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Yunlong Guo, Kangwei Li, Sébastien Perrier, Taicheng An, D. Donaldson, et al.. Spontaneous Iodide Activation at the Air-Water Interface of Aqueous Droplets. Environmental Science and Technology, 2023, 57 (41), pp.15580-15587. 10.1021/acs.est.3c05777 . hal-04262346

# HAL Id: hal-04262346 https://hal.science/hal-04262346v1

Submitted on 21 Aug2024

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1	Spontaneous iodide activation at the air-water interface of aqueous droplets
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#### 24 Abstract

We present experimental evidence that atomic and molecular iodine, I and I<sub>2</sub>, are produced 25 spontaneously in the dark at the air-water interface of iodide-containing droplets without any added 26 catalysts, oxidants or irradiation. Specifically, we observe  $I_3^-$  formation within droplets, and  $I_2$ 27 emission into the gas phase from NaI-containing droplets over a range of droplet sizes in this 28 29 laboratory study. The formation of both products is enhanced in the presence of electron scavengers, either in the gas phase or in solution, and clearly follows a Langmuir-Hinshelwood mechanism, 30 suggesting an interfacial process. These observations are consistent with iodide oxidation at the 31 interface, possibly initiated by the strong intrinsic electric field present there, followed by well-32 known solution-phase reactions of the iodine atom. This interfacial chemistry could be important 33 in many contexts, including atmospheric aerosols. 34

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Keywords: spontaneous iodide activation, air-water interface, iodine, aqueous droplets,
 atmospheric chemistry

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Synopsis: Spontaneous interfacial production of I and I<sub>2</sub> is observed in iodide-containing aerosol droplets. This suggests a previously unknown source of reactive iodine species and could be important in atmospheric marine environment.

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## 43 Table of Contents (TOC)/Abstract Graphic



#### 46 Introduction

Aerosols and cloud droplets are key and ubiquitous components in the atmospheric system, 47 presenting a large air-water interface in the atmosphere.<sup>1,2</sup> Many species display a minimum free-48 energy at the air-water interface of such droplets, thus creating a favored and unique environment 49 for chemical reactions.<sup>3,4</sup> Very recently, it has been found that spontaneous H<sub>2</sub>O<sub>2</sub> and OH radical 50 generation and reduction of organic compounds take place at the air-water interface of 51 microdroplets,<sup>5-9</sup> while those reactions do not occur spontaneously in bulk solution. These 52 intriguing phenomena observed from microdroplets do not rely on additional catalysts, external 53 voltages, reducing agent, oxidants or irradiation, which have received great attention in recent years 54 and were intensively debated.<sup>10-12</sup> In a recent study by Li et al.,<sup>9</sup> an attempt was made to finally 55 recombine all previous observations by underlying the fact that the appearance of such oxidants is 56 not resulting from an oxidation process, but rather a charge separation process involving the ion-57 neutral pair of OH<sup>-</sup>. As initially suggested by Kloss,<sup>13</sup> such interfacial reactivity may be driven by 58 the intrinsic electric field (strength reported at the order of  $\sim 10^9$  V m<sup>-1</sup>) existing at the air-water 59 interface of microdroplets, which is suspected to be strong enough to induce charge separation in 60 ion-pairs or facilitate certain reactions by reducing the activation energy.<sup>14–16</sup> 61

Molecular iodine (I<sub>2</sub>) plays a key role in various important atmospheric cycles,<sup>17,18</sup> especially in 62 marine environments. Iodine photochemistry has been shown to drive new particle formation in the 63 marine boundary layer,<sup>19,20</sup> and to reduce global ozone levels by ca. 15%.<sup>21</sup> Recently, it has been 64 shown that I<sub>2</sub> injections into the stratosphere may even represent a small but non-negligible process 65 inducing O<sub>3</sub> depletion.<sup>22</sup> Gas phase I<sub>2</sub> in the marine environment originates from various sources 66 such as photodissociation of organo-iodine in tidal and open ocean areas, as well as the direct I2 67 release from biota near the coastal areas.<sup>23-25</sup> Currently, the reaction between ozone and iodide 68 anions is believed to dominate abiotic atmospheric iodine production.<sup>18,26,27</sup> 69

Here we report a novel mechanism for I<sub>2</sub> production at the air-water interface of submicron and 70 micron-sized NaI-containing droplets. Both gas phase I<sub>2</sub> and solution phase I<sub>3</sub><sup>-</sup> are experimentally 71 observed in the dark, in the absence of ozone. A very recent publication has also reported on 72 solution-phase I3<sup>-</sup> production (though not gas phase products) within droplets;<sup>28</sup> the present work 73 provides strong evidence that both  $I_{3(aq)}$  and  $I_{2(g)}$  are products of the same water surface-mediated 74 dark oxidation of iodide ions. The spontaneous formation of gas phase I<sub>2</sub>, although perhaps not 75 environmentally significant on Earth due to the low iodide amounts in actual seawater (I-76 concentration of 10–200 nM),<sup>24,29</sup> nevertheless represents another example of redox chemistry that 77 is mediated by the high potentials present at the air-water interface of aqueous droplets. 78

### 79 Materials and Methods

### 80 Quantification of I<sub>3</sub>-

UV-V is absorption spectra of  $I^{-}$ ,  $I_{2(aq)}$  and  $I_{3}^{-}$  standard solutions were measured over a relevant range 81 of concentrations. The I<sub>3</sub><sup>-</sup> standard solution was prepared by adding large excess of NaI to a solution 82 with known amount of  $I_{2(aq)}$ ; due to the large equilibrium constant for  $I_3^-$  formation, the initial  $I_{2(aq)}$ 83 was assumed to fully convert to  $I_3^-$  (see Text S1–S2). As shown in Figure 1A, the absorption peaks 84 for these three species can be identified and clearly distinguished; the I<sub>3</sub><sup>-</sup> showed clear absorption 85 peaks at ~288 and ~352 nm respectively, consistent with previous literature.<sup>30,31</sup> Using the 86 absorption intensity at 352 nm with a molar absorptivity of  $2.38 \pm 0.056 \times 10^4$  L cm<sup>-1</sup> mol<sup>-1</sup> (Figure 87 S1 and Text S2), the  $I_3^-$  concentration could be quantified, with the quantification limit at ~0.20 88 μМ. 89

