

# Palaeoecological reconstruction of Plio-Pleistocene herbivores from the Ahl al Oughlam site (Casablanca, Morocco): Insights from dental wear and stable isotopes

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1	Palaeoecological reconstruction of Plio-Pleistocene herbivores from the
2	Ahl al Oughlam site (Casablanca, Morocco): Insights from dental wear
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4	
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22	
23	Abstract
24	In this article, we use a multiproxy approach based on stable isotope analyses ( $\delta^{18}$ O and

 $\delta^{13}$ C), mesowear, and microwear dental analysis to reconstruct the climate, diet, and

26 habitat of Plio-Pleistocene herbivores at the site of Ahl al Oughlam in western Morocco. 27 This study has been conducted on teeth from several taxa (Sivatherium maurusium, Hipparion pomeli, Metridiochoerus phacochoeroides, Gazella sp., Alcelaphini, and 28 29 Anancus osiris), enabling a robust ecological reconstruction. Our isotopic results suggest a predominance of C<sub>3</sub> plant consumption for all analyzed species, while dental 30 wear indicates a varied diet, with evidence for both browsing and grazing behavior 31 32 being present. The combination of proxies indicates a landscape dominated by open  $C_3$ 33 grasslands in an arid climate, as supported by the Mean Annual Precipitation (MAP) results. These findings contribute to the ongoing debate on the ecological context that 34 35 facilitated the Plio-Pleistocene dispersal, occupation, and evolution of mammals and hominins in northern Africa, for which limited information is available. 36

37

## 38 Keywords

39 North Africa, diet, habitat, stable isotopes, microwear, mesowear.

40

### 41 **1. Introduction**

#### 42 <u>1.2 The archaeopaleontological context during the Plio-Pleistocene in North Africa</u>

43 The richness and diversity of mammalian fossil remains in the Maghreb during the Late 44 Pliocene and Early Pleistocene (Geraads, 2002; Raynal et al., 2002; Sahnouni and Van 45 der Made, 2009; Sahnouni et al., 2011; Van der Made et al., 2021) provides evidence 46 that the Sahara Desert was not always a natural barrier for mammals in the past. The 47 region experienced multiple green and humid phases, known as Green Sahara periods 48 (GSPs), over millions of years, including the Plio-Pleistocene. During these events, the 49 present desert area consisted of river basins and extensive lakes (Van Zinderen Barker, 1962; Larrasoaña et al., 2003, 2013; Castañeda et al., 2009; Grant et al., 2017, 2022; 50

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To date, the oldest artifacts and stone tool-cut marked bones in North Africa associated 55 with the presence of hominins have been found at the site of Aïn Boucherit (~1.9 and 56 57  $\sim$ 2.4 Ma) in the Beni Fouda intramontane basin in northeast Algeria (Sahnouni et al., 58 2018; Duval et al., 2021). In the same area, Acheulean technology has been described at the sites of Aïn Hanech and El Kherba (~1.8 Ma) (Sahnouni et al., 2002, 2012, 2017; 59 Parés et al., 2014; Duval et al., 2021, 2023; Van der Made et al., 2021). In the 60 Constantine Basin (Algeria), biochronologically close to Aïn Hanech, is the little-61 62 known site of Mansourah, with abundant faunal remains and bifacial artifacts associated 63 with the Oldowan industry (Thomas, 1884, 1886; Chaid-Saoudi et al., 2006; Sahnouni and Van der Made, 2009) and the palaeontological site of Aïn Jourdel (~2.5 Ma) 64 (Thomas, 1884; 1886; Duval et al., 2021). The oldest human remains have been found 65 66 in western Algeria at Tighennif, also called Ternifine, ~1 Ma (Arambourg and Hoffstetter, 1963; Geraads et al., 1986; Bocherens et al., 1996; Geraads, 2016; Van der 67 68 Made et al., 2021). In western Morocco, near Casablanca, the earliest archaeological 69 sites date back to the Early and Middle Pleistocene (Raynal et al., 2001, 2002, 2017; 70 Mohib et al., 2019; Gallotti et al., 2022). Up to now, the oldest site with human presence corresponds to unit L of Thomas Quarry I and its abundant ancient Acheulean 71 72 lithic tools (~1.3 Ma) (Raynal and Texier, 1989; Gallotti et al., 2020, 2021, 2023). The former Thomas III Quarry, in which Homo remains were found (Ennouchi, 1975, 73 74 Geraads et al., 1980, 2022a and b; Raynal, 2021) was renamed Oulad Hamida 1 Quarry 75 when the Grotte des Rhinocéros (~0.7 Ma) was discovered and delivered a single Homo

remain associated with stone tools and fauna (Raynal et al., 1993, 2022b; Rhodes et al., 76 77 2006; Daujeard et al., 2012, 2020; Raynal and Mohib, 2016; Gallotti et al., 2022). In Thomas Quarry I, above level L, Hominid Cave (~0.5/0.7 Ma) is also located, where 78 79 early *Homo* remains on the Atlantic shore were discovered (Ennouchi, 1975; Geraads et 80 al., 1980, 2022a; Raynal et al., 2010, 2022a; Daujeard et al., 2016, 2020; Raynal, 2021). On the outskirts of Casablanca sits the site of Ahl al Oughlam (AaO), the richest 81 82 Neogene faunal site in North Africa dated to ~2.5 Ma (Raynal et al., 1990; Geraads, 83 2006; Geraads et al., 2010, 2022b) and whose ecological context is investigated in this 84 work.

85

In eastern Morocco, close to the Algerian border in the Aïn Beni Mathar-Guefaït area 86 (Jerada province), several sites have been discovered and excavated since 2006: the 87 88 Early Pleistocene sites of Aïn Tabouda, Gara Soultana and Charchara 1 with Mode 1 89 lithic tools, and the Middle Pleistocene site of Oued Rabt with Mode 2 lithic tools (Sala-90 Ramos et al., 2022). The oldest site, although palaeontological, is Guefaït. This one is 91 biochronologically located at the Plio-Pleistocene boundary (~3–2.5 Ma; Parés et al., 92 2023), contemporaneous with Ahl al Oughlam, and contains a large diversity of taxa in its fossiliferous levels (Sala-Ramos et al., 2017, 2022; Aouraghe et al., 2019; Piñero et 93 94 al., 2019; Alba et al., 2021; Madurell-Malapeira et al., 2021; Ramírez-Pedraza et al., 95 2023).

96

97 Work based on palaeoecological reconstructions are underrepresented in North Africa, 98 despite the abundance of archaeopalaeontological sites found from the Late Pliocene to 99 the Early Pleistocene in this territory. In Morocco, the studies published to date are 100 focused on the palaeoecology of the Plio-Pleistocene primates of Ahl al Oughlam

(Theropithecus; Fannin et al., 2021) and Guefaït-4.2 (Macaca; Ramírez-Pedraza et al., 101 102 2023). In Algeria, there are more extensive studies available, such as the analysis of 103 carbon and oxygen stable isotopes in herbivores from Tighennif (~1 Ma) (Bocherens et 104 al., 1996; Fannin et al., 2021), and the research based on stable isotopes of pedogenic 105 carbonates from El Kherba (~1.8 Ma) (Sahnouni et al., 2011). However, there still 106 remains a knowledge gap in this area regarding the reconstruction of habitats at the Plio-107 Pleistocene boundary. This contrasts with the numerous and comprehensive studies 108 carried out in eastern (Kingston and Harrison, 2007; Ségalen et al., 2007; Levin et al., 109 2008; Bedaso et al., 2013; Cerling et al., 2015; Robinson et al., 2017, 2021; Lüdecke et al., 2018; Mercader et al., 2021), southern (Lee-Thorp and Sponheimer, 2005; Lee-110 111 Thorp et al., 2007, 2010; Ségalen et al., 2007; Lehmann et al., 2016) and central Africa 112 (Zazzo et al., 2000; Ségalen et al., 2007; Lee-Thorp et al., 2012) focused on the 113 ecological contexts of Plio-Pleistocene hominins and large herbivores.

