM) EDS.

144 **Jean Monnet University (UJM).** One TEM foil was prepared using a Thermo Fisher Scientific
145 FEI 125 Helios Nanolab 600i focused-ion beam scanning electron microscope (FIB/SEM) hosted by the
146 MANUTECH USD platform in Saint-Etienne (France). Before extraction of the sample, the area of
147 interest is protected with a 1 μm carbon coat. A 30 kV 10 nA Ga-ion beam is used to cut out the foil
148 (15x10x1 μm). The foil is then lifted out and deposited on a Cu half-grid, thinned with an ion beam until
149 ~ 100 nm thickness, and then “cleaned up” at low voltage (2 kV) to remove the amorphous part, resulting
150 in a sample ~60 nm thick.

151

152 **Transmission Electron Microscopy**

153 TEM analysis at Curtin University was performed on the FEI Talos FS200X Field Emission Gun TEM
154 equipped with a Super-X EDS detector housed in the John de Laeter Centre. The TEM was operated at
155 200 kV. TEM imaging was conducted in both bright and dark field (BF and DF) modes. The TEM
156 diffraction investigation was conducted with the assistance of Kikuchi patterns that were generated with
157 the convergent beam electron diffraction (CBED). A small spot was used to minimize the beam related
158 damages to the sample, and the sample was tilted to the right diffraction zone axis or two-beam condition
159 to acquire selected area diffraction (SAD) patterns. Both BF and DF TEM imaging were undertaken with
160 the objective apertures after tilting the sample to a diffraction condition within 20°. TEM, STEM, and
161 EDS data acquisition were conducted with the Velox software.

162 Chemical analysis was done by the attached two pairs of super X detectors. The TEM is fitted with four
163 scanning transmission electron microscopy. (STEM) detection systems: High Angle Annular Dark Field
164 (HAADF), upper Dark Field (DF4), lower Dark Field (DF2), and BF. The contrasts in HAADF are

165 predominantly made from chemical/phase differences and the contrast in BF is predominantly made from
166 orientation differences. The DF4 and DF2 detectors reveal both chemical and orientation contrasts at
167 different levels.

168 EDS data was acquired with a beam current of 0.6 nA and collected over 45-90 min for each presented
169 map.

170 HAADF TEM images provide information based on atomic number (Z) contrast with high z values being
171 brighter.

172 TEM characterization in UJM Saint-Etienne was performed with a Cs-corrected TEM (NeoARM200F
173 Cold FEG) operated at 200 kV, operated by the CLYM (Consortium Lyon Saint-Etienne de Microscopie)
174 and hosted within the Hubert Curien Laboratory. The instrument is equipped with a wide-angle energy
175 dispersive X-ray (EDX) spectrometer SDD CENTURIO-X from JEOL, two STEM detectors (annular
176 dark field and annular bright field from JEOL and Gatan) with a CMOS camera Gatan Rio (4kx4k), a
177 CCD camera Ultrascan from Gatan (2kx2k), and an electron energy-loss spectrometer (EELS) (Gatan GIF
178 Quantum ER). The resolution in STEM at 200 kV is 0.78 Å due to the presence of a Cs corrector CEOS
179 ASCOR.

180

181 **Sample preparation for atom probe tomography**

182 For the preparation of APT specimens, monazite was marked by depositing Pt buttons in the middle of
183 the targeted inclusion prior to the deposition of a protective layer (Rickard et al., 2020). APT specimen
184 preparation followed lift-out and sharpening procedures described elsewhere (Rickard et al., 2020;
185 Thompson et al., 2007). A wedge-shaped sample was cut using the ion beam at an angled surface and
186 lifted out using a nanomanipulator. The wedge was brought to the APT specimen holder consisting of a
187 coupon with 22 silicon posts to hold the specimens. The wedge was aligned over the top of a silicon post
188 and a segment was attached using ion beam deposition with a Pt precursor. The wedge above the post was
189 cut leaving ± 2 μm of sample on the silicon post. This process is repeated until all material in the wedge
190 has been placed on posts. Following the mounting of specimens on the silicon posts, the specimens were

191 then sharpened using an annular milling technique until needle-shaped specimens approximately 2.5 μm
192 in length and less than <100 nm wide at the tip were formed. This was followed by a 2 kV clean-up
193 routine to remove the damaged layer. All specimens were imaged by secondary electron (SE) and
194 backscatter electron imaging (BSE) imaging in SEM-mode only at a working distance of 6 mm and an
195 accelerating voltage of 10 kV.

196

197 **Atom probe tomography**

198 APT analyses were conducted at the Geoscience Atom Probe Facility, Curtin University. This system is a
199 Cameca Local Electrode Atom Probe (LEAP) 4000X HR that allows laser-assisted field evaporation and
200 high mass-resolving power. Data were collected using a 125 kHz pulse rate, a laser pulse energy of 100
201 pJ, a base temperature of 70° K and an ion detection rate of 0.3% based on (Fougerouse et al., 2018;
202 Verberne et al., 2019). Post-processing was done using Cameca's Integrated Visualisation and Analysis
203 Software (IVAS) 3.8.0. Reconstructions are based on the density of rutile and using the shank-angle
204 reconstruction mode because of the poly-phased nature of our region of interest (Fougerouse et al.,
205 2021c).

206 Atoms are evaporated as both single ions and molecular species with a range of different ionization states.
207 Peaks in the mass spectra were labelled per individual isotope for specific ionization states and ranged
208 with a constant width of 0.2 Da. For visualization and compositional analyses, isoconcentration surfaces
209 (Hellman et al., 2000) were computed in IVAS using 1-2 at% for the respective element of interest. These
210 isoconcentration surfaces can be used to isolate each of the individual phases in the datasets. System
211 details, reconstruction parameters and other experimental conditions are summarized in (DR-1) following
212 the recommendation of (Blum et al., 2018a). $^{207}\text{Pb}/^{206}\text{Pb}$ ratios are calculated with corrections for
213 background levels and thermal tails. Pb in these samples can be detected at two different ionization states,
214 Pb^+ and Pb^{++} , and corrections are needed to account for the $^{206}\text{Pb}^{++}$ (103 Da) peak tail on the $^{207}\text{Pb}^{++}$
215 (103.5 Da) peak, but also P_3O_7^+ (205 Da) on $^{206}\text{Pb}^+$ (206 Da) and $^{206}\text{Pb}^+$ (206 Da) on $^{207}\text{Pb}^+$ (207 Da).

216 Measurement of isotopic ratios was undertaken using a range of different peak widths and bin sizes
217 following (Blum et al., 2018b).
218 From the prepared submicron monazites specimens (n=11), APT analysis resulted in three successful
219 runs, with a respective yield of 70, 45, and 9 million detected ions. The analysis of interfaces provides a
220 challenge as the interface is a weak spot in the material. The 3D reconstructions show that in all
221 specimens the monazite is encapsulated within rutile indicating that only the rim of the monazite
222 inclusions was analyzed. Monazites not encapsulated by rutile in the APT specimen likely delaminated
223 during the analysis. The datasets were evaluated by generating isoconcentration surfaces that subdivide
224 the specimens into three components - rutile, monazite and the interface. This allowed for the separate
225 interpretation of the mass spectra of each component.

226

227

Results

228 Petrography and analysis of the rutile host

229 The studied rutile grain occurs in a rock consisting of residual layers and leucosomes. Monazite
230 inclusions are found in two rutile grains within this thin section, as well as in the directly surrounding
231 biotite (Fig. 2b, c). No monazite inclusions were found in the garnet or mesoperthite. The rutile grain
232 studied is approximately 250 x 200 μm in size, with at least 46 sub-micron monazite inclusions visible at
233 the surface. EBSD mapping of individual monazite grains shows no evidence for a crystal-preferred
234 orientation of the inclusions in either the rutile or the surrounding biotite (fig 2d).
235 Several different microstructures are revealed by EBSD mapping. Twin boundaries are observed in two
236 orientations corresponding to a 65° rotation and belong to the {101} <100> twin relationship. In addition,
237 a network of low-angle boundaries (Fig. 2d, e) is visible. The low-angle boundaries have a misorientation
238 axis of (010), with the dispersion of 1-2°, corresponding to the {101}<201> slip system (Fig 2e). The
239 LAB are straight (LAB 1) or bent around the monazite inclusions (LAB 2).

240 **Atom probe observations of the rutile host.** Fe and Zr, which are common trace elements in
241 rutile, are detected within the mass spectra. U and Pb are not observed above the background (Fig 6). In

242 the 3D reconstruction, no heterogeneities can be observed within the rutile matrix this includes ranging
243 the theoretical U and Pb peaks. The rutile mass spectra are comparable to those obtained during previous
244 APT studies of rutile e.g., (Verberne et al., 2019; Verberne et al., 2023).

245

246 **Structural and chemical analysis of the monazite inclusions**

247 The examination of two TEM thin-foils, both incorporating a single monazite inclusion, indicates
248 variations in the texture and chemical associations of the outer rim of monazite in the two inclusions. In
249 one monazite, there is a clear core-rim structure (Fig 3). In contrast, the second monazite inclusion only
250 records the ‘core’ structure throughout the entire inclusion (Fig 4). In both monazites, the core has a high
251 density of clusters, while the rim structure has a lower density of clusters but this records a discrete
252 compositional variation (Fig 4, 5).

253 Trace elements observed within the monazite APT mass spectra (Fig. 6) include Ca^{++} (20 Da), Mg^+ (24
254 Da) and Li^+ (6 Da). ^{238}U was detected as $^{238}\text{UO}_2^{++}$ at 135 Da and no peak was present for $^{235}\text{UO}_2^{++}$. Pb in
255 monazite is observed with two different ionization states: $^{206}\text{Pb}^+$ and $^{207}\text{Pb}^+$ (206 and 207 Da) as well as
256 $^{206}\text{Pb}^{++}$ and $^{207}\text{Pb}^{++}$ (103 and 103.5 Da). $^{208}\text{Pb}^+$ and $^{208}\text{Pb}^{++}$ were not observed above the background.
257 Furthermore, no Th-bearing species were identified within the mass spectrum.

258 **Chemical heterogeneities in the monazite core.** The inner region of the monazite exhibits a
259 mottled texture identifiable in BF mode (Fig. 4). In one monazite, this texture is only present in the core-
260 region, whereas this texture is present throughout the second monazite grain imaged by TEM. The
261 mottled texture is due to the presence of dark (in BF mode) and bright (in DF mode) spots. These reflect
262 radiation damage and crystal defects as well as nanoclusters ± 5 nm in size. STEM-EDS revealed some of
263 these nanoclusters are composed of trace elements (U, Pb, and Ca) (Fig. 7a, b).

264 **Trace element distribution in the monazite rim.** This rim is only visible in one of the two
265 imaged monazites (Fig 4b). The monazite rim is ~ 100 nm wide. Here, the matrix of the monazite appears
266 pristine while hosting several nanoclusters 20-30 nm in diameter (Fig 5). Lattice imaging of the clusters
267 shows a continuous lattice consistent with monazite (Fig 5d-f). Here, we differentiate between a cluster

268 and inclusion based on similar properties as mentioned in Turuani et al., 2023. The presence of a grain
269 boundary is indicative of an inclusion while a continuous lattice/absence is an indication of a cluster. The
270 clusters in the monazite consist of two types of darker clusters as seen in the HAADF images (Fig 4b, 7c,
271 d) and sometimes encompass one or two bright spots. One type of darker cluster was shown to be rich in
272 Ca (Fig 7c). However, the second type of nanocluster did not yield an EDS response. This indicates the
273 major constituents of this cluster may have an atomic number that is too low for S/TEM-EDS ($Z < 6$).
274 Alternatively, this could point to a void and represent porosity within the cluster. The bright spots within
275 both these clusters yielded an increased concentration of Pb (Fig. 7c).

