



**HAL**  
open science

**Co-seismic and postseismic deformation, field  
observations and fault model of the October 30, 2020  
Mw=7.0 Samos earthquake, Aegean Sea**

Athanassios Ganas, Panagiotis Elias, Pierre Briole, Sotiris Valkaniotis, Javier Escartin, Varvara Tsironi, Ilektra Karasante, Chrysanthi Kosma

► **To cite this version:**

Athanassios Ganas, Panagiotis Elias, Pierre Briole, Sotiris Valkaniotis, Javier Escartin, et al.. Co-seismic and postseismic deformation, field observations and fault model of the October 30, 2020 Mw=7.0 Samos earthquake, Aegean Sea. *Acta Geophysica*, 2021, 69 (3), pp.999-1024. 10.1007/s11600-021-00599-1 . hal-04116334

**HAL Id: hal-04116334**

**<https://hal.science/hal-04116334>**

Submitted on 3 Jun 2023

**HAL** is a multi-disciplinary open access archive for the deposit and dissemination of scientific research documents, whether they are published or not. The documents may come from teaching and research institutions in France or abroad, or from public or private research centers.

L'archive ouverte pluridisciplinaire **HAL**, est destinée au dépôt et à la diffusion de documents scientifiques de niveau recherche, publiés ou non, émanant des établissements d'enseignement et de recherche français ou étrangers, des laboratoires publics ou privés.

1           **Co-seismic and postseismic deformation, field observations and fault model of the**  
2           **October 30, 2020  $M_w=7.0$  Samos earthquake, Aegean Sea.**

3  
4           **Athanassios Ganas<sup>1</sup>, Panagiotis Elias<sup>2</sup>, Pierre Briole<sup>3</sup>, Sotiris Valkaniotis<sup>4</sup>, Javier Escartin<sup>3</sup>,**  
5           **Varvara Tsironi<sup>1,5</sup>, Ilektra Karasante<sup>1</sup> and Chrysanthi Kosma<sup>1</sup>.**

6           <sup>1</sup>National Observatory of Athens, Institute of Geodynamics, Lofos Nymfon, Thission, 11810  
7 Athens, Greece; [aganas@noa.gr](mailto:aganas@noa.gr) ; [vtironi@noa.gr](mailto:vtironi@noa.gr) ; [hlekara@hotmail.com](mailto:hlekara@hotmail.com) ;  
8 [chrysanthi.kosma.varela@gmail.com](mailto:chrysanthi.kosma.varela@gmail.com)

9           <sup>2</sup>National Observatory of Athens, Institute of Astronomy, Astrophysics, Space Applications  
10 and Remote Sensing, Vas. Pavlou & I. Metaxa, GR-15 236 Penteli, Greece ; [pelias@noa.gr](mailto:pelias@noa.gr)

11           <sup>3</sup>Ecole Normale Supérieure de Paris, PSL Research University, Laboratoire de Géologie -  
12 UMR CNRS 8538, 24 rue Lhomond, 75005 Paris – France [briole@ens.fr](mailto:briole@ens.fr) ;  
13 [escartin@geologie.ens.fr](mailto:escartin@geologie.ens.fr)

14           <sup>4</sup>Department of Civil Engineering, Democritus University of Thrace, 67100 Xanthi, Greece  
15 [valkaniotis@yahoo.com](mailto:valkaniotis@yahoo.com)

16           <sup>5</sup>Department of Geology, University of Patras, 26504 Rio, Greece

17           \*           Correspondence: [aganas@noa.gr](mailto:aganas@noa.gr) Tel.: +30 210 3490186 (A.G.)

18  
19           **Abstract**

20           On 30 October 2020 a large  $M_w = 7.0$  earthquake occurred north of the island of Samos,  
21 Greece. Here we present the characteristics of the seismic fault (location, geometry,  
22 geodetic moment) as inferred from the processing of geodetic data (InSAR and GNSS). We  
23 use the InSAR displacement data from Sentinel-1 interferograms (ascending orbit 29 and  
24 descending 36) and the GNSS offsets from fourteen (14) stations in Greece and Turkey to  
25 invert for the fault parameters. Our inversion modeling indicates the activation of a normal  
26 fault offshore Samos with a length of 40 km, width of 15 km, average slip of 1.7 m, a  
27 moderate dip-angle ( $37^\circ$ ) and with a dip-direction towards North. The inferred fault is  
28 located immediately north of, and adjacent to Samos, with the top of the slip  $\sim 0.6$  km  
29 below surface, and  $\sim 1$  km off-shore at its closest to the island. Near the fault the earthquake  
30 caused the permanent uplift of the island up to 10 cm with the exception of a coastal strip  
31 along the NE part of the northern shore that subsided 2-6 cm. The co-seismic horizontal  
32 motion of GNSS station SAMO was 35.6 cm towards south and 3 cm towards west. A  
33 postseismic signal (22-33% of the co-seismic on the vertical component) was observed at  
34 GNSS stations SAMO and SAMU, with a time constant of 30 days. The effects of the  
35 earthquake included liquefaction, rock falls, rock slides, road cracks and deep-seated

36 landslides, all due to the strong ground motion and associated down-slope mobilization of  
37 soil cover and loose sediments.

38

39 **Keywords:** earthquake, InSAR, GNSS, liquefaction, uplift, Samos, Aegean

40

## 41 **1. Introduction - Tectonic Setting**

42 The Samos island is located in the eastern Aegean Sea, a well-known extensional terrain,  
43 as it is found on the back-arc area behind the Hellenic subduction (Fig. 1; McKenzie, 1978;  
44 Angelier, 1979; Taymaz et al. 1991; Ganas and Parsons, 2009; Meng et al. 2021). The plate  
45 kinematics are determined by the ongoing northward subduction of the African slab under  
46 the Aegean continental lithosphere and the synchronous, westwards motion of the  
47 Anatolian microplate. The GPS velocity field, with respect to Eurasia (fixed reference), is  
48 characterised by a counter-clockwise rotation of the continental plate, with increasing GPS  
49 velocities from eastern Anatolia towards the Hellenic trench (McClusky et al. 2000;  
50 Reilinger et al. 2010; Briole et al. 2021). Within this setting the central Aegean block  
51 (including Samos) moves almost uniformly southwest (S24°W) at an average velocity of ~30  
52 mm/yr (w.r.t to Europe; Briole et al. 2021; revising the 33 mm/yr value of Ganas and  
53 Parsons, 2009; Vernant et al. 2014). This velocity plateau marks the end of the significant  
54 increase in N-S extension of the whole Aegean as observed from North to South (Briole et  
55 al. 2021, their Fig. S29; see also McClusky et al. 2000; D'Agostino et al. 2020). Crustal  
56 extension is accommodated by seismic slip along normal-slip and strike-slip active faults,  
57 that are scattered throughout the north and central Aegean and western Anatolia (Masclé  
58 and Martin, 1990; Taymaz et al. 1991; Ganas et al. 2005; Chatzipetros et al. 2013; Tan et al.  
59 2014; Yolsal-Çevikbilen et al. 2014; Melis et al. 2020). The last seismic activity in the broader  
60 Samos area was the October 2005 seismic sequence in Seferihisar – Sigacik (Benetatos et  
61 al. 2006; Fig. 1) in the Turkish coastal area (North of Samos), involving strike-slip ruptures  
62 with moment magnitudes in the range 5.4-5.8.

63 The geology of Samos comprises metamorphic and carbonate rocks of Mesozoic and early  
64 Cenozoic age that bound two Neogene sedimentary basins (mostly composed of marls and  
65 sandstone formations; Theodoropoulos, 1979). The metamorphic complex of Samos is  
66 composed mainly of marbles and phyllites, serpentinites, blueschists and greenschists  
67 (Okrusch et al. 1984). The active tectonics is dominated by extensional kinematics along  
68 numerous faults (Boronkay and Doutsos, 1994; Ring et al. 1999) with fault-slip data  
69 indicating nearly N-S extension (Boronkay and Doutsos, 1994). A major, north-dipping  
70 normal fault occurs just north of the island (Fig. 1) and it tilted Samos Island to the south  
71 (Ring et al. 1999; Chatzipetros et al. 2013; Nomikou et al. 2021). Geological evidence  
72 (marine biogenic notches now found above mean sea level) suggests that a large part of

73 the northwestern coastline has been uplifted due to a combination of regional  
74 geodynamics and earthquake activity (Stiros et al. 2000; Evelpidou et al. 2021).

75 As the 30 October 2020 epicentre is located offshore Samos (Fig. 1) we investigated the  
76 bathymetric features (submarine scarps, ridges etc.) and other geophysical evidence for  
77 active faulting. The grids publicly available (gmrt <https://www.gmrt.org/>; emodnet  
78 <https://emodnet.eu>) show that overall depths of the basin north of Samos mark a  
79 deepening from ~500 m to 1000 m towards the west. Early seismic profiles in this area of  
80 the Aegean are limited to the Ikaria area, and reveal a half-graben offshore structure  
81 (Mascle and Martin, 1990). Immediately following the Samos earthquake, a bathymetry  
82 survey of the area north of Samos and around the earthquake epicentre revealed ~E-W  
83 seafloor structures consistent with the surface expression of a normal fault scarp, showing  
84 seafloor slopes of 30-50°, and bounding the Samos Basin that is interpreted as a half-graben  
85 (Nomikou et al., 2021). These data are consistent with a normal faulting event nucleating  
86 in the prolongation towards the north of this fault's trace (Nomikou et al., 2021), and also  
87 reveal numerous mass wasting scarps and deposits throughout, that may be linked to the  
88 long-term activity of this fault system. Around Samos – Kuşadası (Fig. 1) Tan et al (2014)  
89 showed that the seismicity is heterogeneously distributed including an activity offshore  
90 Samos in the vicinity of the epicentre of the 2020 Samos earthquake, and that there is  
91 microseismic activity onshore Samos as well as onshore activity in the Kuşadası area.

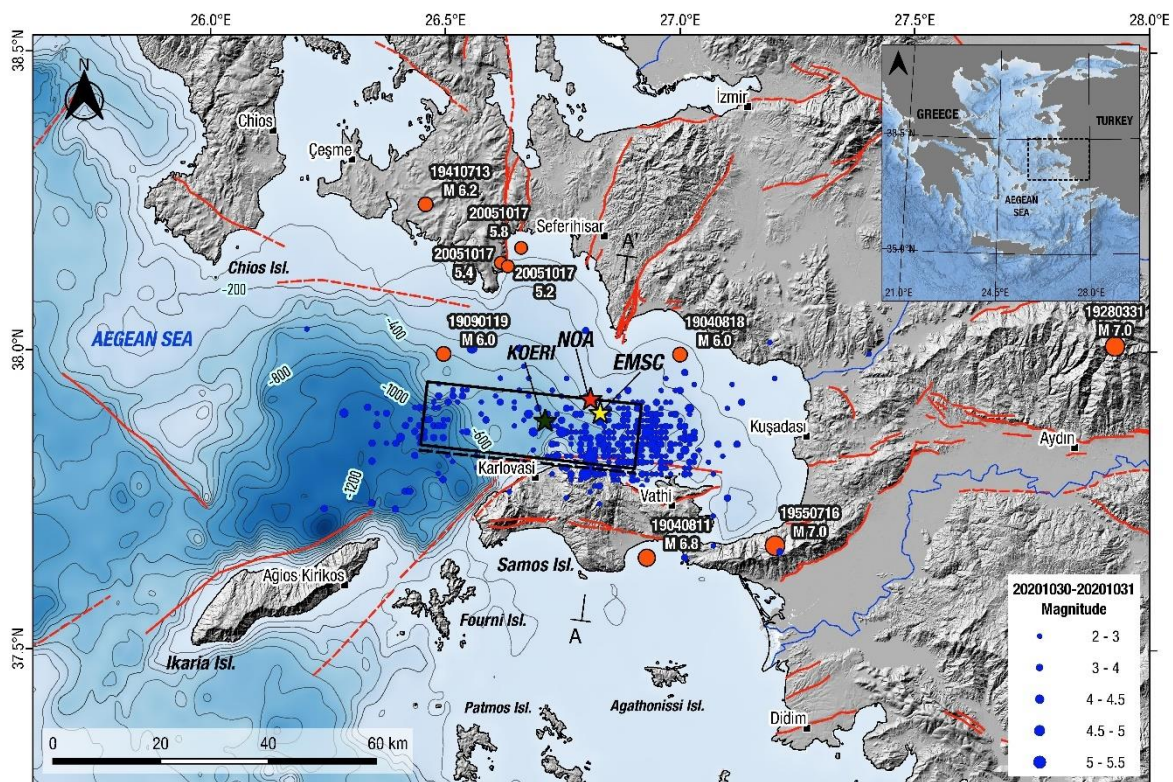
92 In terms of finite strain, the amount of crustal extension between Samos and the Menderes  
93 region of western Anatolia (including the broader Izmir area) ranges from 5.3-7.4 mm/yr  
94 according to Vernant et al. (2014; with 0.8-1.0 mm/yr uncertainty) based on GNSS data  
95 modeling. New GNSS data provide an extension rate of  $6.2 \pm 0.7$  mm/yr between eastern  
96 Samos and Izmir (Fig. 1; Briole et al. 2021). A component of left-lateral slip was also  
97 proposed by Vernant et al. (2014; 4.8-5.0 mm/yr) along with their proposed WNW-ESE  
98 crustal boundary offshore Samos.

99 Previous seismicity around Samos includes a cluster of poorly-located  $M=6.0+$  events in  
100 1904 (Ambraseys, 2001; Fig. 1). A  $M_s=6.1$  (surface magnitude) event occurred on 11 August  
101 1904 (this event is reported as  $M=6.8$  in the online catalog of AUTH's seismological  
102 laboratory) south of Samos, followed by a  $M_s=6.0$  event on 18 August 1904; a third event  
103 with  $M_s=6.0$  occurred on 10 October 1904. All those events caused severe damage to  
104 Samos and settlements along the Turkish coast. Eight (8)  $M \geq 6.0$  size events are reported  
105 by Tan et al. (2014) to have affected Samos during the 19<sup>th</sup> century. A  $M=7.0$  event occurred  
106 in 1955 to the southeast of Samos, probably on the normal fault bounding the northern  
107 side of the Büyük Menderes graben (Fig. 1). During October 2005 three moderate events  
108 ( $5.4 \leq M \leq 5.8$ ) occurred west of Seferihisar, Turkey (Fig. 1) about 40 km north of Samos  
109 (Benetatos et al. 2006; Papadopoulos et al. 2006).

110 In this paper we use geodetic observations (GNSS and InSAR) to model the seismic fault  
111 (sections 3, 4 and 5). The fringes in the interferograms onshore Samos show a pattern that

112 supports more the north-dipping fault plane. Even if the InSAR data cover only a portion of  
 113 the deformation field, it allows us to distinguish between the two patterns (north-dip vs.  
 114 south dip) in terms of gradient, density, curvature and the concentration patches of the  
 115 fringes. We also present field observations (section 6) including liquefaction phenomena  
 116 and InSAR data suggesting massive gravitational movements and discuss our findings in  
 117 terms of footwall deformation patterns (section 7). We show that both InSAR and GNSS  
 118 data provided evidence for co-seismic and postseismic deformation (uplift) of northern  
 119 Samos. We finally discuss the significance of our findings for the active tectonics and strain  
 120 in this part of the Aegean.

121



122

123 **Fig. 1.** Shaded relief map of the 30 October 2020 Samos  $M_w=7.0$  earthquake area showing  
 124 location of the seismic fault determined in this study (black rectangle). Solid stars indicate  
 125 mainshock epicentres from the National Observatory of Athens (NOA), Kandilli Observatory  
 126 and Earthquake Institute (KOERI) and the European-Mediterranean Seismological Center  
 127 (EMSC). Blue dots indicate the first 24-hrs aftershocks from the EMSC catalogue (N=484).  
 128 Initials A-A' denote the end points of the cross section in Fig. 17. Inset box at upper right  
 129 shows study area within Greece. Red lines are active faults from the NOAFAULTs database  
 130 <http://doi.org/10.5281/zenodo.3483136>

131



132 **2. The Samos 30 October 2020 earthquake**

133 On 30 October 2020 11:51 UTC a strong and shallow earthquake rocked the eastern Aegean  
 134 Sea (ITSAK, 2020; Kalogeras et al. 2020; Ganas et al. 2020a; Lekkas et al. 2020; Fomelis et  
 135 al. 2021; Mavroulis et al. 2021). The epicentre was determined offshore the Greek island  
 136 of Samos, about 260 km to the East of Athens (Fig. 1). The local magnitude (Richter scale)  
 137 was measured by NOA as  $M_L=6.7$ , while the moment magnitude was calculated by various  
 138 agencies in the range  $6.9 \leq M_w \leq 7.0$  (see Table 1; Papadimitriou et al. 2020; Kalogeras et  
 139 al. 2020; Ganas et al. 2020a). The Peak ground acceleration at Vathy, Samos (Fig. 1) reached  
 140 0.17 g (Kalogeras et al. 2020). The earthquake triggered a tsunami that caused minor  
 141 damage at the surrounding islands (Samos, Chios, Ikaria; Fig. 1) and especially in the coastal  
 142 towns of Vathy (Samos) and Sigacik (Turkey; Kalogeras et al. 2020; Triantafyllou et al. 2021).  
 143 Immediately after the earthquake field surveys showed coastal uplift at several locations  
 144 of Samos (Lekkas et al. 2020; Evelpidou et al. 2021; Mavroulis et al. 2021), however, in this  
 145 study we provide GNSS and InSAR evidence that there was no uplift along the south coast.  
 146 The focal mechanisms reported by various agencies for the mainshock are listed in Table 1.

147

148 **Table 1.** Focal mechanism solutions reported by various agencies (source: agency websites;  
 149 see supplementary Text S1 for web sources).

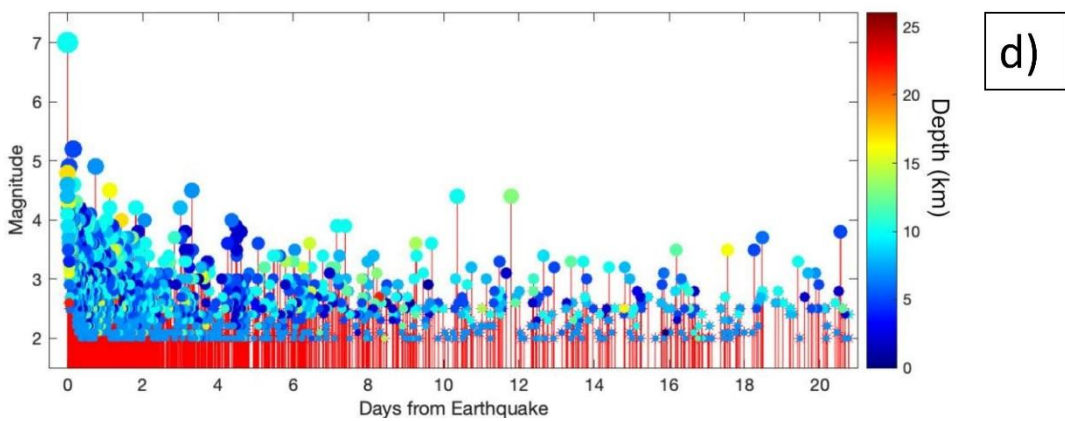
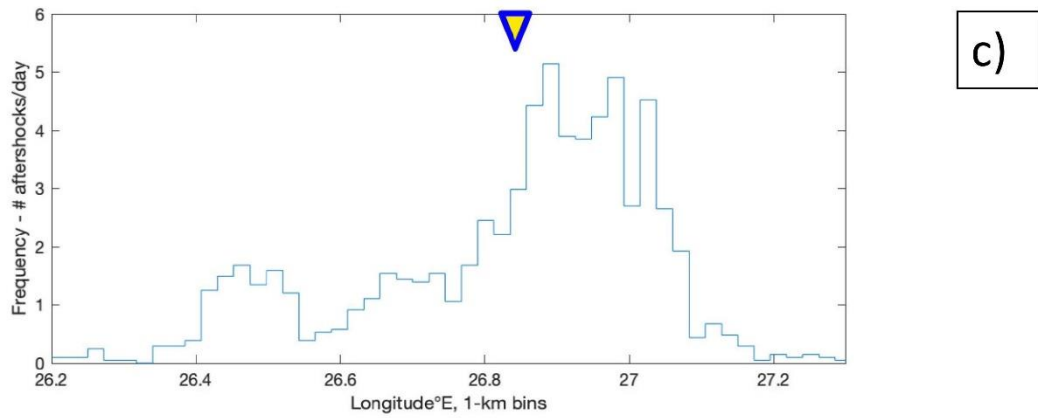
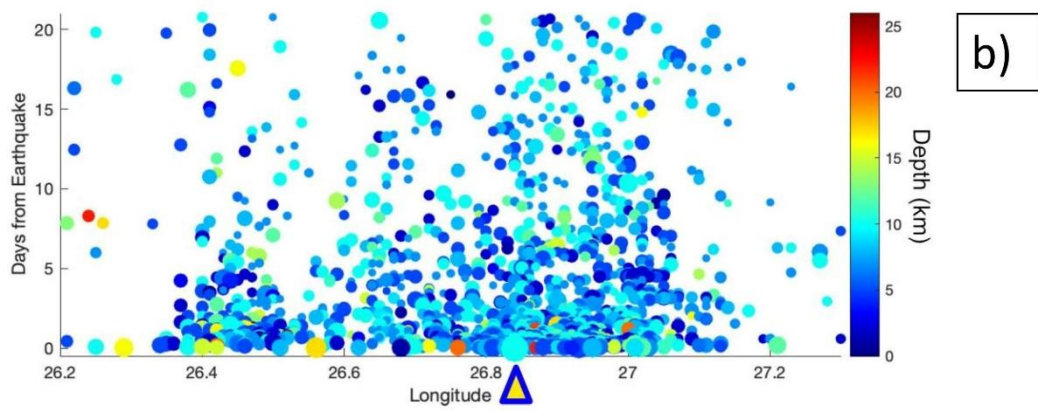
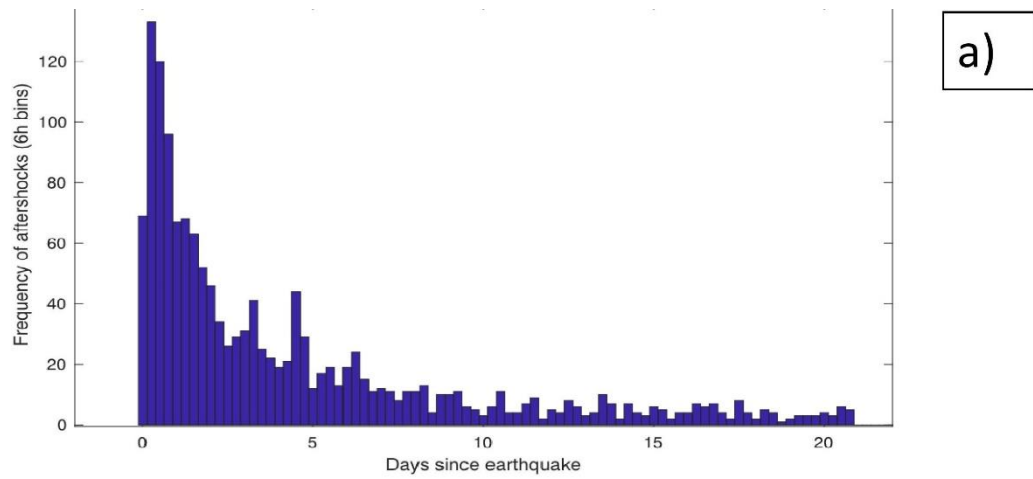
Agency	Time	Lat. (°)	Long. (°)	Depth (Km)	Mag Mw	Strike 1	Dip 1	Rake 1	Strike 2	Dip 2	Rake 2	M <sub>0</sub>
AFAD- ERD	11:51:23	37.888	26.777	16.54	6.6	270	46	-91	95	43	-87	3.26E+19
GFZ	11:51:27	37.900	26.820	15	7	272	48	-93	97	41	-85	3.50E+19
GCMT	11:51:34	37.760	26.680	12	7	270	37	-95	96	53	-86	4.01E+19
INGV	11:51:36	37.836	26.808	10.6	7	289	40	-69	82	53	-107	3.90E+19
IPGP	11:51:26	37.897	26.795	14	7	260	36	-116	111	58	-72	3.76E+19
KOERI- RETMC	11:51:24	37.902	26.794	13.1	6.9	272	55	-93	97	84	-85	3.00E+19
UOA	11:51:26	37.920	26.803	13	6.9	270	50	-81	76	41	-101	2.81E+19
NOA	11:51:24	37.900	26.817	6	6.9	294	54	-65	76	43	-120	2.65E+19
OCA	11:51:26	37.910	26.840	10	7.2	275	45	-96	103	45	-85	-
USGS	11:51:44	37.918	26.790	11.5	7	276	29	-88	93	61	-91	4.09E+19

150

151 The effects of the earthquake were devastating in Greece and Turkey (Vadaloukas et al.  
 152 2020; Cetin et al. 2020). In Greece, two young children lost their lives from a wall collapse  
 153 in the town of Vathy, eastern Samos ([https://edition.cnn.com/europe/live-news/aegean-  
 154 sea-earthquake-dle-intl/h\\_9d313a45fc3463479a6ff1ec4da29755](https://edition.cnn.com/europe/live-news/aegean-sea-earthquake-dle-intl/h_9d313a45fc3463479a6ff1ec4da29755)). In Turkey more than  
 155 100 people died from multiple building collapses in the city of Izmir, ~60-km to the north  
 156 of the epicentre (Fig. 1; [https://en.wikipedia.org/wiki/2020\\_Aegean\\_Sea\\_earthquake](https://en.wikipedia.org/wiki/2020_Aegean_Sea_earthquake)).

