



HAL
open science

Analysis of the microstructure and morphology of disordered kaolinite based on the particle size distribution

Yongjie Yang, Maguy Jaber, Laurent Michot, Baptiste Rigaud, Philippe Walter, Lucie Laporte, Kenan Zhang, Qinfu Liu

► To cite this version:

Yongjie Yang, Maguy Jaber, Laurent Michot, Baptiste Rigaud, Philippe Walter, et al.. Analysis of the microstructure and morphology of disordered kaolinite based on the particle size distribution. *Applied Clay Science*, 2023, 232, pp.106801. 10.1016/j.clay.2022.106801 . hal-03972785

HAL Id: hal-03972785

<https://hal.science/hal-03972785v1>

Submitted on 8 Jan 2025

HAL is a multi-disciplinary open access archive for the deposit and dissemination of scientific research documents, whether they are published or not. The documents may come from teaching and research institutions in France or abroad, or from public or private research centers.

L'archive ouverte pluridisciplinaire **HAL**, est destinée au dépôt et à la diffusion de documents scientifiques de niveau recherche, publiés ou non, émanant des établissements d'enseignement et de recherche français ou étrangers, des laboratoires publics ou privés.



Distributed under a Creative Commons Attribution - NonCommercial 4.0 International License

1 **Analysis of the microstructure and morphology of**
2 **disordered kaolinite based on the particle size**
3 **distribution**

4 **Yongjie Yang^{a,b}, Maguy Jaber^{b,*}, Laurent J. Michot^c, Baptiste**
5 **Rigaud^d, Philippe Walter^b, Lucie Laporte^b, Kenan Zhang^a, Qinfu**
6 **Liu^{a,*},**

7 ^a School of Geoscience and Surveying Engineering, China University
8 of Mining and Technology (Beijing), 100083, Beijing, China

9 ^b Laboratoire d'Archéologie Moléculaire et Structurale (LAMS), UMR
10 8220, CNRS, Sorbonne Université, Case Courrier 225, UPMC 4
11 Place Jussieu, 75005, Paris, France

12 ^c Laboratoire PHENIX, UMR 8234, CNRS, Sorbonne Université,
13 UPMC, 4 Place Jussieu, 75005, F-75252 Paris, France

14 ^d CNRS Institut des Matériaux de Paris Centre (FR2482), Sorbonne
15 Université, 4 Place Jussieu, 75005, Paris, France

16

17 * Corresponding author

18 Email: maguy.jaber@upmc.fr (M. Jaber), lqf@cumtb.edu.cn (Q. Liu)

19

20 **Abstract:**

21 The defected kaolinite is common in sedimentary kaolin, but its
22 structure and disorder mechanism remains unclear. In this paper, three
23 different size distributions of kaolinite particles: 0~300 nm, 200~1000
24 nm, and 500~2000 nm, were separated and classified from high-defect
25 sedimentary kaolin, collected from Quaternary sedimentary kaolin in
26 Guangxi Zhuang Autonomous Region of China. The three particle
27 sizes of kaolinite have been monitored by particle size analyzer, high-
28 resolution magic-angle spinning NMR (MAS NMR), small and wide-
29 angle X-ray scattering (SAXS and WAXS), X-ray diffraction (XRD),
30 BET surface analyzer, nitrogen gas-adsorption, scanning electron
31 microscopy (SEM), and high-resolution transmission electron
32 microscopy (HRTEM). The results revealed that the smallest particle
33 size of kaolinite exhibited high degrees of broken Al-O-Al and Si-O-Si
34 bonds, more disordered structure, and elevated specific surface area.
35 The degree of the structural disorder of kaolinite increases as the
36 particle size decreases, which can be attributed to the more aluminum
37 substituted for silicon in tetrahedron sheet in the finer kaolinite crystal,
38 compared to the coarser kaolinite crystals.

39 **Keywords:** Disordered kaolinite, Particle size distribution,
40 Microstructure, NMR, SAXS, WAXS

41 **1. Introduction**

42 Kaolinite, $\text{Al}_2\text{Si}_2\text{O}_5(\text{OH})_4$, with the unit cell parameters: $a_0=0.517$ nm,
43 $b_0=0.898$ nm, $c_0=0.735$ nm, $\alpha=91.68^\circ$, $\beta=105.12^\circ$, $\gamma=89.75^\circ$ (Richard
44 and Rendtorff, 2019), is a 1:1 layered dioctahedral hydrous
45 aluminosilicate that generally exhibits structural order and disorder in
46 nature (Awad et al., 2018; Bish, 1993; Brindley et al., 1986). Silicon
47 atoms are organized as ditrigonal rings of silicon-oxygen tetrahedra
48 and form the tetrahedral sheet. In contrast, aluminum atoms are
49 organized as aluminum-oxygen octahedra in the octahedral sheet that
50 also contains one-third of empty octahedral voids, which maintains the
51 electrical balance of the whole structure. Kaolinite is then an
52 uncharged structure, even if some samples bear a small permanent
53 negative charge due to the existence of isomorphic substitutions in
54 either the tetrahedral or octahedral sheet. Numerous authors have
55 debated the origin of structural disorder in kaolinite. It was supposed
56 that disorder in kaolinite is caused by substituting cations within the
57 lattice, such as replacing Al^{3+} with Fe^{3+} in alumina octahedral sheet or
58 Si^{4+} substituted by Al^{3+} in silica tetrahedral sheet (Brindley et al., 1986;
59 Newman et al., 1994). The second reason for the structural disorder

60 was suggested that the translation between adjacent layers, which is
61 related to the pseudo-mirror plane, coincident with the long diagonal
62 of the unit cell (Plançon et al., 1989). The third possibility relies on the
63 existence of interlayer shifts between two adjacent layers. There are
64 two types of shifts: t_1 and t_2 , where t_1 is a shift of ($\sim a/3$) and t_2 (\sim
65 $a/3+b/3$). Such alternating t_1 and t_2 shifts were observed by High-
66 Resolution Transmission Electron Microscopy (HRTEM) by Kogure et
67 al. (2010) in sedimentary kaolinite from Brasil. These three types of
68 defects broaden the reflections observed by X-Ray Diffraction (XRD).
69 They particularly affect the (02*l*) and (11*l*) sequences (20-23° 2θ, using
70 CuKα), and the (13*l*) and (20*l*) sequences (35-40° 2θ, using CuKα)
71 (Aparicio, 1999).

72 The question of the link between kaolinite particle size distribution
73 (PSD) and the extent of disorder in the structure remains somehow
74 open. A better understanding of the potential relationships between
75 disorder and particle size would provide relevant information to better
76 assess the geological processes that control the composition of
77 kaolinite formed during weathering and, therefore, yield data about the
78 geological background at the time of mineralization (Varajão et al.,

79 2001).

80 The particle size of kaolinite platelets is generally $< 2 \mu\text{m}$ (equivalent
81 sphere diameter). Suspensions of kaolinite fine particles display low
82 viscosity (Rand and Melton, 1977), better rheological behavior (Gupta
83 et al., 2011; Mporu et al., 2003), and anti-flocculation performance
84 (Du et al., 2010). It is worth pointing out that the size of kaolinite
85 particles used commercially is generally less than $2 \mu\text{m}$, such as paper
86 coating kaolin, where 70~90% of the particles used are less than $2 \mu\text{m}$
87 (Morsy et al., 2014; Silva et al., 2009). Even finer particles less than 1
88 μm used for high-performance thermal shock resistance ceramics
89 (Kobayashi et al., 2000), and 300 nm for high reinforcement of
90 polymer filler (Zhang et al., 2021a; Zhang et al., 2010). Most
91 advanced applications of kaolinite derive from its versatility in terms
92 of particle size, morphology, and microstructure (Jaber et al., 2018;
93 Liu et al., 2021; Yang et al., 2020; Zhang et al., 2017; Zhang et al.,
94 2021b). Microstructural features play a significant role in determining
95 some of the physical and mechanical properties of kaolinite (i.e.,
96 thermal resistance and deformation indices). Still, until now, little
97 research has been carried out to investigate the effect of kaolinite PSD

98 on microstructure and micromorphology. One of the main reasons is
99 likely related to the difficulties associated with controlling natural
100 kaolinite particle size.

101 Wet sieving and pipette methods are the primary ways to assess the
102 particle size distribution of clay minerals (Konert and Vandenberghe,
103 1997; Wen et al., 2002; Yang et al., 2015), but mainly suitable for sizes
104 ranged from 50 to 2000 μm . For smaller particles and especially for
105 particles lower than 2 μm , the sedimentation-based pipette method
106 (SPM) is commonly used. The pipette method defines the particles as
107 the size of a sphere that settles in the liquid at the same sedimentation
108 velocity, supported by Stokes' law (Cheng and Schachman, 1955;
109 Yang et al., 2015). However, based on Stokes' law, it will take 180 h to
110 separate the $< 1 \mu\text{m}$ part for a kaolinite suspension with a viscosity less
111 than 5.7 $\text{mpa}\cdot\text{s}$. It is time-consuming for the ultrafine part, and the
112 results are increasingly unreliable because of the effect of Brownian
113 motion on the rate of sedimentation (Eshel et al., 2004).

