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# CFD simulation of pollutant dispersion using anisotropic models: Application to an urban like environment under neutral and stable atmospheric conditions

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## Abstract

Modeling the atmospheric dispersion of pollutants emitted from different types of sources under various atmospheric conditions is an essential prerequisite for risk assessment studies and emergency preparedness. In this study, we evaluate the performance of different turbulence models and velocity-scalar correlation models implemented in the *Code\_Saturne*. The evaluation is done with observations of four trial cases of the Mock Urban Setting Test (MUST) campaign in an urban type environment and in neutral and stable atmospheric conditions. For all of the trials studied, the CFD model with a first-order closure model ( $k - \varepsilon$ ) predicts 61.1% of the concentrations within a factor of two of the observations, which is higher than the percentage of predicted points (58.8%) when a second-order closure model ( $R_{ij} - \varepsilon$ ) is used. Overall, the CFD model underestimates observed concentrations, regardless of the turbulence model used. For the trial with slightly stable conditions, the results show that the  $k - \varepsilon$  model combined with an algebraic SGDH model predicts 75% concentrations within a factor of two of the observations. The performance of the  $k - \varepsilon$  model is compared to that of the  $R_{ij} - \varepsilon$  model when used with the algebraic SGDH, GGDH models and with the scalar flux transport equation (DFM model). Under slightly stable atmospheric conditions, the DFM model predicts 69% of concentrations within a factor of two of observations, showing promise for modeling under these conditions, despite its relatively high computational cost.

**Keywords:** CFD modeling, MUST experiment, algebraic models, anisotropic, stable conditions

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## 1. Introduction

The dispersion of gaseous pollutants and small particles in an urban area is characterized by complex turbulence phenomena, such as flow detachment and high shear zones (Yu and Thé, 2017). In addition to this dynamic turbulence, thermal air turbulence tends to dampen the turbulent fluctuations resulting from flow shear under conditions of stable or very stable stratification, thus reducing pollutant dispersion. The complexity of the physical phenomena involved in these conditions results in high anisotropy of the atmospheric boundary layer (ABL), which limits the ability of standard isotropic models to correctly predict the dispersion field. Computational fluid dynamics (CFD) models are one of the most promising tools to describe the dispersion of pollutants in the near field of obstacles.

The CFD approaches used to solve the turbulent flow equations in most studies mainly involve large eddy simulation (LES) and Reynolds-averaged Navier-Stokes (RANS) methods. Large eddy simulations (LES) are considered a promising step in atmospheric flows, as they provide a better understanding of the various flow phenomena that occur around obstacles (e.g. (Tominaga and Stathopoulos, 2012, 2011, 2010; Salim et al., 2011)). However, this type of simulation requires higher computing power due to the time dependence of the simulations and the finer grids that have to be used (Lateb et al., 2016). Reynolds-Averaged Navier-Stokes (RANS) simulations use Reynolds decomposition to filter turbulent fluctuations from the mean flow field (Kavian Nezhad et al., 2022; Bahlali et al., 2019; Tominaga and Stathopoulos, 2018, 2017; Yu and Thé, 2017; Milliez and Carissimo, 2008, 2007; Kumar and Feiz, 2016; Kumar et al., 2015). The turbulence on the mean flow field is fully incorporated by a turbulence model and has an effect on both the dynamic and scalar fields. As a result, the accuracy and reliability of field modeling results are highly dependent on the quality of the turbulence model applied. Researchers are constantly investing in RANS simulations to model atmospheric fluxes, trying to improve the sub-models included. In most simulations of atmospheric dispersion in urban areas, the standard  $k - \varepsilon$  model (Launder and Spalding, 1974) is the most widely used. The standard model is usually modified or calibrated to adapt to the problem of dispersion in different terrain topographies and atmospheric stabilities (see for example (Kumar et al., 2015; Milliez and Carissimo, 2007; Alinot and Masson, 2005), etc.). In addition to the standard  $k - \varepsilon$  model, the Reynolds Stress Model (RSM) has been used in numerous test cases for pollutant dispersion in cities and urban areas. This model, originally developed for high Reynolds number flows, has sometimes proved to outperform the standard  $k - \varepsilon$  isotropic model. For instance, Knaus et al. (2018) compared three turbulence models - standard  $k - \varepsilon$ ,  $k - \varepsilon$  renormalization group (RNG) and RSM model- to simulate wind flow over complex terrain. Only the RSM model proved capable of predicting turbulence intensity with a high degree of accuracy. Bahlali et al. (2019) showed that the use of the  $R_{ij} - \varepsilon$  turbulence model developed by Launder et al. (1975) gives a more physical representation of the obstacles' influence on the concentration profiles.

46 Regarding the turbulent flux in the scalar transport equation, it is often  
47 solved using the Simple Gradient Diffusion Hypothesis (SGDH). There are,  
48 however, other models, such as the Generalized Gradient Diffusion Hypoth-  
49 esis (GGDH) (Daly and Harlow, 1970), the Differential Flow Model (DFM),  
50 etc., more commonly used in an industrial context. For atmospheric dispersion  
51 problems, Rossi and Iaccarino (2009) showed that gradient-transport models  
52 are improved by the use of algebraic closures (GGDH and HOGGDH) and pre-  
53 dict atmospheric dispersion in complex flows better than the standard eddy-  
54 diffusivity model (see also (Rossi et al., 2009)).

55 The objective of this work is to perform a comprehensive validation study  
56 of different anisotropic turbulence models for flow and scalar fields. Simula-  
57 tion results are evaluated using observations from the Mock Urban Setting Test  
58 (MUST) campaign (Yee and Biltoft, 2004; Biltoft, 2001), which is a quasi-urban  
59 experiment involving the release of a neutral gas from a point source inside a  
60 set of rectangular shipping containers. Four trials were chosen to simulate flow  
61 and dispersion patterns near the source and at greater distances. The stability  
62 conditions in these trials range from neutral to extremely stable. Numerical sim-  
63 ulations are performed using the three-dimensional (3-D) code *Code\_Saturne*.  
64 The  $k - \varepsilon$  (Launder and Spalding, 1974) and  $R_{ij} - \varepsilon$  LRR (Launder et al.,  
65 1975) models are used and compared to predict the flow field, while the SGDH  
66 and GGDH models are initially adopted to relate the scalar flux to the mean  
67 gradients of the concentration field. An additional analysis of closure for turbu-  
68 lent scalar flux based on the differential flow model (DFM) transport equation  
69 is also presented, and the results are compared with the predictions provided  
70 by the SGDH and GGDH algebraic models. The capabilities and limitations  
71 of the tested models for both wind flow and concentration fields are rigorously  
72 evaluated using multiple validation metrics.

73 The paper is organized as follows: section 2 describes the experimental data  
74 sets and tests selected for the numerical simulations. Section 3 presents the CFD  
75 code, including the solved equations and turbulence models used in this study.  
76 Section 4 presents the numerical setup for the simulations in the code. Sec-  
77 tion 5 presents the numerical results of the simulations and the qualitative and  
78 quantitative analysis of the predictions against the experimental observations.  
79 Section 6 summarizes the conclusions of this study.

## 80 2. Description of the experimental data sets

### 81 2.1. The MUST experiment

82 The experimental campaign MUST (Mock Urban Setting Test) (Yee and  
83 Biltoft, 2004; Biltoft, 2001) was conducted at the US Army Dugway Proving  
84 Ground (DPG), in the Utah desert, in the western United States ( $40^{\circ}12.606'N -$   
85  $113^{\circ}10.635'W$ ) from September 6 to 27, 2001, by the Defence Threat Reduction  
86 Agency. The main objective of this campaign was to obtain an experimental  
87 database on the dispersion of a pollutant at the scale of a pseudo-city in order

88 to develop new models or validate existing numerical models. Figure 1 shows a  
89 photograph of the experiment site.

90 To highlight the presence of obstacles in the urban landscape (buildings,  
91 trees, etc.), the buildings were rectangular containers of dimensions  $12.2\text{m} \times$   
92  $2.42\text{m} \times 2.54\text{m}$ . The urban canopy is thus formed by 120 containers uniformly  
93 distributed in 10 rows of 12 containers. The terrain of the experimental site  
94 was mostly flat, with an average roughness length of  $0.045 \pm 0.0005\text{m}$  (Yee and  
95 Biltoft, 2004).

96 Receptors are placed between rows of containers on four masts, a central  
97 tower, and four sampling lines to collect information on the dynamic field and  
98 concentration field of the pollutant, as well as outside the canopy to measure the  
99 velocity field and temperature. Figure 2 shows the location of the containers,  
100 pollutant source, and measurement devices.

101 The experiment consisted of 63 continuous releases of a non-reactive gas into  
102 the air under different meteorological and atmospheric stratification conditions  
103 (stable, unstable and neutral). The gas released is propylene ( $C_3H_6$ ) and each  
104 release lasts for 15min, from which quasi-stationary periods of 200 seconds were  
105 extracted and used to evaluate the simulated dispersion results in each trial  
106 of the experiment (Yee and Biltoft, 2004). The concentration field is mainly  
107 measured using 40 PID (Photoionization Detectors) placed at 1.6m from the  
108 ground and arranged horizontally along 4 rows. According to Figure 2, the  
109 sampling lines for the concentration measurements are defined as follows:

- 110 • Line 1: receptors 1 through 12,
- 111 • Line 2: receptors 13 to 21,
- 112 • Line 3: receptors 22 to 30,
- 113 • Line 4: receptors 31 to 40.

114 Concentrations were also measured at 8 levels (1, 2, 4, 6, 8, 10, 12, and 16m)  
115 vertically on the 32m tower above the ground surface, and at 6 levels on masts  
116 A, B, C, and D.

## 117 2.2. Selected trials

118 For the present work, 4 trials are selected for the validation of the dispersion  
119 model. The characteristics of each trial, used as input data for the model  
120 simulations, are summarized in Table 1. Average wind speed and direction are  
121 taken at a height of 4m above the ground surface from the south tower upstream  
122 of the urban canopy (Yee and Biltoft, 2004). Observations of surface friction  
123 velocity ( $u^*$ ), Obukhov length ( $L$ ), and turbulent kinetic energy ( $k$ ) are taken  
124 at a height of 4m from a 32m tower (tower T) in the center of the domain. The  
125 selected trials are the following:



Figure 1: Photograph of the MUST experimental campaign site (Biltoft, 2001).

- 126 • Trial No.2681829: this trial was performed under neutral atmospheric  
 127 conditions, i.e. with zero vertical potential temperature gradient for a very  
 128 high value of the Obukhov length ( $L = 28000\text{m}$ ). The ABL in this trial  
 129 has a stability class D according to the Pasquill classification (Pasquill,  
 130 1972). The average wind direction is  $-41^\circ$  and the wind is oriented from  
 131 southwest to northeast (Figure 2). Of the twenty trials in the experimental  
 132 campaign, which are often used to validate models, this trial recorded the  
 133 strongest wind.
- 134 • Trial No.2692157: this trial was conducted under weakly stable atmo-  
 135 spheric conditions, with a moderate wind velocity ( $2.1 < U < 5\text{m/s}$ ). It  
 136 is of stability class F according to the classification of (Pasquill, 1972)  
 137 ( $100 < L < 200\text{m}$ ). In this trial, the wind blows from southeast to north-  
 138 west with a direction angle of  $43^\circ$  (Figure 2).

Table 1: Physical characteristics of the selected trials, obtained at a reference height of 4m:  $U$  is the mean wind speed,  $\alpha$  is the incident wind direction measured from the x-axis,  $k$  is the turbulent kinetic energy,  $Q$  is the pollutant flow rate,  $z_s$  is the height of the release source,  $u_*$  is the frictional velocity,  $L$  is the Obukhov length, and  $\varepsilon$  is the dissipation rate.

Trial No.	$U$ ( $\text{m s}^{-1}$ )	$\alpha$ ( $^\circ$ )	$k$ ( $\text{m}^2 \text{s}^{-2}$ )	$Q$ ( $\text{L min}^{-1}$ )	$z_s$ (m)	$u_*$ ( $\text{m s}^{-1}$ )	$L$ (m)	$\varepsilon$ ( $\text{m}^2 \text{s}^{-3}$ )
2671934	1.63	-48	0.148	200	1.8	0.08	5.8	0.0013
2672033	2.69	-26	0.251	200	1.8	0.17	4.8	0.012
2681829	7.93	-41	1.46	225	1.8	1.1	28000	0.8
2692157	2.98	43	0.505	225	2.7	0.39	130	0.042

- 139 • Trial No.2672033: this trial was conducted under highly stable atmo-  
 140 spheric conditions, with moderate winds ( $2.1 < U < 5\text{m/s}$ ). It is of stabil-  
 141 ity class F according to the classification of (Pasquill, 1972) ( $0 < L < 10\text{m}$ ).  
 142 Among all the trials of the experimental campaign, the lowest value of  
 143 Obukhov length was obtained for this trial. The wind blows from south-  
 144 west to northeast with a direction angle of  $26^\circ$  (Figure 2).