#### 90 Mist chamber and atomizer experiments

The home-made mist chamber has been reported in our recent study<sup>9</sup> and is shown in Figure S2A. 91 It has a volume of ~110 mL. When the gas flow goes through the mist chamber, the reservoir 92 solution at the bottom will be lifted up due to pressure difference and therefore microdroplets are 93 continuously sprayed out. Such microdroplets impacted a hydrophobic PTFE membrane filter (0.2 94 µm pore size, Ref: FGLP04700, Merck Millipore Ltd) mounted at the top of the mist chamber, and 95 then dropped back into the bulk solution for respraying. As shown in Figure S2C, we measured the 96 microdroplet number size distribution (APS, TSI 3321) at the outlet of mist chamber during the 97 98 spraying process, and no particles were observed when a membrane filter was added. For a typical mist chamber experiment, 20 mL bulk NaI solution was added and the total gas flow was adjusted 99 by mass flow controllers at 2.5 L min<sup>-1</sup>. C<sub>2</sub>H<sub>4</sub>ClOH<sub>(g)</sub>, as an electron scavenger with a pKa value 100 of 14.31,<sup>32</sup> was generated by a gentle N<sub>2</sub> flow passing through the headspace of a 120 mL bottle 101 filled with 30 mL pure C<sub>2</sub>H<sub>4</sub>ClOH solution, and the C<sub>2</sub>H<sub>4</sub>ClOH<sub>(g)</sub> concentration was calculated 102 103 based on its vapor pressure. It is expected that the solution in the mist chamber evaporated continuously during the spraying process, and the remaining volume of solution decreased 104 gradually overtime. We calculated the liquid loss rate for each individual experiment, with the range 105 of 1.7–2.6 mL h<sup>-1</sup> (Tables S1–S2). For these electron-scavenging experiments using C<sub>2</sub>H<sub>4</sub>ClOH, 106 the solution pH decreased to 1-2 due HCl formation through hydrolysis of the dissolved 107 C<sub>2</sub>H<sub>4</sub>ClOH<sub>(g)</sub>. Previous studies have shown important role of solution acidity for I<sub>2</sub> production, and 108 iodide activation is favored at lower solution pH.<sup>33,34</sup> Additionally, we performed macroscopic 109

aqueous interfacial experiments as shown in Figure S6, and see whether spontaneous iodideactivation can occur on flat macroscopic surfaces in the dark.

We also produced iodide-containing droplets by atomizing bulk NaI solution using a commercial 112 constant output atomizer (TSI 3076) (Figure S2B), which is a different spraying procedure 113 compared with the mist chamber. Since both spraying procedures used the same bulk NaI solution, 114 we assumed that the composition of the droplets generated from atomizer are the same with 115 microdroplets generated from mist chamber. The droplets produced in the atomizer passed through 116 two tandem 5 L glass bottles, where some large droplets were condensed and collected for UV-Vis 117 measurement. Note that we continuously atomized the bulk solution for about 20 h, until enough 118 solution (i.e., 2–3 mL) accumulated in these bottles to be transferred into a quartz cuvette for UV-119 Vis analysis. Although such atomizers are mostly used for generating small droplets in the 120 121 submicron range (Figure S2E), micron-sized droplets are still produced, but with much lower number concentrations. As shown in Figure S2D, we measured the size distributions of 122 microdroplets produced by such an atomizer, which typically peaked at  $\sim 2 \mu m$  diameter with a 123 broad size range. 124

#### 125 Gas-phase I<sub>2(g)</sub> measurement with CI-Orbitrap

Bromide chemical ionization mass spectrometry (Br-CIMS) has been frequently deployed to 126 directly measure gas-phase iodine species, where dibromomethane (CH<sub>2</sub>Br<sub>2</sub>) is used to produce 127 bromide as reagent ions due to its good affinity to various iodine species.<sup>25,35</sup> An ultrahigh-128 resolution Orbitrap mass spectrometer coupled by a chemical ionization interface (CI-Orbitrap) has 129 emerged as a novel tool, which combines advantages of ultra-high mass resolving power (m/ $\Delta$ m 130  $\approx$ 140 000) and minimal fragmentation from soft atmospheric pressure ionization.<sup>36,37</sup> The 131 experiment setup was shown in Figure S3, where NaI-containing droplets were produced by a 132 commercial atomizer or our home-made mist chamber and then removed by a hydrophobic PTFE 133 membrane filter to allow a particle-free measurement of the remaining gas phase. Since CH<sub>2</sub>Br<sub>2</sub> 134 was used to produce reagent ions,  $I_{2(g)}$  was detected as bromide-adduct such as <sup>79</sup>Br( $I_2$ )<sup>-</sup> and <sup>81</sup>Br( $I_2$ )<sup>-</sup>, 135 with their ratio consistent with the theoretical isotope prediction (Figure S4). We also semi-136 calibrated the instrument by introducing a known quantity of I<sub>2</sub> vapor into the CI-Orbitrap. The 137 instrument showed good linear response for  $I_{2(g)}$  at the ppt level (Text S3 and Figure S5). The data 138 were processed using the Xcalibur 2.2 software (Thermo Scientific), and the selected ions (i.e., m/z139 = 332.7278 for <sup>79</sup>Br(I<sub>2</sub>), m/z = 334.7258 for <sup>81</sup>Br(I<sub>2</sub>)) were exported with a mass tolerance of 10 140 ppm. For consistency, all  $I_{2(g)}$  intensity reported in this study was shown as <sup>79</sup>Br( $I_2$ )<sup>-</sup> and detected at 141 m/z = 332.7278. More detailed information about the CI-Orbitrap can be found elsewhere.<sup>36,37</sup> 142

#### 144 **Results and Discussion**

#### 145 Spontaneous I<sub>3</sub><sup>-</sup> production from I<sup>-</sup>-containing droplets

We carried out experiments in microdroplets  $(1-10 \ \mu m)$ ; a few experiments investigated iodide 146 147 oxidation at macroscopic aqueous interfaces as well. The microdroplet experiments used two droplet sources. In the first one, a home-made mist chamber, NaI-containing microdroplets were 148 produced by spraying a 100 mM NaI bulk solution using various bath gases. The mist chamber was 149 operated in a recirculation mode (Figure S2A), where the microdroplets continuously impacted on 150 a membrane filter and dropped back into the bulk solution before being resprayed. Under such 151 recycling conditions, any possible reaction products in the reservoir bulk water of the mist chamber 152 accumulate over time. The reservoir bulk solution was analyzed for the appearance of  $I_3^-$  by UV-153 Vis spectroscopy, typically every hour. Figure 1A displays the absorption spectrum of I<sub>3</sub><sup>-</sup> and its 154 precursors; Figure 1B clearly indicates an increase of absorbance of  $I_3$  at ~352 nm after spraying a 155 100 mM NaI solution for 7 h using synthetic air  $(20\% O_2 + 80\% N_2)$  as the bath gas. 156

Interestingly, when such experiments are performed under nitrogen instead of synthetic air, Figure 1B shows that the increase of absorbance at  $\sim$ 352 nm is reduced, highlighting a possible role of oxygen in the observed chemistry. Normally, iodide is not oxidized by molecular oxygen in solution or gas phase. To test whether our observation relates to electron capture by oxygen at the air-water interface, 2-chloroethanol (C<sub>2</sub>H<sub>4</sub>ClOH) - a strong electron scavenger - was added to the pure nitrogen gas flow and gave rise to a notable increase of the I<sub>3</sub><sup>-</sup> production, as illustrated in Figure 1B.