114 Mechanical grinding, such as ball and planetary milling, was used to
115 generate ultrafine kaolinite particles (Hamzaoui et al., 2015; Makó et
116 al., 2001; Suraj et al., 1997). However, elevated grinding times will

117 destroy the kaolinite structure. Another option for decreasing particle
118 size is to perform sonication, however, delamination induced by
119 ultrasonics increases the $(-1/3a+1/3b)$ translation disorder (Franco et
120 al., 2004a), and the Hinkley Index (HI) correspondingly decreases
121 from 1.19 to 0.53 (Franco et al., 2004b). Although grinding and
122 ultrasonication can yield very fine kaolinite particles, the products bear
123 numerous non-natural disorder structures. Consequently, such
124 treatments are not adequate for studying the disordered structure of
125 natural ultrafine kaolinite.

126 Bearing this in mind, the objective of the present paper is three-fold: (1)
127 to separate three classes of kaolinite particles under 2 μm , to yield
128 “ultrafine” (< 300 nm), “intermediate” (200~1000 nm), and “coarse
129 particles” (500~2000 nm); (2) to examine the microstructure and
130 micromorphology of the three particle sizes of kaolinite; (3) to
131 postulate a mechanism of the structural disorder and verify it through
132 NMR and XRD experiments.

133 **2. Materials and methods**

134 **2.1 Materials**

135 The sample used for this study was a high-defect (“poor-crystallized”)

136 kaolinite with a Hinckley Index of 0.175, collected from Quaternary
137 sedimentary kaolin in Guangxi Zhuang Autonomous Region, China.
138 Before particle size classification, the original kaolin sample consisted
139 of 82.6% kaolinite, 12.3% quartz, 1.1% anatase, 0.6% rutile, 1.4%
140 illite, and 2% other minerals (**Fig. 1**) as derived from Rietveld
141 refinement of X-Ray Diffractograms.

142 The chemical composition of the raw kaolin sample was measured by
143 X-ray Fluorescence (XRF), and the result is summarized in **Table 1**. It
144 can be noted that the percentages of the main elements (Al_2O_3 and
145 SiO_2) slightly deviate from the theoretical one predicted by the
146 structural formula, confirming the relatively low purity of the kaolin
147 sample inferred from XRD experiments.

148 **2.2 Particle size classification of kaolinite**

149 A series of settlement tests were carried out to obtain particle sizes
150 under 2 μm . Kaolin (solid content: 20%) was dispersed in water with
151 1%wt sodium polyacrylate to obtain a slurry (Yang et al., 2020). The
152 settling velocity v of the kaolinite in the kaolin slurry can be calculated
153 according to Stokes' law:

$$154 \quad v = \frac{gD^2(\rho_1 - \rho_2)}{18\mu} \quad (1)$$

155
$$v = \frac{H}{t} \quad (2)$$

156
$$D = \sqrt{\frac{18\mu H}{gt(\rho_1 - \rho_2)}} \quad (3)$$

157 where v is the settling velocity (mh^{-1}), g is the acceleration due to
158 gravity (9.80ms^{-2}), D is the equivalent sphere diameter of kaolinite (m),
159 ρ_1 is the density of kaolinite (kgm^{-3}), ρ_2 is the density of the liquid, and
160 μ is the viscosity of the slurry system ($\text{kgm}^{-1}\text{s}^{-1}$). H is the settling
161 distance (m), and t is the settling time (h). Based on equation (1), the
162 particle size of kaolinite is related to the settling time in the same
163 viscosity system. Six barrels of kaolin slurry were prepared, and the
164 settling time were fixed at 10 cm/1 h, 10 cm/2 h, 10 cm/3 h, 10 cm/5 h,
165 10 cm/10 h, and 10 cm/15 h (**Fig. S2 Supplementary material**).

166 All supernatants were collected by the siphon method, and part of it
167 was kept for further analysis.

168 For size sorting kaolinite, vacuum filtration was performed as follows:

169 first, a 200 nm membrane was used to filter the $< 2 \mu\text{m}$ kaolin slurry.

170 The filtered solution was referred to as A1. The solids on the

171 membrane were recovered, mixed with an additional aqueous solution,

172 and dispersed again by stirring for 30 min. This process was repeated

173 three times to completely separate the particles from the suspension. A

174 450 nm membrane was then used to repeat the same filtering process.
175 The filtered solution was referred to as A2, while the solids remaining
176 on the membrane were referred to as A3. These three particle size
177 fractions, A1, A2, and A3, were finally oven-dried at 105 °C for 24 h.

178 **2.3 Sample characterization**

179 The particle size distribution of kaolinite was evaluated using a laser
180 particle size analyzer (Mastersizer 2000, Malvern, UK). Prior to
181 analysis, the sample was sonicated for 5 minutes. Additional
182 ultrasound was applied for 1 min in the injection cycle, whereas the
183 total test time was 1~2min.

184 Small- and Wide-angle X-ray Scattering (SAXS and WAXS)
185 experiments were performed to determine the microscale structure of
186 kaolinite platelets. Aqueous suspension samples were sealed in
187 cylindrical glass capillaries of 1.0 ± 0.1 mm diameter (Glas-Technik &
188 Konstruktion, Germany). SAXS patterns were recorded at the SWING
189 beamline of the SOLEIL synchrotron (Orsay, France) using fixed
190 energy of 16keV ($\lambda=0.777$ Å) with two samples-to-detector distances
191 of 0.5 m and 6 m. Bidimensional patterns were recorded on an Eiger
192 detector. The radial integration of the scattering patterns yielded the

193 evolution of the scattered intensity versus scattering vector q , where q

194 is :

$$195 \quad q = \frac{4\pi \sin \theta}{\lambda} \quad (4)$$

196 where: λ is the X-ray wavelength and θ is half the scattering angle. In

197 such conditions, the explored q range extends from 0.0008 \AA^{-1} to 2 \AA^{-1} .

198 XRD patterns of unoriented powders were measured using an X-ray

199 diffractometer (D8 Advance Bruker-AXS, Germany) with $\text{CuK}\alpha$

200 radiation ($\lambda=1.5405 \text{ \AA}$). The diffraction patterns were tested over the

201 range of $2.5\text{-}70^\circ$ (2θ) with a scan rate of $1^\circ/\text{min}$.

202 Brunauer–Emmet–Teller (BET) specific surface areas were deduced

203 from the nitrogen adsorption isotherm at 77 K using the BET surface

204 analyzer (Micromeritics ASAP 2460, U.S.A.). Before measurement,

205 each sample was evacuated under vacuum at 473 K for 2 h. The BET

206 surface area, total pore volume, average pore diameter, particle size,

207 pore width distribution, and micropore area are obtained from the

208 adsorption isotherms. The pore size distribution was acquired from the

209 desorption branch of the isotherm, supported by Kelvin's equation

210 (Kuila and Prasad, 2013).

211 The morphology of the samples was examined by scanning electron

212 microscopy (SEM, Hitachi SU8020, Japan) and high-resolution
213 transmission electron microscopy (HRTEM, JEOL-2010, 200 kV
214 Japan). Before TEM examination, the powder samples were dispersed
215 in ethanol solvent, and a drop of the suspension was deposited onto a
216 carbon film-coated copper grid to let the kaolinite particles precipitate
217 homogeneously. Selected-area electron diffraction (SAED) was
218 recorded with a point resolution of 0.18 nm.

219 ^1H , ^{29}Si , and ^{27}Al Magic angle spinning (MAS) nuclear magnetic
220 resonance (NMR) spectra were acquired on a Bruker Advance 500
221 spectrometer operating at $\Omega_L = 500.50$ MHz (^1H) and 99.43 MHz (^{29}Si)
222 with a 2.5 mm and 4 mm H-X MAS probe. Chemical shifts were
223 calibrated by the signal position of adamantane (1.85 ppm) for ^1H and
224 tetramethylsilane (TMS, 0 ppm) for ^{29}Si as external standards. The ^1H
225 experiment was examined under 90° pulse with a spinning rate of 14
226 kHz, with several scans of 512 and 1s of recycling delay. The ^{29}Si
227 experiment was performed under 90° pulse with a ^1H decoupling
228 spin and was recorded at the same spinning rate. The number of
229 scans was 10240, and the recycling time was 10s. The decomposition
230 of the spectra was performed using Dmfit-2020 software. The spectra

231 were deconvoluted using mixed Gaussian/Lorentzian line fitting
232 (Massiot et al., 2002; Quillard et al., 2011).

