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Key Points:

- A globally enhanced ecosystem C stock (ECS) by land greening has been detected since the 1980s from 40 global models
- The sensitivity of gross primary productivity to leaf area index (LAI) is the largest contributor to the modeled greening effect on ECS uncertainty
- This study estimated the effect of global greening on ECS with the emergent constraint method

Supporting Information:

Supporting Information may be found in the online version of this article.

Correspondence to:

J. Xia, jyxia@des.ecnu.edu.cn

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Uncertainty and Emergent Constraints on Enhanced Ecosystem Carbon Stock by Land Greening

Chenyu Bian^{1,2} , Jianyang Xia^{1,2} , Xuanze Zhang³ , Kun Huang^{1,2} , Erqian Cui^{1,2} , Jian Zhou^{1,2,4} , Ning Wei^{1,2} , Ying-Ping Wang⁵ , Danica Lombardozzi⁶ , Daniel S. Goll⁷ , Jürgen Knauer⁸ , Vivek Arora⁹, Wenping Yuan¹⁰, Stephen Sitch¹¹ , Pierre Friedlingstein¹² , and Yiqi Luo⁴

¹Zhejiang Tiantong Forest Ecosystem National Observation and Research Station, Institute of Eco-Chongming, School of Ecological and Environmental Sciences, East China Normal University, Shanghai, China, ²Research Center for Global Change and Complex Ecosystems, East China Normal University, Shanghai, China, ³Key Laboratory of Water Cycle and Related Land Surface Processes, Institute of Geographical Sciences and Natural Resources Research, Chinese Academy of Sciences, Beijing, China, ⁴School of Integrative Plant Science, Cornell University, Ithaca, NY, USA, ⁵CSIRO Oceans and Atmosphere, Aspendale, VIC, Australia, ⁶Terrestrial Sciences Section, Climate and Global Dynamics, National Center for Atmospheric Research, Boulder, CO, USA, ⁷LSCE/IPSL, CEA-CNRS-UVSQ, Université Paris Saclay, Gif sur Yvette, France, ⁸CSIRO Oceans and Atmosphere, Canberra, ACT, Australia, ⁹Climate Research Division, Environment and Climate Change Canada, Canadian Centre for Climate Modelling and Analysis, Victoria, BC, Canada, ¹⁰Guangdong Province Key Laboratory for Climate Change and Natural Disaster Studies, Zhuhai Key Laboratory of Dynamics Urban Climate and Ecology, School of Atmospheric Sciences, Sun Yat-sen University, Zhuhai, China, ¹¹College of Life and Environmental Sciences, University of Exeter, Exeter, UK, ¹²College of Engineering, Mathematics and Physical Sciences, University of Exeter, Exeter, UK

Abstract Significant land greening since the 1980s has been detected through satellite observation, forest inventory, and Earth system modeling. However, whether and to what extent global land greening enhances ecosystem carbon stock remains uncertain. Here, using 40 global models, we first detected a positive correlation between the terrestrial ecosystem carbon stock and leaf area index (LAI) over time. Then, we diagnose the source of uncertainty of simulated the sensitivities of ecosystem carbon stock to LAI based on a traceability analysis. We found that the sensitivity of gross primary productivity (GPP) to LAI is the largest contributor to the model uncertainty in more than 60% of the vegetated grids. Using the ensemble of four long-term global data sets of GPP and three satellite LAI products from 1982 to 2014, we provided an emergent constraint on the ecosystem carbon stock increase as 0.75 ± 0.46 kg C m⁻² per unit LAI over global land areas. Furthermore, the biome-based results reveal that the tropical forest regions have the highest inter-model variation and model bias. Overall, this study identifies the uncertainty source and provides constrained estimates of the greening effect on ecosystem carbon stock at the global scale.

Plain Language Summary In the last few decades, we can see that there has been more green land from looking at satellite images, checking forest records, and using computer models. However, we do not know if having more green land is helping the ecosystem store more carbon. This study investigated the greening effect on ecosystem carbon stock among the 40 global models, but this effect varied between models. Therefore, we diagnosed the simulated source of uncertainty based on the traceability analysis framework. Based on satellite-derived data sets, we further constrained the greening effect on ecosystem carbon stock at the global scale. Land greening enhances ecosystem carbon stock, although this increase is highly heterogeneous on a global scale.

1. Introduction

The satellite-observed greening of the land surface has been unequivocal since the early 1980s and widely reported in recent studies mainly as a result of CO₂ fertilization (Chen et al., 2019; Mao et al., 2013; Myneni et al., 1997; Pan et al., 2011; Z. Zhu et al., 2013, 2016). Global land greening suggests more carbon (C) is taken up by land ecosystems due to increased absorption of photosynthetically active radiation (PAR) (Chen et al., 2019; Keenan et al., 2016; Pan et al., 2011; Qian et al., 2010). Multiple lines of evidence have shown that global land greening could slow down the increase in atmospheric carbon dioxide (CO₂) concentration by transferring C from the atmosphere to land (Ballantyne et al., 2017; Chen et al., 2019; Graven et al., 2013; Keenan et al., 2016; Pan et al., 2011; Piao et al., 2020; Qian et al., 2010). However, the mean residence time of C in

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Writing – review & editing: Chenyu Bian, Jianyang Xia, Xuanze Zhang, Kun Huang, Erqian Cui, Jian Zhou, Ning Wei, Ying-Ping Wang, Danica Lombardozzi, Daniel S. Goll, Jürgen Knauer, Vivek Arora, Wenping Yuan, Stephen Sitch, Pierre Friedlingstein, Yiqi Luo terrestrial ecosystems ranges from decades in tropical forests (TRFs) to centuries in tundra (J. Wang et al., 2019). Thus, it has been challenging to assess to what extent global land greening has enhanced terrestrial C stock in the past decades.

The ecosystem C stock (ECS) is jointly determined by vegetation productivity and C residence time (Arora et al., 2013; Luo et al., 2003, 2017; Xia et al., 2013; Zhang et al., 2018). However, evidence has revealed contrasting trends in ecosystem productivity (Arora et al., 2013; Friedlingstein et al., 2014; Piao et al., 2020; Winkler, Myneni, Alexandrov, & Brovkin, 2019) and C residence time over the past decades (Brienen et al., 2020; Lu et al., 2018; Wu et al., 2020). On the one hand, land greening is characterized by an increasing leaf area index (LAI), which leads to more active canopy photosynthesis on a global scale. For example, satellite observations and model simulations have shown increasing vegetation productivity (i.e., gross primary productivity, GPP) (Arora et al., 2013; Friedlingstein et al., 2014; Piao et al., 2020; Winkler, Myneni, Alexandrov, & Brovkin, 2019) and peak growth (Huang et al., 2018) since the 1980s. On the other hand, the environmental drivers of land greening, such as rising atmospheric CO₂ concentration and climate warming, could lower C residence time (Brienen et al., 2020; Lu et al., 2018; Wu et al., 2020) in the terrestrial ecosystem by accelerating the turnovers of woody biomass (Yu et al., 2019) and soil organic carbon (Bond-Lamberty et al., 2018; Van Groenigen et al., 2014). Potentially contrasting trends between ecosystem C input and residence time suggest that global land greening could trigger complex plant-soil-atmosphere interactions (Mangan et al., 2010; Piao et al., 2019) and non-uniform effects on ECS across the globe. It remains unclear which processes have dominated ecosystem carbon stock changes due to vegetation greening over the past decades.

Recent advances in data-model integration approaches offer an opportunity to quantify the effect of land greening on ECS on a global scale. By fully coupling the interactions between plants, soil, and atmosphere (Arora et al., 2020), Earth system models (ESMs) are widely used to quantify the greening effect on ECS. Most ESMs currently simulate an increasing trend in carbon uptake (Arora et al., 2013; Friedlingstein et al., 2014; Piao et al., 2020; Winkler, Myneni, Alexandrov, & Brovkin, 2019) and a decreasing trend in ecosystem C residence time (Bond-Lamberty et al., 2018; Van Groenigen et al., 2014; Wu et al., 2020; Yu et al., 2019). However, the simulated magnitude for greening trends and terrestrial C stock by the state-of-the-art ESMs diverged significantly (Luo et al., 2016; Mahowald et al., 2016; Zeng et al., 2016; Z. Zhu et al., 2017). The emergent constraint method developed by Cox et al. (2013) can reduce the inter-model uncertainty by defining an emergent relationship across the multi-model ensemble based on observations. In some previous studies, this approach has been applied to quantify the impacts of global change factors on multiple terrestrial processes, such as GPP (Wenzel et al., 2016), evapotranspiration (Lian et al., 2018), and soil C turnovers (Varney et al., 2020).