114

#### 115 <u>**1.3 The distribution of C**<sub>3</sub>-C<sub>4</sub> plants in North Africa</u>

116 One of the commonly used methods for reconstructing past environments has been the analysis of stable carbon isotope ratios  $({}^{13}C/{}^{12}C)$  in dental enamel bioapatite. This 117 biogenic material is particularly useful due to its resistance to diagenetic processes, 118 119 preserving the original isotopic signal (Wang and Cerling, 1994). The carbon isotope 120 ratios are primarily influenced by the differential fractionation between the two photosynthetic pathways ( $C_3$  and  $C_4$ ) used by plants consumed by animals, and it is 121 122 associated with specific types of landscapes or environmental conditions. The proportion of C3 or C4 plants is not distributed homogeneously throughout Africa, 123 124 varying both geographically and chronologically. Global C<sub>4</sub> ecosystem expansion around 8–10 million years ago, due to the decrease of atmospheric CO<sub>2</sub>, higher 125

temperatures and increasing water-stress represents a significant change in the 126 127 landscape that led to the replacement of  $C_3$  grasses by  $C_4$  grasses in existing open and 128 dry environments in low- to mid-latitude and altitude regions (Ehleringer et al., 1991; 129 Cerling, 1992; Cerling et al., 1997; Edwards et al., 2010; Polissar et al., 2019). During the Plio-Pleistocene, eastern, southern and central Africa were dominated by C4 130 grassland (Zazzo et al., 2000; Ségalen et al., 2007; Strömberg, 2011), relegating C<sub>3</sub> 131 132 grasses to altitudes above 2000 meters and in the forest understory (Tieszen et al., 1979; 133 Young and Young, 1983; Balasse and Ambrose, 2005). The ecological context of these 134 areas, characterized by the high abundance of C<sub>4</sub> grasses, especially in Eastern Africa, has allowed the application of the C<sub>4</sub>-grazers and C<sub>3</sub>-browsers classification, which is 135 136 widely used in studies analyzing the diet of mammals based on the analysis of tooth enamel  $\delta^{13}$ C values (Vogel, 1978; Zazzo et al., 2000; Sponheimer and Lee-Thorp, 2003; 137 138 Lee-Thorp and Sponheimer, 2005; Levin et al., 2008; Bedaso et al., 2013; Cerling et al., 139 2015; Uno et al., 2018).

140

141 This clear distinction between C<sub>4</sub>-grazers and C<sub>3</sub>-browsers is not valid at high latitudes, 142 as the contribution of C<sub>4</sub> plants is non-existent or marginal. This pattern has been 143 observed in multiple studies conducted during the Pliocene and Early Pleistocene in 144 Europe (Edwards et al., 2010; Domingo et al., 2013; Szabó et al., 2022). The same 145 observation applies to studies conducted in North Africa on Pleistocene and Pliocene 146 archaeological and palaeontological sites mentioned above (Bocherens et al., 1996; 147 Sahnouni et al., 2011; Fannin et al., 2021). All of them, based on carbon isotope evidence, reveal a vegetation consisting of  $C_3$  plants including all available resources 148 149 such as trees, woody shrubs, bushes and grasses. This makes it challenging to discern different diets and habitats for mammals and therefore to draw a consistent ecologicalreconstruction with this single method.

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153 To address this challenge, we combine three high-resolution approaches that cover different moments in the animal's life: the  $\delta^{18}$ O and  $\delta^{13}$ C values of tooth enamel 154 bioapatite (which reflect the first years of life related to the crown formation times), 155 tooth mesowear (which reflects the last months/years before death) and tooth occlusal 156 157 microwear (which reflects the last days/weeks before death). By combining these approaches, this work aims at providing a robust ecological reconstruction of the Ahl al 158 Oughlam (AaO) mammal communities, which will shed light on their diets, habitats, 159 160 and seasonality.

161 Up to the present, there is no evidence of hominins at AaO. The earliest known 162 occupation on the Atlantic coast, in western Morocco, corresponds to level TH-L at 1.3 163 Ma (Gallotti et al., 2023). However, 2.5 Ma ago, hominins were widespread in East and 164 South Africa (Hammond and Mongle, 2023), in Chad, at the Koro Toro site (3.5 Ma; 165 Brunet et al., 1995), and in Northern Africa with the oldest occupation at Ain Boucherit 166 at 2.4 Ma (Sahnouni et al., 2018). The absence of hominins in AaO does not necessarily imply their absence during the Plio-Pleistocene in the area, as it is possible that they 167 168 have simply not been discovered in the fossil record. Another possibility could be that 169 the environmental conditions did not facilitate the occupation of this area. In this study, we have also compared the ecology of the taxa from AaO with the same taxa from other 170 171 sites in North, South, East, and Central Africa.

Understanding the role of this territory is crucial in contributing to the debate about thefirst hominin dispersions and occupations, as well as highlighting the importance of

- 174 North Africa in the adaptation and development of mammal behavior during the Late
- 175 Pliocene and Early Pleistocene.
- 176

# 177 **2. Site: Ahl al Oughlam**



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Fig.1. Location map of North African sites mentioned in the text. Plio-Pleistocene boundary sites =
orange circle (Ahl al Oughlam (this study), Guefaït and Aïn Jourdel); Early Pleistocene sites = brown
circle (Thomas Quarry I-Unit L, Tighennif, Aïn Boucherit, Aïn Hanech, El Kherba and Mansourah);
Middle Pleistocene sites = green circle (Grotte des Rhinocéros and Grotte à Hominidés).

183

The fossiliferous locality of Ahl al Oughlam (33°34'11"N, 07°30'44"W) is an
abandoned sandstone quarry located at the southeast boundary of the city of Casablanca
(Morocco), 6.5 km from the present-day Atlantic shore (Fig. 1; Raynal et al., 1990;
Geraads et al., 2022b). Stratigraphically, the quarry cuts through calcarenitic intertidal,

188 supratidal and aeolian deposits correlative to sea-level highstands associated with a platform developed about 105 m a.s.l. which belong to the Ahl al Oughlam 189 190 Morphostratigrafic Unit (Lefèvre and Raynal, 2002; Lefèvre et al., 2021; Geraads et al., 2022b), formerly included in the Messaoudian stage of Biberson's Moroccan 191 192 Quaternary Chronostratigraphy (Biberson, 1961a; Stearns, 1978). The site was 193 investigated in 1953 by Biberson (Biberson, 1961a, 1961b), under the name of Carrière 194 Déprez, who reported the occurrence of some broken pebbles that he considered as 195 representing reworked artifacts of the "archaic pebble culture" representing an initial 196 stage of North African technological evolution. Raynal and Texier (1989) demonstrated 197 that they are actually "geofacts" of natural origin found in marine deposits. In 1985, J.-198 P. Raynal and J.-P. Texier discovered abundant fossil faunal remains in karst formations 199 and in a network of fissures and interconnected galleries between calcarenite blocks, at 200 the foot of what was once a palaeoshoreline.

201 The fillings lack stratification. Given that the populations of the diverse rodent species 202 within the interconnected fissures exhibit no significant distinctions in tooth 203 measurements, they can be regarded as synchronous on a geological scale. Since 204 absolute dating and paleomagnetic analysis are unattainable, the age of the fossiliferous 205 fillings hinges on the identification and evolutionary stages of various taxa (Geraads, 206 2002, 2010). A biostratigraphic analysis of the mammal assemblage dates this locality 207 to the Plio-Pleistocene boundary, indicating an approximate age of 2.5 Ma (Geraads, 208 1997; Geraads et al., 2022b). The genus Equus is absent, suggesting that the site is older 209 than 2.3 Ma (Geraads, 2006). The mammalian fauna includes at least fifty-five species 210 of both small and large mammals. A new species of *Hipparion*, *H. pomeli*, has been 211 described subsequently from the site (Eisenmann and Geraads, 2007). The vertebrate 212 fauna also includes amphibians, reptiles (tortoises, crocodiles, amphisbaenids, lizards

and snakes) and fishes. The most unexpected element of the mammalian fauna is a 213 214 walrus, a pinniped marine mammal. No living or fossil walrus was previously known in 215 the Eastern Atlantic, south of Belgium (Geraads, 2006). The site was formed along the oceanic seashore which explains why both marine and terrestrial forms are found among 216 217 birds. A list of the identified faunal taxa is given in several publications (Raynal et al., 218 1990; Geraads, 1993, 1995, 1996, 1997, 2002, 2004, 2006; Alemseged and Geraads, 1998; Geraads and Amani, 1998; Geraads and Metz-Muller, 1999; Eisenmann and 219 220 Geraads, 2007; Mourer-Chauviré and Geraads, 2008; Geraads et al., 2010; Mourer-221 Chauviré and Geraads, 2010; Gunnell et al., 2011; Sen and Geraads, 2023).

222 **3. Material & methods** 

The material sampled and analyzed is currently stored at the *Institut national des Sciences de l'Archéologie et du Patrimoine* (INSAP) in Rabat (Morocco). The species
included in this research are the following: *Sivatherium maurusium*, *Hipparion pomeli*, *Metridiochoerus phacochoeroides*, *Gazella* sp., Alcelaphini and *Anancus osiris*.

227

# 228 <u>3.1 Stable isotopes (<sup>13</sup>C and <sup>18</sup>O)</u>

229

In this study, we performed stable carbon (<sup>13</sup>C) and oxygen (<sup>18</sup>O) isotope analysis of 82
fossil fauna specimens recovered from AaO (Supplementary Table 1). Sampling has
been conducted on the buccal surface of molars and premolars from adult individuals.
The tooth surfaces were cleaned with a tungsten abrasive drill bit to remove any
adhering external material. Enamel powder for bulk analysis was obtained using gentle
abrasion with a diamond-tipped drill.