233 **3. Results and discussion**

234 **3.1 Particle size classification of kaolinite by membrane**

235 **separation**

236 Filter membrane and vacuum filtration classify the kaolin slurry of < 2
237 μm into three particle size grades. **Fig. 2 (A)** reveals that such a
238 classification is effective as three distinct peaks of particle size
239 distribution can be observed in A1, A2, and A3 curves. A1 displays
240 the smallest particle size, with an equivalent average sphere diameter
241 of around 300 nm. This value is higher than that of the pores of the
242 filter membrane used in the separation (200 nm), which can be
243 assigned to the anisotropic nature of kaolin and/or to the partial
244 aggregation of primary particles. The same trend is observed in the
245 A2 sample, where the equivalent diameter of the largest particles
246 (1000 nm) is higher than the membrane pore size (450 nm). **Fig. 2 (B)**
247 displays the cumulative percentage frequency distributions for the
248 three fractions. A2 particles are distributed in the 200~1000 nm range,
249 whereas the size distribution for A3 particles ranges between 500 and

250 2000 nm.

251 **3.2 Morphology**

252 The reduction in particle size associated with size classification is
253 clearly illustrated in the scanning electron micrographs (**Fig. 3**). A2
254 and A3 particles display a roughly hexagonal shape and significant
255 size heterogeneity, whereas the smallest size fraction displays more
256 homogeneous and rounded particles (**Fig. 3 (a)**). The sub-rounded
257 platy shape suggests the existence of numerous crystalline defects that
258 are particularly common in transported kaolinites. Compared with A2
259 (**Fig. 3 (b)**) and A3 (**Fig. 3 (c)**), the lateral distance between platelets
260 in A1 is lower, and more gaps under 50 nm are visible. The kaolinite
261 lamellae (**Fig. 3 (a)**) of A1 display a thickness of 50-300 nm and the
262 stack of layers are linked to each other forming compact face-face
263 associations after settling. This is in good agreement with Zbik and
264 Gupta's observation (Gupta et al., 2011; Zbik et al., 2008).

265 **Fig. S3 (Supplementary material)** displays more selected TEM
266 images of particles belonging to the three classes of kaolinite particles.
267 In agreement with the SEM observations, the platelet size decreases
268 from A3 to A1. The images also clearly reveal a turbostratic

269 arrangement of stacked crystallites. The selected area electron
270 diffraction (SAED) pattern shown in **Fig. S3 (a)** confirms the disorder
271 structure of kaolinite, and the rings correspond to the (020), (-130), (-
272 113), (014), (3-11), and (-333) lattice plane.

273 **3.3 XRD**

274 XRD patterns of the raw kaolin sample are in good agreement with the
275 standard PDF card of JCPDS 78-2110 (**Fig. 1**) and reveal the presence
276 of some impurities that should be at least partially eliminated by
277 sedimentation. Still, as revealed by the XRD patterns of the different
278 classes (**Fig. 4**), significant amounts of mineral impurities are still
279 present in the coarser fraction A3. Fraction A2 still contains small
280 amounts of quartz, which likely corresponds to very fine particles. In
281 contrast, the smallest fraction with an average particle size of 300 nm
282 is completely devoid of any impurities, and the corresponding XRD
283 pattern only exhibits kaolinite diffraction peaks (**Fig. 4**). Between 19°
284 and 22° (2θ), the reflections (02*l*), (11*l*) are weak and broaden with
285 decreasing particle size. These reflections are very sensitive to the
286 abundance of translation defects, both random and specific
287 displacements of type t_1 ($\sim -a/3$) and t_2 ($\sim -a/3 + b/3$) (Kogure et al.,

288 2010), and their weakening and broadening decrease the HI. The A1
289 sample, therefore, has low crystallinity and is the most disordered
290 among the three fractions. This may be due to either recrystallization
291 in low concentration solution or to weathering of large kaolinite
292 particles. Whereas all patterns display a decrease in the intensities of
293 all XRD reflections, the (060) peak at 61.5° - 63° (2θ) is less affected
294 when kaolinite becomes more disordered. As shown by Brindley and
295 Kurtossy (1961), the $I_{(00l)}/I_{(060)}$ ratio ($l=1$ or 2) allows the evaluation of
296 the primary orientation of the disorder. In the case of sample A1, the
297 narrowing of the (060) reflection and broadening of the (001) indicate
298 a increasing dimensions along the b axis and decreasing along the c
299 axis. This means that the finest particles have a higher aspect ratio.

300 **3.4 SAXS and WAXS**

301 **Fig. 5** displays the SAXS and WAXS results obtained by presenting
302 the evolution of the scattered intensity as a function of the scattering
303 vector q for the three size-sorted kaolinite samples. On a logarithmic
304 scale, the SAXS intensity dependence follows a power-law decay. At
305 high q , the three systems display a power law behavior where I scale
306 as $q^{-3.8}$ for the two larger sizes and as $q^{-3.7}$ for the smaller size. Such

307 behavior is close to Porod's law which states that the evolution for
308 sharp interfaces should scale as q^{-4} . The deviation from Porod's law
309 can be assigned to the surface roughness of the samples. In that
310 context, the smallest particles appear to be slightly rougher than the
311 two other classes. For A1 and A2 particles, a crossover towards a
312 lower slope around q^{-2} is observed around 0.03 \AA^{-1} for A2 and 0.07 \AA^{-1}
313 for A1. This indicates a scattering by slightly aggregated randomly
314 oriented platelets with a high aspect ratio (Duijneveldt et al., 2005;
315 Hotton et al., 2021; Pasco et al., 2021), which agrees with the TEM
316 observations. For the smallest size (A1), the beginning of a Guinier
317 plateau is observed for $q \approx 1.5 \cdot 10^{-3} \text{ \AA}^{-1}$. This suggests that the largest
318 objects in this class of kaolinites have a size of around 4000 \AA ($2\pi/1.5$
319 10^{-3} \AA^{-1}), which concurs with the laser diffraction results (**Fig. 2**).

320 In the high q -range from 0.5 to 2 \AA^{-1} , the WAXS profiles provide
321 information about the crystalline structure. A clear peak at $q \approx 0.88 \text{ \AA}^{-1}$
322 corresponds to a characteristic distance of 7.14 nm is observed. It can
323 be assigned to the $d(001)$ of kaolinite.

324 **3.5 Adsorption isotherms**

325 **Fig. 6** shows the adsorption-desorption isotherms of N_2 at liquid N_2

326 (77 K) temperature for the three particle sizes. The shapes of the
327 isotherm provide a qualitative assessment of the porous structure of
328 the samples. Sample A3 displays an almost reversible Type II nitrogen
329 isotherm. The isotherm shape indicates that this sample does mainly
330 contain macropores, with non-existing micropores and few mesopores.
331 Indeed, the hysteresis is extremely narrow, indicating the presence of a
332 few large mesopores. The situation changes with decreasing particle
333 size. Indeed, both A1 and A2 powders display significant hysteresis
334 patterns without any plateau at high P/P_0 and can then be assigned to
335 Type IV or Type IIB isotherms (Rouquerol et al., 2013). In such cases,
336 the sample contains both mesopores accountable for the hysteresis and
337 macropores, which explain the absence of any plateau at high relative
338 pressure. The volume uptaken at P/P_0 near 1.0 indicates the total
339 porosity of A1 and A2 for pore sizes $< \sim 200$ nm. As shown in **Table 2**,
340 sample A1 has the most pores. Besides, it has more micropores than
341 A2 powder, as shown by a larger adsorbed volume at an extremely low
342 relative pressure ($P/P_0 < 0.01$).

343 The H3 hysteresis loop (according to IUPAC classification) of A1 and
344 A2 are examples of Type II(b) isotherm, indicating the presence of

345 aggregates of plate-like particles.

346 **3.6 Specific surface area and pore size distribution**

347 The measured specific surface area values for A1 and A2 (40.65 and
348 27.13 m²/g, respectively) exceed the range of 10~20 m²/g of refined
349 natural kaolinite reported in the literature (Cases et al., 1986; Gregg
350 and Packer, 1954). The increase in the specific area of A1 and A2 is
351 mainly due to the decreased particle size of these fractions. For all
352 samples, the average particle size calculated from N₂-adsorption
353 (**Table 2**) is close to the size distribution deduced from laser
354 diffraction (**Fig. 2A and B**).

355 **Fig. 7** displays the BJH-HJ P_bSD (partial volume V of each pore
356 diameter D) obtained from the adsorption isotherm of the three
357 powders. The parameter dV/d (logD) is used to access the partial
358 porosity for each pore diameter range (Kuila and Prasad, 2013;
359 Mintova et al., 2015). For kaolinite powder A1, the distribution is
360 bimodal, with a major peak between 50~100 nm and a minor but
361 noticeable peak around 2~3 nm. Compared with A1, A2 displays a
362 similar pore size distribution, with the main peak located around
363 70~100 nm and the peak at 2~3 nm with lower intensity than in the

364 case of A1.

365 Combing the results of B.J.H. pore size distribution (**Fig. 7**) and pore
366 structure parameters (**Table 2**), it is clear that sample A1 had higher
367 BET, larger pore volume, and smaller average pore radius than
368 samples A2 and A3.