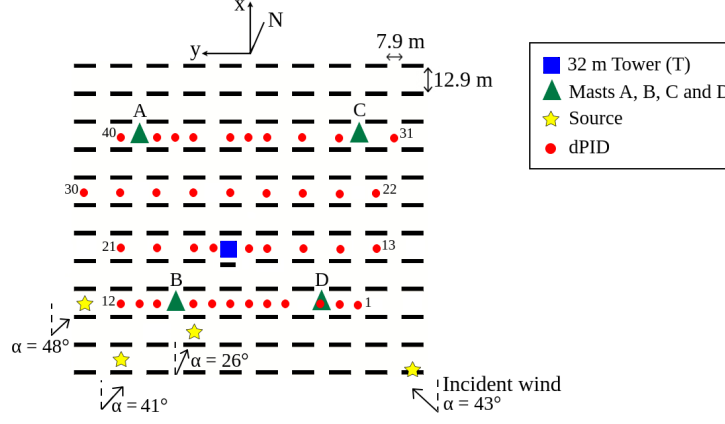


Figure 2: Schematic representation of the MUST experimental site: the triangles represent the masts A, B, C and D, the blue square represents the central tower of 32m height. The position of the sources for the 4 selected trials is represented by yellow stars.

- Trial No.2671934: this trial was carried out under very stable atmospheric conditions, with low winds ( $U < 2.1\text{m/s}$ ). It is of stability class F according to the classification of (Pasquill, 1972) ( $0 < L < 10\text{m}$ ). The lowest velocity value was observed for this trial. The wind blows from southwest to northeast with a direction angle of  $48^\circ$  (Figure 2).

### 3. Description of *Code\_Saturne* model for flow and dispersion

#### 3.1. Governing equations

*Code\_Saturne* (<https://www.code-saturne.org/cms/web/documentation/v6>) is a 3-D CFD parallelized code which solves the Navier-Stokes equations by a stationary or non-stationary algorithm for incompressible or weakly dilatible fluids, isothermal or not, with consideration of passive scalars (Archambeau et al., 2004). The spatial discretization is obtained using a finite volume approach for a collocated arrangement of all variables. In this study, the atmospheric module of the code is used along with an Eulerian RANS approach to solve the average form of the equations for both turbulent flow and scalars. The equations of conservation of mass (continuity), momentum, potential temperature and species concentration are as follows:

- *Continuity equation (with the anelastic approximation):*

$$\frac{\partial \bar{\rho} \bar{U}_i}{\partial x_i} = 0 \quad (1)$$

- *Conservation of momentum:*

$$\frac{\partial \bar{U}_i}{\partial t} + \bar{U}_j \frac{\partial \bar{U}_i}{\partial x_j} = -\frac{1}{\bar{\rho}} \frac{\partial \bar{P}}{\partial x_i} + \frac{\partial}{\partial x_j} \nu \left( \frac{\partial \bar{U}_i}{\partial x_j} + \frac{\partial \bar{U}_j}{\partial x_i} \right) - \frac{(\bar{\rho} - \rho_{ref})}{\bar{\rho}} g \delta_{i3} - \frac{\partial}{\partial x_j} \overline{U'_i U'_j} \quad (2)$$

164 • *Conservation of potential temperature:*

$$\frac{\partial \bar{\Theta}}{\partial t} + \bar{U}_j \frac{\partial \bar{\Theta}}{\partial x_j} = \frac{1}{\bar{\rho} C_p} \frac{\partial}{\partial x_j} \left( \lambda \frac{\partial \bar{\Theta}}{\partial x_j} \right) - \frac{\partial}{\partial x_j} \overline{\Theta' U_j'} + \bar{S}_\Theta \quad (3)$$

165 • *Conservation of concentration (Advection-diffusion equation for a scalar):*

$$\frac{\partial \bar{C}}{\partial t} + \bar{U}_j \frac{\partial \bar{C}}{\partial x_j} = \frac{\partial}{\partial x_j} \left( D \frac{\partial \bar{C}}{\partial x_j} \right) - \frac{\partial}{\partial x_j} \overline{C' U_j'} + \bar{S}_C + \bar{R}_C \quad (4)$$

166 where,  $\rho$  is the air density,  $U$  is the velocity vector,  $p$  is the pressure,  $\Theta$  is the  
 167 potential temperature,  $\rho_{ref}$  the air density at hydrostatic reference state.  $g$  is  
 168 the gravitational acceleration vector,  $C_p$  is the specific heat at constant pressure,  
 169  $\lambda$  is the effective thermal conductivity,  $D$  is the molecular diffusion coefficient  
 170 for the species,  $S_\Theta$  is a thermal source term,  $S_C$  is the species source term,  $R_C$   
 171 is the term of production/destruction of the species by chemical reaction.  
 172 The Reynolds tensor  $\overline{U_i' U_j'}$ , thermal turbulent flux  $\overline{\Theta' U_j'}$  and concentration tur-  
 173 bulent flux  $\overline{C' U_j'}$  are modeled using turbulence closure models.

### 174 3.2. Turbulence models

#### 175 3.2.1. The $k - \varepsilon$ model

176 This model is based on the Boussinesq hypothesis which assumes a alignment  
 177 between the average strain rate tensor and the anisotropic part of the Reynolds  
 178 tensor (Launder and Spalding, 1974). Its formulation is given by the following  
 179 relation:

$$\overline{U_i' U_j'} = \nu_t \left( \frac{\partial \bar{U}_i}{\partial x_j} + \frac{\partial \bar{U}_j}{\partial x_i} \right) - \frac{2}{3} k \delta_{ij} \quad (5)$$

180 The transport equations for the kinetic energy ( $k$ ) and the dissipation rate ( $\varepsilon$ )  
 181 are written as:

$$\frac{\partial k}{\partial t} + U_i \frac{\partial k}{\partial x_i} = \frac{1}{\bar{\rho}} \frac{\partial}{\partial x_j} \left[ \left( \nu + \frac{\nu_t}{\sigma_k} \right) \frac{\partial k}{\partial x_j} \right] + P_t + G - \varepsilon \quad (6)$$

$$\frac{\partial \varepsilon}{\partial t} + U_i \frac{\partial \varepsilon}{\partial x_i} = \frac{1}{\bar{\rho}} \frac{\partial}{\partial x_j} \left[ \left( \nu + \frac{\nu_t}{\sigma_\varepsilon} \right) \frac{\partial \varepsilon}{\partial x_j} \right] + C_{\varepsilon 1} \frac{\varepsilon}{k} (P_t + C_{\varepsilon 3} G) - C_{\varepsilon 2} \frac{\varepsilon^2}{k} \quad (7)$$

182

183 With:

$$P_t = -\bar{\rho} \overline{U_i' U_j'} \frac{\partial \bar{U}_j}{\partial x_i}, \quad G = -\frac{1}{\bar{\rho}} \frac{\beta g}{Pr_t} \nu_t \frac{\partial \bar{\theta}}{\partial z}, \quad \nu_t = \bar{\rho} C_\mu \frac{k^2}{\varepsilon} \quad (8)$$

185 where  $P_t$  and  $G$  are respectively the mechanical and buoyancy production rate of  
 186  $k$ ,  $\nu_t$  is the turbulent eddy viscosity,  $\beta \sim \frac{1}{\Theta}$  is the thermal expansion coefficient,  
 187  $\mu_l$  is the kinematic viscosity,  $\sigma_k$ ,  $\sigma_\varepsilon$  and  $Pr_t$  are the turbulent Prandtl numbers.  
 188  $C_{\varepsilon 1}$ ,  $C_{\varepsilon 2}$ ,  $C_{\varepsilon 3}$  and  $C_\mu$  are empirical constants. The constant  $C_{\varepsilon 3}$  is set to zero  
 189 for unstable stratification and to 1 for stable stratification.

190 It is known that the constants of the standard model may not be suitable  
191 to reproduce the turbulence of the atmospheric flow as they were originally  
192 calibrated for industrial flows (Sadek et al., 2013). Therefore, a new model  
193 with a set of calibrated constants, proposed by Duynkerke (1988), is used in our  
194 calculations and compared to the model with the standard parameterization.  
195 The model constants are given in Table 2. We denote by [LA83] the standard  
196  $k - \varepsilon$  model of Launder and Spalding (1974) while by [DU88] the calibrated  
197 model of Duynkerke (1988).

Table 2:  $k - \varepsilon$  model constants.

	$C_{\varepsilon 1}$	$C_{\varepsilon 2}$	$\sigma_k$	$\sigma_\varepsilon$	$C_\mu$
$k - \varepsilon$ [LA83]	1.44	1.92	1.0	1.3	0.09
$k - \varepsilon$ [DU88]	1.44	1.92	1.0	2.38	0.033

198 In fact, the vertical gradient of potential temperature  $\partial\bar{\theta}/\partial z$  is negative ( $G \geq$   
199  $0$ ) under unstable atmospheric conditions, therefore, the thermal stratification  
200 has an effect of production of the turbulence. Under stable conditions, however,  
201  $\partial\bar{\theta}/\partial z$  is positive ( $G \leq 0$ ), therefore, the thermal stratification has an effect of  
202 destruction of the turbulence.

### 203 3.2.2. The $R_{ij} - \varepsilon$ model

204 This model solves six transport equations for the Reynolds tensors  $R_{ij} =$   
205  $\overline{U_i'U_j'}$  and one equation for the dissipation rate  $\varepsilon$ . The latter is very similar  
206 to that of the  $k - \varepsilon$  model, only the turbulent viscosity and gravity terms  
207 change (Launder et al., 1975). For each  $R_{ij}$  component, the following transport  
208 equation is solved:

$$\frac{\partial R_{ij}}{\partial t} + U_k \frac{\partial R_{ij}}{\partial x_k} = P_{ij} + D_{T,ij} + D_{M,ij} + G_{ij} + \phi_{ij} - \varepsilon_{ij} \quad (9)$$

209 where  $P_{ij}$  is the average shear production term,  $D_{T,ij}$  is the turbulent diffu-  
210 sion term,  $D_{M,ij}$  is the molecular diffusion term,  $G_{ij}$  is the production/destruction  
211 term by thermal effects,  $\phi_{ij}$  is the pressure-strain term and  $\varepsilon_{ij}$  is the dissipation  
212 term.

Table 3:  $R_{ij} - \varepsilon$  LRR model constants.

	$\sigma_k$	$\sigma_\varepsilon$	$C_{\varepsilon 1}$	$C_{\varepsilon 2}$	$C_1$	$C_2$	$C'_1$	$C'_2$	$C_L$	$C_\mu$
$R_{ij} - \varepsilon$ [LA75]	0.82	1.0	1.44	1.92	1.8	0.6	0.5	0.3	0.39	0.09
$R_{ij} - \varepsilon$ [SA13]	1.0	2.01	1.44	1.92	1.8	0.6	0.94	0.03	0.19	0.033

213 We note that the terms  $G_{ij}$ ,  $\varepsilon_{ij}$ ,  $D_{T,ij}$  and  $\phi_{ij}$  are not closed and therefore  
214 must be modeled. The modeling of the  $G_{ij}$  and  $\varepsilon_{ij}$  terms is identical to that of  
215 the  $k - \varepsilon$  model (Launder et al., 1975).

216 Similar to the  $k - \varepsilon$  model, most of the standard empirical constants in the  
 217  $R_{ij} - \varepsilon$  model have been calibrated to achieve the atmospheric turbulence levels  
 218 of the experiment. The methodology for setting the values of the constants is  
 219 presented by [Sadek et al. \(2013\)](#), who also draws on the work of [Panofsky and](#)  
 220 [Dutton \(1984\)](#); [Duynderke \(1988\)](#); [Rodi \(1993\)](#). The new calibrated  $R_{ij} - \varepsilon$   
 221 model is therefore used in the present work along with the standard form. The  
 222 coefficients of the  $R_{ij} - \varepsilon$  model for the standard [LA75] and calibrated [SA13]  
 223 versions are presented in Table 3.