157 The Samos earthquake ranged in magnitude  $6.6 \leq M_w \leq 7.2$  (see Table 1 for estimates;  
 158 excluding AFAD-ERD it becomes  $6.9 \leq M_w \leq 7.2$ ) and it occurred along a normal fault oriented

159 E-W as indicated by the moment tensor solutions of both regional and teleseismic data.  
160 Until 20 November 2020 06:21 UTC more than 1445 aftershocks (with  $2.0 \leq M_L \leq 5.2$ ) were  
161 recorded by the European-Mediterranean Seismological Centre (EMSC; Fig. 2). The  
162 mainshock was followed by one moderate-size aftershock, on 30 October 2020 15:14 UTC  
163 and with magnitude  $M_w=5.2$ . The moment tensor solutions of the mainshock indicate E-W  
164 to ESE-WNW normal faulting in agreement with the regional, extensional tectonics (Table  
165 1; Ring et al. 1999; Chatzipetros et al. 2013). The aftershock sequence extends over a E-W  
166 distance of  $\sim 70$ -km with most of events occurring east of the mainshock and within a depth  
167 range 2-15 km (Fig. 2; see Fig. S1 for a map view). The EMSC aftershock data show an  
168 incipient decay the day after the mainshock (Fig. 2a) while the longitude location vs time  
169 plot (in days from main shock) shows that the sequence is spatially constrained between  
170  $26.4^\circ\text{E} - 27.2^\circ\text{E}$  (Fig. 2b); this plot also shows that the sequence is unevenly distributed  
171 along strike. Indeed, the data show a peak in seismicity between  $26.8^\circ$  and  $27.1^\circ\text{E}$ , a  
172 secondary peak immediately to the west and centered at  $26.7^\circ\text{E}$ , and a third peak centered  
173 at  $26.45^\circ\text{E}$  with a clear seismic gap  $\sim 26.55^\circ\text{E}$  (Fig. 2c). This uneven and segmented  
174 distribution of the aftershocks is established immediately at the start of the aftershock  
175 sequence (see aftershock distribution in Fig. 1), and persists over time, even during the  
176 most recent periods of lower seismic activity (Fig. 2d; a similar feature is reported by  
177 Papadimitriou et al. 2020). This seismicity gap may be due to the complete rupture of the  
178 main asperity along the fault plane and subsequent aftershock activity triggering around  
179 the co-seismic slip patches (e.g., Atzori et al. 2008; Chiaraluce et al. 2011; Ganas et al. 2019;  
180 Melgar et al. 2017, 2020). The E-W extent of the early (24-hr) aftershocks (Fig. 1) is about  
181 55 km, or 1.4 times the size of the modelled fault in this paper (40 km), close to the mean  
182 value for the aftershock zone ratio in California (Neo et al. 2020).





184 **Fig. 2.** Spatiotemporal evolution of the Samos aftershock sequence (source of data: EMSC,  
185 data until 20 November 2020 N=1445) a) histogram of aftershock frequency of occurrence  
186 vs time (6h bins) to show the decay over time b) Daily frequency of occurrence with respect  
187 to longitude (east-west) and depth (circle colours correspond to depth, and size to the  
188 magnitude) c) histogram of aftershocks in frequency per day along strike (east-west; 1 km  
189 bins) d) Magnitude distribution with time and depth. Triangles (yellow with blue outline) in  
190 b and c indicate the location of the mainshock for reference.

191

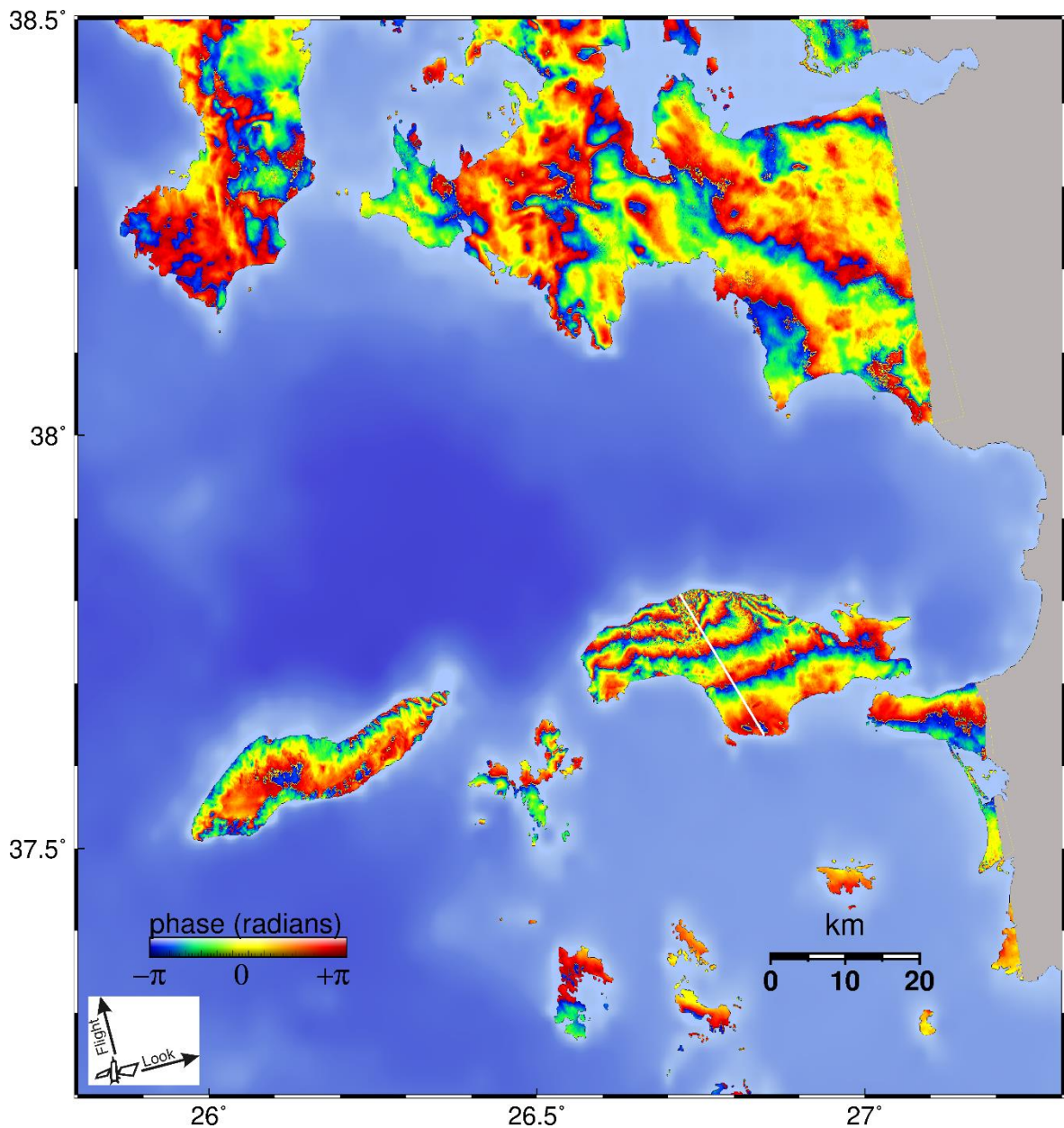
### 192 **3. InSAR data processing**

193 We used Synthetic Aperture Radar (SAR) C-band images acquired over the Samos area by  
194 the Copernicus European satellites Sentinel-1 from the ascending tracks 29 and 131 and  
195 the descending track 36. The first co-seismic, differential interferograms spanning the  
196 dates 24-30 October 2020, and 18-30 October, 2020 with images of track 131 (Fig. S2) were  
197 calculated on the Geohazards Exploitation Platform (<https://geohazards-tep.eu>) using the  
198 SNAP v5.0 software and were presented in Ganas et al (2020a). The acquisition time of the  
199 postseismic image was at 30 October 2020, 16:07 UTC, thus 4 hours & 16 minutes after the  
200 mainshock. Then, we produced 17, 15, 17 co-seismic and 10, 6, 6 post-seismic  
201 interferograms for each of tracks 131 (ascending), 29 (ascending) and 36 (descending)  
202 respectively, using the Automatic Interferometric Processing Station (AIPS; [http://aips-  
203 p.space.noa.gr/Samos](http://aips-p.space.noa.gr/Samos)) at NOA. More information on AIPS can be found in text S2. The east,  
204 north and up unit vectors, from the ground to the satellites, at the range direction (in the  
205 centre of Samos island), are  $[-0.680 -0.157 0.716]$  for track 29,  $[0.629 -0.145 0.764]$  for track  
206 36 and  $[-0.540 -0.125 0.832]$  for track 131. The interferograms captured the ground motion  
207 along the line of sight (LOS; see Look arrows in Fig. S2). These LOS measurements of were  
208 used in the inversion model to retrieve the parameters of the seismic fault.

209 After reviewing all the co-seismic interferograms we decided not to use the track 131 in  
210 this paper, because of the better quality of track 29, which is ascending, too. In Fig. 3 we  
211 present the best co-seismic interferogram belonging to track 29 and spanning the period  
212 23 October – 10 November 2020. Moreover, we selected the four (4) best interferograms  
213 (marked in bold and framed in Table S1) in terms of low tropospheric noise and high  
214 coherence of each of the track 29 and 36 (Fig. 4a, 4d). We then unwrapped the  
215 interferograms, removing the bias by setting each one of them to a zero mean value. Then,  
216 for both the ascending and descending tracks, we averaged the LOS displacement grids  
217 these two averaged grids were used for the inversion modelling in section 5. All the  
218 following interferograms for which their extent is limited only to the Samos island, are at  
219 the notion of post-event minus pre-event, thus phase increasing towards satellite. Fig. S3  
220 demonstrates the increase of signal to noise ratio obtained when using the average  
221 interferogram over the best interferogram of track 36 (D), both wrapped and unwrapped.

222

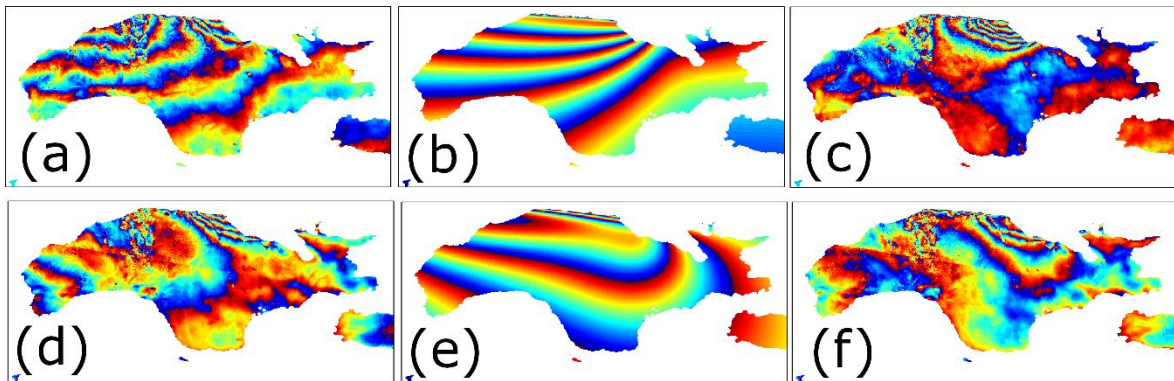
223



224

225 **Fig. 3:** The best co-seismic differential interferogram calculated with images acquired from  
226 ascending track 29 and spanning the period 23 October 2020 – 10 November 2020  
227 (produced by AIPS; grey colour indicates no data). The formation of the interferograms is  
228 at the notion of post-event minus pre-event image, thus phase decreasing towards  
229 satellite. The fringes in the interferogram onshore Samos show a pattern that supports  
230 more the north-dipping fault plane as such concentration requires the vicinity of a fault  
231 edge to them. The InSAR data show that most of Samos moved roughly upwards with the  
232 exception of a coastal strip (from Agios Konstantinos to Kokkari; see Fig. 9) that moved  
233 roughly downwards. Thin white line on Samos indicates location of profile shown in Fig. S9.

234



235

236 **Fig. 4:** (a) and (d): Averaged wrapped interferograms for tracks 29 (A) and 36 (D), averaged,  
237 (b) and (e) the modelled interferograms of track 29 and 36 and (c) and (f) the residuals.

238

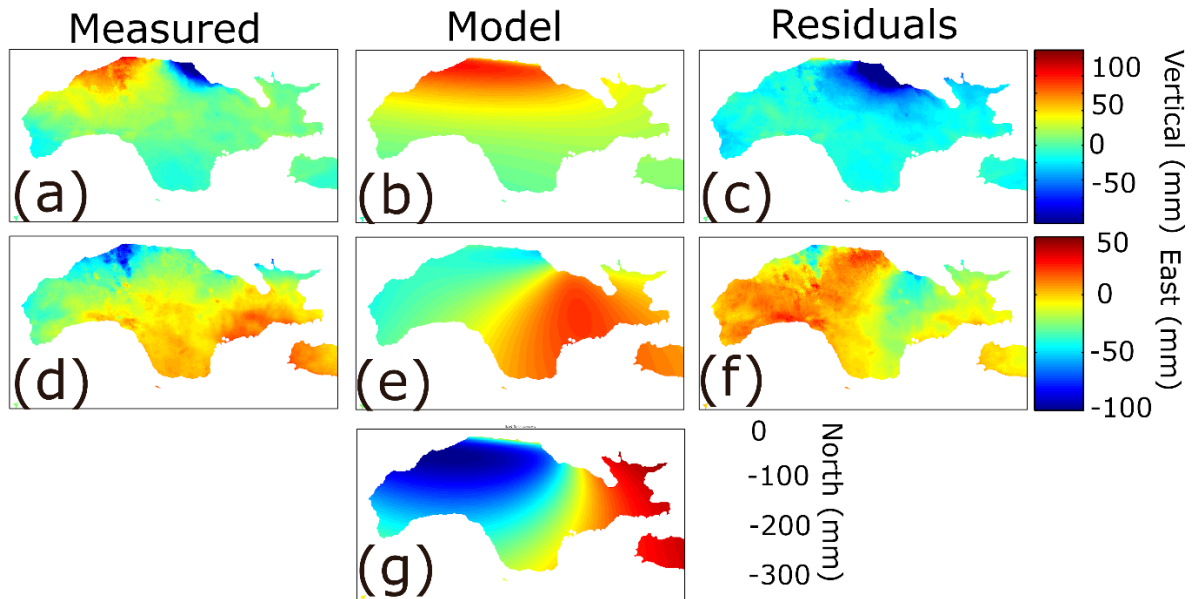
239 The interferograms of Fig. 3, Fig. 4, & Fig. S2 show several fringes corresponding to ground  
240 deformation onshore Samos. The pattern of deformation shows two main features: an area  
241 along the north coast with three fringes oriented ENE-WSW indicating motion away from  
242 the satellite (region between Agios Konstantinos and Kokkari; see Fig. 9 for exact locations)  
243 and another, larger region mostly in the central and west part of the island with a motion  
244 towards the satellite. The absolute value of the interferometric fringes is estimated by  
245 anchoring to the LOS the offsets measured at GNSS station SAMU in Vathy (Fig. 1), that is  
246 +33 mm for the track 29 and +23 mm for track 36. Considering that the sense of the  
247 deformation is the same in both tracks we can infer that the north-northeastern tip of  
248 Samos island shows subsidence and the western part of Samos shows an uplift.  
249 Accordingly, we proceeded to the extraction of the vertical and east components of the  
250 deformation. As InSAR provides only information on two independent LOS measurements  
251 (Fig. 4), in order to retrieve the 3-D motion we use as third observation the north  
252 component derived from the model that is presented in the next section. The north-south  
253 displacements produced by this earthquake are large in the northern Samos, but the weight  
254 of the north-south component to the calculation of the LOS displacement is small (between  
255 10-15%). Therefore, our model is sufficient to provide a first order estimate that is accurate  
256 enough for constituting the third source of displacement. The vertical and east components  
257 of the deformation are shown in Fig. 5 (a and d, respectively). The fact that we can retrieve  
258 an accurate measurement of the vertical displacements is particularly important for  
259 comparison with the field observations along the shoreline (see section 6) and further  
260 geological interpretations.

261 We interpret the deformation of the central and west part of the island as a result of co-  
262 seismic motion along an offshore normal fault, running E-W and dipping to the North. As  
263 the rupture is shallow, and the dip angle (from seismology) is significantly different from  
264 45°, the two conjugate planes produce modelled differential interferograms that are not

265 similar to each other. Therefore, InSAR has the capability to discriminate between the two  
 266 conjugate planes. Our joint inversion (based on InSAR & GNSS displacement data) approach  
 267 has resulted in a robust fault model, providing its location, geometry (strike/dip-angle/rake  
 268 angles and concluding on its dip-direction) and amount of finite slip while assuming a  
 269 simple and planar fault geometry.

270

271



272

273 **Fig. 5:** (a) and (d): The vertical and East components obtained from the averaging of the  
 274 best interferograms, (b), (e) and (g) the modelled vertical, East and North components and  
 275 (c) and (f) the residuals. Units in mm.

276

#### 277 4. GNSS data processing

278 Dual-frequency data from fourteen (14) permanent (continuously recording) Global  
 279 Navigation Satellite System (GNSS) stations were processed. We used GPS observations  
 280 covering approximately 7-10 days before and after the earthquake (23 October 2020 to 10  
 281 November 2020; Table 2 and Fig. 6) in order to obtain position time series and their offsets  
 282 due to the earthquake. The sampling interval was 30 s, and the data comprise 24-hr daily  
 283 files. All station records were complete (rejected epochs 0.00%) thus providing substantial  
 284 observations for mapping the co-seismic displacement field and were included in our  
 285 processing. The stations belong to the Ktimatologio SA (HEPOS), HxGN SmartNET and  
 286 Uranus networks of Greece and the Turkish National Permanent GNSS/RTK Network. The  
 287 azimuthal distribution of the GNSS stations is very good, extending uniformly around the  
 288 epicentral area (Fig. 6). However, most of the GNSS stations are located more than 50 km  
 289 away from the epicentre and therefore captured small displacements of the order of a few



290 cm. The daily coordinates are calculated in ITRF14 (epoch 2020.049) and converted to  
 291 Universal Transverse Mercator UTM (North) zone 35. Their uncertainties were converted  
 292 onto the local geodetic frame, based on the given Cartesian (ECEF; earth-centered, earth-  
 293 fixed) sigma (95%) ones.

294 The processing was made with the Precise Point Positioning (PPP) strategy by means of the  
 295 GIPSY/OASIS II software (version 6.4) developed by the Jet Propulsion Laboratory (JPL;  
 296 <http://gipsy-oasis.jpl.nasa.gov>). We used the JPL final satellite orbits (flinnR) and clocks,  
 297 absolute antenna calibration, random walk troposphere estimation, and the FES2004  
 298 ocean loading model. We calculated the static offsets (and their uncertainties) for the 30  
 299 October 2020 event (Table 2 & Table S2; Fig. 7 and Fig. S4a & S4b). The offsets indicate cm  
 300 size motion mainly along the N-S direction. Station SAMO (Fig. 6) registered the largest  
 301 displacements as it moved 35.6 cm towards south and 3 cm towards west. In terms of Up-  
 302 Down (vertical) displacement we obtained clear trends of co-seismic motion for the three  
 303 stations located onshore Samos (SAMO, 093A & SAMU) with amplitudes +8.2, +2.3 and +2.1  
 304 cm, respectively.. The vertical displacements at the GNSS stations enhance the robustness  
 305 of the north-dipping fault model as they allow to discriminate between the two conjugate  
 306 focal planes (see also synthetic models of Okada-type dislocations; Okada, 1985) in Fig. S5.  
 307 Regarding the large aftershock on 30 Oct. 2020 15:14 UTC (M=5.2; located 9 km to the N  
 308 of Samos, NOA solution), we could not find any evidence of this aftershock in the GNSS  
 309 time series. Furthermore, the GNSS time series of SAMU and SAMO stations show the  
 310 existence of a small post-seismic motion in the days following the mainshock (Fig. 7). This  
 311 postseismic motion leads to drifts of LOS values in our interferograms that we can estimate.  
 312 The values of those drifts is low so we can assess that the interferograms presented in Fig.  
 313 2, 3 & 4 are not biased by post-seismic motions and represent well the co-seismic signal of  
 314 the earthquake.

315

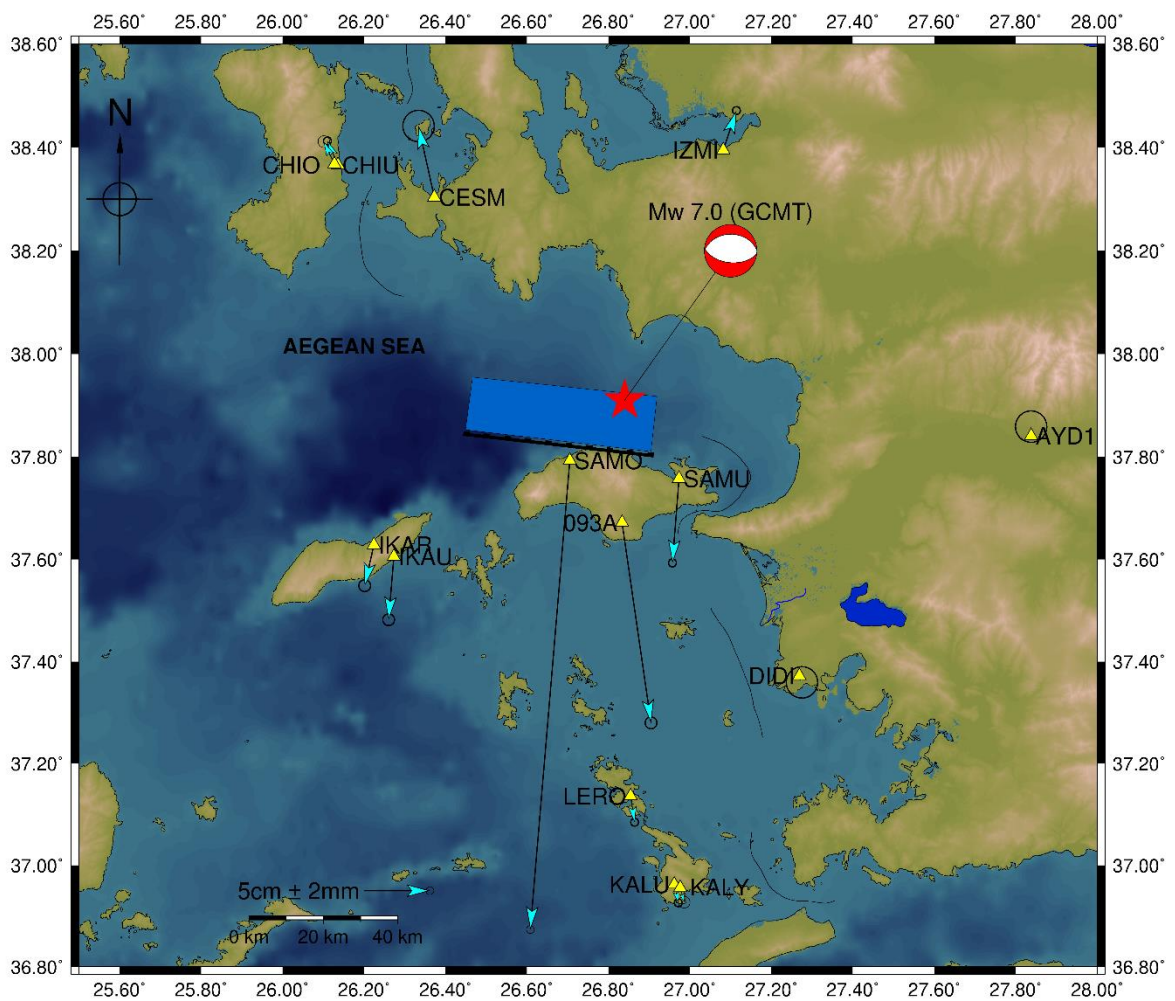
316 **Table 2.** List of co-seismic offsets (in mm) determined by PPP processing. Columns dE, dN,  
 317 dUp show displacements (East, North, Up) and mE, mN, mU the modelled values,  
 318 respectively. The uncertainties of E, N, Up displacements are reported in Table S2. The  
 319 GNSS stations are shown in Fig. 6.

