

Arctic methylmercury cycling

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1 Review 2 Arctic Methylmercury Cycling 3 4 Sofi Jonsson¹⁾*, Michelle Nerentorp Mastromonaco²⁾, Feiyue Wang³⁾, Andrea G Bravo⁴⁾, 5 Warren RL Cairns⁵⁾, John Chételat⁶⁾, Thomas A. Douglas⁷⁾, Gretchen Lescord⁸⁾, Liisa 6 Ukonmaanaho⁹, Lars-Eric Heimbürger-Boavida¹⁰) 7 8 9 10 11 1. Department of Environmental Science, Stockholm University, SE-106 91 Stockholm, Sweden 12 2. IVL, Swedish Environmental Research Institute, SE-411 33, Gothenburg, Sweden 13 3. Centre for Earth Observation Science, Department of Environment and Geography, University of 14 Manitoba, Winnipeg, Manitoba, Canada 15 4. Department of Marine Biology and Oceanography, Institut de Ciències del Mar (ICM-CSIC), 16 Barcelona, Spain 17 5. CNR Institute of Polar Sciences and Ca' Foscari University, Venice, Italy 18 6. Environment and Climate Change Canada, National Wildlife Research Centre, Ottawa, ON Canada 19 7. U.S. Army Cold Regions Research and Engineering Laboratory Fort Wainwright Alaska USA 20 8. Wildlife Conservation Society Canada and Laurentian University (Vale Living with Lakes Center, 21 Sudbury Ontario, Canada 22 9. Natural Resources Institute Finland (Luke), P.O. Box 2, FI-00791, Helsinki, Finland 23 10. CNRS/INSU, Aix Marseille Université, Université de Toulon, IRD, Mediterranean Institute of 24 Oceanography (MIO), Marseille, France 25 26 *corresponding author: sofi jonsson, sofi.jonsson@aces.su.se Abstract 27 28 29 Anthropogenic mercury (Hg) undergoes long-range transport to the Arctic where it is transformed into 30 methylmercury (MeHg), potentially leading to high exposure in some Arctic inhabitants and wildlife. 31 The environmental exposure of Hg is determined not just by the amount of Hg entering the Arctic, but 32 also by biogeochemical and ecological processes occurring in the Arctic. These processes affect MeHg uptake in biota by regulating the bioavailability, methylation and demethylation, bioaccumulation and 33 34 biomagnification of MeHg in Arctic ecosystems. Here, we present a new budget for pools and fluxes 35 of MeHg in the Arctic and review the scientific advances made in the last decade on processes leading 36 to environmental exposure of Hg. Methylation and demethylation are key processes controlling the

pool of MeHg available for bioaccumulation. Methylation of Hg occurs in diverse Arctic

environments including permafrost, sediments and the ocean water column, and is primarily a process carried out by microorganisms. While microorganisms carrying the *hgcAB* gene pair (responsible for Hg methylation) have been identified in Arctic soils and thawing permafrost, the formation pathway of MeHg in oxic marine waters remains less clear. Hotspots for methylation of Hg in terrestrial environments include thermokarst wetlands, ponds, and lakes. The shallow sub-surface enrichment of MeHg in the Arctic Ocean, in comparison to other marine systems, is a possible explanation for high MeHg concentrations in some Arctic biota. Bioconcentration of aqueous MeHg in bacteria and algae is a critical step in the transfer of Hg to top predators, which may be dampened or enhanced by the presence of organic matter. Variable trophic position has an important influence on MeHg concentrations among populations of top predator species such as ringed seal and polar bear distributed across the circumpolar Arctic. These scientific advances highlight key processes that affect the fate of anthropogenic Hg deposited in Arctic environments.

1. Introduction

Mercury (Hg) is a bioaccumulative neurotoxin that is released from the earth's crust through natural
processes and anthropogenic activities (Streets 2019). Since the industrial revolution, direct
anthropogenic emissions have exceeded natural emissions by an order of magnitude and concentration
of Hg in the atmosphere has increased 3-5 times (Driscoll et al. 2013, Li et al 2020). Since Hg is well
mixed globally in Earth's atmosphere these emissions have also reached the Arctic environment. Due
to a variety of physical, chemical, and biological processes, some of which are unique to environments
in cold regions, high levels of Hg exposure have been detected among Arctic communities and
wildlife (AMAP 2021).
The biogeochemical cycle that connects Hg emissions to Hg in Arctic food webs involves inorganic
and organic forms of Hg that undergo a wide range of transformation processes in air, water, and land.
There are three main chemical forms of Hg in Arctic environmental compartments and biota: gaseous
elemental mercury (Hg(0)), inorganic divalent mercury (Hg(II)) and MeHg. Figure 1 shows the
partition of each in the major Arctic environmental compartments and biota.
Industrial activities (e.g., artisanal and small-scale gold mining, coal-fired power stations, waste
incineration, cement manufacture, and others) release both Hg(0) and Hg(II). The main natural Hg
sources to the atmosphere are volcanoes and forest fires. While Hg(II) is typically deposited closer to
the emission source, Hg(0) has a residence time ranging from several months to over a year in the
atmosphere and can be transported from the emission source to remote polar regions within a few days
to weeks (Durnford et al, 2010, Dastoor et al. 2015). Atmospheric mercury depletion events, during
polar sunrise occur in both Arctic and Antarctic regions and lead to the rapid oxidation of Hg(0) to the
more easily deposited Hg(II) species. A large fraction of Hg(0) also dissolves in natural waters where
it is referred to as dissolved gaseous mercury (DGM) (Jiskra et al. 2021). On land, foliar (plant) uptake
of Hg(0) drives the atmospheric deposition of Hg in the Arctic tundra (Obrist et al., 2017; Jiskra et al.
2018; Douglas and Blum, 2019). Through this process, large amounts of Hg have accumulated in
northern permafrost soils from the atmosphere over thousands of years. Today, concerns are raised

about the potential remobilization of this permafrost Hg pool as climate warming leads to increased soil temperatures, thawing permafrost, and changes in the hydrological cycle (Schaefer et al., 2020, Lim et al. 2020). In natural waters, soils and sediments, Hg is primarily found as Hg(II) (Skyllberg, 2012). Due to a high affinity for particles (e.g. mineral surfaces and organic matter), most Hg(II) in sediments and soils is found in the solid phase and these systems commonly act as important sinks for Hg. At the same time, these systems are known to be important sites for Hg methylation. The main form of MeHg existing in natural environments is monomethylmercury (MMHg), although, in some environments, dimethylmercury (DMHg) may become the dominant species, especially in seawater. Since most analytical data on methylated Hg do not differentiate MMHg and DMHg, we will use MeHg to refer to methylated Hg in general from hereon, unless otherwise specified. While both MMHg and DMHg are toxic and lipophilic, only MMHg has been shown to biomagnify in aquatic food webs to concentrations of concern for wildlife and human health (Lavoie et al., 2013). Because of this, the potential risk Hg pollution poses to Arctic ecosystems is not only controlled by the amount of inorganic Hg transported into the system and cycling within the Arctic environment, but also by the extent of the Hg pool that is methylated and accumulated in aquatic food webs. In this review and contribution to the Virtual Special Issue on Arctic mercury, we present a synthesis of work conducted for the AMAP Assessment 2021: Mercury in the Arctic on the processes affecting mercury transformation and biotic uptake in the Arctic (AMAP 2021, Chapter 4). Given the high complexity of the Hg biogeochemical cycle, it is not surprising that significant advances are still being made despite over 50 years of active research. Here, we focus on scientific progress made in the last decade on processes leading to environmental MeHg exposure by presenting (1) a new mass balance for MeHg in the Arctic, and new information on (2) our mechanistic understanding of Hg (de)methylation processes, (3) where hotspots of MeHg are found in the Arctic and, finally, (4) processes affecting Hg accumulating in Arctic biota.