### 224 3.2.3. Models for scalar turbulent fluxes

225 By averaging the transport equations for temperature and concentration, we  
 226 obtain turbulent heat flux and turbulent concentration flux. These terms must  
 227 now be modeled either by a transport equation or by algebraic models (GGDH  
 228 and SGDH). The latter two are obtained from the initial transport equation  
 229 assuming that the turbulence is in equilibrium, such that:  $P_k + G_k = \varepsilon$  and  $P_{ss} =$   
 230  $\varepsilon_{ss}$ , where  $P_k = \frac{1}{2}P_{jj}$  corresponds to half the trace of the production tensor,  
 231 while  $G_k = \frac{1}{2}G_{jj}$  refers to half the trace of the production tensor resulting  
 232 from volumetric forces. In addition,  $P_{ss}$  represents the production term of the  
 233 variance  $S'^2$  of scalar  $S$ , and  $\varepsilon_{ss}$  denotes dissipation of the variance of the same  
 234 scalar quantity. For simplicity, the expression for each model is written only  
 235 for the turbulent concentration flux responsible for the dispersion of the passive  
 236 tracer, knowing that those for the turbulent thermal flux can be deduced in a  
 237 similar manner.

- 238 • Differential Flux Model (DFM)

239 This approach is the most suitable to represent the physical mechanisms  
 240 leading to turbulent fluxes ([Lecocq, 2008](#); [Lecocq et al., 2008](#); [Dehoux,](#)  
 241 [2012](#)). The transport equation for the scalar flux is written as:

$$\begin{aligned}
 \frac{D\overline{C'U'_j}}{Dt} &= -\overline{C'U'_i} \frac{\partial \overline{U'_j}}{\partial x_i} - \overline{U'_i U'_j} \frac{\partial \overline{C}}{\partial x_i} - \beta g_i \overline{C'^2} - \frac{P'}{\rho} \frac{\partial C'}{\rho x_j} - (D + \nu) \frac{\partial C'}{\partial x_i} \frac{\partial U'_j}{\partial x_i} \\
 &+ \frac{\partial}{\partial x_i} \left( -\overline{C'U'_i U'_j} \right) + \frac{\partial}{\partial x_i} \left( D \frac{\partial \overline{C'}}{\partial x_i} U'_j + \nu C' \frac{\partial U'_j}{\partial x_i} \right) + \frac{\partial}{\partial x_i} \left( \frac{C' P'}{\rho} \delta_{ij} \right)
 \end{aligned}
 \tag{10}$$

- 242 • Generalized Gradient Diffusion Hypothesis (GGDH)

243 The GGDH ([Daly and Harlow, 1970](#)) model is obtained from the algebraic  
 244 flux model by removing the production terms due to velocity and buoyancy  
 245 gradients. Thus, the equation of turbulent scalar flux (neglecting the  
 246 molecular diffusion) is written as:

$$\overline{C'U'_j} = -C_\theta \frac{k}{\varepsilon} \left( \overline{U'_k U'_j} \frac{\partial \overline{C}}{\partial x_k} \right)
 \tag{11}$$

247 with  $C_\theta = 0.3$  ([Daly and Harlow, 1970](#); [Dehoux, 2012](#)).

- 248 • Simple Gradient Diffusion Hypothesis (SGDH)  
 249 The SGDH model considers that the turbulent flux of scalar is proportional  
 250 to the gradient of scalar, according to Fick's law (the molecular diffusion  
 251 is neglected). It is given by:

$$\overline{C'U'_j} = -\frac{\nu_t}{Sc_t} \frac{\partial \overline{C}}{\partial x_j} \quad (12)$$

252 where  $Sc_t$  is the turbulent Schmidt number ( $Sc_t = 0.7$ ).

## 253 4. Numerical setup for simulations

### 254 4.1. Computational domain

255 To model the experimental site, we first created the geometry using *Salome*  
 256 software ([www.salome-platform.org](http://www.salome-platform.org)), with the goal of reproducing the exper-  
 257 imental field as closely as possible as well as the containers present. Figure 3  
 258 shows an illustration of the computational domain. As the container grid occu-  
 259 pies a footprint of 193m x 171m, we used a domain with a surface area of 240m  
 260 x 240m in order to comply with the best practice rules proposed in the COST  
 261 Action 732 guidebook (Franke et al., 2007; Schatzmann et al., 2010) for urban  
 262 air dispersion modeling. In particular, the report of Franke et al. (2007) men-  
 263 tions that a minimum distance of  $5H_{\max}$  to  $15H_{\max}$  (where  $H_{\max}$  is the height  
 264 of the tallest building in the domain, here equal to 2.54m) between the domain  
 265 boundaries and the area of interest should be considered, depending on the type  
 266 of boundary face (inlet, outlet, top or side face). Regarding the domain height,  
 267 other studies, such as the one conducted by Reiminger (2020), have proposed a  
 268 minimum height of 96m to be used in the absence of obstacles, so this minimum  
 269 distance is also needed for full-scale modeling in the presence of obstacles. The  
 270 domain height also depends on the thermal stratification of the atmosphere and  
 271 the position of the pollutant source. In this work, we will only deal with disper-  
 272 sion cases in neutral, stable and very stable atmosphere with a source located  
 273 at a height lower than two meters from the ground. Given the position of the  
 274 source (quite close to the ground), and since in these atmospheric conditions the  
 275 vertical dispersion of the pollutant is not as strong as in unstable atmosphere, a  
 276 domain height of 50m is sufficient to reproduce completely the pollutant plume.  
 277 The size of the computational domain is therefore 240m x 240m x 50m.

### 278 4.2. Mesh characteristics

279 In order to enable the numerical solution of the Navier-Stokes and concentra-  
 280 tion transport equations on the created domain, it is necessary to discretize the  
 281 computational domain. To do this, the *NETGEN-2D* algorithm of the *salome*  
 282 software is used to generate an unstructured mesh in the ground plane, with  
 283 increasingly finer cells near the containers and increasingly coarser cells as one  
 284 moves away from the container grid. Then, this two-dimensional (2-D) planar

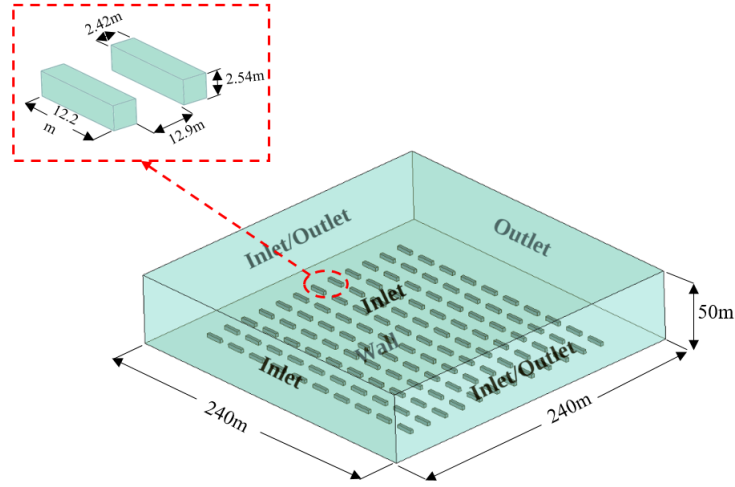


Figure 3: Representation of the computational domain with boundary conditions.

285 mesh is extruded in the vertical direction with a growth factor so that a mesh  
 286 height of 0.25m near the ground and 5.5m at the top of the domain is achieved.  
 287 The total number of cells in the domain is 1,210,582. Figure 4 shows a 3-D rep-  
 288 resentation of the mesh of the computational domain. The mesh is chosen after  
 289 a sensitivity study performed on three mesh sizes for trial No.2692157. Table 4  
 290 summarizes the details of each of the three grid sizes. The results showing the  
 291 sensitivity of the dynamic variables to the mesh size, obtained for three different  
 292 meshes are presented in figures 5 and 6. Overall, the results show no difference,  
 293 except for the first few meters near the ground. Therefore, the reference mesh  
 294 (Mesh2) is chosen for the study.

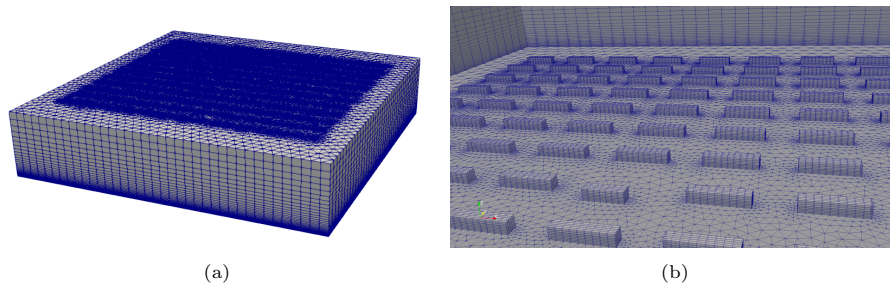


Figure 4: Mesh of the computational domain: (a) External view (b) Internal view.

Table 4: Details of the mesh grids used for the sensitivity study.  $h$  is the height of the first mesh adjacent to the ground wall.

Mesh	Total number of cells	$h$ (m)
Mesh1	1 095 052	0.15
Mesh2	1 210 582	0.25
Mesh3	1 095 052	0.35

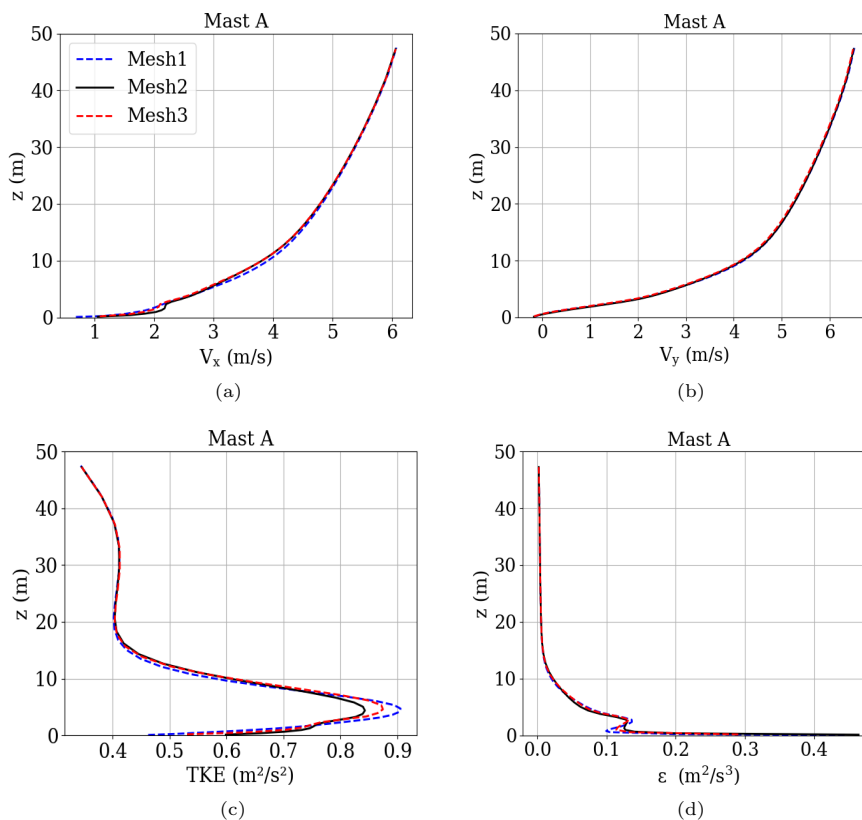


Figure 5: Sensitivity of dynamic variable profiles to mesh size at mast A.

295 *4.3. Boundary conditions and pollutant release condition*

296 In principle, the ABL extends over the entire surface of the earth. However,  
 297 for practical reasons of computational cost and considering the local character  
 298 of the experimental field, we have reduced the extent of the computational  
 299 domain to an area of a few hundred square meters. This computational domain  
 300 is thus an integral part of the ABL and consequently, the conditions applied to  
 301 its borders must ensure continuity with the surrounding ABL. To do this, the  
 302 types of boundary conditions imposed on the faces of the domain are as follows:

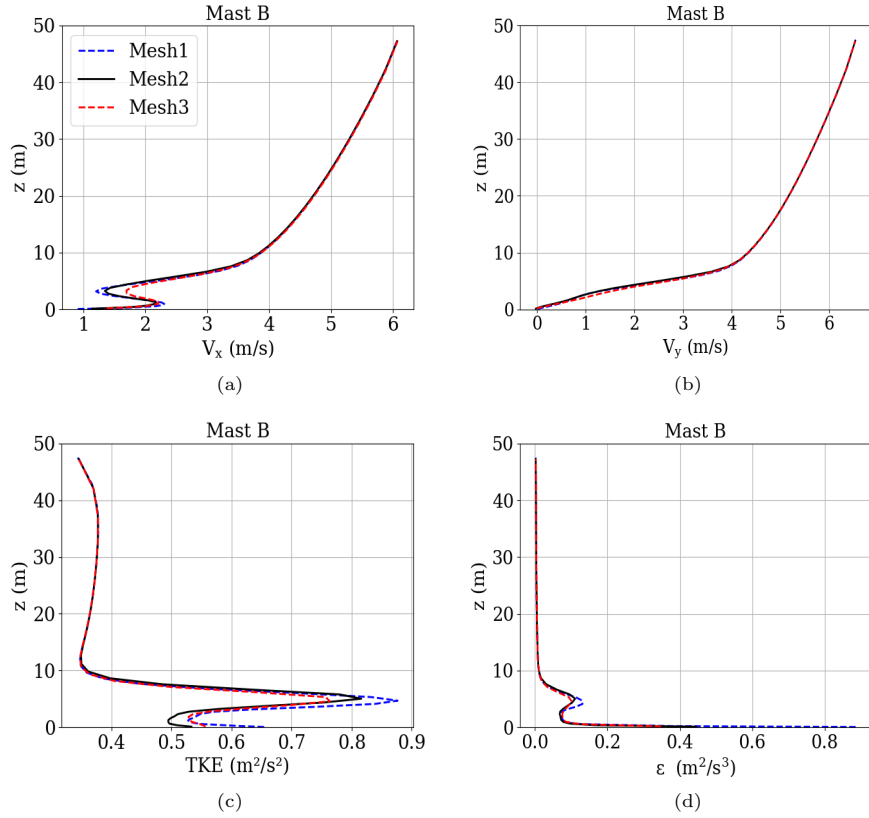


Figure 6: Sensitivity of dynamic variable profiles to mesh size at mast B.