276 The 3D reconstruction of the monazite APT reveals two types of clusters in the monazite rim (Fig. 8). The
277 first type is dominated by Ca. These clusters have a diameter of 20-30 nm containing up to 10 at. % Ca
278 (Fig. 9). Ca-rich inclusions are often accompanied by an increase in Pb of up to 2-3 at. % as well as Y, up
279 to 1.2 at. %. However, not all Ca-rich inclusions show an increase in the Pb signal. The increase in Ca and
280 other trace elements is compensated by a depletion in REE and O with P remaining relatively constant. Pb
281 is spread over four peaks within the mass spectrum, resulting in insufficient counts for meaningful
282 determination of $^{207}\text{Pb}/^{206}\text{Pb}$ ratios after correcting for the static background and the tails of the preceding
283 peaks. Therefore, despite the presence of Pb within the inclusion, no isotopic composition could be
284 obtained.

285 The second inclusion type is enriched in Li^+ , with values up to 30 at. % Li (Fig. 9). These inclusions are
286 ~20 nm in diameter. Li-rich inclusions have Pb concentrations up to 0.6 at. % and therefore incorporate
287 less Pb than the Ca-rich inclusion, yet have elevated levels compared to the monazite matrix. Like the Ca-
288 rich inclusions, the increase of Li and traces is compensated by a depletion in REE and O with P
289 remaining at roughly 10 at. %. (Fig. 9). Pb isotopic compositions could not be obtained from the Li-
290 inclusions for similar reasons as with the Ca-inclusions.

291

292 **The rutile-monazite interface**

293 The interface between rutile and monazite inclusion is sharp and coherent (Fig 4a). Chemical mapping of
294 the interface by TEM shows the enriched zone in the interface between the rutile and monazite is
295 approximately 2 nm wide at the narrowest part of this boundary.

296 STEM-EDS analysis was applied to further characterize the different features present within the
297 monazite. The interface between rutile and monazite with core-rim texture yields 2-4 at. % Pb based on a
298 line profile across the interface (Fig. 7d) and with no signals above the detection limit for other trace
299 elements.

300 The mass spectra obtained from the interface show components of both the monazite and rutile mass
301 spectrum. The boundary is enriched in trace elements including U, Pb, Ca, Li, Mg, Si and Al. Ca and Pb
302 concentrations reach 2-3 at. % in all three specimens. The trace element distribution at the interface of
303 specimen 3 is heterogeneous, most clear for trace elements with a lower overall concentration, e.g. Si and
304 Al (Fig. 10), but similar distributions can be observed for Pb (Fig. 10).

305 The heterogeneous distribution of trace elements in the interface of specimen 3 is visualized in the
306 reconstruction and measured along a 1d profile (fig 10). The combined concentrations of Al + Si are
307 shown and a contrast in concentration is visible from the apex of the tip to the lower part of the
308 reconstructed volume. The transition from low to higher concentration is accompanied by linear features
309 enriched (labelled A, B, and C) in trace elements. Additionally, clusters of Mg are present at this
310 interface. These clusters contain up to 20 at. % Mg and have a Ca concentration of ~4 at% (Fig. 9), which
311 is higher than the average concentration within the interface. The width of the interface was estimated to
312 be ~8 nm based on the Pb concentration profile.

313

314 **Pb isotopic compositions of interfacially segregated Pb**

315 The mass spectra of the interface between rutile and monazite show ^{206}Pb and ^{207}Pb are present as both
316 Pb^+ and Pb^{++} . The 3D reconstruction shows the Pb is homogeneously distributed along the interface of
317 specimens 1 and 2 and heterogeneously in specimen 3 (fig 10). Pb isotopic signals were retrieved from

318 the interfaces of all three specimens. The background and thermal peak tails corrected $^{207}\text{Pb}/^{206}\text{Pb}$ ratios
319 are: 0.0900 ± 0.0198 (2σ) for specimen 1, 0.1016 ± 0.0224 (2σ) for specimen 2 and 0.1733 ± 0.0156 (2σ)
320 in specimen 3. Hence, the Pb isotopic signals and calculated $^{207}\text{Pb}/^{206}\text{Pb}$ ratios for specimens 1 and 2 are
321 identical within error. Specimen 3, however, yields a higher ratio. Calculated model ages for respectively
322 specimens 1, 2 and 3 are 1370 ± 420 Ma (2σ), 1590 ± 420 Ma (2σ) and 2580 ± 150 (2σ). U in the
323 interface was detected as $^{238}\text{UO}_2^{++}$ (135 Da). The concentration of U compared to Pb is very low with a
324 $^{238}\text{U}/^{206}\text{Pb}$ ratio of less than 0.04 in all three specimens. Therefore, it is expected the influence on the
325 $^{207}\text{Pb}/^{206}\text{Pb}$ ratio is negligible.

326 Discussion

327 Monazite inclusions in rutile

328 Inclusions typically occur either as a product of exsolution or incorporation during growth. In the rutile
329 studied here, exsolution is unlikely since monazite is present in different host phases (rutile and biotite).
330 The monazite lacks a crystal-preferred orientation that is expected when monazite forms via an exsolution
331 mechanism. Furthermore, P and REE concentrations in rutile are typically very low, with any significant
332 amounts of REE only reported in extra-terrestrial Niobian-rutile (>6 wt. % Nb_2O_5) (Hlava et al., 1972).
333 Therefore, the monazites are interpreted to be primary inclusions trapped within the rutile.
334 The encapsulation of monazite within rutile potentially shields the monazite from external processes
335 (Daneu et al., 2014; Hart et al., 2016; Hart et al., 2018). However, the measured Pb isotopic compositions
336 of monazite at inclusion interfaces are different. The rutile host is characterized by the presence of several
337 twin boundaries and LABs, and these could aid in the diffusion of elements in rutile (Verberne et al.,
338 2022). As shown, some monazite inclusions are in direct contact with these LAB's (Fig 2). Potentially,
339 these inclusions would no longer be fully shielded with trace elements able to migrate out of the rutile-
340 monazite interface. Depending on the timing, inclusions within a single rutile grain could record different
341 isotopic signatures.

342

343 Formation of trace-element enriched nanoscale clusters

344 **Nanoclusters in the monazite core.** The two types of monazite inclusions are: one with a
345 mottled texture due to radiation damage and one with a distinct core-rim structure (Fig. 4, 5). S/TEM EDS
346 shows that some of the dark and bright spots that lead to the mottled texture are enrichment in Ca, U, and
347 Pb (Fig. 7). The texture and formation of trace element enriched clusters are explained via entrapment in
348 defects generated via radiation damage (Seydoux-Guillaume et al., 2018; Seydoux-Guillaume, 2019). The
349 damaged regions stand out due to structural difference with the host monazite while the bright spot
350 (HAADF) stands out due to being enriched in trace elements, creating this mottled texture as previously
351 observed by others (Seydoux-Guillaume et al., 2019; Turuani et al., 2022; 2023).
352 Clusters within a matrix typically represent a low-energy configuration and are preferable sites for trace
353 elements compared to the crystal lattice of the host. With the rutile present during UHT metamorphism at
354 2580 Ma (Clark et al., 2019), the temperature conditions are met for Pb diffusion into clusters (as
355 observed in Seydoux-Guillaume et al., 2019).
356 The mottled texture therefore is the result of an interplay between the formation and healing of radiation
357 damage sites, and the capture of trace elements in these sites during high-temperature events, which allow
358 for more rapid diffusion of elements.

359 **Nanoscale clusters in the monazite rim.** The rim hosts Ca-rich and Li-rich nanoscale clusters.
360 Given that the lattice of the Ca-rich cluster is the same as the host monazite (Fig 5) with relatively high
361 REE concentrations these clusters are interpreted to be a solid-solution within monazite rather than
362 represent a separate phase.
363 The Li-rich nanoscale cluster has low REE concentrations (<4 at. %) with Li concentrations between 24
364 to 30 at% and P concentrations of ± 10 at. % (fig 9). In the literature, we found no reference for the
365 formation of Li-phosphate in monazite. Nevertheless, the composition of the Li-rich cluster may indicate
366 the formation of a separate Li-phosphate phase.
367 The Mg-rich clusters are only present within the interface and consist of 20 at. % Mg, 10 at% P and, 4
368 at% Ca while showing a depletion in REE concentrations compared to the monazite host. As is the case

369 for the Li-phosphate cluster, the high concentration of Mg suggests the formation of a new phase, via a
370 similar mechanism to the formation of the Ca- and Li-phosphate clusters.

371 **Pb-rich nanoclusters.** Bright spots in HAADF mode associated with the Ca- and Li-rich clusters
372 are rich in Pb based on STEM-EDS (Fig 7). Furthermore, APT based proximity histogram of the Ca- and
373 Li-rich cluster also shows elevated values of Pb. The lack of a ^{204}Pb and ^{208}Pb signal indicated there is no
374 to a negligible ^{232}Th or common Pb component. Potentially, these Pb-rich nanoclusters reflect the same
375 clusters observed within the monazite core. The Pb nanocrystals within the Ca- and Li-phosphate cluster
376 have the same general appearance as Pb present in Si-rich cluster recently described in zircon by (Kusiak.,
377 2015, Kusiak et al., 2019; Lyon et al., 2019), called nanospheres and monazite (Turuani et al., 2022;
378 Turuani et al., 2023) referred to as nanocrystals. Considering we observe a similar presence of Pb in the
379 core that is related associated with U decay, it is unlikely these are primary Pb phases. Furthermore, since
380 we lack data on the structure or valence state of the Pb and associate the Pb with U decay these Pb-rich
381 particles do not fall within the definition of metallic Pb nanospheres (Kusiak 2015, Kusiak et al., 2023).
382 These Pb-rich nanocrystals are intimately linked to the Ca, and Li-rich phosphate clusters, which raises
383 questions about the timing of their formation. One possibility is concurrent formation, with Ca, Li, and Pb
384 migrating at the same time. However, this seems unlikely, as the Pb-rich nanocrystals in the rim are not
385 observed outside of the Ca and Li-rich clusters. If the Pb-rich nanocrystals are indeed the same as those in
386 the core this indicates that during the formation of the rim and the Ca- and Li-rich clusters, the Pb
387 nanocrystals become trapped, preventing small volumes of Pb migrating towards the interface.

388

389 **A Pb mobility model**

390 The texture and compositional variations of the submicron monazite inclusions in rutile provide insight
391 into the tectonothermal history of the studied sample. First, the mottled texture of the monazite (core) is
392 ascribed to both a combination of the presence of radiation damage and the segregation of trace elements
393 into those radiation defects leading to the formation of nanoscale clusters (Seydoux-Guillaume et al.,
394 2019; Seydoux-Guillaume et al., 2002; 2004; 2018; Turuani et al., 2023). Depending on the time of

395 monazite growth and the diffusion of trace elements into defects the clusters rich in Pb can be dated if
396 enough time has passed to accumulate radiogenic Pb. This diffusion process is the first stage of trace
397 element mobility.

398 Second, in some of the monazite crystals, the rim yields a different distribution of clusters. The texture
399 indicates partial recrystallization of the monazite rim and encompasses Ca- and Li-phosphate clusters as
400 well as Pb-rich nanocrystals. The presence of a recrystallized rim in only some of the monazite crystals
401 shows that the reaction is localized.