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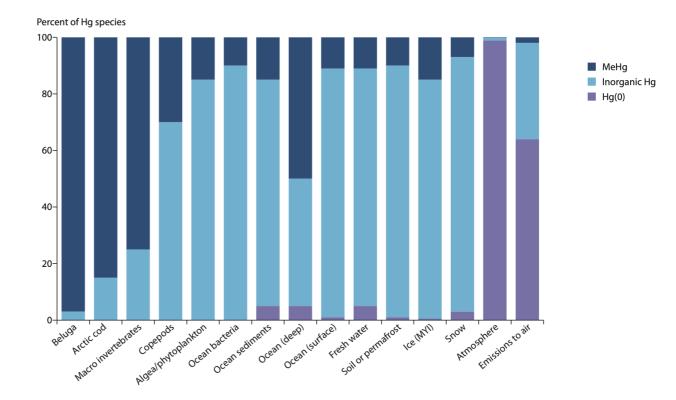


Figure 1. Partition of different Hg species (Hg(0), inorganic divalent Hg(II) and MeHg) in emissions and in various Arctic environmental compartments and biota, expressed as percentages of total Hg (adopted from AMAP 2011, Douglas et al., 2012; Cossa et al., 2014; Schartup et al., 2020).

2. How much methylmercury is circulating in the Arctic environment?

A summary of MeHg concentrations, fluxes and estimated MeHg pools in different environmental compartments used for this mass balance is presented in Table 1, Table 2 and Figure 2. The MeHg budget for the Arctic Ocean was first estimated by Soerensen et al. (2016b) and later revised by Petrova et al. For the AMAP 2021 Hg assessment, THg (Dastoor et al. 2022) and MeHg (this review)

- budgets were calculated for land areas north of 60°N and the Arctic Ocean, including the central basin,
- and the Barents, Kara, Laptev, East Siberian, Chuckhi and Beauford Seas.
- To estimate the stock for the 0-20 m, 20-200 m, 200-500 m and <500 m water masses (Table 2)
- average concentrations of THg and MeHg in seawater calculated from available data (Wang et al.,
- 2012b, 2018; Heimbürger et al., 2015; Agather et al., 2019; Charette et al., 2020; Kim et al., 2020b;
- Petrova et al., 2020; Jonsson et al., 2022, Kohler et al., 2022) were used. For biota, we used the
- estimate previously provided by Soerensen et al., 2016b. Remaining reservoirs (soil, snow, ice and air)
- were estimated from the stocks of Hg (Dastoor et al. 2022; Table 3.8.1 in AMAP (2021), key data for
- this review are also shown in Table S1) and the best estimates of %MeHg for each environmental
- 128 compartment (Table 1 and S1). The MeHg pool in surface soils was estimated by multiplying the total
- Hg pool by the median %MeHg in surface soils (St. Pierre et al., 2015: 0.82%) and wetlands (Tarbier
- et al., 2021: 1.7%, not impacted by recent permafrost thaw) and assuming a 25% coverage of wetlands
- for the Arctic Region (Kåresdotter et al., 2021). Methylmercury reservoirs in snowpack and glaciers
- were estimated using %MeHg measurements for snow (6.1%, range: 0.11-11%; St. Pierre et al., 2019,
- Varty et al., 2021 and Jonsson et al., 2022) and median %MeHg (5.5 %) reported by Zdanowicz et al.,
- 134 (2016) for glaciers, respectively. For sea ice, the mean %MeHg reported by Dimento et al., (2019) was
- used, resulting in an estimated MeHg pool of 0.32 Mg. This value is close to the estimate of MeHg
- 136 (0.33 Mg) previously estimated by Schartup et al. (2020). For the atmosphere, %MeHg were
- calculated from the concentrations of MMHg and DMHg (Table 1) previously reported by Baya et al.
- 138 (2015) and total Hg concentration reported by Dastoor et al. (2022).
- Using geochemical modelling, Soerensen et al., (2016b) estimated MeHg reservoirs in the Arctic
- Ocean (also including the Canadian Arctic Archipelago and the Baffin Bay), sediments and plankton
- resulting in a total of 576 Mg. In the updated MeHg budget presented by Petrova et al. (2020), the total
- pool of MeHg in the Arctic Ocean was estimated to 800 Mg MeHg. Our budget for the Arctic Ocean
- water sums up to 590 ± 341 Mg MeHg and is thus in the same range as earlier estimates.
- Petrova et al. (2020) found a MeHg maxima in the subsurface waters (100–300 m) of the Arctic Ocean
- (in line with observations from Wang et al. (2012b, 2018), Heimbürger et al. (2015), Agather et al.