The emergent constraint method can reduce inter-model uncertainty by combining the emergent relationship and the corresponding observations (Hall et al., 2019). However, the internal mechanism and source of uncertainty for inter-model remain unclear. In addition, the constrained results have their internal uncertainty due to concerns about the emergent constraints (Winkler, Myneni, & Brovkin, 2019). For example, the emergent relationship is derived from the model ensemble, and thus it could be misleading if the model ensemble has some systematic error (Eyring et al., 2019). For the land C cycle, the traceability analysis framework can trace the inter-model uncertainty by decomposing a complex land model to its traceable components based on the C cycle fundamental properties (Luo et al., 2017; Xia et al., 2013). This framework has been widely used for tracing and evaluating the model uncertainty (Wei, Xia, Zhou, et al., 2022; S. Zhou et al., 2018; J. Zhou et al., 2021) at both site and global scales. Therefore, combining the emergent constraint and traceability analysis can help us constrain the inter-model uncertainty and further understand the effects of vegetation on C stock in global models.

In this study, we use the modeling outputs from the fifth and sixth phases of Coupled Model Intercomparison Projects (CMIPs) (i.e., CMIP5 and CMIP6; Tables S1 and S2 in Supporting Information S1) and the Trends in Net Land-Atmosphere Carbon Exchanges project (TRENDY-v9; Table S3 in Supporting Information S1) to investigate the greening effect (i.e., LAI) on ECS since the 1980s. Then, we diagnose the main contributor to the large spread of the simulated greening effect on C stock among different models based on the traceability analysis method. After that, by building an emergent relationship between the greening effects on GPP and ECS, we constrain the estimated greening effect on ECS based on the long-term global observations using an emergent constraint approach. We explore the spatial pattern of the greening effect on ECS using multiple sources of data sets. Lastly, we compare the inter-model variation and model bias of the greening effect on C stock among different biomes.

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2. Materials and Methods

2.1. The Model Intercomparison Projects

In this study, we evaluated the greening effect on ecosystem carbon stock based on the model simulation results among the three MIPs (i.e., CMIP5, CMIP6, and TRENDY-v9). The CMIPs were established to compare the intermodel climate variability and fill the understanding gap of land-atmospheric carbon exchange and feedback with the climate system (Eyring et al., 2016; Taylor et al., 2012). To improve predictions of the terrestrial ecosystem to climate change, the land carbon cycle models have gradually incorporated more and more new processes (Prentice et al., 2015). In the meantime, it also leads to large variations in the simulations and predictions across models and MIPs (S. Zhou et al., 2018). This study uses two generations of CMIPs (CMIP5 and CMIP6) to estimate the greening effects on ECS. There are two primary reasons to include different numbers of models: (a) the variant label of ESMs with "r1i1p1" for CMIP5 and "r1i1p1f1" for CMIP6; (b) a model can be selected if it provides all the diagnostic variables that are required for this study. The variables used to evaluate greening effects include carbon fluxes (i.e., net primary production (npp), gross primary production (gpp), net ecosystem production (nep), and heterotrophic respiration (rh)), leaf area index (lai), carbon pools (cVeg, cLitter, and cSoil), near-surface temperature (tas) and precipitation (pr). Based on these two requirements, a total of 16 ESMs for CMIP5 and 12 ESMs for CMIP6 were selected. More details about the variables used in each model are shown in Tables S1 and S2 in Supporting Information S1. Thus, this study selected outputs of 28 models from the CMIP historical simulations. The CMIP historical simulations of climate change are forced by common data sets (i.e., CO₂ concentration and land use) that largely based on observations. Thus, the historical simulations have included the impact of land-use and land cover change on ECS (Lawrence et al., 2016). In general, they serve as a benchmark for assessing model performance through evaluation with observations (Eyring et al., 2016).

CMIP5. The fifth phase of the Coupled Model Intercomparison Project (CMIP5) mainly focused on the gaps in understanding past and future climate change by performing a suite of climate simulations (Taylor et al., 2012). It contributed from various modeling centers involving the Fifth Assessment Report (AR5) of the International Panel on Climate Change (IPCC). More than 50 models from 20 modeling groups have attended CMIP5. The simulations include two types of climate change experiments: long-term and near-term integrations (Taylor et al., 2012). We only used the long-term history simulation results in this study. The simulation data sets are available from the CMIP5 archive (http://cmip-pcmdi.llnl.gov/cmip5/). We selected 16 models from the historical model scenarios (Table S1 in Supporting Information S1).

CMIP6. The Coupled Model Intercomparison Project Phase 6 (CMIP6) aims to address a wide variety of science questions from a spread of scientific research communities (Eyring et al., 2016; Meehl et al., 2000, 2005, 2014). CMIP6 has adopted a more federated structure and subsequent phase. The core experiments include DECK (Diagnostic, Evaluation, and Characterization of Klima) experiments and historical simulations (Eyring et al., 2016). This study selected 12 models (Table S2 in Supporting Information S1) from the historical simulations based on the data availability (https://pcmdi.llnl.gov/CMIP6/).

TRENDY. This study used an ensemble of 12 processed-based dynamic global vegetation models (DGVMs), including CABLE-POP, CLASSIC, CLM5.0, IBIS, JSBACH, JULES-ES, LPX-Bern, OCN, ORCHIDEE-CNP, ORCHIDEEv3, SDGVM, and VISIT from the Trends in Net Land-Atmosphere Carbon Exchanges project (TRENDY-v9; Friedlingstein et al., 2020) for 1982–2019 (Table S3 in Supporting Information S1). These models have performed the same set of factorial simulations following a standard protocol. Therefore, it can avoid some uncertainties due to climate variability. This study used the output from "S3" simulations under the scenarios with time-varying atmospheric CO_2 and changing climate and land use. The annual time series of global atmospheric CO_2 is derived from ice core CO_2 data merged with NOAA annual resolution for 1700–2019. Either monthly CRU or 6-hourly CRU-JRA55 with a $0.5^{\circ} \times 0.5^{\circ}$ resolution can be used to force these DGVMs. DGVMs are spun-up to steady-state by recycling climate mean and variability from 1901 to 1920 and an atmospheric CO_2 concentration of 276.59 ppm (i.e., 1700 level).

We summed all the carbon pools as ecosystem carbon stock for each model. All the model output variables from three MIPs were linearly interpolated from their native resolution to a standard $0.5^{\circ} \times 0.5^{\circ}$ grid on a global scale. In addition, we only selected the temporal period after 1982 (1982–2005 for CMIP5; and 1982–2014 for CMIP6; 1982–2019 for TRENDY-v9), considering global greening has been detected since the early 1980s and the satellite-observed LAI data sets are available from 1982.

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2.2. Observation Data Sets

Satellite-derived data products provide consistent estimates of LAI and GPP across large spatial and temporal ranges. These data sets have also been widely used for monitoring terrestrial vegetation growth and ecosystem productivity. In this study, three available LAI and four GPP data sets are integrated to access the greening effect on ECS. In addition, two independent data sets, which include eddy-flux observations and the vegetation optical depth (VOD) derived aboveground biomass (AGB), are used to validate the constrained spatial pattern further.

2.2.1. Satellite-Observed LAI Products

The Global Inventory Modeling and Mapping Studies (GIMMS) LAI3g product (Z. Zhu et al., 2013) was generated by Feed-Forward Neural Network (FFNN) algorithm based on NDVI3g data from July 1981 to December 2011. It provides global data at a spatial resolution of $1/12^{\circ}$ and a temporal resolution of half a month. The long-term GLOBMAP LAI product was generated by a combination of AVHRR LAI (1981–2000) and MODIS LAI (2000–2011) (Y. Liu et al., 2012). By establishing the relationship between AVHRR observations and MODIS LAI during the overlapped period (2000–2006), the AVHRR LAI can be retrieved back to 1981. The temporal resolution of this data set is half a month and eight days in 1981–2000 and 2001–2011, respectively. The spatial resolution is 8 km. The GLASS LAI product was estimated from MODIS and AVHRR time-series remote sensing data using the General Regression Neural Networks (GRNNs) approach. The GLASS LAI product provides a geographic projection at the resolution of 0.05° from 1981 to 1999 and a sinusoidal projection at 1 km from 2000 to 2014 (Xiao et al., 2013, 2016). The three satellite-observed products were re-sampled into $0.5^{\circ} \times 0.5^{\circ}$ special resolution using the nearest neighbor algorithm for the period 1982–2014.