402 Rutile and other repository minerals preserve inclusions due to their low compressibility (Mosenfelder et
403 al., 2005). When exhumed, rutile exerts an overpressure on inclusions (Hart et al., 2016). Differences in
404 compressibility between the rutile host and the inclusions will result in a stress acting upon the monazite
405 (Huang et al., 2010; Zaffiro et al., 2019). Therefore, one explanation could be that the formation of the
406 rim in the monazite is the result of localized stress-induced recrystallization. Alternatively, the presence
407 of Li and to a lesser extent Mg could indicate infiltration of a fluid. This too could cause recrystallization
408 of the monazite rim. In general, rutile is resistant to the infiltration of fluids (Triebold et al., 2007; Zack et
409 al., 2004b). As argued in recent studies on the nanoscale and microstructures in rutile (Verberne et al.,
410 2022; Verberne et al., 2023) the presence of twin boundaries and low-angle boundaries could allow for
411 the transport of trace elements. The same structures potentially facilitate fluid infiltration as well.
412 However, we observe Li primarily in clusters and only minor concentration in the interface itself.
413 Therefore, a localized stress-induced recrystallization appears to fit better with these observations.

414 **Low Th monazite.** The absence or very low concentration of Th in monazite is an odd
415 observation as ThO₂ concentration in monazite are mostly reported to be between 1 – 12 wt. % (e.g.,
416 Janots et al., 2012, Seydoux-Guillaume et al., 2018). However, a few studies report on monazite with less
417 than 1 wt. % (e.g., Demartin et al., 1991, and Mannucci et al., 1986). Yet, these studies still show
418 measureable amounts of Th. Closest to our observations are a study by Janots et al., 2008. X-ray maps of
419 monazite show a monazite with high Th in the core and a recrystallized rim with low, but not absent, Th
420 concentrations. Although the analyzed rim is 10 μm wide compared to our 100 nm wide rim, and thus

421 five orders of magnitude larger. The same underlying principle, recrystallization, would offer an
422 explanation. The recrystallization would have occurred at a low-metamorphic grade Poitrasson et al.,
423 1996, Janots et al. 2008 with Th concentration being positively related to the metamorphic grade
424 (Seydoux 2002). This supports our hypothesis of localized stress-induced recrystallization.

425

426 **Trace element systematics at the rutile-monazite interface**

427 The distribution of trace elements at the interface in specimen 3 is heterogeneous (Fig. 10). In Figure 10
428 lower concentrations of trace elements are observed in the upper part of the boundary compared to the
429 lower part of the boundary. The transition occurs across a domain with linear features rich in trace-
430 elements. These linear features are an expression of dislocations at the interface (Dubosq et al., 2019;
431 Fougrouse et al., 2021b; Verberne et al., 2022; Verberne et al., 2023). Heterogeneities at phase
432 boundaries are commonly observed and well-studied within the material sciences (Larere et al., 1988;
433 Sutton et al., 1995; Suzuki et al., 1991; Suzuki et al., 1981). There is extensive evidence that the degree of
434 trace element segregation has an inverse relationship with the interplanar spacing (Sutton et al., 1995). In
435 iron-alloys, it was found that P migrates to boundaries when this spacing decreases (Suzuki et al., 1981).
436 Also, P and S have different degrees of segregation towards boundaries with different orientations in Ni
437 (Larere et al., 1988) and it is apparent that trace elements segregate in different amounts in function to
438 disorientation in olivine (Tacchetto et al., 2021). The segregation into boundaries for ionic systems is
439 more complex since charge compensation needs to be considered (Kliewer, 1965; Sutton et al., 1995). In
440 TiO₂ Ikeda and Chiang (1993); Ikeda et al. (1993) demonstrated differences in segregation based on the
441 valence states of trace elements e.g., Al³⁺ and Nb⁵⁺, boundary type, and the temperature. Here, the
442 curvature of the interface results in a continuous change of opposing crystal faces. Consequently, the
443 spacing between the two phases will vary. This change is reflected in the segregation of trace elements,
444 causing differences in localized trace element concentrations.

445

446 **Implications for geochronology: a geological history from submicron monazite inclusions**

447 The presence of inclusions with a radiogenic component different from the host mineral can result in a
448 mixed chemical and isotopic system (Fougerouse et al., 2018; Turuani et al., 2022). In this study, Pb is
449 found to have segregated in multiple ways that might have occurred at different moments in time. The
450 clusters still preserved within the monazite core are the result of the interplay between the formation of
451 radiation damage, alpha-healing (Seydoux-Guillaume et al., 2018) and volume diffusion (Seydoux-
452 Guillaume et al., 2019). While recrystallization led to the segregation of Pb from the monazite rim into
453 the rutile-monazite interface. We observe that the recrystallization of the monazite rim is local and not
454 pervasive throughout the rutile grain. Also, we see Pb-enriched nanocrystals trapped within clusters in the
455 recrystallized rim. This localized recrystallization combined with trapping of Pb-enriched nanocrystals
456 would suggest the Pb-enriched features formation and recrystallization occurred in two stages, though
457 without Pb isotopic information from these nanocrystals and the clusters in the core the timing of these
458 stages remains elusive.

459 Regardless, the mobility results in several Pb reservoirs with, potentially unique, Pb isotopic signatures
460 that record the timing of their formation and the decoupling of the U-Pb system. In ideal situations, this
461 Pb signature can be independently studied (Fougerouse et al., 2018; Peterman et al., 2016; Seydoux-
462 Guillaume et al., 2019; Turuani et al., 2022; Valley et al., 2014; Verberne et al., 2020).

463 In this study, the only reservoir with a significant amount of Pb, measurable by APT, is the interface
464 between rutile and monazite. Age determination in APT data is based on analyzing the $^{207}\text{Pb}/^{206}\text{Pb}$ ratios
465 to derive the timing of Pb segregation in contrast to a direct age. The assumption is that this ratio is
466 preserved after segregation. This also means that the presence of U or Th compared to total Pb needs to be
467 minimal. Important to note is that any additional influx of radiogenic Pb derived from the same source
468 material will obscure the determination of the first event. However, such event could possibly also record
469 the latter, preserving both (Taylor et al., 2023). In our results, the absence of ^{204}Pb and ^{208}Pb supports the
470 assumption that no common Pb is present and thus all Pb is of radiogenic origin (Peterman et al., 2016;
471 Valley et al., 2014; Verberne et al., 2020). The U/Pb ratios in the boundary are <0.05 . This indicates large

472 amounts of Pb are unsupported and that there is negligible production of radiogenic Pb via continued U
473 decay after the Pb migration into the rim.
474 The $^{207}\text{Pb}/^{206}\text{Pb}$ ratio varied between the interfaces in the three APT specimens. The obtained ratios are
475 0.0900 ± 0.0198 (2σ) for specimen 1, 0.1016 ± 0.0224 (2σ) for specimen 2 and 0.1733 ± 0.0156 (2σ) in
476 specimen 3. Specimens 1 and 2 have $^{207}\text{Pb}/^{206}\text{Pb}$ ratios with large uncertainties, however, specimen 3
477 yields a more precise ratio which can be evaluated. Let us reiterate that the monazite is enclosed within
478 rutile. Radiogenic is Pb produced by radioactive decay of U and the Pb will be retained in monazite. Later
479 metamorphic and recrystallization events will mobilize the Pb, isolating it from U. Thus the $^{207}\text{Pb}/^{206}\text{Pb}$
480 ratio at the time of Pb segregation is preserved. The rocks in the Napier Complex have a complex history.
481 The oldest potential time of formation is suggested by Black et al., 1984, with monazite crystallization
482 around 3070 Ma. If we assume an initial crystallization age of 3070 Ma, the current $^{207}\text{Pb}/^{206}\text{Pb}$ of 0.1733
483 ± 0.0156 (2σ) indicates Pb production for a duration of 2580 Myr before segregation. This would
484 correspond with the decoupling of U-Pb occurring around 550 Ma, which would match the late
485 hydrothermal event in the area (Black, 1983 #18). This shows that analyses of grain boundaries, which
486 are often avoided during geochronological studies can provide additional information about the geological
487 history.

488

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504 **Conflict of Interest Declarations**

505 The authors declare there are no conflicts of interest.

506

507 **References**

- 508 Bachmann, F., Hielscher, R., and Schaeben, H., Texture analysis with MTEX—free and open source
509 software toolbox, *in* Proceedings Solid State Phenomena2010, Volume 160, Trans Tech Publ, p.
510 63-68.
- 511 Bingen, B., and Van Breemen, O., 1998, U-Pb monazite ages in amphibolite-to granulite-facies
512 orthogneiss reflect hydrous mineral breakdown reactions: Sveconorwegian Province of SW
513 Norway: *Contributions to Mineralogy and Petrology*, v. 132, no. 4, p. 336-353.
- 514 Black, L., James, P., and Harley, S., 1983, The geochronology, structure and metamorphism of early
515 Archaean rocks at Fyfe Hills, Enderby Land, Antarctica: *Precambrian Research*, v. 21, no. 3-4, p.
516 197-222.
- 517 Black, L. P., Fitzgerald, J. D., and Harley, S. L., 1984, Pb isotopic composition, colour, and
518 microstructure of monazites from a polymetamorphic rock in Antarctica: *Contributions to*
519 *Mineralogy and Petrology*, v. 85, no. 2, p. 141-148.

- 520 Black, L., et al. (1986). "Four zircon ages from one rock: the history of a 3930 Ma-old granulite from
521 Mount Sones, Enderby Land, Antarctica." *Contributions to Mineralogy and Petrology* 94(4): 427-
522 437.
- 523 Blum, T. B., Darling, J. R., Kelly, T. F., Larson, D. J., Moser, D. E., Perez-Huerta, A., Prosa, T. J.,
524 Reddy, S. M., Reinhard, D. A., and Saxey, D. W., 2018a, Best practices for reporting atom probe
525 analysis of geological materials: *Microstructural Geochronology: Planetary Records Down to*
526 *Atom Scale*, v. 232, p. 369-373.
- 527 Blum, T. B., Reinhard, D. A., Chen, Y., Prosa, T. J., Larson, D. J., and Valley, J. W., 2018b, Uncertainty
528 and Sensitivity Analysis for Spatial and Spectral Processing of Pb Isotopes in Zircon by Atom
529 Probe Tomography: *Microstructural Geochronology: Planetary Records Down to Atom Scale*, p.
530 327-350.
- 531 Carson, C. J., Ague, J. J., and Coath, C. D., 2002, U–Pb geochronology from Tonagh Island, East
532 Antarctica: implications for the timing of ultra-high temperature metamorphism of the Napier
533 Complex: *Precambrian Research*, v. 116, no. 3, p. 237-263.
- 534 Clark, C., Taylor, R. J., Johnson, T. E., Harley, S. L., Fitzsimons, I. C., and Oliver, L., 2019, Testing the
535 fidelity of thermometers at ultrahigh temperatures: *Journal of Metamorphic Geology*.
- 536 Clark, C., Taylor, R. J., Kylander-Clark, A. R., and Hacker, B. R., 2018, Prolonged (> 100 Ma) ultrahigh
537 temperature metamorphism in the Napier Complex, East Antarctica: A petrochronological
538 investigation of Earth's hottest crust: *Journal of Metamorphic Geology*, v. 36, no. 9, p. 1117-
539 1139.
- 540 Cramer, F., 2018, Scientific colour maps: *Zenodo*, v. 10.
- 541 Dache, F., Simons, P., and Roy, R., 1968, Pressure-temperature studies of anatase, brookite, rutile and
542 TiO₂-II: *Am Mineral*, v. 53, p. 1929-1939.
- 543 Daneu, N., Recnik, A., and Mader, W., 2014, Atomic structure and formation mechanism of (101) rutile
544 twins from Diamantina (Brazil): *American Mineralogist*, v. 99, no. 4, p. 612-624.

