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High pressure dissociation of CaTiO₃ perovskite into CaO and CaTi₂O₅

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Highlights:

- Phase diagram of the calc un. titanate perovskite CaTiO₃ up to 170 GPa and 4500 K.
- Thermal equation of state of the orthorhombic (*Pbnm*) perovskite CaTiO₃.
- Dissociation of the CaTi J₃ perovskite into CaO and CaTi₂O₅ without going through a post-perovskite place

Key words:

Perovskite, CaTiO₃, Equation of State, Dissociation, High-pressure

Abstract

In this study, we investigated phase transformations of CaTiO₃ perovskite using x-ray diffraction at high pressure and high temperature up to 170 GPa and 4500 K in a laser-heated diamond-anvil cell. We report a high-pressure dissociation of CaTiO₃ into CaO-B2 and CaTi₂O₅ with a monoclinic P2/m structure, instead of the expected transformation of the orthorhombic distorted perovskite structure into a post-perovskite phase. We propose that this transition may be favored by the *B1* to *B2* phase change of CaO at around 60 GPa. In order to

provide additional information on the high pressure properties of CaTiO₃ perovskite, we measured its melting temperature using CO₂ laser heated diamond anvil cell up to 55 GPa yielding a fit of the melting curve to a Kraut-Kennedy empirical law according to : $Tm(K) = 2188 * [1 + 4.23 * (\Delta V/V_0)]$. To provide some further insight into the thermodynamic properties of CaTiO₃, we determined the P-V-T equation of state of the orthorhombic mineral perovskite, fitted by using a third order Birch-Murnaghan equation of state and a Berman thermal expansion model. The fit of the data yields to K₀ = 180.6(4) GPa, K'₀ = 4 (fixed), $\partial K/\partial T = -0.022(1)$ GPa K⁻¹, $\alpha_1 = 3.25(5) \times 10^{-5}$ K⁻¹, $\alpha_2 = 1.3(1) \times 10^{-8}$ K⁻²

Plain language summary

Among 5000+ discovered exoplanets, some of then are classified as super-Earths. Based on their average densities derived from mass/rad us observations, these planets are expected to be rocky but have a larger radius than that of the Earth and thus higher internal pressures. Bridgmanite, the orthorhombic perovskite structured compound with (Mg,Fe)SiO₃ composition, considered to be the most abundant mineral in the Earth's rocky mantle is likely abundant in super-Earths as well. Atomistic calculations have predicted that MgSiO₃ dissociates between 1 to 3 terapascal (1 7. Pa = 10^{12} Pa = 10 Mbar) into oxide phases but such pressures remain outside of the realm of high pressure experiments. In this work, we focus on the archetypal perovskite structure $C_a T_i O_i$, and investigate its phase diagram using the laserheated diamond anvil cell technicas coupled with in situ x-ray diffraction. At pressures above 60 GPa and temperatures up to 450° K, we show evidence of the dissociation of CaTiO₃ into CaO and CaTi₂O₅ instead of a transformation into a denser post-perovskite phase. Similar decomposition schemes nave been proposed for MgSiO₃ at ultra-high pressures in super-Earth's interiors; the decomposition reaction discovered here in CaTiO₃ could allow to better understand their consequences by providing a model amenable to experiments under accessible pressures.

1 Introduction

Bridgmanite, the orthorhombic (*Pbnm* space group) perovskite-structured compound with (Mg,Fe)SiO₃ composition is considered to be the most abundant mineral in the Earth's rocky mantle. This mineral undergoes a well-documented transition to a *Cmcm* CaIrO₃-type post-perovskite phase at approximately 120 GPa (Murakami et al. 2004; Oganov & Ono, 2004). For super-Earth modelling, it is important to investigate MgSiO₃ behaviour at much higher pressures and temperatures. Atomistic calculations have predicted that MgSiO₃

dissociates between 1 to 3 TPa into simple oxides MgO and SiO₂ or more complex assembly of denser silicates and oxides (Mg₂SiO₄, MgSi₂O₅) (Umemoto et al., 2011, 2017, 2019 and references therein; Wu et al. 2014; Niu et al. 2015). However, such pressures remain outside of the realm of static high-pressure experiments and difficult to achieve together with relevant temperatures in dynamic compression experiments. Therefore, to experimentally acquire better knowledge about these important mechanisms and transitions, another approach is to study analogue compounds. An analogue material usually mimics the crystallographic structure of the material while having a different composition. Several compounds such as CaRhO₃, MgGeO₃, NaMgF₃, GdFeO₃ and much more have been identified as model structures to explore high-density high-pressure evolution of $..BX_3$ compounds (Xu et al., 2015 and references herein). In this work, we focus on the mineral called perovskite (and archetypal perovskite structure) CaTiO₃, which can be interesting to compare with the perovskite-structured materials of rocky planets such a. MgSiO₃ and CaSiO₃ perovskite. Since CaTiO₃ has the *Pbnm* orthorhombic perovskies structure at ambient conditions, it is expected to undergo structural transitions at more moderate pressures than bridgmanite or CaSiO₃. It has been also predicted that CaTiC - may undergo a series of transition perovskite (Pv) – post-perovskite (pPv) – post-post per Jvskite (ppPv) (Xu et al. 2015) that makes of this compound a very interesting case to study.

Perovskite and post-perovsk.tc s ructures have been extensively studied during these past few years both by *ab initic* calculations and experiments (e.g. Murakami et al., 2004; Hirose et al., 2017 and references therein; Dutta et al., 2019; Umemoto et al., 2011, 2017, 2019; Xu et al., 2015; Tsuchine et al., 2004). For CaTiO₃, first-principles calculations by Xu and al. (2015) proport that the *Pbnm* perovskite structure transforms to a post-perovskite *Cmcm* CaIrO₃ phase at $^22.7$ GPa, followed by another transition to a post-post-perovskite structure at 71.7 GPa. Using diamond anvil cells at room-temperature, it has been demonstrated by Guennou et al. (2010) that CaTiO₃ remains under the *Pnma* (or equivalent *Pbnm*) orthorhombic perovskite structure up to 60 GPa. With CaTiO₃, we focus on a class of materials in which the structural distortion is dominated by tilts of BO₆ octahedra (Zhao et al. 2011). It is interesting to identify the point where this structure destabilizes, for instance when the tilt angle ceases to evolve and bond compression becomes the main structural modification. The absence of heating in Guennou et al. (2010) may have, possibly, retained a metastable structure at the highest pressures. A dynamic compression study (Trunin et al.