2.2.2. Satellite-Derived GPP Data Sets

This study uses four satellite data-driven long-term global GPP data sets, including EC-LUE GPP, GPP_{Enh}, P-Model GPP, and RS-LUE GPP (described below). We linearly interpolated all the four GPP data sets from their native resolution to a standard $0.5^{\circ} \times 0.5^{\circ}$ spatial resolution and further selected the data sets from 1982 to 2014 to match the LAI data sets.

EC-LUE GPP. The GPP data set was generated from the revised light-use efficiency (LUE) model, EC-LUE, by integrating the regulations of several environmental variables, which include CO_2 concentration, radiation, and atmospheric vapor pressure deficit (VPD), and driven by four variables: normalized difference vegetation (NDVI), PAR, air temperature, and the Bowen ratio of sensible to latent heat flux (Yuan et al., 2007, 2010). The EC-LUE model is based on two fundamental assumptions: (a) the fraction of absorbed PAR (fPAR) has a linear relationship with NDVI; (b) the LUE is limited by air temperature or soil moisture (Yuan et al., 2007). The revised EC-LUE model performed well in the simulation of global GPP and was validated and compared by some other studies (Zheng et al., 2020). The temporal resolution of this data set is eight days for 1982–2018, and the spatial resolution is 0.05° .

 GPP_{Enh} . The GPP_{Enh} data set was generated based on an enhanced remote sensing LUE model (Madani et al., 2017, 2020). The LUE model which provides enhanced GPP estimates in this study as follows: $GPP_{Enh} = F$ PAR × PAR × LUE_{opt} × fVPD × fSM × fT_{min} , where fVPD, fSM, and fT_{min} represents the water and temperature environmental constraint functions. The optimized LUE was extrapolated to the global domain based on a generalized additive model GAM framework (Hastie & Tibshirani, 1986). Global monthly GIMMS3g FPAR (Z. Zhu et al., 2013), meteorological data sets from Modern-Era Retrospective analysis for Research and Applications, Version2 (MERRA-2; Gelaro et al., 2017), and geospatial information are the primary input to drive the LUE model. More details about this data set can be found in Madani et al. (2017, 2020). It provides global data at a spatial resolution of 8 km and a monthly temporal resolution from 1982 to 2016.

P-model GPP. The P-model is also a LUE model based on an optimality principle that predicts stomatal conductance and foliar photosynthesis (H. Wang et al., 2017). This study used the estimated GPP on a global scale developed by B. D. Stocker et al. (2018, 2019) with the Seasonal optimization of fixation and uptake of nitrogen (SOFUN v1.1.0, B. Stocker, 2018) modeling framework. In the P-model, GIMMS3g FPAR data sets drive the simulations of GPP on a global scale from 1982 to 2016. A set of empirical soil moisture stress functions (GPP = $\beta(\theta)$ GPP_{mod}) are used to correct the simulated GPP. The simulations of s1b results are used in this study. Where s1b represents using parameters of β_b (the intermediate sensitivity and distinguishing

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parameters between woody and herbaceous vegetation) to estimate the uncertainty in the sensitivity of $\beta(\theta)$. The temporal resolution of the GPP data set provided by the P-model is daily for 1982–2016, and the spatial resolution is 0.5° .

RS-LUE GPP. The remote-sensing-based LUE (RS-LUE) model can estimate GPP fluxes based on the MODIS GPP algorithm (Kolby Smith et al., 2016; Running et al., 2004) as GPP = FPAR × PAR × LUE_{max} × $f(T_{min})$ × f(VPD). The FPAR is a satellite-derived fraction of PAR absorbed by vegetation, which is based on the GIMMS3g using a neural network algorithm (Z. Zhu et al., 2013). PAR is incoming PAR. LUE_{max} represents the maximum LUE. The functions of $f(T_{min})$ and f(VPD) represent the environmental function for minimum temperature and VPD, respectively. In this study, we used an updated version extending for 1982–2016 (Version 4) with the CRUNCEP P1 Standard run (O'Sullivan et al., 2020). The spatial resolution is $0.5^{\circ} \times 0.5^{\circ}$ grid, and the temporal resolution is monthly.

2.2.3. Data-Oriented Aboveground Biomass Carbon and FLUXNET Data

The passive microwave-based global AGB carbon product (version 1.0; Y. Y. Liu et al., 2015) is derived from a harmonized time series of VOD from 1993 to 2012. The passive microwave-based AGB data set provided annual estimates with a spatial resolution of 0.25° , respectively. The reported uncertainty of the passive microwave-based AGB in each grids cell was less than 30% (Saatchi et al., 2011). In this study, we re-sampled the AGB data from the spatial resolution of 0.25° – 0.5° .

The daily eddy covariance NEE data were obtained from FLUXNET2015 (Tier1) data sets (http://fluxnet.flux-data.org/data/fluxnet2015-dataset/), which is an openly available database that offers an observation of carbon fluxes between ecosystems and the atmosphere. These data included in the data sets are collected at sites from multiple flux networks worldwide (https://fluxnet.fluxdata.org/) and were pre-filtered to exclude the low turbulence conditions to meet the requirement of the eddy-covariance technique, as described in Papale et al. (2006). A total of 82 flux sites (Table S4 in Supporting Information S1 for a full list, which has at least 5-year consistent data) with 925 site-years of NEE data were used in this study. However, considering flux sites did not include corresponding LAI data sets, we extracted the satellite observed LAI values within a 3×3 pixel window around the site based on the original resolution for each of the three LAI products (Table S5 in Supporting Information S1) as similar to the previous study (Huang et al., 2019). Then we use the average of 3×3 pixels based on the ensemble of three LAI products to represent the LAI values in each site. Here we use the positive sign convention for land carbon uptake, that is: NEP = -NEE, and calculated the sensitivity of NEP to LAI-based observed data sets (Table S4 in Supporting Information S1).

2.2.4. Biome Classification

We evaluated the greening effect on C stock at the biome-level by constructing a common biome mask. The biome map was based on the MODIS/TERRA-AQUA land cover product (MCD12C1) Collection 5.1 (Friedl et al., 2010). We masked water, permanent snow and ice, and barren regions first. We assigned one of the remaining land cover types to a $0.5^{\circ} \times 0.5^{\circ}$ grid cell by taking the most common land cover from the original underlying $0.05^{\circ} \times 0.05^{\circ}$. Then each 0.5° grid was further classified into one of eight biomes, following the details in Todd-Brown et al. (2013) as tundra, boreal forest (BOF), TRF, temperate forest (TEF), shrubland, grassland and savannas, cropland and urban, or wetlands (Figure S1 in Supporting Information S1).

2.3. Traceability Analysis on the Modeled Sensitivity of C Stock to LAI

2.3.1. Traceability Analysis Framework on C Stock

We conducted a traceability analysis to further explain the wide spread of sensitivities of ecosystem carbon stock to LAI among the 40 global models from three MIPs based on the Traceability analysis system for Model Evaluation (TraceME v1.0; J. Zhou et al., 2021). The TraceME system is based on the traceability analysis framework (Xia et al., 2013) to identify the inter-model source of uncertainties in ecosystem carbon stock to its components. This traceability framework has been widely used and developed in diagnosing the causes of differences in simulated terrestrial carbon cycle among models (Jiang et al., 2017; Luo et al., 2017; Wei, Xia, Zhou, et al., 2022; Xia et al., 2013; S. Zhou et al., 2018). The traceability analysis method used in this study is briefly described below.

