

# Exploration of changes in the chemical composition of sedimentary organic matter and the underlying processes during biodegradation through advanced analytical techniques

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1	Exploration of changes in the chemical composition of
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#### 25 Abstract

26 Rationale. Although the scientific community widely investigated the organic matter 27 biodegradation processes, only a limited number of studies have explored the molecular 28 changes of this material while its structure, composition, and origin play a key in these 29 processes. Methodology. In this context, we decided to examine the effects of the 30 biodegradation on the chemical composition of sedimentary organic matter and to explore the 31 underlying mechanisms. We conceived a laboratory-based degradation experiment utilizing 32 organic-rich sediments artificially composed of two contrasting organic matter end-members 33 (i.e., soil and algae) under two oxygen conditions. The sediment samples before and after 34 incubation were then analyzed via laser desorption/ionization Fourier transform ion cyclotron 35 resonance mass spectrometry for molecular characterization and via thermally assisted 36 hydrolysis and methylation gas chromatography-mass spectrometry in order to offer insights 37 into the mechanisms driving biodegradation processes. Results. Our results from molecular 38 characterization unveiled distinct pathways of biodegradation contingent upon the source 39 material. Moreover, they hinted at a predilection for altering high molecular weight compounds 40 like lignin/CRAM and CAS compounds, manifesting as a conversion into lower molecular 41 weight counterparts. Furthermore, the complementary findings from biomarker analyses 42 underscored the influence of environmental factors - specifically oxygen conditions and 43 microbial communities - on organic matter decomposition. **Discussion.** Although this study is 44 a controlled laboratory experiment and more studies are needed, it demonstrated the intricate 45 interplay among chemical, biological, and environmental factors, profoundly shaping the 46 reactivity of organic matter. This study underscores the critical need for persistent inquiry, 47 aimed at unraveling the factors and conditions governing the diverse pathways of biodegradation. 48

49 Keywords: Sedimentary organic matter; Biodegradation; Molecular characterization; LDI FT–
50 ICR MS; Biomarkers.

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#### 52 **1. Introduction**

53 Sediments operate as sources and sinks of nutrients and pollutants in aquatic systems. 54 They also represent a large reservoir of organic matter (OM) from diverse sources (e.g., terrestrial, aquatic, estuarine, marine, etc.) and in various proportions according to the 55 56 environment, region, and hydrology. Indeed, sedimentary OM is derived from aquatic material 57 such as algae, bacteria, plankton, macrophytes, and nekton formed in situ, but also receives 58 terrestrial OM such as soil, vascular plants, leaves, root exudates and anthropogenic OM such 59 as soil OM, which is easily transported from the upstream catchment into rivers through 60 hydrological processes, finally ending up in sediments (Briand et al. 2015; Jeanneau et al. 2018; 61 van der Meij et al. 2018). Sediments are also a reactive compartment where diagenetic 62 processes occur, inducing physical, chemical, and/or biological changes to the sedimentary OM 63 (Henrichs 1992; Kuznetsova et al. 2019; Milliken 2003). Among the diagenetic processes, 64 biodegradation plays a key role in OM biogeochemistry as it is one of the main processes 65 causing changes in its amount, composition, and properties in aquatic systems (Moran et al., 66 2000; Arndt et al., 2013; Sankar et al., 2019; Derrien et al., 2019a).

OM biodegradation is based on the common preconception that OM is represented as a mix of recalcitrant materials usually associated with terrestrial origin and labile materials associated with aquatic OM (Hansell 2013). OM biodegradation is often oversimplified and/or restricted to preferential degradation of labile and/or low molecular weight (LMW) organic compounds leading to the rapid loss of them rather than the recalcitrant material like lignin and high molecular weight (HMW) organic compounds (Curtis-Jackson et al. 2009; Hach et al. 2020; Hansen et al. 2016; Moran and Zepp 1997; Ward et al. 2019). However, since the early 74 2000s, new paradigms on the stability of terrestrial OM and more specifically on soil OM have been proposed (Wershaw 2004). They conceptualized soil OM stabilization as the result of a 75 76 combination of physical factors controlling its accessibility (Lützow et al. 2006) as well as 77 chemical factors such as its oxidizing degree controlling the energy available for 78 microorganisms (LaRowe and Van Cappellen 2011). Even recalcitrance as an intrinsic property 79 of molecules has been questioned arguing that under favorable environmental conditions, all 80 organic molecules would be biodegradable (Kleber 2010). These paradigms are in line with 81 what was observed in several studies where a preferential conversion of high molecular weight 82 (HMW) compounds into reactive and bioavailable LMW compounds were observed 83 (Kellerman et al. 2015; Kothawala et al. 2020; Liu et al. 2019). In the literature, biodegradation 84 processes have also been associated with the production of HMW aromatic material through 85 alteration and recombination of existing compounds (e.g., large organic biomolecules or 86 humification processes) (Hach et al. 2020; Kleber and Lehmann 2019; Lee et al. 2016). These 87 various scenarios of biodegradation processes and our difficulties in finding a consensus on the 88 biodegradation mechanisms and pathways are related not only to the characteristics of OM 89 itself such as composition and sources but also to the characteristics and properties of the 90 ecosystem (i.e., physiochemical conditions and biological communities) (D'Andrilli et al. 91 2019; Kothawala et al. 2020; McCallister et al. 2018). Furthermore, priming effects represent 92 a potentially important mechanism underlying the biodegradation of OM in aquatic systems 93 complexifying our understanding (Bianchi 2011; Bianchi and Ward 2019; Zhuang et al. 2021). 94 Priming effects refer to quantitative and qualitative changes in the microbial decomposition of 95 recalcitrant OM upon the addition of labile OM (Guenet et al. 2014, 2010a). They are 96 delineated as negative, neutral, or positive effects depending on the consequences in the ways 97 that they may reduce or exert no effect on or stimulate OM biodegradation (B Guenet et al., 98 2010; Gontikaki et al., 2013; Ward et al., 2016, 2019; Bengtsson et al., 2018; Derrien et al.,