Powdered enamel samples were chemically treated at the biomolecular laboratory at the *Institut Català de Paleoecologia Humana i Evolució Social* (IPHES - CERCA).
Samples weighed from 3.5 to 9.5 mg. Chemical treatment of samples was based on
protocols originally proposed by Koch et al. (1997) and later modified in Tornero et al.
(2013). Samples were treated for 4 h in 0.1 M acetic acid [CH3COOH] (0.1 ml
solution/0.1 mg of sample), neutralized with distilled water and freeze-dried.

243

The <sup>18</sup>O and <sup>13</sup>C of carbonates were measured using an automated carbonate preparation 244 245 device (KIEL-III) coupled with a gas-ratio mass spectrometer (Finnigan MAT 252) at 246 the environmental isotope laboratory at the University of Arizona, US. Powdered 247 samples were allowed to react with dehydrated phosphoric acid under a vacuum at 248 70°C. Accuracy and precision of the measurements were checked and calibrated using 249 calcium carbonate international standards (NBS-19 and NBS-18). Replicate measurements of the standards during analysis had errors of  $\pm 0.1\%$  for  $\delta^{18}$ O and 250  $\pm 0.08\%$  for  $\delta^{13}$ C (1 sigma). Carbon isotope composition is reported in  $\delta$  notation, where 251  $\delta^{13}C = [(Rsample/Rstandard) - 1], R = {}^{13}C/{}^{12}C$  and  $\delta$  is expressed in per mil,  $\infty$ . Carbon 252 253 and oxygen values are reported relative to PDB (Vienna Pee Dee Belemnite).

254

The modern plants using a C<sub>3</sub> photosynthetic pathway (including most trees, woody shrubs, bushes, herbs and grasses) have an average isotopic  $\delta^{13}$ C value of -28.5‰ (global range -20 to -37‰), with -23‰ as the maximum value for typical C<sub>3</sub> plants and a cut-off of -31.5‰ for closed-canopy forests (Kohn, 2010). In this work, we use the average  $\delta^{13}$ C value of -27‰ for modern C<sub>3</sub> plants, excluding analyses from the understory of closed-canopy forests below -31.5‰ (Kohn, 2010). Modern plants that employ the C<sub>4</sub> photosynthetic pathway (primarily tropical grasses and sedges) have an average isotopic composition of -12‰ (global range -10 to -14‰) (Cerling et al., 1997;
Cerling and Harris, 1999; Kohn and Cerling, 2002). The bivariate graphs were made
with the Past 4.02 software (Hammer et al., 2001). All mean values in the text are
followed by their standard deviation.

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- 267

# 268 <u>3.1.2 Estimation of enrichment factor, $%C_{4,}\delta^{13}C_{diet-meq}$ and MAP</u>

269

270 Abundance of atmospheric  $CO_2$  has been changing over time (Tipple et al., 2010) with a consequent influence on the carbon values of plants during its fixation in the 271 photosynthesis process (Farquhar et al., 1989). To make a comparison between the 272 273 carbon values of present-day plants with those consumed by animals in the past, it is crucial to perform corrections. In our work, we have corrected the  $\delta^{13}C$  values of 274 modern plants by +1.5‰, due to the difference between the pre-Industrial  $\delta^{13}C_{CO2}$  value 275 of ~ -6.5% established for the Plio-Pleistocene (Tipple et al., 2010), and the  $\delta^{13}C_{CO2}$ 276 value of  $\sim -8\%$  at the time the reference collection of modern plants was collected 277 278 (Kohn, 2010). When plants are consumed, a fixation occurs in the tooth enamel of the animals, a process of isotopic fractionation (Krueger and Sullivan, 1984; Koch, 1998; 279 280 Cerling and Harris, 1999; Passey et al., 2005). In this study, we use formula (1) in 281 Tejada et al. 2018 which determines that there is a positive correlation between body mass and enamel-diet enrichment fractionation (expressed as  $\varepsilon^*_{enamel-diet}$ ). The  $\varepsilon^*_{enamel-diet}$ 282 283 has been calculated from the estimation of the body mass of the species analyzed: S. 284 maurusium (Basu et al., 2016), H. pomeli (Bernor et al., 2021), Gazella sp. and 285 Alcelaphini (Bibi and Kiessling, 2015), M. phacochoeroides (Rannikko, 2019) and A. 286 osiris (Larramendi, 2016).

287

288 (1) 
$$\epsilon^* = 2.4 + 0.034$$
 (Body Mass)

289

The  $\delta^{13}$ C values in plants and animals have also been proposed to estimate mean annual precipitation (MAP) in the fossil record. We use the formulae (2 and 3) developed in Kohn (2010) and Kohn et al. (2015) to estimate the MAP, which applies corrections for altitude, latitude, and  $\delta^{13}$ C of atmospheric CO<sub>2</sub> from the average modern equivalent of diet composition ( $\delta^{13}$ C<sub>diet-meq</sub> =  $\delta^{13}$ C<sub>enamel</sub> -  $\epsilon^*_{enamel-diet}$  - Plio-Pleistocene atmospheric CO<sub>2</sub>):

296

297 (2) MAP = 
$$10 \wedge [\Delta^{13}C - 2.01 + 0.000198 \times elev - 0.0129 \times |lat| / 5.88] - 300$$

298

299 where *elev* is elevation in meters, *lat* is latitude in degrees, and

300

301 (3) 
$$\Delta^{13}C = (\delta^{13}C_{atm} - \delta^{13}C_{diet-meq})/(1 + \delta^{13}C_{diet-meq}/1000)$$

302

where  $\delta^{13}C_{atm}$  is the carbon isotope of atmospheric CO<sub>2</sub> (-6.5‰ in the Plio-Pleistocene; Tipple et al., 2010) and  $\delta^{13}C_{diet-meq}$  is the modern equivalent of diet composition ( $\delta^{13}C_{diet-meq}$ ) estimated from  $\delta^{13}C_{enamel}$  after  $\epsilon^*_{enamel-diet}$  correction (see formula (1)).

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We estimated (4) the fraction of C<sub>4</sub> vegetation consumed by AaO herbivores using  $\delta^{13}$ C modern plant values, -24‰ for C<sub>3</sub> vegetation ecosystems (Kohn, 2010) and -12‰ for C<sub>4</sub> vegetation (Cerling et al., 1997). The fossil  $\delta^{13}$ C<sub>enamel</sub> values have been corrected for  $\delta^{13}$ C of Plio-Pleistocene atmospheric CO<sub>2</sub> and  $\epsilon^*_{enamel-diet}$  to obtain the  $\delta^{13}$ C<sub>diet-meq</sub>.

312 (4) 
$$%_{C_4} = 100 \times (\delta^{13}C_{diet-meq} - \delta^{13}C_{C_3})/(\delta^{13}C_{C_4} - \delta^{13}C_{C_3})$$

313

# 314 <u>3.1.3 Obligate drinkers vs. non-obligate drinkers taxa</u>

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316 To calculate relative aridity in AaO, we relied on the fundamental principle of different 317 drinking behaviors exhibited by mammals (Nicholson, 1985; Kohn, 1996; Levin et al., 318 2006; Blumenthal et al., 2017). In arid environments, both open water sources and plant 319 water and tissues experience evaporation of <sup>16</sup>O and enrichment of <sup>18</sup>O (Gonfiantini et 320 al., 1965). However, this evaporative effect is more pronounced in plants. As a result, it 321 is anticipated that animals primarily obtaining water through plant consumption (non-322 obligate drinkers) will display higher  $\delta^{18}O_{enamel}$  values compared to obligate drinkers 323 relying on direct access to surface water (Bocherens et al., 1996; Kohn, 1996). The 324 greater the difference in  $\delta^{18}$ O values between these two groups, the higher the aridity 325 (Lehmann et al., 2016; Roberts et al., 2018; Lüdecke et al., 2022). For this comparison, 326 we selected the most reliable species from AaO to classify the mentioned categories. 327 Specifically, we chose S. maurusium and Gazella sp. as non-obligate drinkers, and H. 328 pomeli and A. osiris as obligate drinkers. Generally, grazers among ungulates have a 329 higher dependence on surface water than browsers (Western, 1975; Faith, 2018). This 330 pattern is supported by the dental wear results from AaO, which classified S. maurusium 331 and Gazella sp. as browsers and H. pomeli and A. osiris as grazers.