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ECS is jointly determined by ecosystem carbon influx (e.g., NPP) and ecosystem C residence time (τ_E) (Xia et al., 2013):

$$ECS = \tau_E \times NPP \tag{1}$$

The ecosystem C residence time (Lu et al., 2018) in Equation 1 can be calculated by the following equation:

$$\tau_E = ECS/R_h \tag{2}$$

where ECS is the current C stock, which equals the sum of all C pools in each model. R_h represents the carbon losses through heterotrophic respiration. Generally, τ_E are directly available based on each model structure and the environmental scalars (i.e., ξ):

$$\tau_E = \tau_E' \xi^{-1} \tag{3}$$

where τ_E' represents ecosystem baseline C residence time. The term τ_E' usually preset based on the model-specific structure and parameters (i.e., C transfer coefficients among different C pools, C allocation coefficients, and decomposition coefficients); ξ is an environmental scalar that reflects the environmental regulators on τ_E' , especially the temperature (ξ_T) and water (ξ_W) scalar. In this study, we only considered the effect of temperature (ξ_T) and water (ξ_W) as similar to the previous studies (Jiang et al., 2017; Wei, Xia, Zhou, et al., 2022; Xia et al., 2013; S. Zhou et al., 2018; J. Zhou et al., 2021):

$$\xi = \xi_T \xi_W \tag{4}$$

The term ξ_T and ξ_W can be estimated by mean annual temperature (*T*) and annual total precipitation (*W*) (S. Zhou et al., 2018) as:

$$\xi_T = Q_{10}^{\left[(T - T_0)/10 \right]} \tag{5}$$

$$\xi_W = \frac{W}{W_0} \tag{6}$$

where Q_{10} is a parameter that represents the temperature sensitivity of ecosystem respiration. The terms T_0 and W_0 represent the reference temperature and precipitation, respectively. In this study, T_0 and W_0 were set as the maximum values of annual temperature and precipitation in the simulation period for each model.

In addition, NPP in Equation 1 can be decomposed into gross primary production (GPP) and carbon use efficiency (CUE) (Xia et al., 2017):

$$NPP = CUE \times GPP \tag{7}$$

2.3.2. Source of Uncertainty on Modeled Sensitivity of ECS to LAI

Applying the traceability analysis framework, we first decomposed the ECS to its traceable component. Then, we quantified the contributions of each traceable component to the uncertainty of modeled sensitivity of ECS to LAI based on the hierarchical partitioning method (Chevan & Sutherland, 1991; Murray & Conner, 2009). This method can be used to calculated the independent effect of each explanatory variable on a single dependent variable. For a given dependent variable (y) and k explanatory variables ($x_1, x_2, x_3, ..., x_k$), there would be 2^k different combinations of explanatory variables. For example, when k = 3, there are $2^3 = 8$ numbers are required. The eight numbers are indexed as $X_0, X_1, X_2, X_3, X_{12}, X_{13}, X_{23}$, and X_{123} . Each subscript represents the combination of the concluded explanatory variables (i.e., X_{12} represents the combination of x_1 and x_2 , while X_0 is a null number). The independent effect of each explanatory variable x_1 (I_{xp} I = 1, 2, 3, ..., k) on a single dependent variable y can be calculated by comparing the fit of all nested models (2^k) lacking that variable based on the hierarchical ordering (Chevan & Sutherland, 1991; Murray & Conner, 2009):

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$$I_{x_{l}} = \sum_{i=0}^{k-1} \frac{\sum (r_{y,x_{l}x_{h}}^{2} - r_{y,x_{h}}^{2}) / \binom{k-1}{i}}{k}$$
(8)

where x_h is any subset of explanatory variables excluding x_l (Chevan & Sutherland, 1991; Murray & Conner, 2009). The relative independent contribution (RI_{xl}) can be further calculated as:

$$RI_{x_{l}} = \frac{I_{x_{l}}}{\sum_{l=1}^{k} I_{x_{l}}} \tag{9}$$

We converted all the traceable components to their logarithmic form to ensure the formulations could be expressed in sum form. Then, we calculated the sensitivity of each component to LAI (e.g., $\partial \ln(ECS)/\partial LAI$, $\partial \ln(NPP)/\partial LAI$, $\partial \ln(CUE)/\partial LAI$, $\partial \ln(GPP)/\partial LAI$, $\partial \ln(\tau_E)/\partial LAI$, $\partial \ln(\tau_E)/\partial LAI$, $\partial \ln(\xi_E)/\partial LA$

2.4. Emergent Constraint Method

The simulation results of ESMs often exhibit significant differences due to their complexity and incorporation of many processes. It is difficult to understand the climate system or predict future changes accurately. The terrestrial ecosystem, which sequesters about one-third of anthropogenic CO2 emissions each year, acts as a substantial carbon sink and plays a vital role in climate change (Friedlingstein et al., 2020; Le Quéré et al., 2018). Therefore, improving our understanding of the terrestrial ecosystem is important to reduce current uncertainties in projections of future change. In a previous study, Cox et al. (2013) have developed an "emergent constraint" method to reduce the uncertainty of future climate impact on the tropical C cycle. This approach defines an emergent relationship among multi-model outputs and uses one observable variable to constrain another that could not be observed or measured directly (Eyring et al., 2019; Hall et al., 2019). Here we use a similar constrained approach to evaluate the greening effect on ECS. To obtain model-specific greening effect on GPP and C stock, the monthly model output data were first time-averaged to yearly over the study period. Then, for each model, the sensitivities of GPP and C stock to LAI are calculated based on the globally averaged values. We established an emergent relationship between the sensitivities of GPP and C stock to LAI based on the ensemble of global models in CMIP5, CMIP6, and TRENDY (1982-2005 for CMIP5; and 1982-2014 for CMIP6; 1982-2019 for TRENDY-v9). We compared the different GPP-LAI sensitivities by analyzing all the combinations of four GPP and three LAI products (n = 12; Figure S2 in Supporting Information S1). The constrained results produced by different GPP and LAI combinations are highly uncertain (Figure S3 in Supporting Information S1). Thus, the multiple data sets of GPP and LAI were integrated as a constrain to reduce the possible observational uncertainty. In addition, we also calculated the greening effect on GPP and C stock for each pixel for each model to do the spatial emergent constraint.

A linear least-squares regression model was first built between the partial sensitivity of ECS to LAI (∂ ECS/ ∂ LAI) and the partial sensitivity of GPP to LAI (∂ GPP/ ∂ LAI) across the models:

$$\hat{y}_i = a + bx_i + \varepsilon_i \tag{10}$$

Let the least-squares for y_i involves minimizing, then the least-squares error of the regression model can be calculated as:

$$s^{2} = \frac{1}{N-2} \sum_{i=1}^{N} \{ y_{i} - \hat{y}_{i} \}^{2}$$
 (11)

Then for a given x, the "prediction error" is:

$$\sigma_f(x) = s\sqrt{1 + \frac{1}{N} + \frac{\left\{x - \overline{x}\right\}^2}{N\sigma_x^2}} \tag{12}$$

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$$\sigma_x^2 = \frac{\sum\limits_{n=1}^N \left\{ x_i - \overline{x} \right\}^2}{N} \tag{13}$$

The probability density function of the predicted variable \hat{y} given x is provided by the Cox et al. (2013) as:

$$P(\hat{y}|x) = \frac{1}{\sqrt{2\pi\sigma_f^2}} \exp\left\{-\frac{(y_i - \hat{y}_i)^2}{2\sigma_f^2}\right\}$$
(14)

where $\sigma_f = \sigma_f(x)$, as described above.

Finally, for a given observation-based probability distribution function (PDF) P(x) and a model-based PDF $P(\hat{y}|x)$, the PDF of the emergent constraint on y is:

$$P(y) = \int_{-\infty}^{\infty} P(\hat{y}|x)P(x)dx \tag{15}$$

3. Results

3.1. Inter-Model Variation of the Greening Effect on Ecosystem C Stock

The multi-model ensemble of global ECS was positively correlated with global mean LAI in CMIP5 (n=16, $R^2=0.80$, P<0.001; Figure 1a) from 1982 to 2005, CMIP6 (n=12, $R^2=0.95$, P<0.001; Figure 1b) from 1982 to 2014, and TRENDY (n=12, $R^2=0.96$, P<0.001; Figure 1c) from 1982 to 2019. Such a significantly positive sensitivity was found in 11 out of 16 models in CMIP5 and all models in CMIP6 and TRENDY (Figure 1d). The sensitivity of C stock to per unit of LAI changes greatly varied among the models, ranging from -0.59 to 4.35 kg C m⁻² per m² m⁻² in CMIP5, from 0.12 to 3.79 kg C m⁻² per m² m⁻² in CMIP6, and from 0.36 to 4.88 kg C m⁻² per m² m⁻² in TRENDY (Figure 1d; Table S6 in Supporting Information S1). Thus, the CMIP5 models have the largest variation of the sensitivity of C stock to per unit of LAI change (with a standard deviation of 1.50 kg C m⁻² per m² m⁻²), followed by the TRENDY (a standard deviation of 1.20 kg C m⁻² per m² m⁻²) and CMIP6 (a standard deviation of 1.11 kg C m⁻² per m² m⁻²) models (Figure 1d). The spatial distribution with greater inter-model variation of the sensitivity of C stock to LAI among the 40 models were mainly located in TRF, boreal, and shrubland regions (Figure 1e).