99 2019c). All these discrepancies in the observed biodegradation mechanisms coupled with the 100 no consensus on the occurrence and role of priming effects result in a critical need for more 101 studies focused on the biodegradation processes and mechanisms through approaches 102 combining analyses providing quantitative and qualitative information (D'Andrilli et al. 2019). 103 Most of the biodegradation studies in the literature aim at quantifying the OM degradation and 104 mineralization, and its kinetic through microbial biomass measurement, dissolved organic 105 carbon or CO<sub>2</sub> concentrations/fluxes (Bowen et al. 2009; Kalbitz et al. 2003; Moran et al. 2000; 106 Sankar et al. 2019; Vähätalo et al. 2010). Only a few studies target the qualitative aspect of the 107 biodegradation processes (i.e., exploration of the molecular changes) while the structure and composition of OM play a key in these processes (Arndt et al. 2013; Herzsprung et al. 2017; 108 109 Kellerman et al. 2015; Kothawala et al. 2020). Yet, investigating these molecular changes in 110 similar conditions will not only allow us to lighten our knowledge of biodegradation concepts 111 but also provide a chemical fingerprint that helps the understanding of OM reactivity and 112 metabolism. Additionally, the few studies exploring the molecular composition during 113 biodegradation are limited to the dissolved fraction while two fractions exist (i.e., particulate 114 and dissolved) and both are intimately linked to each other (Derrien et al. 2019a). Hence, we 115 decided to tackle this knowledge gap by investigating the compositional changes of 116 sedimentary OM during biodegradation in this study.

Due to the high heterogeneity and complexity of the OM composition related to its dynamics, involvement in a large number of biogeochemical processes, and the type of environment, we have decided to head toward a laboratory experiment for this study. Indeed, controlled laboratory experiments offer the benefits of providing in-depth insight into the effects of selected parameters have on OM composition, in contrast with field samples that only represent a snapshot in space and time. The controlled laboratory degradation experiment was conducted using organic-rich sediments artificially composed of two contrasting OM end124 members (i.e., soil and algae) at known mixing ratios and under two oxygen conditions (oxic 125 and anoxic). The molecular changes induced by biodegradation in sediments and the 126 mechanisms behind them were investigated through two advanced analytical tools: laser desorption/ionization Fourier transform ion cyclotron resonance mass spectrometry (LDI FT-127 128 ICR MS) and thermally assisted hydrolysis and methylation gas chromatography-mass 129 spectrometry (THM–GC–MS). Currently, high-resolution mass spectrometry techniques such 130 as FT-ICR MS are likely considered as the most powerful tool used to characterize the OM 131 molecular properties (Kellerman et al. 2015; Valle et al. 2018). The LDI FT-ICR MS technique 132 was chosen as it represents a direct and highly sensitive analysis at the molecular level to 133 explore the biodegradation-induced changes in chemical composition in sedimentary OM. The 134 second technique was selected to explore the biomarker fingerprints to obtain clues to 135 understand the underlying processes during biodegradation (Meyers and Ishiwatari 1993; 136 Wakeham and Ertel 1988). Both techniques were also selected as they present the advantage 137 of being able to directly analyze artificial sediments without any pre-treatment step such as 138 chemical extraction and thus limiting modifications of the sample (Aubriet and Carré 2019). 139 The main goal of this study was to provide a detailed exploration of the molecular composition 140 changes of sedimentary OM induced by biodegradation and evaluate the role of factors such as 141 type of sources and oxygen on it. The study specifically aimed at answering the following 142 questions: (i) what molecules or groups of molecules are the most affected by biodegradation 143 processes, (ii) what is the potential effect of OM source or oxygen conditions (i.e., oxic versus 144 anoxic) on biodegradation processes? and (iii) what are the underlying processes behind 145 biodegradation?

146

#### 147 **2. Material and methods**

#### 148 2.1. Experimental design of the incubation experiment

149 Two contrasted end-members (i.e., soil and algae) were mixed at soil-to-algae ratios 150 (S:A) of 100:0, 75:25, 50:50, 25:75, and 0:100, based on the organic carbon (OC) 151 concentrations (Derrien et al. 2019b). In natural aquatic systems, ratios between both OM 152 sources greatly vary according to the season and the hydrology. Therefore, we decided to 153 perform the study at several ratios. The soil end-member (topsoil, 0-10cm) was collected at Bukhansan National Park, South Korea (37°43'37.0" N 127°00'50.9"), a fully forested area 154 155 composed of broadleaved trees such as *Ouercus mongolica*, *O. serrata*, the coniferous tree, 156 Pinus densiflora, and the Robinia pseudoacacia plantation (Choi and Yang 2012). A 157 commercial unicellular green alga (Chlorella vulgaris), which is commonly found in 158 freshwaters and marine environments, was purchased from Aquanet Co., Ltd. in 159 Gyeongsangnam-do, South Korea, and used as the algae end-member. Further details regarding 160 the end-members and the experiment can be found in Derrien et al. 2019c. Briefly, incubation 161 experiments in oxic and anoxic conditions were performed in pre-washed and pre-combusted 162 (450 °C for 4 hours) 125 ml Wheaton® amber glass bottles with Teflon screwcaps. Samples 163 for anoxic incubation were prepared and sampled using a sterile Aldrich® AtmosBag two-hand 164 glove bag under nitrogen atmosphere. A mass (12 g) of end-member mixed samples were 165 mixed with ultrapure water (Barnstead<sup>™</sup> Easypure<sup>™</sup> RoDi, Thermo Scientific) at a solid-to-166 solution ratio of 1:4 and equilibrated for 48 h in the dark at room temperature after shaking at 167 100 rpm for one hour. River water from the Jungnang River, a tributary of Seoul's Han River 168 (37°40'16"N, 127°04' 47"E), was used as inoculum for microbial incubation. Before 169 incubation, all artificial sediment samples were spiked with 3% (v/v) of the prepared inoculum, 170 and a sufficient amount (1%, v/v) of nutrients (NH<sub>4</sub>NO<sub>3</sub> and K<sub>2</sub>HPO<sub>4</sub> at 10 mM) was added to 171 each sample to avoid nutrient limitations during the incubation. The low inoculum-to-sample ratio had a negligible effect on the final OC concentrations of the end member mixtures (<1%) 172 173 and on the molecular signature. Therefore, we concluded that the concentration and signature

of the inoculum could be safely ignored. Finally, the samples in oxic and anoxic conditions were incubated in the dark at 25 °C for 60 days (Guenet et al. 2014; Navel et al. 2012). Samples were collected for sampling on day 0 and day 60, with "day 0" corresponding to the day when the samples were inoculated. Degradation experiments were carried out in duplicate.