1971) suggested a phase change in $CaTiO_3$ at 60 GPa without giving any structural information for the new state.

In the present study, we investigate the phase diagram of calcium titanate $CaTiO_3$ perovskite using the laser-heated diamond anvil cell technique coupled with *in situ* x-ray diffraction. We report the results of experiments conducted up to 170 GPa and temperatures up to 4500 K. We measure the thermal equation of state of the *Pbnm* orthorhombic perovskite phase. We also observe a gradual dissociation of $CaTiO_3$ into oxides instead of transformations into post-perovskite or post-perovskite phases.

2 Methods

2.1 X-ray diffraction experiment

The *in-situ* experiments were carried out at the 127 beamline at the European Synchrotron Radiation Facility (ESRF) using lase -heated diamond anvil cell (LH-DAC) methods. We characterised the samples using powder a ray diffraction in angle-dispersive mode using an MCCD detector (Mezouar et al. 200); Mezouar et al, 2017). In this setup, the laser heating system is composed of two math-power diode-pumped Nd-YAG infrared lasers entering the DAC from both sides with a angle of 20°. Temperatures were measured by multi-wavelength spectral radiometry methods (Schultz et al. 2005). The working wavelength was 0.3738 Å (i.e., 33 keV) and r mained the same during the whole experiment. We conducted eight runs in total (1.1 le 1) with two types of starting materials. The first three runs were done using commercial synthetic CaTiO₃ (purchased from Kurt J. Lesker Company with a purity of 99.9%). Finied and loaded without any pressure transmitting medium but with platinum black product that served as IR laser heating absorber as well as pressure calibrant (series #1, #2 and #3). The following four runs were carried out with a natural CaTiO₃ sample (Collection of Minerals, Sorbonne University, courtesy from J.C. Boulliard) compressed between KCl disks acting as thermal insulating pressure medium and pressure calibrant (series #4, #5, #6 and #7). X-ray diffraction patterns collected at ambient conditions confirmed the orthorhombic structure and the purity of the CaTiO₃ sample, with an excellent match with reference diffraction patterns obtained on a single crystal by Guennou et al. (2010). In the natural CaTiO₃ samples, trace elements in the crystal structure enabled significant YAG laser (1.064 µm) absorption and heating. The last experimental run was conducted with a natural CaTiO₃ sample loaded with neon as a pressure medium and a small piece of gold was added for pressure measurements (series #8).

X-ray diffraction images were collected on a MAR165 2D-CCD detector with exposure time of the order of 30 seconds at ambient temperature conditions, and up to 90 seconds at the most extreme conditions. Each CCD image was then processed with the DIOPTAS software (Prescher & Prakapenka, 2015) to obtain intensity versus 2-theta angle integrated patterns. These patterns were subsequently treated using the FullProf Suite software (Rodriguez-Carvajal, 2001) with a Le Bail refinement. Pressures corresponding to each diffractogram were calculated using different equations of state of platinum (Matsui et al. 2009), gold (Takemura & Dewaele, 2008) and KCl (Dewaele et al. 2012), as described in supplementary materials. The EoSFit program (Angel et al., 2014; Gonzales-Plata et al., 2016) was used to fit the thermal equation of state parameters \odot CaTiO₃ with the perovskite *Pbnm* structure.

2.2 Melting experiment

To provide additional information on the 1-jeh-pressure behaviour of CaTiO₃, we performed in situ laser-heated diamond anvil cell experiments to determine the melting curve of CaTiO₃ up to 60 GPa. These experiments were carried out at IMPMC laboratory, Sorbonne University. Single pieces of synthetic C² fiC₃ (same sample as in the run #1-3 Table 1) were loaded in DACs equipped with anvils of 200 and 300 µm culet sizes to reach pressures ranging between 16 and 55 GPa. The same le chamber drilled in the pre-indented Re-gasket of the DAC assembly was loaded with sample pieces of about 10-20 µm thicknesses which were sandwiched between the pressure medium discs of KCl of 10-20 µm thickness. KCl acted as pressure transmitting medium, 'ut also as a thermal insulator and chemical barrier between the sample and the diar ion,¹ anvils. A ruby sphere was placed at the edge of the sample chamber to measure prossure with ruby fluorescence. Pressures were measured before and after each heating run. Our experimental setup is equipped with one-sided CO_2 laser heating system which is built with a Synrad Evolution 125 CO₂ laser delivering a gaussian beam (wavelength 10.6 μ m, M² < 1.2) focused by a ZnSe lens on the sample through the diamond in an axial geometry (focal length 25 mm). The use of a CO₂ laser permits to heat nonmetallic samples (*i.e.*, wide bandgap insulating materials) such as $CaTiO_3$ without any additional laser absorber such as Pt black. Heating is also effective across the entire thickness of the sample, with lesser axial temperature gradient (Smith et al., 2018). Temperature measurements were performed by spectro-radiometry, *i.e.* the analysis of thermal emission of the sample in the visible region (from 500 to 800 nm). The imaging and pyrometry system was made of a two-mirrors Schwarzschild reflective objective lens (focal 850 mm) used to

make an image of the hot spot on the entrance slit of the spectrometer (Acton 500i) equipped with an Andor-CDD camera. Optical pyrometry in the 500-800 nm range typically yields Planck's law fits within 10 K (uncertainty of the fit), which corresponds to an actual temperature uncertainty of about 130 K when emissivity changes and temperature variation are taken into account. Melting was detected by observation of a discontinuity in the curve of temperature vs laser power. This discontinuity in the absorption of the laser radiation, appears as a sudden increase of temperature at a given laser power, temperature which reaches then a so called 'plateau' attributed to melting. In few cases the laser power was slowly increased until the temperature reached a stable value ("plateau") due to melting at a given pressure (*e.g.* Boehler, 2000; Pradhan et al., 2015; Geballe & Jeanloz, -012). We managed to fit the melting points using both the Kraut-Kennedy equation (K1 ut & Kennedy, 1966) and the Simon equation (Simon, 1937). The formulation of the Kr. ut-k ennedy equation is:

(1)

$$T_m = T_m^0 \left(1 - C \frac{\Delta \nu}{V_0} \right)$$

with T_m the melting temperature at a given preseure and T_m^{0} is the melting temperature at 1 atm. C is a constant depending on the mater. I. The isothermal dilatation $\Delta V/V_0$ is defined as: (2)

$$\frac{\Delta V}{V_0} = \frac{V_0 - V}{V}$$

where V_0 is the volume at amplicate pressure and temperature.