Station	Lat. (°)	Long. (°)	dE	dN	dU	mE	mN	mU
SAMO	37.7927	26.7053	-30	-356	82	-38	-348	75
SAMU	37.7575	26.9734	-5	-64	21	4	-57	30
093A	37.6716	26.8334	22	-152	23	11	-167	8
CHIO	38.3679	26.1272	-8	17	2	-7	19	2
CHIU	38.3665	26.1359	-8	18	5	-7	19	2
IKAR	37.6282	26.2243	-7	-31	7	-18	-34	4
IKAU	37.6055	26.2733	-4	-48	4	-21	-43	3

KALU	36.9624	26.9617	3	-14	-3	1	-15	-2
LERO	37.1364	26.8545	3	-20	-	1	-24	-
KALY	36.9558	26.9762	3	-11	-1	1	-15	-2
AYD1	37.8407	27.8378	0	7	-	0	1	-
CESM	38.3038	26.3726	-12	54	-	-12	48	-
IZMI	38.3948	27.0819	10	30	2	13	36	-2
DIDI	37.3721	27.2687	2	-5	-	5	-15	-

320

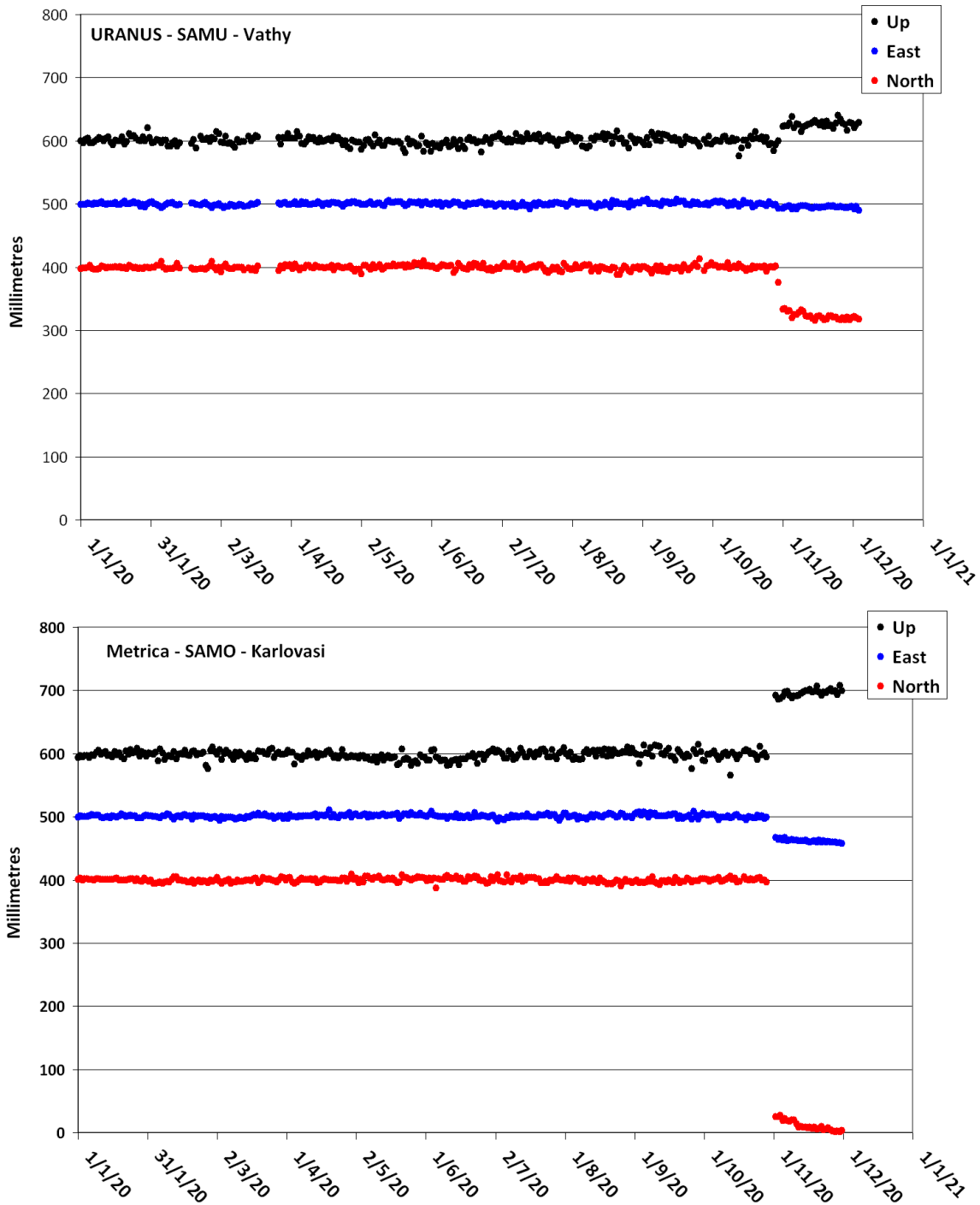
321



322

323 **Fig. 6:** GNSS co-seismic displacement map showing also topography/bathymetry, the focal  
 324 mechanism (beachball; compressional quadrants in red) and the epicentre of the Samos  
 325 October 30, 2019 earthquake. Blue rectangle indicates the surface projection of the seismic  
 326 fault (thick black line is the upper edge with boxes on the downthrown side). Triangles  
 327 indicate permanent GNSS station locations. Vectors indicate horizontal displacement of  
 328 GNSS stations with 67% confidence scaling.





329

330 **Fig. 7.** Position time-series of stations SAMU (Vathy) and SAMO (Karlovasi) (E, N, Up in blue,  
 331 red, black dots, respectively). The time-series have been offset for clarity on the Y-axis. The  
 332 small post-seismic signal is visible mainly in the North-South component. Station locations  
 333 are shown in Fig. 6.

334

335

## 336 **5. Joint inversion of geodetic data – fault model**

337 We inverted the geodetic data (InSAR and GNSS) assuming a half-space elastic model with  
338 uniform slip along a rectangular fault surface, using the code *inverse6* (Brìole, 2017). For  
339 the modelling of the InSAR data, the same weight was given to the ascending and  
340 descending picks. The ascending and descending observations (fringe pickings), here 64 for  
341 the ascending track 29 and 33 for the descending track 36, thus the weight ratio descending  
342 over ascending is 5/10. The 99 data points are compared against the model in Table S3 (see  
343 Fig. S6 for location of picked fringes). We then used the 14 sets of GNSS offsets (horizontal  
344 and vertical; Table 2) weighting 10 times more the horizontal offsets than the vertical ones  
345 and the InSAR data. The root mean square (r.m.s) scatter between data and model is 6.4  
346 mm for the GPS horizontal data, 9.8 mm for the GPS vertical data and between 23 - 59 mm  
347 for the InSAR data (see Table 4 for model performance), respectively. We tuned the model  
348 in favor of GNSS horizontal with the ratio 6:1 because of the higher reliability of the  
349 horizontal GNSS displacements and their broad coverage, while InSAR is confined inland  
350 Samos, which is, in particular, the reason for the weak constraint on the fault length from  
351 InSAR only.

352 The geodetic observations presented above imply clearly and unambiguously that the  
353 seismic fault plane is the one dipping towards north (Fig. 8). Indeed, there is no solution  
354 with the antithetic plane that can fit correctly both GNSS and InSAR at the same time.  
355 Moreover, the north-dipping plane is consistent with the submarine geomorphology (e.g.  
356 Mascle and Martin, 1990; Nomikou et al. 2021) and the tectonic context of the area.. At a  
357 local scale there are two to three fringes showing motion away from the satellite in the  
358 coastal area between Agios Konstantinos and Kokkari; (Fig. 4; see Fig. 9 for locations) that  
359 cannot be modelled with our simple model of uniform slip on a single fault plane. We  
360 suggest that they were formed because of a “local effect” due to the complexity and  
361 possibly the non-planarity of the mainshock rupture near its eastern termination. We also  
362 considered the case of a “deep-seated” gravitational instability that was induced due to  
363 strong ground shaking but the size of the feature (nearly 6-km E-W) and the shape of the  
364 fringes seem to rule it out. Moreover, there are no clearly defined boundaries of the local  
365 deformation visible in the interferograms, and the coastal area covered by this pattern  
366 changes in different orbit interferograms, thus hinting at a tectonic source.

367 We started the inversion for the East, North, Up coordinates of the fault-top centre and for  
368 the fault length. We did not invert for the dip, and rake angles, assuming 37°, and -90°  
369 respectively, i.e., using the GCMT solution which is consistent ( $\pm 2^\circ$ ) to our geodetic  
370 solutions for previous events in the Aegean Sea (i.e. Kos, 2017; Zakynthos, 2018; Durres,  
371 2019; Ganas et al. 2019, 2020b, 2020c). The 37° dip-angle is close to the average (39°) of  
372 the teleseismic determinations of the focal mechanism parameters (see Table 1; also, the  
373 37° angle fits the hypocentre determinations from regional networks, see Fig. 17 below).

374 The InSAR data alone do not allow to invert for the fault azimuth because of their  
375 insufficient spatial coverage but thanks to the existence of the GNSS data we inverted for  
376 the azimuth, too. We have not tried to constrain the rake of the slip vector, as we know  
377 that in the modelling of InSAR (and also GNSS), there is a strong trade-off between rake  
378 and azimuth angles. Therefore, we considered that it was more rigorous to rely on the  
379 angles provided by seismology through the focal mechanisms; rake angle between  $[-116^\circ /$   
380  $-65^\circ]$ , median  $-92^\circ$ , and azimuth  $[260^\circ / 294^\circ]$ , median  $272^\circ$ . Considering the fault as pure  
381 normal (rake  $-90^\circ$ ) we inverted only for the azimuth (strike) with small relaxation. In Fig. S7  
382 we see that it is inverted to  $276^\circ \pm 1^\circ$ . Indeed, there is a strong agreement between the  
383 azimuth of the modelled fault and orientation of the known active structures (see active  
384 faults in Fig. 1).

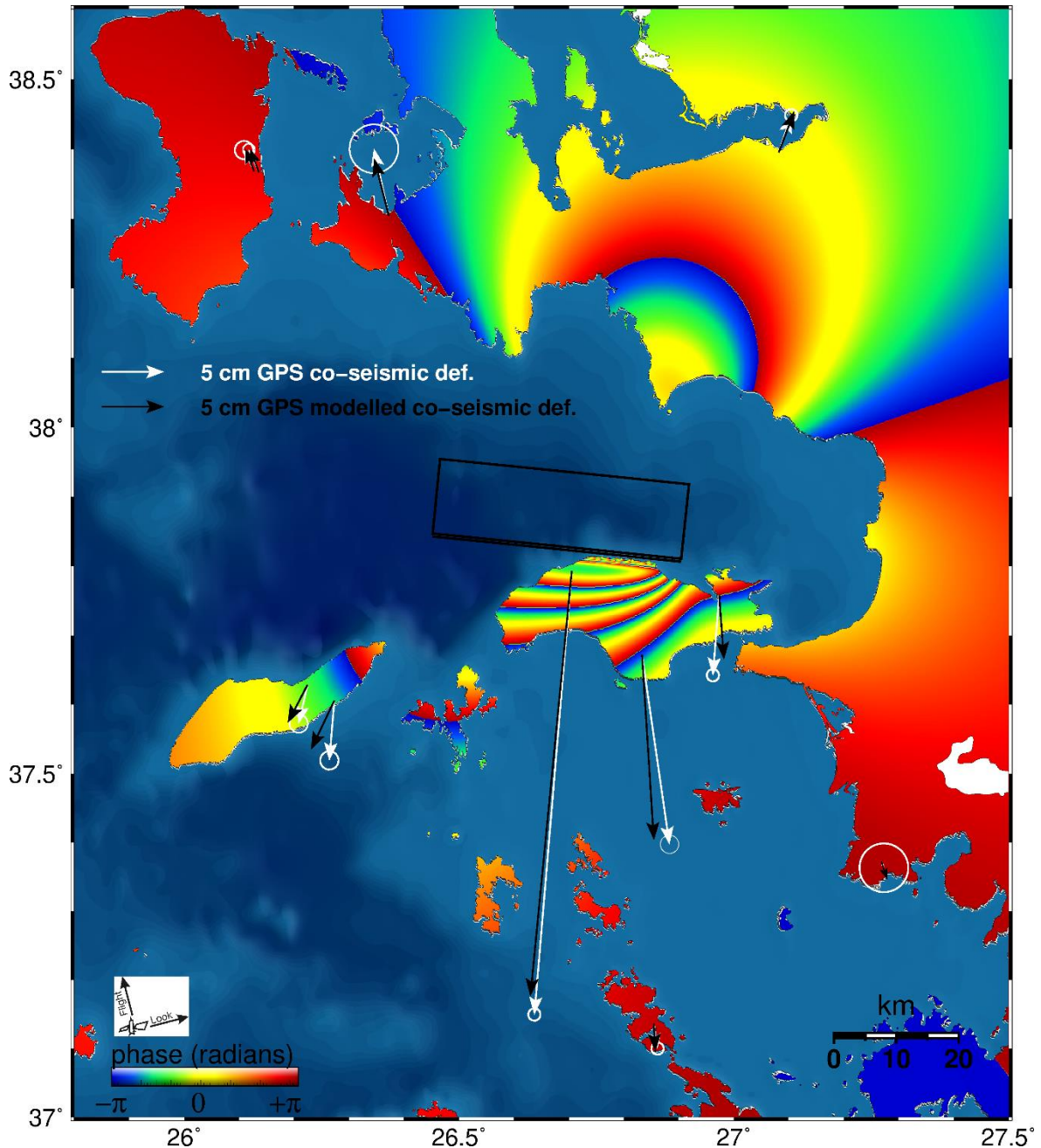
385 Concerning the location of the fault, we obtained robust estimates on the 3D position of  
386 the top-fault centre (uncertainty of 0.7 km, 0.5 km and 0.3 km in east, north, vertical,  
387 respectively), i.e., the horizontal coordinates are well constrained by InSAR displacement  
388 data and the vertical coordinates are also well constrained by the gradient of the InSAR  
389 fringes. The best-fitting length was  $40 \pm 3$  km (Fig. S7). Like for azimuth and rake, there is  
390 also a trade-off between the amount of slip and the fault width. The best-fitting depth to  
391 top-fault was  $0.6 \pm 0.3$  km and the fault width  $15 \pm 3$  km. Non-planar fault solutions were not  
392 tested as the GNSS and InSAR data do not have sufficient coverage in the near field of this  
393 offshore fault..

394 The fault parameters were used to construct a synthetic interferogram which is shown in  
395 Fig. 8 (ascending orbit) and Fig. S8 (descending orbit), respectively (also shown are the fault  
396 projection to the surface and the measured and modelled horizontal co-seismic GPS  
397 vectors). The modelled GNSS displacements agree well with the observed GPS data (Table  
398 2, Fig 6, Fig. 8, Fig. S8). We also present the residuals of the InSAR vertical and east  
399 components and the north component model In Fig. 5c and Fig. 5e, respectively. As  
400 mentioned earlier the negative residuals in Fig. 5e are due to the unmodelled "rupture  
401 complexity" effect along part of the northern coastline (Agios Konstandinos to Vourliotes;  
402 Fig. 9). We also drew a 21-km long NNW-SSE displacement profile (Fig. S9; the profile  
403 location is shown in Fig. 3) over central Samos in order to compare LOS InSAR vs modelled  
404 values and we obtained an average difference of 5.5 mm between the two sets (the  
405 standard deviation is 8 mm). In terms of earthquake size, the geodetic moment ( $M_0$ ) is  
406 calculated to  $3.06 \cdot 10^{19}$  N m assuming a rigidity of the medium of  $3.0 \cdot 10^{10}$  Pa. This value is  
407 close to the median of the seismological one which is  $3.5 \cdot 10^{19}$  N m. The fault parameters  
408 are presented in Table 3 in comparison to the parameters obtained using only ascending  
409 orbit data (Ganas et al. 2020a).

410 The geodetic data does not allow to constrain multiple patches on the fault plane. This is  
411 due to the fact that the fault is offshore, thus without geodetic constraints in the near field  
412 (on its hangingwall). However, the aftershock distribution along the fault (in an E-W

413 direction, Figs. 2b and c), shows a clear differentiation that may be related to either a)  
 414 a fault geometry that is more complex than the simple planar fault that we have adopted  
 415 here, and/or b) that the finite displacement is unevenly distributed along strike. Seismicity  
 416 data alone, however, do not have the required precision in location to clearly delineate the  
 417 geometry of the seismogenic fault (see also seismicity cross-sections in Papadimitriou et al.  
 418 2020 and Fomelis et al. 2021).

419



420

421 **Fig. 8:** Map showing a synthetic interferogram (ascending orbit). The black rectangle  
 422 indicates the surface projection of the modelled fault (dip-direction is towards north).  
 423 Arrows indicate GPS offset vectors (observed vs. predicted; see Table 2). The formation of

424 the synthetic interferogram is at the convention of post-event minus pre-event, thus phase  
 425 decreasing towards satellite.

426

427 **Table 3.** List of fault parameters after inversion of geodetic data.

	<b>Ganas et al. 2020a</b>	<b>This study</b>
Number of inversions	2267	5610
$M_0$ (N-m) – geodetic	3.5e+19	3.06e+19
Mw	7.0	7.0
R.m.s (total, weighted)	16.3	13.0
Centre of Top-fault LON (°)	26.707	26.6784
Centre of Top-fault LAT (°)	37.827	37.8290
Azimuth – strike (°)	N276E	N276E $\pm$ 1
Centre of Top-Fault Depth – km	0.9	0.6 $\pm$ 0.3
Length – km	36	40 $\pm$ 3
Width – km	18	15 $\pm$ 3
Dip (°)	37	37 (not inverted)
Normal slip (m)	1.8	1.7 $\pm$ 0.4
Strike slip (mm)	0	0 (not inverted)
rms InSAR ascending (mm) on pickings		50.2
rms InSAR descending (mm) on pickings		30.5
rms InSAR east component		23
Rms InSAR vertical component		23
Rms InSAR LOS ascending (mm) on whole Samos		50
Rms InSAR LOS descending (mm) on whole Samos		59
rms GPS vertical (mm)		9.8
rms GPS horizontal (mm)		6.4

428

429 With a fault dip-angle assumed to be 37°, the top-edge of the fault at 0.6  $\pm$  0.3 km, and a  
 430 width of 15  $\pm$  3 km, the root of the fault is located at a depth of 9.6  $\pm$  2 km, and the geodetic

431 centroid depth is at  $5.1 \pm 2$  km. This depth of the fault root is consistent with the 10-13 km  
432 range of hypocentre depths determined by USGS, OCA, KOERI, INGV (see Table 1) and  
433 Papadimitriou et al. (2020). For the Kos 2017 M6.6 event with a similar north-dipping  
434 geometry we estimated the fault root at 10 km (Ganas et al. 2019).

435

## 436 **6. Earthquake Environmental Effects**

437 The environmental effects of the 30 October 2020 include rock falls, rock slides, landslides,  
438 road-fill failures, coastal uplift and small-size liquefaction features (see also Lekkas et al.  
439 2020; Foumelis et al. 2021; Mavroulis et al. 2021). The data were collected during a field  
440 survey that took place on 6-8 November 2020, one week after the occurrence of the  
441 earthquake. We were able to timely record the triggered geo-environmental effects  
442 (mainly slope failures and liquefaction). The areas affected by these phenomena are mainly  
443 found on the northern coastline of the island (Fig. 9) such as a) the villages of Avlakia,  
444 Kokkari and the town of Vathy on the east and b) by the coastal zone from Agios  
445 Konstantinos to Karlovasi to the west (Fig. 9). Regarding the other parts of the island, the  
446 earthquake triggered only sparse and small size rock falls and rock slides on sites very prone  
447 to slope failures. In particular, liquefaction phenomena such as sand craters and ejection  
448 of fine-grained sand through ground cracks were documented at the coastal site of  
449 Malagari beach, on the NW edge of the Vathy bay (Fig. 10a). Additional environmental  
450 effects were mapped over Samos island using Copernicus Sentinel-2 & Maxar WorldView  
451 optical satellite imagery. A map showing the spatial distribution of earthquake-induced  
452 failures is shown in Fig. 9, while the coordinates of the field-documented failures are listed  
453 at Table S4.

454 We observed that the dominant geological effects triggered by the earthquake are related  
455 to slope failures. In particular, rock falls and slides, and shallow and deep-seated landslides  
456 were mapped on the northern and western part of the island and at the Karlovassi  
457 hinterland area, on both natural and cut slopes. In several cases we observed that the  
458 failures are associated with pre-existing tectonic discontinuities (faults, joints) and steep  
459 slopes within heavily fractured sedimentary rocks e.g., limestones (as is the case in Avlakia  
460 where sliding occurred along a steep, northeast-facing slope with northeast-dipping  
461 limestone strata; the dip angle was measured by AG between  $51$  and  $53^\circ$ ). Moreover, no  
462 primary fault surface ruptures were observed in the field which is consistent with the off-  
463 shore fault location determined by geodesy (Fig. 8).

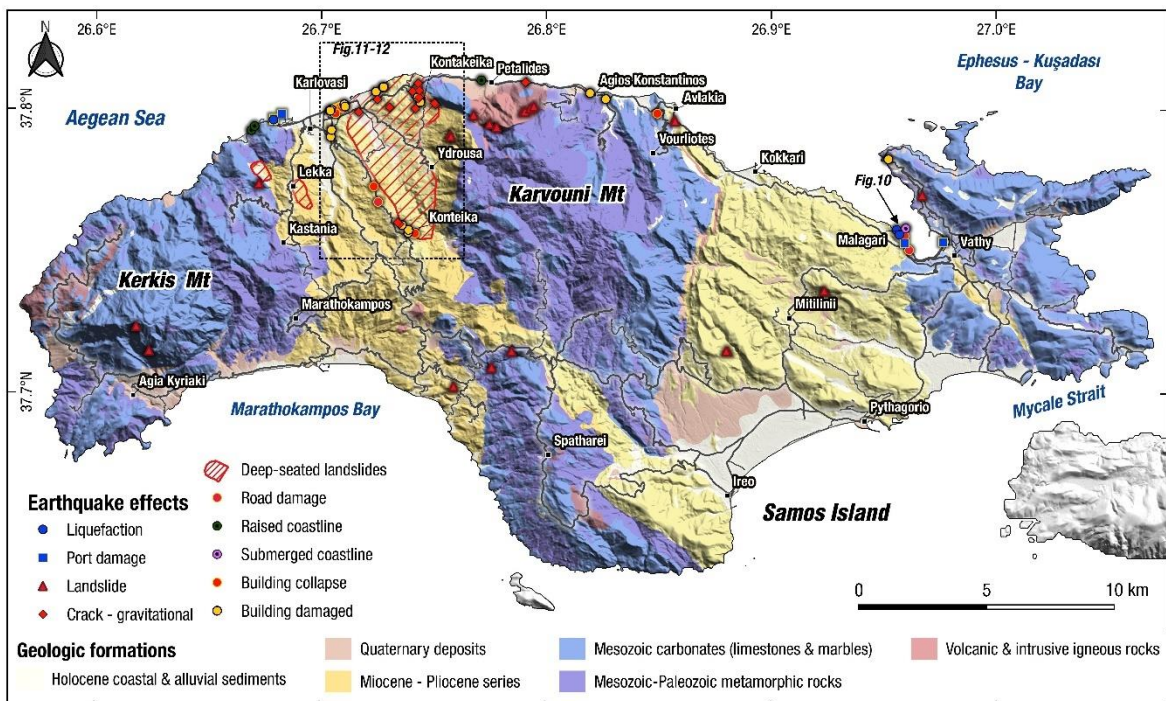
464 The liquefaction site at Malagari beach (Fig. 10a) is a few metres away to a coastal site  
465 which subsided as it is evidenced by the partial submergence of a coastal tree line, that was  
466 located along the beach (approximately oriented N-S). Part of the beach is now under water  
467 with preliminary measurements indicating a maximum subsidence of 53 cm (Fig. 10b, c, d).



468 We suggest that this is due to strong ground shaking that induced lateral spreading of the  
 469 surface soil horizons.

470 Additional data were provided by the InSAR images that showed extensive decorrelation in  
 471 the hilly area both east and south of Karlovasi (Fig. 11). The InSAR images indicate chaotic  
 472 structures that formed inside valleys (Fig. 12). As the structure is clear on all co-seismic pair  
 473 combinations (including the 30 October 2020 scene) it is evident that it does not represent  
 474 a tropospheric effect between the SAR acquisitions or an artefact of the DEM used to make  
 475 the interferograms. We mapped a convergent pattern of motion (Fig. 11a) inside the main  
 476 valley to the east of Neo Karlovassi (Fig. 12) that we interpret as a gravitational feature than  
 477 a primary (seismic) one.