369 This variety of pore structures can be attributed to the presence of
370 more internal defects in ultrafine nano-kaolinite (**Fig. S3**).

371 **3.7 Solid-state NMR**

372 Solid-state NMR was used for probing the local environment and
373 short-range order in size-sorted kaolinites. ¹H MAS NMR spectra of
374 the three particle sizes of kaolinite were obtained at 500.5 MHz (**Fig.**
375 **8**). They can be decomposed into 6 or 7 Gaussian/Lorentzian peaks.
376 Peak 1 located at ~6.40 ppm, indicates the presence of residual water
377 in the three kaolinite samples (Zhu et al., 2016). A second peak at
378 4.3~4.5 ppm can be assigned to physically adsorbed water on the
379 external kaolinite surface. Both kinds of water will be easily
380 dehydrated overnight under a vacuum at 100 °C. In the range of
381 around 2 ppm, two peaks are observed at 2.79~2.65 ppm and at
382 1.81~1.80 ppm for the three samples of A1, A2 and A3. Peak 3 at

383 2.79~2.65 ppm is assigned to inner surface hydroxyls that form a
384 hydrogen bond with oxygens from the SiO₄ tetrahedra, whereas peak 4
385 at 1.80 ppm corresponds to inner hydroxyls, with no linkage with any
386 chemical groups in agreement with previously reported data (Fafard et
387 al., 2017; Zhu et al., 2016). The integral areas (S) of peak 3 of A1, A2,
388 and A3 are 38.48%, 21.31%, and 17.02%, respectively. $S_{3A1} > S_{3A2} >$
389 S_{3A3} indicate that the degree of broken Al-O-Al and Si-O-Si bonds of
390 particles evolve as A1 > A2 > A3. This seems to concur with the
391 appearance of samples as observed by SEM and TEM. It must be
392 pointed out that for an ideal Al₂Si₂O₅(OH)₄ kaolinite structure (Fafard
393 et al., 2017), the ratio between the integrated signals of inner surface
394 hydroxyl protons and inner hydroxyl protons should be 3:1 (**Fig. 8 (h)**).
395 The ratios obtained for A1, A2, and A3 are significantly lower (0.97:1,
396 0.75:1, and 0.63:1, respectively), which may be attributed to the highly
397 disordered structure and fine particle size of the studied kaolinite. In
398 addition, these ratios decrease with increasing particle size, which
399 suggests that the smaller the particles, the higher the number of
400 hydroxyl groups.

401 In the range 0~2 ppm, the ¹H resonance peak displays two chemical

402 shifts (i) peak 5 around 1.3 ppm (1.38, 1.35 and 1.32 ppm for A1 to
403 A3, respectively) assigned to $(Al_{VI})_2-OH$; (ii) peak 6 between 0 and 1
404 ppm (δ 0.85, 0.56 and 0.60 ppm for A1 to A3, respectively) assigned
405 to the singly coordinated hydroxyl of $Al_{VI}-OH$ (Huittinen et al., 2013).
406 The integral areas of this latter signal for A1, A2, and A3 are 15.5%,
407 8.68%, and 9.97%, respectively, which proves again that A1 has the
408 most edge hydroxyl groups. For the two large of A2 and A3, an
409 additional strong, sharp signal exists at 1.01 ppm, which can be
410 attributed to small isolated SiO_2 particles bearing hydrogen-bonded
411 hydroxyl groups ($Si-OH-Si$) (Viani et al., 2017).
412 The ^{29}Si NMR spectra between -90~ -92.5 ppm can be decomposed
413 into three components. A weak one located at about -89 ppm can be
414 assigned to $Q^3(1Al)$ (Cadars et al., 2012; Magi et al., 1984; Thompson,
415 1984) and corresponds to the presence of one tetrahedral Al connected
416 to three tetrahedral silicon atoms via bridging oxygen atoms, and to
417 the octahedral sheet by an apical oxygen atom (see details in **Fig. 10**).
418 The relative intensity of this signal increases with diminishing particle
419 size. Considering that the radius of $Al^{3+(IV)}$ (0.39 Å) is larger than
420 that of Si^{4+} (Radii=0.26 Å), the presence of more Al for Si substitution

421 in the finest particles could explain at least partially the high degree of
422 disorder displayed by A1 particles. At a slightly lower chemical shift,
423 all samples exhibit a wide signal that can be decomposed into two
424 components with almost equal intensities (see **Fig. 8 (b), (d), and (f)**)
425 that are both ascribed to Q³(0Al). The fact that the signal is split in two
426 reveals that Si sites in SiO₄ tetrahedrons, i.e., are not strictly equivalent,
427 which agrees with previous studies (Magi et al., 1984; Thompson et al.,
428 1992). In addition, samples A2 and A3 display additional resonances
429 at δ -107.23 ppm and -107.17 ppm that correspond to Q₄ signals
430 assigned to the presence of silica impurities.

431 The ²⁷Al NMR spectra (**Fig. 9**) display a strong peak at δ =3 ppm
432 assigned to six-coordinated Al (Al(VI)) and a much weaker signal at
433 δ =70 ppm assigned to Al(IV) (Newman et al., 1994; Thompson, 1984).
434 This latter signal is rather weak in agreement with the structure of
435 kaolinite, but its intensity increases with decreasing particle size,
436 which confirms the results obtained from the analysis of ²⁹Si spectra,
437 i.e., that Al for Si substitution is more important in smaller particles.

438 **4. Conclusion**

439 Three different particle sizes of kaolinite were separated from high-

440 defect raw sedimentary kaolin through sedimentation and vacuum
441 membrane filtration, that is, A1 (<300 nm), A2 (200~1000 nm), and
442 A3 (500~2000 nm). The comprehensive studies reveal that the degree
443 of the structural disorder of kaolinite increases as the particle size
444 decreases. The smallest particle size of kaolinite possessed high
445 degrees of broken Al-O-Al and Si-O-Si bonds, a more disordered
446 structure, and an elevated specific surface area. The high disorder
447 degree of tiny kaolinite crystals can be attributed to the more
448 aluminum substituted for silicon in the tetrahedron sheet of their
449 structure, compared to the coarser kaolinite crystals.

450

451

452 **Acknowledgment**

453 This work thanks the financial support of the National Key R&D
454 Program of China (No. 2019YFC1904903) and the Science and
455 Technology Major Projects of Shanxi Province of China (No.
456 20181101003). Yongjie Yang thanks the support provided by the
457 China Scholarship Council (CSC No. 202006430032) during his visit
458 to Sorbonne Université.
459

460 **References**

- 461 Aparicio, P., 1999. Mineralogical Interference on Kaolinite Crystallinity Index
462 Measurements. *Clays and Clay Minerals* 47, 12-27.
- 463 Awad, M.E., López-Galindo, A., Sánchez-Espejo, R., Sainz-Díaz, C.I., El-Rahmany,
464 M.M., Viseras, C., 2018. Crystallite size as a function of kaolinite structural order-
465 disorder and kaolin chemical variability: Sedimentological implication. *Applied Clay*
466 *Science* 162, 261-267.
- 467 Bish, D.L., 1993. Rietveld Refinement of the Kaolinite Structure at 1.5 K. *Clays and*
468 *Clay Minerals* 41, 738-744.
- 469 Brindley, G., Kao, C.-C., Harrison, J., Lipsicas, M., Raythatha, R., 1986. Relation
470 between structural disorder and other characteristics of kaolinites and dickites. *Clays*
471 *and Clay Minerals* 34, 239-249.
- 472 Brindley, G., Kurtossy, S.S., 1961. Quantitative determination of kaolinite by X-ray
473 diffraction. *American Mineralogist: Journal of Earth Planetary Materials* 46, 1205-
474 1215.
- 475 Cadars, S., Guégan, R., Garaga, M.N., Bourrat, X., Le Forestier, L., Fayon, F.,
476 Huynh, T.V., Allier, T., Nour, Z., Massiot, D., 2012. New Insights into the Molecular
477 Structures, Compositions, and Cation Distributions in Synthetic and Natural
478 Montmorillonite Clays. *Chemistry of Materials* 24, 4376-4389.
- 479 Cases, J., Cunin, P., Grillett, Y., Poinsignon, C., Yvon, J., 1986. Methods of analysing
480 morphology of kaolinites: relations between crystallographic and morphological
481 properties. *Clay minerals* 21, 55-68.
- 482 Cheng, P., Schachman, H., 1955. Studies on the validity of the Einstein viscosity law
483 and Stokes' law of sedimentation. *Journal of polymer science* 16, 19-30.
- 484 Du, J., Morris, G., Pushkarova, R.A., Smart, R.S., 2010. Effect of surface structure
485 of kaolinite on aggregation, settling rate, and bed density. *Langmuir* 26, 13227-
486 13235.
- 487 Duijneveldt, J.S.v., Klein, S., Leach, E., Pizzey, C., Richardson, R.M., 2005. Large
488 scale structures in liquid crystal/clay colloids. *Journal of Physics: Condensed Matter*
489 17, 2255-2267.
- 490 Eshel, G., Levy, G., Mingelgrin, U., Singer, M., 2004. Critical evaluation of the use
491 of laser diffraction for particle-size distribution analysis. *Soil Science Society of*
492 *America Journal* 68, 736-743.
- 493 Fafard, J., Terskikh, V., Detellier, C., 2017. Solid-state ¹H and ²⁷Al NMR studies of
494 DMSO-kaolinite intercalates. *Clays and Clay Minerals* 65, 206-219.
- 495 Franco, F., Perez-Maqueda, L.A., Perez-Rodriguez, J.L., 2004a. The effect of
496 ultrasound on the particle size and structural disorder of a well-ordered kaolinite.