The Simon equation is to mulated as:

(3)

$$P_m = A\left[\left(\frac{T_m}{T_m^0}\right)^B - 1\right]$$

with T_m the melting temperature at a given pressure and T_m^{0} is the melting temperature at 1 atm. A and B are constants that depend on the material.

3 Results and discussion

3.1 CaTiO₃ thermal equation of state

To establish the thermal equation of state of CaTiO₃ perovskite, we used different data sets. Guennou et al. (2010) provide a reliable equation of state at ambient temperature up to 60 GPa obtained by single crystal x-ray diffraction using fluorescence of ruby as pressure gauge and helium as pressure-transmitting medium. However, above 40 GPa, they noticed a broadening of the peaks in the diffraction pattern attributed to a non-hydrostatic compression of the single crystal sample. Accordingly, we only included their data from 0 to 40 GPa in the refinement. In this work, we present a room-temperature dataset that can be compared to Guennou et al. (2010) in the same pressure range. The two datasets agree up to 47 GPa and we discuss below the evolution at higher pressures.

We coupled the Guennou et al. (2010) results with our data obtained up to temperatures in excess of 3000 K at 60 GPa. Most of the dotate comes from the series with platinum (see "Data EoS" in repository 10.17632/nmkzfc($h/5^{-2}$). In the refinement process, we avoided to include data that showed signs of structural transition to other phases (typically above 60 GPa). We also used the most recent data from Yashima et al. (2009) in the refinement. This work established the high-term erature thermal expansion of CaTiO₃ perovskite and temperature-induced transition processes at ambient pressure up to 1720 K by neutron diffraction and Rietveld analysis returnent.

The refinement at $T_0 = 300$ K was using a with the software EoSFit (Angel et al., 2014) using a third order Birch-Murnaghan equation of state:

(4)

$$P(V,T_0) - \frac{3}{2} \sum_{r=0}^{V} \left[\left(\frac{V}{V_0} \right)^{-\frac{7}{3}} - \left(\frac{V}{V_0} \right)^{-\frac{5}{3}} \right] \left[1 + \frac{3}{4} (K'_0 - 4) \left[\left(\frac{V}{V_0} \right)^{-\frac{2}{3}} - 1 \right] \right]$$

V₀ is the lattice volume at 300 K, V is the volume at a pressure point, K₀ is the bulk modulus and K'₀ is the first pressure derivative of the bulk modulus ($\partial K/\partial P$).

In order to determine the pressure considering the temperature we introduced a thermal pressure term, according to:

(5)

$$P^{tot}(V,T) = P(V,T_0) + P^{th}(V,T)$$

To determine the thermal pressure, we used a Berman model:

(6)

$$\alpha \approx \alpha_0 + \alpha_1 (T - T_0)$$

 α is the thermal expansion. Once the thermal expansion is calculated, we have to consider a temperature dependent bulk modulus written as:

(7)

$$K_T = K_0 + \frac{\partial K}{\partial T} (T - T_0)$$

Then, the thermal pressure can be expressed as:

(8)

$$P^{th}(V,T) = \int_{-\infty}^{T} \alpha k_1 T$$

The results are presented in Table 2 and Figure 1a. Our preferred model is $V_0 = 33.62$ cm³/mol/f.u., $K_0 = 180.6$ GPa, $K'_0 = 4$. $\alpha_0 = 3.245*10^{-5}$ K⁻¹, $\alpha_1 = 1.254*10^{-8}$ K⁻² and $\partial K/\partial T = -0.022$ GPa.K⁻¹. All parameters were refined simultaneously except V_0 which was fixed to the value reported by Guen icult al. (2010) which was obtained with high precision measurement on a single crystal. The refined K'₀ was found very close to 4 (4.06). We decided to fix it to 4 in the final refinement process.

We also calculated use relative error between the volume measured during the experiments and obtained by refining the diffraction patterns and the volume calculated with the obtained equation of state. The minimum and maximum error we obtained were 0.7% and -1.2%, respectively, which corresponds to a volume difference of approximately 0.199 and - 0.334 cm³/mol/f.u. (Figure 1b). The median error of our data is -0.07%. Thus, we can consider that the parameters of the thermal equation of state of CaTiO₃ are well defined with a χ^2 of 1.03.

By comparing our P-V-T data to other thermal equation of states, we extend the $CaTiO_3$ dataset to higher pressure and temperature and find a good agreement with other perovskite structures (Table 3). Our results on the thermal equation of state are comparable to thermoelastic parameters from other CaTiO₃ perovskite studies (Guennou et al. 2010, Webb et al. 1999, Kung & Rigden 1997, Fischer et al. 1993) done with different techniques (i.e.

ultrasonic interferometer, single crystal XRD in LH-DAC and powder XRD in LH-DAC) and extend the studied pressure range. The bulk modulus obtained in the present study is slightly higher but in good agreement with other studies (Webb et al. 1999). They found the same dK/dT as in the present study; however, they refined their thermal equation of state with constant thermal expansion (α_0) at ambient pressure whereas we include a temperature dependence term (α_0 and α_1) for a better fit. When refined with a constant thermal expansivity at 1 bar for comparison, we obtain a value of the same order (α =3.80 10⁻⁵ K⁻¹) as in Webb et al. (1999). Compared to other titanate perovskites (such as SrTiO₃ or BaTiO₃) (Guennou et al. 2010, Pruzan et al. 2002, Webb et al. 1999, Fischer et al. 1993), CaTiO₃ is less compressible than SrTiO₃ and BaTiO₃, which can be at first order a consequence of the cation size.