478



479

480 **Fig. 9.** Locations of mapped earthquake environmental effects of the October 30  
 481 earthquake over Samos island (6-8 November 2020 field survey). Arrow shows location of  
 482 Fig. 10. Box shows extent of Fig. 11 & 12. Background geology is a simplified version of the  
 483 geological map of Theodoropoulos (1979).



484

485 **Fig. 10.** Field photographs showing earthquake-induced failures at Malagari, Samos (see  
 486 Fig. 9 for its location, opposite to the town of Vathy) a) liquefaction b) coastal subsidence  
 487 of 53 cm c) submergence of former beach d) submerged tree line.

488

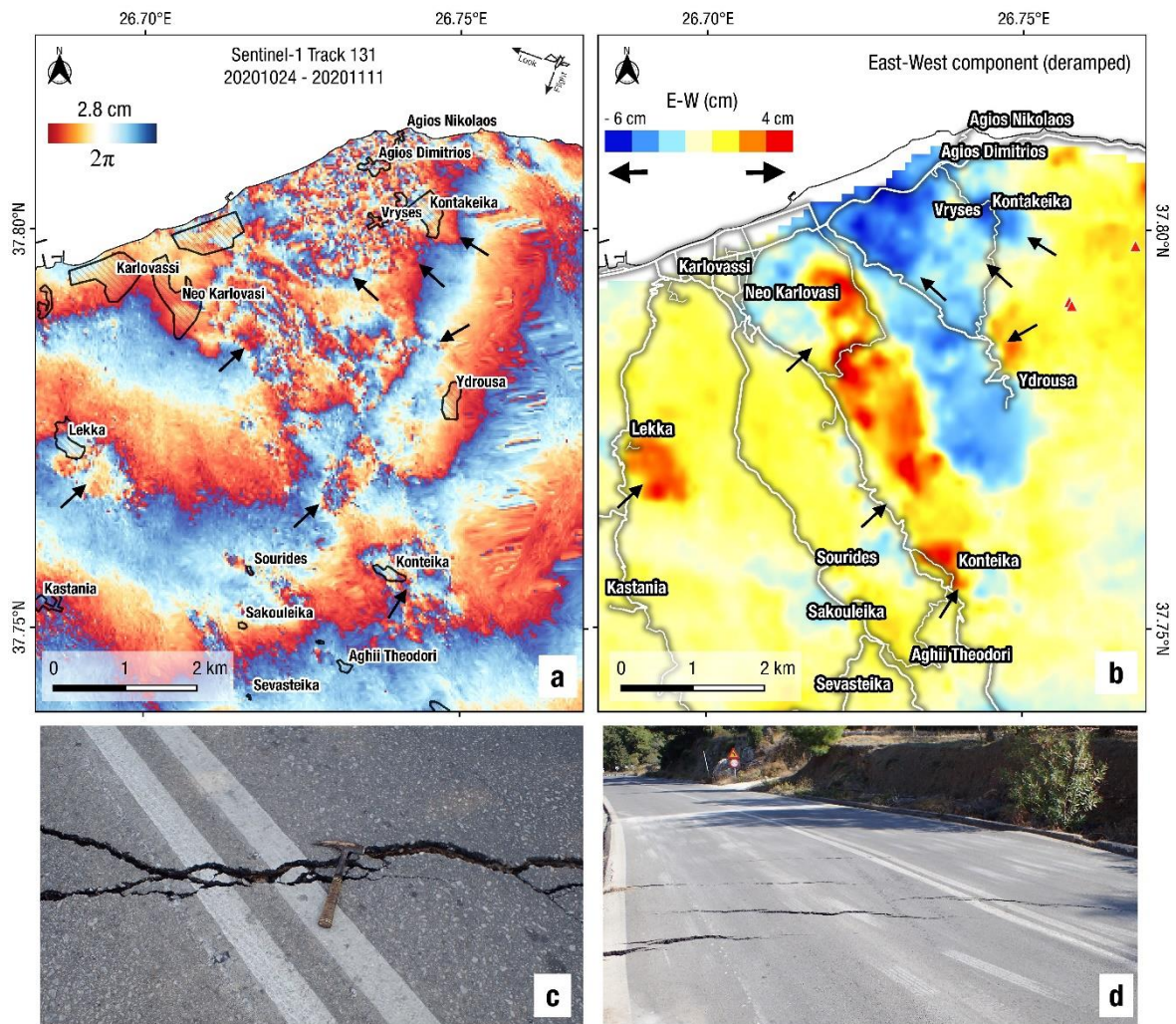
489 Both InSAR and optical satellite data support this interpretation as a set of massive, deep-  
 490 seated gravitational displacements (DSGD; e.g. Moro et al. 2011), distributed along the  
 491 western and eastern slopes of the Neo Karlovassi valley, that were triggered during the 30  
 492 October 2020 earthquake. The decomposed unwrapped interferograms (East – West  
 493 component; Fig. 11b) document the converging motion of the ground surface and show  
 494 the extent of triggered landslides. The valley also contains huge (1-4 km) paleo-landslide  
 495 complexes visible in optical satellite images and the digital elevation model, parts of which  
 496 were activated (Fig. 12). Most clear earthquake-induced landslide cases are the one up-  
 497 stream from Neo Karlovasi (interpreted as a Pliocene limestone cap that slid along ENE-  
 498 dipping beds; Fig. 13), and the area north of village Kontakeika towards the coast, on the  
 499 other side of the valley (Fig. 12). The deformation and offsets from the main activated  
 500 landslide complex near Kontakeika are also visible along the coastal road network and they  
 501 were confirmed by field inspection on November 6, 2020 (Fig. 11 c,d). Due to the small  
 502 amount of displacement of the triggered DSGDs, field evidence is limited to local tensional  
 503 fractures perpendicular to the slide movement axis, and ground cracks along its boundaries  
 504 (Fig. 11d). Co-seismic triggering of large deep-seated landslides is a common occurrence



505 during shallow earthquakes, and are frequently observed through mainly geodesic imaging  
506 due to the small amount of displacement involved (Moro et al. 2011; Lacroix et al. 2014;  
507 Lacroix et al. 2020).

508

509

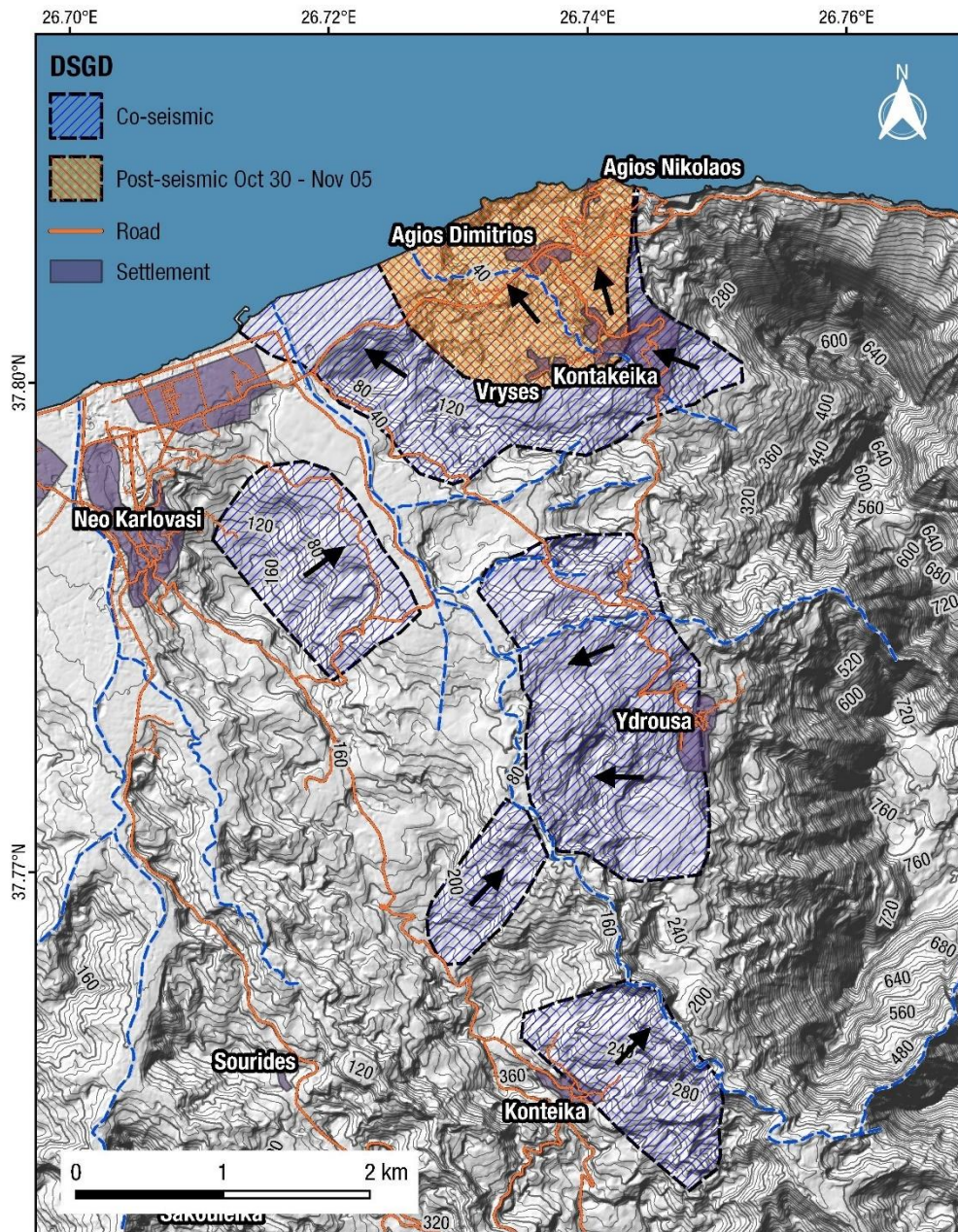


510

511 **Fig. 11.** Maps showing co-seismic landslide activity around Karlovasi, evident in Sentinel-1  
512 differential interferograms. De-correlated areas and fringe patterns mark the location of  
513 DSGDs (black arrows) in descending (a) interferogram. Decomposed unwrapped  
514 interferograms (East – West component) show the extent of triggered landslides (b). The  
515 E-W displacement raster image was deramped around the landslide area in order to  
516 diminish the co-seismic displacement effect and to enhance movement along the eastern  
517 and western slopes. c & d) Field photos of ground cracks along the boundaries of activated  
518 landslides (see Table S4 for coordinates).

519



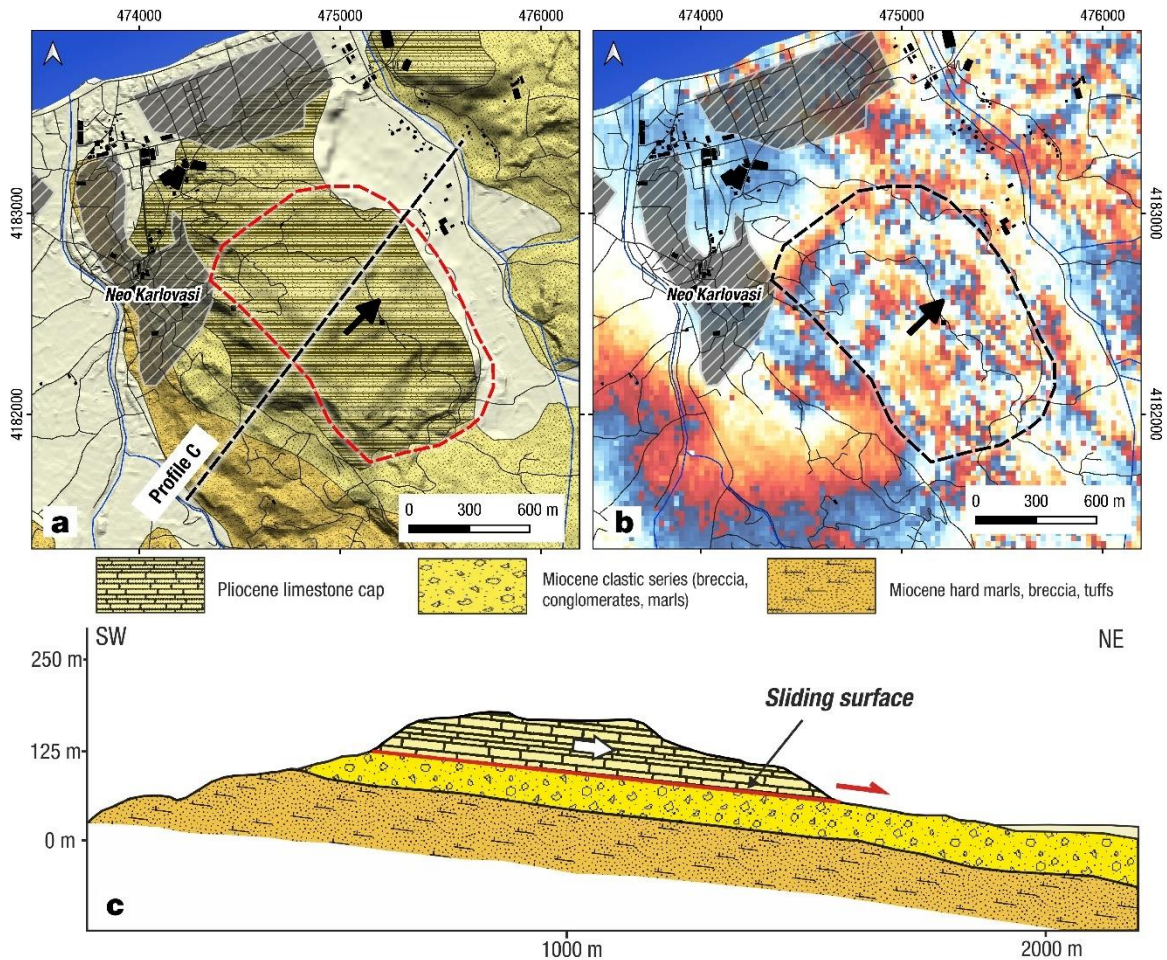


520

521 **Fig. 12.** Shaded relief map of the valley to the east of Karlovasi where deep-seated  
 522 gravitational displacements (DSGD) were observed both in InSAR images (Fig. 11) and in  
 523 the field. The activated slides are delineated by polygons, along both sides of the valley.  
 524 Northernmost structure (orange) showed displacement during the early post-seismic  
 525 period (30 October – 5 November 2020). Contour interval is 10 m.

526





527

528 **Fig. 13.** Geological map (a; Theodoropoulos, 1979), InSAR interferogram of the image pair  
 529 23 October – 4 November 2020 (b) and cross-section (c) of earthquake-induced slide to  
 530 the SE of Neo Karlovassi, Samos. The sliding surface is located at the base of the Pliocene  
 531 limestone beds that overlie the Miocene clastic series of western Samos.

532

## 533 7. Discussion-Conclusions

### 534 7.1 Postseismic deformation of the footwall and relation of co-seismic to post-seismic 535 slip

536

537 One of the important aspects of the Samos earthquake is that it is one of a few cases of  
 538 normal-slip earthquakes where we can accurately measure postseismic deformation of the  
 539 footwall block. This is possible because of the magnitude of the mainshock ( $M_w=7.0$ ) and  
 540 of the existence of sufficient geodetic data. The two GNSS stations (SAMU & SAMO) that  
 541 were situated on the footwall close to the fault (Fig. 6) have provided accurate information  
 542 on ground deformation during both the co-seismic and the post-seismic phases (during the  
 543 first 10-days; Fig. 14 and Table 4). The earthquake produced a significant amount of post-

544 seismic displacement in the near field (a few km from the top-fault edge; Fig. 14), detected  
545 by the two stations; larger in SAMO with a few cm amplitude, a horizontal motion of -4.4  
546 cm (North-South component) and 1.8 cm of uplift; Table 4). In SAMU, located 24 km  
547 towards East, the postseismic motion was 1.6 cm and 0.7 cm, respectively. The ratio of the  
548 postseismic to co-seismic motion per station is 12-25% (north-south component) and 22-  
549 33% (on the vertical component). Interestingly we observe that the co-seismic and post-  
550 seismic vectors are not aligned; the SAMO station shows a motion towards N202°E during  
551 the postseismic period as opposed to N185°E during the co-seismic one. The fastest part of  
552 the postseismic motion lasted until 10 November 2020, that is 11 days (Fig. 14 inset). Then,  
553 the decay was slower until the 1<sup>st</sup> of December 2020 where we cannot see it anymore  
554 because of data being comparable to noise (Fig. 14 inset). Thus, the duration of the  
555 detectable post-seismic signal is 30 days which is measured on the horizontal component  
556 of station SAMO. This value is close to the 21 days we found in the case of the  $M_w=6.7$   
557 Zakynthos earthquake (Ganas et al. 2020b). It is reasonable to suggest that the cause of the  
558 deformation is shallow afterslip along the rupture plane of the mainshock as it has been  
559 documented in other similar ruptures (e.g., Cheloni et al. 2014; Pousse-Beltran et al. 2020;  
560 Briole et al. 2021). Shallow afterslip near the top-edge of the rupture plane may be  
561 responsible for the early aftershocks (24-hr) that are aligned along Samos’s north coast (Fig.  
562 1). For capturing the full pattern of the fast post-seismic motion that lasts just a few weeks  
563 it would have been necessary to deploy in the field additional GNSS stations which were  
564 not available. This would be also necessary to fully capture the flexure of the footwall away  
565 from the rupture (e.g., Fernandez-Blanco et al. 2020), as our data (Fig. 15) plot on the flat  
566 part of the vertical displacement profile.

567

568 **Table 4.** Estimation of seismic displacements at stations SAMU and SAMO and their InSAR  
569 equivalents. Ratio is the post-seismic displacement divided by the cos-seismic one. The  
570 post-seismic data refer to the first 10-day period following the mainshock.

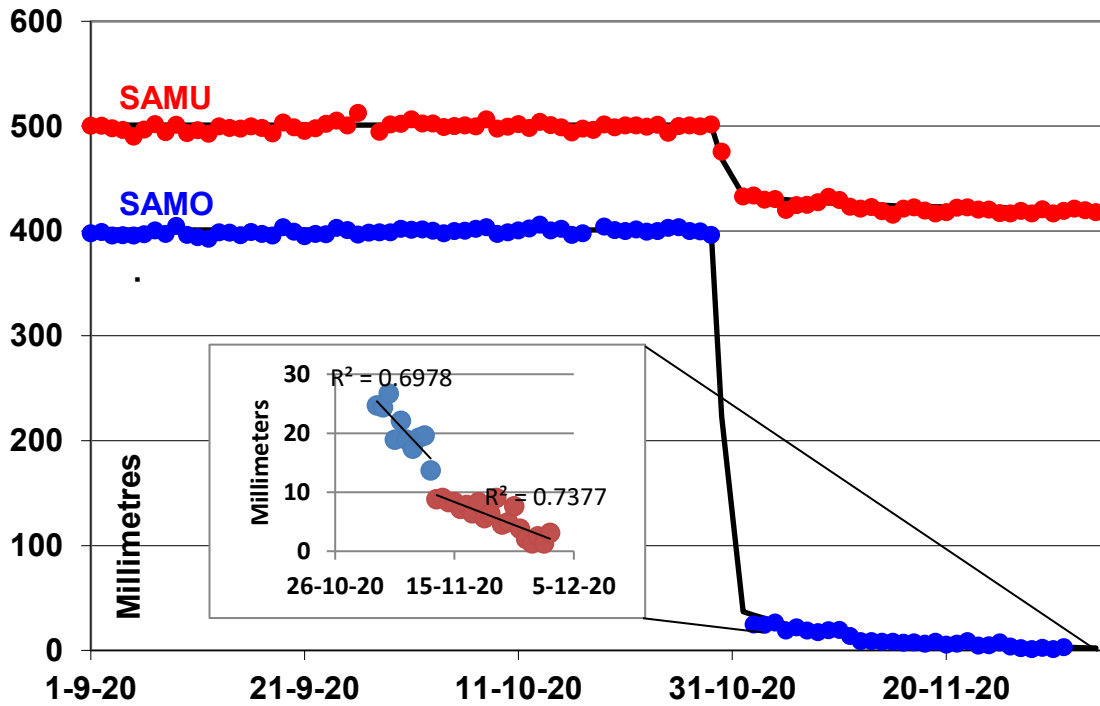
571

Station	dE (mm)	dN (mm)	dU (mm)	I.o.s (mm) track 29	I.o.s (mm) track 36
SAMU co-seismic	-5	-64	21	33	23
SAMU post-seismic	0	-16	7	8	8
SAMO co-seismic	-30	-356	82	155	82
SAMO post-seismic	-11	-44	18	27	13
SAMU ratio	0	0.25	0.33	0.24	0.35
SAMO ratio	0.36	0.12	0.22	0.17	0.16

572

573





574

575 **Fig. 14.** Time series of North coordinates of station SAMO (Karlovasi; blue dots) and SAMU  
 576 (Vathy; red dots). The inset box shows the post-seismic period of SAMO with a break in  
 577 slope around 11 November 2020, thus signifying the end of the fast term of the post-  
 578 seismic motion.

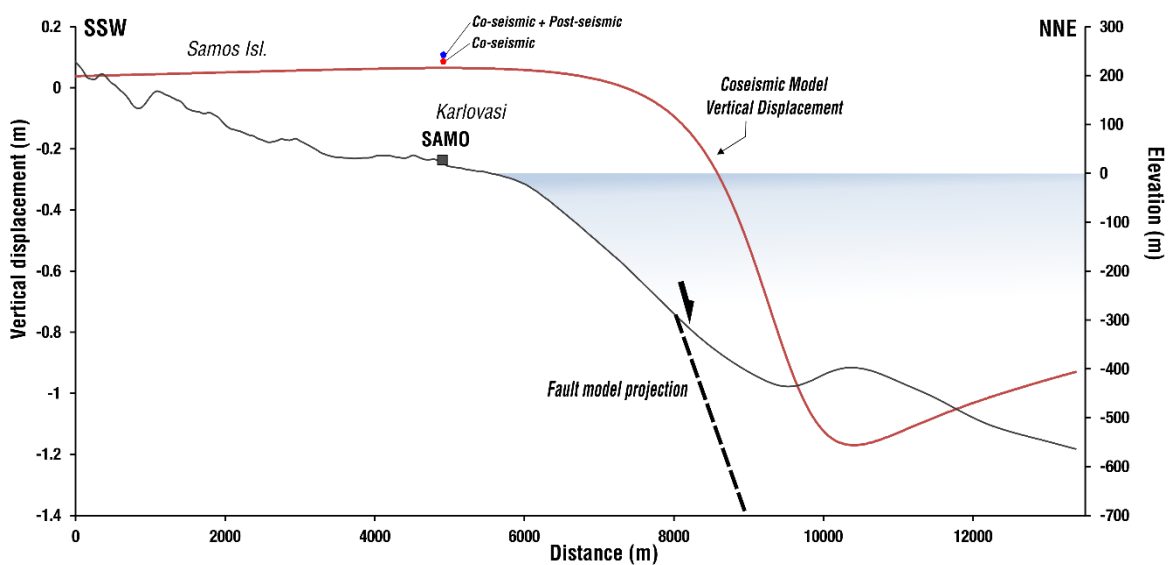
579

580 We also detected postseismic motions in the InSAR data (as did Foumelis et al. 2021), in  
 581 particular a post-seismic signal is visible on track 131, that was acquired on the same day,  
 582 ~4h15' after the mainshock. In three out of five postseismic interferograms having as  
 583 primary image (master) this particular acquisition, one fringe is visible, spread out in the  
 584 north of the island, with the same direction of motion (LOS) with that of the co-seismic  
 585 ones. To compare with the GPS data (Table 4), the postseismic difference in the LOS of  
 586 pixels enclosing the two GPS stations of the island, for the ascending pass are 19 mm (5  
 587 mm for the LOS. of the descending pass), and in terms of percentage the postseismic LOS  
 588 motion is about 15-34 % of the co-seismic one. Our observations support almost one whole  
 589 fringe on the north coast of the island (ending northern than SAMO), however the same  
 590 area has been partially subjected to gravitational motions (see Fig. 12 and Fig. 13) so at this  
 591 stage we didn't proceed further into differentiating those signals. Moreover, the  
 592 deformation mapped in post-seismic interferograms (secondary images on 11, 17 and 23  
 593 November 2020), demonstrate the continuation of the post-seismic process, in agreement  
 594 with the GPS data (Fig. 13). These post-seismic interferograms are stored in the AIPS (see  
 595 text S2 and <http://aips-p.space.noa.gr/Samos/products/131/post>).