- 545 Demartin, F., Pilati, T., Diella, V., Donzelli, S., & Gramaccioli, C. M. (1991). Alpine monazite; further
546 data. *The Canadian Mineralogist*, 29(1), 61-67.
- 547 Dubosq, R., Rogowitz, A., Schweinar, K., Gault, B., and Schneider, D. A., 2019, A 2D and 3D
548 nanostructural study of naturally deformed pyrite: assessing the links between trace element
549 mobility and defect structures: *Contributions to Mineralogy and Petrology*, v. 174, no. 9, p. 72.
- 550 Erickson, T., Pearce, M., Taylor, R., Timms, N. E., Clark, C., Reddy, S., and Buick, I., 2015, Deformed
551 monazite yields high-temperature tectonic ages: *Geology*, v. 43, no. 5, p. 383-386.
- 552 Fougrouse, D., Reddy, S., Seydoux-Guillaume, A.-M., Kirkland, C., Erickson, T., Saxey, D., Rickard,
553 W., Jacob, D., Leroux, H., and Clark, C., 2021a, Mechanical twinning of monazite expels
554 radiogenic lead: *Geology*, v. 49, no. 4, p. 417-421.
- 555 Fougrouse, D., Reddy, S. M., Aylmore, M., Yang, L., Guagliardo, P., Saxey, D. W., Rickard, W. D., and
556 Timms, N., 2021b, A new kind of invisible gold in pyrite hosted in deformation-related
557 dislocations: *Geology*, v. 49, no. 10, p. 1225-1229.
- 558 Fougrouse, D., Reddy, S. M., Saxey, D. W., Erickson, T. M., Kirkland, C. L., Rickard, W. D. A.,
559 Seydoux-Guillaume, A. M., Clark, C., and Buick, I. S., 2018, Nanoscale distribution of Pb in
560 monazite revealed by atom probe microscopy: *Chemical Geology*, v. 479, p. 251-258.
- 561 Fougrouse, D., Saxey, D. W., Rickard, W. D., Reddy, S. M., and Verberne, R., 2021c, Standardizing
562 Spatial Reconstruction Parameters for the Atom Probe Analysis of Common Minerals:
563 *Microscopy and Microanalysis*, p. 1-10.
- 564 Grand'Homme, A., Janots, E., Seydoux-Guillaume, A.-M., Guillaume, D., Bosse, V., and Magnin, V.,
565 2016, Partial resetting of the U-Th-Pb systems in experimentally altered monazite: Nanoscale
566 evidence of incomplete replacement: *Geology*, v. 44, no. 6, p. 431-434.
- 567 Harley, S. L., 2016, A matter of time: the importance of the duration of UHT metamorphism: *Journal of*
568 *Mineralogical and Petrological Sciences*, v. 111, no. 2, p. 50-72.

- 569 Harley, S. L., and Motoyoshi, Y., 2000, Al zoning in orthopyroxene in a sapphirine quartzite: evidence
570 for > 1120 C UHT metamorphism in the Napier Complex, Antarctica, and implications for the
571 entropy of sapphirine: *Contributions to Mineralogy and Petrology*, v. 138, no. 4, p. 293-307.
- 572 Harley, S. L., & Black, L. P. (1987). The Archaean geological evolution of Enderby Land, Antarctica. In
573 R. G. Park & J. Tarney (Eds.), *Evolution of the Lewisian and comparable Precambrian high grade*
574 *terrains* (27th edn, pp. 285–296). Oxford: Geological Society of London.
- 575 Hart, E., Storey, C., Bruand, E., Schertl, H. P., and Alexander, B. D., 2016, Mineral inclusions in rutile: A
576 novel recorder of HP-UHP metamorphism: *Earth and Planetary Science Letters*, v. 446, p. 137-
577 148.
- 578 Hart, E., Storey, C., Harley, S. L., and Fowler, M., 2018, A window into the lower crust: Trace element
579 systematics and the occurrence of inclusions/intergrowths in granulite-facies rutile: *Gondwana*
580 *Research*, v. 59, p. 76-86.
- 581 Hellman, O. C., Vandenbroucke, J. A., Rusing, J., Isheim, D., and Seidman, D. N., 2000, Analysis of
582 Three-dimensional Atom probe Data by the Proximity Histogram: *Microsc Microanal*, v. 6, no. 5,
583 p. 437-444.
- 584 Hlava, P., Prinz, M., and Keil, K., 1972, Niobian rutile in an Apollo 14 KREEP fragment: *Meteoritics*, v.
585 7, no. 4, p. 479-486.
- 586 Hokada, T., 2001, Feldspar thermometry in ultrahigh-temperature metamorphic rocks: Evidence of crustal
587 metamorphism attaining ~ 1100 C in the Archean Napier Complex, East Antarctica: *American*
588 *Mineralogist*, v. 86, no. 7-8, p. 932-938.
- 589 Huang, T., Lee, J.-S., Kung, J., and Lin, C.-M., 2010, Study of monazite under high pressure: *Solid State*
590 *Communications*, v. 150, no. 37-38, p. 1845-1850.
- 591 Ikeda, J. A. S., and Chiang, Y. M., 1993, Space charge segregation at grain boundaries in titanium
592 dioxide: I, relationship between lattice defect chemistry and space charge potential: *Journal of the*
593 *American Ceramic Society*, v. 76, no. 10, p. 2437-2446.

- 594 Ikeda, J. A. S., Chiang, Y. M., Garratt-Reed, A. J., and Sande, J. B. V., 1993, Space charge segregation at
595 grain boundaries in titanium dioxide: II, model experiments: *Journal of the American Ceramic*
596 *Society*, v. 76, no. 10, p. 2447-2459.
- 597 James, P., and Black, L., 1981, *Archaean Geology*.
- 598 Janots, E., Berger, A., Gnos, E., Whitehouse, M., Lewin, E., & Pettke, T. (2012). Constraints on fluid
599 evolution during metamorphism from U–Th–Pb systematics in Alpine hydrothermal monazite.
600 *Chemical Geology*, 326, 61-71.
- 601 Janots, E., Engi, M., Berger, A., Allaz, J., Schwarz, J. O., & Spandler, C. (2008). Prograde metamorphic
602 sequence of REE minerals in pelitic rocks of the Central Alps: implications for allanite–
603 monazite–xenotime phase relations from 250 to 610 C. *Journal of Metamorphic Geology*, 26(5),
604 509-526.
- 605 Kliewer, K., 1965, Space charge in ionic crystals. II. The electron affinity and impurity accumulation:
606 *Physical Review*, v. 140, no. 4A, p. A1241.
- 607 Kooijman, E., Mezger, K., and Berndt, J., 2010, Constraints on the U-Pb systematics of metamorphic
608 rutile from in situ LA-ICP-MS analysis: *Earth and Planetary Science Letters*, v. 293, no. 3-4, p.
609 321-330.
- 610 Kusiak, M., Kovaleva, E., Wirth, R., Klötzli, U., Dunkley, D., Yi, K., and Lee, S., 2019, Lead oxide
611 nanospheres in seismically deformed zircon grains: *Geochimica et Cosmochimica Acta*, v. 262, p.
612 20-30.
- 613 Larere, A., Guillope, M., and Masuda-Jindo, K., 1988, Segregation energies of phosphorus and sulfur in
614 various relaxed structures of grain boundaries in Ni by tight-binding approximation: *Le Journal*
615 *de Physique Colloques*, v. 49, no. C5, p. C5-447-C445-450.
- 616 Lloyd, G. E., Farmer, A. B., and Mainprice, D., 1997, Misorientation analysis and the formation and
617 orientation of subgrain and grain boundaries: *Tectonophysics*, v. 279, no. 1-4, p. 55-78.

- 618 Lyon, I., Kusiak, M., Wirth, R., Whitehouse, M., Dunkley, D., Wilde, S., Schaumlöffel, D., Malherbe, J.,
619 and Moore, K., 2019, Pb nanospheres in ancient zircon yield model ages for zircon formation and
620 Pb mobilization: *Scientific Reports*, v. 9.
- 621 Mannucci, G., Diella, V., Gramaccioli, C. M., & Pilati, T. (1986). A comparative study of some
622 pegmatitic and fissure monazite from the Alps. *The Canadian Mineralogist*, 24(3), 469-474.
- 623 Meinhold, G., 2010, Rutile and its applications in earth sciences: *Earth-Science Reviews*, v. 102, no. 1-2,
624 p. 1-28.
- 625 Mitchell, R. J., and Harley, S. L., 2017, Zr-in-rutile resetting in aluminosilicate bearing ultra-high
626 temperature granulites: Refining the record of cooling and hydration in the Napier Complex,
627 Antarctica: *Lithos*, v. 272, p. 128-146.
- 628 Mosenfelder, J. L., Schertl, H.-P., Smyth, J. R., and Liou, J. G., 2005, Factors in the preservation of
629 coesite: The importance of fluid infiltration: *American Mineralogist*, v. 90, no. 5-6, p. 779-789.
- 630 Ni, Y., Hughes, J. M., and Mariano, A. N., 1995, Crystal chemistry of the monazite and xenotime
631 structures: *American Mineralogist*, v. 80, no. 1-2, p. 21-26.
- 632 Pape, J., Mezger, K., and Robyr, M., 2016, A systematic evaluation of the Zr-in-rutile thermometer in
633 ultra-high temperature (UHT) rocks: *Contributions to Mineralogy and Petrology*, v. 171, no. 5.
- 634 Peterman, E. M., Reddy, S. M., Saxey, D. W., Snoeyenbos, D. R., Rickard, W. D., Fougereuse, D., and
635 Kylander-Clark, A. R., 2016, Nanogeochronology of discordant zircon measured by atom probe
636 microscopy of Pb-enriched dislocation loops: *Sci Adv*, v. 2, no. 9, p. e1601318.
- 637 Poitrasson, F., Chenery, S., & Bland, D. J. (1996). Contrasted monazite hydrothermal alteration
638 mechanisms and their geochemical implications. *Earth and Planetary Science Letters*, 145(1-4),
639 79-96.
- 640 Prior, D. J., Wheeler, J., Peruzzo, L., Spiess, R., and Storey, C., 2002, Some garnet microstructures: an
641 illustration of the potential of orientation maps and misorientation analysis in microstructural
642 studies: *Journal of Structural Geology*, v. 24, no. 6-7, p. 999-1011.