Importantly, when compared to other perovskite structures present in the Earth mantle (*i.e.*, MgSiO₃, CaSiO₃) (Tange et al. 2012, Ricolleau et al. 2009, Fiquet et al. 2000, Shim et al. 2000), CaTiO₃ is more compressible but its compressible view diminishes less than MgSiO₃ and CaSiO₃ as temperature increases.

These data have implications for planetry science. They confirm that perovskitestructured titanates are less incompressible than perovskite-structured silicates (table 2) opening an interesting crystal-chemical perspective to better understand the effect of substitutions of silicon in the perovskite octahedral site. Moreover, such measurements open the way to a future thermodynamic modeling of CaTiO₃ perovskite at high pressure and temperature which might have some relevance to better understand the behaviour of titanium at lower mantle conditions.

3.2 Orthorhombin and cubic perovskite structures of CaTiO₃

At ambient pressure and high temperature, the orthorhombic *Pbnm* phase of CaTiO₃ is stable up to 1500K (\pm 30K) (Naylor & Cook, 1946; Guyot et al., 1993; Redfern, 1996; Yashima & Ali, 2009) and further transforms into a cubic phase (space group *Pm-3m*) at higher temperature. At ambient temperature and high-pressure, Guennou et al. (2010) observed that the *Pbnm* structure was stable at least up to 60 GPa, a pressure range that we have extended up to 110 GPa in the present study (Figure 2a). However, we observe a discontinuity of the c axis in the compression behaviour of the orthorhombic perovskite *Pbnm* structure at 60 GPa with a change in slope of the evolution of the unit cell parameter as a function of pressure (Figure 2b). Each cell parameter was obtained from a LeBail fit to our diffractograms using the FullProf software. In figure 2b, cell parameters are presented in a

pseudo-cubic lattice geometry for easier comparison. This discontinuity in compression cannot be ascribed to non-hydrostaticity in the Ne pressure medium which has no reported rheological nor phase change in this pressure range, but rather reveals a change in compression mechanism, possibly associated to a thermodynamic path that progressively drives the perovskite structure outside its stability field. At room-temperature though, it is likely that energetic barriers cannot be easily overcome to reach a new thermodynamic stable state. Up to 60 GPa in Figure 2b, CaTiO₃ is still stable. Above 90 GPa, a axis becomes almost incompressible. Also, b axis remains constant on nearly the whole compression up to 100 GPa. However, it is clear that c axis experiences a change in compression above 60 GPa. It is striking to note that these changes in the compression slopes co... spond to pressures at which a major transition is observed in CaTiO₃ at high pressure an ¹ ter perature (see below). Such behaviour has been described in other systems: a stro-gly distorted metastable phase is described for instance in cold compressed forsterite (Finke'stein et al., 2014).

The structure of the orthorhombic *Pbnm* per *pyshite* is made of tilted TiO₆ octahedra and of strongly distorted CaO₁₂ dodecahedral sites. It is known from many studies (Glazer, 1972; Andrault & Poirier, 1991; Angel ., al., 2005) that tilt and distortions change as a function of pressure and temperature and us it the orthorhombic structure tends to transform to a cubic structure with increasing temperature. At pressure below 60 GPa, as temperature increases, we indeed observe a transformation of the Pbnm orthorhombic CaTiO₃ perovskite into a more symmetric cubic structure (Figure 3). However, in most diffraction patterns, we still detect relics of the orther horribic structure mixed with the cubic phase. These relics are seen in the form of a doubling of the peaks corresponding to the cubic phase with no appearance of the orthorhombic peaks between 2.5 Å (10°2 θ) and 1.9 Å (11.5°2 θ) and can be attributed to a gradient of emperature during the experiment. A tetragonal phase (I4/mcm) has been reported in literature (e.g., Guyot et al., 1993; Redfern, 1996; Yashima & Ali, 2009) at intermediate temperatures between orthorhombic and cubic phases. Its restricted temperature stability range may however explain why it was not observed in the present study. According to Redfern (1996) or Yashima & Ali (2009), the existence of a tetragonal phase (I4/mcm) between the perovskite and the cubic domain is shown in the diffractograms by the disappearance of the Pbnm reflections 120, 210, 121 and 103 in the d-spacing range between 2.37 and 2.10 Å, which only leaves the 211 reflection. It is clear from our diffractograms that the 211 reflection was never observed, so we conclude that either the tetragonal phase only exists in a narrow range of temperature that we missed here, or alternatively that the

tetragonal domain no longer exists at high pressure. A similar transition detailed by Komabayashi et al. (2007) is observed in CaSiO₃ perovskite which transforms to a $Pm\overline{3}m$ cubic phase, from a *I4/mcm* tetragonal phase, with a slightly positive dT/dP slope (1.1 K/GPa) and at relatively low temperature (≈ 500 K) compared to CaTiO₃ in the same range of pressure. The study also shows that the incorporation of aluminium stabilizes the tetragonal phase to higher temperature (Kurashina et al. 2004). Thomson et al. (2019) found the same stabilization process with the incorporation of titanium in CaSiO₃ which brings an interesting analogy between CaTiO₃ and CaSiO₃.

By comparing the measured unit cell parameters of the cubic perovskite phase of CaTiO₃ and those calculated using the thermal equation of state of the orthorhombic perovskite, we found that the cubic phase is slightly smaller than the *Pbnm* phase extrapolated at these temperatures (≈ 3500 K). Indeed, the difference because the calculated and measured volume of the cubic phase, using the perovskite the nal equation of state in the same condition, is around -8% in volume (Table S1).

3.3 Melting of CaTiO₃ perovskite

Several studies have attempted to retermine the melting temperature of $CaTiO_3$ at ambient pressure providing estimates ranging between 2188 K and 2262 K (Coughanour et al., 1954; DeVries et al., 1954; R. S. J.C.h, 1958, Jongejean & Wilkins, 1970; V. Daněk & I. Nerád, 2002). To anchor the melting curve at 1 bar, we have retained the value provided by Coughanour et al., i.e. 2188 K which was also selected by Robie et al. 1978.

In Figure 4, we present the *ittel* solid-liquid line using two models for comparison, one from Kraut & Kennedy, 1956, nd one from Simon, 1937.