178 The sampling was performed as follows. First, samples were taken from the incubator, 179 and the overlying water was carefully removed so as not to disturb the sediment. The artificial 180 sediment samples were then centrifuged at 5000 rpm for 20 min to remove the porewater. 181 Finally, the porewater-free sediment samples were freeze-dried, ground, and homogenized for 182 further analyses. All the analyses were performed on both replicates except for the LDI FT-183 ICR MS where both replicates were carefully mixed in order to obtain a representative sample 184 for the different modalities. Total organic carbon (TOC) contents of sediments were measured 185 using an elemental analyzer coupled with an isotope ratio mass spectrometer (EA-IRMS; 186 EuroEA-Isoprime IRMS, GV instruments, UK) for all samples in order to estimate the 187 percentage of removal or production of OC after the 60 days of incubation (Supplementary 188 Material, Table A.1).

189 2.2. Laser desorption ionization Fourier transform ion cyclotron resonance mass spectrometry
190 (LDI FT–ICR MS)

191 Molecular analyses were performed with a (-) LDI solariX 2XR FT ICR mass 192 spectrometer (Bruker Daltonic, Bremen, Germany) in Mass Spectrometry Based Convergence 193 Research Center at Kyungpook National University. The instrument was equipped with a 7 T 194 refrigerated actively shielded superconducted magnet and a dual-mode ESI/matrix-assisted 195 laser desorption/ionization ion source with a frequency-tripled Nd:YAG (neodymium-doped 196 yttrium aluminum garnet) laser emitting at 355 nm. Briefly, between 2 and 5 µg of freeze-dried 197 artificial sediment samples were directly loaded onto preplaced double-sided tape on a MALDI 198 (matrix-assisted laser desorption/ionization) plate. Before analysis, the plate was dried under vacuum for 2 hours. The broadband spectra were acquired between m/z 100 and 1000, with a 4MW free induction decays (2 Ω), and 300 spectra were averaged to improve the signal-tonoise ratio. For data processing, only peaks with an S/N ratio of ≥ 4 were considered. A detailed description of the analytical method is given in Solihat et al. (2019).

203 Molecular formulae were assigned to peaks in the mass range of m/z 200-600 with a 204 mass error range of  $\pm$  1ppm. Molecular formulae were identified considering the following elemental combinations:  ${}^{12}C_{0-\infty}{}^{14}H_{0-\infty}{}^{16}O_{0-\infty}{}^{14}N_{0-5}{}^{32}S_{0-2}$ . Only the formulae presented the 205 206 following elemental ratio criteria,  $0.3 \le H/C \le 2.25$ , O/C < 1.2, N/C < 0.5, and S/C < 0.2, and 207 the double bond equivalent rule (i.e.,  $DBE = 1/2(2C-H+N) \ge 0$ ) was considered (Koch et al. 208 2007). Several selected indices, including intensity-weighted average (wa) molecular masses, 209 elemental ratios, DBE, AImod (modified aromatic index), and MLB<sub>L</sub> (molecular lability 210 boundary for labile contributions) were also calculated (D'Andrilli et al. 2015; Koch and 211 Dittmar 2006). Finally, the assigned molecular formulae were examined using the van 212 Krevelen diagram and were categorized into eight different compound classes (Supplementary 213 Material, Table A.2) (Derrien et al. 2017; Kim et al. 2003).

214 2.3. Thermally assisted hydrolysis and methylation (THM) gas chromatography-mass
215 spectrometry (THM-GC-MS)

216 THM-GC-MS analyses were performed according to Jeanneau et al. (2014). Briefly, we introduced between 0.3 mg (algae) and 5 mg (soil) of freeze-dried solid residue into an 80 217 218 µL aluminum reactor with an excess of solid tetramethylammonium hydroxide (ca. 10 mg). 219 The THM reaction was performed on-line using a vertical micro-furnace pyrolyzer PZ-2020D 220 (Frontier Laboratories, Japan) operating at 400 °C. The products of this reaction were injected 221 into a gas chromatograph GC-2010 (Shimadzu, Japan) equipped with a SLB 5MS capillary 222 column in the split mode (60 m  $\times$  0.25 mm ID, 0.25  $\mu$ m film thickness). The temperature of 223 the transfer line was 321°C, and the temperature of the injection port was 310 °C. The oven 224 was programmed to maintain an initial temperature of 50 °C for 2 minutes, then rise to 150 °C at 15 °C min<sup>-1</sup>, and then rise to 310 °C at 3 °C min<sup>-1</sup> where it stayed for 14 minutes. Helium 225 226 was used as the carrier gas at a flow rate of 1.0 ml/min. Compounds were detected using a 227 QP2010+ mass spectrometer (Shimadzu, Japan) operating in full scan mode. The temperature 228 of the transfer line was set at 280 °C, the ionization source at 200 °C, and molecules were 229 ionized by electron impact using an energy of 70 eV. The list of analyzed compounds and m/z230 ratios used for their integration are given in the Supplementary Material (Table A.2). 231 Compounds were identified based on their full-scan mass spectra by comparison with the NIST 232 library and with published data (Nierop and Verstraten 2004; Nierop et al. 2005).