- 643 Reddy, S. M., Saxey, D. W., Rickard, W. D., Fougereuse, D., Montalvo, S. D., Verberne, R., and Van
644 Riessen, A., 2020, Atom Probe Tomography: Development and Application to the Geosciences:
645 Geostandards and Geoanalytical Research, v. 44, no. 1, p. 5-50.
- 646 Reddy, S. M., Timms, N. E., Pantleon, W., and Trimby, P., 2007, Quantitative characterization of plastic
647 deformation of zircon and geological implications: Contributions to Mineralogy and Petrology, v.
648 153, no. 6, p. 625-645.
- 649 Rickard, W. D., Reddy, S. M., Saxey, D. W., Fougereuse, D., Timms, N. E., Daly, L., Peterman, E.,
650 Cavosie, A. J., and Jourdan, F., 2020, Novel Applications of FIB-SEM-Based ToF-SIMS in Atom
651 Probe Tomography Workflows: Microscopy and Microanalysis, p. 1-8.
- 652 Seydoux-Guillaume, A.-M., Fougereuse, D., Laurent, A., Gardés, E., Reddy, S., and Saxey, D., 2019,
653 Nanoscale resetting of the Th/Pb system in an isotopically-closed monazite grain: A combined
654 atom probe and transmission electron microscopy study: Geoscience Frontiers, v. 10, no. 1, p. 65-
655 76.
- 656 Seydoux-Guillaume A.-M., Deschanel X., Baumier C., Neumeier S., Weber W.J., and Peugot S. (2018).
657 Why natural monazite never becomes amorphous: experimental evidence for alpha self-healing.
658 *American Mineralogist*, 103, 824-827. doi.org/10.2138/am-2018-6447
- 659 Seydoux-Guillaume A.M., Wirth R., Deutsch A. and Schärer U. (2004). Microstructure of 24 – 1928 Ma
660 concordant monazites: implications for geochronology and nuclear waste deposits. *Geoch.*
661 *Cosmochim. Acta*, Vol. 68, 2517-2527.
- 662 Seydoux-Guillaume, A.-M., Montel, J.-M., Bingen, B., Bosse, V., De Parseval, P., Paquette, J.-L., Janots,
663 E., and Wirth, R., 2012, Low-temperature alteration of monazite: Fluid mediated coupled
664 dissolution–precipitation, irradiation damage, and disturbance of the U–Pb and Th–Pb
665 chronometers: Chemical Geology, v. 330, p. 140-158.
- 666 Seydoux-Guillaume A.M., Goncalves P., Wirth R. and Deutsch A. (2003). TEM study of polyphasic and
667 discordant monazites: site specific specimen preparation using the Focused Ion Beam technique.
668 *Geology*, Vol. 31, 973-976.

- 669 Seydoux-Guillaume A.M., Wirth R., Nasdala L., Gottschalk M., Montel J.M. and Heinrich W. (2002b).
670 An XRD, TEM and Raman study of experimentally annealed natural monazite. *Phys. Chem.*
671 *Minerals*, Vol. 29, 240-253.
- 672 Seydoux-Guillaume, A.-M., Paquette, J.-L., Wiedenbeck, M., Montel, J.-M., and Heinrich, W., 2002,
673 Experimental resetting of the U–Th–Pb systems in monazite: *Chemical geology*, v. 191, no. 1-3,
674 p. 165-181.
- 675 Seydoux-Guillaume A.M., Wirth R., Heinrich W. and Montel J.M. (2002). Experimental determination
676 of the Th partitioning between monazite and xenotime using Analytical Electron Microscopy and
677 X-ray Diffraction Rietveld analysis. *Eur. J. Mineral.*, Vol. 14, 869 — 878.
- 678 Sheraton, J. W., Thomson, J. W., and Collerson, K. D., Mafic dyke swarms of Antarctica 1987,
679 Geological Association of Canada.
- 680 Smye, A. J., and Stockli, D. F., 2014, Rutile U-Pb age depth profiling: A continuous record of
681 lithospheric thermal evolution: *Earth and Planetary Science Letters*, v. 408, p. 171-182.
- 682 Sutton, A. P., Balluffi, R. W., and Sutton, A., 1995, Interfaces in crystalline materials.
- 683 Suzuki, K., Ichihara, M., and Takeuchi, S., 1991, High-resolution electron microscopy of lattice defects in
684 TiO₂ and SnO₂: *Philosophical Magazine A*, v. 63, no. 4, p. 657-665.
- 685 Suzuki, S., Abiko, K., and Kimura, H., 1981, Phosphorus segregation related to the grain boundary
686 structure in an Fe–P alloy: *Scripta Metallurgica*, v. 15, no. 10, p. 1139-1143.
- 687 Swope, R. J., Smyth, J. R., and Larson, A. C., 1995, H in rutile-type compounds: I. Single-crystal neutron
688 and X-ray diffraction study of H in rutile: *American Mineralogist*, v. 80, no. 5-6, p. 448-453.
- 689 Tacchetto, T., Reddy, S. M., Saxey, D. W., Fougereuse, D., Rickard, W. D. A., and Clark, C., 2021,
690 Disorientation control on trace element segregation in fluid-affected low-angle boundaries in
691 olivine: *Contributions to Mineralogy and Petrology*, v. 176, no. 7, p. 59.
- 692 Taylor, R. J., Reddy, S. M., Saxey, D. W., Rickard, W. D., Tang, F., Borlina, C. S., Fu, R. R., Weiss, B.
693 P., Bagot, P., and Williams, H. M., 2023, Direct age constraints on the magnetism of Jack Hills
694 zircon: *Science Advances*, v. 9, no. 1, p. eadd1511.

- 695 Thompson, K., Lawrence, D., Larson, D. J., Olson, J. D., Kelly, T. F., and Gorman, B., 2007, In situ site-
696 specific specimen preparation for atom probe tomography: Ultramicroscopy, v. 107, no. 2-3, p.
697 131-139.
- 698 Triebold, S., von Eynatten, H., Luvizotto, G. L., and Zack, T., 2007, Deducing source rock lithology from
699 detrital rutile geochemistry: An example from the Erzgebirge, Germany: Chemical Geology, v.
700 244, no. 3-4, p. 421-436.
- 701 Turuani, M. J., Laurent, A. T., Seydoux-Guillaume, A. M., Fougereuse, D., Saxey, D., Reddy, S. M.,
702 Harley, S. L., Reynaud, S., and Rickard, W. D. A., 2022, Partial retention of radiogenic Pb in
703 galena nanocrystals explains discordance in monazite from Napier Complex (Antarctica): Earth
704 and Planetary Science Letters, v. 588.
- 705 Turuani, M. J., Seydoux-Guillaume, A.-M., Laurent, A. T., Reddy, S., Harley, S., Fougereuse, D., Saxey,
706 D., Gouy, S., Parseval, D., and Reynaud, S., 2023, Nanoscale features revealed by a multiscale
707 characterization of discordant monazite highlight mobility mechanisms of Th and Pb.
- 708 Valley, J. W., Cavosie, A. J., Ushikubo, T., Reinhard, D. A., Lawrence, D. F., Larson, D. J., Clifton, P.
709 H., Kelly, T. F., Wilde, S. A., and Moser, D. E., 2014, Hadean age for a post-magma-ocean
710 zircon confirmed by atom-probe tomography: Nature Geoscience, v. 7, no. 3, p. 219-223.
- 711 Verberne, R., Reddy, S., Saxey, D., Fougereuse, D., Rickard, W., Plavsa, D., Agangi, A., and Kylander-
712 Clark, A., 2020, The geochemical and geochronological implications of nanoscale trace-element
713 clusters in rutile: Geology.
- 714 Verberne, R., Reddy, S. M., Saxey, D. W., Fougereuse, D., Rickard, W. D., Quadir, Z., Evans, N. J., and
715 Clark, C., 2022, Dislocations in minerals: Fast-diffusion pathways or trace-element traps?: Earth
716 and Planetary Science Letters, v. 584, p. 117517.
- 717 Verberne, R., Saxey, D. W., Reddy, S. M., Rickard, W. D., Fougereuse, D., and Clark, C., 2019, Analysis
718 of Natural Rutile (TiO₂) by Laser-assisted Atom Probe Tomography: Microscopy and
719 Microanalysis, p. 1-8.

- 720 Verberne, R., van Schrojenstein Lantman, H. W., Reddy, S. M., Alvaro, M., Wallis, D., Fougrouse, D.,
721 Langone, A., Saxey, D. W., and Rickard, W. D., 2023, Trace-element heterogeneity in rutile
722 linked to dislocation structures: Implications for Zr-in-rutile geothermometry: Journal of
723 Metamorphic Geology, v. 41, no. 1, p. 3-24.
- 724 Watson, E. B., Wark, D. A., and Thomas, J. B., 2006, Crystallization thermometers for zircon and rutile:
725 Contributions to Mineralogy and Petrology, v. 151, no. 4, p. 413-433.
- 726 Zack, T., Moraes, R., and Kronz, A., 2004a, Temperature dependence of Zr in rutile: empirical calibration
727 of a rutile thermometer: Contributions to Mineralogy and Petrology, v. 148, no. 4, p. 471-488.
- 728 Zack, T., von Eynatten, H., and Kronz, A., 2004b, Rutile geochemistry and its potential use in
729 quantitative provenance studies: Sedimentary Geology, v. 171, no. 1-4, p. 37-58.
- 730 Zaffiro, G., Angel, R. J., and Alvaro, M., 2019, Constraints on the Equations of State of stiff anisotropic
731 minerals: rutile, and the implications for rutile elastic barometry: Mineralogical Magazine, v. 83,
732 no. 3, p. 339-347.
- 733

734 **Figure captions**

735 **Fig. 1** a) Simplified geological map of the Napier complex after (Carson et al., 2002). b)

736 Photomicrography showing the location of the analyzed rutile grain within the thin section. The rutile is
737 in contact with a coarse garnet that hosts rutile, spinel and sapphirine inclusions and resides in a matrix of
738 mesoperthite, quartz and minor biotite. C and d) Field images of the outcrop at Mt. Hardy. The photos
739 show the large scale foliation present at the sample locality with close up photo of a leucosome rich in
740 coarse-grained garnet and quartz from which the sample was prepared. Sample coordinates 66°49'04.2"S
741 50°42'60.0"E. Fsp = feldspar, Grt = garnet, Sph = sapphirine, Sp = spinel, Rt = rutile, Msp =
742 mesoperthite. Red arrow indicates studied rutile.

743 **Fig. 2:** (a): BSE images of analyzed rutile (Rt). Rutile is surrounded garnet (Grt), mesoperthite (Msp), k-
744 feldspar (Kfs) and biotite (Bt). The lift-out locations for TEM and APT specimens are indicated on the
745 image. TEM-1 corresponds to the lift-out taken at Curtin University and this monazite has a recrystallized
746 rim. TEM-2 corresponds with the monazite that does not show the recrystallized rim (b): BSE image
747 showing the presence of the submicron monazites (Mnz) in rutile, as well as corundum (Crn) and Fe-
748 oxide exsolutions. (c): BSE image of biotite directly in contact with rutile, also hosting submicron
749 monazite. (d) Mean-misorientation map showing the subtle differences in orientations that are linked to
750 low-angle boundaries. The black boundaries show the location of twin boundaries (cf. (Verberne et al.,
751 2022)). (e) Mean-misorientation and local misorientation maps for two low-angle boundaries. In both
752 maps a monazite is present, with LAB 1 being in close vicinity and LAB 2 being pinned on the monazite
753 inclusion. The corresponding slip system for the low-angle boundaries is $\{101\}\langle 201\rangle$.

754

755 **Fig 3.** TEM BF cross-section to illustrate the structure of the monazite inclusion. From outward inward:
756 the rutile host, the interface between rutile and the monazite, the recrystallized rim including 5-20nm
757 clusters and nano-inclusions of Li, Ca and Pb, and the core of the monazite exhibiting a mottled texture.

758 **Fig. 4:** TEM/STEM characterization of the TEM foils. a) From the left to the right: TEM-Bright Field
759 (BF) image of one entire monazite inclusion in rutile, Selected Area Electron Diffraction (SAED) pattern
760 from the monazite inclusion seen in the [12-1] zone axis and associated with a crystallographic model
761 from CePO₄-monazite in the same orientation, and high-resolution STEM-Annular Dark Field (ADF)
762 image of the monazite-rutile interface. b-Top) Left STEM-ADF and right STEM-ABF images from part
763 of the Mnz inclusion seen in A. b-below) STEM BF and HAADF images from a Mnz inclusion with an
764 alteration rim.