497 Journal of colloid and interface science 274, 107-117.

498 Franco, F., Pérez-Maqueda, L.A., Pérez-Rodríguez, J.L., 2004b. Influence of the
499 particle-size reduction by ultrasound treatment on the dehydroxylation process of
500 kaolinites. *J Therm Anal Calorim* 78, 1043-1055.

501 Gregg, S., Packer, R., 1954. The production of active solids by thermal
502 decomposition. Part IV. Vermiculite. *Journal of the Chemical Society*, 3887-3893.

503 Gupta, V., Hampton, M.A., Stokes, J.R., Nguyen, A.V., Miller, J.D., 2011. Particle
504 interactions in kaolinite suspensions and corresponding aggregate structures. *Journal*
505 *of colloid and interface science* 359, 95-103.

506 Hamzaoui, R., Muslim, F., Guessasma, S., Bennabi, A., Guillin, J., 2015. Structural
507 and thermal behavior of proclay kaolinite using high energy ball milling process.
508 *Powder Technology* 271, 228-237.

509 Hotton, C., Sirieix-Plenet, J., Ducouret, G., Bizien, T., Chenneviere, A., Porcar, L.,
510 Michot, L., Malikova, N., 2021. Organisation of clay nanoplatelets in a
511 polyelectrolyte-based hydrogel. *Journal of colloid and interface science* 604, 358-
512 367.

513 Huittinen, N., Sarv, P., Lehto, J., 2013. A ^1H and ^{27}Al NMR investigation of
514 yttrium(III) and europium(III) interaction with kaolinite. *Applied Clay Science* 80-81,
515 182-188.

516 Jaber, M., Lambert, J.-F., Balme, S., 2018. Protein adsorption on clay minerals,
517 *Surface and Interface Chemistry of Clay Minerals*, pp. 255-288.

518 Kobayashi, Y., Sumi, K., Kato, E., 2000. Preparation of dense cordierite ceramics
519 from magnesium compounds and kaolinite without additives. *Ceramics International*
520 26, 739-743.

521 Kogure, T., Elzea-Kogel, J., Johnston, C.T., Bish, D.L., 2010. Stacking Disorder in a
522 Sedimentary Kaolinite. *Clays and Clay Minerals* 58, 62-71.

523 Konert, M., Vandenberghe, J., 1997. Comparison of laser grain size analysis with
524 pipette and sieve analysis: a solution for the underestimation of the clay fraction.
525 *Sedimentology* 44, 523-535.

526 Kuila, U., Prasad, M., 2013. Specific surface area and pore-size distribution in clays
527 and shales. *Geophysical Prospecting* 61, 341-362.

528 Liu, L., Liu, Q., Algeo, T., Zhang, H., Yang, Y., Peng, G., Zhang, S., Hong, H., Liu,
529 D., 2021. Mineralogy, Geochemistry, and Genesis of Kaolinitic Claystone Deposits
530 in the Datong Coalfield, Northern China. *Clays and Clay Minerals* 69, 68-93.

531 Magi, M., Lippmaa, E., Samoson, A., Engelhardt, G., Grimmer, A., 1984. Solid-state
532 high-resolution silicon-29 chemical shifts in silicates. *The Journal of Physical*
533 *Chemistry* 88, 1518-1522.

534 Makó, É., Frost, R.L., Kristóf, J., Horváth, E., 2001. The Effect of Quartz Content on

535 the Mechanochemical Activation of Kaolinite. *Journal of colloid and interface*
536 *science* 244, 359-364.

537 Massiot, D., Fayon, F., Capron, M., King, I., Le Calvé, S., Alonso, B., Durand, J.-O.,
538 Bujoli, B., Gan, Z., Hoatson, G., 2002. Modelling one- and two-dimensional solid-
539 state NMR spectra. *Magnetic Resonance in Chemistry* 40, 70-76.

540 Mintova, S., Jaber, M., Valtchev, V., 2015. Nanosized microporous crystals:
541 emerging applications. *Chemical Society Reviews* 44, 7207-7233.

542 Morsy, F.A., El-Sherbiny, S., Hassan, M.S., Mohammed, H.F., 2014. Modification
543 and evaluation of Egyptian kaolinite as pigment for paper coating. *Powder*
544 *Technology* 264, 430-438.

545 Mpofo, P., Addai-Mensah, J., Ralston, J., 2003. Investigation of the effect of polymer
546 structure type on flocculation, rheology and dewatering behaviour of kaolinite
547 dispersions. *International Journal of Mineral Processing* 71, 247-268.

548 Newman, R., Childs, C., Churchman, G., 1994. Aluminium coordination and
549 structural disorder in halloysite and kaolinite by ^{27}Al NMR spectroscopy. *Clay*
550 *Minerals* 29, 305-312.

551 Pasco, H., de Viguerie, L., Faustini, M., Coelho-Diogo, C., Ersen, O., Gervais, C.,
552 Gobeaux, F., Ihiawakrim, D., Jaber, M., Walter, P., Sanchez, C., 2021. Shedding
553 Light on Functional Hybrid Nanocomposites 19th Century Paint Medium. *Advanced*
554 *Functional Materials*.

555 Plançon, A., Giese, R., Snyder, R., Drits, V., Bookin, A., 1989. Stacking faults in the
556 kaolin-group minerals: Defect structures of kaolinite. *Clays and Clay Minerals* 37,
557 203-210.

558 Quillard, S., Paris, M., Deniard, P., Gildenhaar, R., Berger, G., Obadia, L., Boulter,
559 J.M., 2011. Structural and spectroscopic characterization of a series of potassium-
560 and/or sodium-substituted beta-tricalcium phosphate. *Acta Biomater* 7, 1844-1852.

561 Rand, B., Melton, I.E., 1977. Particle interactions in aqueous kaolinite suspensions: I.
562 Effect of pH and electrolyte upon the mode of particle interaction in homoionic
563 sodium kaolinite suspensions. *Journal of colloid and interface science* 60, 308-320.

564 Richard, D., Rendtorff, N.M., 2019. First principles study of structural properties and
565 electric field gradients in kaolinite. *Applied Clay Science* 169, 67-73.

566 Rouquerol, J., Rouquerol, F., Llewellyn, P., Maurin, G., Sing, K.S., 2013. Adsorption
567 by powders and porous solids: principles, methodology and applications. Academic
568 press.

569 Silva, F.A.N.G., Luz, A.B., Sampaio, J.A., Bertolino, L.C., Scorzelli, R.B., Duttine,
570 M., da Silva, F.T., 2009. Technological characterization of kaolin: Study of the case
571 of the Borborema–Seridó region (Brazil). *Applied Clay Science* 44, 189-193.

572 Suraj, G., Iyer, C., Rugmini, S., Lalithambika, M., 1997. The effect of micronization

573 on kaolinites and their sorption behaviour. *Applied Clay Science* 12, 111-130.

574 Thompson, J., 1984. 29 Si and 27 Al nuclear magnetic resonance spectroscopy of 2:
575 1 clay minerals. *Clay minerals* 19, 229-236.

576 Thompson, J.G., Uwins, P.J., Whittaker, A.K., Mackinnon, I.D., 1992. Structural
577 characterisation of kaolinite: NaCl intercalate and its derivatives. *Clays and Clay*
578 *Minerals* 40, 369-380.

579 Varajão, A.F.D.C., Gilkes, R.J., Hart, R.D., 2001. The relationships between kaolinite
580 crystal properties and the origin of materials for a Brazilian kaolin deposit. *Clays and*
581 *Clay Minerals* 49, 44-59.

582 Viani, A., Mali, G., Mácová, P., 2017. Investigation of amorphous and crystalline
583 phosphates in magnesium phosphate ceramics with solid-state 1 H and 31 P NMR
584 spectroscopy. *Ceramics International* 43, 6571-6579.

585 Wen, B., Aydin, A., Duzgoren-Aydin, N.S., 2002. A comparative study of particle
586 size analyses by sieve-hydrometer and laser diffraction methods. *Geotechnical*
587 *Testing Journal* 25, 434-442.

588 Yang, X., Zhang, Q., Li, X., Jia, X., Wei, X., Shao, M.a., 2015. Determination of Soil
589 Texture by Laser Diffraction Method. *Soil Science Society of America Journal* 79,
590 1556-1566.