In order to test whether the observed iodide activation is dependent on the spraying procedure, 164 we employed a commercial atomizer (TSI 3076), rather than the mist chamber, to nebulize the same 165 NaI solution into microdroplets, which were subsequently trapped in two tandem glass bottles for 166 167 UV-Vis measurement. Figure S2B illustrates the experimental arrangement used for these experiments. As shown in Figure 1C, we also observed the peak at ~352 nm in this case, indicating 168 formation of  $I_3$  in the microdroplets collected using four different spraving gases. Again, adding 169 oxygen or 2-chloroethanol to the nitrogen gas yielded much greater amounts of I<sub>3</sub>- product than 170 using pure nitrogen. Compared with the first vessel, the microdroplets trapped in the second vessel 171 generally showed a higher absorbance at  $\sim$ 352 nm due to a longer residence time. 172

Finally, we demonstrated that iodide anions are oxidized even on flat macroscopic surfaces in the dark. For this purpose, one glass bottle (30 mL size) was filled with 3 mL volume of NaI- containing solution with a macroscopic air-water interface (surface area of ca.  $8.3 \text{ cm}^2$ ), and another identical glass bottle was fully filled with the same solution with no head-space (i.e., with no airwater interface) (Figure S6). Under such conditions, I<sub>3</sub><sup>-</sup> production was observed after several days from the solution that had a macroscopic air-water interface, while no I<sub>3</sub><sup>-</sup> formation was observed in the solution with no interface, even days after preparing the stock solution (Figure S6). These experiments confirm that iodide oxidation does not occur in bulk solution, but requires an interface, and that it is not driven by impurities present in our water supply or in the chemicals used.

The enhanced iodide activation due to the addition of electron scavengers suggests the involvement of free electrons in the mechanism. This in turn implies that I<sup>-</sup> may be oxidized to produce a free electron according to (E1), in a manner similar to that previously suggested for the dissociation of hydroxide anions (OH<sup>-</sup>) into OH radicals and electrons under the naturally formed electric field (~10<sup>9</sup> V m<sup>-1</sup>) at the air-water interface of microdroplets.<sup>6,9,13–15,38</sup>

Similar to OH<sup>-</sup> as initially proposed by Kloss,<sup>13</sup> the forward step in E1 is suggested to be a charge 188 separation process induced by an interfacial electric field, and hence it should not be regarded as 189 an oxidation process that costly in energy. For instance, iodide is serving a typical example for 190 investigating charge-transfer-to-solvent (CTTS) transitions and the corresponding solvation 191 transitions have been investigated.<sup>39</sup> However, the transition is typically triggered by appropriate 192 UV light, in contrast to what is suggested here. The electron affinity of the iodine atom is 193 determined as 3.059 eV.<sup>40</sup> It is known that solvation shells do differ between the bulk and the 194 interfacial region, and that this may facilitate or promote some redox reactions.<sup>41-43</sup> Here 195 specifically, modified solvation processes at the interface will affect the equilibrium E1. If the 196 electron is scavenged, the equilibrium in (E1) will be shifted to the right and enhance iodide 197 activation. Once produced, the atomic iodine may react through a sequence of different reactions 198 (E2-E5) producing I<sub>2</sub>, I<sub>2</sub><sup>-</sup> and finally I<sub>3</sub><sup>-</sup>, of which many I<sub>2</sub><sup>-</sup> may decay back to I<sup>-</sup> (E6),<sup>44</sup> according 199 200 to:

$$I \cdot + I \cdot \rightleftharpoons I_{2(aq)}$$
(E2)

$$I \cdot + I^- \rightleftharpoons I_2^-$$

$$I_2^- + I^- \rightleftharpoons I_3^- \tag{E4}$$

$$I_{2(aq)} + I \quad \rightleftharpoons \quad I_3 \tag{E5}$$

(E3)

$$l_2^{-} + l_2^{-} \rightleftharpoons l_3^{-} + l^{-}$$
(E6)

Under pure nitrogen, the solvated electron generated in E1 can recombine on an approximately 206 hundred picosecond timescale with the nearby I radical due to a solvent cage effect,<sup>45</sup> resulting in 207 a null cycle. By contrast, enhanced I<sub>3</sub><sup>-</sup> formation was observed when an electron scavenger such as 208 C<sub>2</sub>H<sub>4</sub>ClOH<sub>(g)</sub>, O<sub>2(g)</sub> or N<sub>2</sub>O<sub>(g)</sub> was added to the pure N<sub>2</sub> gas spraying flow,<sup>46,47</sup> either in the mist 209 chamber (Figure 1B and Figure S7A) or using a commercial atomizer (Figure 1C and Figure S8). 210 Note that C<sub>2</sub>H<sub>4</sub>ClOH<sub>(g)</sub> has high solubility in water, which can lower the solution pH by forming 211 HCl through hydrolysis of the dissolved C<sub>2</sub>H<sub>4</sub>ClOH<sub>(g)</sub>, as illustrated in Figure S7D-E. The lowered 212 pH may also affect the electron lifetime in solution. In addition, by dissolving over time in the 213 aqueous solution, C<sub>2</sub>H<sub>4</sub>ClOH may also change the physical characteristics of the nebulized aerosol 214 droplets due to changes in colligative properties and specific molecular interactions. However, our 215 observation of enhanced I<sub>3</sub><sup>-</sup> production in the presence of O<sub>2(g)</sub> and N<sub>2</sub>O<sub>(g)</sub>, neither of which is 216 217 strongly soluble, gives strong support to electron capture at the air-aqueous interface as being a key driver of I<sub>3</sub><sup>-</sup> production (see Text S4 and Figure S8). 218