233 The target compounds (i.e., 136 biomarkers) were classified into eight chemical 234 families: small organic acids (SOAc), phenolic compounds (PHE) including lignin and tannin 235 markers, carbohydrates (Car), low molecular weight fatty acids (LMWFA), high molecular 236 weight fatty acids (HMWFA), fatty alcohols (Falcohol), chlorophyll markers (Phytol) 237 including phytol and neophytadiene, and N-containing compounds (Ncompound). The peak 238 area of the selected m/z (mass/charge) for each compound was integrated and corrected by a 239 mass spectra factor calculated as the reciprocal of the proportion of the fragment used for the 240 integration and the entire fragmentogram provided by the NIST library (Supplementary 241 Material, Table A.3) (Denis et al. 2017; Jeanneau et al. 2018). The proportion of each 242 compound class was calculated by dividing the sum of the areas of the compounds in this class 243 by the sum of the peak areas of all analyzed compounds expressed as a percentage.

- 244 2.4. Statistical analyses
- 245

- All the statistical analyses were performed with R (version 3.4.3) using the packages 246 Vegan (Vegan 2.4-1) and FactoMineR.
- 247 2.4.1. LDI FT–ICR MS data

The Bray-Curtis dissimilarity matrix was computed for the whole set of the LDI FT– ICR MS data (Supplementary Material, Table A.4 and A.5) to quantify the intra- and intervariability of the samples (Bray and Curtis 1957). Principal coordinates analysis (PCoA) was then performed based on this dissimilarity matrix.

252 2.4.2. THM–GC–MS data

253 A first principal component analysis (PCA) was performed using the relative 254 proportions of the 136 biomarkers as variables, which allows the direct comparison of the 255 different samples without the concentration effect (Supplementary Material, Table A.6 and 256 A.7). This PCA allows identification of the correlated variables based on a modulus of the 257 Pearson coefficient > 0.8. When the two variables were correlated, the least abundant was 258 removed. Then a second PCA was performed. The variables with a correlation lower than 0.5 259 with the two first factors were removed, resulting in a new set of 67 variables. A third PCA 260 and the associated hierarchical ascendant classification (HAC) were calculated. A last PCA 261 and HAC were performed using specific microbial biomarkers (i.e., microbial FA: nC12:0, 262  $nC_{13:0}$ ,  $nC_{14:0}$ ,  $aC_{15:0}$ ,  $nC_{15:0}$ ,  $brC_{16:0}$ ,  $nC_{16:1}$ ,  $nC_{17:0}$ ,  $nC_{18:1}$  (2 isomers), and  $nC_{19:1}$ ) in order to 263 explore the role of the microbial activity during biodegradation.  $nC_{16:0}$  and  $nC_{18:0}$  were not 264 considered since these biomarkers are produced by the THM of algae and derive from higher 265 plant inputs, respectively.

266

267 **3. Results and discussion** 

All the detailed results of LDI FT–ICR MS and THM–GC–MS are presented in Supplementary
Material (Tables A.4-7),

270 **3.1.** Identification of the biodegradation-induced changes in the pure end-member artificial

271 *sediments* 

272 At day 0, no differences were observed based on the H/C and O/C ratios, but a clear 273 distinction was observed for the distribution between the CHO, CHON, CHOS, and CHONS 274 elemental formulae, common indices, and compound classes between both pure artificial 275 sediments (S:A, 100:0 and S:A, 0:100) (Table A4-5). S:A, 100:0 mainly contained carbon-rich 276 and heavily carboxylated compounds with a part of them rich in nitrogen (e.g., CHO of 55.2% 277 and 39.3% and CHON of 33.6% and 38.4% in oxic and anoxic conditions respectively). By 278 contrast, S:A, 0:100 exclusively consisted of CHO compounds (e.g., 98.9% and 96.5% in oxic 279 and anoxic conditions respectively). Higher values of DBE and AI<sub>mod</sub> and a much lower value 280 for the MLB<sub>L</sub> for S:A, 0:100 than for S:A, 100:0 indicated the presence of highly aromatic and 281 unsaturated compounds in the composition of the algae and the near absence of labile 282 compounds. This was supported by the identification and distribution of the compounds classes 283 as the S:A, 0:100 mainly consisted of unsaturated hydrocarbons (e.g., 36.3% and 37.5% for 284 oxic and anoxic conditions, respectively), lignins/CRAM (e.g., 30.7% and 30.4% for oxic and 285 anoxic conditions, respectively) and condensed aromatic structures (CAS, e.g., 30.1% and 286 27.2% for oxic and anoxic conditions, respectively) (Fig. 1). Algae material is commonly 287 associated with labile material, however, it was also demonstrated that this kind of material can 288 be characterized by a high amount of hydrocarbons (Youngblood et al. 1971), lignins (Alzate-289 Gaviria et al. 2020; Martone et al. 2009) and by biopolymers such as the algaenans, which represent the major constituent of the cell walls of green algae, for instance (Derenne et al. 290 291 1992; Gelin et al. 1996). S:A, 100:0 also presented a distribution in compounds class rich in

CAS (e.g., 51.2% and 37.2% for oxic and anoxic conditions, respectively) and lignins/CRAM (e.g., 31.1% and 38.2% for oxic and anoxic conditions, respectively). However, the rest of the distribution consisted of lipids (e.g., 7.1% and 7.3% for oxic and anoxic conditions, respectively), unsaturated hydrocarbons and/or tannins (e.g.,  $\sim$ 5%), and a small amount of labile material such as proteins (e.g.,  $\sim$ 3%) and carbohydrates and amino sugars (e.g,  $\sim$ 2%) (Fig. 1).

298 After 60 days of incubation, the main observed differences were related to the 299 distribution in compound classes. A decrease of the highest molecular weight compounds (i.e., 300 CAS and Lignins/CRAM) in S:A, 0:100 was observed in favor of an increase of the percentage 301 of unsaturated hydrocarbons of 10.8 and 21.7 points with values reaching 47.1% and 592.2% 302 for oxic and anoxic conditions, respectively (Fig. 1). A slight increase in labile compounds was 303 also observed. Regarding the S:A, 100:0, the few changes were mainly observed in oxic 304 conditions, with a decrease in the observed percentage of the CAS (e.g., from 51.2% to 38.9% 305 at day 0 and day 60, respectively). In parallel, a slight increase in the percentage of the 306 lignins/CRAM and labile compounds was also noted for both oxygen conditions. Classically, 307 biodegradation is associated with a rapid loss of low molecular weight (LMW) compounds 308 (Hansen et al. 2016; Wetzel 2001). However, in the present study, the results suggest 309 preferential biodegradation of the high molecular weight (HMW) compounds associated with 310 conversion into lower molecular weight compounds (Curtis-Jackson et al. 2009; Hach et al. 311 2020; Liu et al. 2019; Moran and Zepp 1997). These observations agree with the paradigm in 312 which all organic molecules could be biodegradable irrespective of their recalcitrance if 313 favorable physical, chemical, and biological factors are present (Kleber 2010). This 314 observation can be also the result of two-step processes. As it was previously observed in the literature, LMW compounds could have been consumed primarily generating a rapid loss of 315 316 labile and LMW materials followed by the consumption of HMW compounds (e.g., CAS and lignins/CRAM) and their transformations in LMW compounds (D'Andrilli et al. 2019; Hansen
et al. 2016). Unfortunately, the sampling design does not allow us to identify which processes
(one or two steps processes) occurred.