- 332
- 333 <u>3.2 Dental wear analysis</u>
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335 <u>3.2.1 Dental mesowear analysis</u>

337 A total of 74 molars were studied using mesowear analysis, excluding 28 teeth that were 338 unworn, extremely worn, or had broken or damaged cusp apices (Supplementary 339 
**Table 2**). The sensitivity of cusp sharpness to ontogenetic age has been documented for
 340 both young and old individuals. Conversely, mesowear displays a weaker correlation 341 with age in intermediate age groups, and demonstrates a stronger relationship with 342 dietary factors (Rivals et al., 2007). The mesowear analysis method assesses the relief 343 and sharpness of worn cusp apices, which are related to attritive and abrasive wear on 344 dental enamel. These analyses were performed on the buccal side of upper molars and 345 the lingual side of lower molars, with a preference for the paracone of the upper M2. To 346 evaluate wear on teeth from P3 to M3, we used the extended mesowear method 347 proposed by Kaiser and Solounias (2003). In this study, we applied the method 348 introduced by Mihlbachler et al. (2011), which uses seven cusp categories ranging from 349 high and sharp (stage 0) to completely blunt with no relief (stage 6). A low-abrasion 350 diet such as browsing, which induces high attrition, preserves sharpened apices on 351 buccal cusps as teeth wear down, while a high-abrasion diet results in more rounded and 352 blunted buccal cusp apices. The bivariate graph was performed with the Past 4.02 software (Hammer et al., 2001). 353

354

#### 355 <u>3.2.2 Dental microwear analysis</u>

356

The investigation of dental microwear has a rich history in elucidating the dietary preferences and ecological niches of fossil species. Starting with early studies in the late 1970s, which employed scanning electron microscopy (SEM) at high magnifications (~100–500x) (Rensberger, 1978; Walker et al., 1978), to subsequent research employing low-magnification light microscopy (MacFadden et al., 1999; Solounias and Semprebon, 2002), and more recently, dental microwear texture analysis (DMTA) utilizing three-dimensional surface measurements acquired through white-light confocal microscopy and scale-sensitive fractal analysis (Ungar et al., 2003; Scott et al., 2006). Despite all methods being valid and complementary, for this study, we opted for at lowmagnification microwear approach due to the accessibility of an extensive microwear reference collection of extant species, particularly those of African origin (Solounias and Semprebon, 2002).

369 In this work, a total of 109 molars and premolars were selected and molded. After 370 examination of the molds under a stereomicroscope, 49 specimens were discarded due to poor preservation or taphonomic defects (Supplementary Table 2) (El-Zaatari, 371 372 2010; King et al., 1999). Examination of enamel microwear features was conducted 373 through standard light microscopy, using a Zeiss Stemi 2000C stereomicroscope at a 374 magnification of 35x on high-resolution epoxy casts of the teeth. The cleaning, molding, 375 casting, and examination protocols, as well as feature classification criteria, were 376 consistent with those established by Semprebon and Solounias (2002) and Semprebon et 377 al. (2004). Upper and lower teeth, spanning from P4 to M3, were included in the 378 analysis following Xafis et al. (2017). The occlusal surface of each specimen was first cleaned with acetone and 96% ethanol before being molded with high-resolution 379 380 silicone (Heraeus Kulzer, PROVIL novo Vinylpolysiloxane, Light C.D. 2 regular set) 381 and transparent casts were created using clear epoxy resin (C.T.S. Spain, EPO 150 + K151). Scoring of microwear features, including small and large pits (round scars), 382 383 scratches (elongated scars with parallel sides), gouges (irregular edges and much larger 384 and deeper than the large pits), puncture pits (deepest at their centers, symmetrical, 385 craterlike features with regular margins), and cross scratches (oriented perpendicularly 386 to the majority of scratches), was performed using a standard 0.16  $\text{mm}^2$  ocular reticle. In 387 order to achieve a more precise categorization of grazers, browsers, and mixed feeders, 388 we used the percentage of individuals in a sample possessing scratch numbers that fall 389 between 0 and 17 (Semprebon and Rivals, 2007). The scratch width score (SWS) was calculated based on the ratio of fine to coarse scratches ("0" predominantly fine 390 scratches per tooth surface, "1" to those with a mixture of fine and coarse scratches, and 391 "2" to those with predominantly coarse scratches). The results were compared with a 392 393 database constructed from extant ungulate taxa (Solounias and Semprebon, 2002; Rivals 394 and Semprebon, 2011). To reduce inter-observer error, dental microwear analysis was 395 conducted by a single experienced researcher (IRP).

396

397 **4. Results** 

398

# 399 <u>4.1 Stable isotopes ( $\delta^{13}$ C and $\delta^{18}$ O)</u>

400

The  $\delta^{13}C$  values of all taxa measured in AaO ranged from -14.1 to -8‰ with a mean and 401 402 standard deviation of  $-11.4 \pm 1.2\%$  (n = 82) (Fig. 2). The data for each species are 403 described below: *H. pomeli* (n = 16; range = -13.5 to -10.6%; mean = -11.7  $\pm 0.9$ %), Alcelaphini (n = 18; range = -13 to -8‰; mean = -10.3 ± 1.3‰), Gazella sp. (n = 18; 404 405 range = -12.8 to -10.6%; mean = -11.6  $\pm$  0.6%), S. maurusium (n = 6; range = -14.1 to -10.8%; mean =  $-12.4 \pm 1.4\%$ ), *M. phacochoeroides* (*n* = 23; range = -13.2 to -8.4%; 406 407 mean =  $-11.7 \pm 1.2\%$ ) and A. osiris (n = 1; -13.3%). The two taxa with the highest 408 variability and the most positive values are Alcelaphini and M. phacochoeroides. Gazella sp. and H. pomeli are the species with the lowest variability and S. maurusium 409 410 is the species with the most negative values.



412

413

414 Fig. 2. Jitter boxplots for  $\delta^{13}$ C values of the ungulates from the Plio-Pleistocene site of AaO. Boxes = first 415 and third quartile for each taxon, whiskers = full range for each taxon and hashes within boxes = median 416 value.

417

425

The  $\delta^{18}$ O values ranged from -3 to 2.9‰ with a mean and standard deviation of -0.6 ± 1.2‰ (n = 82) (**Fig. 3**). The data for each species are described below: *H. pomeli* (n =16; range = -3 to 0.2‰; mean = -1.8 ± 0.9‰), Alcelaphini (n = 18; range = -2.2 to 0.6‰; mean = -0.8 ± 0.7‰), *Gazella* sp. (n = 18; range = -2.3 to 2.1‰; mean = -0.2 ± 1.3‰), *S. maurusium* (n = 6; range = -0.9 to 2.9‰; mean = 1.2 ± 1.3‰), *M. phacochoeroides* (n = 23; range = -1.8 to 2‰; mean = -0.4 ± 0.9‰) and *A. osiris* (n =1; -1.4‰). The whole isotopic dataset is provided in **Supplementary Table 1**.



427

428 Fig. 3. Jitter boxplots for  $\delta^{18}$ O values of the ungulates from the Plio-Pleistocene site of AaO. Boxes = first 429 and third quartile for each taxon, whiskers = full range for each taxon and hashes within boxes = median 430 value.

431

# 432 <u>4.1.2 Estimation of enrichment factor, $%C_{4,}\delta^{13}C_{diet-meq}$ and MAP</u>

433

The  $\varepsilon^*_{enamel-diet}$  of the AaO species ranges from 12.4 to 14.7‰, depending on the body 434 435 mass. C<sub>4</sub> plants are virtually absent in the diet of AaO species. They are only present, in low percentages, in four individuals of Alcelaphini ( $%C_4 = 4, 5, 8 \text{ and } 11\%$ ) and in one 436 437 individual of *M. phacochoeroides* ( $%C_4 = 6\%$ ) (Supplementary Table 1). The MAP results are estimated in Table 1. In order to calculate it, the five individuals with C<sub>4</sub> 438 presence have been excluded based on the criteria established by Kohn (2010). The 439 average  $\delta^{13}C_{diet\text{-meq}}$  value (-26.1‰) and MAP (385 mm/yr) suggest a relatively arid 440 441 environment for AaO, defined by Kohn (2010) as ecosystems with a MAP  $\leq$  500 mm/yr and  $\delta^{13}C_{diet-meq}$  values between -26 and -24‰ (Table 1; Supplementary Table 1). 442

**444 Table 1.** Summary of  $\delta^{13}C$  (‰), Body Mass (Kg),  $\epsilon^*_{enamel-diet}$  (‰),  $\delta^{13}C_{diet-meq}$  (‰), %C<sub>4</sub> and

Mean

445 Annual Precipitation (MAP - mm/yr).

446

**Body Mass**  $\delta^{13}C_{diet-meq}$  (%)  $\delta^{13}C$  (‰) **Species**  $\epsilon^{*}_{enamel-diet}$  (%) %C<sub>4</sub> n MAP (mm/yr) (Kg) S. maurusium 6  $-12.4 \pm 1.4$ 1236 +14 $-27.9 \pm 1.4$ 0  $1041\pm760$ Alcelaphini  $-10.3 \pm 1.3$  $-23.9\pm1.3$ 18 200 +13.2 $1.5\pm3.2$  $177\pm210$  $\textbf{-11.6} \pm 0.6$ *Gazella* sp. 30 +12.4 $\textbf{-24.1} \pm 0.6$ 0  $139 \pm 112$ 18  $-11.7 \pm 0.9$  $-26.5\pm0.9$ 0  $400\pm269$ H. pomeli 16 250 +13.3 $-11.7\pm1.2$ 297 +13.4 $-26.5\pm1.2$  $0.3\pm1.3$  $455\pm274$ 23 *M. phacochoeroides* A. osiris 1 -13.3 5000 +14.7-29.6 0 1996  $-11.4 \pm 1.2$  $-26.1 \pm 1.4$ 82 -+13.2 $0.4\pm1.7$  $385 \pm 419$ All species

447

#### 448 4.1.3 Obligate drinkers vs. non-Obligate drinkers taxa

449

The ANOVA statistical test and Tukey's pairwise comparison show significant differences in the  $\delta^{18}O_{\text{enamel}}$  values among the obligate drinkers (n = 17; *A. osiris* and *H. pomeli*) and non-obligate drinkers (n = 24; *S. maurusium* and *Gazella* sp.) with a *p*value of 1.819. The non-obligate drinkers mean is 0.15‰ and the obligate drinker mean is-1.77‰, implying a difference between the two groups of 1.9‰.