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- At the inlet of the domain: a Dirichlet-type condition is applied, based on velocity, temperature and turbulence profiles whose expressions are developed in the next section.
  - At the outlet of the domain : a Neumann type condition is applied, except for the pressure (free outflow).
  - On the ground and the containers : a logarithmic rough wall function with aerodynamic roughness of 0.04m is applied.
  - On the side faces: automatic inlet or outlet condition is applied depending on the wind direction of the flow.
  - At the top of the domain: an inlet condition (Dirichlet) is imposed, allowing a maximum wind speed to be maintained at the center of the highest cell in the domain. A condition of symmetry is inadequate in this case since the fluxes of heat and non-zero momentum created at ground level will tend towards a zero value at the top. Consequently, this behavior

317 is in disagreement with the physics of the surface boundary layer (SBL)  
 318 which requires a constancy of the flow with the altitude. Figure 3 shows  
 319 the type of boundary conditions on each face of the domain.

#### 320 4.3.1. Input profiles and turbulence models

321 For this work, a series of simulations is carried out on selected trials using the  
 322 atmospheric module of *Code\_Saturne*. The profiles of wind speed, temperature  
 323 and turbulence variables are specified at the input boundary. Table 5 represents  
 324 the dynamic input profiles as well as the models used for each of the selected  
 325 simulations.

326 The inlet wind velocity is well-described by a logarithmic vertical profile,  
 327 derived from the Monin-Obukhov similarity theory (Stull, 1988). It is used to  
 328 determine the mean velocity in the surface layer. It is given by the following  
 329 equation:

$$U(z) = \frac{u^*}{\kappa} \left[ \ln \left( \frac{z}{z_0} \right) - \psi_m(\zeta) + \psi_m(\zeta_0) \right] \quad (13)$$

330 Where  $\kappa$  (0.41) is the von-Kármán constant (Dyer, 1974);  $\zeta$  ( $z/L$ ) is the Monin-  
 331 Obukhov stability parameter;  $\psi_m$  is the momentum stability correction function.  
 332 In neutral stability regime,  $\psi_m = 0$ ; in stable regime,  $\psi_m(\zeta) = -5\zeta$  (Hicks, 1976;  
 333 Businger et al., 1971; Dyer and Hicks, 1970).

334 In all simulations, the Monin-Obukhov vertical temperature profile is set as  
 335 the inflow thermal boundary condition, such as:

$$T(z) = \Theta(z) + z\lambda_{adia} \quad (14)$$

With:

$$\Theta(z) = \Theta(z_r) + \frac{\Theta^*}{\kappa} \left[ \sigma_n \ln \left( \frac{z}{z_r} \right) - \psi_h(\zeta) + \psi_h(\zeta_0) \right]$$

336 Where  $\lambda_{adia}$  ( $-0.009766^\circ C/m$ ) is the adiabatic lapse rate;  $z_r$  is the reference  
 337 (anemometer) level at which  $\Theta$  is determined;  $\Theta^*$  is the temperature scale;  $\psi_h$   
 338 is the stability correction function for potential temperature. In neutral stability  
 339 regime,  $\Theta^* = 0$ ; in stable regime,  $\sigma_n = 0.74$  and  $\psi_h = -5\zeta$  (Hicks, 1976;  
 340 Businger et al., 1971; Dyer and Hicks, 1970).

341 For the trials conducted under stable conditions, the following wind velocity  
 342 profiles are also used:

- 343 • Profile of Gryning et al. (2007) :

$$\frac{U(z)}{u_*} = \frac{1}{\kappa} \left( \ln \left( \frac{z}{z_0} \right) + \frac{5z}{L} \left( 1 - \frac{z}{2z_i} \right) + \frac{z}{L_{MBL}} - \frac{z}{z_i} \left( \frac{z}{2L_{MBL}} \right) \right) \quad (15)$$

344 With :

$$\frac{u_*}{fL_{MBL}} = \left( -2 \ln \left( \frac{u_*}{fz_0} \right) + 55 \right) \exp \left( -\frac{(u_*/fL)^2}{400} \right)$$

- 345 • Profile of [Peña et al. \(2009\)](#) (see also [Konow, 2014](#)):

$$U(z) = \frac{u_*}{\kappa} \left[ \ln \left( \frac{z}{z_0} \right) + b \frac{z}{L} \left( 1 - \frac{z}{2z_i} \right) + \frac{1}{d} \left( \frac{\kappa z}{\eta} \right)^d - \left( \frac{1}{1+d} \right) \frac{z}{z_i} \left( \frac{\kappa z}{\eta} \right)^d - \frac{z}{z_i} \right] \quad (16)$$

346 With

$$\eta = \frac{\kappa z_i}{[d(1+d)^{1/d}]} \left[ \left( \left[ \ln \left( \frac{u_*}{fz_0} \right) - A \right]^2 + B^2 \right)^{1/2} + \ln \left( \frac{z_i}{z_0} \right) \right]^{-1/d}$$

347  $A = 1.5$ ,  $B = 5.2$  et  $d = 1.25$

- 348 • Profile of [Beljaars and Holtslag \(1991\)](#):

$$U(z) = \frac{u_*}{\kappa} \left[ \ln \left( \frac{z}{z_0} \right) - \Psi_m \left( \frac{z}{L} \right) \right] \quad (17)$$

349 With

$$\Psi_m = -a \left( \frac{z - z_0}{L} \right) - b \left( \frac{z}{L} - \frac{c}{d} \right) e^{-d \frac{z}{L}} + b \left( \frac{z_0}{L} - \frac{c}{d} \right) e^{-d \frac{z_0}{L}}$$

350

351  $a = 1$ ,  $b = 0.667$ ,  $c = 5$  et  $d = 0.35$

352

353 Where  $L_{MBL}$  is the length scale in the middle layer of the ABL,  $f_c = 10^{-4} \text{ s}^{-1}$   
 354 is the Coriolis parameter,  $z_i \approx 0.1u_*/f$  is the height of the ABL ([Zilitinkevich](#)  
 355 [and Baklanov, 2002](#)).

356 For the turbulent kinetic energy and its dissipation rate, the following profiles  
 357 are used:

- 358 • Profiles of [Gorlé et al. \(2009\)](#):

359 The evolution of  $k(z)$  with altitude  $z$  is modeled by a square root math-  
 360 ematical relationship whose parameters  $C_1$  and  $C_2$  are determined by a  
 361 least squares method ([Gorlé et al., 2009](#)). The experimental measurements  
 362 to which the  $k(z)$  profile is fitted are those taken by three 2-D horizontal  
 363 sonic anemometers at levels 4, 8 and 16m from a 16m telescopic pneumatic

364 mast located 30.5m upstream of the obstacle array. The input profile of  
 365 turbulent dissipation is then deduced from the smoothing of  $k(z)$ . The  
 366 values of  $C_1$  and  $C_2$ , obtained for each trial, are given in Table ???. The  
 367 expressions of  $k$  and  $\varepsilon$  are as follows:

$$k(z) = \sqrt{C_1 \ln(z + z_0) + C_2} \quad (18)$$

$$\varepsilon(z) = \frac{\sqrt{C_\mu} u_*}{\kappa(z + z_0)} \sqrt{C_1 \ln(z + z_0) + C_2} \quad (19)$$

368  
 369

370 • Profiles of [Kerschgens et al. \(2000\)](#)

371 These profiles are used to model the turbulence in the ABL under stable  
 372 atmospheric conditions. Under these conditions, their analytical expres-  
 373 sions are given by the following relations:

$$k(z) = \frac{1}{2} \left( \overline{U'^2} + \overline{V'^2} + \overline{W'^2} \right) = \frac{1}{2} (\sigma_U^2 + \sigma_V^2 + \sigma_W^2) \quad (20)$$

$$\varepsilon(z) = \frac{u_*^3}{\kappa z} \left( 1 + 4 \frac{z}{L} \right) \quad (21)$$

With :

$$\begin{aligned} \sigma_U &= 2.4u_* \exp\left(-0.3 \frac{z}{z_i}\right), \\ \sigma_V &= 2.0u_* \exp\left(-0.3 \frac{z}{z_i}\right), \\ \sigma_W &= 1.3u_* \exp\left(-0.3 \frac{z}{z_i}\right) \end{aligned}$$

374 As for the turbulence models used for the simulations, the  $k - \varepsilon$  model is  
 375 applied with the SGDh approach. In addition, the  $R_{ij} - \varepsilon$  model is applied  
 376 with the SGDh and GGDh approaches as well as with the DFM model (Table  
 377 5).

378 In order to evaluate the performance of different algebraic models (SGDH  
 379 and GGDH), as well as that of the transport equation (DFM model), we select  
 380 trial No.2692157 to perform the simulations. The results obtained with this trial,  
 381 for modeling scalar concentration flux, are developed and analyzed in section  
 382 5.3 of this article.

383 For each turbulence model used, numerous simulations are performed by  
 384 changing the parameterization of the dynamic profiles (wind speed and turbulent

Table 5: Input profiles and turbulence models used for all simulations on selected trials of the MUST experiment.

Trial No.	Stability condition	Velocity profiles	Temperature profiles	Turbulence profiles	Turbulence models	Scalar fluxes turbulence models
2671934	Very stable	Eqn (15) Eqn (17)	Eqn (14)	Eqn (20 & 21) Eqn (18 & 19)	$k-\varepsilon$ [DU88] $R_{ij}-\varepsilon$ [SA13]	SGDH
2672033	Highly stable	Eqn (15) Eqn (17)	Eqn (14)	Eqn (20 & 21) Eqn (18 & 19)	$k-\varepsilon$ [DU88] $R_{ij}-\varepsilon$ [SA13]	SGDH
2681829	Neutral	Eqn (13)	Eqn (14)	Eqn (18 & 19)	$k-\varepsilon$ [DU88] $R_{ij}-\varepsilon$ [LA75]	SGDH
2692157	Weakly stable	Eqn (13)	Eqn (14)	Eqn (18 & 19)	$k-\varepsilon$ [DU88] $R_{ij}-\varepsilon$ [LA75]	SGDH
2692157	Weakly stable	Eqn (16)	Eqn (14)	Eqn (20 & 21)	$R_{ij}-\varepsilon$ [SA13]	GGDH
2692157	Weakly stable	Eqn (15)	Eqn (14)	Eqn (20 & 21)	$R_{ij}-\varepsilon$ [SA13]	DFM

variables) at the domain entrance each time. For instance, the profile of [Businger et al. \(1971\)](#) is applied with the  $k-\varepsilon$  and  $R_{ij}-\varepsilon$  models (both with SGDH). The profiles of [Na et al. \(2009\)](#) and [Gryning et al. \(2007\)](#) are applied with the model  $R_{ij}-\varepsilon$  (with GGDH and DFM, respectively). As for the turbulent variables, they are computed from the profiles of [Gorlé et al. \(2009\)](#) for the  $k-\varepsilon$  and  $R_{ij}-\varepsilon$  models (both with SGDH), and from the profiles of [Kerschgens et al. \(2000\)](#) for the  $R_{ij}-\varepsilon$  (with GGDH and DFM). In what follows, the analysis of the simulation results will be based only on the turbulence models used in each simulation.