320 Regarding the biomarker distributions, two distinctive fingerprints were identified between the 321 2 pure end-member artificial sediments but closely the same irrespective of the oxygen 322 condition (Fig. 2). In S:A, 100:0, the most abundant chemical family was HMWFA with 323 aliphatic side chain containing 20 and more C atoms ( $60.6 \pm 3.3\%$  and  $60.0 \pm 7.4\%$  for oxic 324 and anoxic conditions, respectively). These HMWFA included linear n-alkanoic acids from n-325  $C_{20:0}$  to *n*- $C_{30:0}$  with an even-over-odd predominance characteristic of plant-derived inputs 326 (Eglinton and Hamilton 1967), linear  $\omega$ -hydroxyacids and  $\alpha, \omega$ -diacids from *n*-C<sub>16</sub> to *n*-C<sub>26</sub>, 327 10,16-dihydroxy $C_{16:0}$  and 9,10,18-trihydroxy  $C_{18:0}$  characteristic of plant-derived aliphatic 328 biopolymers cutin and suberin (Armas-Herrera et al. 2016; Kolattukudy 2001). The main 329 molecules were 9,10,18-trihydroxyC<sub>18:0</sub> (21.4  $\pm$  0.9% and 14.5  $\pm$  7.6% for oxic and anoxic 330 conditions, respectively) and  $\omega$ OH-C22 (7.9  $\pm$  0.7% and 10.7  $\pm$  0.7% for oxic and anoxic 331 conditions, respectively), which indicated a higher proportion of suberin from roots than cutin 332 from leaves (Mueller et al. 2012) (Table A6-7). The second most abundant chemical family 333 was PHE produced by the THM of lignins and tannins ( $11.5 \pm 0.0\%$  and  $11.7 \pm 3.2\%$  for oxic 334 and anoxic conditions, respectively). This chemical family was dominated by 3,4-335 dimethoxybenzoic acid methyl ester and 3,4,5-trimethoxybenzoic acid methyl ester, which are 336 typical of the THM of woody plants (Challinor 1995) and applied to investigate the transfer 337 and reactivity of OM (Denis et al. 2017; Williams et al. 2016). The third chemical family was 338 LMWFA with  $8.9 \pm 0.4\%$  and  $11.4 \pm 3.1\%$  for oxic and anoxic conditions, respectively. This 339 chemical family includes n-alkanoic acids from  $n-C_{8:0}$  to  $n-C_{19:1}$ , iso and anteiso  $C_{15:0}$  and  $C_{17:0}$ , iso  $C_{14:0}$  and  $C_{16:0}$  and n-alkenoic acids *n*- $C_{16:1}$ , *n*- $C_{18:1}$  and *n*- $C_{19:1}$ . The LMWFA with less than 340 341 13 C atoms can be derived from microbial or plant-derived inputs, while the LMWFA with

342 more than 13 C atoms are known as phospholipid fatty acids and are microbial biomarkers 343 (Frostegård et al. 1993) except for  $n-C_{16:0}$  and  $n-C_{18:0}$ , which can derive from plant-derived 344 inputs. In S:A, 0:100, the most abundant chemical family was LMWFA with  $67.8 \pm 3.1\%$  and 345  $61.3 \pm 3.8\%$  for oxic and anoxic conditions, respectively. The distribution of the compounds was different from that analyzed in S:A, 100:0, with a large predominance of  $n-C_{16:0}$  (19.3 ± 346 347 0.2% and 22.3  $\pm$  0.3% for oxic and anoxic conditions, respectively) and a large proportion of 348  $n-C_{16:2}$  (12.1 ± 2.3% and 9.5 ± 0.7% for oxic and anoxic conditions, respectively) and  $n-C_{18:2}$ 349  $(32.9 \pm 2.3\%$  and  $25.7 \pm 2.8\%$  for oxic and anoxic conditions, respectively). The second most 350 abundant chemical family was the chlorophyll biomarkers with two compounds: phytol and 351 neophytadiene. They represented  $18.3 \pm 3.2\%$  and  $21.5 \pm 7.8\%$  of oxic and anoxic conditions, 352 respectively. Finally, the third chemical family was composed of N-containing compounds, 353 which represented  $10.6 \pm 0.1\%$  and  $13.6 \pm 3.2\%$  (i.e., for oxic and anoxic conditions, 354 respectively). Among this family, 1,3-dimethyluracil was the dominant compound  $(4.3 \pm 0.6\%)$ 355 and  $7.3 \pm 1.1\%$  for oxic and anoxic conditions, respectively).

After 60 days of incubation, the characteristic chemical fingerprints of each pure endmember artificial sediments remained the same, and no significant differences were observed (Fig. 2).