765 **Fig. 5 STEM-ADF images** of clusters in the core and recrystallization rim. a-c) within the core there is a
766 high density of clusters (bright spots) which are trapped within zones of radiation damage (black spots).
767 d-f) The recrystallized rim yields far fewer bright spots, however different inclusions are visible. Note
768 that despite the presence of clusters and inclusions the lattice is almost not affected

769 **Fig 6:** Mass spectrum of rutile, monazite and the interface with highlighted trace elements that are
770 enriched in the interface. Li, Ca and U are observed in one location in the mass spectrum. Pb is observed
771 at both Pb⁺ and Pb⁺⁺. ²⁰⁶Pb is visible at 103 Da and 206 Da, and ²⁰⁷Pb is observed at 103.5 Da and 207 Da,
772 ²⁰⁸Pb is not present above background.

773

774 **Fig 7: TEM EDS spectra and maps**

775 a) TEM EDS analyses conducted on a monazite without recrystallized rim. EDS analyses was performed
776 on the matrix and a cluster to show the difference in composition and the presence of Ca, Pb and U in the
777 cluster. b) TEM EDS maps of the core of the monazite with a recrystallized rim showing the clusters are
778 rich in Ca, Pb, and U. c) TEM EDS maps of clusters present in the recrystallized rim. Ce is shown as an
779 indication for decreased concentrations of major elements. d) TEM EDS interface analysis, showing an
780 EDS map to highlight that Pb is enriched at the interface with a Pb concentration profile (white arrow).

781 **Fig. 8** (a, b): Ion maps of specimen 1 and 2 showing monazite (green) is encapsulated within rutile (grey).
782 The interface is enriched in Ca and Pb in both specimens. Also, the curvature over the monazite interface

783 is clearly visible in the ion maps of specimen 2. (c): Ion maps of specimen 3 not only yields a Pb and Ca
784 enriched interface but also reveal the presence of Ca (d) and Li-rich (e) inclusions. Within the Li-rich
785 inclusion an isoconcentration surface for Pb is shown to demonstrate the presence and equivalence of the
786 inclusions in Fig 4b

787

788 **Fig 9** a) Proximity histograms for the three inclusions observed within the APT data. All inclusions show
789 a slight increase in Pb concentration and a strong increase in one of three cations (Ca, Li, Mg) indicating
790 the formation of a new phase or solid-solution. Ce is shown as a proxy for the REE, which show similar
791 trends at lower concentrations.

792

793 **Fig. 10** Ion map showing the heterogeneous distribution of Si + Al within the interface of specimen 3.
794 The 1D concentration profile was taken along the interface as indicated in the top panel. The enrichment
795 of Al, Si and Pb clearly varies within each section. The linear features near the middle are interpreted as
796 dislocation present at the interface

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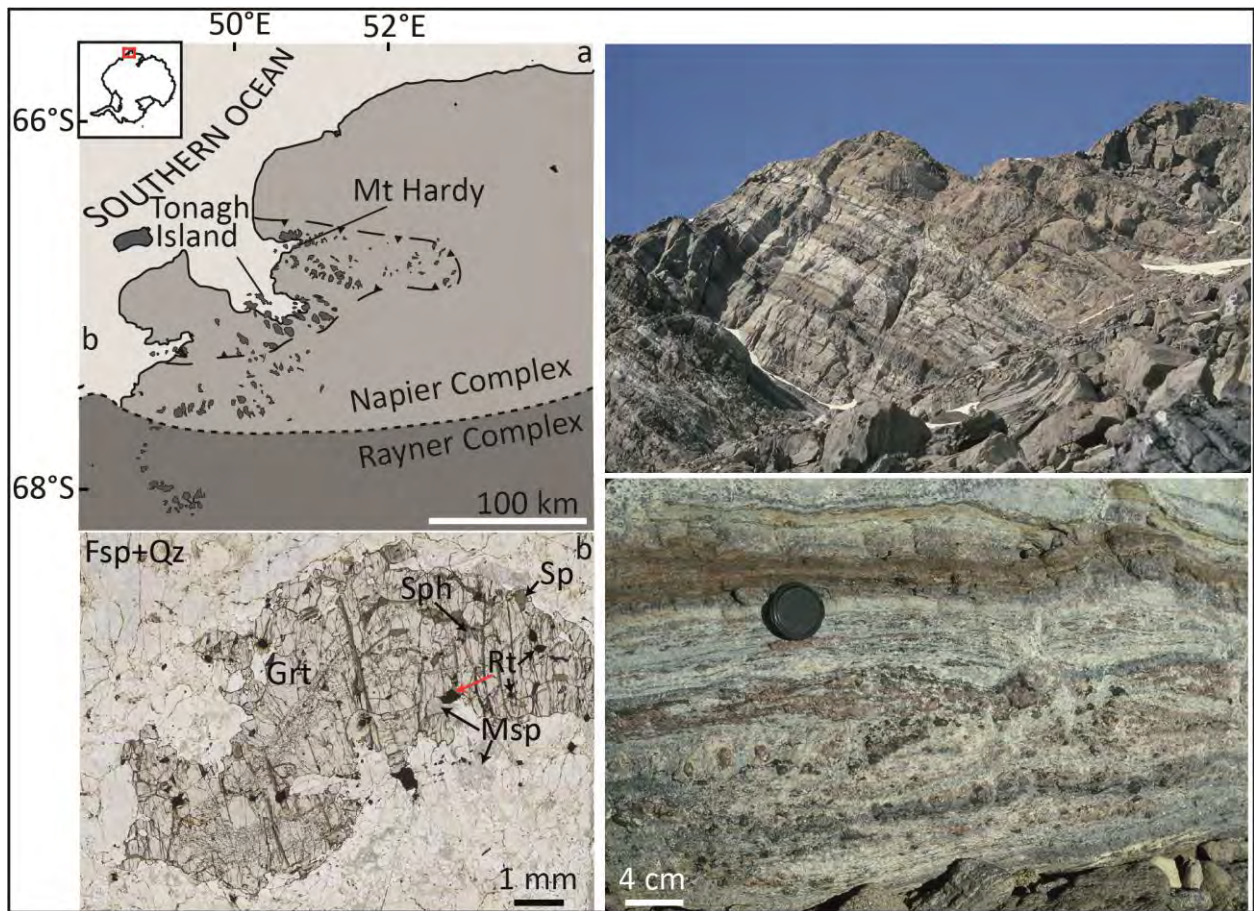
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805 Figure 1



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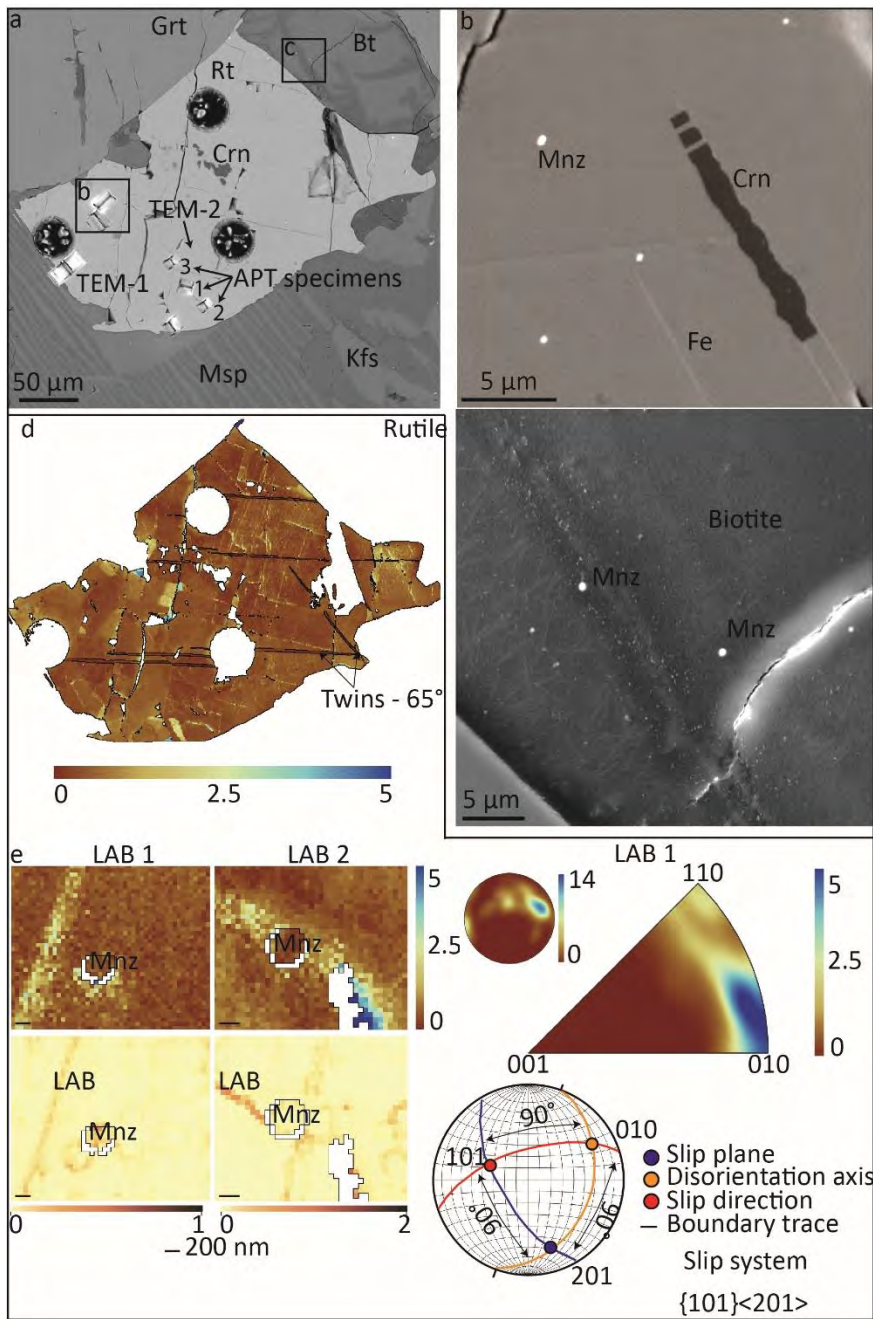
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815 Figure 2