455

#### 456 <u>4.2 Dental wear analysis</u>

457

Table 2. Summary of mesowear and microwear data of the fossil assemblages from AaO. Abbreviations: *n*, number of specimens; MWS, mesowear score; NS, average number of scratches; NP, average number
of pits; SD, standard deviation; SWS, scratch width score; %PP, percentage of specimens with puncture
pits; %HC, percentage of specimens with hyper-coarse scratches; %LP, percentage of specimens with

462	large pits; %G, percentage of specimens with gouges; %XS, percentage of specimens with cross
463	scratches; %0–17, percentage of specimens with number of scratches ranging between 0 and 17.
464	
465	
466	
467	
468	
469	

Species	Mesowear		Microwear											
•	n	MWS	n	NS	SD	NP	SD	SWS	%PP	%HC	%LP	%G	%XC	<b>%0–17</b>
S. maurusium	4	1	4	12.8	1.9	24.75	2.7	1	0	0	50	100	25	100
Alcelaphini	22	2.2	17	20.0	2.6	20.1	3.4	1.1	11.8	0	82.4	29.4	11.8	17.6
Gazella sp.	17	1.2	8	9.5	1.6	23.4	4.6	1	0.0	0	100	50	12.5	100
H. pomeli	6	4.3	13	20.6	3	19.1	4	1	15.4	0	84.6	7.7	0	0
M. phacochoeroides	-	-	18	19.7	4.6	23	3.6	0.9	5.6	0	83.3	22	100	33.3
A. osiris	-	-	1	26	-	25.5	-	1	0	100	100	0	100	0

#### 470

#### 471 <u>4.2.1 Dental mesowear analysis</u>

472 Some authors have observed differences in the mesowear score between the upper and

473 lower molars of the same individuals (Kaiser and Fortelius, 2003). Since the *Gazella* sp.

and Alcelaphini teeth analysed include both upper and lower molars, we conducted an

475 ANOVA statistical test and Tukey's pairwise comparison to enhance the analysis

476 conducted. The statistical results did not reveal significant differences in the mesowear

477 score values between lower molars (n = 5) and upper molars (n = 12) of *Gazella* sp.,

478 with a *p*-value of 0.9694, nor between lower molars (n = 13) and upper molars (n = 9)

479 of Alcelaphini, with a *p*-value of 0.9826.

480 The mesowear pattern of AaO specimens is characterized by occlusal surfaces ranging 481 from sharp cusps and high relief (low abrasion) to low relief with rounded cusps (high 482 abrasion). It corresponds to mesowear scores with values from 0 to 5 (Table 2; Fig. 4). All three Bovidae have sharp cusps, which means that attrition predominates strongly 483 484 over abrasion. The S. maurusium has the lowest level of abrasion with a MWS of 1, 485 falling in the MWS average of extant leaf browsers. The Gazella sp. has also a low average MWS (1.2), being in the range of MWS of extant mixed feeders but with a 486 487 tendency towards a browsing diet. The Alcelaphini, with variability in abrasion levels, 488 has an average MWS of 2.2 and is categorized as a grass-dominated mixed feeder. In 489 the case of the Equidae H. pomeli, the MWS indicates high levels of abrasion, with a 490 MWS average of 4.3. This pattern places *H. pomeli* among extant grazers.





494 Fig. 4. Jitter boxplot with mesowear scores (MWS) of the ungulates from AaO. Boxes = first and third
495 quartile for each taxon; whiskers = full range for each taxon and hashes within boxes = median value.
496 Extant ungulates database (Fortelius and Solounias, 2000).

- 497
- 498 4.2.2 Dental microwear analysis
- 499

500 At AaO, all herbivores display browsing or mixed feeding dietary traits (Fig. 5). The 501 species analyzed are classified into two groups. The first group includes the S. maurusium and Gazella sp. and plot among extant browser ungulates. The second 502 503 group, which includes Alcelaphini, H. pomeli, M. phacochoeroides and A. osiris, plot 504 among extant mixed feeders. The first group has microwear patterns characterized by a 505 high number of pits (NP = 24.1) and a small number of scratches (NS = 11.2) and the 506 second group has a microwear pattern defined by a similar number of pits (NP = 21.9) 507 and scratches (NS = 21.6). Gazella sp. (NS = 9.5) and S. maurusium (NS = 12.8) are the 508 species with a lower number of scratches, and H. pomeli (NS = 20.6) and A. osiris (NS 509 = 26) have a higher number of scratches. All species have a similar and a relatively high number of pits. The A. osiris (NP = 25.5) and S. maurisium (NP = 9.5) have the highest 510 511 abundance of pits. Although the Alcelaphini, H. pomeli, M. phacochoeroides and A. 512 osiris in Figure 5 are classified among extant mixed feeders, the high number of scratches (see Table 2, NS and 0–17%) indicates a high proportion of grasses in their 513 514 diet.

515

The SWS, being between 0.9 and 1.1, indicates a mixture of both fine and coarse
scratches for all species. Hyper-coarse scratches have been found only in *A. osiris*.
Cross scratches are absent in *H. pomeli* and present in a low percentage in Alcelaphini, *Gazella* sp. and *S. maurusium*. *A. osiris* and *M. phacochoeroides* have them in 100% of

the individuals. This trait is characteristic of animals with a bunodont dental
morphology as in the case of the latter two species. Gouges have been observed in all
species except in *A. osiris* and their presence is remarkable in the four individuals of *S. maurusium*. Puncture pits are present in a low percentage in Alcelaphini, *H. pomeli* and *M. phacochoeroides*, and large pits are present in all species in proportions varying
from 50% to 100% of the individuals.

526

527





**Fig. 5.** Bivariate plot of the average numbers of pits and scratches for the species from AaO. Error bars correspond to standard deviation ( $\pm$  1 SD) for the fossil samples. Plain ellipses correspond to the Gaussian confidence ellipses (p = 0.95) on the centroid for the extant browsers and grazers based on the database from Solounias and Semprebon (2002).

534

## 535 **5. Discussion**

536

#### 537 <u>5.1 Palaeodiet</u>

538

539 The combination of the three approaches allows us to characterize the dietary 540 preferences of the different taxa, as well as to reconstruct the general ecological context (Table 3). In AaO, all species have a predominantly plant-based C<sub>3</sub> diet with values 541 ranging from -14 to -8‰, with a small consumption of  $C_4$  plants in Alcelaphini and M. 542 phacochoeroides (Table 1). In some contexts, interpreting the carbon isotopic signal 543 544 can be complex, especially when there are no extreme values that can indicate high consumption of C<sub>4</sub> plants linked to grass consumption (ca.  $\delta^{13}C > 0$ %; Cerling and 545 Harris, 1999) or the presence of  $C_3$  closed canopy forests (ca.  $\delta^{13}C < -16\%$ ; Kohn, 546 547 2010), types of environments in which the signal is obvious. In AaO, there is both inter-548 and intra-species variability in carbon values. Recent isotopic studies on extant species 549 in sub-Saharan Africa have shown a weak relationship between intra-taxon variation in 550  $\delta^{13}C_{enamel}$  (C<sub>3</sub> and C<sub>4</sub>) and corresponding changes in herbivore diet and vegetation 551 structure (woody cover). Therefore, adopting a community-level approach when making 552 inferences about vegetation using isotopic data, focusing on the relationships between 553 different taxa is recommended. This approach allows for a more accurate understanding 554 of the vegetation dynamics and ecological context (Robinson et al., 2021). In AaO, the 555 interspecific difference between Alcelaphini and S. maurusium is striking and could 556 represent different ecological niches and diets (Fig. 2). Dental wear analysis (micro- and 557 mesowear) shows variable feeding traits that range from browser to mixed-feeder to 558 grazer behavior, including Alcelaphini with a grass-dominated, mixed-feeding diet or S.

559 maurusium as the most browsing taxon (Table 3). Oxygen isotope data are consistent 560 with the dental wear feeding pattern, with browser species having more positive values 561 (S. maurusium and Gazella sp.) than the rest, and grass-dominated mixed feeders having 562 more negative values. The correlation in oxygen is higher between the most grazing 563 species, according to dental wear pattern, such as H. pomeli, and the most browsing 564 ones such as S. maurusium, showing very distant values between them. This could 565 suggest the consumption of different resources, such as the preference for grasses by the 566 former and leaves by the latter.