359 3.2. Identification of the biodegradation-induced changes in mixed samples

The mixed artificial sediments (e.g., S:A, 75:25; S:A, 50:50 and S:A, 25:75) did not show notable differences in their molecular profile at day 0 (Table A.4- 5). A slight increase for the H/C<sub>wa</sub> was observed compared to the end members with values of ~1.1 instead of 0.8 or 0.9, suggesting a lower unsaturation for these samples that was confirmed by the values of DBE, which are much lower than for both pure end-member artificial sediments (e.g., values ranged from 10.8 to 12.6 irrespective of the oxygen conditions). Likewise, the values of AI<sub>mod</sub>> 0.5 and inferior to both pure end-member artificial sediments demonstrated the occurrence also 367 of less aromatic compounds. The elemental formula distributions of CHO and CHON are also 368 quite close to each other, whatever the proportions of algae; even so, a slightly higher 369 proportion of CHON formula compounds is observed in anoxic conditions for the S:A, 25/75 370 mixed sediment which reaches 49.2% for instance. The percentages of formulae with sulfur 371 (i.e., CHOS and CHONS) were higher than for either end members with values ranging from 372 ~10% to ~17% for CHOS and ~7% to 14.3% for CHONS. Regarding the distribution in 373 compound classes, similarities in values were also observed (Fig. 1). The two main groups of 374 compound classes were lignins/CRAM and CAS with percentages ranging from 38.6% to 375 44.2% and 22.2% to 30.2% respectively, followed by lipids and unsaturated hydrocarbons with 376 percentages ranged between 9% and 17.9%. Proteins and carbohydrates and aminosugars 377 occurred at less than 5%, respectively. This distribution was also reflected in the MLB<sub>L</sub> index 378 (percentage of labile material), which reaches values 1.5- to 7-fold those observed for the soil 379 or pure algae artificial sediment, respectively (Table A.4 and A.5).

380 After 60 days of incubation, each mixed artificial sediment displayed a unique 381 molecular profile compared to the initial according to its ratio of end-members and/or its 382 oxygen condition. S:A, 50:50, irrespective of the oxygen conditions, strictly showed the same 383 distribution in compounds class, suggesting a homogeneous effect of the biodegradation on the 384 whole pool of OM an. A similar observation was made for S:A, 75:25 in oxic conditions. By 385 contrast, molecular changes were observed in S:A, 25:75 for both oxygen conditions and 386 specifically for S:A, 75:25 in anoxic conditions (Fig. 1). In S:A, 75:25, a decrease of the CAS 387 and lignin/CRAM was noted in favor of the unsaturated hydrocarbons becoming, thus, the main 388 compound class (i.e., 33.7%). For S:A, 25:75, similar changes were also observed, however, 389 the lignin/CRAM compound class remained the main one (i.e., values of 39% and 33% for oxic and anoxic conditions, respectively). Similar changes to the changes observed for the end-390 391 member in higher proportion would have been expected, however, results displayed different trends as 3 different scenarios of biodegradation according to the mixing ratios were observed (e.g., no changes or homogeneous changes for all types of compounds, consumption of LMW compounds associated to production of LMW compounds and preferential degradation of one type of HMW compounds). These differences and similarities in the compound distribution for the mixed samples suggest the occurrence of underlying mechanisms.

397 At day 0, the biomarker distribution of biomarkers into chemical families illustrated a 398 gradual modification of the biomarker distribution induced by the mixing of end-members (Fig. 399 2). This gradual modification was characterized by the progressive increase in the proportion 400 of LMWFA from 24.0  $\pm$  5.2% (S:A,75:25; 26.3  $\pm$  6.0% and 21.6  $\pm$  4.7% in oxic and anoxic 401 conditions, respectively) to  $55.4 \pm 3.8\%$  (S:A, 25:75;  $57.8 \pm 4.6\%$  and  $53.0 \pm 0.2\%$  in oxic and 402 anoxic conditions, respectively). An increase in the proportion of chlorophyll biomarkers was 403 also observed with the increase of algae proportions with values from  $3.0 \pm 1.3\%$  (S:A, 75:25, 404  $3.8 \pm 1.4\%$  and  $2.2 \pm 0.9\%$  in oxic and anoxic conditions, respectively) to  $11.0 \pm 3.5\%$  (S:A, 405 25:75,  $11.0 \pm 1.4\%$  and  $11.0 \pm 5.8\%$  in oxic and anoxic conditions, respectively). At the same 406 time, the proportion of HMWFA decreased from  $45.0 \pm 6.4\%$  (S:A, 75:25;  $41.6 \pm 7.5\%$  and 407  $48.3 \pm 4.6\%$  in oxic and anoxic conditions, respectively) to  $19.6 \pm 3.9\%$  (S:A, 25:75;  $18.5 \pm$ 408 4.1% and  $20.6 \pm 4.8\%$  in oxic and anoxic conditions, respectively). The PHE proportion also 409 decreased regularly from  $9.5 \pm 1.1\%$  (S:A, 75:25;  $8.8 \pm 0.2\%$  and  $10.2 \pm 1.0\%$  in oxic and 410 anoxic conditions, respectively) to  $2.3 \pm 0.4\%$  (S:A, 75:25;  $2.0 \pm 0.2\%$  and  $2.6 \pm 0.3\%$  in oxic 411 and anoxic conditions, respectively).

412 At day 60, various changes were observed in the distribution of biomarkers. The most 413 significant changes were observed for S:A, 75,25, and S:A, 50:50, and specifically in anoxic 414 conditions, where a significant (p-value < 0.05) diminution was observed for almost all the 415 types of chemical families. By contrast, an increase in the proportion of LMWFA was observed 416 for all mixing ratios and oxygen conditions, but this increase was significant only for S:A, 417 75;25 and S:A, 50:50 in anoxic conditions. Regarding S:A, 25:75 only the diminution of SOAc
418 and Ncompound were significant.