591 Yang, Y., Zhang, H., Zhang, K., Liu, L., Ji, L., Liu, Q., 2020. Vulcanization,
592 interfacial interaction, and dynamic mechanical properties of in-situ organic amino
593 modified kaolinite/SBR nanocomposites based on latex compounding method.
594 *Applied Clay Science* 185.

595 Zbik, M.S., Smart, R.S., Morris, G.E., 2008. Kaolinite flocculation structure. *Journal*
596 *of colloid and interface science* 328, 73-80.

597 Zhang, K., Zhang, H., Liu, L., Yang, Y., Liu, L., Liu, Q., 2021a. Dispersibility of
598 Kaolinite-Rich Coal Gangue in Rubber Matrix and the Mechanical Properties and
599 Thermal Stability of the Composites. *Minerals* 11.

600 Zhang, S., Liu, Q., Yang, Y., Wang, D., He, J., Sun, L., 2017. Preparation,
601 morphology, and structure of kaolinites with various aspect ratios. *Applied Clay*
602 *Science* 147, 117-122.

603 Zhang, S., Liu, Q., Yang, Y., Zhang, H., Liu, J., Zeng, S., LaChance, A.M., Barrett,
604 A.T., Sun, L., 2021b. An efficient method to prepare aluminosilicate nanoscrolls
605 under mild conditions. *Chem Commun (Camb)* 57, 789-792.

606 Zhang, Y., Liu, Q., Zhang, Q., Lu, Y., 2010. Gas barrier properties of natural
607 rubber/kaolin composites prepared by melt blending. *Applied Clay Science* 50, 255-
608 259.

609 Zhu, X., Zhu, Z., Lei, X., Yan, C., 2016. Defects in structure as the sources of the
610 surface charges of kaolinite. *Applied Clay Science* 124-125, 127-136.

611

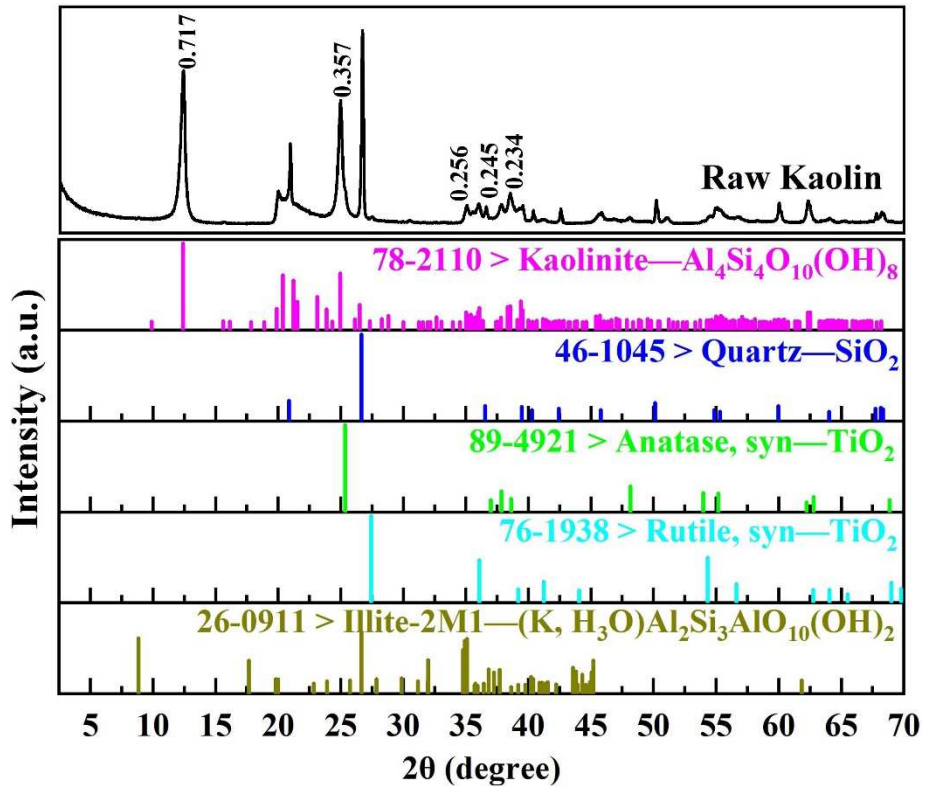


Fig. 1. XRD patterns of raw kaolin sample.

612

613

614

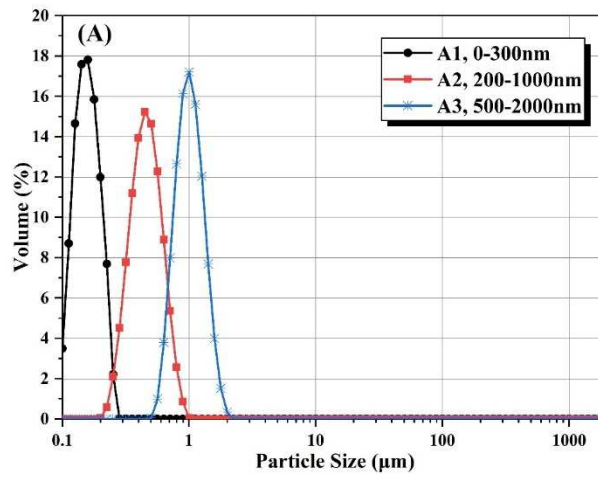
615

616

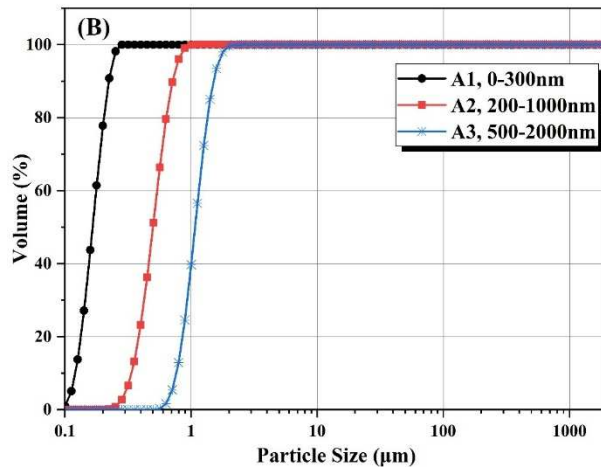
617

618

619



620



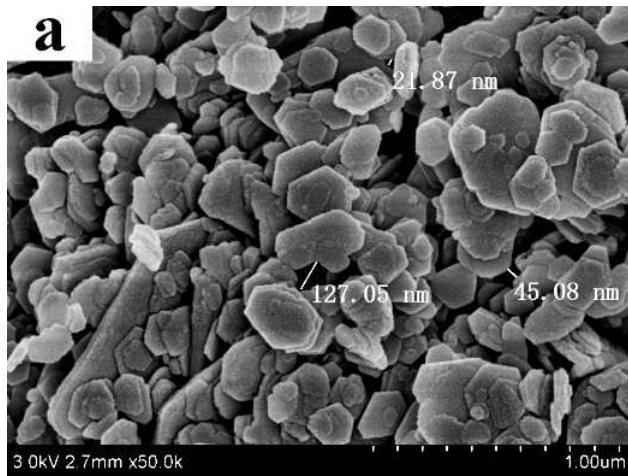
621

622 **Fig. 2.** Particle size distribution (A) and cumulative particle size distribution (B) of

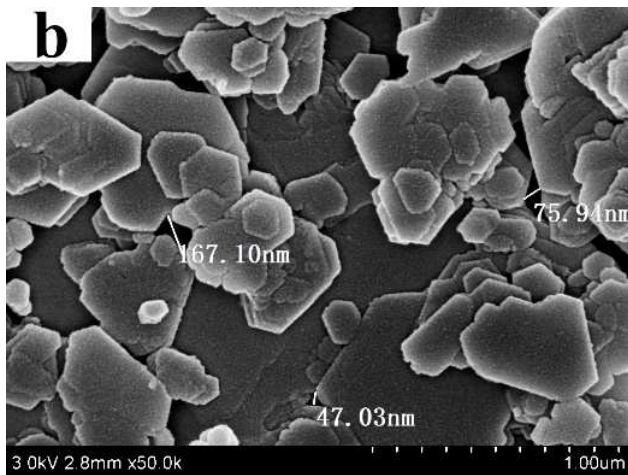
623 kaolinite after vacuum filtration of the $< 2 \mu\text{m}$ kaolin suspension.