#### Notes:

- Regardless of the type of turbulence model (isotropic or anisotropic) used in the simulations, the meteorological turbulence input profiles are those that describe the evolution of kinetic energy, isotropically, with altitude. In other words, in the case of the  $R_{ij}-\varepsilon$  model, only the diagonal terms of the  $R_{ij}$  tensors are considered to construct the TKE, and the other terms are considered to be zero. We assume that this is a reasonable assumption for operational applications.
- For trials involving stable atmospheric conditions, thermal turbulent flux is always modeled using the SGDH hypothesis, while concentration turbulent flux is modeled using different algebraic models (SGDH or GGDH) or

405 the concentration scalar transport equation (DFM model). Consequently,  
406 only the effect of the model from the point of view of dispersion is stud-  
407 ied, focusing on the passive scalar, and the study excludes the influence  
408 of thermal turbulent flux.

#### 409 4.3.2. Pollutant release condition

410 Regarding the release condition, the pollutant is injected through a scalar  
411 source term. In the injection cell, the pollutant flow rate in L/min (Table 1) is  
412 imposed for the studied trials.

#### 413 4.4. CFD solver options

414 The simulations are performed in stationary mode in order to solve the wind  
415 and dispersion fields in the domain. Thus, two calculations have been imple-  
416 mented. The first calculation is used to determine the dynamic field, i.e. only the  
417 wind field, and stops when the steady state is reached. The second calculation is  
418 used to compute the dispersion of the passive scalar, in this pre-calculated wind  
419 field with frozen velocity, pressure and temperature. To ensure convergence of  
420 the results, the accuracy of the solver is set to  $10^{-4}$  for all the computed vari-  
421 ables. The spatial numerical scheme for each variable (convection scheme) is  
422 set automatically by the solver. The Semi-Implicit Method for Pressure-Linked  
423 Equations (SIMPLE) algorithm is used for the pressure-velocity coupling with  
424 a reduction of the relaxation coefficient.

## 425 5. Results and discussion

### 426 5.1. Comparison between $k - \varepsilon$ and $R_{ij} - \varepsilon$ models: dynamic field in neutral 427 condition

428 The neutral case is chosen to study the effect of the anisotropic turbulence  
429 model  $R_{ij} - \varepsilon$  both on the atmospheric flow and on the dispersion of the passive  
430 tracer. For this purpose, the scalar turbulent flux is modeled only by the SGDH  
431 model.

432 Figure 7 illustrates the velocity and TKE fields obtained using both the  $k - \varepsilon$   
433 and  $R_{ij} - \varepsilon$  models at  $z = 4m$ . Figure 8 shows the lateral velocity and turbulent  
434 kinetic energy fields, obtained with the  $k - \varepsilon$  and  $R_{ij} - \varepsilon$  models at  $x = 130m$ .  
435 We find that the presence of the containers significantly influences the flow, even  
436 at heights greater than the height of a container ( $H = 2.54$  m). The container  
437 grid acts as a rough surface, causing a deceleration of the flow downstream and  
438 amplifying turbulence generation through the shearing of the mean flow and the  
439 wake interference produced behind each container.

440 By comparing the two turbulence models, we find that the  $R_{ij} - \varepsilon$  model  
441 generates relatively low levels of turbulence in the near field of the container  
442 grid as well as in the far field (Figure 8b). Analysis of the numerical results

443 for the dynamic field revealed the recirculation phenomena in the flow, which  
 444 both models ( $k - \varepsilon$  and  $R_{ij} - \varepsilon$ ) effectively capture at the rear of the containers.  
 445 However, the length of the recirculation as well as the size of the recirculation  
 446 bulbs are larger with the  $k - \varepsilon$  model (result not shown here).

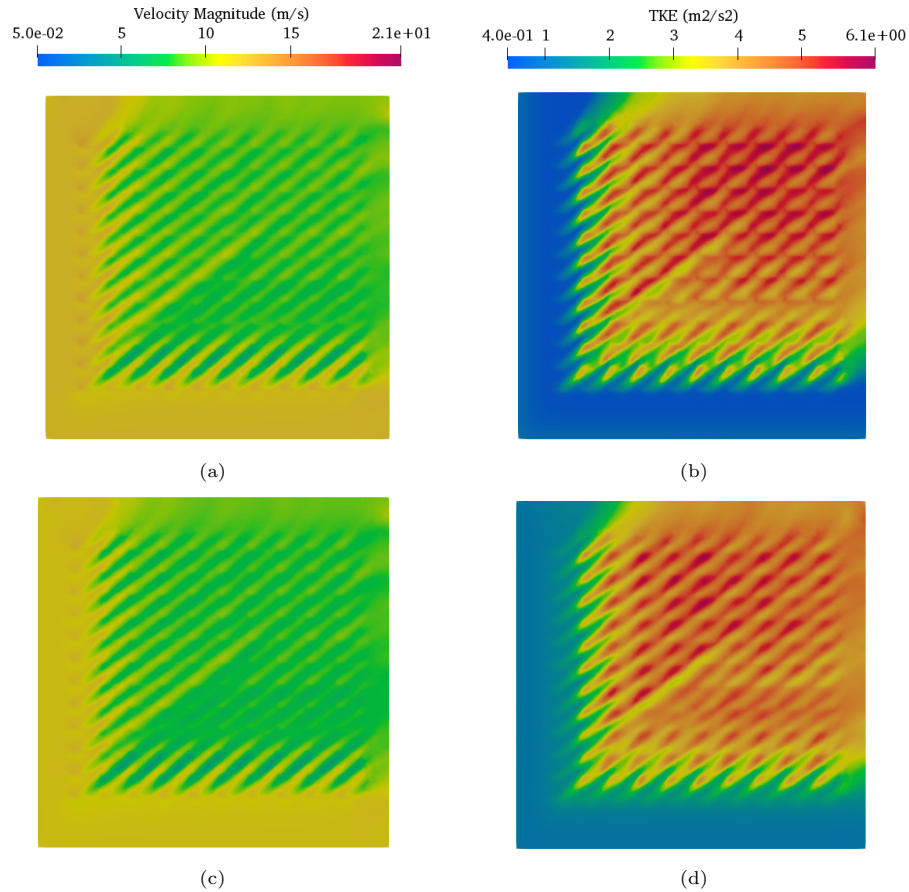


Figure 7: Horizontal wind speed (left) and TKE (right) fields at  $z = 4m$  for the neutral atmospheric stability trial No. 2681829 using the  $k - \varepsilon$  model (a, b) and the  $R_{ij} - \varepsilon$  model (c, d).

447 Figures 9 and 10 show the velocity and TKE profiles, respectively, on masts  
 448 A and D, and tower T. At masts A and D, the velocity and TKE profiles are  
 449 essentially identical between the  $k - \varepsilon$  and  $R_{ij} - \varepsilon$  models and in agreement  
 450 with the available experimental measurement points. The peak production of  
 451 the TKE, a few meters above the canopy, is consistent with the TKE contours  
 452 observed in Figure 8b, for both turbulence models. However, the small number  
 453 of measurement points (two measures) does not guarantee a good correlation  
 454 between the numerical and experimental profiles over the whole height of the

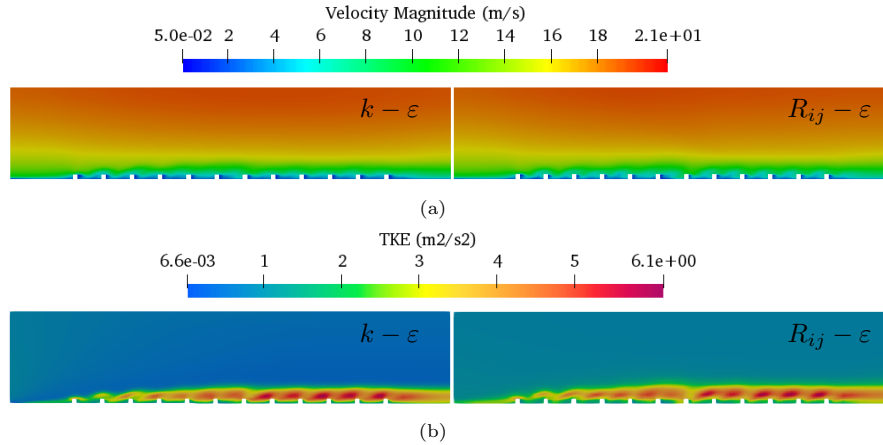


Figure 8: Lateral dynamic fields at  $x = 130m$  (a) wind speed and (b) TKE for the neutral atmospheric stability trial No.2681829 with  $k - \epsilon$  (left) and  $R_{ij} - \epsilon$  (right) models.

455 computational domain. On the central tower for which we have four mea-  
 456 surement points, we observe an overestimation of the velocity by the numerical  
 457 models and an underestimation of the TKE compared to the experimental data.

458 *5.2. Comparison between  $k - \epsilon$  and  $R_{ij} - \epsilon$  models: dispersion field in neutral*  
 459 *condition*

460 The pollutant release in this trial is from a 1.8m high source with a flow rate  
 461 of 225L/min. Figure 11 represents the pollutant concentration field at ground  
 462 level for the  $k - \epsilon$  and  $R_{ij} - \epsilon$  models.

463 From a qualitative standpoint, we have noted a minor impact of the turbu-  
 464 lence model on pollutant dispersion. The plume described by the  $k - \epsilon$  model  
 465 appears to be slightly larger in comparison to the one predicted by the  $R_{ij} - \epsilon$   
 466 model. This distinction is primarily due to the slight overestimation of TKE  
 467 by the  $k - \epsilon$  model, as depicted in Figure 10. Since the turbulent fluctuations  
 468 are larger, the turbulent concentration flux is also larger. Therefore, we ob-  
 469 serve a more extensive plume with the  $k - \epsilon$  model. We can also note that the  
 470 plume guideline does not coincide with the wind direction ( $-41^\circ$  with respect  
 471 to the x-axis). This is due to the presence of a channeling effect imposed by  
 472 the containers on the plume, making it deviate from the wind direction. This  
 473 plume deflection phenomenon has also been observed by Milliez and Carissimo  
 474 (2007); Kumar et al. (2015); Bahlali et al. (2019) etc., who explained it by the  
 475 channeling effect that the containers impose on the plume, causing it to deflect  
 476 from the wind direction. In other words, the aspect ratio of the containers also  
 477 influences plume dispersion.

478 Figure 12 shows the concentration profile over all 48 receptors, with the

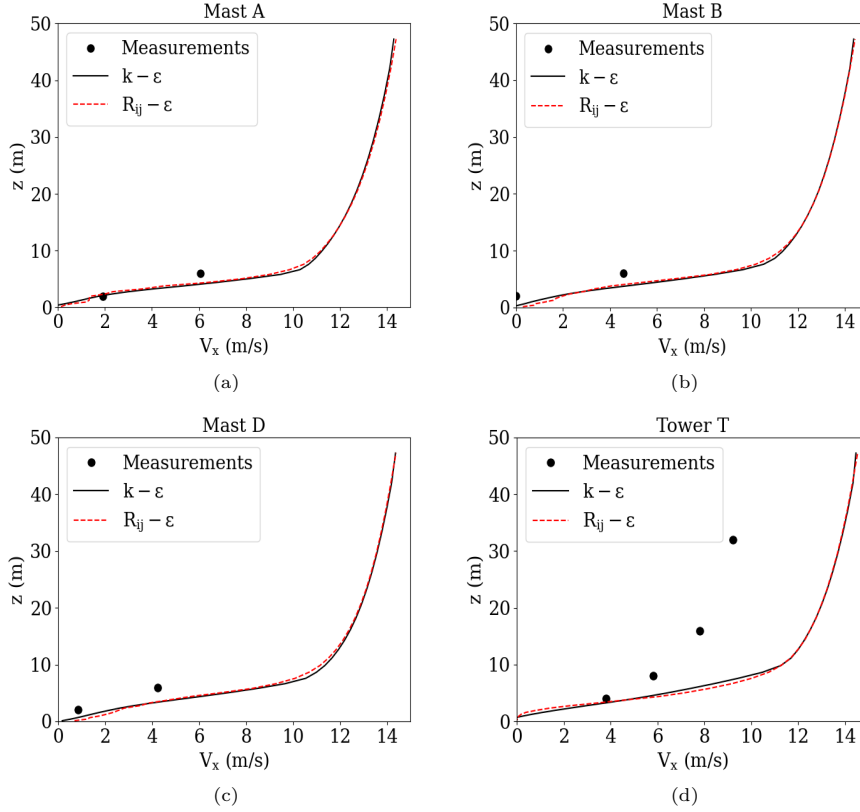


Figure 9: Comparison between the observed velocity and the one calculated with the  $k - \epsilon$  and  $R_{ij} - \epsilon$  models for the neutral case No.2681829 at the position of (a) Mast A (b) Mast B (c) Mast D and (d) Tower T.