146 (2019), Jonsson et al. (2022)), leading to a net export of MeHg from the Arctic Ocean to the Nordic 147 Seas and the North Atlantic of 7.5 Mg/yr. A part of the MeHg produced in the subsurface ocean layer 148 (20–200 m; 17 Mg/yr) is converted to DMHg, which diffuses up to the polar mixed layer (PML; 0–20 149 m), where some of it evades to the atmosphere. Here DMHg is proposed to rapidly photochemically degrade back to MMHg and deposit back to the PML. For the AMAP 2021 report, MeHg (MMHg + 150 151 DMHg) evasion was estimated to 4.5 (2.8-6.3) Mg/yr using a DMHg to Hg⁰ evasion ratio of 0.16 (Soerensen et al, 2016b) and the new Hg⁰ evasion estimate (32 (20-45) Mg/yr (Table 3.8.1 in AMAP 152 153 (2021)). The MeHg deposition was then calculated to 2.6 Mg/yr assuming the evasion to deposition 154 ratio (0.57) presented by Soerensen et al. (2016b). In addition, DMHg may be photochemically 155 degraded in surface waters to MMHg (West et al. 2022). However, as our budget deals with the total 156 stock of MeHg, transformation reactions between methylated Hg species within compartments are not 157 considered here. Riverine MeHg inputs (0.8 Mg/yr) were estimated from Sonke et al., (2018) and 158 equal inputs by coastal erosion was assumed (as done for Hg, Table 3.8.1 in AMAP (2021)). The benthic flux of MeHg to subsurface waters was estimated to 0.9-4.6 Mg/yr (Soerensen et al., 2016b, 159 160 Kim et al., 2020b). The in situ net water column methylation (5.1 Mg/yr) was estimated by calculating 161 the difference between estimated inputs (total input: 18.4 Mg/yr; rivers: 0.8 Mg/yr; coastal erosion: 0.8 Mg/yr; atmospheric deposition: 2.6 Mg/yr; oceans: 9 Mg/yr; sediment: 4.7 Mg/yr, sea ice: 0.5 Mg/yr) 162 163 and outputs (total output: 23.5 Mg/yr; atmospheric evasion: 4.5 Mg/yr; ocean inputs: 16.5 Mg/yr; 164 sediment: 1.4 Mg/yr, sea ice: 0.4 Mg/yr; plankton uptake 0.7 Mg/yr). Our in situ net water column 165 methylation for the Arctic Ocean is close to the in situ net water column methylation estimates 166 provided from earlier models (Kim et al., 2020: 10 Mg/yr; Soerensen et al., 2022: 15 Mg/yr for the 167 upper 200 m). Our estimated in situ production corresponds to a daily net Hg methylation potential of 7.5 10⁻⁶ d⁻¹ (calculated using the water column total Hg budget estimate (1,870 Mg Hg) from Dastoor 168 169 et al., 2022). This calculated net Hg methylation potential (net result of in situ Hg methylation and in 170 situ MeHg demethylation processes) is three orders of magnitude lower than the potential Hg methylation rate constants (0.6-13 10⁻³ d-1) reported from incubation studies by Lehnherr et al., (2011). 171

The Arctic MeHg budget highlights the importance of several processes that control MeHg reservoirs. Specifically, MeHg production in the water column and in shelf sediments are dominant inputs to the Arctic Ocean (Figure 2). Atmospheric deposition and marine evasion also represent large fluxes of MeHg, with a net deposition flux. Similarly, fluxes of MeHg via oceanic currents are large, with an estimated net export from the Arctic Ocean to the North Atlantic. The modelling of estimates for pools and fluxes is based on existing data, which have varying degrees of uncertainty. In particular, our *in situ* net Hg methylation potential is associated with high uncertainty, although the much lower *in situ* net methylation used in our model (in comparison to potential methylation rates (k_m) values available) suggests a rapid turnover of MeHg in Arctic waters (Lehnherr et al., 2011 reported potential demethylation rates of 0.2-0.6 d⁻¹). Some budget estimates are based on a "best estimate" approach (when no data was available) or a very limited set of observations. Therefore, improved spatial and temporal coverage of measurements is important to improve estimates of MeHg fluxes, including from meltwater and erosion.

Table 1. Total MeHg concentrations, fluxes and pools in the Arctic environment (excluding aquatic biota larger than plankton). The total MeHg concentrations reported cover both freshwater and the marine compartments, while the fluxes focus on those relevant for the Arctic Ocean.

Environmental Compartment	MeHg Concentration	%MeHg	Description	Region	References
Atmosphere	2.9±3.6 pg/m³ MMHg	<1%		Hudson Bay and CAA	Baya et al., 2015
	3.8±3.1 pg/m ³ DMHg				
Soil	0.02-2.11 ng/g (d.w)	0.82%		Canada	St. Pierre et al., 2015
	0.5 and 3.4 ng/g (d.w)	1%	Tundra (organic horizon) (n=2)	Alaska	Olson et al., 2018
	0.1 and 1.0 ng/g (d.w)	3%	Tundra (B-horizon) (n=2)	Alaska	Olson et al., 2018
	0.17±0.55 ng/g (d.w)	0.48±1.7%	Peat plateaus underlain by permafrost (upper 50 cm, n=4)	Subarctic Sweden	Tarbier et al., 2021
	1.8±4.7 ng/g (d.w)	6.1±6.5%	Wetlands forming due to permafrost thaw (upper 50 cm, n=5)	Subarctic Sweden	Tarbier et al., 2021
	0.78±1.2 ng/g (d.w)	1.7±2.4%	Wetlands not impacted by permafrost thaw (upper 50 cm, n=5)	Subarctic Sweden	Tarbier et al., 2021
	0.11-6.71 ng/g (d.w)	0.3-17%	Wetland soils	Canada	Varty et al., 2021
	0.08±0.01 ng/g (d.w)	0.1%	Organic layer	Alaska	Yang et al., 2016
	0.07±0.01 ng/g (d.w)	0.1%	Mineral layer	Alaska	Yang et al., 2016
Soil pore water/wetland and stream water	0.43 ng/L	11	Poor fens (recent permafrost thaw)		Gordon et al., 2016
	0.01–0.05 ng/L		Palsas underlain by permafrost	Subarctic Sweden	Fahnestock et al., 2019
	1.0–2.0 ng/L		Wetlands formed due to permafrost thaw	Subarctic Sweden	Fahnestock et al., 2019
	0.23 and 0.51 ng/L	Up to 20%	Wetlands	Canada	Varty et al.,2021
	$0.216 \pm 0.020 \text{ ng/L to}$	0.47-9.9%	Downstream permafrost thaw slump	Canada	St Pierre et al., 2018
	3.45 ± 1.45 ng/L				
Rivers	0.04-0.18 ng/L				Table S2
Lakes	0.02-0.65 ng/L				Table S2
Ponds	<0.02-0.53 ng/L				Table S2

Glaciar ice and firn	0.1-0.7 ng/L	1-17% (median 5.5%)	Firn and ice Core Samples	Canada	Zdanowicz et al., 2016
Glacial meltwater	<0.010-0.235 ng/L		Unfiltered		Table S3
	<0.010-0.075 ng/L		Filtered		Table S3
Lichen	1.41–17.1 ng/g		High Arctic coast	Canada	St. Pierre et al., 2015
Seawater	0.018±0.016 ng/L MMHg: 0.018±0.010 ng/L (n=162) DMHg: 0.008±0.006 ng/L (MMHg:DMHg ~ 1.38)	12 ± 6%		Western Arctic Ocean	Agather et al., 2019
	0.036 ±0.024 ng/L			Beaufort Sea	Wang et al., 2012b (only stations included for this review)
	0.048 ±0.028 ng/L	18±12%		Canadian Arctic	Wang et al., 2018 (only stations included for this review)
	0.045±0.017 ng/L			Western Arctic Ocean	Heimbüger et al., 2015
	0.034±0.022 ng/L	30-33%		Eastern Arctic Ocean	Heimbüger et al., 2015
	0.027±0.019 ng/L MMHg 0.018±0.016 ng/L DMHg 0.006±0.007 ng/L (MMHg:DMHg ~ 1:2)	41±38%		Western and eastern Arctic Ocean	Jonsson et al., 2022
	0.013±0.004 ng/L			Barents Sea Opening/shelf	Petrova et al., 2020
	0.049±0.03 ng/L MMHg 0.019±0.013 ng/L DMHg 0.039±0.028 ng/L (MMHg:DMHg ~ 2:1)	24±14%		Fram Strait	Petrova et al., 2020
	0.075±0.019 ng/L			Fram Strait outflow (East Greenland Current)	Petrova et al., 2020
	0.020±0.012 ng/L (n=122)	9%	0-20 m	Arctic Ocean (this review)	