624

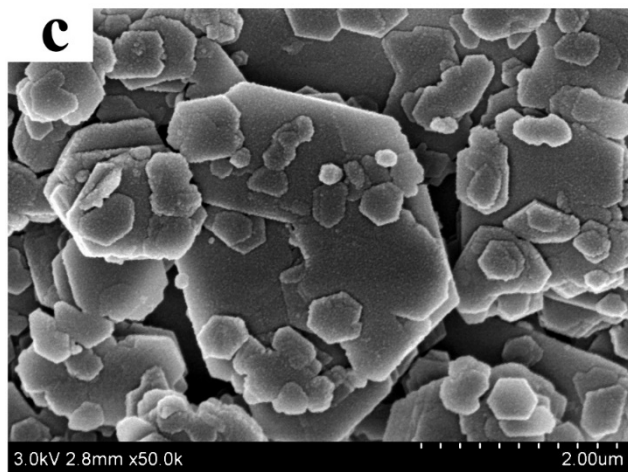
625



626



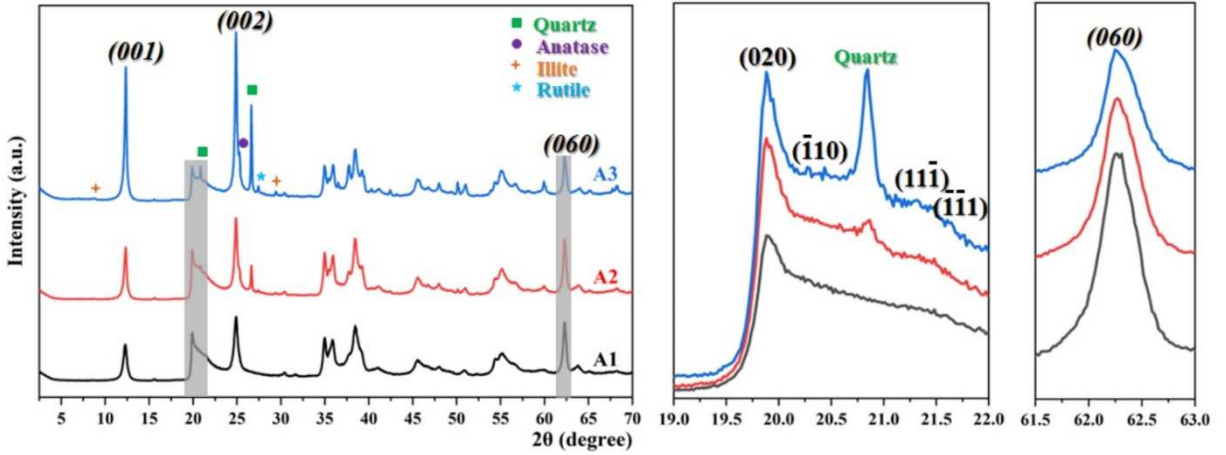
627



628

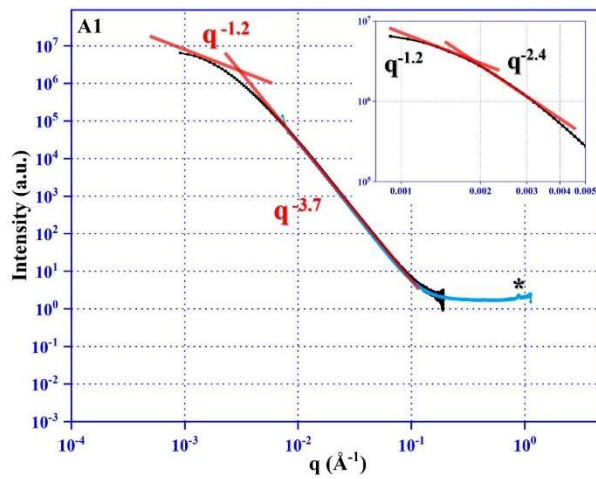
629 **Fig. 3.** SEM observations of three different particle sizes of kaolinite; (a) A1; (b) A2;
630 (c) A3.

631
632
633

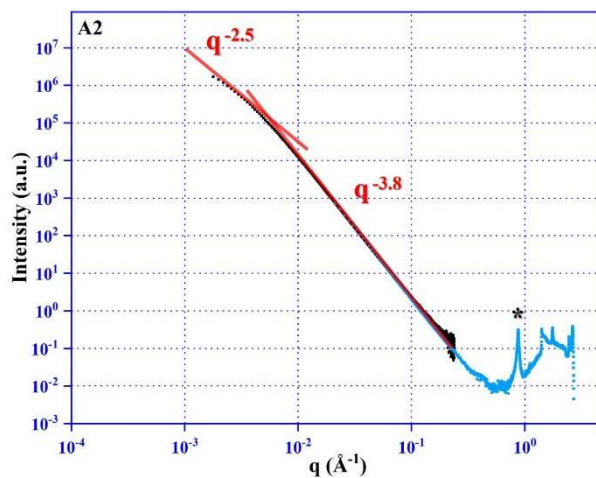


635 **Fig. 4.** XRD patterns of kaolin samples classified into different particle sizes; A1
636 ≤ 300 nm, A2 200~1000 nm, A3 500~2000 nm.

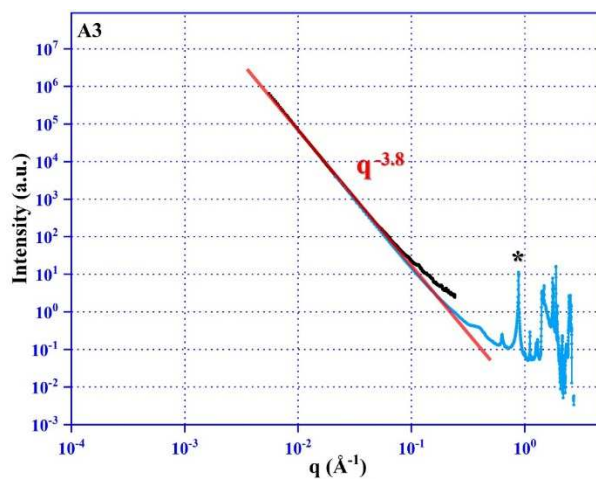
637
638
639
640
641
642
643



644



645

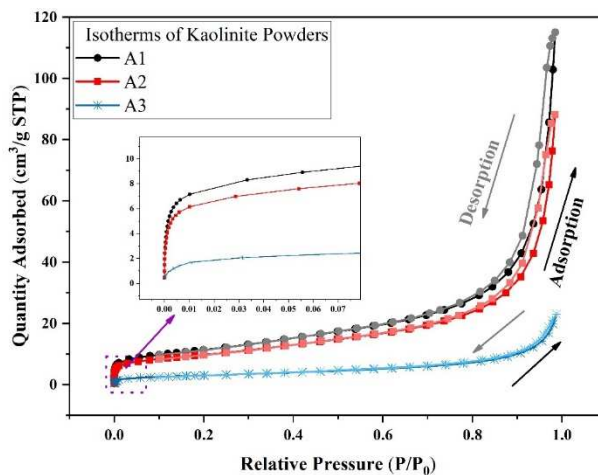


646

647

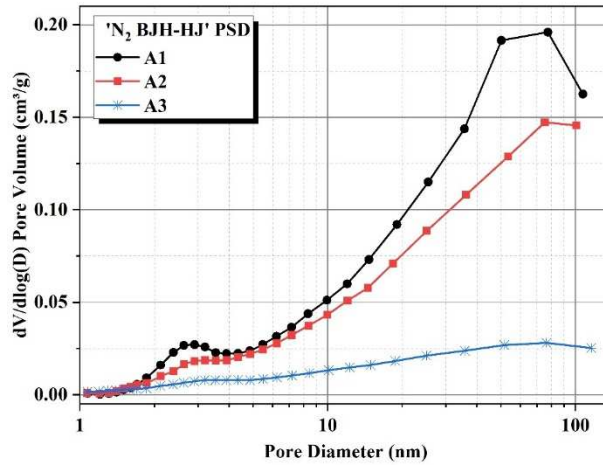
648

Fig. 5. Small- and Wide-angle scattering curves were obtained from SAXS and WAXS.