479 last eight (41-48) placed vertically on the central tower T. For both models,  
 480 we observe a main concentration peak measured at receptor No.8, followed by  
 481 four secondary peaks recorded by the receptors placed in the direction of the  
 482 plume. As the plume moves away from the source, it is diffused and its concen-  
 483 tration decreases. This is the reason why the concentration at the secondary  
 484 peaks evolves in a decreasing way. This phenomenon is also observed in the  
 485 experiment, but with a main peak located at receptor No.9, followed by four  
 486 secondary peaks. The correspondence between the numerical and experimental  
 487 results is globally satisfactory for the position of the peaks. It can be noted  
 488 that the underestimation of the concentration increases with distance from the  
 489 source.

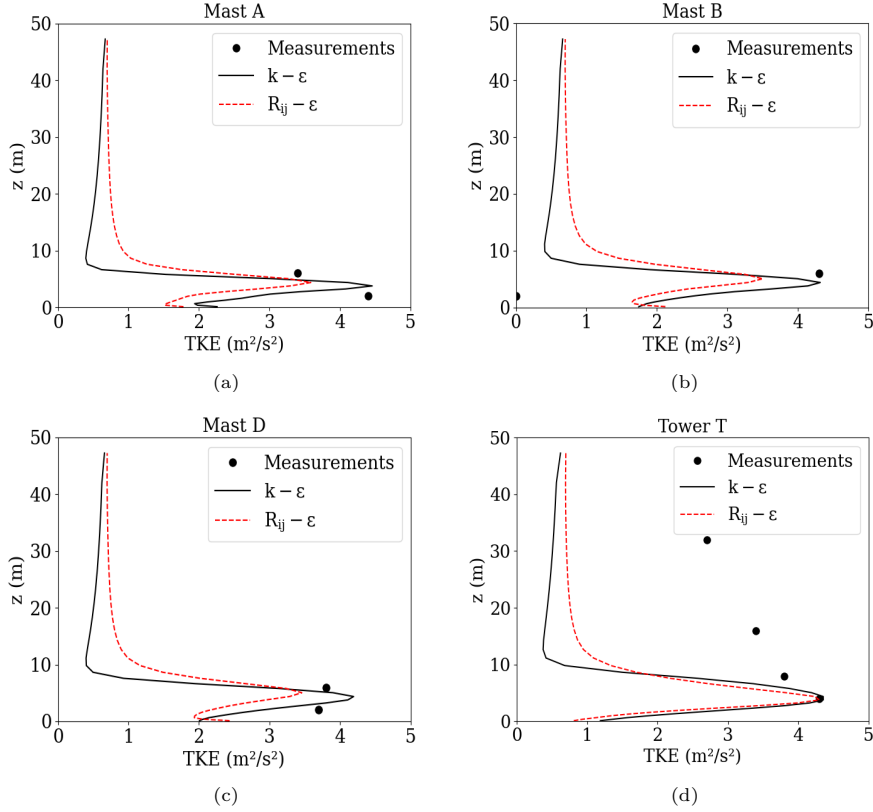


Figure 10: Comparison between the observed TKE and the one calculated with the  $k - \epsilon$  and  $R_{ij} - \epsilon$  models for the neutral case No.2681829 at the position of (a) Mast A (b) Mast B (c) Mast D and (d) Tower T.

490 *5.3. Evaluation of anisotropic models for the stable case*

491 *5.3.1. Dynamic field*

492 Figure 13 shows the velocity profiles computed at different masts using the  
 493 CFD code and compared to experimental observations. We note that the ve-  
 494 locity profiles at the masts vary slightly when the velocity and turbulence input  
 495 profiles are varied. However, the modification of the turbulence models leads to  
 496 more noticeable variations of the velocity profiles, in particular in the first five  
 497 meters above the ground, where the  $R_{ij} - \epsilon$  model leads to a variation of the sign  
 498 of the vertical velocity gradient. Since this model is anisotropic, it is supposed  
 499 to better capture the physical phenomena present in the near wall (ground and  
 500 containers). However, due to the small number of experimental measurements  
 501 between 0 and 6 meters from the ground, it is quite difficult to identify if this  
 502 change in sign of the gradient captured by the  $R_{ij} - \epsilon$  model reflects the reality  
 503 of the experiment.

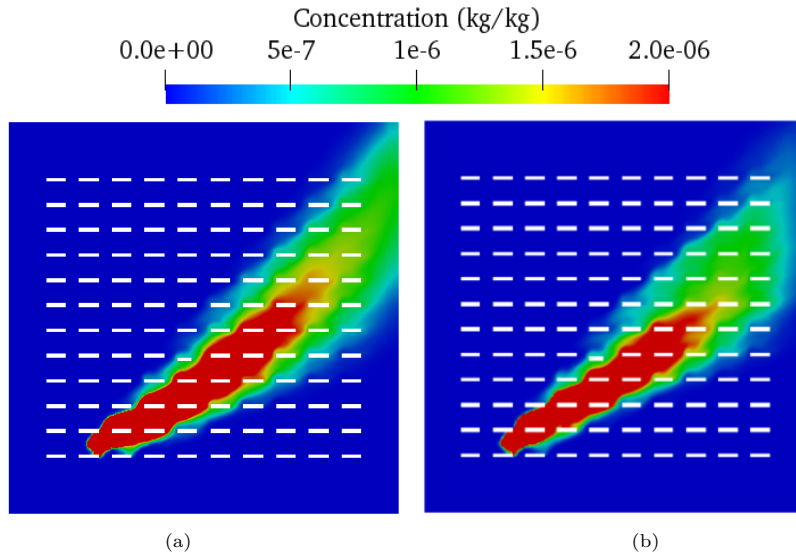


Figure 11: Average pollutant concentration field (kg/kg) at ground level ( $z = 1.6m$ ), with the models (a)  $k - \epsilon$  (b)  $R_{ij} - \epsilon$  (With SGDH) for Trial No.2681829.

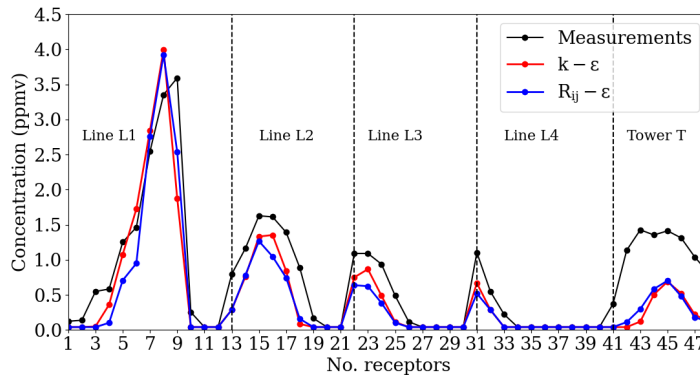


Figure 12: Comparison between the observed concentration profile and the one calculated using the  $k - \epsilon$  and  $R_{ij} - \epsilon$  models across all PID sensors for Trial No.2681829.

504 Figure 14 shows the TKE profiles calculated at different masts using the CFD  
 505 code, in comparison with the experimental measurements. A first observation  
 506 can be made on the shape of the profiles, which is more or less the same, from  
 507 a qualitative point of view. The profiles show the presence of three large zones  
 508 in the flow:

- 509 • From 0m to about 5m above the ground (almost 2 times the height of the  
 510 containers): the vertical magnitude of the TKE is positive. Turbulence

511 is generated by the flow's shearing action against the ground and the  
512 containers' walls, making it a zone of turbulence production.

- 513 • From 5m to 20m: the vertical gradient of the TKE is negative, the turbu-  
514 lent kinetic energy decreases gradually under the action of the viscosity of  
515 the air. It is thus a zone of dissipation of the TKE.
- 516 • From 20m to 50m: the vertical gradient of TKE is very weak (almost null  
517 in certain cases). The production and dissipation of turbulence compen-  
518 sate each other, thus the turbulence is in equilibrium.

519 Moreover, we notice that the  $k-\varepsilon$  model tends to predict an overproduction  
520 of the turbulence. For this model, the turbulence peak on all masts is the  
521 highest as well as the vertical gradient of TKE in the production area (steeper  
522 slope than that with the  $R_{ij}-\varepsilon$  model). This result highlights the difficulty  
523 of the  $k-\varepsilon$  model in predicting the anisotropy of the turbulence. The  $R_{ij}-\varepsilon$   
524 model, being anisotropic, takes better into account the presence of the ground  
525 and the containers and allows to obtain a profile of the TKE which fits better  
526 the experimental measurements.

### 527 5.3.2. Dispersion field

528 Figure 15 shows the predicted mean concentration field at ground level with  
529 the  $k-\varepsilon$  and  $R_{ij}-\varepsilon$  models. Note that the plume size is significantly larger when  
530 employing the  $k-\varepsilon$  model, except in the case of the  $R_{ij}-\varepsilon$  model combined  
531 with DFM. This can be explained by the fact that the  $k-\varepsilon$  model predicts an  
532 overproduction of turbulence (Figure 14) close to the ground which increases  
533 the turbulent transport flux and thus leads to a larger lateral diffusion of the  
534 pollutant over the extent of the grid. As in the case of neutral stratification, we  
535 observe that the directional line of the plume does not coincide with the wind  
536 direction ( $43^\circ$  from x-axis).

537 Figure 16 shows the comparison between observed and predicted concentra-  
538 tions at the receptors using different turbulence models. We observe a main  
539 peak followed by four secondary peaks. We also find that the concentration  
540 intensity at the main peak (7ppmv) is about twice as high as the main peak  
541 observed in trial No.2681829 with a neutral stability condition (3.5ppmv). This  
542 increase in concentration is primarily due to the atmospheric conditions under  
543 which the trial is conducted. Since trial No.2692157 is conducted under mod-  
544 erately stable atmospheric conditions, turbulence generation is attenuated by  
545 thermal stratification of the air, which reduces turbulent transport of the pollu-  
546 tant; therefore, concentrations are locally higher. We also observe, in this case,  
547 a good match between the numerical results and the measurements concerning  
548 the position of the peaks.

549 As for the magnitude of the highest peak, all the used models overestimate  
550 it, except the  $R_{ij}-\varepsilon$  model (with GGDH) which is the only one to predict the

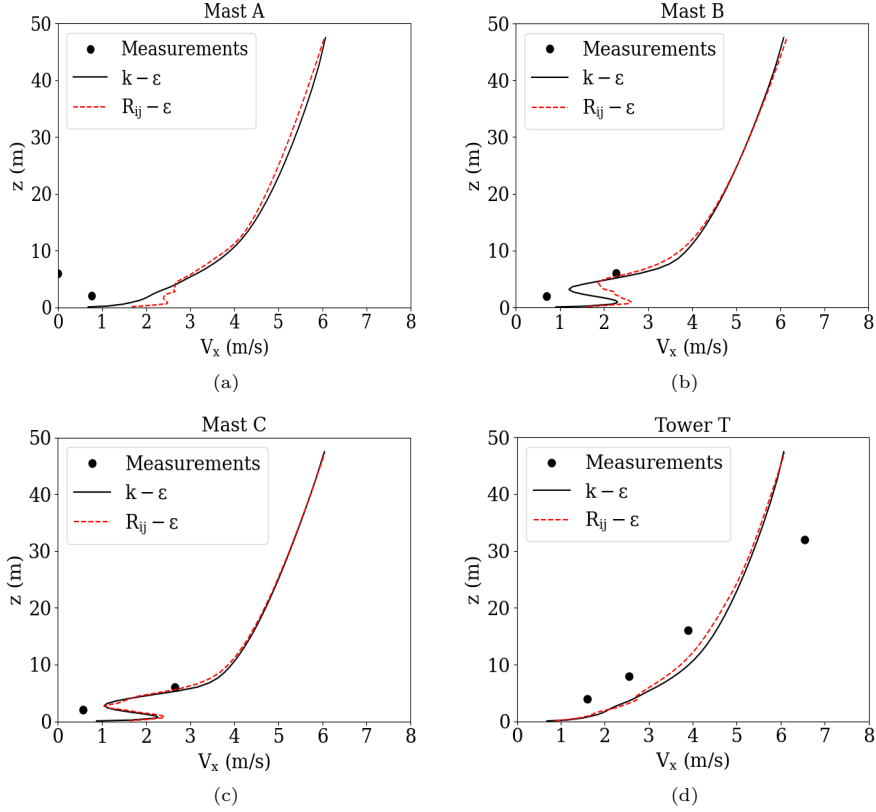


Figure 13: Comparison between the observed velocity and the one calculated with the  $k - \epsilon$  and  $R_{ij} - \epsilon$  models for the stable case No.2692157 at the position (a) Mast A (b) Mast B (c) Mast C and (d) Tower T.