	0.040±0.026 ng/L (n=439)	21%	20-200 m		Wang et al., 2012, 2018;
	0.046±0.028 ng/L (n=177)	32%	200-500 m		Heimbürger et al., 2015;
	0.042±0.021 ng/L (n=288)	30%	>500 m		Agather et al., 2019; Charette et al., 2020; Kim et
					al., 2020b; Petrova et al., 2020, Jonsson et al., 2022; Kohler et al., (2022)
					Korner et al., (2022)
Shelf sediment	0.15-0.37 ng/g	0.43±0.17%		Beaufort Sea and Chukchi Sea	Fox et al., 2014
	0.2 ng/g			Western and eastern Arctic Ocean (data synthesis)	Soerensen et al., 2016b
	0.019-2.4 ng/g	0.06%-2.6%	High MeHg observed close to the Lena river delta	Laptev Sea	Liem-Nguyen et al., 2022
	0.002-0.168 ng/g	0.003%– 0.27%		Eastern Siberian Sea	Kim et al., (2020)
Basin sediment	0.05 ng/g			Western Arctic Ocean	Soerensen et al., 2016b
Plankton (freshwater)	0.03 ng/g		Phytoplankton	Lake Melville, Canada	Schartup et al., 2015a
	<0.4-30 ng/g		Water column seston		Chételat et al., 2018
	10-99 ng/g		Zooplankton	Skeleton Lake, Canada	Varty et al. (2021)
Plankton (marine)	<1.5 ng/g		Phytoplankton	Chukchi Sea	Fox et al., 2017
	8–96 ng/g			Circumpolar Arctic	Foster et al., 2012; Pućko et al., 2014; Pomerleau et al., 2016
Biofilms/microbial mats/Periphyton	2–12 ng/g		Shoreline rock biofilm, lakes	Lakes in East Canadian Arctic	Chételat et al., 2018
	0.3–1.3 ng/g		Microbial mats, lakes	Lakes in East Canadian Arctic	Camacho et al., 2015
	2–56 ng/g		Periphyton, filamentous algae	Lakes in East Canadian Arctic	van der Velden et al., 2013a
Snow	0.02-0.07 ng/L	0.2-10%	Surface snow	data synthesis	Soerensen et al., 2016b
	<0.01–0.074 ng/L	21-31%	Snowpack	data synthesis	Soerensen et al., 2016b

	0.14-9.1 ng/L	1.1-0.74%	Snow (unfiltered)	Lake Hazen, Canada	St. Pierre et al., 2019
	0.039-1.2 ng/L	13-42%	Snow (filtered)		
	0.002±0.0002 ng/L	0.41±49%	Surface snow from sea ice	Arctic Ocean	DiMento et al., 2019
	0.19-0.67 ng/L	8-11%		Canada (High Arctic)	Varty et al.,2021
	0.005-0.008 ng/L	1.5-5.1%	Surface snow from sea ice (unfiltered)	Arctic Ocean	Jonsson et al., 2022
Sea ice	0–0.5 ng/L	0 - 40%	multi-year	Eastern Beaufort Sea and	Beattie et al., 2014
				McClure Strait	
	0.017±0.024 ng/L	3.5 ± 4.3%		Central Arctic Ocean	DiMento et al., 2019
	0.006±0.004 ng/L	1.0-2.3%	first-year	Central Arctic Ocean	Schartup et al., 2020
	0.022±0.015 ng/L	2.2-5.0%	multi-year	Central Arctic Ocean	Schartup et al., 2020
	0.014±0.011 ng/L	8.2%		Central Arctic Ocean	Jonsson et al., 2022
		(media)			
Air	MMHg: $2.9 \pm 3.6 \text{ pg/m}^3$	range: <0.5 (L	.OQ) to 12.7 pg/m3	Hudson Bay, Canadian	Baya et al. 2015
	DMHg: $3.8 \pm 3.1 \text{ pg/m}^3$	range: <0.5 (L	LOQ) to 11.3 pg/m3	Arctic Archipelago	

Table 2. Total MeHg fluxes and pools in the Arctic environment (excluding aquatic biota larger than plankton). While the MeHg concentrations reported (Table 1) cover both freshwater and the marine compartments, the fluxes focus on those relevant for the Arctic Ocean. Details on how the MeHg pools (and their uncertainty) were calculated is provided in Table S1.

Environmental	MeHg Fluxes	References
Compartment		
Atmospheric deposition	8.5 (8-10) Mg/yr	Soerensen et al., 2016b
	2.6 Mg/yr	this review
Atmospheric evasion	14 Mg/yr	Soerensen et al., 2016b
	4.5 (2.8-6.3) Mg/yr	this review
Sea ice uptake	0.4 Mg/yr	Petrova et al., 2020
Sea ice discharge	0.5 Mg/yr	Soerensen et al., 2016b
Coastal erosion	0.07-0.46 Mg/yr	Soerensen et al., 2016b
	1.5 Mg/yr	Petrova et al., 2020
Wetland ponds	1.8–40 ng/m², day	Lehnherr et al., 2012a
Glacial meltwater	0.03-14.9 g/y	St. Pierre et al., 2019
(unfiltered), Ellesmere Island	0.49-11.7 mg/km ²	
Riverine inputs	0,8 Mg/yr	Sonke et al. 2018
Inflow to Arctic Ocean	3-13 Mg/yr	Soerensen et al., 2016b
—of Atlantic deep and mid- depth waters	6.8 Mg/yr	Soerensen et al., 2016b
Inflow to Arctic Ocean	9±1.5 Mg/yr	Petrova et al., 2020
—from Fram Strait	7.5±2 Mg/yr	Petrova et al., 2020
—from Barents Sea Opening	1 Mg/yr	Petrova et al., 2020
—from Bering Strait	0.5 Mg/yr	Petrova et al., 2020
Outflow to North Atlantic	8–27 Mg/yr	Soerensen et al., 2016b
Outflow to North Atlantic	16.5±2 Mg/yr	Petrova et al., 2020
—via Fram Strait	14±3 Mg/yr	Petrova et al., 2020
—via Davis Strait	2.5±1 Mg/yr	Petrova et al., 2020
Sedimentation load (shelf sediment)	1.3 Mg/yr	Soerensen et al., 2016b
	0.3 Mg/yr	Petrova et al., 2020
Sedimentation load (basin sediment)	0.1 Mg/yr	Soerensen et al., 2016b
Sediment diffusion to water (shelf sediment)	0.9 Mg/yr	Soerensen et al., 2016b
	4.6±0.2 Mg/yr	Kim et al., 2020b
Sediment diffusion to water (basin sediment)	0.1 Mg/yr	Soerensen et al., 2016b