(9)

$$T_m = 2188 * \left[1 + 4.23 * \left(\frac{\Delta V}{V_0} \right) \right] \quad (Kraut - Kennedy)$$

(10)

$$Pm = 65.66 * \left[\left(\frac{T_m}{T_0} \right)^{1.05} - 1 \right] \quad (Simon)$$

Residuals of the two fits are given in Table S2.

Those melting results are compared with published experimental or calculated melting curves of MgSiO₃ (Zerr & Boehler, 1993; DiPaola & Brodholt, 2016) and CaSiO₃ (Zerr et al., 1997; Braithwaite & Stixrude, 2019) silicate perovskites. As shown in Figure 4, in the pressure range between 20 and 60 GPa, the melting curve of bridgmanite determined using a liquid-solid interface atomistic simulations (DiPaola & Brodholt, 2016) is indeed very close to the experimental melting curve of CaTiO₃. However, the experimental data from Zerr & Boehler (1993) suggest a different behaviour of MgSiO₃ with a steeper slope and higher melting temperatures. CaSiO₃ perovskite has a higher melting temperature than other perovskites and it has been determined both experimentally using CO₂ LH-DAC (Zerr et al., 1997) and using DFT calculations (Braithwaite & Stixrude, 2C¹9). Experimental slopes of melting curves of MgSiO₃ and CaSiO₃ appear steeper than in C₆ TiO₃. Unfortunately, to our best knowledge, no other melting studies on titanates (i.e. Ba TiO₃, SrTiO₃ MnTiO₃) at high pressure and temperature have been conducted. Therefore, the present work provides the first insights on the melting of titanate perovskites at high pressures.

It is interesting to note that our results or rue ting of $CaTiO_3$ perovskite under pressure suggest that titanate and silicate perovskite have melting temperatures of the same order of magnitude which could imply relatively teak dependencies of the melting temperature of mantle perovskites upon substitution of silicon in the octahedral sites. But these speculations remain to be confirmed experimentally and by ab initio simulations.

3.4 Dissociation of CaTiO in CaO and CaTi₂O₅.

At pressures exceeding 50-60 GPa, x-ray diffraction patterns of CaTiO₃ perovskite drastically change upon reating (Figure 5), which suggests a transition to another phase or an assemblage of new phase. We first analysed the high pressure diffractograms ranging from 90 to 170 GPa which showed no traces left of the former *Pbnm* orthorhombic initial perovskite structure. We used the tool DICVOL (Louër & Boultif, 2004) in the FullProf Suite to test several possible structures (space groups *Pmmm*, *Pnma*, *P2*₁/*m*, *P2*₁/*c*, *Pcca*, etc). We also tried to index those diffractograms with a post-perovskite CaIrO₃-type *Cmcm* phase, as observed for instance in MgSiO₃, but we could not find any satisfactory match with the data. We also extended our research by testing other structures found in high-pressure ABX₃ systems before or after dissociation (cotunnite, brookite, baddeleyite, pseudobrookite, stibnite, CaSi₂O₅, Ca₂SiO₄, SrSi₂O₅, Fe₂P, etc) but again, without any satisfying results. The best fit was obtained using a *P2/m* space group for all diffractograms over the whole pressure range

including high temperature and quenched to ambient temperature data. This lattice volume is of around 300 Å³. This volume is much larger than any unit cell volumes of CaTiO₃ extrapolated to those conditions (≈ 170 Å³), thus indicating that the stoichiometry is different from CaTiO₃. Along these lines, this volume of 300 Å³ would perfectly match a stoichiometric volume of CaO + CaTi₂O₅ in those conditions.

In all high-temperature diffractograms at pressures above 90 GPa, two reflections are also consistent with a CaO-*B2* phase, according to the equation of state for *B2*-phase of CaO (Yamanaka et al., 2002) corrected for temperatures. One example of the refinement is given in Figure 6, where the two co-existing phases can be identified in that pressure domain.

Interestingly enough, CaO transforms from NaCl-*B1* to CsCl-*B2* structures at around 60 GPa at 300 K (Richet et al. 1988, Yamanaka et al 2072), which corresponds to the pressure under which CaTiO₃ starts to transform into the *P2/m* phase. Accordingly, we think the dissociation is likely driven by the change in structure of CaO.

Following these observations reinforced by an analogy with Wu et al. (2009) who have suggested dissociation of a FeTiO₃ pervskite into $(Fe_{1-\delta},Ti_{\delta})O$ and $Fe_{1+\delta}Ti_{2-\delta}O_5$, we propose that above 60 GPa and at high variable relations. (*B2*) and CaTi₂O₅ instead of going through a pPv phase, following the dissociation reaction:

(11)

$$CaTiO_3 \rightarrow \frac{1}{2}CaO + \frac{1}{2}CaTi_2O_5$$

To validate this hypothesis, we notice that in addition to those of the P2/m phase and of B2 - CaO, we cannot observe any diffraction line of TiO₂. The summary of the CaTiO₃ phases observed at different pressures and temperatures are all represented in a data-plot diagram in Figure 7a and a tentative phase diagram Figure 7b.

This dissociation of the perovskite $CaTiO_3$ without going through a post-perovskite phase is an important experimental demonstration of the importance of simple oxides for the thermodynamic evolution of ABX₃ systems. Tsuchiya et al. (2015) predicted a new Fe₂P-type phase for cotunnite-SiO₂ around 0.9 TPa which could impact ABX₃ structures by favoring their dissociations in the silicate mantles (MgSiO₃, CaSiO₃) of super-Earths. Also, there is still no evidence of a CaIrO₃-type post-perovskite phase in CaSiO₃ up to 149 GPa (Fujino et al. 2009). Though, the appearance of such a complex oxide as CaTi₂O₅ is a transformation

that has similarities with the prediction made for $MgSiO_3$ and the stabilization of $MgSi_2O_5$ at extreme conditions, according to Umemoto et al. (2011, 2017).