The inter-model variation in the sensitivity of C stock to LAI was jointly contributed by C stock and LAI. For example, among the 16 ESMs in CMIP5, the global multi-year mean of C stock ranges from 5.33 kg C m⁻² for CESM1-WACCM to 23.28 kg C m⁻² for GFDL-ESM2G (Figure S4a in Supporting Information S1), while LAI varies from 1.53 m² m⁻² for CanESM2 to 3.38 m² m⁻² for MPI-ESM1 (Figure S5a in Supporting Information S1). For the 12 ESMs in CMIP6, C stock varies from 5.36 to 65.09 kg C m⁻² (Figure S4a in Supporting Information S1), while LAI changes from 1.33 to 2.21 m² m⁻² (Figure S5a in Supporting Information S1). For the 12 DGVMs in TRENDY models, C stock varies from 7.22 to 58.58 kg C m⁻² (Figure S4a in Supporting Information S1), while LAI changes from 1.46 to 4.03 m² m⁻² (Figure S5a in Supporting Information S1). The ensemble means of global multi-year averaged LAI in CMIP6 (1.79 m² m⁻²) is much less than that for CMIP5 (2.30 m² m⁻²) or TRENDY (2.45 m² m⁻²). Only IPSL-CM5A-LR and NorESM1-ME simulated a negative correlation between LAI and C stock in CMIP5 (Figure 1d). The simulated LAI by IPSL-CM5A-LR (Figure S5b in Supporting Information S1) and the simulated C stock by CESM1-WACCM and NorESM1-ME (Figure S4b in Supporting Information S1) decreased over time.

3.2. Trace the Source of Uncertainty in Modeled Sensitivity of Ecosystem Carbon Stock to LAI

The variation in modeled sensitivity of ECS to LAI is decomposed into several traceable components' sensitivity to LAI based on the traceability analysis framework in temporal and spatial scales (1982–2005 for CMIP5; and 1982–2014 for CMIP6; 1982–2019 for TRENDY-v9; Figure 2). The terms $\partial \ln(\text{NPP})/\partial \text{LAI}$ and $\partial \ln(\tau_E)/\partial \text{LAI}$ contribute equally (50%) to the across-model variance in the sensitivity of C stock to LAI in all three MIPs

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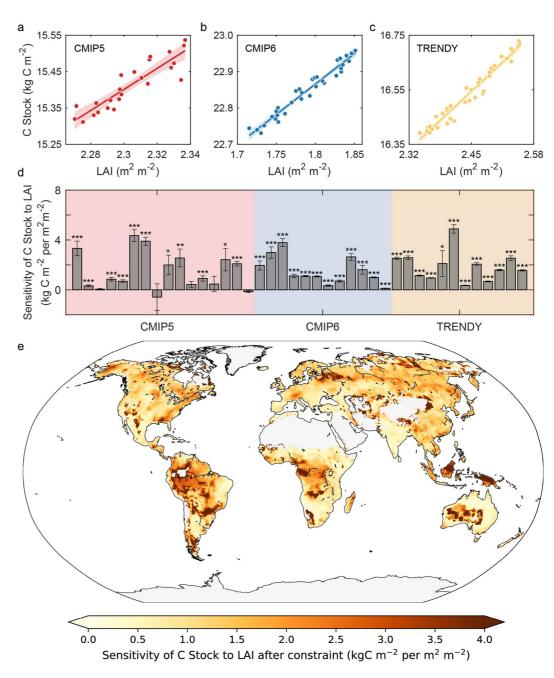


Figure 1. Comparison of the sensitivity of ecosystem carbon stock to leaf area index (LAI) among Coupled Model Intercomparison Projects and TRENDY models. (a–c) Temporal relationship between ecosystem C stock and LAI based on the model ensemble of Coupled Model Intercomparison Project (CMIP5) (a), Coupled Model Intercomparison Project Phase 6 (CMIP6) (b) and TRENDY (c). Each scatter represents an individual year of LAI and C stock for the model ensemble results and the straight lines represent the linear regression for the model ensemble with red, blue, and yellow indicating CMIP5, CMIP6, and TRENDY, respectively. Panel (d) shows the sensitivities of C stock to per unit LAI change in 40 models for the three MIPs. Each bar represents one model. The error bars indicate the mean \pm standard error (i.e., SE) of the sensitivity of C stock to LAI. The asterisks indicate the trend is statistically significant using a *t*-test, *** represents P < 0.001, ** represents P < 0.01, * represents P < 0.05. (e) The spatial distribution of standard deviation in the sensitivity of C stock to LAI among the 40 models.

based on the results of temporal dynamics. The modeled variation of $\partial \ln(\text{NPP})/\partial \text{LAI}$ can be decomposed into $\partial \ln(\text{GPP})/\partial \text{LAI}$ and $\partial \ln(\text{CUE})/\partial \text{LAI}$. The contributions of $\partial \ln(\text{GPP})/\partial \text{LAI}$ to the $\partial \ln(\text{NPP})/\partial \text{LAI}$ variation is 71% for CMIP5, 77% for CMIP6, 99% for TRENDY, and 90% for the three MIPs, while the contributions of

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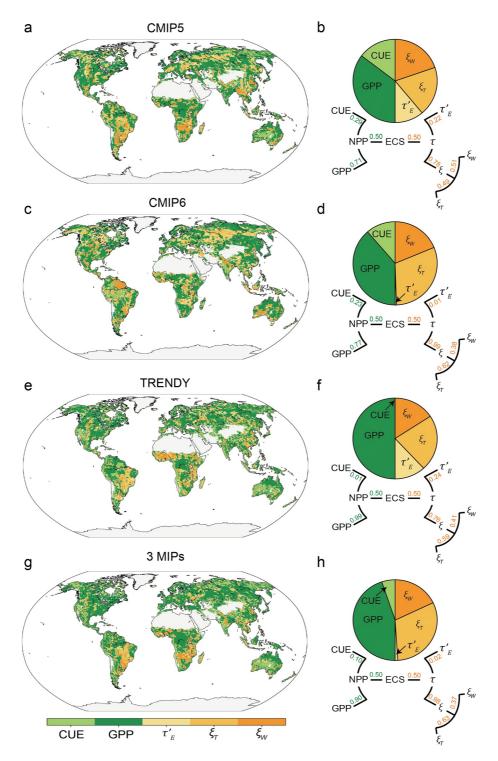


Figure 2. Variance decomposition of the sensitivity of ecosystem carbon stock to leaf area index (LAI) for Coupled Model Intercomparison Project (a, b), Coupled Model Intercomparison Project Phase 6 (c, d), TRENDY (e, f), and all the models from three MIPs (g, h). The left global maps show the spatial distribution of the dominant variable for the variation. The pie plots represent the attribution of variance for each component sensitivity to LAI. The right panels show the hierarchical decomposition of ecosystem carbon stock into its components. We first converted all the traceable components to its logarithmic form, and further quantified the contributions of each traceable component to the uncertainty of modeled sensitivity of ecosystem C stock to LAI (see Section 2.3.2). The variation of $\partial \ln(\text{ECS})/\partial \text{LAI}$ is decomposed to $\partial \ln(\text{NPP})/\partial \text{LAI}$ and $\partial \ln(\tau_E)/\partial \text{LAI}$. And then, variation of $\partial \ln(\text{NPP})/\partial \text{LAI}$ is decomposed into $\partial \ln(\text{GPP})/\partial \text{LAI}$ and $\partial \ln(\text{CUE})/\partial \text{LAI}$, while $\partial \ln(\tau_E)/\partial \text{LAI}$ could be decomposed into $\partial \ln(\tau_E)/\partial \text{LAI}$ and $\partial \ln(\xi)/\partial \text{LAI}$. $\partial \ln(\xi)/\partial \text{LAI}$ could further decomposed to $\partial \ln(\xi_T)/\partial \text{LAI}$ and $\partial \ln(\xi_W)/\partial \text{LAI}$. The symbols in the figure is: C residence time (τ), baseline C residence time (τ'_E), environmental scalar (ξ), temperature scalar (ξ_T) and water scalar (ξ_W).