567

568**Table 3.** Summary of dietary interpretations based on tooth wear and stable carbon isotope data.569Abbreviations: B = browser, MF-B = browsing-dominated mixed feeder, MF = mixed feeder, MF-G =570grass-dominated mixed feeder, and G = grazer. The parentheses indicate the consumption of C<sub>4</sub> plants (in571very low percentage) in some species.

572

<b>G</b>	Stable					
Species	isotopes	Mesowear	Microwear			
S. maurusium	C <sub>3</sub>	В	В			
Alcelaphini	C <sub>3</sub> (C <sub>4</sub> )	MF-G	MF-G			
Gazella sp.	C <sub>3</sub>	MF-B	В			
H. pomeli	C <sub>3</sub>	G	MF-G			
M. phacochoeroides	C <sub>3</sub> (C <sub>4</sub> )	-	MF-G			
A. osiris	C <sub>3</sub>	-	MF-G			

573

574 <u>5.1.1 Comparison of AaO with other African sites</u>

The *Sivatherium maurusium* from AaO has the most negative C<sub>3</sub> carbon values ( $\delta^{13}$ C 576 mean =  $-12.4 \pm 1.4\%$ ) and the most positive oxygen values (Fig. 2 and 3) compared to 577 578 other AaO species. These oxygen and carbon values could be attributed to the preferential consumption of leaves from the upper parts of trees or shrubs which is 579 580 consistent with dental wear patterns suggesting a browsing diet for AaO S. maurusium. The same isotopic pattern in  $\delta^{13}C$  and  $\delta^{18}O$  values has also been observed in 581 Sivatherium from Hadar and Aramis (Ethiopia; ~4-3 Ma) (Wynn et al., 2016) and 582 583 Makapansgat (South Africa; ~3 Ma (Lee-Thorp et al., 2010). Although this species is 584 known from various North African sites such as Aïn Boucherit, Aïn Hanech and El-Kherba (Early Pleistocene; Algeria), Aïn Jourdel (Late Pliocene; Algeria), and Ichkeul 585 586 and Aïn Brimba (Late Pliocene; Tunisia), this is the first time that data on its diet has 587 been obtained from this area.

588

589 In East Africa, the Giraffidae were represented by two genera: Giraffa and Sivatherium, 590 from the early Pliocene onwards. Sivatherium became extinct during the Middle 591 Pleistocene. Overall, its diet underwent changes in the Turkana Basin (Kenya), being a browser (C<sub>3</sub> plants) at ~4 Ma and switching to grazing (C<sub>4</sub> plants) between ~2.5 and 1 592 Ma (Cerling et al., 2015). In the Chiwondo beds (Malawi; ~2.5 Ma), Sivatherium 593 exhibited less negative  $\delta^{13}$ C values (n = 1;  $\delta^{13}$ C = -7.9‰) than the purely browsing fossil 594 *Giraffa* (n = 2;  $\delta^{13}$ C mean = -12.1 ± 1.1‰), indicating that it incorporated some C<sub>4</sub> grass 595 596 in its diet (Bocherens et al., 2011). In Olduvai (Tanzania; ~1.8 Ma), a multiproxy approach suggested a mixed diet for S. maurusium, C4-dominated or mixed diets, a 597 598 mesowear characteristic of grazers, but with microwear indicating a browsing-599 dominated mixed diet, likely due to seasonal changes in its diet (Uno et al., 2018).

In South Africa, *Sivatherium* maintained its preference for C<sub>3</sub> plants from ~5 to 3 Ma. In 601 the Langebaanweg site (~5 Ma), the  $\delta^{13}$ C mean value (n = 26) is -11.5 ± 1.1‰ (Franz-602 Odendaal et al., 2002). However, hypsodonty indices, dental mesowear, and dental 603 microwear suggested mixed feeding dietary traits with a browsing tendency for this 604 605 species (Franz-Odendaal and Solounias, 2004). In the Makapansgat Limeworks Member 3 (~3 Ma). Sivatherium exhibited carbon values (n = 2:  $\delta^{13}$ C mean = -11.4 ± 606 0.1‰) similar to those of *Giraffa* (n = 8;  $\delta^{13}$ C mean = -11 ± 0.9‰) at the same site 607 (Sponheimer and Lee-Thorp, 2009). The diet of AaO Sivatherium follows the same C<sub>3</sub>-608 609 browser pattern as that of the rest of Africa during the ~3 to 2.5 Ma period, but further 610 multiproxy studies are needed in more recent North African sites such as Aïn Hanech or 611 El-Kherba (~1.8 Ma) to determine whether this dietary pattern remains the same or 612 changes, as in the case of East Africa.

613

614 The Alcelaphini from AaO is the taxon with the least negative carbon values, with a predominance of C<sub>3</sub> plant consumption values (mean  $\delta^{13}C = -10.3 \pm 1.3\%$ ), indicating 615 even some  $C_4$  consumption (Table 2). The results of micro- and mesowear tooth 616 617 analysis indicate a mixed diet, but with a high content of grasses in its feeding ecology. In general, this taxon is indicative of open dry environments, both during the African 618 619 Pliocene and Pleistocene and in extant populations. Regardless of the consumption of 620  $C_3$  or  $C_4$  plants, their feeding behavior is predominantly grazing (Vrba, 1984; Sponheimer et al., 1999; 2003; Bobe and Behrensmeyer, 2004; Cerling et al., 2015). 621

Different species of Alcelaphini were present in North Africa from the Late Pliocene
(Arambourg, 1979; Geraads & Amani, 1998) until they were declared extinct in this
area in 1994 (Mallon and Kingswood, 2001), but the only fossil record of their diet has

been carried out at the site of Tighennif (1 Ma; Algeria). At this site, the Alcelaphini tribe is represented by the species *Connochaetes taurinus prognu*. The pattern of  $\delta^{13}$ C values (n = 6; mean  $\delta^{13}$ C = -10.5 ± 1.4‰) is similar to that observed in the Alcelaphini from AaO (Bocherens et al., 1996; Fannin et al., 2021). Compared to the carbon values of different ungulates analyzed in Tighennif, this species has the most positive carbon values, indicating, as in the case of AaO, a small consumption of C<sub>4</sub> plants in some individuals.

633

In central Africa, 2500 km west of the Rift Valley, at the Koro Toro site (Chad; ~3.5–3

635 Ma), Alcelaphini have a grazing diet based on C<sub>4</sub> plants with an average  $\delta^{13}$ C value (*n* =

636 6) of  $1.4 \pm 0.8\%$  (Zazzo et al., 2000).

637

In eastern Africa (Turkana Basin), the remarkable predominance and similarity in the  $C_4$ isotopic signature of Alcelaphini from approximately ~4.4 Ma to present-day specimens is particularly notable (White et al., 2009; Cerling et al., 2015). In Olduvai (Tanzania; ~1.8 Ma), the carbon isotope data indicate C<sub>4</sub>-dominated diets, while dental wear classifies this species as a grass-dominated mixed feeder (Uno et al., 2018).

643 In southern Africa, Wonderwerk Cave has a longer temporally stratified sequence from 644 the Early Pleistocene to the Holocene, with the presence of Alcelaphini in all layers. 645 They show a progressive increase in C<sub>4</sub> consumption from 50% in the Early Pleistocene 646 to 100% in the Holocene (Ecker et al., 2018). In the Makapansgat Limeworks Member 3 (~3 Ma), Alcelaphini (Parmularius, Damaliscus and Connochaetes) have C4 values, in 647 contrast to the Tragelaphini from the same level, which have a clear C<sub>3</sub> values. 648 649 Ecomorphological and microwear data confirm a grazing diet for the former and a 650 browsing for the latter (Sponheimer et al., 1999; Schubert et al., 2006). In Haasgat cave

(~2.6-1.9 Ma), as in Makapansgat, Alcelaphini have a C<sub>4</sub> consumption while 651 652 Tragelaphini have C<sub>3</sub> plant consumption values (Adams et al., 2013). In Swartkrans 653 cave (~1.9-1.8 Ma), the Alcelaphini Connochaetes sp. also has a C<sub>4</sub>-based diet (Lee-Thorp et al., 1989). However, in other South African sites, the Alcelaphini carbon 654 655 values are clearly C<sub>3</sub>, although attributed to grass consumption. This is the case of Hoedjiespunt (~0.35–0.25 Ma) with a mean  $\delta^{13}$ C (n = 16) of -9.8 ± 1.3‰, Elandsfontein 656  $(\sim 1.0-0.6 \text{ Ma})$  with a mean  $\delta^{13}$ C (n = 19) of  $-10.3 \pm 1.2\%$ , and Langebaanweg ( $\sim 5$  Ma) 657 with a mean  $\delta^{13}$ C (*n* = 7) of -10.8 ± 1‰ (Lehmann et al., 2016). 658

659

The *Hipparion pomeli* from AaO has a diet based on C<sub>3</sub> plants ( $\delta^{13}$ C mean = -11.7 ± 660 661 0.9%). It is the species with the highest mesowear, indicative of a grazing diet. 662 Microwear classifies it as a mixed feeder, but the high number of scratches is related to 663 a high consumption of grasses. The presence of *Hipparion* between ~3.5 and 1.7 Ma is 664 well documented in North Africa in sites such as Aïn Brimba and Lake Ichkeul in 665 Tunisia, Aïn el Bey, Aïn el Hadj Baba, Aïn Jourdel, Mansourah, Aïn Boucherit, Aïn 666 Hanech, Puits Carouby, Carrière Brunie, Carrières Quartier St-Pierre in Algeria, and 667 Fouarat, Oued Akrech, and Guefaït in Morocco. The dating of the youngest North African hipparion record occurs at Aïn Hanech (~1.7 Ma) (Van der Made et al., 2022). 668 669 Regarding studies on the feeding habits of this species in North Africa, there are only 670 preliminary results for the Hipparion sp. from the Guefaït site (~3-2.5 Ma; Morocco), which show a predominance of C<sub>3</sub> plant consumption (n = 4;  $\delta^{13}$ C mean = -9.5 ± 1‰), 671 672 although with less negative carbon values than those of the H. pomeli from AaO 673 (Ramírez-Pedraza et al., 2023).