Environmental Compartment	MeHg Pool	References
Atmosphere	1.6 (0–3.2) Mg	Baya et al., 2015; Table 3.8.1 in AMAP (2021)
Surface soils	510 (270–749) Mg	St. Pierre et al., 2015; Tarbier et al., 2021
Snowpack	2.4 (0.43-4.3) Mg	St. Pierre et al., 2019, Varty et al., 2021; Jonsson et al., 2022

Glaciers	133 (±1.2) Mg	Zdanowicz et al., 2016
Sea ice	0.32 (0.04-1.0) Mg	Dimento et al., 2019;
		Schartup et al., 2020
Shelf sediments	0.002-2.4 ng/g	Fox et al., 2014; Soerensen
		et al., 2016b; Liem-Nguyen
		et al., 2022; Kim et al.,
		(2020)
Deep ocean sediments	0.05 ng/g	Soerensen et al., 2016b
Water column		Wang et al., 2012, 2018;
0–20 m depth	4 (±2) Mg	Heimbürger et al., 2015; Agather et al., 2019;
20-200 m depth	52 (±34) Mg	
200-500 m depth	73 (±44) Mg	Charette et al., 2020; Kim et
>500 m depth-bottom)	461 (±232) Mg	al., 2020b; Petrova et al.,
,	. (- , 5	2020, Jonsson et al., 2022;
		Kohler et al., 2022
Plankton	0.7 Mg	Soerensen et al., 2016b

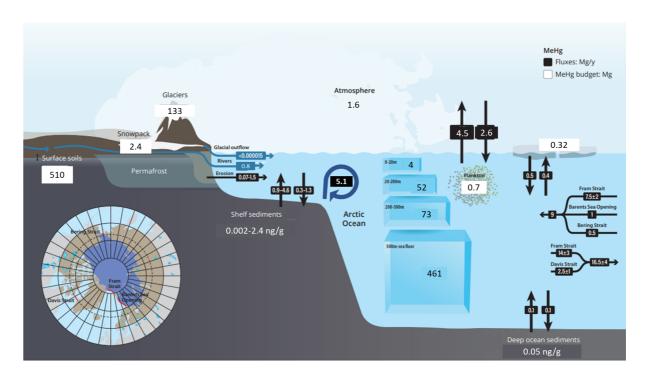


Figure 2. MeHg fluxes and budget in the Arctic environment. Updated from Soerensen et al. (2016b) and Petrova et al. (2020) with data shown in Table 1 and Table 2.

3. Mechanistic understanding of Hg (de)methylation processes

3.1. Formation and degradation pathways of MeHg

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Given that MeHg is the bioaccumulative form of Hg, understanding the rates of both methylation and demethylation as well as Hg(II) reduction, which might limit the availability of this species for methylation, have implications for biotic uptake and exposure of Hg. While photochemical processes might govern Hg chemistry at the surface of the waters, biological processes are the key to understand Hg cycling in freshwater, seawater and in sediments. Methylation of Hg is primarily thought to be a biotic process carried out by various microorganisms, though evidence for abiotic methylation exists. In 2013, the gene pair involved in biological Hg methylation (hgcAB genes) were discovered (Parks et al., 2013). This opened up new possibilities to search for Hg methylators (Gilmour et al., 2013, Jones et al., 2019; Capo et al., 2020; Gionfriddo et al., 2016; McDaniel et al., 2020) and uncover the intracellular processes involved in Hg methylation (Smith et al., 2015). Carriers of the *hgcAB* genes that are known to methylate Hg include a wide range of sulfate reducing bacteria, iron reducing bacteria, methanogens, and some fermentative microorganisms (Gilmour et al., 2013, 2018). A global analysis of metagenomes was first to reveal the presence of hgcAB in the Arctic environment by identifying the genes in Deltaproteobacteria, Firmicutes, and particularly in methanogenic archaea from thawing permafrost soil (Podar et al., 2015). Fahnestock et al. (2019) also showed that concentrations of MeHg in porewater increased across a thaw gradient along with the abundance and diversity of potential Hg methylators. It is important to note that not all Hg methylating microorganisms have the same capacity to methylate Hg and that this appears to depend on individual strains rather than taxonomic ranking (Ma et al., 2019). Marine microorganisms with a greater oxygen tolerance and broader habitat range might be important Hg methylators in the seawater (Villar et al. 2020; Lin et al., 2021). In the last decade, several studies have highlighted the importance of Hg methylation in suboxic and even oxic waters for MeHg bioaccumulation and biomagnification in aquatic food webs (Sunderland et al., 2009; Heimbürger et al. 2010; Lehnher et al., 2011; Mason et al., 2012; Wang et al., 2012b, 2018; Blum et al., 2013; Heimbürger et al., 2015; Schartup et al., 2015a; Soerensen et al., 2016a; Gascón Díez et al., 2016; Rosati et al. 2018; Agather et al., 2019; Petrova et al., 2020). While

anaerobic biological Hg methylation is mediated by the hgcAB genes that are abundant in marine sediments (Podar et al., 2015), the aerobic pathway, if it exists, remains unclear. HgcA or hgcAB-like genes have recently been identified in global ocean oxic seawater (Villar et al., 2020) as well as in Antarctic sea ice (Gionfriddo et al. 2020), though not in the Arctic Ocean (Bowman et al., 2020a). Further studies are needed to identify key players for Hg methylation in polar marine waters and sea ice. It is generally agreed that primarily dissolved complexes of Hg(II) and MMHg are taken up by Hgmethylating bacteria at the base of aquatic food webs (Drott et al., 2007; Schaefer and Morel, 2009; Chiasson-Gould et al., 2014; Mangal et al., 2019). Adsorption of Hg(II) and MMHg to the solid phase are thus important processes limiting their biological uptake (Schartup et al., 2015a; Mazrui et al., 2016). However, as shown in previous work from other marine or brackish water systems, different dissolved and solid chemical forms of Hg differ in their bioavailability (Jonsson et al., 2012, 2014; Mazrui et al., 2016). Although recent terrestrial inputs of Hg(II) and MeHg did not contribute much to the stock of MeHg in the estuarine system, 40 to 70% of the bioaccumulated Hg pool originated from these recent terrestrial Hg inputs. In these studies, refractory pools of Hg(II) and MMHg were not fully accounted for. MeHg demethylation can be biotically or abiotically mediated. Abiotic demethylation is mostly photochemical degradation in surface waters by different mechanisms including direct photolysis and indirect photolysis via reactions with reactive oxygen species such as hydroxyl radicals. Research in the Arctic, subarctic, and temperate areas has shown photochemical demethylation from ultraviolet (UV) radiation is an important aspect of freshwater Hg budgets (Seller et al., 1996; Lehnherr et al., 2012a; Poste et al., 2015; Klapstein and O'Driscoll, 2018). Because organic matter compounds in surface waters readily absorb UV radiation, they mediate the photodemethylation of aqueous MeHg (Klapstein et al., 2017). A recent study on ponds with high DOC concentrations resulting from permafrost thaw slumps found higher rates of photodemethylation when compared to an oligotrophic lake in the polar desert with low DOC concentrations (Girard et al., 2016). Photolytically produced reactive oxygen species from organic matter have been shown to mediate MeHg demethylation in

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Alaskan lacustrine systems (Hammerschmidt and Fitzgerald, 2010). In the Arctic, photodemethylation rates are likely regulated by prolonged ice coverage (Poste et al., 2015; Klapstein and O'Driscoll 2018). Non-photochemical abiotic demethylation may also occur via reactions with selenoamino acids (Khan and Wang, 2010), but the importance of such reactions remains to be further examined outside of the laboratory.