419 *3.3. Factors controlling the biodegradation* 

420 *3.3.1. Source mixing factor* 

421 In order to understand the role and the weight of the factor "source", a PCoA was 422 performed based on the Bray-Curtis dissimilarity matrix computed all the FT-ICR MS 423 parameters (Fig. 3). Unfortunately, the ordination did not show any clear distinction either 424 according to the sources or mixing of sources and/or the incubation time. Although FT-ICR 425 MS analysis has emerged as a reliable tool for the in-depth molecular characterization of 426 complex mixtures as OM with the capacity of identifying thousands of molecules per sample 427 and assigning associated elemental formulae with a high level of confidence (Repeta 2015; 428 Stubbins and Dittmar 2014), this technique can also present some inherent limitations. The 429 main limitations of the FT-ICR MS tool relate to the inconsistent ionization efficiency for 430 different sample matrices, potentially leading to a bias of detection toward certain groups of 431 compounds or even to un-characterization of material (Blackburn et al. 2017; D'Andrilli et al. 432 2020; Dittmar et al. 2008; Hawkes et al. 2019; Mopper et al. 2007; Patriarca et al. 2020). In the 433 present study, it is noticed the number of assigned formulae decreased gradually with the 434 increase of the algae end-member, which could be assimilated to lower efficiency of the 435 ionization for this kind of matrix (Table A.4 and A.5) (Blackburn et al. 2017; Kew et al. 2018). 436 A lower ionization efficiency and preferential ionization for some group of compounds might 437 constrain the data interpretation and make it harder for disentangling the complexity of the OM 438 molecular composition (Derrien et al. 2019a; Hawkes et al. 2018). This could be one of the 439 reasons for the failure of the ordination to identify sub-groups according to the sources, oxygen conditions, and/or time of incubations. Another reason could be related to the non-440 441 discriminatory character of the data. FT-ICR MS is a technique able to identify with high

442 accuracy hundreds to several thousands of ions with an m/z range typically from 200 to 1000 443 Da. However, this level of sensitivity is not that discriminant for observing an ordination 444 between our set of samples.

445 By contrast, we clearly identified an effect of the factor "source" through the use of the 446 biomarkers. The PCA based on the distribution of the 67 variables displayed a clear separation 447 on axis-PC1 according to the proportion of algae (Fig. 4). The HAC performed on the results 448 of this latter PCA identified 4 groups. The 2 pure artificial sediments at day 0 and day 60 were 449 grouped with their respective endmember sample in two different groups (Groups 1 and 3). 450 The biomarker fingerprint modifications induced by the mixing of end-members were fairly 451 well illustrated by the projection of the samples at day 0 on PC1, representing 62.7% of the 452 variance of the dataset. They corresponded almost exactly to the theoretical points calculated 453 via the end-member mixing approach (Fig. 4). The HAC grouped S:A, 75:25 at day 0 samples 454 with the S:A, 100:0 (Group 3) and the rest of the mixed samples (S:A, 50:50 and 25:75) from 455 day 0 as a distinct group (Group 2). The last group identified by the HAC consisted of all three 456 S:A mixed samples from day 60 (Group 4). This group was characterized by proportions of *n*-457  $C_{14:0}$ , *n*- $C_{15:0}$ , *n*- $C_{16:0}$ , *a* $C_{15:0}$ , and *n*- $C_{18:0}$  FA higher than the mean of the samples and by 458 proportions of N-containing compounds, succinic and fumaric acids, n-C<sub>18:1</sub>, n-C<sub>18:2</sub>, n-C<sub>16:2</sub> 459 FA, and  $\alpha, \omega$ -diacid *n*-C<sub>9:0</sub> lower than the mean of the samples. 60 days' incubation resulted in 460 a specific decrease in the proportion of compounds that were specific to algae namely N-461 containing compounds and di-unsaturated FA. However, this decrease was not associated with 462 an increase in soil-specific markers such as HMWFA or PHE whose proportions tend to 463 decrease or to remain stable during the incubation (Fig. 2). This decrease was significant 464 (Wilcoxon test) only for PHE for S:A, 75:25, and 50:50. In the same time, the proportion of 465 microbial FA increased, which could be interpreted as stimulation of microbial biomass and 466 therefore the occurrence of positive priming effects (Guenet et al. 2014).

#### 467 *3.3.2. Oxygen factor*

468 Oxygen is one of the most important reactants in biogeochemical cycles as it impacts 469 the redox potential of the environment as well as the energetic situation of the microorganisms 470 (Brune et al. 2000). Therefore, its occurrence or absence may affect the biodegradation 471 processes and pathways. In the molecular profiles obtained by LDI FT-ICR MS, the main 472 differences were observed in S:A, 25:75, S:A, 75:25, and S:A, 0:100. The absence of oxygen 473 led to a higher production of unsaturated hydrocarbons of 1.5 to ~3 times higher at day 60 than 474 at day 0 with values from 12.0% to 33.7% and from 17.9% to 28.3% for the S:A, 25:75 and 475 S:A, 75:25 samples, respectively (Fig. 1 and Table A.5). In parallel, for the same samples, a 476 higher distribution was noticed in labile compounds after biodegradation than in oxic 477 conditions. These observations suggest a preferential biodegradation through the conversion of 478 HMW molecules into smaller compounds. By contrast, no differences were observed for the 479 distribution in biomarkers (Fig. 2). This was also confirmed with the PCA as the oxic and 480 anoxic samples are gathered together with no exceptions (Fig. 4). From our results, it seems 481 difficult to state on the effect of the oxygen condition on the observed biodegradation pathways. 482 3.4. Behind the chemical mechanisms of biodegradation

483 In order to further investigate the potential stimulation of the microbial biomass and 484 determine if there was a modification of the composition of the microbial community, the distribution of microbial FA (e.g., FA with a low molecular weight ( $< C_{19:0}$ ) except for  $C_{16:0}$ 485 486 and C<sub>18:0</sub>) was investigated using a PCA and a HAC, resulting in four groups depicted by grey 487 areas in Fig. 5. At day 0, the distribution of microbial FA in soil was modified by the addition 488 of algae. The evolution of this distribution was almost regular along the segment formed by the 489 end members. For algae, samples from days 0 and 60 remained in the same group (Group 1), 490 which highlighted that the distribution of FA remained almost stable during the incubation. On 491 the contrary for the soil, samples from day 60 were plotted in a distinct group (Group 3). This

492 indicates a modification of the distribution of microbial FA during the experiment, which could 493 be interpreted as a modification of the composition of the microbial community (Frostegård et 494 al. 1993). The S:A mixed samples at day 60 were also classified as a specific group (Group 4). 495 The direction of the modifications of the distribution of microbial FA was almost orthogonal 496 (i) to the segment formed by the end members and (ii) to the direction of the modifications in 497 the soil during the experiment. Consequently, these changes in the distribution of FA were not 498 only the result of the mixture but were also due to the adaptation of the composition of 499 microbial communities to the mixture of soil and algae.