551 maximum concentration measured by the third receptor. On the other hand,  
 552 moving away from the source, this same model seems to underestimate the  
 553 concentration the most among the other models. In general, the predictions  
 554 obtained by all the models are close to the experimental measurements when  
 555 moving away from the source. However, no model is able to correctly reproduce  
 556 the profile corresponding to the concentrations observed by receptors 41 to 48  
 557 on the central tower.

558 The statistical assessment involves the measured concentrations from 40  
 559 receptors distributed along four sampling lines, as well as eight additional recep-  
 560 tors located on central tower T, along with the corresponding average concen-  
 561 trations predicted by the simulations. The performance measures calculated in  
 562 this study are the normalized mean square errors (NMSE), the fractional bias  
 563 (FB), the correlation coefficient (COR), mean geometric bias (MG), the geo-  
 564 metric variance (VG), and the fractional accuracy factor 2 (FAC2), which is the

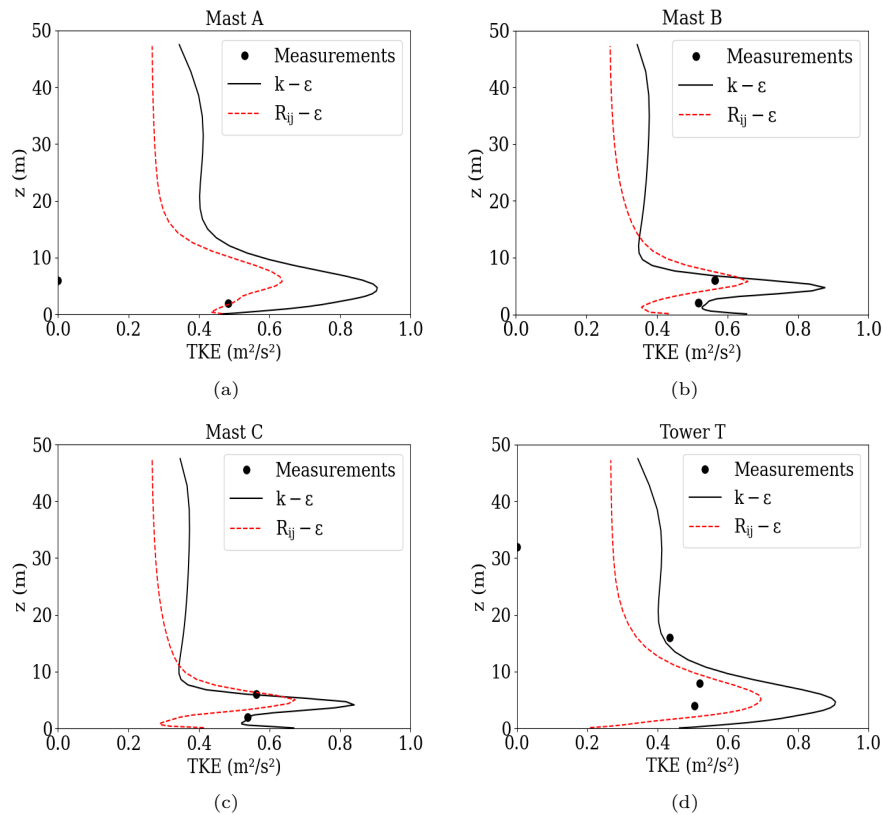


Figure 14: Comparison between the observed TKE and the one calculated with the  $k - \varepsilon$  and  $R_{ij} - \varepsilon$  models for the stable case No.2692157 at the position of (a) Mast A (b) Mast B (c) Mast C and (d) Tower T.

565 fraction of predictions within a factor of two of field measurements. The results  
 566 of the statistical study, to assess the accuracy of each model, are presented in  
 567 Table 6. A model is considered perfect if:  $NMSE = FB = 0$  and  $COR = FAC2$   
 568  $= MG = VG = 1$  (Chang and Hanna, 2004). A model is considered sufficiently  
 569 efficient if it satisfies the following criteria:  $-0.3 < FB < 0.3$ ,  $0.7 < MG < 1.3$ ,  
 570  $NMSE < 4$ ,  $VG < 1.6$  and  $FAC2 \geq 0.5$  (Chang and Hanna, 2004). We recall  
 571 that these criteria are applicable to dispersion within urban areas.

572 All FB values are positive, indicating that all models tend on average to  
 573 underestimate the concentration, with the  $k - \varepsilon$  model being the least under-  
 574 estimating. Examination of the MG values in Table 6 shows that no model  
 575 succeeds in satisfying the acceptance criterion, with the  $k - \varepsilon$  model being the  
 576 closest to the tolerable range. The same is true for the VG values. As for the  
 577 NMSE values, they are all within the accepted range. Therefore, all models  
 578 perform well for the statistical measures (FB and NMSE), but appear to have

579 a systematic bias at low concentrations based on the values for the logarithmic  
580 measures (MG and VG). All models succeed in having more than 50% of the  
581 predicted concentrations within a factor of two (FAC2) of the observations. The  
582  $k - \varepsilon$  model is the best performing with 75% of the predictions within a factor  
583 of two of the observations. As for the  $R_{ij} - \varepsilon$  model, it performs significantly  
584 worse with about 63%, 60% and 69% with the SGD, GGDH and DFM models,  
585 respectively.

586 Finally, good one-to-one correlations can be observed for all models through  
587 the high values of the correlation coefficient, the highest of which (COR =  
588 0.92) is for the  $R_{ij} - \varepsilon$  model (with GGDH and DFM). Comparing the  $k - \varepsilon$   
589 and  $R_{ij} - \varepsilon$  models, we find that the latter, known to be anisotropic, does  
590 not necessarily improve the predictions over the former, regardless of the scalar  
591 turbulent flux model with which it is associated. However, a study by [Bahlali  
592 et al. \(2019\)](#) shows that using the  $R_{ij} - \varepsilon$  model provides a more physical rep-  
593 resentation of the influence of obstacles on concentration profiles. The three  
594 velocity-scalar correlation models (SGD, GGDH, and DFM), when employed  
595 alongside the  $R_{ij} - \varepsilon$  model, achieve a score of three (indicating that three sta-  
596 tistical measures meet their criteria) in the statistical evaluation, and therefore,  
597 all three models can be considered for retention. The GGDH hypothesis, which  
598 uses the Reynolds stress tensor does not improve the predictions compared to  
599 the isotropic SGD model (lower FAC2 and more under-prediction). In fact,  
600 the GGDH model is obtained by considering assumptions of weak equilibrium  
601 ( $\frac{D\overline{U_i'\theta'}}{Dt} = 0$ ) and dynamic and thermal equilibrium of turbulence, compared to  
602 the DFM model. In addition, the production terms due to the velocity gradient  
603  $\overline{U_j'\theta'}\frac{\partial\overline{U_i}}{\partial x}$  and buoyancy  $g_i\overline{\Theta'^2}$  (Refer to Equation 10) are suppressed in the for-  
604 mulation of [Daly and Harlow \(1970\)](#). This model is not suitable for flows with  
605 natural convection and buoyancy, and does not give satisfactory results in the  
606 case of very strong stratification ([Dehoux, 2012](#)). By analogy, the same applies  
607 to atmospheric flows under stable stratification.

608 The DFM model, on the other hand, is more suitable for stratified at-  
609 mospheric flows and gives better results than the GGDH model (less under-  
610 prediction and higher FAC2). This is probably due to the presence of a gravity  
611 term with scalar variance in the scalar turbulence transport equation.

Table 6: Statistical performance measures evaluated for trial No.2692157.

	FB	MG	NMSE	VG	FAC2(%)	COR
$k - \varepsilon + \text{SGD}$	0.13	1.39	0.37	2.23	75	0.91
$R_{ij} - \varepsilon + \text{SGD}$	0.27	1.66	0.45	3.3	62.5	0.90
$R_{ij} - \varepsilon + \text{GGDH}$	0.5	1.81	0.59	2.98	60.4	0.92
$R_{ij} - \varepsilon + \text{DFM}$	0.37	1.86	0.45	4.47	68.8	0.92
Fluidyn-Panache	0.2	1.28	0.34	1.5	87.5	0.91

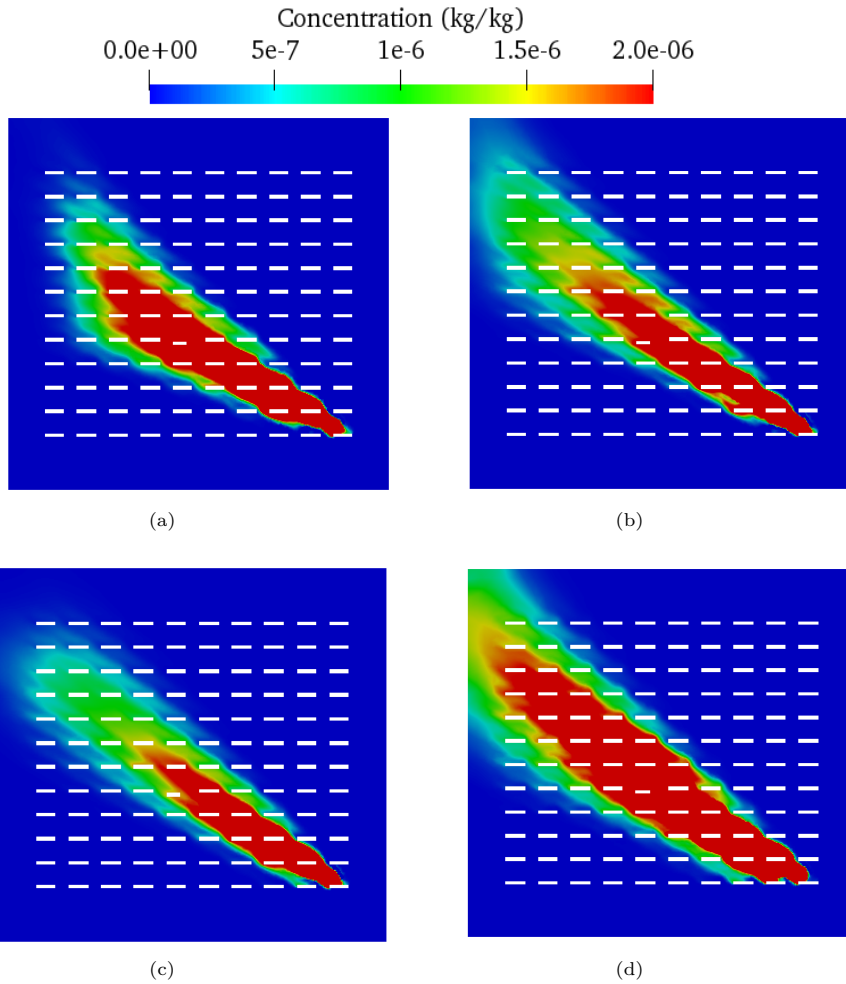


Figure 15: Average pollutant concentration field (kg/kg) at  $z = 1.6m$ , with models (a)  $k - \varepsilon$  (with SGDH) (b)  $R_{ij} - \varepsilon$  (with SGDH) (c)  $R_{ij} - \varepsilon$  (with GGDH) and (d)  $R_{ij} - \varepsilon$  (with DFM) for trial No.2692157.

#### 612 5.4. Overall model performance

##### 613 5.4.1. Quantitative evaluation of Code\_Saturne models

614 In this section, we present a statistical evaluation study of the model for the  
 615 selected set of simulated trials. Figure 17 shows the scatter plots of observed and  
 616 predicted concentrations using the  $k - \varepsilon$  and  $R_{ij} - \varepsilon$  models. Table 7 presents  
 617 the values of the statistical indicators evaluated for all trials with the  $k - \varepsilon$  and  
 618  $R_{ij} - \varepsilon$  models.

619 In general, the CFD model underestimates the observations, this underes-  
 620 timation being slightly more pronounced when the anisotropic model is used.

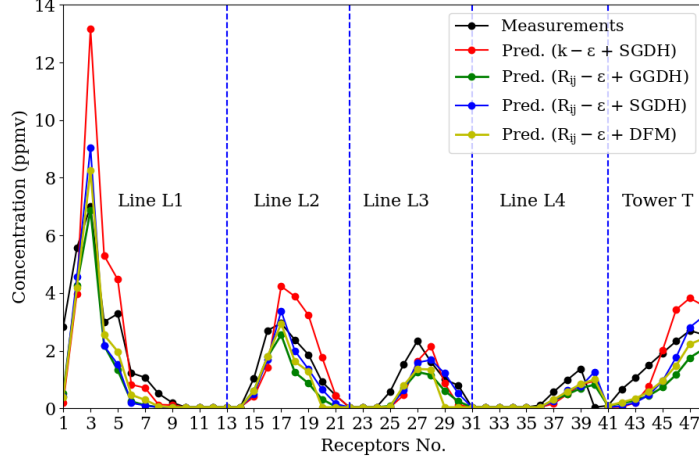


Figure 16: Comparison between the observed concentration profile and the one calculated with the  $k - \varepsilon$  and  $R_{ij} - \varepsilon$  models on all the PID receptors for trial No.2692157.