In every high-pressure transition from a perovskite to a post-perovskite phase or to an assemblage containing oxides, a negative volume variation of around 1-2% has been observed (e.g., Hirose et al., 2005 and references therein). Considering the stoichiometry of equation (11), the volume of CaTiO₃, extrapolated to high-pressure and high-temperature (\approx 164Å), is slightly higher than that of $\frac{1}{2}$ CaO + $\frac{1}{2}$ CaTi₂O₅ (\approx 156Å) in the same conditions. Using our high-pressure high-temperature equation of state of CaTiO₃ and that of CaO as B2 (Yamanaka et al., 2002), we infer that the volume variation of reaction (11) is between -1% and -5% depending on the temperature. Although a high-pressure ₁ olymorph of CaTiO₃ with CaIrO₃ structure was predicted by first-principles calculations (Ya et al., 2015), the results of the present experimental study suggest that breakdown a coording to reaction (8) bypasses the expected post-perovskite phase and that dissociation to two other phases is favoured.

Between 60 and 90 GPa, the CaTiO₃ perovekute structure can still be recognized in our patterns in coexistence with an assemblage of CaO-B2 and $CaTi_2O_5-P2/m$. The CaTiO₃ reflections weaken while the reflections attributed to $CaTi_2O_5$ grow with increasing pressure. In this transition pressure range, however, te v peaks remain unidentified. They could be due to yet unidentified defect structures $dot v_1 d d$ from CaTiO₃ perovskite, in a manner possibly similar to those reported in MgSiO₂ d v_1 Schauner et al. (2008). Another important aspect of this transition mechanism is the effect of temperature on the dissociation process given in equation (11). At ambient temperature, the cold compression of CaTiO₃ is unable to overcome energy barriers needed to drive the transition into CaO and CaTi₂O₅. We observe instead a change in the compression regime of the lattice parameters. As temperature is increased, the main peaks of the *P2/m* phase start to appear but this new phase still coexists with CaTiO₃ perovskite for kinetic reasons over a 30 GPa pressure range. It is important to note that the complex mixed phase regime appearing in Figure 7 is prohibited according to thermodynamics and thus, draws its origin in kinetic effects.

Other authors have also addressed the complexity of different transition processes in ABO₃ compounds. Wu et al. (2009) have shown that ilmenite FeTiO₃ goes through a perovskite phase but then dissociates into a wüstite-type (Fe_{1- δ},Ti_{δ})O and a monoclinic *C2/c* Fe_{1+ δ}Ti_{2- δ}O₅ structure, thus omitting the post-perovskite phase. More recently, Akaogi et al. (2016) proposed a dissociation of FeTiO₃ into Fe₂TiO₄ + TiO₂ or Fe₂TiO₄ + FeTi₂O₅

depending on the temperature. They also highlighted the dissociation process of ZnTiO₃ (Akaogi et al. 2015) where the ilmenite structure transforms into a perovskite and then dissociates directly into its simpler oxides with a rocksalt-type ZnO and a baddeleyite-type TiO₂. Kojitani et al. (2005) studied the dissociation of SrSiO₃ into Sr₂SiO₄ and SrSi₂O₅ at 16 GPa and proceeded with a Rietveld analysis of the obtained species. In 2010, Okada et al. (2010) found in MnTiO₃ the same process of decomposition as shown in the present study for CaTiO₃ but with an orthorhombic MnTi₂O₅ structure with lattice parameters indeed close to those of FeTi₂O₅ from Akaogi et al. (2016). In addition, Kondo et al. (2000) described a high-pressure phase transition in MnO from *B1* to a distorted *B1* at 30 GPa, which echoes with the dissociation of MnTiO₃ into MnO and MnTi₂O₅.

These data show that titanates tend to dissociate more enaily without going through a post-perovskite phase whereas germanates adopt more casily post perovskite phases. (Ito & Matsui, 1979). Tateno et al. (2006) describe a transition from perovskite to a CaIrO₃-type post-perovskite phase in MnGeO₃ at 60 GPa and '80° K, exactly the same process as in MgGeO₃ at 63 GPa and 1800 K (Hirose et al., 1005). More recently, ZnGeO₃ also joined the Pv-pPv transition group with a similar process as MgSiO₃ as it transforms around 114 GPa (Yusa et al., 2014). In 2017, However, Dunca et al. triggered the dissociation of FeGeO₃ into FeO and GeO₂ at high temperature without adopting a pPv structure, following the same transition path as in FeSiO₃ (Dutta even, 2017).

Based on those literative observations and on the experimental data of the present study, we thus confirm that whom chemical bonding characteristics are met as it is the case in titanates, ABO₃ systems often break down into different oxides compounds. As the dissociation of CaTiO₃ (Pv) without transition to a post-perovskite phase is a phenomenon that has already been seen in other ABO₃ systems, it further brings us to the conclusion that this type of transition is common, especially in ATiO₃ systems (table 4). Moreover, according to Akaogi et al. (2016) (and references therein), it seems that the titanate perovskites do not transform into post-perovskite but instead dissociate into oxides, a behaviour seemingly related to a small degree of covalency of Ti-O bonds.

4 Implications for planetary interiors

It is interesting to compare our observation of the dissociation of CaTiO₃ with the prediction by Umemoto et al. (2011) using atomistic calculations of a dissociation of MgSiO₃ into MgO and MgSi₂O₅ phase with a $P2_1/c$ -type structure at pressures of the order of 900

GPa. Moreover, Tsuchiya & Tsuchiya (2015) predicted the dissociation of $CaSiO_3$ and $MgSiO_3$ into their corresponding oxides, with the change in structure of SiO_2 as a possible cause, which echoes with the dissociation of $CaTiO_3$ likely favored by the CaO *B1-B2* transformation.

To summarize, the decomposition of several perovskites (titanates, germanates and silicates) under pressure is often observed into a combination of their simplest constitutive oxides (AO and BO₂). However, a thorough examination of the literature reveals more complex stoichiometries of these breakdowns. With CaTiO₃, these transformations and dissociations can be observed in a pressure range accessible to accurate experimental studies which could usefully shed light on certain aspects of the evolution of MgSiO₃ or CaSiO₃ compounds at ultra-high pressures in exoplanets. Moreover, further detailed study of the CaTi₂O₅ phase could usefully shed light on some properties of ASi₂O₅ compounds in super-Earth mantles

5 Conclusion

In conclusion, we explored the high-pressure high temperature transformations of CaTiO₃ perovskite up to 170 GPa and 4.5% K. We measured the thermal equation of state in the stability domain of CaTiO₃ perovskite. We also measured the melting curve of this perovskite up to 55 GPa. At higher predicted and temperatures, at odds with what had been predicted by theoretical calculations, we show that CaTiO₃ does not transform into a post-perovskite CaIrO₃ *Cmcm* structure but instead dissociates into *B2*-CaO and monoclinic CaTi₂O₅. The transition process starts at 60 GPa and is spread, for kinetic reasons, over a large range of pressure. The ransition toward *B2*-CaO and CaTi₂O₅ is complete at pressures around 100 GPa.