The presence of an electron scavenger will shift the equilibrium in E1, allowing I atoms to live 219 sufficiently long to produce I<sub>2</sub>, I<sub>2</sub><sup>-</sup> as well as I<sub>3</sub><sup>-</sup>, via E2–E6 above. This reaction scheme gives a 220 plausible explanation for the observation of I<sub>3</sub><sup>-</sup> formation. As many of such intermediates decay 221 back to iodide, our observations represent only a lower limit to the actual activation of the iodide 222 anion. It is worth mentioning that a recent work by Xing et al.<sup>28</sup> also reported similar I<sub>3</sub><sup>-</sup> formation 223 224 from spraying NaI bulk solution into microdroplets, and they proposed that the oxidation of iodide was explained by spontaneous production of OH radicals at the interface and subsequent OH-225 oxidation chemistry. Given the ubiquitous interfacial production of OH radicals,<sup>9</sup> this pathway is 226 certainly taking place as well under our conditions, as the experimental approach is overall similar. 227 However, we do suggest that this OH-initiated iodide oxidation should be regarded as minor. In 228 229 fact, the addition of an organic compound, in large excess and acting as an OH radical scavenger, should prevent the formation of  $I_2$  or  $I_3$ . By contrast, the addition of 2-chloroethanol did not prevent 230 the formation of these products but rather enhanced their appearance, showing indeed that the key 231 element in this chemistry is the electron trapping process that displaces equilibrium E1 to the right. 232 Additionally, our observation of gas phase I<sub>2</sub> formation and Langmuir-Hinshelwood behaviour 233 (vide infra) suggest that spontaneous iodide activation is more likely driven by a surface-mediated 234 process involving the loss of free electrons (E1). 235

 $I_3^-$  production rates (in nmol h<sup>-1</sup>; see Text S5) measured under different spraying gases were 236 summarized in Figure 2A. The clear evidence for activation of iodide seen in the droplets was not 237 observed in bulk solution, and was enhanced in the presence of gas phase electron scavengers. 238 These observations are consistent with a surface-mediated mechanism taking place at the air-water 239 interface. This conjecture is supported by the dependence of I<sub>3</sub><sup>-</sup> production on the bulk iodide 240 concentration. Figure 2B shows that the I<sub>3</sub><sup>-</sup> production rate does not depend linearly on bulk iodide 241 concentration (as expected for a first order bulk process), but exhibits limiting behavior, with a 242 linear dependence at lower I<sup>-</sup> concentrations (0.1–20 mM), and reaching a plateau at higher I<sup>-</sup> 243 concentrations (20-100 mM). The detailed results for each individual mist chamber experiment are 244 summarized in Figure S9 and Table S1. Such a relationship between the I<sub>3</sub><sup>-</sup> production rate and the 245 NaI bulk concentration is well represented by a Langmuir-Hinshelwood model, which is specific 246 for surface driven processes.<sup>48–51</sup> This idea is explored further below. Altogether, these indications 247 are strongly supportive of an interfacial oxidation of iodide anion giving rise to  $I_3^-$  production. 248

#### 249 Release of interfacial I<sub>2</sub> into the gas phase

Molecular iodine is only moderately soluble in water, with a Henry's law constant of ca. 2.8 M atm<sup>-1</sup> at room temperature.<sup>52</sup> As a consequence, if I<sub>2</sub> production occurs at the air-water interface, one would expect a significant fraction of that compound to desorb into the gas phase. To test whether gaseous iodine is produced at the droplet surface, we performed direct measurement of  $I_{2(g)}$ using an online chemical ionization orbitrap (CI-Orbitrap) instrument, with bromide as reagent ion.

As shown in Figure 3A, we observed clear steady-state  $I_{2(g)}$  signal (detected as bromide-adduct 255 such as  $^{79}Br(I_2)$ ) when atomizing a 100 mM NaI solution under N<sub>2</sub> or pure air, while a 100 mM 256 257 (NH<sub>4</sub>)<sub>2</sub>SO<sub>4</sub> solution nebulized by the same procedure led to signal at the level of instrumental background. Compared with the results using N<sub>2</sub> as carrier gas, enhanced I<sub>2(g)</sub> signal was observed 258 using pure air (Figure 3A–B and Figure S10A–B), consistent with the observations of enhanced I<sub>3</sub><sup>-</sup> 259 production in the presence of O<sub>2(g)</sub> discussed above. This comparison reinforces the fact that electron 260 261 scavengers such as oxygen can play an important promotion role for the interfacial iodide activation. Figure 3C and Figure S10C show the I<sub>2(g)</sub> signal as a function of pH of a 10 mM NaI bulk solution. 262 The production of I<sub>2</sub> seems to be more efficient under acidic conditions, again similar to what we 263 see for I3<sup>-</sup> production in Figure S7C. Such pH dependence agrees with previous findings that 264 abundant protons (H<sup>+</sup>) at lower pH could enhance the production of  $I_{3}(aq)$  and  $I_{2(g)}$ .<sup>33,53</sup> As shown in 265 Figure 3D, the I<sub>2(g)</sub> signal also shows a non-linear dependence on the bulk NaI concentration, similar 266 to what is shown above in Figure 2B for  $I_3^-$  production. This is compelling evidence to support the 267 proposal that both spontaneous  $I_{2(g)}$  production and  $I_3^-$  production via iodide activation take place 268

- as an interfacial process. Such Langmuir–Hinshelwood behaviour also rules out the potential iodide activation route from interfacial production of  $H_2O_2$ . For instance, if I<sub>2</sub> would arise from the titration of iodide by  $H_2O_2$ , then the amount of I<sub>2</sub> produced should not depend on the I<sup>-</sup> concentration due to the great excess of I<sup>-</sup> to  $H_2O_2$ , which does not agree with the above observation in this study.
- The CI-Orbitrap instrument demonstrated a linear response to different  $I_{2(g)}$  concentrations, and 273 here was only calibrated for semi-quantitative purposes (see Figure S5 and Text S3 for details). The 274 corresponding semi-quantitative I<sub>2(g)</sub> concentrations for all experimental conditions tested here are 275 summarized in Table S3. Total surface area concentrations could be calculated based on the droplet 276 size distribution measured at the output of the atomizer or mist chamber (Figure S2C-E), 277 respectively. As shown in Table S4, the surface area normalized  $I_{2(g)}$  concentrations were calculated 278 to be in the range of  $2-8 \times 10^{15}$  molec m<sup>-2</sup> for droplets produced from atomizer and mist chamber 279 respectively, which agrees in between within a factor of 4. 280