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 $\partial \ln(\text{CUE})/\partial \text{LAI}$ to the $\partial \ln(\text{NPP})/\partial \text{LAI}$ variation is 29% for CMIP5, 23% for CMIP6, 1% for TRENDY, and 10% for all the global models from the three MIPs. It suggests that the variation in $\partial \ln(\text{GPP})/\partial \text{LAI}$ represents the largest uncertainty source in the modeled sensitivity of NPP to LAI. By decomposing the variation of $\partial \ln(\tau_E)/\partial \text{LAI}$ into $\partial \ln(\tau_E')/\partial \text{LAI}$ and $\partial \ln(\xi)/\partial \text{LAI}$, we found that the contribution of $\partial \ln(\tau_E')/\partial \text{LAI}$ to $\partial \ln(\tau_E)/\partial \text{LAI}$ is smaller than that of $\partial \ln(\xi)/\partial \text{LAI}$. Specifically, the contributions of $\partial \ln(\tau_E')/\partial \text{LAI}$ to the $\partial \ln(\tau_E)/\partial \text{LAI}$ variation is 22% for CMIP5, 1% for CMIP6, 24% for TRENDY, and 2% for the three MIPs, while the contributions of $\partial \ln(\xi)/\partial \text{LAI}$ to the $\partial \ln(\tau_E)/\partial \text{LAI}$ variation is 78% for CMIP5, 99% for CMIP6, 76% for TRENDY, and 98% for the three MIPs. In addition, $\partial \ln(\xi_T)/\partial \text{LAI}$ has a greater contribution to the $\partial \ln(\xi)/\partial \text{LAI}$ compared with $\partial \ln(\xi_W)/\partial \text{LAI}$ for CMIP6 (62%), TRENDY (59%), and all the three MIPs combined (63%), but a smaller contribution for CMIP5 (49%) (Figures 2b, 2d, 2f, and 2h).

We further detect each traceable component's contribution to the variation of the greening effect on ECS across models on a spatial scale. The contribution of each component was calculated for each grid. The component which made the greatest contribution to the variation in $\partial \ln(\text{ECS})/\partial \text{LAI}$ was identified as the dominant uncertainty source (Figures 2a, 2c, 2e, and 2g). The results showed that $\partial \ln(\text{GPP})/\partial \text{LAI}$ was the main uncertainty contributor whether for CMIP5 (at 62.0% vegetated land grid), CMIP6 (60.1%), TRENDY (56.3%), or all three MIPs (60.4%). The spatial distribution of the other contributors varies greatly among the three MIPs. For CMIP5 and CMIP6, the contribution of $\partial \ln(\tau_E')/\partial \text{LAI}$ was mainly located in central Siberia; while for TRENDY and the ensemble of three MIPs, the contribution of $\partial \ln(\tau_E')/\partial \text{LAI}$ was mainly located in Brazil regions.

3.3. Observational-Based Greening Effect on Gross Primary Production

We examined the inter-annual covariance between satellite-derived GPP (Figure S6 in Supporting Information S1) and LAI (Figure S7 in Supporting Information S1) in each vegetated land grid cell between 1982 and 2014. About 68.8% of vegetated land grids showed a significant positive correlation between GPP and LAI (Figure 3; Figure S8 in Supporting Information S1). The mean sensitivity of GPP to per-unit change of LAI (i.e., the slope of GPP-LAI correlations) across all land grids was 297.91 ± 267.41 g C m⁻² yr⁻¹ per m² m⁻², which was significantly larger than zero (*t*-test, P < 0.05; Figure 3b). Then, we classified the observed sensitivity of GPP to LAI change into eight biomes based on the MODIS/TERRA-AQUA land cover product (MCD12C1; Friedl et al., 2010) Collection 5.1 (Figure S1 in Supporting Information S1). Among the eight biomes, as shown in Figure 3c, shrubland had the largest sensitivity of GPP to per unit change in LAI (482.01 \pm 322.33), followed by grassland (403.11 \pm 292.10), cropland (379.72 \pm 244.75), tundra (250.25 \pm 204.63), BOF (185.74 \pm 126.25), TEFs (154.93 \pm 112.21), wetlands (148.34 \pm 140.35), and TRFs (14.16 \pm 93.09 g C m⁻² yr⁻¹ per m² m⁻²).

3.4. Emergent Constraints of Greening Effect on Ecosystem Carbon Stock

Among the 40 models, there was a strong positive correlation between the sensitivities of GPP and C stock to LAI ($R^2 = 0.63$, P < 0.001; Figure 4a) since the 1980s. Combining the satellite-derived sensitivities of GPP to LAI and the model-based sensitivities of C stock to LAI, we constrained the magnitude of the greening effect on C stock. The constrained greening effect on C stock (0.75 ± 0.46 kg C m⁻² per m² m⁻²) was lower than the model ensemble mean value (1.63 ± 1.29 kg C m⁻² per m² m⁻²), with a reduction of 64.54% in the range of uncertainty (Figure 4b). Additional emergent constraint analyses based on the rigorous matches of temporal duration from 1982 to 2005 in all model simulations and satellite-derived observations were performed to overcome the possible bias due to temporal variations (Winkler, Myneni, & Brovkin, 2019). The constrained results (0.79 ± 0.50 kg C m⁻² per m² m⁻²) by strict temporal period matching (Figure S9 in Supporting Information S1) were slightly higher than those with non-strict time-period matching (0.75 ± 0.46 kg C m⁻² per m² m⁻²). A further sensitivity analysis of the emerged relationship to the number of included models showed that the emerged relationship is not sensitive to the number of included models (Figure S10 in Supporting Information S1). Additionally, considering a similar result produced within the same modeling center (Todd-Brown et al., 2013, 2014), the emerged relationship which includes more models can generate a more robustness result (Figure S10 in Supporting Information S1).

The greening effect on C stock varied among the three MIPs (Figures 1d and 2e and Figure 4a). TRENDY had the highest greening sensitivity of C stock (1.91 \pm 1.20 kg C m⁻² per m² m⁻²), followed by CMIP6 (1.54 \pm 1.11 kg C m⁻² per m² m⁻²) and CMIP5 (1.48 \pm 1.50 kg C m⁻² per m² m⁻²). The constrained sensitivities

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5000

-500

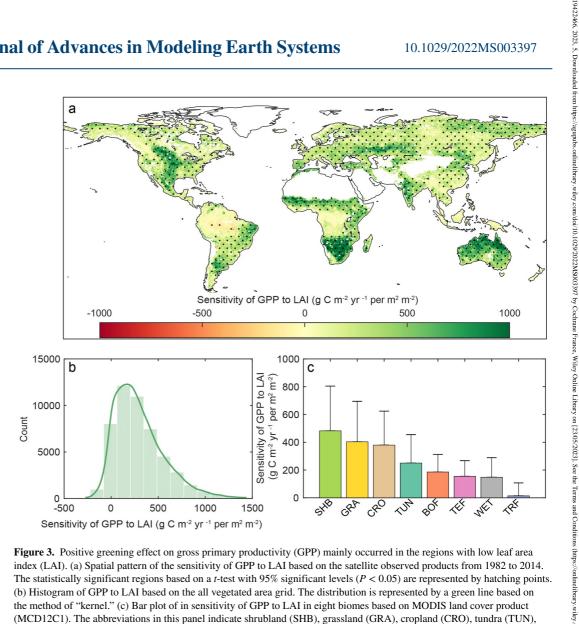


Figure 3. Positive greening effect on gross primary productivity (GPP) mainly occurred in the regions with low leaf area index (LAI). (a) Spatial pattern of the sensitivity of GPP to LAI based on the satellite observed products from 1982 to 2014. The statistically significant regions based on a t-test with 95% significant levels (P < 0.05) are represented by hatching points. (b) Histogram of GPP to LAI based on the all vegetated area grid. The distribution is represented by a green line based on the method of "kernel." (c) Bar plot of in sensitivity of GPP to LAI in eight biomes based on MODIS land cover product (MCD12C1). The abbreviations in this panel indicate shrubland (SHB), grassland (GRA), cropland (CRO), tundra (TUN), boreal forest (BOF), temperate forest (TEF), wetland (WET), and tropical forest (TRF), respectively. The error bars in this panel indicate the standard deviation.

200 **g**

of C stock to LAI were 1.22 ± 0.30 , 0.28 ± 0.69 , and 0.29 ± 0.70 kg C m⁻² per m² m⁻² for TRENDY, CMIP6, and CMIP5, respectively. The uncertainty spread was reduced by 75%, 54%, and 38% for TRENDY, CMIP5, and CMIP6, respectively (Figure S11 in Supporting Information S1).