Bacterial resistance to MeHg (and Hg(II)) is one of the most widely observed phenotypes in eubacteria (Barkay et al., 2003) and is mediated by the *mer* operon, which has been found in a diverse range of habitats including soils, sediment and water, both freshwater and marine systems (Barkay et al., 2003 and references therein). Within the operon (which contains different types of genes), attention has been focused on the *merA* gene, which catalyzes reaction transforming Hg(II) into Hg(0) and *merB*, which codifies for the alkaline mercury lyase, a protein that demethylates MeHg into Hg(II). Little is known about the occurrence, diversity and distribution of *these genes* in the Arctic region. Of potential importance for Hg demethylation, a recent study, however, unveiled the presence of *merB* in deep waters of the Arctic Ocean (Bowman et al., 2020a). Also, a study performed in lake sediment cores showed the presence of hgcA in Arctic lakes (Ruuskanen et al., 2020). The diversity of both Hg methylators and MeHg demethylators in the Arctic needs to be further explored.

3.2. Formation and degradation pathways of DMHg

The main formation and degradation pathways of the DMHg remain unknown in marine waters, including the Arctic Ocean (Agather et al., 2019; Petrova et al., 2020; Bowman et al., 2020b, Jonsson et al. 2022). Although early studies identified synergies between the activity of organisms and production of DMHg (Baldi et al.,1993; Sommar et al., 1999; Pongratz and Heumann, 1998), there is currently no experimental data to support that it is a biological product. Baldi et al. (1993) suggested the formation of DMHg in bacterial cultures was the result of an abiotic reaction between dissolved sulfide and MeHg (via the formation of bismethylmercurysulfide). Recent work has also shown abiotic formation of DMHg from MMHg when MMHg is adsorbed onto organic or inorganic surfaces

containing reduced sulfur sites (Jonsson et al., 2016). Other known pathways of DMHg formation of potential relevance in natural environments include the reaction of MeHg with selenoamino acids (Khan and Wang, 2010) and methylation with methylcobalamin (Imura et al., 1971). Volatilization of DMHg from the marine boundary layer and the subsequent atmospheric demethylation of DMHg to MMHg has been suggested as a source of MeHg in surface snow (Kirk et al., 2008; St. Louis et al., 2007), marine waters (Baya et al., 2015) and coastal terrestrial environments (St. Pierre et al., 2015). Baya et al. (2015), who were first to quantify DMHg in ambient air, also suggested the evasion of DMHg from marine waters in Hudson Bay and the Canadian Arctic Archipelago was enough to sustain the atmospheric concentrations of MMHg observed. In addition to degradation of DMHg in air, DMHg could potentially also be degraded in marine waters, although there have been contradictory claims in the literature on whether the rates are high enough for the process to play a role in the cycling of DMHg (Mason and Sullivan, 1999; Black et al., 2009). West et al. (2022), however, recently demonstrated that DMHg in arctic seawater readily degrades under natural sunlight. The photodecomposition rates observed in tested waters were in the same range as the rate observed for MMHg, suggesting that a significant fraction of the DMHg that is transported to surface waters may be degraded to MMHg directly in surface waters rather than in the atmosphere. It is also possible that DMHg is degraded through other pathways which do not involve photochemical reactions. DMHg has recently been shown to degrade in aqueous solutions to MeHg in the presence of reduced sulfur as dissolved sulfide or as the metastable iron sulfide mackinawite (West et al., 2020).

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4. Where are methylmercury hotspots in the Arctic?

4.1. Terrestrial Environments

4.1.1 Arctic Tundra

Mercury has a strong affinity to organic substances, and organic soils therefore play an important role
in the mobilization and transportation of Hg from soils to water bodies. As such, surface vegetation,
soils, and organic matter are likely locations where elevated Hg is present and available for
transformation. St. Pierre et al. (2015) reported that soil MeHg concentrations varied by two orders of
magnitude (0.02–2.11 ng/g) across several sites in the Canadian Arctic Archipelago, with a median of
0.12 ng/g. A tundra study (Olson et al., 2018) reported concentrations of MeHg between 0.5 and 3.4
ng/g in organic horizons and 0.1 to 1.0 ng/g in B horizons, with MeHg on average accounting for 1%
of THg (THg with a maximum of 3% in the lower B horizon). A study from Fennoscandia further
found MeHg accounting for $0.48 \pm 1.7\%$ of THg in peat plateau soils, respectively (Tarbier et al.
2021). As discussed further below, higher %MeHg values has been reported from subarctic and Arctic
soils saturated with water (Tarbier et al. 2021: $6.1 \pm 6.5\%$, Varty et al. 2021: $>10\%$).
In addition to driving the deposition of Hg to Arctic tundra by foliar uptake (Obrist et al., 2017), plants
also bioaccumulate Hg from the soil via the roots (Zhou and Obrist, 2021). Generally, very few data
are available on MeHg concentrations in Arctic vegetation. Lichens in some coastal areas of the
Canadian High Arctic (MeHg: 1.41-17.1 ng/g) were more enriched in MeHg (and THg) relative to
underlying soils than at inland sites. This suggests greater uptake/deposition of MeHg may originate
from the evasion of DMHg and MMHg from open marine waters in coastal areas (St. Pierre et al.,
2015). Further research is needed to identify sources and pathways of MeHg uptake in terrestrial
vegetation.