#### 281 Langmuir-Hinshelwood behaviour

The dependence of product formation on bulk iodide concentration is consistent between both aqueous phase  $I_3^-$  production (Figure 2B) and gas phase  $I_{2(g)}$  production (Figure 3D). This implies that the production of both species has the same dependence on the presence of iodide anions at the air-water interface. In order to describe this process by a Langmuir–Hinshelwood kinetic model, we normalize all the production rates to their maxima, and plot these as a function of bulk NaI concentration in Figure 4, fitting the data to a Langmuir–Hinshelwood equation (1):

normalized I<sub>3</sub> and I<sub>2(g)</sub> production 
$$= \frac{A[I_{\bar{a}q}]}{B + [I_{\bar{a}q}]} = \frac{A[I_{\bar{a}q}]}{\frac{1}{K_{ads} + [I_{\bar{a}q}]}}$$
 (1)

In this equation,  $[I_{aq}]$  is the bulk concentration of iodide; parameter *A* represents the maximum observed rate; parameter *B* represents the ratio of the desorption to adsorption rate constants for I<sup>-</sup> adsorption from the bulk solution to the interface, which is equivalent to the inverse of equilibrium constant of I<sup>-</sup><sub>aq</sub> for surface adsorption (*K*<sub>ads</sub>).<sup>49,50</sup>

By applying a Langmuir–Hinshelwood fit to the combined data in Figure 4, we obtain an adsorption constant  $K_{ads}$  of 295±100 M<sup>-1</sup> for I<sup>-</sup> partitioning from bulk to the interface. Note that this does not represent a true thermodynamic equilibrium constant, as we have ignored activity coefficient and the surface standard state is not defined. However, comparing this value to other similar analyses shows reasonable agreement within a factor of 4–13, with the reported  $K_{ads}$  values (i.e., 23±3 M<sup>-1</sup> and 70±16 M<sup>-1</sup>), from analysis of the heterogeneous reaction between O<sub>3(g)</sub> and NaI solution.<sup>50,51</sup>

#### 300 Atmospheric implications

It is worth mentioning that some previous studies also reported  $I_{3}(aq)$  and  $I_{2}(g)$  production as a result 301 of the addition of an external trigger i.e., ozone, irradiation, catalyst, etc.<sup>33,34,53,54</sup> We observed 302 spontaneous  $I_3^-$  and  $I_{2(g)}$  production from NaI droplets through an iodide activation process that does 303 not require catalysts, external voltages, oxidants or irradiation. This represents a previously 304 unknown source of reactive iodine species. Our results suggest that such spontaneous dark 305 production of I<sub>3</sub><sup>-</sup> and I<sub>2(g)</sub> is an interfacial process, which can be well described using a Langmuir– 306 Hinshelwood formalism. The production of  $I_3^-$  and  $I_{2(g)}$  is strongly affected by the presence of 307 electron scavengers such as N<sub>2</sub>O, O<sub>2</sub> or C<sub>2</sub>H<sub>4</sub>ClOH, consistent with oxidation of iodide at the 308 interface. 309

We propose that this spontaneous iodide activation could be induced by the strong electric field 310  $(\sim 10^9 \text{ V m}^{-1})$  which is thought to be formed naturally at the air-water interface of microdroplets.<sup>14,15</sup> 311 However, molecular dynamics simulations show large fluctuations across the interface due to 312 partial solvation effect and the associated variation of the solvent electrostatic potential, which 313 suggests the incomplete and large uncertainty in the current understanding of interfacial electric 314 field.<sup>55,56</sup> Similar to the recent observation of spontaneous H<sub>2</sub>O<sub>2</sub> and OH radical production at the 315 interface of microdroplets,<sup>9,38</sup> we hypothesize that such a high electric field can dissociate iodide 316 anions into iodine atoms and electrons (e<sup>-</sup>) (E1), following by the formation of reactive iodine 317 species (I and I<sub>2</sub>), which can further accommodate into the aqueous phase or be released into the 318 gas phase and then potentially trigger more iodine-related reactions with other compounds. For 319 instance, the aqueous iodine species may react with dissolved organic matter in the sea-surface 320 microlayer to produce organo-iodine species,  $^{21,23}$  while the interfacial released  $I_{2(g)}$  could deplete 321 tropospheric O<sub>3</sub>, initiate new particle formation and oxidize hydrocarbons.<sup>17,57,58</sup> The Supporting 322 Information (Text S6 and Figures S11-S12) gives details of some preliminary experiments that 323 324 show new particle formation by reaction of the I<sub>2</sub> released from droplet surfaces with ozone under visible irradiation. 325

#### **Supporting Information**

The Supporting Information includes additional experimental details, materials and methods, and further results from mist chamber and atomizer experiments,  $I_{2(g)}$  calibration with CI-Orbitrap and aerosol flow tube (AFT) experiments.

330 Acknowledgment

- 331 This study was supported by the European Research Council under the Horizon 2020 research and
- 332 innovation programme / ERC Grant Agreement 101052601 SOFA. TA and YG acknowledge
- the financial support from the National Natural Science Foundation of China (42020104001 and
- 334 42377221).

## 335 Author Information

- 336 Yunlong Guo and Kangwei Li contributed equally to this work.
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## 341 **References**