596 In terms of vertical motions, we also observe a fast postseismic uplift in the data of footwall  
 597 stations SAMO (Fig. S3) and SAMU. The postseismic uplift ranges between 12-28% of the

598 co-seismic value measured at the GNSS stations (Table 4) that are located about 3.7 km  
 599 south of the top-fault line surface projection (Fig. 15). As already mentioned this  
 600 deformation is due to fault afterslip, so it will be sustained over months or years (Ergintav  
 601 et al. 2009; Briole et al. 2021) because of the size of the earthquake ( $M_w=7.0$ ). We also note  
 602 that the vertical postseismic motion is on the same direction as the co-seismic motion (Fig.  
 603 15), i.e., positive thus signifying further uplift of the footwall. We suggest that postseismic  
 604 uplift is a process that should be included in long-term deformation studies in both fore-  
 605 arc (e.g., Crete; Robertson et al. 2019; Ott et al. 2021) and back settings (e.g., Corinth;  
 606 Houghton et al. 2003; Fernandez-Blanco et al. 2020).

607



608

609 **Fig. 15.** Cross-section across the seismic fault showing vertical displacement (according to  
 610 the model of Table 3; red line). GNSS station SAMO is shown on the footwall. The vertical  
 611 displacement (in m) is shown on the left-Y axis. Elevation and bathymetry are shown on the  
 612 right-Y axis. Vertical exaggeration is times 6 the horizontal (X-axis).

613

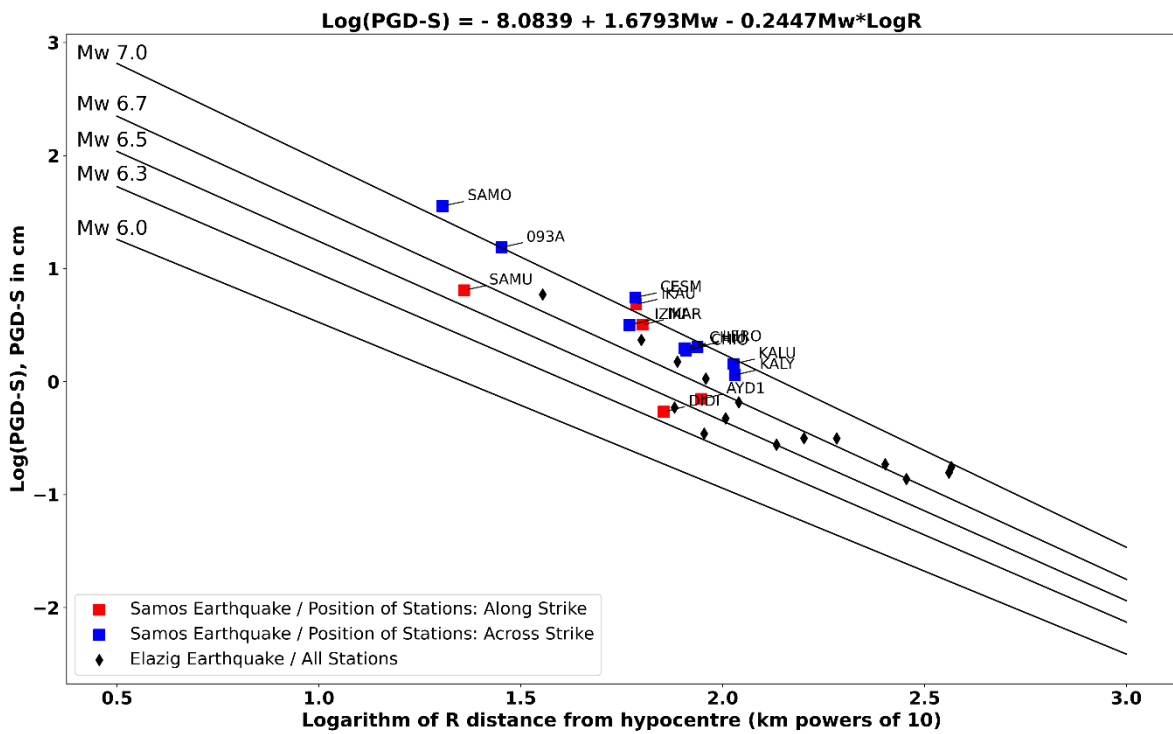
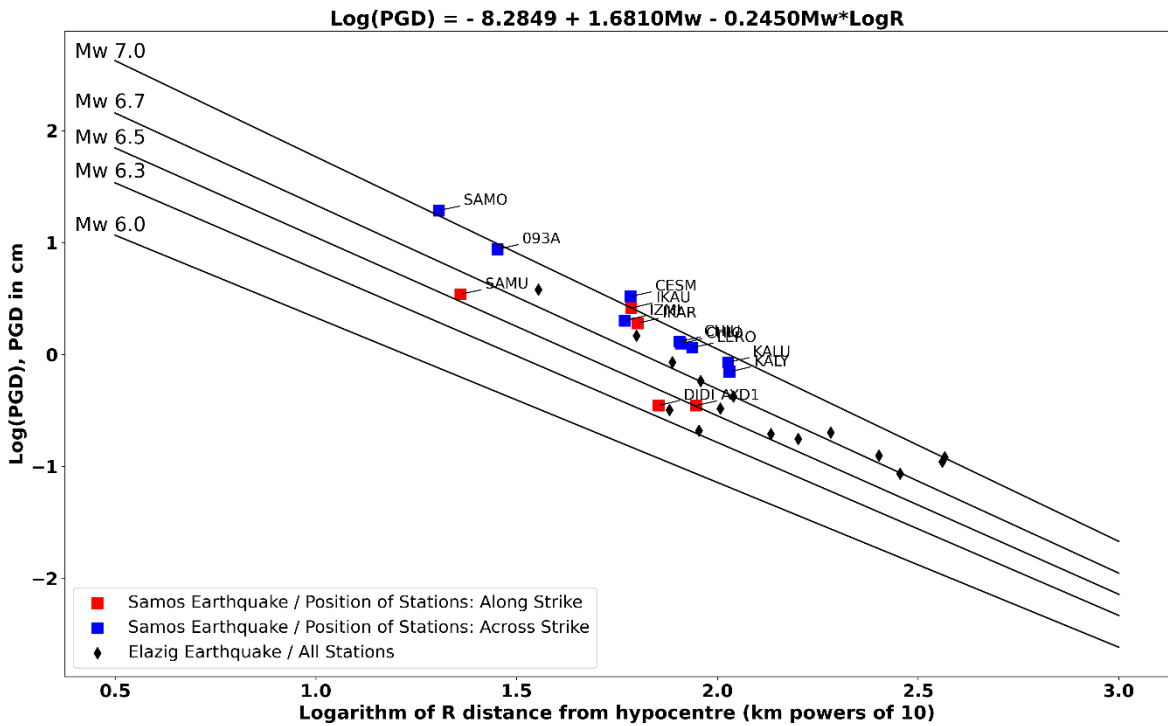
## 614 7.2 Asymmetric Elastic response of the crust

615 Geodetic processing of GNSS observations can supply reliable and unsaturated  
 616 measurements of ground motion displacement during earthquakes (e.g., Melgar et al.  
 617 2015; Huang et al. 2017; Ruhl et al. 2019). The displacement data are extremely useful for  
 618 earthquakes with moment magnitudes larger than  $M=5.0$  and originating from stations at  
 619 the near-field (equal to or less than one fault length) up to regional distances (200-300 km  
 620 or even more). The displacements comprise measurements (offsets) of GNSS antenna  
 621 positions due to tectonic strain of the crust imposed by the passage of seismic waves. The  
 622 peak ground displacements (PGD) are defined on the horizontal plane (sum of offsets of  
 623 the North-South and East-West component of GNSS; e.g., Ganas et al. 2018) and/or all

624 three directions of motion (3D; Melgar et al. 2015). The PGDs are related to earthquake  
625 magnitude ( $M_w$ ) and distance to hypocentre  $R$  by a relation of the form  $\text{LOG}(\text{PGD}) = a + b *$   
626  $M_w + c * \text{LOG}R * M_w$  (Ganas et al. 2018). The Samos event data (Table 2) show that GNSS  
627 horizontal displacements increase closer to the epicentre in a systematic way, in agreement  
628 with the response of an elastic body to tectonic strain (Fig. 6; Fig. 16). First, we observe that  
629 GNSS offset data from “near-field” stations (one fault length) SAMO, SAMU, 093A, (Fig. 6)  
630 are in agreement with the PGD scaling relationships where offset data are correlated to  
631 earthquake magnitude and distance to hypocentre (Melgar et al. 2015; Ganas et al. 2018b;  
632 Ruhl et al. 2019). As with other shallow earthquakes in the Aegean (period of data 1997-  
633 2017; Ganas et al. 2018) we mapped a linear displacement field (in Log-Log space) as  
634 expected from an elastic response of the crust to the permanent strain caused by  
635 instantaneous seismic slip along the offshore fault. However, the stations located across  
636 strike (blue squares in Fig. 16) align close to the  $M_w=7.0$  regression line (in agreement with  
637 seismology estimates) while the stations located along strike ( $\pm 45^\circ$ ; red squares) are  
638 aligned close to the  $M_w=6.5$  regression line. We attribute this difference in magnitude  
639 scaling with distance to the azimuthal position of the GNSS station with respect to the strike  
640 of the rupture plane. The pattern of tectonic strain is asymmetric in elastic earth (see the  
641 Okada-type model of Fig. S5) and this is reflected on the co-seismic geodetic data  
642 (horizontal GNSS offsets). Moreover, there are two stations onshore Ikaria island (IKAR &  
643 IKAU; Fig. 6) that are located to the SW of the rupture at a distance of  $\sim 60$  km from the  
644 hypocentre. Those stations scale with the  $M_w=7.0$  regression line despite their azimuthal  
645 position ( $\sim 45^\circ$  off-strike); we believe that their offsets are impacted by a possible  
646 (westward) directivity effect of the rupture. In case of a bilateral rupture, we would  
647 anticipate lower PGD at IKAU & IKAR and higher PGD in SAMU, DIDI and ADY1, as the latter  
648 three stations are located towards east of the epicentre. The Samos case demonstrates  
649 that (when available) GNSS stations across strike of dip-slip ruptures should be considered  
650 first in magnitude-scaling estimates. .

651

652



655 **Fig. 16:** (top panel) two-dimensional graph of geodetic data (Y-axis; cm) that uses  
 656 logarithmic scales to show a linear decrease of PGD (peak ground displacement) with  
 657 hypocentral distance (X-axis; km) for the Elazig (black rhombs) and Samos (coloured  
 658 squares) earthquakes2020). Blue squares show the Samos earthquake data from GNSS  
 659 stations across strike. Red stars indicate the data from GNSS stations along strike. Date  
 660 format in legend (upper right) is YYYYMMDD. (bottom panel) Graph as above but showing  
 661 decrease of PGD-S (cm) with hypocentral distance (km). The oblique blue lines are the

662 predicted scaling values for a particular earthquake magnitude from the relations above  
663 each graph (from Ganas et al. 2018). The Elazig earthquake data (24 Jan. 2020;  $M_w=6.7$ ) are  
664 from Melgar et al. (2020).

665

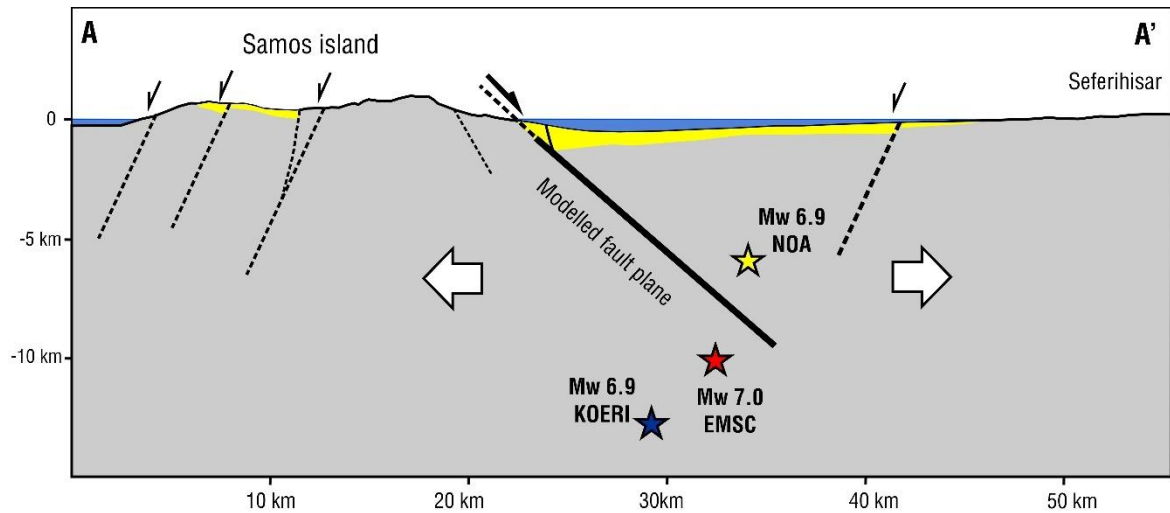
### 666 **7.3 Extension across the Samos – Kuşadası Bay**

667 In the eastern Aegean crustal extension is accommodated by a combination of normal-slip  
668 and strike-slip faulting as demonstrated by geological, seismological and geodetic data (i.e.,  
669 Boronkay and Doutsos, 1994; Ocakoglou et al. 2004; Benetatos et al. 2006; Ozkaymak and  
670 Sozbilir, 2012; Chatzipetros et al. 2013; Tan et al. 2014; Yolsal-Çevikbilen et al. 2014; Ganas  
671 et al. 2019), however, in the offshore area of Samos detailed knowledge was lacking.  
672 According to our joint-inversion (GNSS & InSAR) model, the main structure is the 40-km  
673 long seismic fault (activated by the Samos 2020 earthquake) that dips towards north at an  
674 angle of  $37^\circ$  and reaches the depth of 10 – 11 km (Fig. 17). The root of the fault is at the  
675 base of the upper crust which comprises about 40-50% of the total crustal thickness (about  
676 30-km thick in the area of Samos; Grigoriadis et al. 2016). As such this fault comprises a  
677 major Quaternary structure of the Aegean Sea given that fault length is about 3.3 times the  
678 seismogenic part of the crust. It seems that fault growth was more efficient along strike of  
679 the fault, although this is not reflected in sea-bottom morphology (that deepens towards  
680 the western part of the footwall), similarly to the 1949 earthquake offshore Chios (Melis et  
681 al. 2020) and the 2019 earthquake offshore Kos (Ganas et al. 2019).

682 The Samos fault throw assuming a) pure normal slip and b) flat pre-rift topography, should  
683 exceed 2 km (by summing footwall relief south of Karlovassi and offshore basin depth), thus  
684 providing a preliminary estimate for a mean throw rate of 1 mm/yr or about 1.4 mm/yr  
685 mean slip rate. This fault accommodates a significant portion of tectonic strain  
686 accumulation in the Samos – Kuşadası Gulf (Fig. 17) where a crustal block boundary has  
687 been suggested by Vernant et al (2014) and Briole et al. (2021). Both fault size and  
688 geometry indicate that this is a relatively young structure developed during the Quaternary  
689 and represents the main source of seismic hazard in this region. We note that our geodetic  
690 data inversion is fit by a planar fault (40 by 15 km), although we are well aware that in  
691 nature faults are non-planar.

692 The surface projection of the fault is found near the northern shore of Samos (Fig. 15; Fig.  
693 17), about 1.5-km offshore the port of Karlovassi, assuming a  $37^\circ$  dip-angle. In addition, the  
694  $N276^\circ E$  orientation of the fault is roughly similar to that of the active faults and/or  
695 seismogenic sources in this area (Chatzipetros et al. 2013; Caputo and Pavlides, 2013;  
696 Ganas et al. 2013). All databases contain an offshore north-dipping normal fault that is  
697 described as the main seismic source for the Samos Gulf.

698



699

700 **Fig. 17.** Tectonic model across the Samos Gulf (See Fig. 1 for location A – A'). Black lines are  
 701 active faults (half arrow indicates downthrown side). The bold black line corresponds to the  
 702 model fault, and dashed lines are inferred normal faults from literature. The yellow colour  
 703 indicates the recent sediment cover. Solid stars indicate mainshock hypocentres, projected  
 704 onto the cross-section. White solid arrows indicate long-term extensional strain within the  
 705 crust.

706

707 Although it is crucial for short-term seismic hazard assessment, we cannot infer the  
 708 structure towards east, i.e., towards the Kuşadası area where large active normal faults  
 709 exist (Meng et al. 2021). It is possible that the 2020 rupture plane continues eastwards for  
 710 another 10-15km or that another, north-dipping fault segment exists towards the east of  
 711 the 2020 rupture. Preliminary aftershock data (Papadimitriou et al. 2020; Foumelis et al.  
 712 2021) are not conclusive on the geometry of the structure towards east. The eastern cluster  
 713 of the 2020 aftershocks (offshore NE Samos and onshore eastern Samos; Fig. S1) probably  
 714 reflects triggered seismicity on optimally-oriented faults to N-S extension because of  
 715 Coulomb stress transfer (Papadimitriou et al. 2020). More offshore data are needed to  
 716 better define the fault zone geometry and size, and shed light on a possible co-planar (to  
 717 2020) configuration of seismic sources or not. Towards the west, the recent offshore  
 718 seismic survey data by Nomikou et al. (2021), confirm that the Samos earthquake seismic  
 719 fault has its surface trace roughly projected near the Samos coast (in Fig. 8) following the  
 720 planar geometry assumed in our inversion. In addition, the main rupture patch extends for  
 721 ~25-km west of Karlovassi (Fig. 1) as evidenced by the visible E-W gap in the early  
 722 aftershock (24-hr) distribution. In summary, the proposed fault characteristics (Table 3) as  
 723 revealed by the 2020 earthquake in terms of location, geometry and seismic potential are  
 724 robust and they should update the fault databases of active structures of the eastern  
 725 Aegean Sea (e.g., Papazachos et al., 2001; Ganas et al., 2013; Caputo and Pavlides, 2013;  
 726 Sakellariou et al., 2013; Emre et al., 2016).

727

## 728 **8. Conclusions**

- 729 1. Our joint inversion (GNSS and InSAR) modeling indicates that the Samos M7.0  
730 earthquake occurred along a north-dipping normal fault. The fault is located  
731 offshore Samos occupying a length of 40 km, width of 17 km, an average coseismic  
732 slip of 1.7 m, and a moderate dip-angle (37°).
- 733 2. The root mean square (r.m.s) error between geodetic data and model is 6.4 mm for  
734 the GPS horizontal data, 9.8 mm for the GPS vertical data and between 23 - 59 mm  
735 for the InSAR data (see Table 3 for model performance), respectively.
- 736 3. The earthquake caused the permanent uplift of the island in the range 1-10 cm with  
737 the exception of a coastal strip along part of the northern shore (between Agios  
738 Konstantinos – Kokkari) that subsided 2-6 cm.
- 739 4. We observed a fast postseismic slip (uplift) ranging between 22-33% of the co-  
740 seismic slip value. This is primarily due to fault shallow afterslip as it is recorded  
741 within the first 10-days since the mainshock. Postseismic uplift is an important  
742 process that should be included in earthquake deformation studies.
- 743 5. The effects of the earthquake included liquefaction, rock falls, rock slides, road  
744 cracks and deep-seated landslides, all due to the strong ground motion and  
745 associated down-slope mobilization of soil cover and loose sediments.
- 746 6. The Samos earthquake demonstrates that (when available) GNSS station data  
747 (offsets) across strike of dip-slip ruptures should be considered first in magnitude-  
748 scaling estimates.

749

750 **Acknowledgements:** We thank Brendan Crowell, Nicolas Chamot-Rooke, Ross Stein, Marco  
751 Meschis, Tuncay Taymaz, Diego Melgar, George Papathanassiou, Evi Nomikou, Isidoros  
752 Kampilis, Ioannis Kassaras, Antonio Avallone, Eirini Efstathiou, Thekla Thoma and Efthimios  
753 Lekkas for comments and discussions. Sentinel-1 and Sentinel-2 satellite imagery was  
754 acquired through European Space Agency and Copernicus, and WorldView imagery  
755 through Maxar Open Data program. We are indebted to ESA, Geohazards Lab and Terradue  
756 for providing access to Geohazards Exploitation Platform (GEP) for InSAR cloud processing.  
757 GNSS data were provided by Ktimatologio SA, Hexagon SmartNET and Uranus (Tree)  
758 networks of Greece. We thank Tuncay Taymaz and Semi Ergintav for sharing Turkish GNSS  
759 data. We thank the Samos local authorities for their help. Several figures were created by  
760 use of GMT (Wessel et al. 2019).

761 **Conflict of interest:** The authors declare no conflict of interest

762

763



764 **References**

- 765 Angelier J. 1979. Néotectonique de l' arc Egéen. Soc. Géol. Nord, 3, pp. 1-417.
- 766 Atzori, S., M. Manunta, G. Fornaro, A. Ganas, S. Salvi, 2008. Postseismic displacement of  
767 the 1999 Athens earthquake retrieved by the Differential Interferometry by Synthetic  
768 Aperture Radar time series, Journal of Geophysical Research, 113, B09309,  
769 doi:10.1029/2007JB005504
- 770 Benetatos, C., A. Kiratzi, A. Ganas, M. Ziazia, A. Plessa, and G. Drakatos 2006. Strike-slip  
771 motions in the Gulf of Siğaçık (western Turkey): Properties of the 17 October 2005  
772 earthquake seismic sequence, Tectonophysics 426, 3-4, 263-279, DOI:  
773 10.1016/j.tecto.2006.08.003.
- 774 Boronkay, K. and Doutsos, T. 1994. Transpression and transtension within different  
775 structural levels in the central Aegean region. Journal of Structural Geology, 16 (11), 1555-  
776 1573, [https://doi.org/10.1016/0191-8141\(94\)90033-7](https://doi.org/10.1016/0191-8141(94)90033-7)
- 777 Briole, P., 2017. Modelling of earthquake slip by inversion of GNSS and InSAR data assuming  
778 homogeneous elastic medium, Zenodo. <http://doi.org/10.5281/zenodo.1098399>
- 779 Briole, P., A. Ganas, P. Elias, D. Dimitrov, 2021. The GPS velocity field of the Aegean. New  
780 observations, contribution of the earthquakes, crustal blocks model, Geophysical Journal  
781 International, ggab089, <https://doi.org/10.1093/gji/ggab089>
- 782 Caputo, R. S. Pavlides, 2013. The Greek Database of Seismogenic Sources (GreDaSS),  
783 version 2.0.0: A compilation of potential seismogenic sources (Mw>5.5) in the Aegean  
784 Region. DOI:10.15160/UNIFE/GREDASS/0200.
- 785 Cetin, K. O., Mylonakis, G., Sextos, A. Stewart, J. P. Seismological and engineering effects of  
786 the Mw 7.0 Samos Island (Aegean Sea) Earthquake, 2020. Hellenic Association of  
787 Earthquake Engineering. Report 2020/02. <https://doi.org/10.18118/G6H088>.
- 788 Chatzipetros, A., Kiratzi, A., Sboras, S., Zouros, N., Pavlides, S., 2013. Active faulting in the  
789 north-eastern Aegean Sea Islands. Tectonophysics, 597-598, pp. 106-122.
- 790 Cheloni, D. R. Giuliani, E. D'Anastasio, S. Atzori, R.J. Walters, L. Bonci, N. D'Agostino, M.  
791 Mattone, S. Calcaterra, P. Gambino, F. Deninno, R. Maseroli, G. Stefanelli, 2014. Coseismic  
792 and post-seismic slip of the 2009 L'Aquila (central Italy) MW 6.3 earthquake and  
793 implications for seismic potential along the Campotosto fault from joint inversion of high-  
794 precision levelling, InSAR and GPS data. Tectonophysics, 622, 168-185,  
795 <https://doi.org/10.1016/j.tecto.2014.03.009>.
- 796 Chiaraluce, L., Valoroso, L., Piccinini, D., Di Stefano, R., and De Gori, P. 2011. The anatomy  
797 of the 2009 L'Aquila normal fault system (central Italy) imaged by high resolution foreshock  
798 and aftershock locations, J. Geophys. Res., 116, B12311, doi:10.1029/2011JB008352.