3.5. Spatial Variation of Greening Effect on Ecosystem Carbon Stock

1000

Sensitivity of GPP to LAI (g C m⁻² yr ⁻¹ per m² m⁻²)

Using the emergent constraint approach, we then constrained the greening effect on ECS on a global scale for each pixel. Compared with the multi-model ensemble means (Figure 4c), the constrained results show reduced spatial heterogeneity and a dampened greening effect on ECS, with 87.5% of global vegetated pixels showing a positive greening effect (Figure 4d). Our results showed that the constrained greening effect on ECS average over only vegetated grid cells of 0.40 ± 0.71 is 44% lower than that of the multi-model ensemble mean of 0.72 ± 0.81 kg C m⁻² per m² m⁻². Then, we classified the greening effect on ECS before and after constraint into eight biomes. As shown in Figure S12 in Supporting Information S1, the constrained greening effect on C stock decreases most in TRFs with 87.9%, followed by tundra (61.0%), TEFs (53.7%), BOF (44.0%), wetlands (33.7%), cropland (23.9%), grassland (21.6%), and shrubland (6.4%). We examined the constrained spatial pattern by applying multiple lines of data sets, such as eddy-flux observations and AGB C based on harmonized VOD data. First, we examined the relationship between annual NEP and annual mean LAI using flux-tower measurements from FLUXNET at each flux site with at least 5 years of data (Figure S13 and Table

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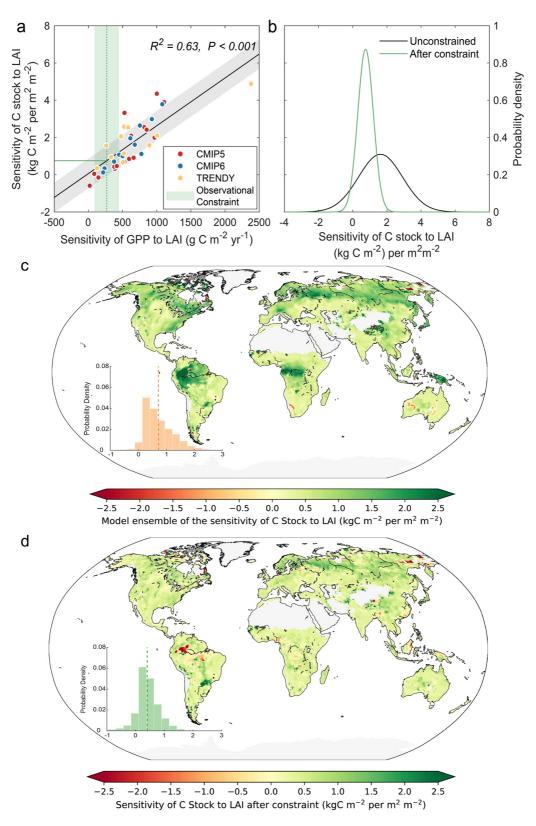


Figure 4.

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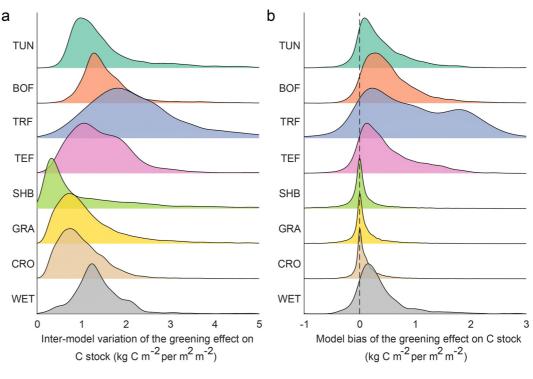


Figure 5. Ridgeline probability density plots show the comparisons of the inter-model variation (a) and model bias (b) of the greening effect on C stock among different biomes. Inter-model variation is represented by the standard deviation of the greening effect on C stock among all the 40 global models used in this study, while the model bias used here comes from the difference between multi-model ensemble results and the constrained results.

S4 in Supporting Information S1). Results show that the mean slope of all sites is 0.23 ± 0.11 kg C m⁻² yr⁻¹, which significantly differs from zero (P < 0.001) (Figure S14a in Supporting Information S1). Although trends are rarely significant (13.4%) and highly variable across the sites, 63.41% of all sites showed a positive correlation between NEP and LAI (Figure S14b in Supporting Information S1). In addition, the mean values of greening effects on NEP were 0.33 (n = 24), 0.25 (n = 51), and -0.27 (n = 7) kg C m⁻² yr⁻¹ for the LAI ranges of <1.5, from 1.5 to 3, and >3, respectively (Figure S14c in Supporting Information S1; all P < 0.05). Additional greening effect analyses based on the AGB C generated from VOD data show the most extensive significant positive greening effect (t-test, P < 0.05) in 32% of vegetated land grids (Figure S15 in Supporting Information S1). The spatial pattern of the greening effect on AGB is similar to that of GPP. The results show that shrubland had the highest sensitivity of AGB C to per unit of LAI change with 5.05 \pm 4.79 Mg C ha⁻¹ per m² m⁻², but TRFs had a negative response with -0.29 ± 4.34 Mg C ha⁻¹ per m² m⁻² (Figure S15c in Supporting Information S1).

We further compared the difference between the inter-model variation and model bias for the eight biomes (Figure 5). The results showed that the inter-model variation and the model of the greening effect on C stock have a good spatial consistency. Especially, the TRFs have the highest inter-model variation and model bias. Tundra, BOF, TEF, and wetland have mediated inter-model variation and model bias, while shrubland, grassland, and cropland have low inter-model variation and model bias (Figure 5).

Figure 4. Emergent constraint on the sensitivity of ecosystem carbon stock to leaf area index (LAI). (a) The correlation between the sensitivity of gross primary productivity (GPP) to LAI and the sensitivity of C stock to LAI. The green dash-line and shading areas indicate the mean and the range of the observational uncertainty of the sensitivity of GPP to LAI, respectively. The black straight line shows the linear regression relationship across the three MIPs ensemble together with the prediction error (gray shading areas). The green horizontal line shows the constrained sensitivity of C stock to LAI. Each scatter represents a model, red, blue, and yellow indicating Coupled Model Intercomparison Project, Coupled Model Intercomparison Project Phase 6, and TRENDY, respectively. (b) The probability distribution function for the LAI sensitivity of C stock. The black line indicates the Gaussian distribution of the original unweighted model ensemble of C stock to LAI. The green line is the satellite-derived data-constrained probability distribution of the sensitivity of C stock to LAI. Spatial distribution of the multi-model ensemble means of the sensitivity of C stock to LAI (c) and the constrained sensitivity of C stock to LAI (d). The insert panel shows the corresponding frequency distribution of the sensitivity, and the dashed line represents the mean value.

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4. Discussion

This study shows a positive impact of global land greening on terrestrial ECS during the past three decades. This finding is consistent with recent empirical evidence on the associated increases in vegetation biomass (Fang et al., 2014; C. Zhu & Xia, 2020) and soil C stock (G. Zhou et al., 2006; J. Zhu et al., 2020) in terrestrial ecosystems since the 1980s. The enhanced net primary productivity with land greening can be an important mechanism underlying the increase in terrestrial ECS (Figure S16a in Supporting Information S1). Based on the satellite data, this study constrains the positive greening effect on global terrestrial ECS to 0.75 ± 0.46 kg C m⁻² per unit LAI increase (Figure 4a). Hajima et al. (2014) used the sensitivity of GPP to LAI (GPP/LAI) as a proxy to represent leaf-level photosynthesis. Similarly, the ECS divided by LAI (ECS/LAI) could represent leaf-level carbon sequestration. These results suggest that increases in leaf area have led to a proportional increase in photosynthesis and C stock. The constrained results in Figure 4a represent leaf-level carbon sequestration, which is overestimated by most global models. Additionally, all models in the three MIPs have shown a shortening of ecosystem C residence time during the past three decades. Thus, the positive greening effect on ECS (Figure 1d) mainly results from the considerable increase in net primary productivity rather than the decrease in ecosystem C residence time (Figure S16 in Supporting Information S1).