- (1) George, C.; Ammann, M.; D'Anna, B.; Donaldson, D. J.; Nizkorodov, S. A. Heterogeneous
  photochemistry in the atmosphere. *Chem. Rev.* 2015, *115*, 4218-4258.
- 344 (2) Pillar-Little, E.; Guzman, M. An overview of dynamic heterogeneous oxidations in the 345 troposphere. *Environments* **2018**, *5*, 104.
- 346 (3) Ruiz-Lopez, M. F.; Francisco, J. S.; Martins-Costa, M. T. C.; Anglada, J. M. Molecular 347 reactions at aqueous interfaces. *Nat. Rev. Chem.* **2020**, *4*, 459-475.
- 348 (4) Rossignol, S.; Tinel, L.; Bianco, A.; Passananti, M.; Brigante, M.; Donaldson, D. J.; George,
- C. Atmospheric photochemistry at a fatty acid-coated air-water interface. *Science* 2016, *353*, 699 702.
- (5) Lee, J. K.; Samanta, D.; Nam, H. G.; Zare, R. N. Micrometer-sized water droplets induce spontaneous reduction. *J. Am. Chem. Soc.* **2019**, *141*, 10585-10589.
- 353 (6) Lee, J. K.; Walker, K. L.; Han, H. S.; Kang, J.; Prinz, F. B.; Waymouth, R. M.; Nam, H. G.;
- Zare, R. N. Spontaneous generation of hydrogen peroxide from aqueous microdroplets. *P. Natl.*
- 355 Acad. Sci. USA **2019**, 116, 19294-19298.
- (7) Lee, J. K.; Han, H. S.; Chaikasetsin, S.; Marron, D. P.; Waymouth, R. M.; Prinz, F. B.; Zare,
  R. N. Condensing water vapor to droplets generates hydrogen peroxide. *P. Natl. Acad. Sci. USA* **2020**, *117*, 30934-30941.
- (8) Gong, C.; Li, D.; Li, X.; Zhang, D.; Xing, D.; Zhao, L.; Yuan, X.; Zhang, X. Spontaneous
  reduction-induced degradation of viologen compounds in water microdroplets and its inhibition by
  host-guest complexation. *J. Am. Chem. Soc.* 2022, *144*, 3510-3516.
- (9) Li, K.; Guo, Y.; Nizkorodov, S. A.; Rudich, Y.; Angelaki, M.; Wang, X.; An, T.; Perrier, S.;
  George, C. Spontaneous dark formation of OH radicals at the interface of aqueous atmospheric
  droplets. *P. Natl. Acad. Sci. USA* 2023, *120*, e2220228120.
- (10) Musskopf, N. H.; Gallo Jr, A.; Zhang, P.; Petry, J.; Mishra, H. The air-water interface of
   water microdroplets formed by ultrasonication or condensation does not produce H<sub>2</sub>O<sub>2</sub>. *J. Phys. Chem. Lett.* 2021, *12*, 11422-11429.
- (11) Nguyen, D.; Nguyen, S. C. Revisiting the effect of the air-water interface of ultrasonically
   atomized water microdroplets on H<sub>2</sub>O<sub>2</sub> formation. J. Phys. Chem. B 2022, 126, 3180-3185.
- (12) Nguyen, D.; Lyu, P.; Nguyen, S. C. Experimental and thermodynamic viewpoints on claims
  of a spontaneous H<sub>2</sub>O<sub>2</sub> formation at the air-water interface. *J. Phys. Chem. B* 2023, *127*, 23232330.
- (13) Kloss, A. I. Electron-radical dissociation and the water activation mechanism. *Doklady Akademii Nauk Sssr* 1988, *303*, 1403-1407.
- (14) Xiong, H.; Lee, J. K.; Zare, R. N.; Min, W. Strong electric field observed at the interface of
  aqueous microdroplets. *J. Phys. Chem. Lett.* 2020, *11*, 7423-7428.
- (15) Hao, H.; Leven, I.; Head-Gordon, T. Can electric fields drive chemistry for an aqueous
   microdroplet? *Nat. Commun.* 2022, *13*, 280.
- (16) Heindel, J. P.; Hao, H.; LaCour, R. A.; Head-Gordon, T. Spontaneous formation of hydrogen
  peroxide in water microdroplets. *J. Phys. Chem. Lett.* 2022, *13*, 10035-10041.
- (17) Carpenter, L. J. Iodine in the marine boundary layer. *Chem. Rev.* **2003**, *103*, 4953-4962.

- 382 (18) Carpenter, L. J.; MacDonald, S. M.; Shaw, M. D.; Kumar, R.; Saunders, R. W.; Parthipan,
- R.; Wilson, J.; Plane, J. M. C. Atmospheric iodine levels influenced by sea surface emissions of inorganic iodine. *Nat. Geo.* **2013**, *6*, 108-111.
- (19) O'Dowd, C. D.; Hoffmann, T. Coastal new particle formation: A review of the current state of-the-art. *Environ. Chem.* 2005, *2*, 245-255.
- 387 (20) Huang, R.-J.; Hoffmann, T.; Ovadnevaite, J.; Laaksonen, A.; Kokkola, H.; Xu, W.; Xu, W.;
- Ceburnis, D.; Zhang, R.; Seinfeld, J. H.; O'Dowd, C. Heterogeneous iodine-organic chemistry fasttracks marine new particle formation. *P. Natl. Acad. Sci. USA* **2022**, *119*, e2201729119.
- 390 (21) Carpenter, L. J.; Chance, R. J.; Sherwen, T.; Adams, T. J.; Ball, S. M.; Evans, M. J.; Hepach,
- H.; Hollis, L. D. J.; Hughes, C.; Jickells, T. D.; Mahajan, A.; Stevens, D. P.; Tinel, L.; Wadley, M.
- R. Marine iodine emissions in a changing world. *Proc. R. Soc. A* **2021**, *477*, 20200824.
- 393 (22) Koenig, T. K.; Baidar, S.; Campuzano-Jost, P.; Cuevas, C. A.; Dix, B.; Fernandez, R. P.;
- Guo, H.; Hall, S. R.; Kinnison, D.; Nault, B. A.; Ullmann, K.; Jimenez, J. L.; Saiz-Lopez, A.;
  Volkamer, R. Quantitative detection of iodine in the stratosphere. *P. Natl. Acad. Sci. USA* 2020,
- *396 117*, 1860-1866.
- 397 (23) Simpson, W. R.; Brown, S. S.; Saiz-Lopez, A.; Thornton, J. A.; Glasow, R. Tropospheric
  398 halogen chemistry: Sources, cycling, and impacts. *Chem. Rev.* 2015, *115*, 4035-4062.
- (24) Carpenter, L. J.; Nightingale, P. D. Chemistry and release of gases from the surface ocean. *Chem. Rev.* 2015, *115*, 4015-4034.
- 401 (25) Tham, Y. J.; He, X.-C.; Li, Q.; Cuevas, C. A.; Shen, J.; Kalliokoski, J.; Yan, C.; Iyer, S.;
- 402 Lehmusjarvi, T.; Jang, S.; Thakur, R. C.; Beck, L.; Kemppainen, D.; Olin, M.; Sarnela, N.; Mikkila,
- 403 J.; Hakala, J.; Marbouti, M.; Yao, L.; Li, H.; Huang, W.; Wang, Y.; Wimmer, D.; Zha, Q.; Virkanen,
- J.; Spain, T. G.; O'Doherty, S.; Jokinen, T.; Bianchi, F.; Petaja, T.; Worsnop, D. R.; Mauldin III, R.
- L.; Ovadnevaite, J.; Ceburnis, D.; Maier, N. M.; Kulmala, M.; O'Dowd, C.; Maso, M. D.; Saiz-Lopez, A.; Sipila, M. Direct field evidence of autocatalytic iodine release from atmospheric aerosol.
- 407 *P. Natl. Acad. Sci. USA* **2021**, *118*, e2009951118.
- 408 (26) Vogt, R.; Sander, R.; von Glasow, R.; Crutzen, P. J. Iodine chemistry and its role in halogen
  409 activation and ozone loss in the marine boundary layer: A model study. *J. Atmos. Chem.* 1999, *32*,
  410 375-395.
- 411 (27) Pillar-Little, E. A.; Guzman, M. I.; Rodriguez, J. M. Conversion of iodide to hypoiodous acid
  412 and iodine in aqueous microdroplets exposed to ozone. *Environ. Sci. Technol.* 2013, 47, 10971413 10979.
- (28) Xing, D.; Yuan, X.; Liang, C.; Jin, T.; Zhang, S.; Zhang, X. Spontaneous oxidation of I<sup>-</sup> in
  water microdroplets and its atmospheric implications. *Chem. Commun.* 2022, *58*, 12447-12450.
- 416 (29) Saiz-Lopez, A.; Plane, J. M. C.; Baker, A. R.; Carpenter, L. J.; Glasow, R. V.; Martín, J. C.
- 417 C.; McFiggans, G.; Saunders, R. W. Atmospheric chemistry of iodine. *Chem. Rev.* 2011, *112*, 1773418 1804.
- 419 (30) Fournier, M. C.; Falk, L.; Villermaux, J. A new parallel competing reaction system for
  420 assessing micromixing efficiency-Experimental approach. *Chem. Eng. Sci.* 1996, *51*, 5053-5064.
- 421 (31) Kireev, S. V.; Shnyrev, S. L. Study of molecular iodine, iodate ions, iodide ions, and triiodide
- ions solutions absorption in the UV and visible light spectral bands. *Laser Phys.* **2015**, *25*, 075602.
- 423 (32) Dixon, S. L.; Jurs, P. C. Estimation of pKa for organic oxyacids using calculated atomic
- 424 charges. J. Comput. Chem. **1993**, 14, 1460-1467.