621 This underestimation is illustrated by the values of FB in Table 7 and is clearly  
 622 highlighted in Figure 17 which shows that the scatter plot tends to spread out  
 623 below the acceptable zone of FAC2.

624 Regarding the values of the statistical indicators presented in Table 7, it can  
 625 be seen that the numerical calculations predict with a good precision (correlation  
 626 higher than 0.8) the physical phenomena present during these experimental  
 627 trials. As for the average values of MG and VG, they exceed the acceptable  
 628 margins whatever the turbulence model used. However, their values remain  
 629 close to the threshold when a  $k - \varepsilon$  is used. The  $R_{ij} - \varepsilon$  model, although a  
 630 second-order model, does not differ significantly from the  $k - \varepsilon$  (first-order)  
 631 model in terms of modeling dispersion for this trial. Both models obtain a score  
 632 of 3, which means that the criterion for only three of the six statistical measures  
 633 is satisfied.

634 The analysis of the results of the stable trials, in particular the two tri-  
 635 als No.2671934 (with an Obukhov length  $L=5.8m$ ) and No.2672033 ( $L=4.8m$ ),  
 636 allows to make several observations:

- 637 • For trial No.2671934, no turbulence model is able to perfectly predict the  
 638 position and the amplitude of the maximum concentrations, even if the  
 639 shape of the predicted concentration profile is similar to the one observed  
 640 in the trial (profile not shown here). Consequently, none of the models  
 641 used achieves acceptable performance (score  $\geq 3$ ). Overall, a shift in peak  
 642 position and an underestimation of the concentration measured by the re-  
 643 ceptors are observed across this profile, as shown in Figure 17. This shift  
 644 in concentration values is not detected with the trial No.2672033, even

645 though it has a more stable atmospheric stability than trial No.2671934.  
646 However, this last one took place in the presence of wind at a speed  
647 (1.6m/s) lower than that of the first one (2.69m/s) at the entrance of the  
648 domain. This finding confirms the conclusions of [Degrazia et al. \(2011\)](#)  
649 that the usual Eulerian RANS dispersion models are not suitable for very  
650 stable atmospheric conditions for two main reasons: (i) these models have  
651 limitations in terms of modeling when the advection velocity in the trans-  
652 port equation approaches very low values ( $U < 2\text{m/s}$ ) due to the meander-  
653 ing phenomenon; (ii) these models take into account pollutant transport  
654 by convection, molecular and turbulent diffusion, but not the large-scale  
655 phenomena that are dominant during low wind sequences. This may also  
656 explain the large underestimation of concentration values observed by re-  
657 ceptors placed far from the source.

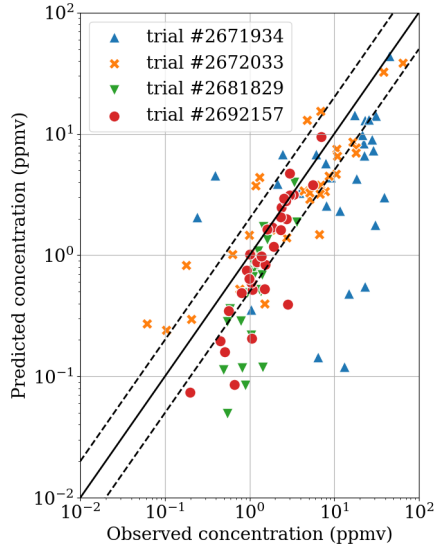
- 658 • It appears from the predictions of the trial No.2672033 that the use of the  
659 constants suggested by [Sadek et al. \(2013\)](#) for the  $R_{ij} - \varepsilon$  model (retained  
660 as a result) significantly modifies the concentration profile obtained with  
661 the standard model and allows to obtain results closer to the experimental  
662 observations.

#### 663 5.4.2. Comparison with numerical results of the Fluidyn-Panache code

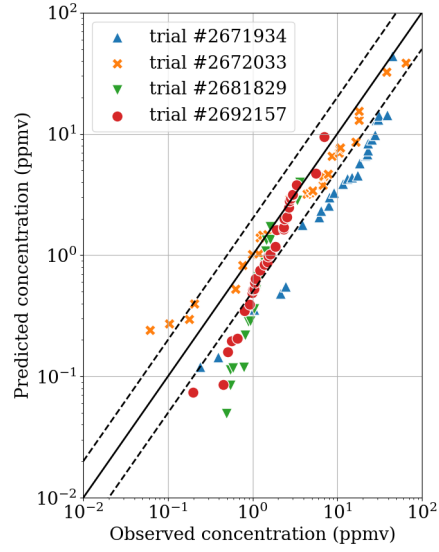
664 In another study on the MUST experiment, [Kumar et al. \(2015\)](#) carried  
665 out an evaluation of CFD simulation results using the Fluidyn-Panache code  
666 ([Fluidyn-Panache, 2010](#)), which consists of a three-dimensional flow model and  
667 an Eulerian dispersion model. Many combinations of wind inflow boundary con-  
668 ditions for the wind and turbulence profiles were considered for the simulations.  
669 The turbulence model employed in the simulations is the isotropic  $k - \varepsilon$  model,  
670 which includes a modified value of the  $C_{\varepsilon 3}$  coefficient to account for buoyancy  
671 effects under stable atmospheric conditions. In order to compare the evaluation  
672 results of the present study, statistical measures of performance are computed  
673 for trial No.2692157 (Table 6) as well as for the 4 trials on average (Table 7).  
674 Only the best performing combinations, i.e. the simulations with the highest  
675 score, are taken into account in this statistical study.

676 For trial No.2692157, we find that the Fluidyn-Panache model provides  
677 equally good predictions as the *Code\_Saturne* model and that the criteria for  
678 all six statistical indicators are met. However, an even higher percentage of  
679 predicted concentrations (87.5%) are within a factor of two of the observations  
680 (Table 6).

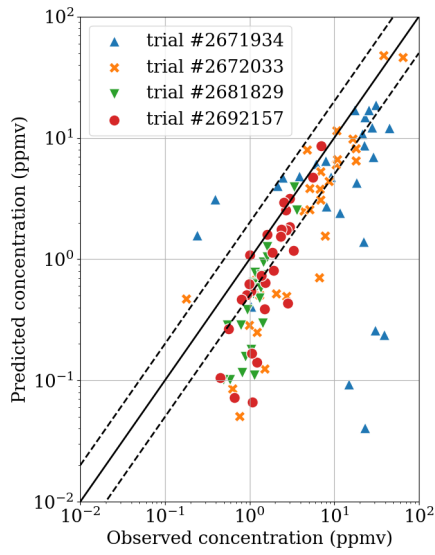
681 For all trials, the Fluidyn-Panache model predicts 77.1% of the points in a  
682 FAC2 of the observations, while the present model simulates 61.6%, 58.8% of  
683 the points in a FAC2 of the observations with the  $k - \varepsilon$  and  $R_{ij} - \varepsilon$  models,  
684 respectively. As for the FB values, only the Fluidyn-Panache model satisfies  
685 the criterion (0.138) while the present model underestimates the concentration



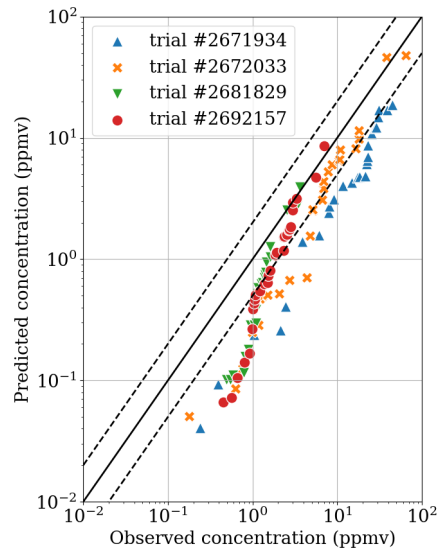
(a)



(b)



(c)



(d)

Figure 17: Scatter plots (on the left) and q-q plots (on the right) illustrating observed and predicted concentrations using the  $k - \epsilon$  model (a,b) and the  $R_{ij} - \epsilon$  model (c,d) for the four trials.

686 values with a value of 0.473 with the  $k - \varepsilon$  model and 0.528 with  $R_{ij} - \varepsilon$ . The  
 687 values of NMSE for the simulated concentrations at all receptors for the four  
 688 cases are respectively 1.513 and 1.493 with the  $k - \varepsilon$  and  $R_{ij} - \varepsilon$  models. These  
 689 values are almost 3.75 times higher than the value (0.41) reported by Kumar  
 et al. (2015) with the Fluidyn-Panache model.

Table 7: Statistical performance measures assessed for all trials studied.

	FB	MG	NMSE	VG	FAC2(%)	COR
<i>Code_Saturne</i> , $k - \varepsilon$	0.473	1.815	1.513	30.92	61.6	0.833
<i>Code_Saturne</i> , $R_{ij} - \varepsilon$	0.528	2.215	1.493	106	58.8	0.835
Fluidyn-Panache, $k - \varepsilon$	0.138	1.118	0.41	1.895	77.1	0.933

690

## 691 6. Conclusions

692 This study presents 3-D simulations of the dispersion of a pollutant under  
 693 neutral and stable atmospheric conditions. The *Code\_Saturne* model is evalu-  
 694 ated with concentration measurements obtained from four trials of the MUST  
 695 experimental campaign in a complex urban field. One of these tests is per-  
 696 formed under neutral conditions, while the other tests are performed under  
 697 relatively stable to very stable atmospheric conditions. Simulations are done  
 698 using two turbulence models,  $k - \varepsilon$  and  $R_{ij} - \varepsilon$  LRR, and scalar turbulent fluxes  
 699 are evaluated using SGDH and GGDH algebraic models, as well as the differ-  
 700 ential transport equation (DFM). Qualitative and quantitative analyses show  
 701 that the performance of the models is reasonably acceptable in terms of cap-  
 702 turing multiple concentration peaks and the distribution of the hourly average  
 703 concentration.

704 Numerical results from simulations on the moderately stable test show that  
 705 the turbulence models have an overall underestimation of the mean concen-  
 706 tration. The comparison between the  $k - \varepsilon$  and  $R_{ij} - \varepsilon$  models shows that  
 707 the former tends to generate an overproduction of turbulence near the ground,  
 708 resulting in a significantly wider plume spread. The isotropic model ( $k - \varepsilon$ )  
 709 predicts approximately 75% of the concentration values within a factor of two,  
 710 whereas the  $R_{ij} - \varepsilon$  model predicts 62.5% and 60.4% when used with the SGDH  
 711 and GGDH models, respectively. The use of the differential transport equation  
 712 (DFM) to solve scalar turbulent fluxes improves predictions with nearly 69% of  
 713 the concentration values within a factor of two (FAC2) of the observations. It  
 714 also provides a better correlation with experimental measurements, compared  
 715 to algebraic models such as SGDH and GGDH. Regardless of its computational  
 716 cost, this model proves to be very adequate to study atmospheric dispersion in  
 717 complex terrain with moderate to low stable stability conditions.

718 In all simulation cases, the CFD model consistently underestimates the ob-  
 719 servations. Overall, when using the isotropic first-order  $k - \varepsilon$  model, the CFD  
 720 model predicts a slightly higher percentage of concentrations within a factor of  
 721 two (FAC2) of the observations. Additionally, it is noteworthy that the model's

722 performance tends to decrease with lower wind speeds, regardless of atmospheric  
723 stability.

724 Further studies will be carried out on the different algebraic models and on  
725 the scalar transport equation, in order to better evaluate the efficiency of these  
726 models, especially under extremely stable atmospheric conditions. It might also  
727 be beneficial to fine-tune the various constants employed in the formulation of  
728 these models to enhance their compatibility with dispersion-related challenges.

729 As the Reynolds stress tensor has been used in its isotropic form, relying  
730 solely on the tensor trace or kinetic energy, this may explain the discrepancies  
731 obtained with the  $R_{ij} - \varepsilon$  model. Therefore, future investigations will include  
732 a re-evaluation of calculations using this model by essentially incorporating the  
733 other components of the Reynolds tensor into the boundary conditions, particu-  
734 larly the momentum flux  $\overline{U'W'}$ , to better account for anisotropy and enhance  
735 model performance.

736 As the approach used to model thermal turbulent flux is the simple gradient  
737 model (SGDH), it will be necessary to test other algebraic models or even the  
738 transport equation to verify the effect of thermal turbulence anisotropy on the  
739 turbulent flux of the concentration scalar.

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