500 A second point is that the Euclidean distance between day 0 and 60 for the different 501 S:A ratios seems to increase with the proportion of algae from  $2.9 \pm 0.9$  (S:A, 75:25) to  $3.9 \pm$ 502 0.3 (S:A, 50:50) and  $4.7 \pm 0.3$  (S:A, 25:75). The number of experimental replicates was not 503 sufficient to test the significance of this increase. However, it could suggest that the magnitude 504 of the modification of the composition of microbial communities increased with the proportion 505 of algae added to the soil. Moreover, S:A mixed samples at day 60 were grouped without clear 506 differentiations of mixture ratios, which could indicate that the modification of the microbial 507 communities was independent of the S:A ratio. Finally, this group of S:A mixed samples from 508 day 60 were characterized by proportions of  $n-C_{12:0}$  and  $n-C_{14:0}$  higher than the average and 509 proportion of n-C<sub>18:1</sub> (isomer 1), n-C<sub>17:0</sub> and n-C<sub>19:1</sub> lower than the average. Moreover, n-C<sub>18:1</sub> 510 (isomer 1) correlated with  $iC_{17:0}$ . The differences between the average proportion in group 4 511 and the average proportion in the entire dataset were significant for  $iC_{17:0}$  (*p*-value < 0.001). 512 This group may then be associated with a decrease in the proportion of terminally branched FA 513  $(iC_{17:0})$  that has been proposed as a marker of Gram-positive bacteria (Willers et al. 2015). 514 These results could signify that this microbial group was disadvantaged by the mixing of soil with algae. 515

516

#### 517 **4.** Conclusion

518 This study aimed at gaining insights into the biodegradation pathways and molecular 519 modifications of sedimentary organic matter, as well as to fathom the underlying mechanisms 520 via a laboratory experiment. The FT-ICR MS analysis allowed us to identify the group of 521 compounds preferentially degraded during early microbial diagenesis. In both pure end-522 members, CAS and lignins/CRAM were identified as the molecules the most affected by 523 biodegradation. For the mixed artificial sediments (i.e., S:A, 75:25, S:A, 50:50, and S:A, 524 25:75), distinct biodegradation alterations emerged for each ratio, underscoring the potent 525 influence of the "source" factor. The divergent shifts in compositional changes for different 526 mixing ratios post-incubation also hint at the potential involvement of underlying mechanisms, 527 such as priming effects. Notably, with the present study design, no discernible correlation was 528 established between the proportions of labile material and the intensity of priming effects. 529 Furthermore, the combination of the biomarker and molecular analyses allowed us to suggest 530 the capability of the microbial communities to adapt or shift in response to the varying types 531 and proportions of organic matter sources, facilitating the biodegradation of diverse materials. 532 However, additional studies and complementary qualitative analyses are imperative to test this 533 hypothesis conclusively. Nevertheless, this latter observation underscores the compelling need 534 to delve into and elucidate the intricate interplay between microbial communities and organic 535 matter. Addressing this challenge necessitates a holistic analytical approach, one that combines 536 organic geochemistry and meta-omic methods. This holistic approach, as underlined by 537 Bianchi and Ward (2019), would enable the identification of both the composition and function 538 of microbial communities. While this study indeed stems from a controlled laboratory 539 experiment, resulting in the formulation of questions and hypotheses for future testing, it 540 fundamentally demonstrates the intrinsic significance of chemical, biological, and 541 environmental factors in shaping the reactivity of OM.

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549	
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551	The data that support this study are available in the article and accompanying online
552	supplementary material.
553	
554	Conflict of Interest statement
555	The authors declare no conflicts of interest.
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560	
561	Figure' captions

Fig. 1 Distribution in compounds classes and elemental formulas of all the artificial sediments
for both oxygen conditions (i.e., oxic and anoxic). Characteristics of compound classes are
presented in Table A.2.

565 Fig. 2 Relative proportions of the chemical families identified by the THM–GC–MS analysis

of the end members at t = 0 day and t = 60 days in all the artificial sediments for both oxygen

567 conditions (i.e., oxic and anoxic). Uncertainties are the standard deviation of two experimental

replicates. Asterisks was added on data where a significant difference (p-value < 0.05) is observed between day 0 and day 60.

570 Fig. 3 Principal coordinate analysis ordination of the Bray-Curtis dissimilarity matrix

571 computed for the LDI FT–ICR MS data.

572 **Fig. 4** Principal component analysis performed with the distribution of biomarker (67). Purple

573 pentagons are the theoretical points of the S:A mixed samples calculated using an end-member

574 mixing approach. Grey shapes highlight the groups identified by the hierarchical analysis 575 classification.

576 **Fig. 5** Principal component analysis performed with the distribution of microbial fatty acids 577 biomarker (11). Grey shapes highlight the groups identified by the hierarchical analysis 578 classification.

579

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Fig. 1 Distribution in compounds classes of all the artificial sediments for both oxygen conditions
(i.e., oxic and anoxic). Characteristics of compound classes are presented in Table A.2.

- **Fig. 2** Relative proportions of the chemical families identified by the THM–GC–MS analysis of the
- 885 end members at t = 0 day and t = 60 days in all the artificial sediments for both oxygen conditions
- 886 (i.e., oxic and anoxic). Uncertainties are the standard deviation of two experimental replicates.
- 887 Asterisks was added on data where a significant difference (p-value < 0.05) is observed between day 0 888 and day 60.
- 889





Fig. 3 Principal coordinate analysis ordination of the Bray-Curtis dissimilarity matrix computed for
 the LDI FT–ICR MS data.



903 **Fig. 4** Principal component analysis performed with the distribution of biomarker (67). Purple

904 pentagons are the theoretical points of the S:A mixed samples calculated using an end-member 905 mixing approach. Grey shapes highlight the groups identified by the hierarchical analysis

- 906 classification.
- 907 43





912 biomarker (11). Grey shapes highlight the groups identified by the hierarchical analysis classification.