799 D'Agostino, N., M. Métois, R. Koci, L. Duni, N. Kuka, A. Ganas, I. Georgiev, F. Jouanne, N.  
800 Kaludjerovic, R. Kandić, 2020. Active crustal deformation and rotations in the southwestern  
801 Balkans from continuous GPS measurements. *Earth and Planetary Science Letters*, 539,  
802 116246, <https://doi.org/10.1016/j.epsl.2020.116246>

803 Ergintav, S., McClusky, S., Hearn, E., Reilinger, R., Cakmak, R., Herring, T., Ozener, H., Lenk,  
804 O., and Tari, E. 2009. Seven years of postseismic deformation following the 1999, M = 7.4  
805 and M = 7.2, Izmit-Düzce, Turkey earthquake sequence, *J. Geophys. Res.*, 114, B07403,  
806 doi:10.1029/2008JB006021

807 Evelpidou, N., Karkani, A., Kampolis, I. 2021. Relative Sea Level Changes and  
808 Morphotectonic Implications Triggered by the Samos Earthquake of 30th October 2020. *J.*  
809 *Mar. Sci. Eng.*, 9, 40. <http://doi.org/10.3390/jmse9010040>

810 Fernández-Blanco, D., de Gelder, G., Lacassin, R., & Armijo, R. 2020. Geometry of flexural  
811 uplift by continental rifting in Corinth, Greece. *Tectonics*, 39, e2019TC005685.  
812 <https://doi.org/10.1029/2019TC005685>

813 Fomelis, M., Papazachos, C., Papadimitriou, E. et al. 2021. On rapid multidisciplinary  
814 response aspects for Samos 2020 M7.0 earthquake. *Acta Geophys.*  
815 <https://doi.org/10.1007/s11600-021-00578-6>

816 Ganas, A., G. Drakatos, S. B. Pavlides, G. N. Stavrakakis, M. Ziazia, E. Sokos, and V. K.  
817 Karastathis, 2005. The 2001 Mw = 6.4 Skyros earthquake, conjugate strike-slip faulting and  
818 spatial variation in stress within the central Aegean Sea, *J. Geodyn.*, 39(1), 61 – 77,  
819 doi:10.1016/j.jog.2004.09.001.

820 Ganas, A., Drakatos, G., Rontogianni, S., Tsimi, C., Petrou, P., Papanikolaou, M., Argyrakis,  
821 P., Boukouras, K., Melis, N., Stavrakakis, G. 2008. NOANET: the new permanent GPS  
822 network for Geodynamics in Greece. European Geophysical Union, Geophysical Research  
823 Abstracts 10, EGU2008-A-04380.

824 Ganas, A., Parsons, T., 2009. Three-dimensional model of Hellenic Arc deformation and  
825 origin of the Cretan uplift. *J. Geophys. Res.: Solid Earth* 114 (B6)  
826 <https://doi.org/10.1029/2008JB005599> .

827 Ganas, A., Oikonomou, A. and Tsimi, Chr, 2013. NOAfaults: a digital database for active  
828 faults in Greece. *Bulletin Geological Soc. Greece*, 47, 518-530,  
829 <http://dx.doi.org/10.12681/bgsg.11079>

830 Ganas, A., Elias, P., Kapetanidis, V, Valkaniotis, S., Briole, P., Kassaras, I., Argyrakis, P.,  
831 Barberopoulou, A., Moshou, A., 2019. The July 20, 2017 M6.6 Kos Earthquake: Seismic and  
832 Geodetic Evidence for an Active North-Dipping Normal Fault at the Western End of the Gulf  
833 of Gökova (SE Aegean Sea), *Pure and Applied Geophysics*, 176 (10), 4177-4211  
834 <https://doi.org/10.1007/s00024-019-02154-y>

835 Ganas, A., Elias, P., Briole, P., Tsironi, V., Valkaniotis, S., Escartin, J., Karasante, I., and  
836 Efstathiou, E., 2020a, Fault responsible for Samos earthquake identified, Temblor,  
837 <http://doi.org/10.32858/temblor.134>

838 Ganas A, Briole P, Bozionelos G, Barberopoulou A, Elias P, Tsironi V, Valkaniotis S, Moshou  
839 A, Mintourakis I, 2020b. The 25 October 2018 Mw= 6.7 Zakynthos earthquake (Ionian Sea,  
840 Greece): a low-angle fault model based on GNSS data, relocated seismicity, small tsunami  
841 and implications for the seismic hazard in the west Hellenic Arc, Journal of Geodynamics  
842 doi: <https://doi.org/10.1016/j.jog.2020.101731>

843 Ganas, A., Elias, P., Briole, P., Cannavo, F., Valkaniotis, S., Tsironi, V., Partheniou, E.I. 2020c.  
844 Ground Deformation and Seismic Fault Model of the M6.4 Durres (Albania) Nov. 26, 2019  
845 Earthquake, Based on GNSS/INSAR Observations. Geosciences, 10 (6), 210  
846 <https://www.mdpi.com/2076-3263/10/6/210/htm>

847 Goldstein, R. M.; Werner, C. L. 1998. Radar interferogram filtering for geophysical  
848 applications. Geophys. Res. Lett. 25(21), 4035-4038.

849 Grigoriadis, V.N., Tziavos, I.N., Tsokas, G.N. et al. 2016. Gravity data inversion for Moho  
850 depth modeling in the Hellenic area. Pure Appl. Geophys. 173, 1223–1241.  
851 <https://doi.org/10.1007/s00024-015-1174-y>

852 Houghton, S. L., G. P. Roberts, I. D. Papanikolaou, J. M. McArthur, and M. A. Gilmour, 2003.  
853 New 234U-230Th coral dates from the western Gulf of Corinth: Implications for extensional  
854 tectonics, Geophys. Res. Lett., 30(19), 2013, doi:10.1029/2003GL018112.

855 Huang, Y., Yang S., Qiao, X., Lin, M., Zhao, B., Tan, K. 2017. Measuring ground deformations  
856 caused by 2015 Mw7.8 Nepal earthquake using high-rate GPS data, Geodesy and  
857 Geodynamics, 8, 4, pp. 285-291, <https://doi.org/10.1016/j.geog.2017.03.003>

858 Institute of Engineering Seismology and Earthquake Engineering, 2020. Earthquake North  
859 of Samos Island (Greece) of 30/10/2020, Preliminary Report, Thessaloniki.

860 Kalogeras I, Melis NS, Kalligeris N 2020. The earthquake of October 30th, 2020 at Samos,  
861 Eastern Aegean Sea, Greece. Preliminary Report, National Observatory of Athens, Institute  
862 of Geodynamics. <https://accelnet.gein.noa.gr/2020/11/09/the-earthquake-of-october-30th-2020-at-samos-eastern-aegean-sea-greece-preliminary-report>  
863 Lacroix, P.,  
864 Perfettini, H., Taïpe, E. and Guillier, B., 2014. Coseismic and postseismic motion of a  
865 landslide: Observations, modeling, and analogy with tectonic faults. Geophys. Res. Lett. 41,  
866 6676–6680 <http://doi.org/10.1002/2014GL061170>

867 Lacroix, P., Handwerger, A.L. and Bièvre, G., 2020. Life and death of slow-moving landslides.  
868 Nat. Rev. Earth. Environ., 1, 404–419. <https://doi.org/10.1038/s43017-020-0072-8>

869 Lekkas, E., Mavroulis, S., Gogou, GA, et al 2020. The October 30, 2020, Mw 6.9 Samos  
870 (Greece) earthquake. Newsletter of Environmental, Disaster and Crises Management  
871 Strategies, 21, ISSN 2653-9454, Athens, Greece.

872 Mascle, Jean, and Laure Martin, 1990. Shallow structure and recent evolution of the  
873 Aegean Sea: A synthesis based on continuous reflection profiles, Marine Geology, Volume  
874 94, Issue 4, Pages 271-299.

875 Mavroulis, S., Triantafyllou, I., Karavias, A., Gogou, M., Katsetsiadou, K.-N., Lekkas, E.,  
876 Papadopoulos, G.A., Parcharidis, I. 2021. Primary and Secondary Environmental Effects  
877 Triggered by the 30 October 2020, Mw = 7.0, Samos (Eastern Aegean Sea, Greece)  
878 Earthquake Based on Post-Event Field Surveys and InSAR Analysis. Appl. Sci., 11, 3281.  
879 <https://doi.org/10.3390/app11073281>

880 McClusky, S., et al. 2000, Global Positioning System constraints on plate kinematics and  
881 dynamics in the eastern Mediterranean and Caucasus, J. Geophys. Res. 105, B3, 5695-5719,  
882 <http://doi.org/10.1029/1999JB900351>

883 McKenzie, D. 1978. Active tectonics of the Alpine–Himalayan belt: the Aegean Sea and  
884 surrounding regions, Geophys. J. Roy. Astr. Soc. 55, 1, 217-254,  
885 <http://doi.org/10.1111/j.1365-246X.1978.tb04759.x>.

886 Melgar, D., B. W. Crowell, J. Geng, R. M. Allen, Y. Bock, S. Riquelme, E. M. Hill, M. Protti,  
887 and A. Ganas, 2015. Earthquake magnitude calculation without saturation from the scaling  
888 of peak ground displacement, Geophys. Res. Lett., 42,  
889 <http://onlinelibrary.wiley.com/doi/10.1002/2015GL064278/full> .

890 Melgar, D., A. Ganas, J. Geng, C. Liang, E. J. Fielding, and I. Kassaras, 2017. Source  
891 characteristics of the 2015 Mw6.5 Lefkada, Greece, strike-slip earthquake, J. Geophys. Res.  
892 Solid Earth, 122, <http://dx.doi.org/10.1002/2016JB013452>

893 Melgar, D. A. Ganas, T. Taymaz, S. Valkaniotis, B. W Crowell, V. Kapetanidis, V. Tsironi, S.  
894 Yolsal-Çevikbilen, T. Öcalan, 2020. Rupture kinematics of 2020 January 24 Mw 6.7  
895 Doğanyol-Sivrice, Turkey earthquake on the East Anatolian Fault Zone imaged by space  
896 geodesy, Geophysical Journal International, Volume 223, Issue 2, Pages 862–874,  
897 <https://doi.org/10.1093/gji/ggaa345>

898 Melis, N.S., Okal, E.A., Synolakis, C.E. et al. 2020. The Chios, Greece Earthquake of 23 July  
899 1949: Seismological Reassessment and Tsunami Investigations. Pure Appl. Geophys. 177,  
900 1295–1313. <https://doi.org/10.1007/s00024-019-02410-1>

901 Meng, J., Sinoplu, O., Zhou, Z. et al. 2021. Greece and Turkey Shaken by African tectonic  
902 retreat. Sci Rep 11, 6486. <https://doi.org/10.1038/s41598-021-86063-y>

903 Moro, M., M. Chini, M. Saroli, S. Atzori, S. Stramondo, and S. Salvi, 2011. Analysis of large,  
904 seismically induced, gravitational deformations imaged by high-resolution COSMO-SkyMed  
905 synthetic aperture radar, *Geology*, 39(6), 527–530, <https://doi.org/10.1130/G31748.1>

906 Mountrakis, D., Kiliyas, A., Vavliakis E., Psilovikos, A. Sc Thomaidou, E., 2003. Neotectonic  
907 map of Samos island (Aegean Sea, Greece): Implication of geographical information  
908 systems in the geological mapping. — *Proceedings Volume 1, 4<sup>th</sup> European Congress on*  
909 *Regional Geoscientific Cartography and Information Systems*, 11-13, Bologna, Italy.

910 Neo, J.C. Huang, Y. Yao, D. and Wei, S., 2020. Is the Aftershock Zone Area a Good Proxy for  
911 the Mainshock Rupture Area?. *Bulletin of the Seismological Society of America*, 111 (1):  
912 424–438. doi: <https://doi.org/10.1785/0120190200>

913 NOAFAULTs KMZ layer Version 2.1 (2019 update) (Version V2.1) [Data set]. Zenodo.  
914 <http://doi.org/10.5281/zenodo.3483136>

915

916 Nomikou, P., Evangelidis, D., Papanikolaou, D., Lampridou, D., Litsas, D., Tsaparas, Y.,  
917 Koliopoulos, I. 2021. Morphotectonic Analysis along the Northern Margin of Samos Island,  
918 Related to the Seismic Activity of October 2020, Aegean Sea, Greece. *Geosciences*, 11, 102.  
919 <https://doi.org/10.3390/geosciences11020102>

920 Ocakoglu, N., Demirbag, E. & Kuscu, I. 2004. Neotectonic structures in the area offshore of  
921 Alaçatı, Doğanbey and Kuşadası (western Turkey): evidence of strike-slip faulting in the  
922 Aegean extensional province. *Tectonophysics* 391, 67–83.

923 Okada, Y. 1985. Surface deformation due to shear and tensile faults in a half-space. *Bull*  
924 *Seismol Soc Am*, 75:1135-1154.

925 Okrusch, M., P. Richter and G. Katsikatsos, 1984. High-pressure rocks of Samos, Greece.  
926 *Geological Society, London, Special Publications*, 17, 529-536,  
927 <https://doi.org/10.1144/GSL.SP.1984.017.01.39>

928 Ott, R., Wegmann, K.W., Gallen, S.F., Pazzagila, F.J., Brandon, M., Ueda, K., Fassoulas, C.,  
929 2021. Reassessing Eastern Mediterranean tectonics and earthquake hazards from the AD  
930 365 earthquake: AGU Advances, in press.

931 Özkaymak, C., and Sözbilir, H., 2012. Tectonic geomorphology of the Spildağı High Ranges,  
932 western Anatolia, *Geomorphology*, 173–174, 128-140.

933 Papadimitriou P., Kapetanidis V., Karakonstantis A., Spingos I., Kassaras I., Sakkas V.,  
934 Kouskouna V., Karatzetzou A., Pavlou K., Kaviris G. and Voulgaris N. 2020. First Results on  
935 the Mw=6.9 Samos Earthquake of 30 October 2020. *Bulletin Geological Society of Greece*,  
936 56, 251-279 <https://doi.org/10.12681/bgsg.25359>

- 937 Papadopoulos, G. A., Latoussakis, I., Daskalaki, E., Diakogianni, G., Fokaefs, A., Kolligri, M.,  
938 Liadopoulou, K., Orfanogiannaki, K., and Pirentis, A. 2006. The East Aegean Sea strong  
939 earthquake sequence of October–November 2005: lessons learned for earthquake  
940 prediction from foreshocks, *Nat. Hazards Earth Syst. Sci.*, 6, 895–901,  
941 <https://doi.org/10.5194/nhess-6-895-2006>.
- 942 Papanikolaou D. 1979. Unités tectoniques et phases de déformation dans l'île de Samos,  
943 Mer Egée, Grèce. *Bull. Soc. Géol. France*, 21, 745–752.
- 944 Papazachos, B. C., Mountrakis, D. M., Papazachos, C. B., Tranos, M. D., Karakaisis, G. F., and  
945 Savvaidis, A. S., 2001. The faults which have caused the known major earthquakes in Greece  
946 and surrounding region between the 5<sup>th</sup> century BC and today. In *Proceedings of 2nd*  
947 *National Conference Anti-Seismic Engineering and Technical Seismology*, 17–26, Technical  
948 Chamber of Greece, Thessaloniki, Greece.
- 949 Pousse-Beltran, L., Socquet, A., Benedetti, L., Doin, M.-P., Rizza, M., & D'Agostino, N. 2020.  
950 Localized afterslip at geometrical complexities revealed by InSAR after the 2016 Central  
951 Italy seismic sequence. *Journal of Geophysical Research: Solid Earth*, 125, e2019JB019065.  
952 <https://doi.org/10.1029/2019JB019065>
- 953 Reilinger, R. S. McClusky, D. Paradissis, S. Ergintav, P. Vernant, 2010. Geodetic constraints  
954 on the tectonic evolution of the Aegean region and strain accumulation along the Hellenic  
955 subduction zone, *Tectonophysics*, Volume 488, Issues 1–4, Pages 22-30,  
956 <https://doi.org/10.1016/j.tecto.2009.05.027>.
- 957 Ring, U., S. Laws, M. Bernet, 1999. Structural analysis of a complex nappe sequence and  
958 late-orogenic basins from the Aegean Island of Samos, Greece, *Journal of Structural*  
959 *Geology*, 21, 11, 1575-1601, [https://doi.org/10.1016/S0191-8141\(99\)00108-X](https://doi.org/10.1016/S0191-8141(99)00108-X).
- 960 Robertson, J., Meschis, M., Roberts, G.P, Ganas A., Gheorgiou, D., 2019. Temporally  
961 constant Quaternary uplift rates and their relationship with extensional upper-plate faults  
962 in south Crete (Greece), constrained with <sup>36</sup>Cl exposure dating. *Tectonics*. 38 (4), 1189-  
963 1222. <https://doi.org/10.1029/2018TC005410>
- 964 Ruhl, C.J., D. Melgar, J. Geng, DE Goldberg, BW Crowell, RM Allen, Y. Bock, S. Barrientos, S.  
965 Riquelme, JC Baez, E. Cabral-Cano, X. Pérez-Campos, EM Hill, M. Protti, A. Ganas, M. Ruiz,  
966 P. Mothes, P. Jarrín, JM Nocquet, JP Avouac, E. D'Anastasio, 2019. A Global Database of  
967 Strong-Motion Displacement GNSS Recordings and an Example Application to PGD Scaling,  
968 *Seismological Research Letters*, 90 (1), 271-279 <https://doi.org/10.1785/0220180177>
- 969 Sakellariou, D., Mascle, J., & Lykousis, V. 2013. Strike slip tectonics and transtensional  
970 deformation in the Aegean region and the Hellenic arc: Preliminary results. *Bulletin of the*  
971 *Geological Society of Greece*, 47, 647-656 <http://dx.doi.org/10.12681/bgsg.11098>



972 Stiros, S.C., J. Laborel, F. Laborel-Deguen, S. Papageorgiou, J. Evin, P.A. Pirazzoli, 2000.  
 973 Seismic coastal uplift in a region of subsidence: Holocene raised shorelines of Samos Island,  
 974 Aegean Sea, Greece, *Marine Geology*, Volume 170, Issues 1–2, 41-58,  
 975 [https://doi.org/10.1016/S0025-3227\(00\)00064-5](https://doi.org/10.1016/S0025-3227(00)00064-5) .

976 Tan, O., Papadimitriou, E.E., Pabucçu, Z. et al. 2014. A detailed analysis of microseismicity  
 977 in Samos and Kusadasi (Eastern Aegean Sea) areas. *Acta Geophys.* 62, 1283–1309.  
 978 <https://doi.org/10.2478/s11600-013-0194-1>

979 Taymaz, T., J. Jackson, and D. McKenzie 1991. Active tectonics of the north and central  
 980 Aegean Sea, *Geophys. J. Int.* 106, 2, 433-490, DOI: 10.1111/j.1365-246X.1991.tb03906.x

981 Theodoropoulos, D. 1979. Geological Map of Greece, 1:50,000 scale; Neon Karlovasi and  
 982 Limin Vatheos Sheets. IGME, Athens, Greece.

983 Triantafyllou, I.; Gogou, M.; Mavroulis, S.; Lekkas, E.; Papadopoulos, G.A.; Thravalos, M.  
 984 2021. The Tsunami Caused by the 30 October 2020 Samos (Aegean Sea) Mw7.0 Earthquake:  
 985 Hydrodynamic Features, Source Properties and Impact Assessment from Post-Event Field  
 986 Survey and Video Records. *J. Mar. Sci. Eng.* 9, 68. <https://doi.org/10.3390/jmse9010068>

987 Vadaloukas, G., Vintzileou, E., Ganas, A., Giarlelis, C., Ziotopoulou, K., Theodoulidis, N.,  
 988 Karasante, I., Margaris, B., Mylonakis, G., Papachristidis, A., Repapis, C., Psarropoulos, P.  
 989 N., & Sextos, A. G. 2020. Samos Earthquake of 30th October, 2020. Preliminary Report of  
 990 the Hellenic Association for Earthquake Engineering, Athens, Greece.  
 991 <https://doi.org/10.13140/RG.2.2.22609.76644>

992 Vernant, Philippe; Robert Reilinger, Simon McClusky, 2014. Geodetic evidence for low  
 993 coupling on the Hellenic subduction plate interface, *Earth and Planetary Science Letters*,  
 994 Volume 385, Pages 122-129, <https://doi.org/10.1016/j.epsl.2013.10.018> .

995 Wessel, P., Luis, J. F., Uieda, L., Scharroo, R., Wobbe, F., Smith, W. H. F., & Tian, D. 2019.  
 996 The Generic Mapping Tools version 6. *Geochemistry, Geophysics, Geosystems*, 20, 5556–  
 997 5564.

998 Yolsal-Çevikbilen, S., Taymaz, T., and Helvacı, C. 2014. Earthquake mechanisms in the Gulfs  
 999 of Gökova, Siğacık, Kuşadası, and the Simav Region (western Turkey): Neotectonics,  
 1000 seismotectonics and geodynamic implications, *Tectonophysics*, 635, 100-124,  
 1001 <https://doi.org/10.1016/j.tecto.2014.05.001>.

1002

1003

1004 **Supplementary material**

1005 **Text S1.**

AFAD

<https://depem.afad.gov.tr/downloadDocument?id=2064>

GFZ

<http://geofon.gfz-potsdam.de/eqinfo/event.php?id=gfz2020vimx>

GCMT

<https://www.globalcmt.org/CMTsearch.html>

INGV

<http://autorcmt.bo.ingv.it/quicks.html>

IPGP

<http://geoscope.ipgp.fr/index.php/en/catalog/earthquake-description?seis=us7000c7y0>

KOERI

<http://udim.koeri.boun.edu.tr/zeqmap/osmapen.asp>

UOA

[http://www.geophysics.geol.uoa.gr/stations/gmaps3/eventpage\\_leaf.php?scid=nkua2020vimx&lng=en](http://www.geophysics.geol.uoa.gr/stations/gmaps3/eventpage_leaf.php?scid=nkua2020vimx&lng=en)

NOA

<http://bbnet.gein.noa.gr/HL/seismicity/mts/revise-moment-tensors>

OCA

[http://sismoazur.oca.eu/focal\\_mechanism\\_emsc?start=2020-10-29&end=2020-10-31&minlat=0.00&minlon=0.00&maxlat=38.00&maxlon=27.00&mindepth=0&maxdepth=20&minmag=5&maxmag=7.5&eventtype=all](http://sismoazur.oca.eu/focal_mechanism_emsc?start=2020-10-29&end=2020-10-31&minlat=0.00&minlon=0.00&maxlat=38.00&maxlon=27.00&mindepth=0&maxdepth=20&minmag=5&maxmag=7.5&eventtype=all)

USGS

<https://earthquake.usgs.gov/earthquakes/eventpage/us7000c7y0/executive>

1006

1007

1008 **Text S2:**

1009 AIPS (<http://aips.space.noa.gr>) is a platform for routine (updated in every new acquisition)  
1010 and on-demand SENTINEL-1 interferogram processing service. It provides freely the  
1011 original, the filtered, using the adaptive power spectrum filter of Goldstein and Werner  
1012 (1998) with a coherence threshold of 0.3, interferograms as well as its amplitude in  
1013 UTM/WGS Transverse Mercator projection, at full resolution (25 m), in GEOTIFF format. It  
1014 is based on ESA's SNAP (<http://step.esa.int/main/>) software version 6.0. The multiblock  
1015 factor is 6 in range and 2 in azimuth. Moreover, it produces figures and raw images in PNG  
1016 format as quick-look products. The digital elevation model (DEM) used for the processing  
1017 is the Shuttle Radar Topography Mission (SRTM) 1 Arc-Second Global (doi number:  
1018 /10.5066/F7PR7TFT). During the stage of processing, the interferogram was formed by  
1019 cross-multiplying the post-event acquisition image) with the complex conjugate of the pre-  
1020 event image). The resulting phase represents the difference between the two images.  
1021 Currently the routine processing includes four tracks of central Greece and one-track form  
1022 Etna in Sicily, Italy but there is the flexibility to extend the processing upon request. It keeps  
1023 a rolling-archive with old data stored offline. The on-demand service is triggered upon a  
1024 moderate or big earthquake in the Euro-Mediterranean region. For the case of the Samos  
1025 earthquake the interferometric products can be found at [http://aips-  
p.space.noa.gr/Samos/products/](http://aips-<br/>1026 p.space.noa.gr/Samos/products/)

1027

1028

1029 **Table S1.**

1030 List of the pre-seismic, co-seismic and post-seismic interferograms of the track 131, 29  
 1031 and 36, produced by AIPS and stored at <http://aips-p.space.noa.gr/Samos/products/>. The  
 1032 interferograms in bold and framed are used to produce the average interferograms of  
 1033 track 29 and 36 that were used in the inversion modelling.