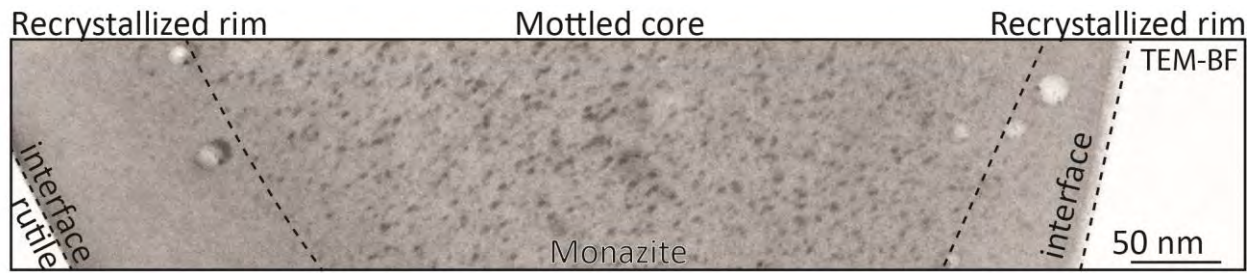
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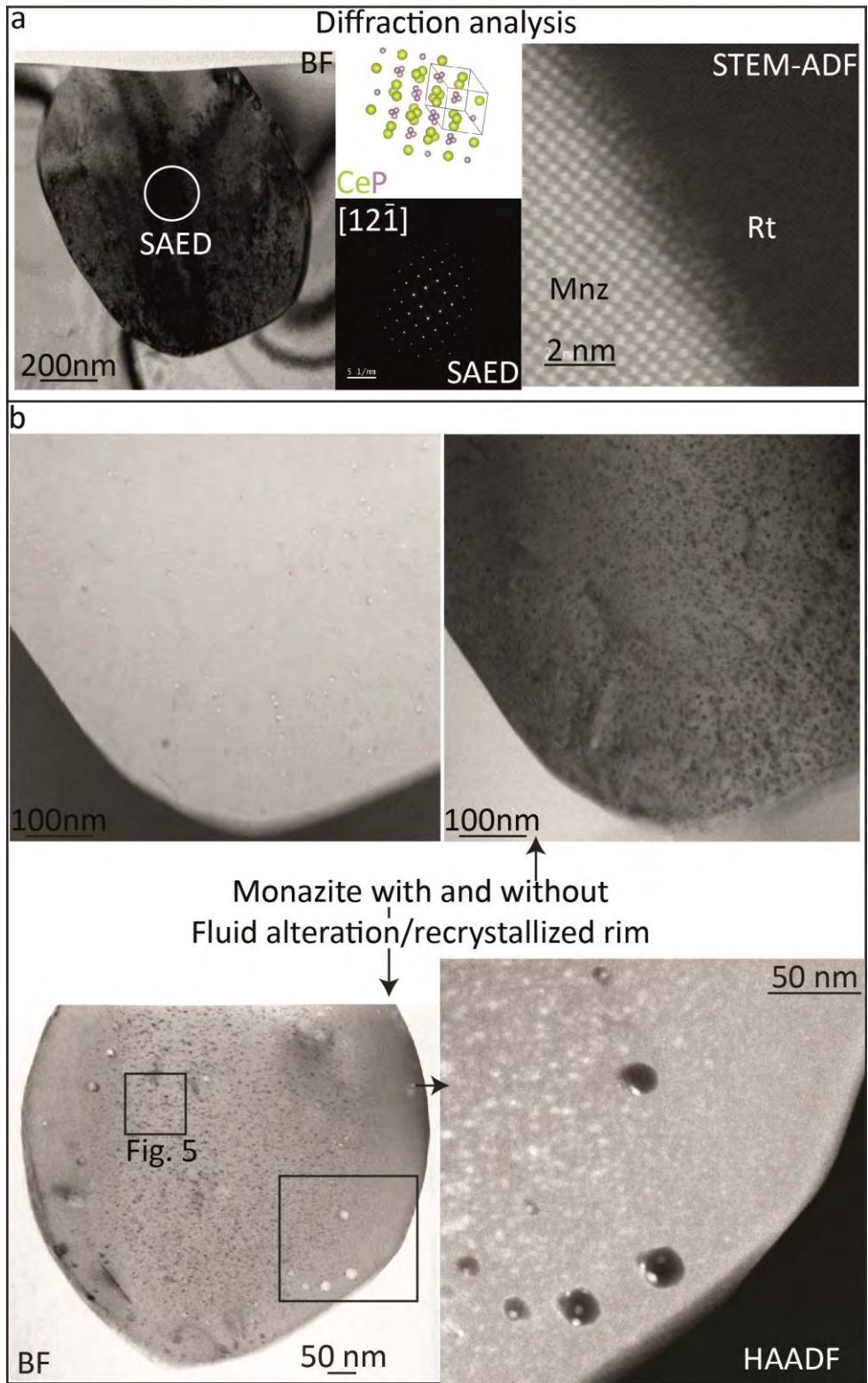
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820 Figure 3