The magnitude of the greening effect on C stock has a sizable inter-model variation, ranging from -0.59 to 4.88 kg C m⁻² per m² m⁻² (Figure 1d) among three MIPs. Compared with CMIP5 and CMIP6, the TRENDY project shows a higher constrained sensitivity of C stock to LAI with less uncertainty (Figure S11 in Supporting Information S1). The model simulations in the TRENDY project are forced by the same set of climate reanalysis data sets, which avoids the uncertainty from climate models. Most models have simulated a higher positive impact of land greening on ECS than the data-constrained result (Figure 4a). Some missing processes in current models could lead to the highly variable greening effect on C stock. First, nearly half of the models (20 out of 40) used in this study do not consider nutrient limitations (Tables S1–S3 in Supporting Information S1), especially the widespread phosphorus limitation (38 out of 40 models) throughout the global terrestrial ecosystems (Du et al., 2018; Elser et al., 2007; Wieder et al., 2015) (The simulated greening effect on C stock was 1.84 ± 1.37 for only C models, 1.42 ± 1.20 for CN models, and 1.31 ± 0.91 kg C m⁻² per m² m⁻² for CNP models; Table S6 in Supporting Information S1). Second, one critical limiting factor for land carbon uptake is soil moisture (Green et al., 2019; Humphrey et al., 2018, 2021), but many ESMs have high degrees of uncertainty in plant hydraulic processes, such as the xylem embolism (Franks et al., 2017; Konings et al., 2017) and drought legacy effect (Anderegg et al., 2015) or lack representation of such processes. Third, most ESMs used in this study did not represent the shifts in plant community structure and composition (Medlyn et al., 2015), which could weaken the positive response of ecosystem productivity and enhance the ecosystem turnover rates to increasing atmospheric CO₂ (Kolby Smith et al., 2016), nitrogen deposition (Bonan & Doney, 2018; Drewniak & Gonzalez-Meler, 2017; Wei, Xia, Wang, et al., 2022), and climate change (Richardson et al., 2018). Lastly, some global models used here did not include the processes of disturbance (e.g., deforestation and fire, Table S3 in Supporting Information S1, Figure S17 in Supporting Information S1), which reduce ECS and increase land C losses (Achat et al., 2015; Walker et al., 2019). Those models thus likely have overestimated the greening effect on ECS. An additional reason for the simulated overestimation in most models is that some satellite-derived GPP data sets do not incorporate increasing CO₂ on the LUE in C3 vegetation. Thus, we selected the GPP data sets accounting for the CO₂ fertilization effect (i.e., EC-LUE and P-model) to validate the constrained results. The constrained greening effect on C stock shows a higher value $(0.82 \pm 0.58 \text{ kg C m}^{-2} \text{ per m}^2 \text{ m}^{-2} \text{ based})$ on EC-LUE GPP and 1.15 ± 0.68 kg C m⁻² per m² m⁻² based on P-model GPP) when considering the CO₂ effect on LUE in GPP data sets (Figure S18 in Supporting Information S1). These findings indicate that models and GPP data sets need improvements for a more accurate constrained effect of global land greening on terrestrial C stock.

The large uncertainties in the simulated greening effect on C stock almost equally stem from the sensitivities of residence time and NPP to LAI in the three MIPs (Figures 2b, 2d, 2f, and 2h). By decomposing the modeled residence time into its components, we found that the sensitivities of environmental scalars to LAI have a larger uncertainty contribution than that in baseline C residence time (78% for CMIP5, 99% for CMIP6, 76% for TRENDY, and 98% for 3MIPs). This may be due to the changes in vegetation cover altering the radiative of the land surface (Duveiller et al., 2018) and thus further influencing the energy (Alkama et al., 2022; Forzieri et al., 2017; Zeng et al., 2017) and water vapor (Yuan et al., 2019) allocation via biophysical processes. Additionally, the discrepancies in disturbance processes (such as fire and land use change) among global models are an important challenge. In addition, our results also show the sensitivity of GPP to LAI is the largest contributor to this uncertainty in more than 60% of the vegetated grids (Figure 2g). One possible reason is the difference

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across canopy structures used by each model in scaling up from leaf to canopy, such as big-leaf, two-leaf, or multiple-layer (Cui et al., 2019; Li et al., 2018). Furthermore, we recognized that in the regions where other components (except for GPP) are the main contributors, such as southeastern Africa, central and southern South America, central Australia, etc., using the satellite-observed sensitivity of GPP to LAI as constrain may cause some systemic errors inevitably.

This study shows that the greening effect on ECS is highly variable across different regions or biomes. The positive and the most significant impact of vegetation greening on ECS mainly occurred in the areas with low LAI (Figure 3, Figures S14 and S15 in Supporting Information S1). This is because areas of low LAI (often also water-limited) are generally more sensitive to elevated CO₂ (Z. Zhu et al., 2016). Elevated CO₂ can decrease stomatal conductance (Field et al., 1995) and reduce water loss and leaf-scale transpiration (Ukkola et al., 2016). While in high LAI regions, especially in TRFs, further increases in LAI can negatively influence ecosystem C uptake and thus affect C stock (Figures S14c and S15 in Supporting Information S1). Such a spatial pattern indicates that the global greening effect on land C stock strongly depends on the spatial distributions of the canopy structure (e.g., LAI; Figure S7 in Supporting Information S1). Our results from satellite-derived data sets indicate a minor greening effect on GPP in tropical and high LAI regions (Figure 3). The eddy-tower measurements support such patterns in the tropics, which have shown that shaded leaves (Li et al., 2018) and respiratory C losses (Srinivasan et al., 2017) increase rapidly with increasing LAI, resulting in the reverse of the initial positive greening effects on ecosystem carbon uptake (see Figure S14, for the flux sites with LAI larger than 3 m² m⁻², carbon sinks are negatively sensitive to LAI increase). Another possible reason is that many models overestimate the response of LAI to elevated CO₂ (Kolby Smith et al., 2016), especially in the mature closed-canopy forest (Duursma et al., 2016; Norby et al., 2003). Moreover, the greening effect on AGB C shows that the positive greening effects are mainly located in shrubland regions with less model bias, but the adverse effects are in tropical areas with the largest model bias (Figure S15 in Supporting Information S1, Figure 5b). We also acknowledged that the biome classification based on MODIS land cover product might lead to some uncertainty inevitably because of the mismatch between the plant functional types used in the model and observations in some grids.

5. Conclusions

In summary, the constrained results in this study reveal a positive greening effect on ECS on a global scale since the 1980s. However, the ability of vegetation to absorb CO₂ and mitigate climate change may not be as optimistic as most models predicted. In addition, there is a high spatial heterogeneity of greening impact on ECS globally. The pervasive and positive greening effect in low-LAI ecosystems implies a land C stock shift toward high-latitude regions in the North Hemisphere under further vegetation greening. This study also reveals a great uncertainty of the positive greening effects on land carbon stock in the current generation of global models, mainly caused by the ecosystem productivity at the leaf level. This finding emphasizes the need to strengthen the simulation of vegetation canopy structure in land surface models and ESMs. Furthermore, the biome-based results reveal that the TRF regions have the highest inter-model variation and model bias. Our study identifies the uncertainty source and provides constrained estimates of the greening effect on ecosystem carbon stock. These findings underscore the importance of linking vegetation structural changes to ECS dynamics in the projections of future climate-carbon cycle feedbacks.

Data Availability Statement

The model simulations of CMIP5 and CMIP6 were obtained from the Earth System Grid Federation (ESGF) data archive (https://esgf-node.llnl.gov/projects/esgf-llnl/). The raw model output of TRENDY is available at the following sftp site: trendy-v9@trendy.ex.ac.uk. Access will be contacting Stephen Sitch (S.A.Sitch@exeter.ac.uk). All data used in this research are publicly available from the cited literature and the following links: The GIMMS LAI are available at https://daac.ornl.gov/VEGETATION/guides/Mean_Seasonal_LAI.html. The GLOBMAP LAI data set is available at https://zenodo.org/record/4700264#.YyLcBnbP0b8. The GLASS LAI data set is available at https://www.glass.umd.edu/Download.html. The EC-LUE GPP data are available at https://doi.org/10.6084/m9.figshare.8942336.v3. The GPPEnh data sets can be obtained online at https://daac.ornl.gov/cgi-bin/dsviewer.pl?ds_id=1789. The GPP based on the P-model is available at https://zenodo.org/record/1423484#.YyLesHbP1Gq. The FLUXNET2015 data set is available online at https://fluxnet.org/data/fluxnet2015-dataset/. The code and

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metadata used in this study for associated analysis and producing the main figures can be accessed at https://doi.org/10.6084/m9.figshare.21118711.v1.

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