Ascending track 131			Ascending track 29		Descending track 36		
pre-seismic	co-seismic	post-seismic	co-seismic	post-seismic	pre-seismic	co-seismic	post-seismic
20201018- 20201024	20201018-	20201117-	20201017-	20201110-	20201024- 20201030	<b>20201024-</b>	20201111-
	20201111	20201123	20201104	20201116		<b>20201111</b>	20201123
	20201024-	20201030-	<b>20201023-</b>	20201104-		<b>20201006-</b>	20201105-
	20201030	20201105	<b>20201116</b>	20201110		<b>20201111</b>	20201123
	20201018-	20201105-	20201029-	20201110-		20201030-	20201105-
	20201030	20201123	20201122	20201122		20201111	20201117
	20201024-	20201030-	<b>20201023-</b>	20201116-		<b>20201024-</b>	20201111-
	20201123	20201117	<b>20201110</b>	20201122		<b>20201117</b>	20201117
	20201018-	20201030-	<b>20201011-</b>	20201104-		20201006-	20201105-
	20201105	20201123	<b>20201110</b>	20201116		20201105	20201111
	20201018-	20201111-	20201029-	20201104-		20201018-	20201117-
	20201117	20201117	20201104	20201122		20201105	20201123
	20201006-	20201111-	20201023-			20201012-	
	20201111	20201123	20201122			20201117	
	20201024-	20201030-	20201017-			20201012-	
	20201117	20201111	20201110			20201105	
	20201018-	20201105-	20201011-			20201018-	
	20201123	20201111	20201104			20201123	
	20201012-	20201105-	20201029-			20201018-	
	20201030	20201117	20201116			20201117	
	20201024-		20201023-			20201030-	
	20201111		20201104			20201105	
	20201006-		20201017-			20201030-	
	20201105		20201116			20201123	
	20201024-		<b>20201011-</b>			20201030-	
	20201105		<b>20201116</b>			20201117	
	20201012-		20201029-			20201024-	
	20201117		20201110			20201117	
	20201006-		20201017-			20201024-	
	20201030		20201122			20201123	
20201012-				<b>20201024-</b>			
20201111				<b>20201123</b>			
20201012-				20201018-			
20201105				20201111			

1034

1035

1036

1037

1038

1039 **Table S2.** List of GNSS displacement data (mm; E, N, Up components) and their  
 1040 uncertainties.

			dE	dN	dU	sE	sN	sU		
	<b>SAMO</b>	26.7053	37.7927	-30	-356	82	2	2	4	
	<b>SAMU</b>	26.9734	37.7575	-5	-64	21	2	2	4	
	<b>093A</b>	26.8334	37.6716	22	-152	23	3	3	6	
	<b>CHIO</b>	26.1272	38.3679	-8	17	2	3	3	6	
	<b>CHIU</b>	26.136	38.3665	-8	18	5	2	2	4	
	<b>IKAR</b>	26.2243	37.6282	-7	-31	7	3	3	6	
coseismic	<b>IKAU</b>	26.2733	37.6055	-4	-48	4	3	3	6	
	<b>KALU</b>	26.9617	36.9624	3	-14	-3	2	2	4	
	<b>LERO</b>	26.8545	37.1364	3	-20	-	2	2	-	
	<b>KALY</b>	26.9762	36.9558	3	-11	-1	3	3	6	
	<b>AYD1</b>	27.8378	37.8407	0	7	-	8	8	-	
	<b>CESM</b>	26.3726	38.3038	-12	54	-	8	8	-	
	<b>IZMI</b>	27.0819	38.3948	10	30	2	2	2	4	
	<b>DIDI</b>	27.2687	37.3721	2	-5	-	8	8	-	
	<hr/>									
		<b>SAMO</b>	26.7053	37.7927	-41	-400	100	2	2	4
		<b>SAMU</b>	26.9734	37.7575	-5	-80	28	2	2	4
		<b>CHIO</b>	26.1272	38.3679	-8	22	2	3	3	6
		<b>CHIU</b>	26.136	38.3665	-10	23	5	2	2	4
	<b>IKAR</b>	26.2243	37.6282	-5	-41	7	3	3	6	
Coseismic + 10 days	<b>IKAU</b>	26.2733	37.6055	-14	-50	4	3	3	6	
	<b>KALU</b>	26.9617	36.9624	3	-14	-3	2	2	4	
postseismic	<b>LERO</b>	26.8545	37.1364	3	-11	-	2	2	-	
	<b>KALY</b>	26.9762	36.9558	3	-23	-1	3	3	6	
	<b>AYD1</b>	27.8378	37.8407	0	7	-	8	8	-	
	<b>CESM</b>	26.3726	38.3038	-12	54	-	8	8	-	
	<b>IZMI</b>	27.0819	38.3948	12	36	2	2	2	4	
	<b>DIDI</b>	27.2687	37.3721	2	-5	-	8	8	-	

1041

1042



1043 **Table S3.**

1044 List of the picked LOS values from the averaged interferograms. The deformation  
 1045 direction for the LOS. is at the notion of pre-event minus post-event, thus values  
 1046 increasing towards satellite. Lon: Longitude (°), lat is latitude (°). InSAR and modelled  
 1047 values are in mm. The picks are displayed in **Fig. S6**.

1048

Pick No	ASCENDING track 29				DESCENDING track 36			
	lon	lat	InSAR	model	lon	lat	InSAR	model
1	26.644	37.772	114	112.6	26.761	37.678	28	31.6
2	26.668	37.769	114	110.7	26.800	37.688	28	38.7
3	26.695	37.768	114	111.3	26.825	37.695	28	43.1
4	26.717	37.769	114	112.8	26.852	37.706	28	48.8
5	26.735	37.775	114	118.5	26.868	37.722	28	55.5
6	26.745	37.784	114	128.7	26.861	37.752	28	68.5
7	26.751	37.796	114	134.9	26.842	37.766	28	72.4
8	26.746	37.808	114	108.5	26.826	37.777	28	75.0
9	26.740	37.812	114	69.7	26.813	37.791	28	75.7
10	26.651	37.705	86	63.6	26.798	37.806	28	39.6
11	26.668	37.716	86	69.3	26.807	37.806	0	35.9
12	26.690	37.728	86	76.0	26.818	37.796	0	69.6
13	26.713	37.738	86	82.4	26.865	37.776	0	75.6
14	26.736	37.748	86	90.2	26.882	37.771	0	74.5
15	26.751	37.757	86	98.0	26.902	37.768	0	68.3
16	26.766	37.766	86	106.9	26.586	37.748	28	36.3
17	26.776	37.776	86	119.4	26.621	37.728	28	31.4
18	26.777	37.790	86	132.7	26.642	37.700	28	23.6
19	26.775	37.801	86	127.3	26.841	37.783	0	75.1
20	26.762	37.809	86	93.7	26.977	37.766	0	29.2
21	26.755	37.811	86	71.8	26.992	37.758	0	27.7
22	26.758	37.691	58	46.0	27.006	37.754	0	25.6
23	26.778	37.699	58	46.9	27.026	37.749	0	22.4
24	26.797	37.707	58	47.3	26.815	37.808	-28	-5.4
25	26.815	37.715	58	48.2	26.833	37.795	-28	66.5
26	26.834	37.723	58	47.9	26.851	37.787	-28	72.6
27	26.852	37.731	58	47.5	26.871	37.786	-28	71.7
28	26.863	37.742	58	50.9	26.920	37.769	-28	58.5
29	26.863	37.758	58	66.6	27.042	37.746	0	20.2
30	26.849	37.773	58	96.6	26.824	37.805	-56	26.9
31	26.826	37.783	58	119.9	26.839	37.799	-56	54.2
32	26.810	37.790	58	129.7	26.857	37.795	-56	58.3
33	26.797	37.798	58	129.2	26.959	37.772	0	32.6
34	26.784	37.809	58	62.9				
35	26.782	37.659	30	29.2				
36	26.799	37.652	30	24.8				
37	26.810	37.648	30	22.2				

38	26.827	37.647	30	19.5
39	26.851	37.647	30	16.1
40	26.867	37.652	30	14.7
41	26.882	37.659	30	13.9
42	26.965	37.696	30	11.3
43	26.986	37.707	30	12.6
44	26.997	37.717	30	14.3
45	27.000	37.729	30	17.1
46	26.993	37.742	30	21.0
47	26.989	37.758	30	27.2
48	26.980	37.772	30	34.3
49	26.965	37.780	30	41.3
50	26.949	37.782	30	45.1
51	26.875	37.783	30	93.6
52	26.853	37.784	30	111.5
53	26.834	37.790	30	125.5
54	26.822	37.795	30	128.4
55	26.806	37.802	30	113.0
56	26.797	37.809	30	60.1
57	26.860	37.794	2	117.4
58	26.842	37.796	2	121.9
59	26.832	37.799	2	119.8
60	26.821	37.803	2	106.2
61	26.813	37.808	2	52.8
62	26.856	37.802	-26	91.1
63	26.846	37.802	-26	96.9
64	26.834	37.804	-26	84.4

1049

1050

1051 **Table S4.** List with locations of field surveyed earthquake effects.

Date	LAT	LON	Location	Feature
06-11-20	37.782800	26.952165	NW Vathy	building damage
06-11-20	37.782800	26.952170	NW Vathy	building damage
06-11-20	37.782865	26.951988	NW Vathy	building damage
06-11-20	37.770015	26.967010	Vathy	Landslide
06-11-20	37.753943	26.976883	Vathy	building damage
06-11-20	37.753864	26.976861	Vathy	building damage
06-11-20	37.753657	26.976442	Vathy	cracks
06-11-20	37.753572	26.976501	Vathy	cracks
06-11-20	37.753573	26.976502	Vathy	cracks
06-11-20	37.751111	26.959167	Kokkari	Camp
06-11-20	37.796880	26.857543	Avlakia	Landslide
06-11-20	37.797113	26.857372	Avlakia	Landslide
06-11-20	37.796700	26.857658	Avlakia	Landslide
06-11-20	37.796515	26.857833	Avlakia	Landslide
06-11-20	37.809698	26.791153	Livadaki	road cracks
06-11-20	37.809697	26.791152	Livadaki	road cracks
06-11-20	37.809696	26.791151	Livadaki	road cracks
06-11-20	37.809647	26.790563	Livadaki	road cracks
06-11-20	37.809684	26.791061	Livadaki	road cracks
06-11-20	37.809584	26.790536	Livadaki	road protection blocks
06-11-20	37.792673	26.705121	Karlovasi	School damage
06-11-20	37.792722	26.705295	Karlovasi	GNSS Antenna SAMO
06-11-20	37.792741	26.705278	Karlovasi	GNSS Antenna SAMO
06-11-20	37.792778	26.705278	Karlovasi	GNSS Antenna SAMO
06-11-20	37.790618	26.704542	Karlovasi	Church Theotokou
06-11-20	37.790800	26.704401	Karlovasi	Church Theotokou
06-11-20	37.790339	26.704289	Karlovasi	building damage – yellow mark
06-11-20	37.790438	26.704197	Karlovasi	Church Theotokou
06-11-20	37.790525	26.703200	Karlovasi	Church Theotokou
06-11-20	37.790524	26.703200	Karlovasi	Church Theotokou
06-11-20	37.794683	26.680910	Karlovasi	Port
06-11-20	37.795409	26.678122	Karlovasi	Jetty - view coastline
06-11-20	37.796130	26.678811	Karlovasi	Jetty - corner - crack
06-11-20	37.796130	26.678809	Karlovasi	Jetty - corner
06-11-20	37.796110	26.678809	Karlovasi	Jetty - corner - crack
06-11-20	37.796308	26.678824	Karlovasi	Jetty - centre -crack
06-11-20	37.796308	26.678824	Karlovasi	Jetty - centre - liquefaction
06-11-20	37.796589	26.678989	Karlovasi	Jetty - centre -crack
06-11-20	37.796635	26.678836	Karlovasi	Jetty wall -crack
06-11-20	37.796695	26.678981	Karlovasi	Jetty - corner - pop-up
06-11-20	37.796694	26.678981	Karlovasi	Jetty - corner - pop-up
06-11-20	37.798332	26.682602	Karlovasi	Jetty-crack
06-11-20	37.798235	26.682622	Karlovasi	Jetty-crack
07-11-20	37.750716	26.961429	Vathy	road cracks
07-11-20	37.750719	26.961334	Vathy	road cracks

07-11-20	37.773889	26.897222	Kokkari	collapse
07-11-20	37.778270	26.884915	Kokkari	collapse
07-11-20	37.796413	26.857920	Avlakia	Landslide
07-11-20	37.796610	26.857723	Avlakia	Landslide
07-11-20	37.798812	26.849460	Road to Vourliotes	collapse
07-11-20	37.799737	26.852413	Road to Vourliotes	Fault
07-11-20	37.799925	26.852378	Road to Vourliotes	Fault
07-11-20	37.803827	26.826588	Agios Konstantinos	building damage
07-11-20	37.803861	26.826472	Agios Konstantinos	building damage
07-11-20	37.799722	26.851667	Agios Konstantinos	building damage
07-11-20	37.799722	26.851667		Raised beach
07-11-20	37.799722	26.851667		Raised beach
07-11-20	37.810219	26.771204	Petalides	raised coastline
07-11-20	37.809229	26.743253	Agios Nikolaos	road cracks
07-11-20	37.809187	26.742978	Agios Nikolaos	road cracks
07-11-20	37.803605	26.724465	Kontakaiika north	road cracks
07-11-20	37.803592	26.724498	Kontakaiika north	road cracks
07-11-20	37.803588	26.724529	Kontakaiika north	road cracks
07-11-20	37.803583	26.724529	Kontakaiika north	road cracks
07-11-20	37.803582	26.724528	Kontakaiika north	road cracks
07-11-20	37.793980	26.670743	Potami	raised coastline
07-11-20	37.793825	26.670602	Potami	raised coastline
07-11-20	37.792297	26.668608	Potami	Raised beach
07-11-20	37.792455	26.668855	Potami	salt pans
07-11-20	37.792457	26.668863	Potami	raised beach
07-11-20	37.792517	26.668987	Potami	beach rock
07-11-20	37.798747	26.706262	Karlovasi	collapse
07-11-20	37.800574	26.709502	Karlovasi	collapse
07-11-20	37.800532	26.709530	Karlovasi	collapse
07-11-20	37.801080	26.708335	Karlovasi beach front	collapse
07-11-20	37.801087	26.708115	Karlovasi beach front	collapse
07-11-20	37.800833	26.708056	Karlovasi beach front	collapse
07-11-20	37.800385	26.705687	Karlovasi beach front	view to the east
07-11-20	37.800306	26.705436	Karlovasi beach front	collapse
07-11-20	37.800325	26.705346	Karlovasi beach front	collapse
07-11-20	37.773310	26.723260°	road to Kontaiika	collapse
07-11-20	37.767487	26.725427°	road to Kontaiika	wall damage (fence)
07-11-20	37.767429	26.725481	road to Kontaiika	wall damage (fence)
07-11-20	37.756800	26.738886	Kontaiika	collapse
08-11-20	37.753122	26.95935	Vathy new port	jetty damage
08-11-20	37.753122	26.95935	Vathy new port	jetty damage
08-11-20	37.753122	26.95935	Vathy new port	jetty damage
08-11-20	37.753122	26.95935	Vathy new port	jetty damage
08-11-20	37.756406	26.957641	Malagari	Building damage
08-11-20	37.756828	26.957874	Malagari	soil liquefaction
08-11-20	37.756853	26.958147	Malagari	road cracks
08-11-20	37.756729	26.958311	Malagari	tree line (submerged)
08-11-20	37.756804	26.957965	Malagari	soil liquefaction

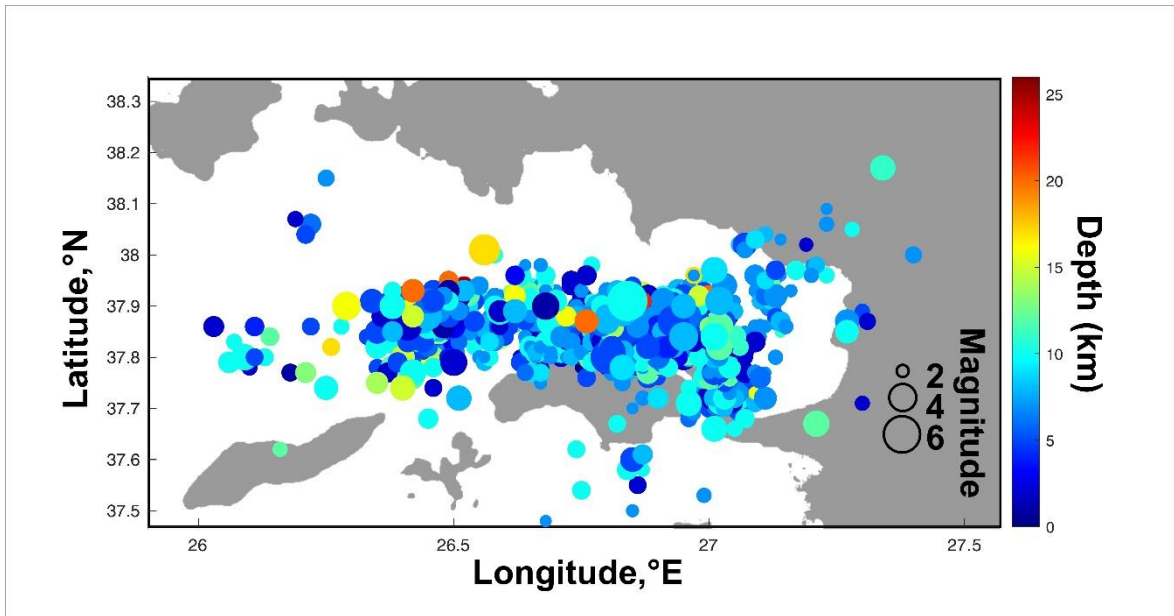
08-11-20	37.756803	26.957965	Malagari	soil liquefaction
08-11-20	37.756827	26.957926	Malagari	soil liquefaction
08-11-20	37.756827	26.957924	Malagari	soil liquefaction
08-11-20	37.756729	26.958313	Malagari	tree line submerged
08-11-20	37.756729	26.958314	Malagari	tree line submerged
08-11-20	37.756728	26.958314	Malagari	tree line submerged
08-11-20	37.756727	26.958277	Malagari	ground offset
08-11-20	37.756727	26.958276	Malagari	ground offset
08-11-20	37.756729	26.958276	Malagari	ground offset
08-11-20	37.796431	26.678894	Karlovasi jetty	cracks
08-11-20	37.79643	26.678893	Karlovasi jetty	cracks
08-11-20	37.796433	26.67889	Karlovasi jetty	cracks
08-11-20	37.796432	26.678889	Karlovasi jetty	cracks
08-11-20	37.79643	26.678885	Karlovasi jetty	cracks
08-11-20	37.796432	26.678882	Karlovasi jetty	cracks
08-11-20	37.795377	26.678302	Karlovasi shipyard	tsunami upper limit
08-11-20	37.795377	26.678302	Karlovasi shipyard	tsunami upper limit
08-11-20	37.795154	26.678336	Karlovasi shipyard	tsunami extent limit

1052

1053



1054 **Fig. S1:** Seismicity map (EMSC data) showing both magnitude and depth distribution of  
1055 events.

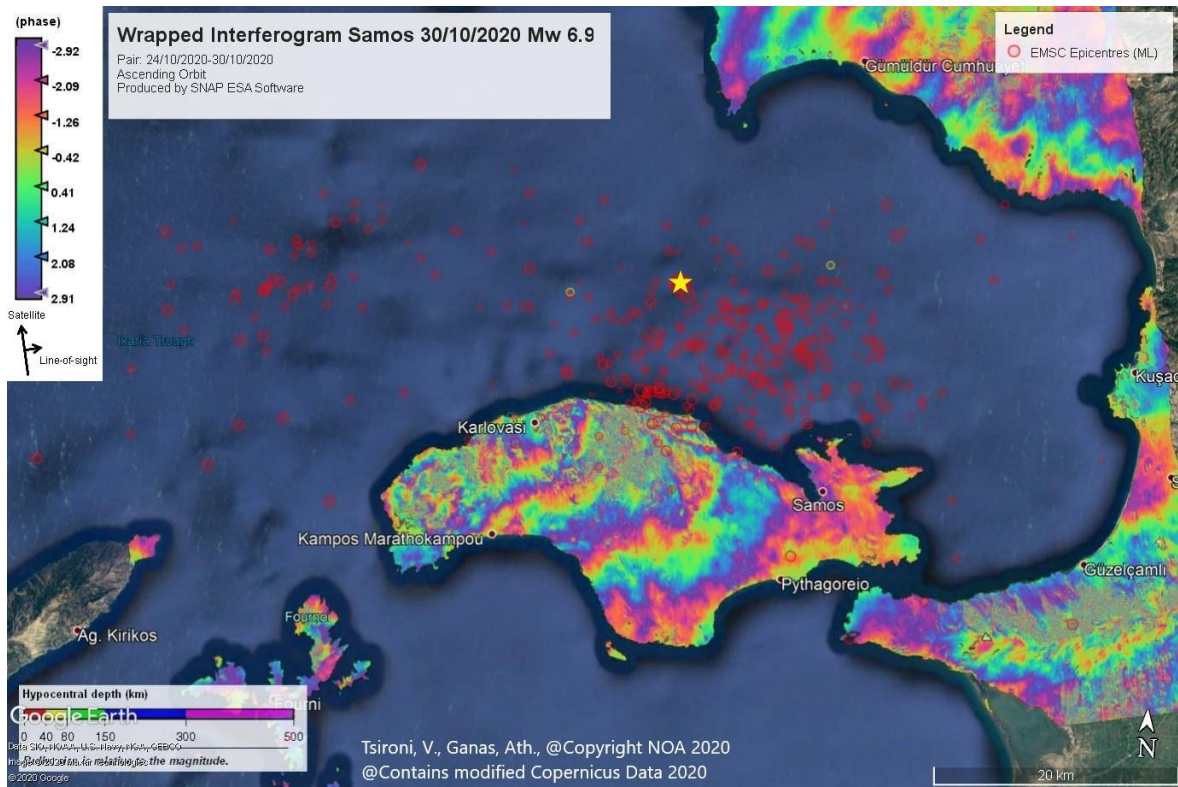


1056

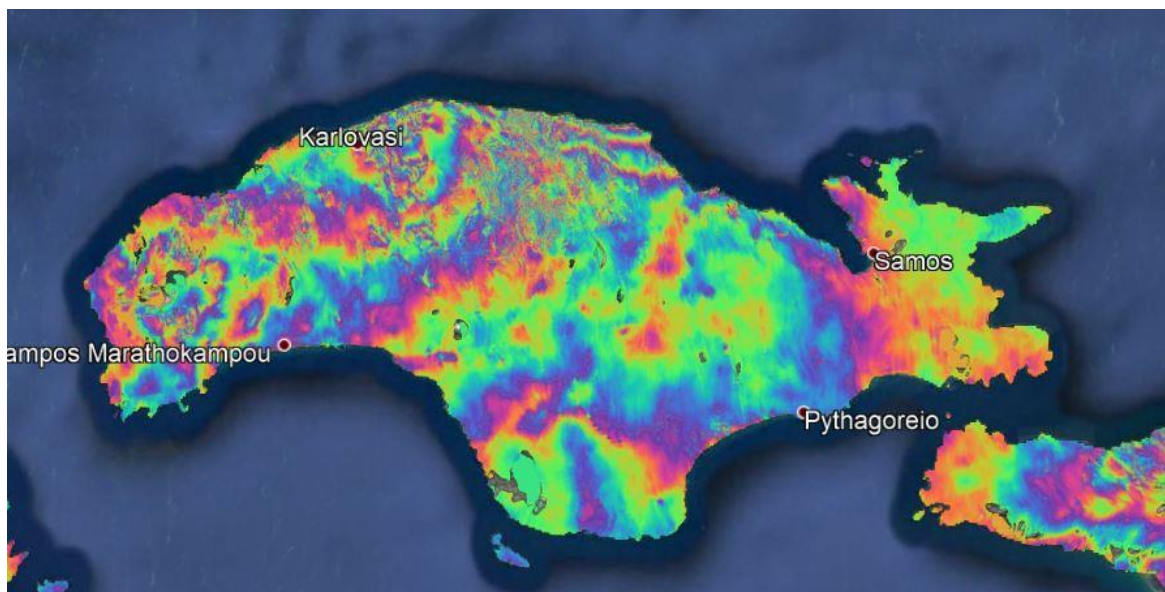
1057

1058

1059 **Fig. S2:** The initial co-seismic differential interferograms (wrapped phase; cropped swath  
 1060 of ascending track 131; Ganas et al. 2020a) over Samos island, produced exploiting the  
 1061 Geohazards Exploitation Platform. Top panel shows the image pair 24 Oct. – 30 Oct., 2020.  
 1062 Bottom panel: 18 Oct. – 30 Oct., 2020. The fringes in the differential interferogram show a  
 1063 pattern that supports more the north-dipping plane as the density and the curvature of the  
 1064 fringes indicates. These data show that Samos moved towards the satellite in the LOS with  
 1065 the exception of a coastal strip from Agios Konstantinos to Kokkari that moved away. The  
 1066 interferograms are draped over Google Earth imagery.