A major consequence of climate warming in the Arctic and Subarctic has been the thawing and degradation of permafrost (Jorgenson et al., 2001, 2006; Lantz and Kokelj, 2008; Walvoord et al., 2019). Given the large amounts of Hg that potentially could be released from Arctic soils in a warmer climate (Schaefer et al., 2020, Lim et al. 2020), methylation rates in the Arctic terrestrial environment have gained more attention. Permafrost degrades through two dominant mechanisms; first is the top down thaw of near-surface permafrost through the downward expansion of the seasonally thawed active layer (Douglas et al., 2021). This mechanism, which typically encompasses a broad areal extent but is not vertically extensive, generates wetter and more anoxic conditions that are potentially favorable for Hg methylating bacteria (Yang et al., 2016). While increased Hg methylation under such conditions has been demonstrated in temperate and tropical zones (Hsu-Kim et al., 2018), few studies have covered Hg methylation in zones with thawing permafrost (Fahnestock et al., 2019; Tarbier et al., 2021; Varty et al., 2021). The second mechanism for permafrost degradation is the development of thermokarst features through thaw subsidence of ice-rich frozen ground (Vonk et al., 2015; Olefeldt et al., 2016; Olefeldt et al., 2021). This typically leads to creation of a variety of features including scarps, pits, thaw slumps, and ponding of surface water that do not encompass large regions but are hotspots for geomorphologic processes. Thermokarst wetlands, ponds and lakes have been repeatedly identified as Hg methylation hotspots across the Arctic due to the concurrent mobilization of carbon and nutrients, creating conditions suitable for methylation. In previously formed thaw ponds, Lehnherr et al. (2012a, 2012b) measured potential Hg methylation rates (1.8 to 40 ng/m²/day) similar to rates in already identified Hg methylation hotspots. High concentrations of dissolved MeHg (ranging up to 22 ng L⁻¹), likely from either increased export of MeHg from peat soils, or in situ methylation in pond sediments, have also been found in eastern Canadian thaw ponds (MacMillan et al., 2015). Gordon et al. (2016) observed up to an order of magnitude higher MeHg concentrations in the poor fens associated with recent permafrost thaw (0.43 ng/L, filtered) than in the hydrologically isolated bogs (0.04 to 0.1 ng/L, filtered) in the Scotty Creek catchment of the discontinuous permafrost zone (in Canada's Northwest Territories). In the Stordalen Mire in Subarctic Sweden, MeHg concentrations increased by two orders

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of magnitude along the permafrost thaw gradient, from 0.01–0.05 ng/L in palsas underlain by permafrost to 1.0-2.0 ng/L in thawed fens (Fahnestock et al., 2019). The thawed fens were associated with heightened Hg accumulation, reducing conditions, higher sulfur concentrations and a greater proportional representation of putative Hg methylators (i.e. microbial strains carrying the Hg methylation gene pair but have not yet have been confirmed to methylate Hg in laboratory studies) within the microbial communities (from 0.5% to 7.8%), compared to the dry frozen palsas and bogs at intermediate stages of thaw (Fahnestock et al., 2019). In another Fennoscandian permafrost peatland study, an order of magnitude higher concentration of MeHg (and fraction of THg as MeHg) was found in topsoil (upper 50 cm) from collapse fens (edge of the fen receiving peat from the collapsing peat plateau; 1.8 ± 4.7 ng g⁻¹; $6.1 \pm 6.5\%$ of THg) compared to the peat plateau soils $(0.17 \pm 0.55$ ng g⁻¹; $0.48 \pm 1.7\%$ of THg; Tarbier et al. 2021). Median MeHg and %MeHg in the soil from the collapse fen $(1.8 \pm 4.7 \text{ ng g}^{-1}; 1.7 \pm 2.4\% \text{ of THg})$ was also 3.6 times higher than in the nearby fens assumed to have no recent history of permafrost, suggesting higher methylation of Hg in the recently formed fens. In line with these studies, Varty et al. (2021) observed higher %MeHg in soils saturated with water in comparison to drier soils collected on Ellesmere Island (Canada). Permafrost thaw has also been shown to influence microbial processes (Mackelprang et al., 2011) and biogeochemical cycling (Frey and McClelland, 2009) in downstream freshwater systems. Several recent studies have reported sites with permafrost slumps (i.e. slope failures or mass-wasting features resulting from the thawing of ice-rich permafrost) are abundant with both organic matter compounds and Hg. In fact, the highest concentrations of THg and MeHg (1200 and 7 ng/L unfiltered, respectively) ever recorded in natural flowing freshwater were measured downstream of a permafrost thaw slump in the Northwest Territories of Canada (St. Pierre et al., 2018). Other studies show similarly elevated levels of both THg and MeHg in water and sediments from areas influenced by permafrost degradation (MacMillan et al., 2015; Gordon et al., 2016; Houben et al., 2016). In stratified thaw ponds, elevated MeHg concentrations and percentages in bottom waters suggest in situ methylation is occurring in anoxic waters (MacMillan et al. 2015). In addition to slump age, many factors may be contributing to the reported differences in organic matter and Hg levels downstream of

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permafrost slumps, highlighting the need for further study of how permafrost degradation will affect interactions between organic matter and Hg across various Arctic and Subarctic landscapes. Factors such as the depth of the permafrost active layer, peatland buffering and interferences as well as ground water influences may affect organic matter-Hg dynamics downstream of thermokarst processes (Frey and McClelland, 2009; Lim et al., 2019). Furthermore, thaw slumps are likely affecting a wide range of other biogeochemical processes that will either directly or indirectly alter Hg methylation or bioaccumulation in receiving water bodies (Pokrovsky et al., 2018).

4.1.2. Freshwater systems

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Arctic freshwater systems support several species of fish, which can be important in the traditional diet of some Indigenous peoples (Lockhart et al, 2005). Mercury enters lakes and streams primarily via terrestrial runoff and atmospheric deposition, but it can also originate from groundwater passing through unfrozen parts of permafrost below the water body (i.e. taliks). Inorganic Hg(II) entering Arctic freshwater systems can either be photolytically reduced to Hg(0) and evade back to the atmosphere, be methylated to MeHg or bind to particles and sink to the bottom sediments. Where methylation of Hg dominates in freshwater systems has been shown to vary between different Arctic freshwater bodies (see Table S2; Chételat et al., 2015). Mercury and MeHg dynamics in freshwater systems change with the seasons, having low or no mobilization during freezing conditions (November-May) which increases with higher river flows and flooding during the melting season from May to June. Methylmercury in terrestrial snow, for example, enters High Arctic lakes during snowmelt via melt streams, having higher concentrations during spring melt than during summer (Loseto et al., 2004). Methylmercury concentrations in Canadian Arctic freshwater systems are generally low ranging from 0.04 to 0.18 ng/L in unfiltered samples in rivers, 0.02 to 0.10 ng/L in lakes and greater than 0.02 to 0.53 ng/L in ponds (see Table 1). Compared to deeper lakes, shallow water bodies generally have higher MeHg due to warmer temperatures and higher organic matter available for microbial metabolism and are probable methylation hotspots (Lehnherr et al., 2012a; Lehnherr, 2014; MacMillan et al., 2015). Varty et al. (2021) showed that small, shallow Arctic lakes can reach higher concentrations of water column MeHg (up to 0.72 ng/L)

