## **SUPPLEMENTARY MATERIAL**

## How to measure the ESR intensity of the Al centre in optically bleached coarse quartz grains for dating purpose?

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**Figure S1.** ESR DRCs based on the Al signal measured in sample #1, whose intensity has been obtained from the four methods presented in this paper.





**Figure S2.** ESR DRCs based on the Al signal measured in sample #2, whose intensity has been obtained from the four methods presented in this paper.



**Figure S3.** ESR DRCs based on the AI signal measured in sample #3, whose intensity has been obtained from the four methods presented in this paper.





**Figure S4.** ESR DRCs based on the AI signal measured in sample #4, whose intensity has been obtained from the four methods presented in this paper.

**Figure S5.** ESR DRCs based on the AI signal measured in sample #5, whose intensity has been obtained from the four methods presented in this paper.





**Figure S7.** ESR DRCs based on the Al signal measured in sample #6, whose intensity has been obtained from the four methods presented in this paper.

**Figure S8.** ESR DRCs based on the AI signal measured in sample #7, whose intensity has been obtained from the four methods presented in this paper.



**Figure S8.** ESR DRCs based on the AI signal measured in sample #8, whose intensity has been obtained from the four methods presented in this paper.





## Figure S9. Comparison of AI signal measured in various aliquots of the eight quartz samples.

**Figure S10.** Comparison of the number of peaks per sample. Only the natural aliquot spectrum of each quartz sample is represented. All spectra were aligned with respect to the first peak. For sample #2, the signal intensity has been multiplied by a factor 6 and by 4 for the sample #4.



**Figure S13.** Comparison of the natural spectra of the sample #1 at low temperature (A), at room temperature with 2 mW and 10 mW (B). A basic visualization of the signal at room temperature is given in C.





**Figure S14.** Comparison of the natural spectra of the sample #2 at low temperature (A), at room temperature with 2mW and 10 mW (B). A basic visualization of the signal at room temperature is given in C.



**Figure S13.** Comparison of the natural spectra of the sample #3 at low temperature (A), at room temperature with 2mW and 10 mW (B). A basic visualization of the signal at room temperature is given in C.



**Figure S14.** Comparison of the natural spectra of the sample #4 at low temperature (A), at room temperature with 2 mW and 10 mW (B). A basic visualization of the signal at room temperature is given in C. For the identification of  $SO_3$  centre, see in detail lkeda et al. (1992) and Ryabov et al. (1983).



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Magnetic field (G) **Figure S15.** Comparison of the natural spectra of the sample #5 at low temperature (A), at room temperature with 2 mW and 10 mW (B). A basic visualization of the signal at room temperature is given in C.

**Figure S16.** Comparison of the natural spectra of the sample #6 at low temperature (A), at room temperature with 2mW and 10 mW (B). A basic visualization of the signal at room temperature is given in C. For the identification of SO<sub>3</sub> centre, see in detail lkeda et al. (1992) and Ryabov et al. (1983).











**Table S1**. Number of positive (+) and negative (-) peaks identified for various quartz samples, including the eight of the present study (\*). For sample #4, the spectrum is very noisy, inducing a possible bias on the peaks counting.

| Sample            | Context            | Peaks<br>+ | Peaks - | Total | Fraction (µm) | References                        |
|-------------------|--------------------|------------|---------|-------|---------------|-----------------------------------|
| Terra Amata       | Aeolian<br>(dune)  | 12         | 12      | 24    | 40-160        | Yokoyama et al. (1985)            |
| Pure quartz       | Artificial         | 12         | 12      | 24    | 30-50         | Lin et al. (2006)                 |
| ROX 1.14          | Aeolian<br>(loess) | 11         | 11      | 22    | 125-180       | Kabacińska and Timar-Gabor (2022) |
| STY 1.10          | Aeolian<br>(loess) | 10         | 10      | 20    | 125-180       | Kabacińska and Timar-Gabor (2022) |
| 2 MV 80           | Aeolian<br>(loess) | 8          | 8       | 16    | 63-90         | Kabacińska and Timar-Gabor (2022) |
| OUC1102-300 (#1)* | Fluvial            | 14         | 14      | 28    | 100-200       | This study                        |
| KT04-2-2 (#2)*    | Aeolian<br>(dune)  | 16         | 16      | 32    | 180-250       | Ben Arous et al. (2022)           |
| SF05-2-4 (#3)*    | Aeolian<br>(dune)  | 12         | 12      | 24    | 180-250       | Ben Arous et al. (2022)           |
| CAC1203 (#4)*     | Fluvial            | 12         | 12      | 24    | 100-200       | García-Vadillo et al (2021)       |
| CUB1004 (#5)*     | Fluvial            | 14         | 14      | 28    | 100-200       | Duval et al. (2017)               |
| ALC1201 (#6)*     | Fluvial            | 14         | 14      | 28    | 100-200       | Duval et al. (2015)               |
| GD1405 (#7)*      | Fluvial<br>(karst) | 13         | 13      | 26    | 100-200       | Duval et al. (2022)               |
| LC09-1-2 (#8)*    | Aeolian<br>(dune)  | 14         | 14      | 28    | 180-250       | Ben Arous et al. (2022)           |

## References

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