In Central Africa, diachronic (~6–3 Ma) isotopic studies carried out on the species *Hipparion* cf. *hasumense* in Chad show a clear predominance of C<sub>4</sub> plant consumption. In Kossom Bougoudi (~6–5 Ma), the average  $\delta^{13}$ C (n = 6) is -1.3 ± 0.8‰, and in the Koro Toro site (~3.5–3 Ma) the average  $\delta^{13}$ C value (n = 2) is 0.9 ± 1.1‰ (Zazzo et al., 2000).

680

In East Africa, as was the case with the Alcelaphini, the hipparion equid *Eurygnathohippus* has continuity in C<sub>4</sub>-grazer plant consumption from ~4 to 1.5 Ma (Bocherens et al., 2011; Cerling et al., 2015; Manthi et al., 2020). Other methods such as microwear and mesowear also indicate a grazing diet for the *Eurygnathohippus* from Olduvai (Tanzania; ~1.8 Ma) (Uno et al., 2018).

686

In southern Africa, the *Hipparion* from Langebaanweg (~5 Ma) has a  $\delta^{13}$ C mean value 687 688 (n = 6) of -11.3  $\pm$  0.9‰, implying a consumption of C<sub>3</sub> plants (Franz-Odendaal et al., 689 2002). Other lines of evidence, such as mesowear analysis, in this species are consistent 690 with grazing behavior (Franz-Odendaal et al., 2003). Therefore, it would have a diet 691 based on C<sub>3</sub> grasses, like the *H. pomeli* from AaO. In contrast, at the Makapansgat site (~3 Ma), the *Hipparion* indicates  $C_4$  plant consumption (Sponheimer and Lee-Thorp, 692 693 2009; Lee-Thorp et al., 2010). In general, during the African Plio-Pleistocene, 694 Hipparion had a diet mainly based on the consumption of C<sub>4</sub> grasses. Grass 695 consumption has also been observed in AaO, although with the intake of  $C_3$  plants. 696 Despite the low presence of C<sub>4</sub> plants in North Africa during the study period, some 697 species in AaO show a consumption of C<sub>4</sub> plants, but this is not the case for *H. pomeli*, 698 which may indicate slightly different dietary preferences compared to its relatives in 699 other parts of Africa.

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The *Metridiochoerus phacochoeroides* from AaO shows a predominance of  $C_3$  plant carbon values (mean  $\delta^{13}C = -11.7 \pm 1.2\%$ ), but with significant variability, including a small percentage of  $C_4$  consumption in one individual as observed in Alcelaphini. Microwear analysis shows a mixed diet, although with significant consumption of grasses, like other species classified as mixed feeders. The combination of both proxies indicates a preferential diet based on  $C_3$  grass consumption.

707

The genus *Metridiochoerus* first appeared in Africa ~3.5 Ma. *Metridiochoerus phacochoeroides* is endemic to northern Africa and is distinct from the east and South
African representatives of the genus. Besides AaO, *Metridiochoerus* is rare in North
Africa, but is also known from its type-locality Aïn Jourdel (Thomas, 1884).

712

In East Africa, *Metridiochoerus* between ~4 and 1 Ma were predominantly grazers with high C<sub>4</sub>-grazer proportions (Harris and Cerling, 2002; Cerling et al., 2015; Uno et al., 2018). Dental microwear analysis conducted on *Metridiochoerus* from Olduvai (~1.8 Ma, Tanzania) points to mixed feeding and grazing behavior (Uno et al., 2018). In the Chiwondo Beds in Malawi (~2.5–1.5 Ma), *M. andrewsi* and *M. compactus* exhibit a predominant consumption of C<sub>3</sub> plants (Lüdecke et al., 2016).

719

Little is known about the diet of *Metridiochoerus* in South Africa. Some populations of *M. andrewsi* from Elandsfontein (~1.0–0.6 Ma) were consuming mostly  $C_3$  plants (Luyt et al., 2000), but due to the overall predominance of  $C_3$  plants in this area, as in the Chiwondo Beds in the Malawi Rift, could have been including grasses in their diet as well.

The Gazella sp. from AaO has carbon values that correspond to a C3 plant-based diet 726 (mean  $\delta^{13}C = -11.6 \pm 0.6\%$ ). Microwear analysis shows a browser pattern; however, 727 mesowear analysis indicates a mixed diet with a browsing tendency. This could imply a 728 729 more purely browsing diet over a relatively brief period and a more varied consumption 730 of plants, including grasses in their diet, for most of the year. As with Alcelaphini, Gazella has been present in North Africa since the Pliocene (Sahnouni et al., 2011; Van 731 732 der Made et al., 2021) until the present day in Morocco, Algeria, and Tunisia (Schmid et 733 al., 2005).

734

In North Africa, isotopic studies have been carried out on this species in both Tighennif (~1 Ma; Algeria) with a mean  $\delta^{13}$ C value (n = 5) of -11.9 ± 0.9‰ and Guefaït (~3–2.5 Ma; Morocco) with a mean  $\delta^{13}$ C value (n = 4) of -10.9 ± 1.1‰. At both sites, there is a clear predominance of C<sub>3</sub> plant consumption by *Gazella* (Fannin et al., 2021; Ramírez-Pedraza et al., 2023).

740

741 In East Africa, Gazella janenschi from the Laetolil Beds (~4.3-3.5 Ma) has carbon values ranging from -1.9 to -9.8‰, implying a variable diet from  $C_4$ -grazing to  $C_3$ -742 743 browsing (Kingston and Harrison, 2007). This dietary flexibility has also been observed 744 in different Antilopini analyzed in the Turkana Basin (Kenya), with extracted proportions of the isotopic signal of 17% C<sub>4</sub>-grazing, 75% mixed C<sub>3</sub>-C<sub>4</sub>, and 8% C<sub>3</sub>-745 browsing for the period of ~4.3 to 2.3 Ma, and 54% C<sub>4</sub>-grazing, 46% mixed C<sub>3</sub>-C<sub>4</sub>, and 746 747 0% C<sub>3</sub>-browsing for the period  $\sim$ 2.3 to 1.0 Ma (Cerling et al., 2015). The multiproxy 748 study (isotopes, microwear, and mesowear) of Antilopini from Olduvai (1.8 Ma) also 749 indicates a mixed feeder diet (Uno et al., 2018). The Gazella sp. from AaO, as well as

those from the Pleistocene and Pliocene in Africa in general, seem to have the plasticity to consume distinct types of plants, adapting to seasonal availability. Even in East Africa before ~2.3 Ma, with a predominance of  $C_4$  grassland in the area, they maintain a percentage of contribution from  $C_3$  plant consumption.

754

In South Africa, *Gazella vanhoepeni* from Makapansgat (~3 Ma) is considered either a browser or a mixed-feeder preferring browsing based on isotopic, ecomorphological, microwear, and mesowear dental studies (Schubert et al., 2006). The diachronic and multiproxy study (isotopes, dental microwear texture, and mesowear analysis) of Antilopini (*Antidorcas*) from Sterkfontein Member 4 (~2.8–2 Ma) and Swartkrans Members (~2–0.8 Ma) also indicates a mixed diet with a certain browsing tendency for this genus (Sewell et al., 2019).

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#### 763 5.2 Palaeohabitat and palaeoclimate

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The results of the combination of proxies suggest for AaO a landscape with a predominance of open and dry habitat, composed of C<sub>3</sub> woody plants and grasses and a low presence of C<sub>4</sub> plants. The closed canopy forest is absent in the surroundings of AaO, since no species has  $\delta^{13}C_{diet-meq}$  values below -31.5‰, which is established as a limit for this habitat.