- 425 (33) Kim, K.; Yabushita, A.; Okumura, M.; Saiz-Lopez, A.; Cuevas, C. A.; Blaszczak-Boxe, C.
- 426 S.; Min, D. W.; Yoon, H.-I.; Choi, W. Production of molecular iodine and tri-iodide in the frozen
- 427 solution of iodide: Implication for polar atmosphere. *Environ. Sci. Technol.* **2016**, *50*, 1280-1287.
- 428 (34) Halfacre, J. W.; Shepson, P. B.; Pratt, K. A., pH-dependent production of molecular chlorine,
  429 bromine, and iodine from frozen saline surfaces. *Atmos. Chem. Phys.* 2019, *19*, 4917-4931.
- 430 (35) Wang, M.; He, X.-C.; Finkenzeller, H.; Iyer, S.; Chen, D.; Shen, J.; Simon, M.; Hofbauer,
- 431 V.; Kirkby, J.; Curtius, J.; Maier, N.; Kurtén, T.; Worsnop, D. R.; Kulmala, M.; Rissanen, M.;
- 432 Volkamer, R.; Tham, Y. J.; Donahue, N. M.; Sipilä, M. Measurement of iodine species and sulfuric
- 433 acid using bromide chemical ionization mass spectrometers. *Atmos. Meas. Tech.* 2021, 14, 4187-
- 434 4202.
- (36) Riva, M.; Ehn, M.; Li, D.; Tomaz, S.; Bourgain, F.; Perrier, S.; George, C. CI-Orbitrap: An
  analytical instrument to study atmospheric reactive organic species. *Anal. Chem.* 2019, *91*, 94199423.
- 438 (37) Riva, M.; Brüggemann, M.; Li, D.; Perrier, S.; George, C.; Herrmann, H.; Berndt, T.
- Capability of CI-Orbitrap for gas-phase analysis in atmospheric chemistry: A comparison with the
  CI-APi-TOF technique. *Anal. Chem.* 2020, *92*, 8142-8150.
- 441 (38) Zhao, L.; Song, X.; Gong, C.; Zhang, D.; Wang, R.; Zare, R. N.; Zhang, X. Sprayed water 442 microdroplets containing dissolved pyridine spontaneously generate pyridyl anions. *P. Natl. Acad.*
- 443 *Sci. USA* **2022**, *119*, e2200991119.
- 444 (39) Bhattacharyya, D.; Mizuno, H.; Rizzuto, A. M.; Zhang, Y.; Saykally, R. J.; Bradforth, S. E.
- New insights into the charge-transfer-to-solvent spectrum of aqueous iodide: Surface versus bulk. *J. Phys. Chem. Lett.* 2020, *11*, 1656-1661.
- (40) Peláez, R. J.; Blondel, C.; Delsart, C.; Drag, C. Pulsed photodetachment microscopy and the
  electron affinity of iodine. *J. Phys B: At. Mol. Opt.* 2009, *42*, 125001.
- 449 (41) Kong, X.; Castarede, D.; Thomson, E. S.; Boucly, A.; Artiglia, L.; Ammann, M.; Gladich, I.;
- Pettersson, J. B. C. A surface-promoted redox reaction occurs spontaneously on solvating inorganic
  aerosol surfaces. *Science* 2021, *374*, 747-752.
- (42) Ammann, M.; Artiglia, L. Solvation, surface propensity, and chemical reactions of solutes at
  atmospheric liquid-vapor interfaces. *Acc. Chem. Res.* 2022, *55*, 3641-3651.
- (43) Fauré, N.; Chen, J.; Artiglia, L.; Ammann, M.; Bartels-Rausch, T.; Li, J.; Liu, W.; Wang, S.;
  Kanji, Z. A.; Pettersson, J. B. C.; Thomson, E. S.; Kong, X. Unexpected behavior of chloride and
  sulfate ions upon surface solvation of martian salt analogue. *ACS Earth Space Chem.* 2023, 7, 350-
- 457 359.
- (44) Nishiyama, Y.; Terazima, M.; Kimura, Y. Ultrafast relaxation and reaction of diiodide anion
  after photodissociation of triiodide in room-temperature ionic liquids. *J. Phys. Chem. B* 2012, *116*,
  9023-9032.
- (45) Sagar, D. M.; Bain, C. D.; Verlet, J. R. R. Hydrated electrons at the water/air interface. *J. Am. Chem. Soc.* 2010, *132*, 6917-6919.
- (46) Zechner, J.; Köhler, G.; Getoff, N.; Tatischeff, I.; Klein, R. Wavelength effects on the
  photoprocesses of indole and derivatives in solution. *Photochem. Photobiol.* 1981, *34*, 163-168.
- (47) Zepp, R. G.; Braun, A. M.; Hoigne, J.; Leenheer, J. A. Photoproduction of hydrated electrons
  from natural organic solutes in aquatic environments. *Environ. Sci. Technol.* 1987, *21*, 485-490.
- 467 (48) Pöschl, U.; Letzel, T.; Schauer, C.; Niessner, R. Interaction of ozone and water vapor with 468 spark discharge soot aerosol particles coated with  $benzo[a]pyrene: O_3$  and H<sub>2</sub>O adsorption,