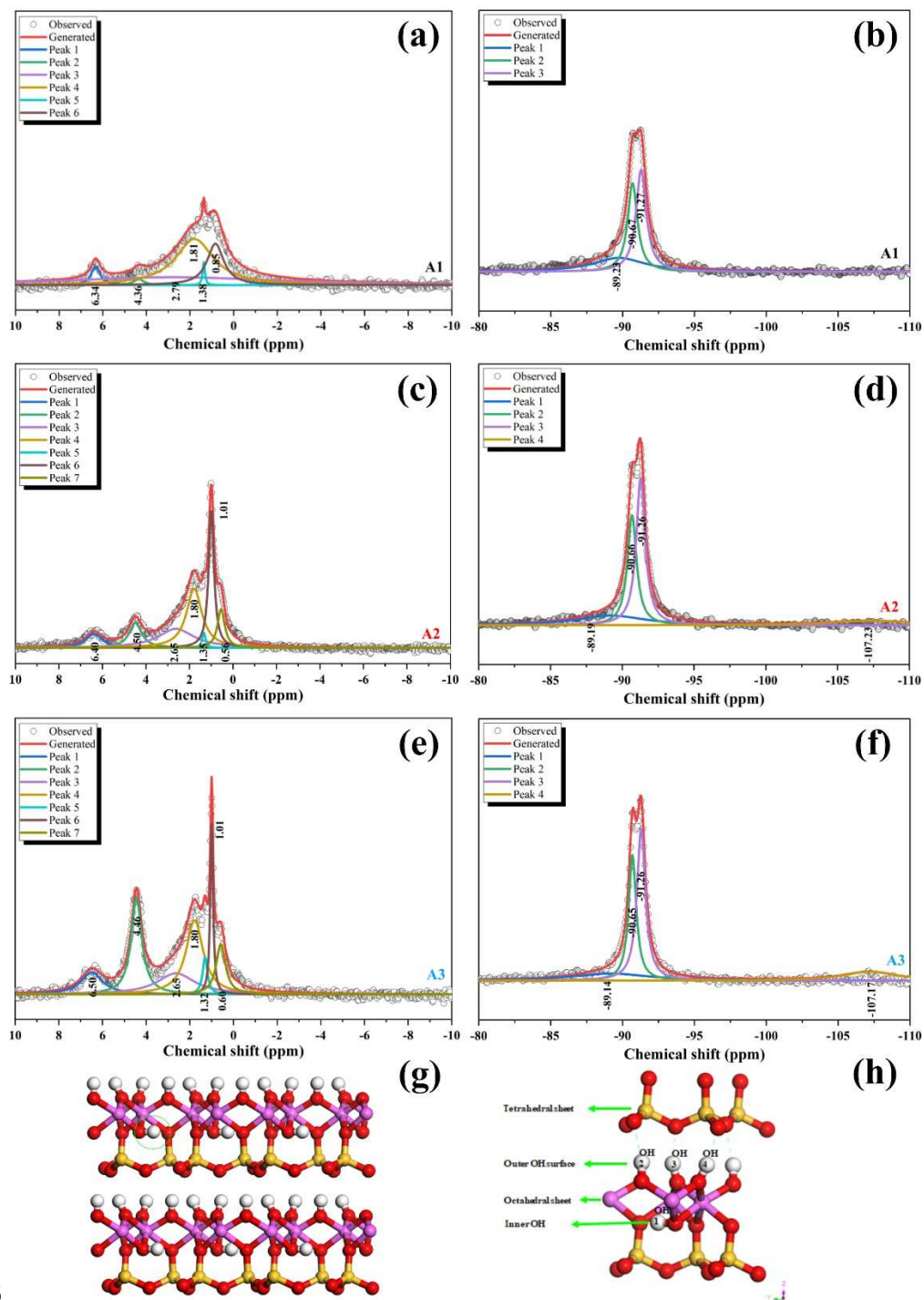


649

650 **Fig. 6.** Adsorption-desorption isotherms of N₂ at liquid N₂ (77 K) temperature of
651 three particle size distributions of A1, A2, and A3.
652
653

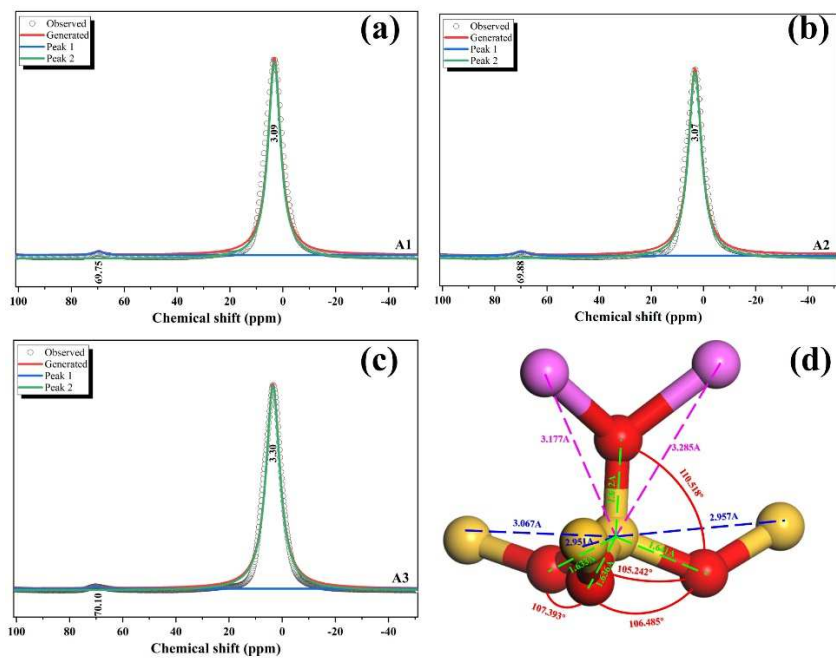


654 **Fig. 7.** BJH-HJ P_bSD of A1, A2, and A3 acquired from the N₂ adsorption isotherm;
655 Sample A1 and A2 have bimodal distribution with a minor peak around 2~3 nm,
656 while A1 has a major peak at 50-100 nm and A2 at 70~100 nm. A3 shows a
657 unimodal distribution with a very weak peak around 100 nm.
658
659



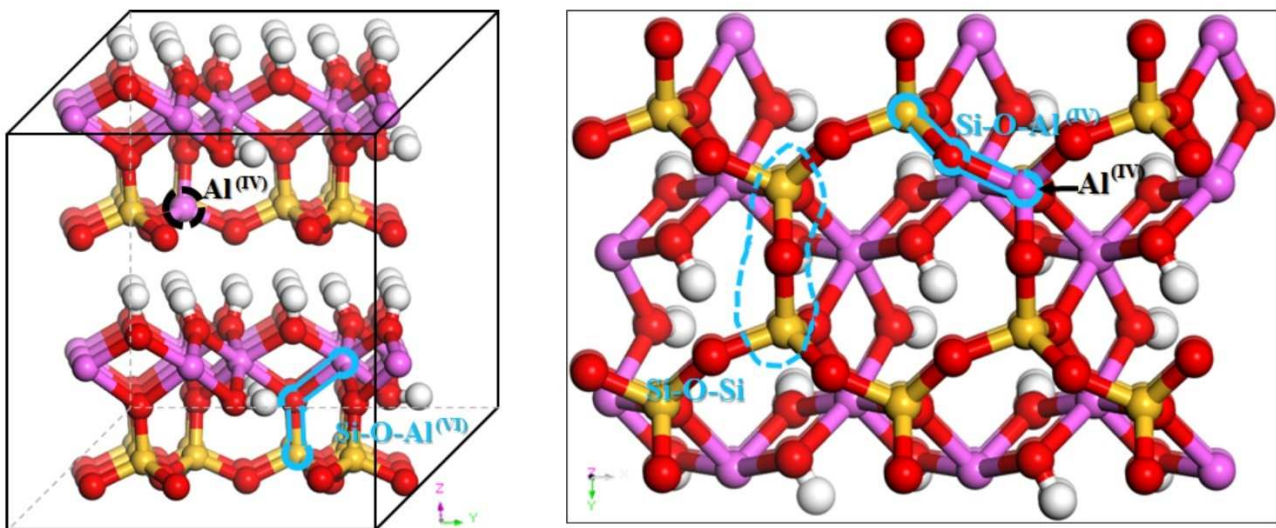
6
661
662

Fig. 8. ^1H (left) and ^{29}Si (right) MAS NMR spectra of A1, A2, and A3.



663
664
665

Fig. 9. ^{27}Al MAS NMR spectra of A1, A2, and A3.



667
668
669
670
671

Fig. 10. Projection of the structure of kaolinite from the (100) direction, which contains one $\text{Q}^3(1\text{Al})$ Si moieties per unit cell.

672 **Table 1** Chemical composition of original kaolin in the study (wt%).

Sample	SiO ₂	Al ₂ O ₃	TiO ₂	Fe ₂ O ₃	P ₂ O ₅	K ₂ O	CaO	MgO	SO ₃	Cr ₂ O ₃	LOI
Kaolin	49.92	31.29	1.79	1.15	0.16	0.12	0.09	0.09	0.06	0.027	15.1

673

674

675

676

677 **Table 2** BET and pore structure parameters of three different particle

678 sizes of kaolinite.

	^c BET Surface Area (m ² /g)	^a Pore Volume (cm ³ /g)	^b Pore size (nm)	^c Nanoparticle Size (Average Particle Size, nm)
A1	40.6572	0.178006	17.51289	147.5753
A2	27.1262	0.136377	15.67295	430.963
A3	11.0345	0.035655	10.92476	1087.4934

679

680 a: Single point adsorption total pore volume of pores less than 127.7428 nm
681 diameter at p/p° of A1= 0.984750664;

682 Single point adsorption total pore volume of pores less than 122.8734 nm
683 diameter at p/p° of A2= 0.984134123;

684 Single point adsorption total pore volume of pores less than 160.7880 nm
685 diameter at p/p° of A3 = 0.987934275;

686 b: Adsorption average pore diameter (4V/A by B.E.T.);

687 c: Average Particle Size.

688