770

#### 771 <u>5.2.1 Dry environment and seasonality</u>

The Mean Annual Precipitation (MAP) in AaO measured 385 mm/year, indicating a dry
climate. The aridity of the region is supported by the fact that environments dominated
by C<sub>3</sub> plants with a MAP of 500 mm/year or less are typically associated with arid

775 conditions. This aridity pattern is also evident in the Tighenif site (1 Ma; Algeria), 776 recording a MAP of 297 mm/year (Supplementary Table 3). This reinforces the 777 enduring prevalence of dry climates even with established hominin presence in North 778 Africa. Kohn (2010) argues that the indicator's reliability depends on the mean MAP of 779 all analyzed taxa. Thus, the average MAP remains the critical factor for assessing the robustness of this indicator. However, the variability in  $\delta^{13}C_{enamel}$ , MAP, and  $\delta^{13}C_{diet-med}$ 780 781 values, both inter- and intra-species (Table 1), suggests the presence of microhabitats 782 (Patalano et al., 2021). In arid regions like those surrounding AaO and Tighenif, there 783 can be localized areas of higher humidity, particularly in close proximity to water sources. Vegetation in these zones may exhibit lower  $\delta^{13}$ C values than expected for the 784 785 surrounding dry environment, especially in a particularly wet year or season (Kohn, 786 2010). This variability in the values could also be explained by the dynamics of the 787 assemblage accumulation. Although the formation data of the site suggest a rapid accumulation over time (Geraads et al., 2022), we cannot precisely determine how long 788 789 it took to form, potentially encompassing many years with diverse climates and 790 resources for consumption.

791

In dry environments, C<sub>4</sub> plants exhibit greater efficiency than C<sub>3</sub> plants under conditions 792 793 of low atmospheric CO<sub>2</sub> concentration, high temperatures, and moisture stress 794 (Ehleringer et al., 1991; Cerling et al., 1993). Therefore, the presence of C<sub>4</sub> plants in the 795 diet of Alcelaphini and *M. phacochoeroides*, even in small proportions, aligns with arid environments. The variability in  $\delta^{13}C_{enamel}$  values within species, especially in the 796 797 aforementioned species, may suggest the existence of pronounced seasonal fluctuations. Relative to aridity, the difference observed in the  $\delta^{18}$ O values of obligate and non-798 799 obligate drinkers in AaO (1.9%) is slightly higher than the difference found in 800 herbivores from the Tsavo (1.7‰) region (Kenya), which has been defined as a dry 801 savannah with vegetation characterized by semi-desert bushland with riparian woodland 802 and a MAP of 500 mm/yr (Cerling et al., 2003; 2015). The species S. maurusium exhibits the highest  $\delta^{18}$ O values, even in comparison to *Gazella* sp. Given that dental 803 wear patterns indicate a browsing diet for both species, the elevated  $\delta^{18}$ O values in S. 804 805 maurusium may be attributed to its consumption of leaves from tall plants that it has 806 access to due to its large size. Sun-exposed leaves, particularly in arid conditions, are highly sensitive to evaporative fractionation, resulting in an enrichment of <sup>18</sup>O 807 808 (Gonfiantini et al., 1965; Barbour, 2007).

In this line, the combination of tooth microwear (short-term) and tooth mesowear (long-809 810 term) proxies has been used as a tool to identify seasonality (Sánchez-Hernández et al., 811 2016). Considering that microwear reflects the signal of a short-term diet, such as a 812 single season, while mesowear reflects a longer period including probably all seasons, 813 discrepancies in the results between both proxies would be indicative of dietary changes 814 due to seasonal availability of resources. In AaO, S. maurusium exhibits a browsing 815 pattern with the two proxies, indicating that its diet does not change throughout the 816 year. For Gazella sp., tooth mesowear indicates browse-dominated mixed feeding, 817 while microwear suggests browsing. This species is likely to have a more diverse diet, 818 characterized by a higher consumption of grasses throughout the year, along with a 819 significant increase in browsing during a particular season. In the case of Alcelaphini 820 and *H. pomeli*, tooth mesowear indicates grass-dominated mixed feeding and grazing, 821 respectively, while microwear suggests mixed feeding. It indicates a predominant 822 consumption of grasses throughout the year with the introduction of other food items 823 during specific periods of the year.

826

827 In all species from AaO, the most discriminant microwear feature is the number of 828 scratches, which is different in browsers and mixed feeders (Fig. 5). The high number 829 of scratches in mixed feeders indicates significant consumption of grasses (see %0-17 830 in **Table 2**). However, the number of pits is very similar between browsers and mixed 831 feeders with grazing tendencies. The relatively high number of pits and the presence of 832 large pits in all species could be caused by incidental ingestion of abrasive grit resulting 833 from feeding close to the ground in open and arid habitats (Semprebon and Rivals, 834 2007; Hoffman et al., 2015). This enables us to emphasize two environmental aspects: 835 (1) The prevalence of taxa with a  $C_3$  plant-based diet, many of which consume a 836 significant amount of grasses, indicating the significant presence of  $C_3$  grasslands in the 837 study area. This assertion is corroborated by previous research on the feeding habits of 838 Theropithecus atlanticus in AaO. In that study, they employed time- and graminivore-839 calibrated carbon stable isotope values to illustrate a progressive transition towards an 840 increased dependence on high-throughput bulk-feeding graminivory and proposed a consumption of grasses exhibiting a distinct  $C_3$  isotopic signature (Fannin et al., 2021). 841 842 (2) The presence of  $C_3$  grassland, feeding close to the ground, and the presence of T. 843 atlanticus (Alemseged and Geraads, 1998; Fannin et al., 2021) are indicators of open 844 habitats. It is noteworthy that no other cercopithecine species have been discovered at 845 the AaO site in Morocco. Additionally, while no fossil remains of Theropithecus have 846 been found at the contemporaneous site of Guefaït in Morocco, the presence of M. cf. 847 sylvanus has been noted (Alba et al., 2021). Previous ecological studies suggest that this 848 macaque inhabited a woodland ecosystem, where it had a folivorous-frugivorous diet with hard item consumption which was primarily based on  $C_3$  plants (Ramirez-Pedraza 849

850 et al., 2023). This indicates a low wood cover in the landscape of western Morocco during the Plio-Pleistocene boundary. In the same line, the relative frequency of two 851 852 species of bovids found in Guefaït and AaO may provide insight into the predominant 853 environmental conditions. Specifically, Tragelaphini are associated with wooded and 854 humid habitats, while Alcelaphini are linked to more open environments (Bobe et al., 855 2007). Notably, the abundance of Alcelaphini and the scarcity of Tragelaphini in AaO is 856 significant (Geraads and Amani, 1998; Geraads et al., 2022a), while the opposite trend 857 is observed in Guefaït (Sala-Ramos et al., 2022; Ramírez-Pedraza et al., 2023). Previous 858 works on AaO have used a method for predicting palaeohabitats using measurements 859 (height/width proportions) of bovid astragali (DeGusta and Vrba, 2003). The proportions of this bone in the bovids from AaO confirm the absence of heavy wood 860 cover and the predominance of open habitat (Geraads et al., 2022a). 861

862

## 863 **6.** Conclusion

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865 This research highlights the importance of integrating different proxies for 866 palaeoecological reconstructions, being particularly useful in areas where there is a 867 predominance of C<sub>3</sub> plants such as in North Africa, where it is more challenging to 868 define the diet and habitat of mammals from stable isotope analysis alone. In this work, 869 the complementarity of dental wear analysis has allowed us to determine diets that 870 range from grazing to browsing behavior. Our multiproxy analysis indicates that during 871 the Plio-Pleistocene boundary in the western coastal region of Morocco, open  $C_3$ 872 grasslands and shrubland predominated in an arid climate.

The most notable distinction between taxa from East, South, and Central Africa in comparison to those from the North is their extensive reliance on  $C_4$  plants, which are 875 associated with tropical grasses from open arid habitats. In the AaO area, although there 876 was a consumption of C<sub>3</sub> plants, the environment was similarly arid and characterized 877 by open grasslands. This aridity was also evident in Tighenif, even with the presence of 878 hominins in the region. Therefore, the dry environment should not have been a 879 limitation to hominins occupation in North Africa. Despite AaO being one of the most 880 abundant fossil fauna sites in North Africa during the Plio-Pleistocene, taxon diversity 881 remains relatively low compared to contemporary sites in the East and South. This 882 could be attributed to the limited diversity of habitats in AaO, such as the absence of 883 woody vegetation or, for instance, the possible lack of permanent water sources, as 884 indicated by the absence of taxa like Hippopotamidae, crocodiles, or freshwater birds 885 (Geraads et al., 2022).

However, further research in different archaeological and palaeontological sites from
the Early Pleistocene and Late Pliocene in the Maghreb is necessary to gain a holistic
understanding of the ecological and climatic dynamics that contributed to the dispersal,
occupation, and evolution of mammals and early hominins.

890

## 891 **CRediT author statement**

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898

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## 927 Appendix A. Supplementary data

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