- 469 benzo[a]pyrene degradation, and atmospheric implications. J. Phys. Chem. A 2001, 105, 4029470 4041.
- (49) Mmereki, B. T.; Donaldson, D. J. Direct observation of the kinetics of an atmospherically
  important reaction at the air-aqueous interface. *J. Phys. Chem. A* 2003, *107*, 11038-11042.
- 473 (50) Reeser, D. I.; Donaldson, D. J. Influence of water surface properties on the heterogeneous 474 reaction between  $O_{3(g)}$  and  $I_{(aq)}$ . *Atmos. Environ.* **2011**, *45*, 6116-6120.
- 475 (51) Wren, S. N.; Donaldson, D. J. Glancing-angle Raman spectroscopic probe for reaction
  476 kinetics at water surfaces. *Phys. Chem. Chem. Phys.* 2010, *12*, 2648-2654.
- 477 (52) Sander, R. Compilation of Henry's law constants (version 4.0) for water as solvent. *Atmos.*478 *Chem. Phys.* 2015, *15*, 4399-4981.
- 479 (53) Hayase, S.; Yabushita, A.; Kawasaki, M.; Enami, S.; Hoffmann, M. R.; Colussi, A. J. Weak
  480 acids enhance halogen activation on atmospheric water's surfaces. *J. Phys. Chem. A* 2011, *115*,
  481 4935-4940.
- 482 (54) Watanabe, K.; Matsuda, S.; Cuevas, C. A.; Saiz-Lopez, A.; Yabushita, A.; Nakano, Y.
- 483 Experimental determination of the photooxidation of aqueous I<sup>-</sup> as a source of atmospheric I<sub>2</sub>. *ACS* 484 *Earth Space Chem.* **2019**, *3*, 669-679.
- 485 (55) Martins-Costa, M. T. C.; Anglada, J. M.; Francisco, J. S.; Ruiz-Lopez, M. F. Reactivity of
- atmospherically relevant small radicals at the air–water interface. *Angew. Chem. Int. Ed.* 2012, *51*,
  5413–5417.
- 488 (56) Martins-Costa, M. T. C.; Ruiz-López, M. F. Electrostatics and chemical reactivity at the air–
   489 water interface. J. Am. Chem. Soc. 2023, 145, 1400-1406.
- 490 (57) Sherwen, T.; Schmidt, J. A.; Evans, M. J.; Carpenter, L. J.; Großmann, K.; Eastham, S. D.;
- 491 Jacob, D. J.; Dix, B.; Koenig, T. K.; Sinreich, R.; Ortega, I.; Volkamer, R.; Saiz-Lopez, A.; Prados-
- 492 Roman, C.; Mahajan, A. S.; Ordóñez, C. Global impacts of tropospheric halogens (Cl, Br, I) on
- 493 oxidants and composition in GEOS-Chem. *Atmos. Chem. Phys.* **2016**, *16*, 12239-12271.
- 494 (58) McFiggans, G.; Bale, C. S. E.; Ball, S. M.; Beames, J. M.; Bloss, W. J.; Carpenter, L. J.;
- 495 Dorsey, J.; Dunk, R.; Flynn, M. J.; Furneaux, K. L.; Gallagher, M. W.; Heard, D. E.; Hollingsworth,
- 496 A. M.; Hornsby, K.; Ingham, T.; Jones, C. E.; Jones, R. L.; Kramer, L. J.; Langridge, J. M.; Leblanc,
- 497 C.; LeCrane, J. P.; Lee, J. D.; Leigh, R. J.; Longley, I.; Mahajan, A. S.; Monks, P. S.; Oetjen, H.;
- 498 Orr-Ewing, A. J.; Plane, J. M. C.; Potin, P.; Shillings, A. J. L.; Thomas, F.; von Glasow, R.; Wada,
- 499 R.; Whalley, L. K.; Whitehead, J. D. Iodine-mediated coastal particle formation: An overview of
- the Reactive Halogens in the Marine Boundary Layer (RHaMBLe) Roscoff coastal study. *Atmos.*
- 501 *Chem. Phys.* **2010**, *10*, 2975-2999.
- 502
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**Figure 1.** Spontaneous  $I_3^-$  production from NaI droplets based on UV-Vis measurement. (A) The absorption spectra of NaI,  $I_{2(aq)}$  and  $I_3^-$  standard solutions at selected concentrations. (B) Timedependent absorption spectra for bulk solution (100 mM NaI, pH=6.2) in mist chamber experiment under different spray gases (see inset for enlarged spectra), including pure N<sub>2</sub>, 20% O<sub>2</sub> + 80% N<sub>2</sub>, and 0.72% C<sub>2</sub>H<sub>4</sub>ClOH<sub>(g)</sub> in N<sub>2</sub>. (C) The absorption spectra for droplets produced by atomizing the same 100 mM NaI bulk solution under different carrier gases, which were collected as condensing liquid in two tandem bottles as shown in Figure S2B.



**Figure 2.** I<sub>3</sub><sup>-</sup> production rate from mist chamber experiments under different conditions. (A) Fixed NaI bulk solution (100 mM) under different spraying gases. (B) Varied NaI bulk solution (0.1–100 mM) under the same spraying gas (0.72% C<sub>2</sub>H<sub>4</sub>ClOH<sub>(g)</sub> in N<sub>2</sub>). Error bars represented standard deviation (1 $\sigma$ ) from 2–10 measurements from different spraying time (Tables S1–S2).

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Figure 3. Direct  $I_{2(g)}$  measurement from NaI-containing droplets under various experimental conditions. (A) Time-dependent raw  $I_{2(g)}$  intensity measured at the output of atomizer by atomizing 100 mM NaI under N<sub>2</sub> and air, respectively. (B)  $I_{2(g)}$  intensity measured at the output of atomizer or mist chamber with 100 mM or 10 mM NaI under air vs. N<sub>2</sub>. (C)  $I_{2(g)}$  intensity measured at the output of atomizer with 10 mM NaI under different solution pH. (D)  $I_{2(g)}$  intensity measured at the output of atomizer as a function of bulk NaI concentration.

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Figure 4. Observed  $I_3^-$  and  $I_{2(g)}$  production from spontaneous iodide activation as a function of bulk NaI concentration. The data from Figure 2B and Figure 3D were normalized to their respective maxima and the combined data fit with a Langmuir–Hinshelwood model.