1067



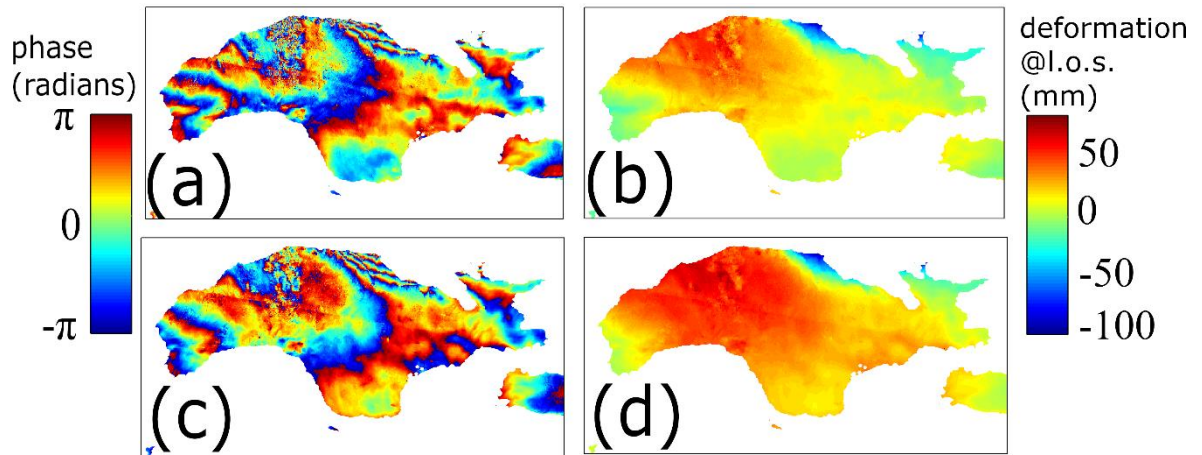
1068

1069 **Fig. S3:** (a) Best interferogram of Track 36 (D) spanning the dates 10 June – 11 November  
1070 2020, (b) the unwrapped one, (c) the average interferogram of the best four and (d) the  
1071 unwrapped one.

1072

1073

1074



1075

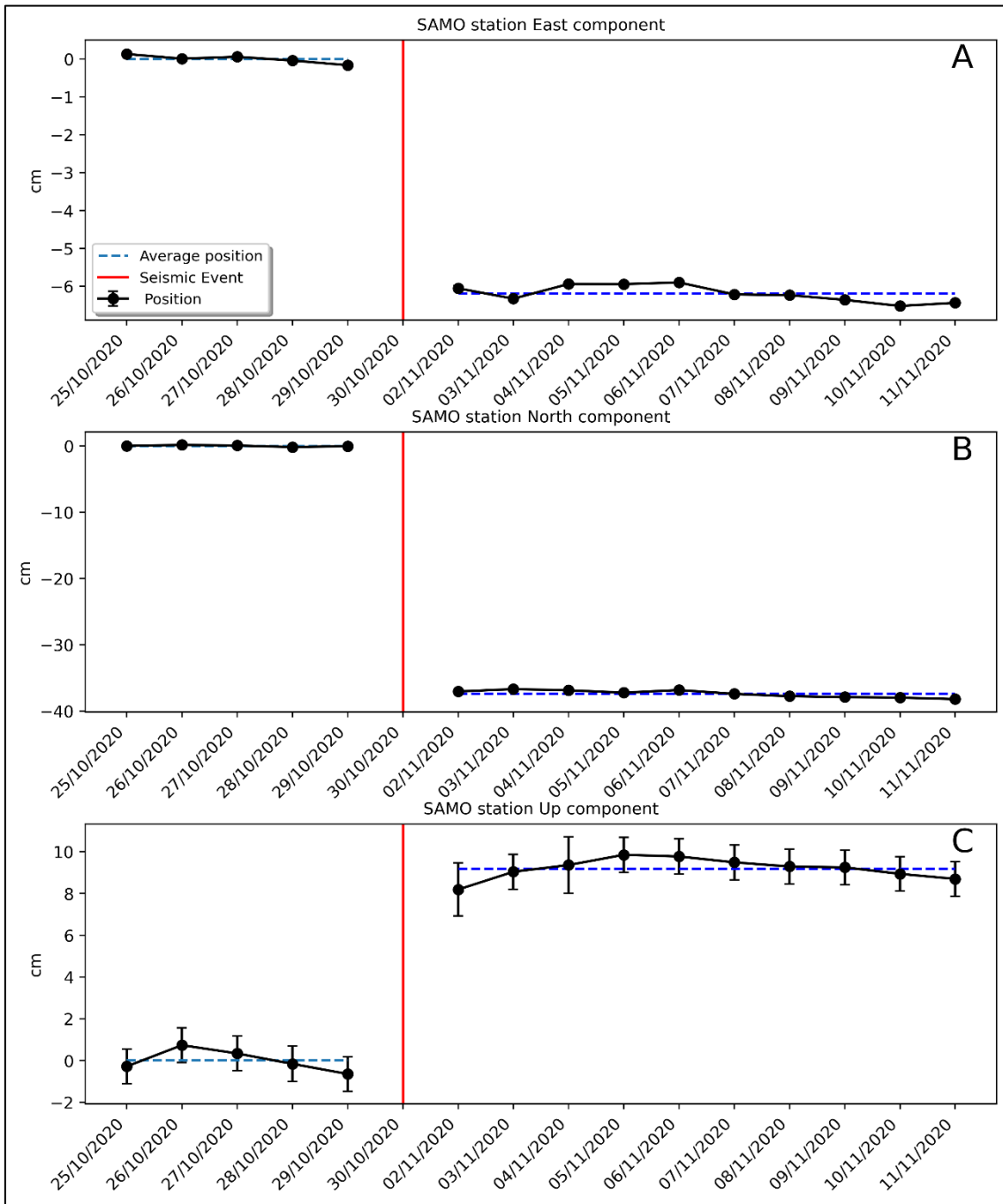
1076

1077

1078

1079

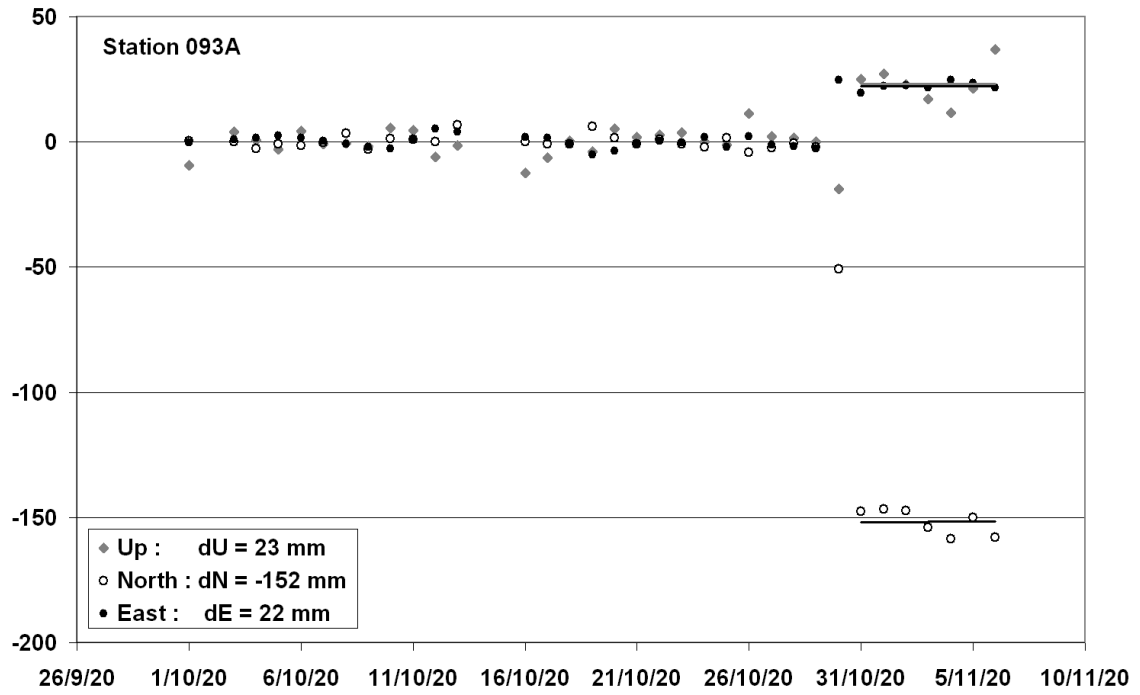
1080 **Fig. S4:** Graphs with time series of GNSS coordinates of the SAMO (a) and 093A (b) GNSS  
1081 station across the day of the earthquake (red vertical line in a).



1082

1083

(a)



1084

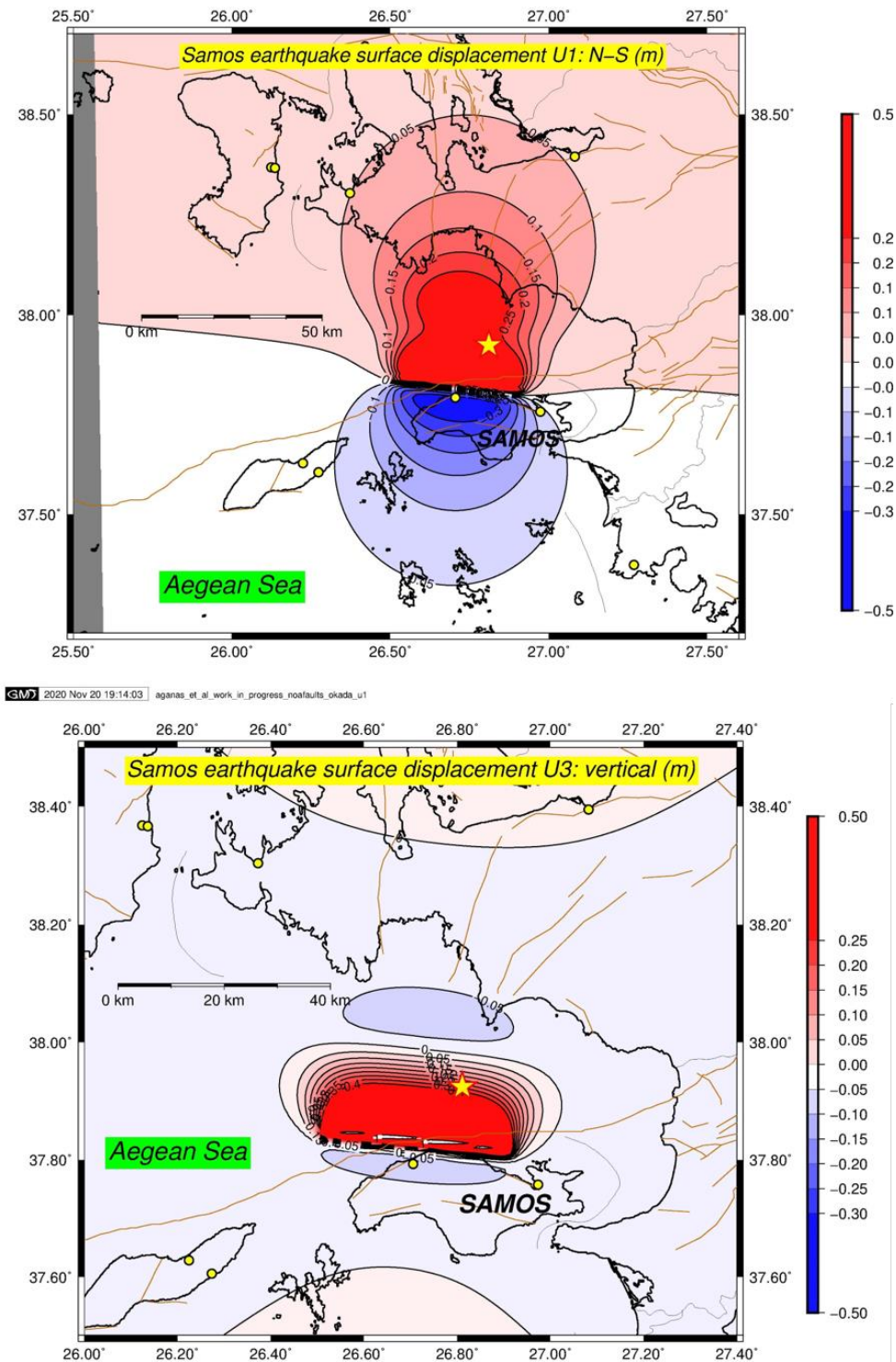
1085

(b)



1086 **Fig. S5:** Synthetic models of ground surface displacement using Okada-type dislocations in  
 1087 elastic half-space. The modeled dislocation (from Ganas et al. 2020a) has length/width of  
 1088 36 km / 18 km, a finite slip of 1.82 m and a strike/dip/rake: 276°/37°/-90° (see GCMT and  
 1089 USGS moment tensor solutions from Table 1 on the main paper). Top panel) North-South  
 1090 displacement bottom panel) Up-Down displacement. Values are in m. Yellow star is the  
 1091 epicentre of the mainshock.

1092



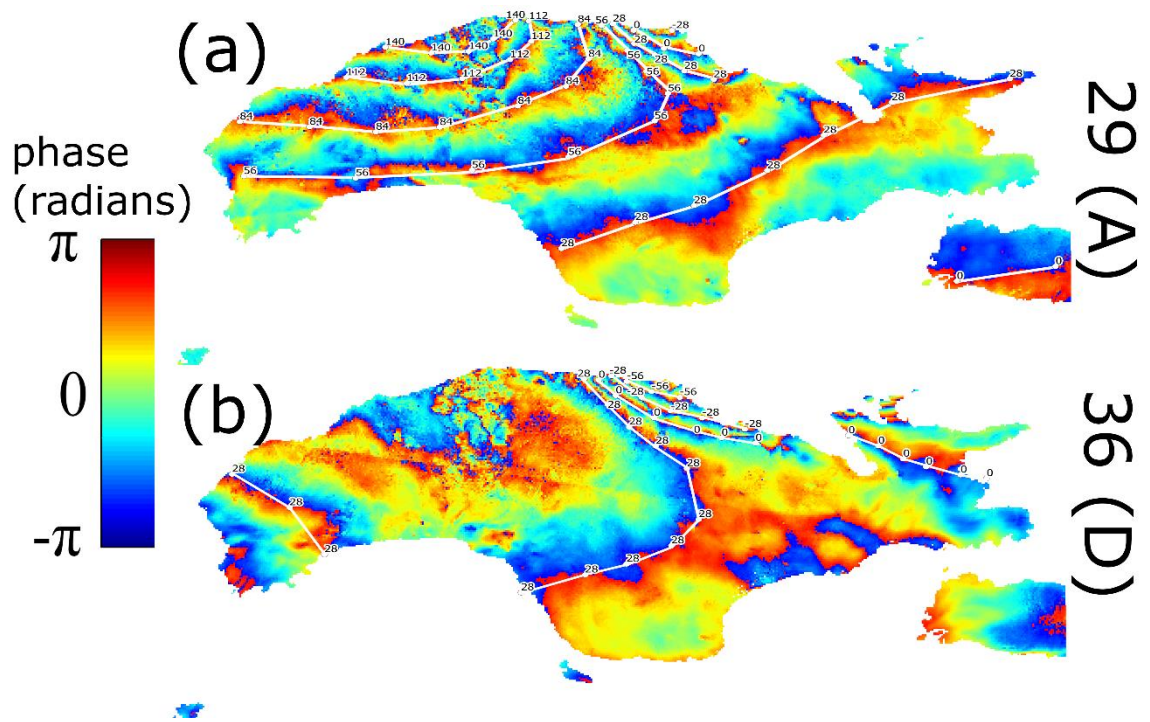
1093

1094 **Fig. S6:** Picked values in the two averaged interferogram (a) ascending track 29 and b)  
1095 descending track 36/

1096

1097

1098



1099

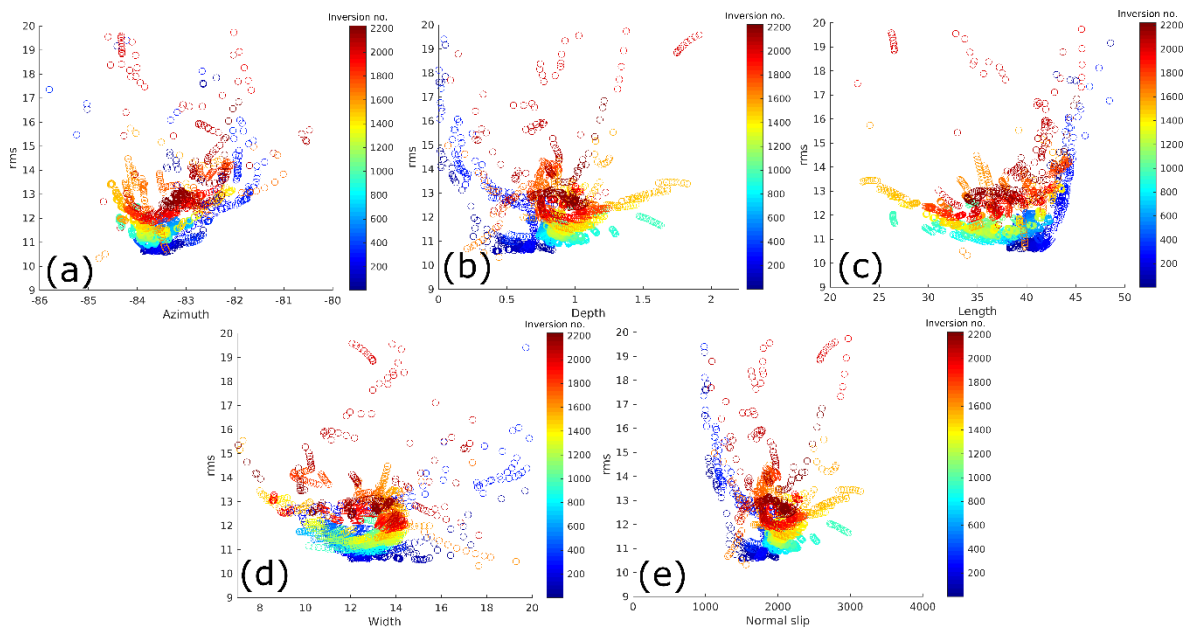
1100

1101 **Fig. S7.** Divergence diagrams of the set of inversions for a) fault azimuth (strike), b) depth  
1102 to top-fault, c) length, d) width and e) normal slip.

1103

1104

1105



1106

1107

1108

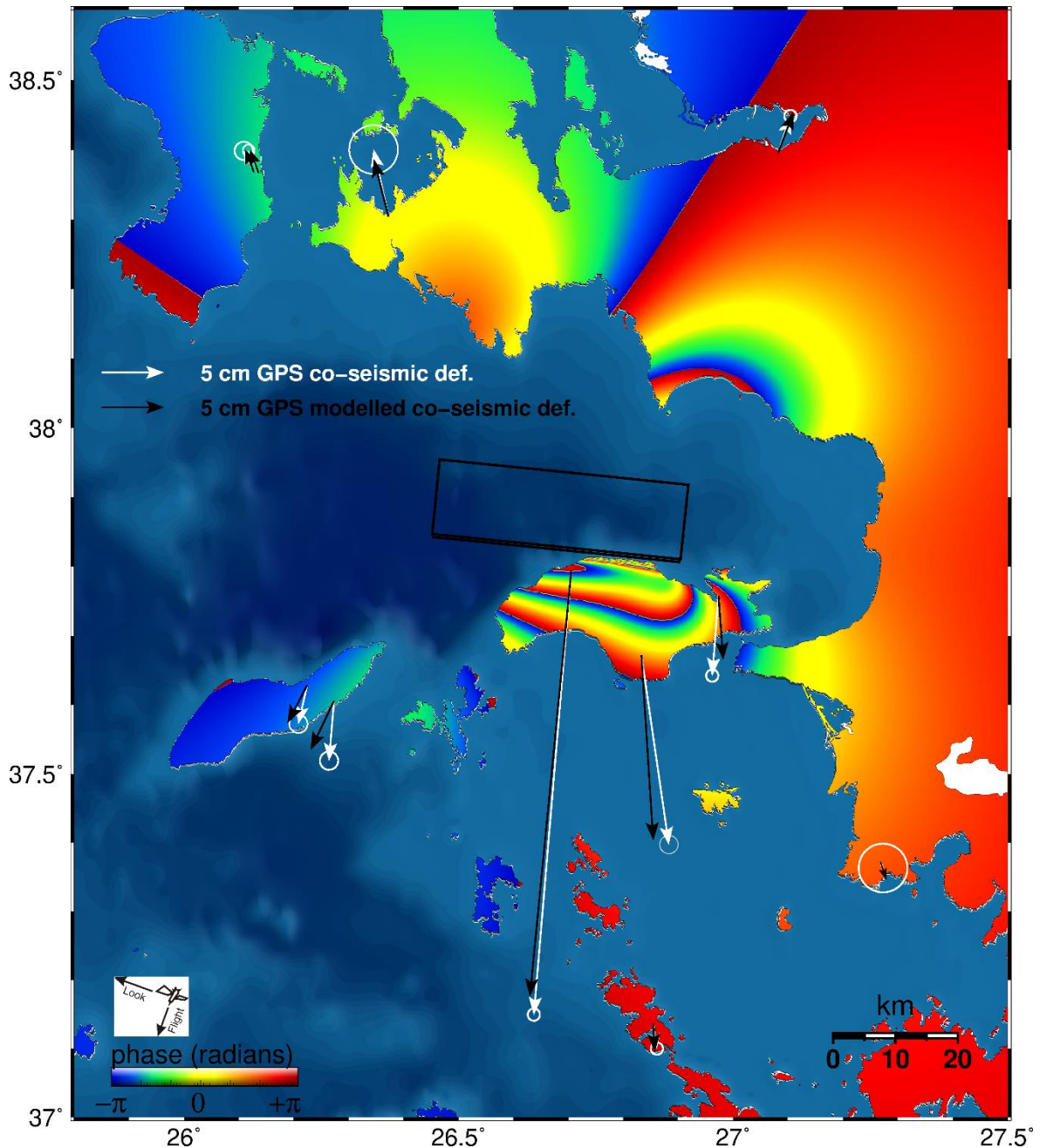
1109

1110

1111 **Fig. S8.** Fault model from the InSAR data (synthetic model; descending orbit). The dip-  
1112 direction is towards north. Arrows indicate GPS offset vectors (observed vs. predicted). The  
1113 formation of the interferogram is at the notion of post-event minus pre-event image, thus  
1114 phase decreasing towards satellite.

1115

1116

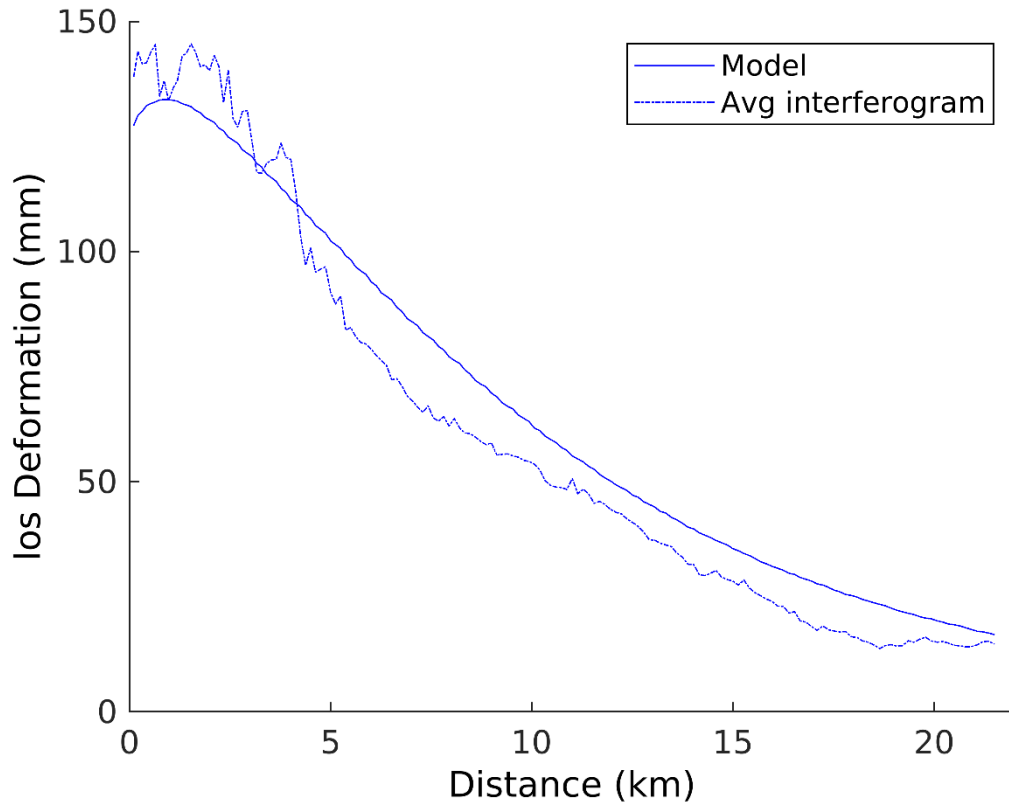


1117

1118

1119

1120 **Fig. S9.** Displacement profile onshore Samos showing a comparison of the InSAR observed  
1121 values (LOS; track 29) versus the modelled values (see Fig. 8 of the main paper). The  
1122 location of the profile is shown in Fig. 3.



1123