Strengthening knowledge of the evolution of ABX₃ structures at high pressures is of first importance to infer the mineralogical state of exoplanetary mantles. We propose that future studies of the CaTi₂O₅ phase discovered in this work in a pressure range amenable to precise experimental studies will provide insights into exoplanetary ultra-high pressure phenomena that are still beyond reach of experimental mineral physics studies.

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Figures



Figure 1: a) Thermal equation of state of $CaTiO_3$ reconskite with isotherms from 300 K to 2600 K, along with experimental data used in the refinement. This dataset includes: Guennou et al. (2010) data from 0 to 40 GPa (black do s), data from Yashima et al. (2008) at ambient pressure (represented by the coloured trigges), our own data collected at room-temperature from 40 to 47 GPa (black diamonds) and data collected at high-temperature from 30 to 60 GPa (coloured diamonds) b) Relative error between the volume calculated with the thermal EoS and the volume measured on each diffractograms. The colour scale corresponds to the measured temperatures in K.



Figure 2: a) Room temperature X-ray diffraction pattern? of CaTiO₃ perovskite as a function of pressure with peak positions indicated in orange; here we pressure medium peak is marked in green and main peaks are masked so as to evidence details among lower intensity reflections; b) Room-temperature CaTiO₃ cell (*x*), evolution as a function of pressure. Cell parameters are recalculated within a pseudo-cubic cell from the *Pbnm* structure where $a_{pc}=a_{pv}/\sqrt{2}$, $b_{pc}=b_{pv}/\sqrt{2}$, $c_{pc}=c_{pv}/2$ (pc: pse add -cubic; pv: perovskite). The straight dotted lines are guide to the eye to evidence the change in the compression behavior. We observe a change in the slope at 90 GPa par axis a and b, and a change at 60 GPa for axis c.

5000



Figure 3: Orthorhombic perovskite to c² bic structure transition. At the bottom, the perovskite (blue ticks) diffractogram with c² en circles corresponding to the platinum reflections. On top, the diffractogram of the cubic ² hase (orange ticks) with open triangles corresponding to the KCl reflections.



Figure 4: Our melting curve of C ΣL_3 plotted along with melting curves of CaSiO₃ and MgSiO₃ for comparison. The study from Zerr & Boehler (1993) and Zerr et al. (1997) have been also performed with DAC experiments, while the study from DiPaola & Brodholt (2016) and Braithwaite & Stixrude (2016) are from atomistic simulations. The black circular dots are our experimental data and the green and black lines are Simon's and Kraut-Kennedy's fits of our data points. The data at a nbient pressure come from Coughanour et al. (1954) for CaTiO₃ and from Basmajian & $\Box \nabla$ Vries (1957) for SrTiO₃ and BaTiO₃.



Figure 5: Examples of diffractograms t pressures from 41 to 124 GPa and temperature between 1500 K and 2500 K sho *vi* ng the evolution from pure CaTiO₃ at low pressures (orange ticks) to CaTi₂O₅ (blue ti :ks, + CaO (red ticks) mixture at the highest pressures. Open black circles represent the plati un. Open red circles show regions where new reflections are first observed, until we observe the full crystallisation of CaTi₂O₅ mixed with CaO at pressures above 100 GPa. The diffractogram at 76 GPa shows a coexistence between the CaTiO₃ perovskite phase and the appearance of what we assume are the reflections from the CaTi₂O₅ + CaO phase. The diffractogram at 91 GPa shows the reflections of the CaTi₂O₅ + CaO phase in formation b it not fully formed due to kinetic reasons.



Figure 6: Full profile LeBail fn of a high-pressure diffraction pattern recorded at 122 GPa and 2900 K (series #2, Table 1). The green ticks correspond to P2/m-CaTi₂O₅ structure (a=7.78 Å, b=5.36 Å, c=1.18 Å and β =91.24°), the orange ticks to B2-CaO (a=2.53 Å) and the grey ticks correspond to the platinum (a=3.65 Å). $\chi^2 = 0.055$. Formular unit Z=4. Data available here : 10.17632/r.nkzfcdh75.3



Figure 7: Summary of the $CaTiO_3$ phases observed at different pressures and temperatures.

a) Summary of XRD data collected at different pressures and temperatures. Yellow dots: CaTiO₃ orthorhombic (Pbnm) perovskite phase. Yellow triangles: neutron diffraction data from Yashima et al. (2009) at ambient pressure and up to 1500K. Green dots: CaTiO₃ cubic perovskite phase Pm $\overline{3}$ m. Red symbols: mixed phase regime of kinetic

origin containing $CaTiO_3$ perovskite (Pbnm) together with a higher-pressure assemblage marked as blue dots when devoid of perovskite signal. This high-pressure assemblage consists of CaO (B2) and $CaTi_2O_5$ P2/m. Each pressure condition was measured either with KCl, platinum or gold pressure standards.

b) Tentative phase diagram of CaTiO₃, according to X-ray diffraction data collected in this study at different pressure-temperature and to previous literature data. The hypothetical phase boundaries have been drawn without fit, using data of figure 7a. The melting curve of CaTiO₃ perovskite was fitted with a Simon model. The tetragonal phase intermediate between orthorhombic and cubic perovskite structures was not represented.

Series #	CaTiO ₃ type	Origin	Pressure transmitting medium	Pressure scale	Pressure domain (GPa)	Temperature domain (K)	Diamond culets
1	Powdered	Synthetic	$CaTiO_3 + Pt$	Pt	30 - 78	1500 - 4000	150/300
2	Powdered	Synthetic	$CaTiO_3 + Pt$	Pt	83 - 120	2000 - 3000	100/300
3	Powdered	Synthetic	$CaTiO_3 + Pt$	Pt	121 - 155	2000 - 4400	250
4	Polycrystalline	Natural	KCl	KC1	35 - 75	2100 - 3500	200
5	Polycrystalline	Natural	KCl	KC1	47 - 115	2100 - 4400	100/300
6	Polycrystalline	Natural	KCl	KCl	127 - 161	1400 - 4000	100/300
7	Polycrystalline	Natural	KCl	KCl	154 - 168	2800 - 4500	150/300
8	Polycrystalline	Natural	Ne	Au	28 - 109	300	150/300

Tables

Table 1: Samples used for the different experimental runs.