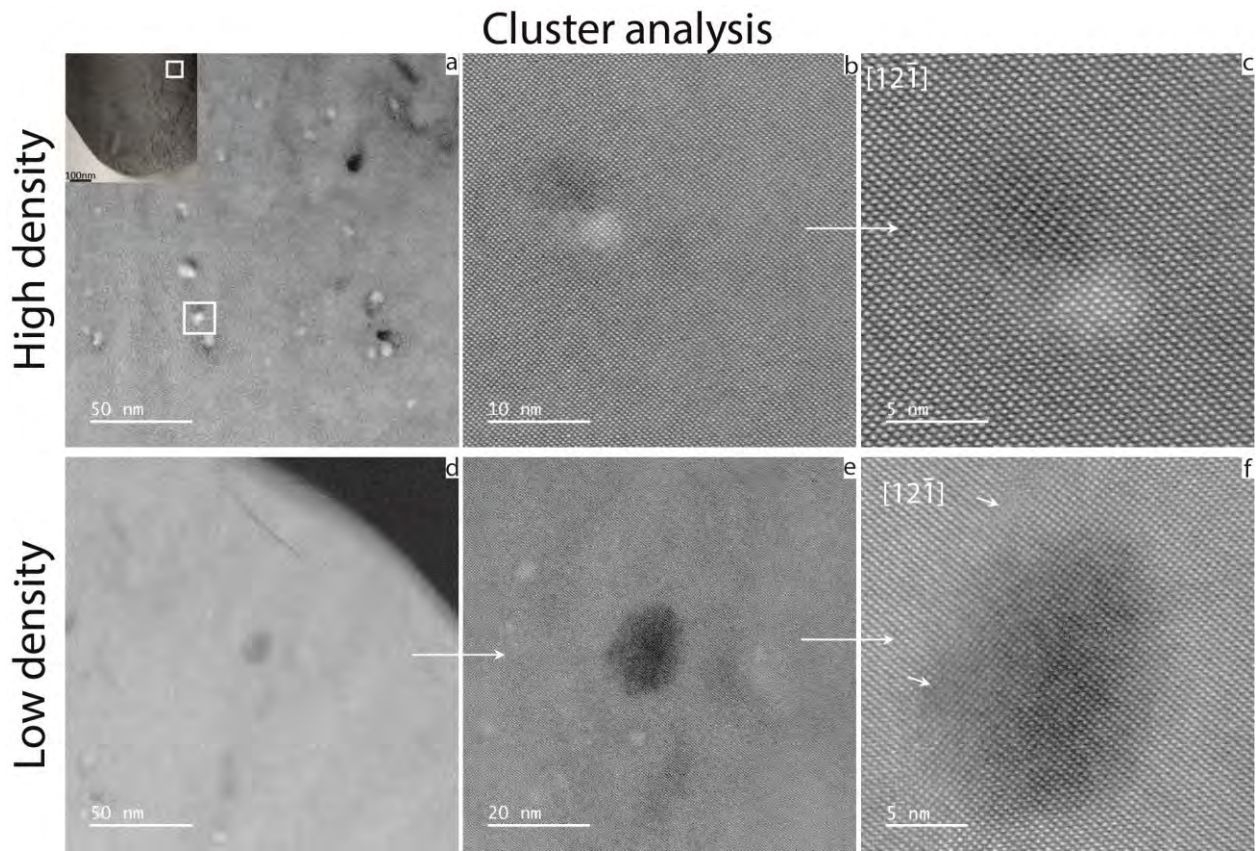


823 Figure 4



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825 Figure 5

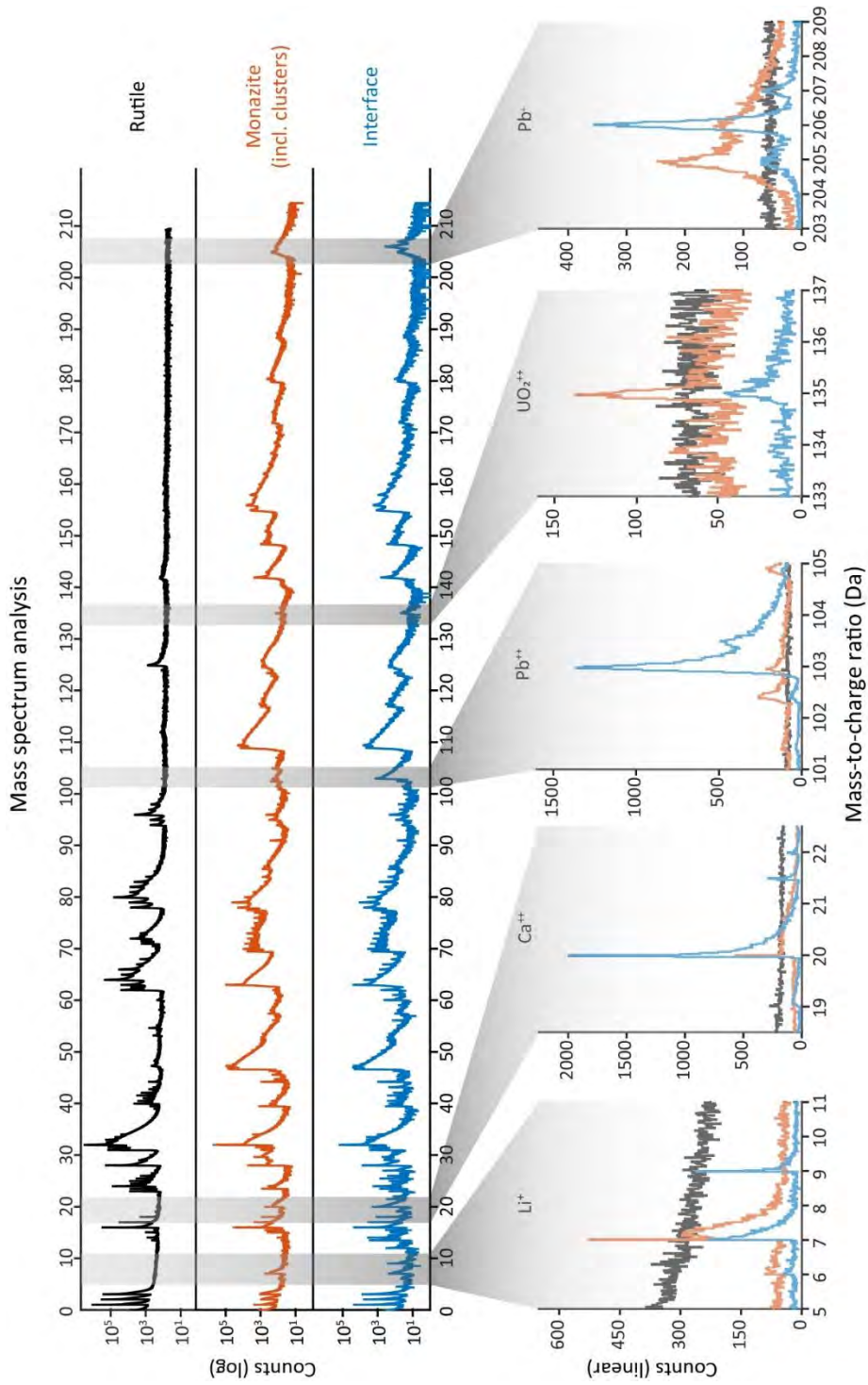


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828 Figure 6

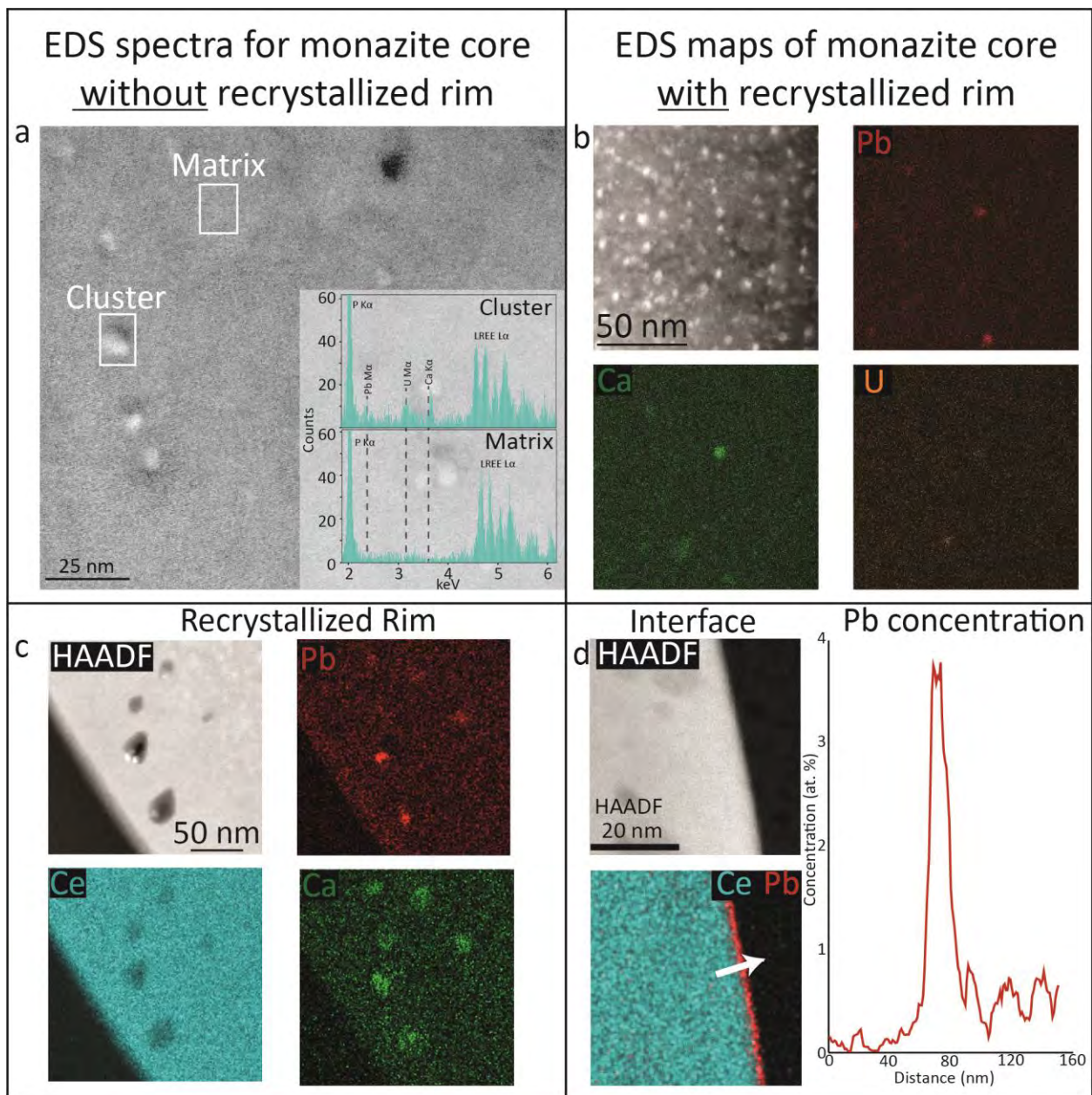
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832 Figure 7



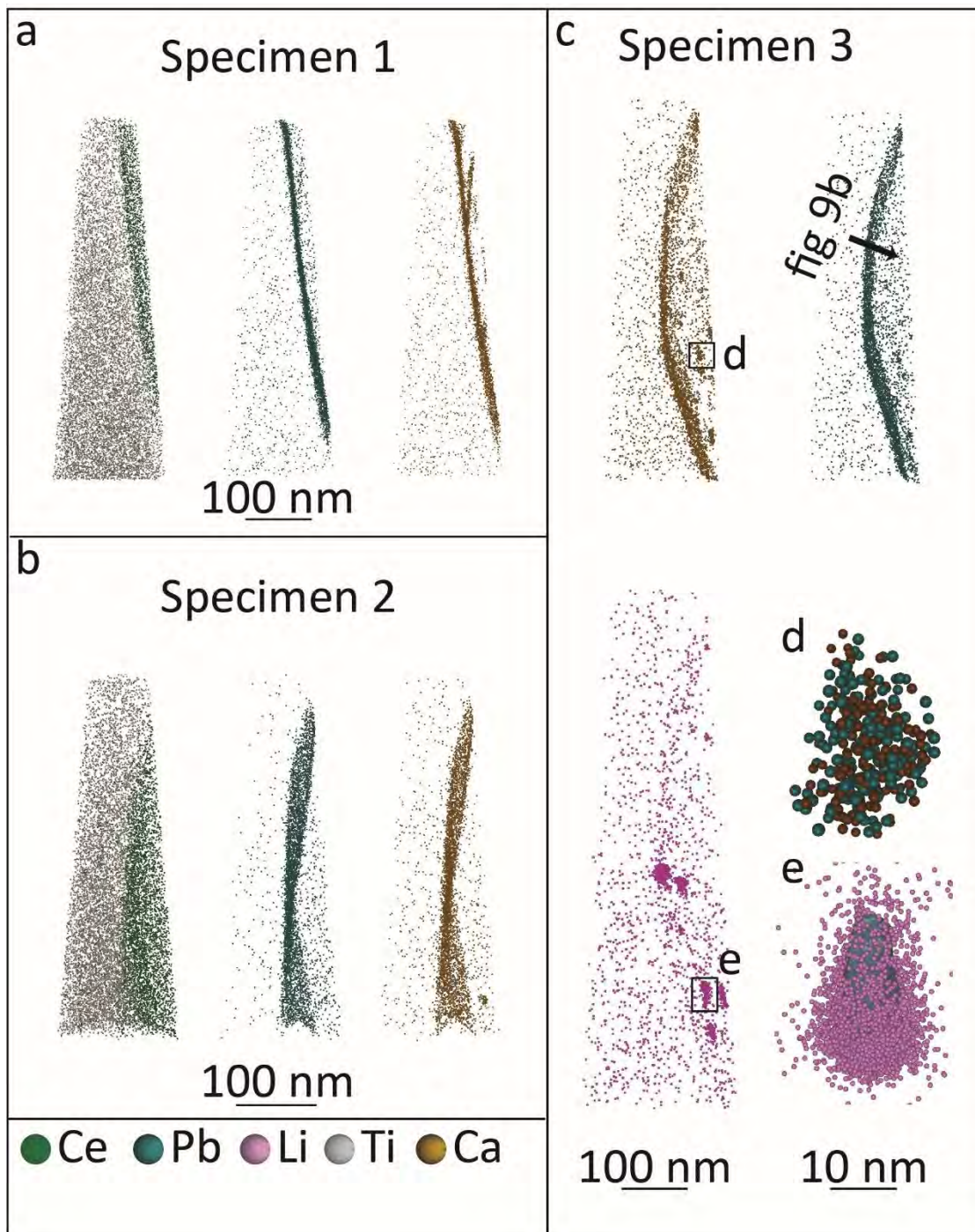
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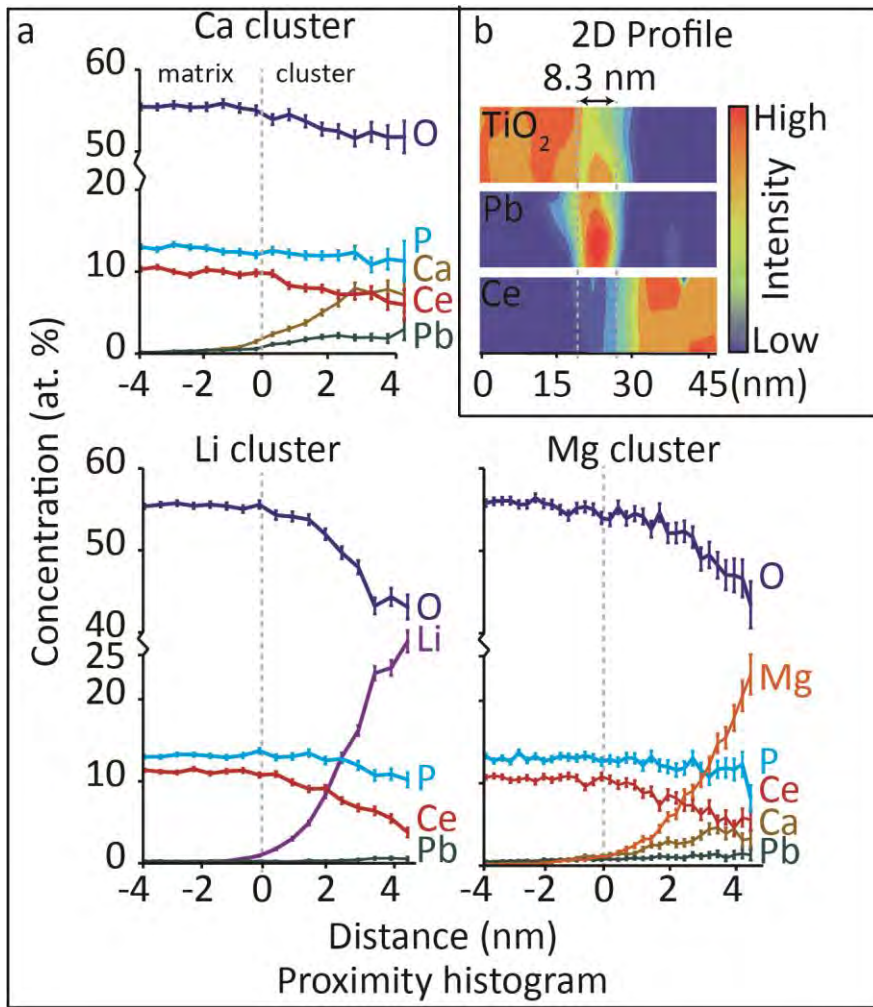
837 Figure 8



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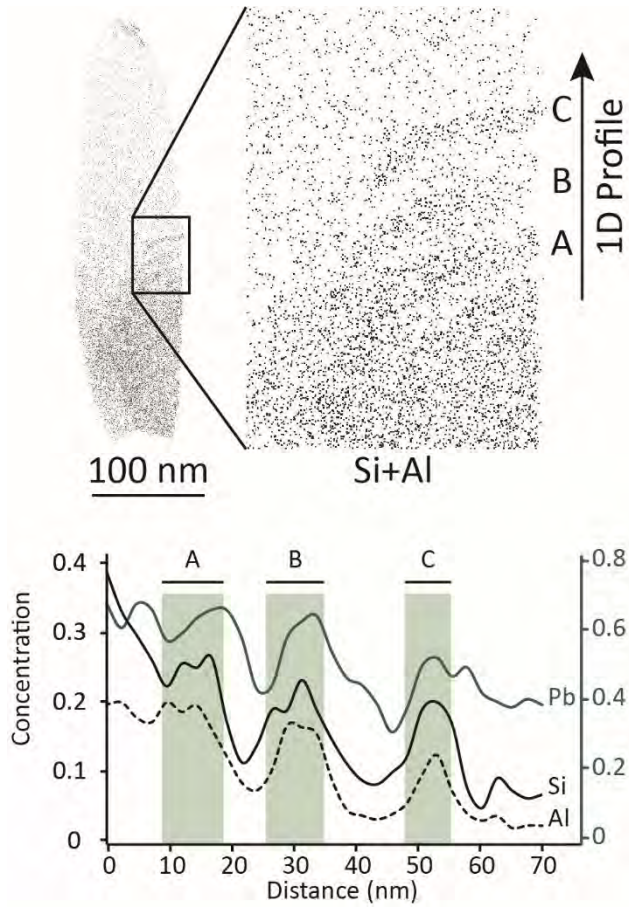
840 Figure 9



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843 Figure 10



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