V_0 (cm ³ /mol/f.u.)	33.622 (fixed)
K_0 (GPa)	180.6 (4)
K'0	4 (fixed)
dK/dT (GPa.K ⁻¹)	-0.022 (1)
$\alpha_0 (x 10^{-5} \text{ K}^{-1})$	3.25 (5)
$\alpha_1 (x 10^{-8} \text{ K}^{-2})$	1.3 (1)

Table 2: Parameters of the $\alpha_1(\chi_{10} \times \chi) = 1.5(1)$ thermal equation of state of CaTiO₃ ($\chi^2 = 1.03$), with V₀ fixed and set to the value determined by Guennou et al. (2010) and K'_0 fixed to 4.

	V ₀	K. (Gpa) K'.		dK/dT (GPa K	$\alpha_0 (10^{-5} \text{ K}^{-1})$	$\alpha_1 (10^{-8} \text{ K}^{-1})$	
	(cm ³ /mol/f.u.)	K ₀ (Opa)	\mathbf{K}_{0}	1)	1)	²)	
CaTiO ₃	_						
This study	33.62 (7) ^a	180.6 (4)	4 (fixed)	-0.022 (1)	3.25 (5)	1.3 (1)	
Guennou							
(2010)	33.62 (7)	181.0 (6) 172.1	4 (fixed)	-	-	-	
Webb (1999)	33.82	(10)	6.09^{b}	-0.0225 (7)	3.87	-	
Kung (1997)	34.17 (2)	175	5.78	-	-	-	
Fischer (1993)	33.87	177 (3)	5.1 (8)	-	-	-	
SrTiO ₃ *	_						
Guennou	25.02 (1)	1.55 (2)					
(2010)	35.93 (1)	165 (3)	6.4 (8)	-	-	-	
Webb (1999)	35.89	172.0 (1)	5.36 ^b	-0.031 (1)	$3.23(2)^{c}$	-	
$E_{1}^{(1)} = \frac{1}{2} \left(\frac{1}{2} - \frac{1}{2} \right)^{-1}$	25.07	178.8	4 21 (1)				
Fischer (1993)	35.87	(46)	4.31 (1)		-	-	
PaT:O **							
$\frac{\text{Darron}}{\text{Darron}} (2002)$	- 20 70	125	61				
FIUZall(2002)	36.12	139 2	0.4		-	-	
Fischer (1993)	38.74	(31)	9 ^d	_	-	-	
1		(01)					
MgSiO ₃							
	-	256.7					
Tange (2012)	24.45 (fixed)	(15)	-` 09 (6)	-	-	-	
Ricolleau							
(2009)	24.69 (fixed)	245 、1	• 4 (fixed)	-0.036 (1)	3.19 (17)	0.88 (16)	
Fiquet (2000)	24.43 (1)	259.5 (9)	3.69 (4)	-0.017 (2)	2.18 (12)	0.11 (8)	
CaSiO ₃	_						
K1colleau	27 16 (fine d)	244(1)	1 (fine 1)	0.025(2)	2.06(10)	0.97(19)	
(2009)	27.40 (f1xed)	244 (1)	4 (11xea)	-0.035 (2)	3.00 (19)	0.87 (18)	
Shim (2000)	21.45 (3)	236 (4)	3.9 (2)	-0.028 (11)	2.2 (3)	-	

^a From Guennou et al. (2010) – Room temperature DAC with x-ray diffraction.
^b From Fischer et al. (1993) – Extrapolated from 500 MPa. Piston cylinder apparatus and ultrasonic interferometric technique.

^c From de Ligny & Richet (1996) – Calorimetry and energy-dispersive x-ray diffraction.
^d From Fischer & Polian (1987) – Brillouin scattering in DAC.
* Data for the cubic phase (0-10 GPa)

** Data for the tetragonal phase (0-2 GPa)

Table 3: Thermal equation-of-state parameters of different titanate, bridgmanite and calcium silicate perovskite.

Material	Pressure	geometry	a (Å)	b (Å)	c (Å)	β (°)	$V(Å^3)$
CaTi ₂ O ₅	120 GPa	Monoclinic <i>P2/m</i>	7.776	5.362	7.167	91.29	298.75
	155 GPa	Monoclinic <i>P2/m</i>	7.662	5.268	7.049	91.422	284.43
NaMg ₂ F ₅ (Dutta, 2019)	88 GPa	Monoclinic $P2_{l}/c$	5.618	4.977	7.729	89.23	197.08
	137 GPa	Monoclinic $P2_1/c$	5.423	4.732	7.312	89.2	187.62
FeTi ₂ O ₅ (Akaogi, 2016)	31 GPa HT°	Orthorhombic	8.567	5.753	5.257	-	259.10
$\begin{array}{c} (\text{Intage}, 2010) \\ \text{MnTi}_2\text{O}_5 \\ (\text{Okada, 2010}) \end{array}$	61 GPa HT°	Orthorhombic	8.263	5.743	5.179	_	245.77
FeTi ₂ O ₅ (Wu, 2009)	40 GPa HT°	Monoclinic C2/c	10.135	4.305	10.023	141.48	272.37
SrSi ₂ O ₅ (Kojitani, 2005)	16 GPa HT°	Orthorhombic Cmca	5.239	9.280	13.441	_	653.47

Table 4: Comparison of the AB_2X_5 structures and lattice paral. etcrs.

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Declaration of interests

The authors declare that they have no known competing financial interests or personal relationships that could have appeared to influence the work reported in this paper.

Graphical abstract



Highlights:

- Phase diagram of the calcium titanate perovskite CaTiO₃ up to 170 GPa and 4500 K.
- Thermal equation of state of the orthorhombic (*Pbnm*) perovskite CaTiO₃.
- Dissociation of the CaTiO₃ perovskite into CaO and CaTi₂O₅ without going through a post-perovskite phase.