in winter when ice covered. In addition, snow MeHg concentrations in spring (0.19-0.67 ng/L) suggest that snowpack may be an important MeHg source during the melt period. Water residence time may also play a role, with less within-lake processing of aqueous MeHg in smaller lakes (Richardson et al., 2021). In Alaskan lakes, sediment methylation has been shown to be the major source of MeHg to those ecosystems (Hammerschmidt and Fitzgerald, 2006). Varty et. al. (2021) also suggested sediment MeHg production is important for the annual MeHg budget of Skeleton Lake and can explain the elevated concentrations of MeHg observed during the ice-covered season. Generally, in freshwater ecosystems, the main carrier for MeHg is DOC (Tsui and Finlay, 2011), which has the double role of promoting the solubility and transport of MeHg in surface waters and attenuating dissolved MeHg uptake by aquatic food webs by limiting transport across cell membranes. High Arctic lakes typically have low biological productivity, limiting Hg uptake by algae and suspended detrital organic matter. At the same time, many Arctic lakes have low DOC concentrations, often below the suggested threshold of 8 to 11 mg/L DOC, above which DOC is hypothesized in some studies to counteract MeHg bioaccumulation (French et al., 2014; Braaten et al., 2018). For MeHg production, methylation of Hg in High Arctic wetland ponds was suggested to be negatively correlated with DOC levels by decreasing Hg(II) availability to methylating microorganisms (Lehnherr et al., 2012a). While some studies suggest DOC concentrations may be declining in Arctic lakes (Saros et al., 2015), others show increases in DOC because of climate change, land use practices and other water chemistry interactions (Macdonald et al., 2005; Isles et al., 2020). During melting, most transported Hg is bound to particles. A total flux of (dissolved) MeHg flux from all Arctic rivers was estimated by Sonke et al. (2018) to be 584 kg/y. The larger variability among different flux estimates for rivers, e.g. the Mackenzie River in various studies (72 kg/yr of dissolved MeHg (Sonke et al. 2018), 14 kg/yr MeHg (Mu et al. 2019) and 13 and 15 kg/y MeHg (unfiltered) (Emmerton et al, 2013)) stresses the need for additional studies to better constrain these fluxes. The total Hg flux from larger Arctic rivers estimated by Sonke et al. (2018) was 27.6 Mg/yr (13.4 Mg/yr as dissolved Hg and 13.9 as particulate Hg). The fate of Hg exported to the Arctic Ocean is less clear. Coastal sediments and oligotrophic lake sediments in a High Arctic region of the

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Canadian Arctic Archipelago were found to methylate Hg, though at a low rate (St. Pierre et al., 2014;

434 Hudelson et al., 2020).

4.1.3 Glaciers

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Methylmercury has been shown to be released in glacier-fed streams of the Canadian High Arctic (Ellesmere Island; St. Louis et al., 2005; St. Pierre et al., 2019), Yukon (Halm and Dornblaser, 2007) and Alaska (Nagorski et al., 2014; Vermilyea et al., 2017). Reported MeHg concentrations in unfiltered glacial meltwater vary from <0.010 to 0.26 ng/L and account for 1 to 26% of THg, while concentrations of MeHg in filtered meltwater vary from <0.010 to 0.075 ng/L and account for 4 to 32% of THg (see Table S3). Estimates of unfiltered MeHg fluxes in glacier-fed streams of Ellesmere Island are between 0.03 and 14.9 g/y, which translate to annual yields ranging from 0.49 to 11.7 mg/km²/y in basins with 21 to 82% glacier coverage (St. Pierre et al., 2019; see Table S3). An increase in the glacial sediment Hg flux in response to enhanced glacier melt rates might lead to higher exposure risks to inorganic Hg or MeHg in downstream aquatic ecosystems in some catchments. On land, changes might occur with the development of seasonally wet areas in newly deglaciated forelands where methylation of sediment-bound Hg could occur. In southeastern Alaska, for example, watersheds with greater wetland abundance have higher Hg concentrations, in both water and aquatic fauna, than glacierized basins (Nagorski et al., 2014). An increased supply of labile and bioavailable DOM from terrestrial sources such as wetlands or newly formed organic soils (e.g., Wietrzyk-Pełka et al., 2020) might also increase the risk of riverine Hg uptake by biota in these watersheds. It could also promote Hg methylation by sulfur-reducing bacteria in catchments where subglacial oxidation of sulfides causes enhanced sulfidization of ponds, wetlands and lakes (Graham et al., 2012). Cryoconite holes (dirt-filled melt cavities on glacier surfaces) are abundant over glacier margins or stagnant ice and can accumulate atmospherically deposited impurities released by ice melt, as well as microbes (Cook et al., 2016). These micro-environments could potentially act as local hotspots of Hg methylation in glacial environments, but direct evidence supporting this is presently limited to Tibetan glaciers (Zhang et al., 2020).

In the marine environment, meltwater plumes from tidewater glaciers carrying both dissolved and particulate Hg may be injected into zones of locally high primary productivity in coastal surface waters, where subsequent methylation and biological uptake can occur (Schartup et al., 2015a). Some glaciers of southern Alaska are sources of old (up to ~3900 y) and highly labile DOM, of possible subglacial bacterial origin, which is discharged into coastal waters (Hood et al., 2009, 2015). Likewise, retreating glaciers in fjords in southern Spitsbergen, Svalbard, discharge sediments containing up to 2.8% weight of OC (Kim et al., 2020a). Seafloor sediments in these fjords have THg concentrations of up to 84 ng/g. Hawking et al. (2021) recently reported dissolved MeHg concentrations from Greenland glacial catchments that were similar to values previously reported from wetlands. As discussed elsewhere (Dastoor et al., 2022) for THg, further studies are warranted to confirm the high levels of MeHg in these glacial catchments, particularly since the source(s) and processes leading to high MeHg yield remain unknown. Hg concentrations correlate closely with OC and are thought to be largely derived from glacially eroded sedimentary rocks (e.g., shales) in the glacierized basin. The discharge of Hg-bearing glacial sediments containing organic matter might promote Hg methylation and uptake by marine biota within the water column, which has not been well studied. However, as a recent and exhaustive review highlighted, the local effect of glacial runoff on coastal marine primary productivity is likely limited to a few kilometers beyond glacier termini, and whether runoff promotes or limits productivity varies between land- or marine-terminating glaciers (Hopwood et al., 2020). For example, observations in fjords in northwestern Svalbard suggest that glacial meltwater plumes could reduce the phytoplankton biomass (Piquet et al., 2014), and this might effectively limit MeHg production and uptake in the marine food chain. Furthermore, the high deposition rates of inorganic matter near tidewater glaciers, or at the mouth of proglacial channels, dilute the organic matter content in the seafloor sediments of Svalbard fjords (Liu et al., 2015), and this might limit benthic Hg methylation. On the whole, therefore, the environmental impact of Hg releases from melting glaciers and ice caps is likely to be local, rather than widespread across the Arctic. A study on the East Greenland shelf found no enhanced Hg or MeHg levels near the 79°N glacier, precluding large inputs or locally enhanced MeHg production (Petrova et al. 2020).

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4.2. Marine systems

4.2.1 Seawater

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488	MeHg concentrations in the western Arctic Ocean have been reported to be 0.011 ± 0.010 ng/L in
489	filtered seawater (Agather et al., 2019, Figure 3) and 0.046±0.024 ng/L in unfiltered seawater (Wang
490	et al., 2018). Heimbürger et al. (2015) found similar levels of 0.045 ± 0.017 ng/L and 0.034 ± 0.022 ng/L and 0.034 ± 0.022 ng/L and 0.034 ± 0.022 ng/L
491	in unfiltered seawater in the Eastern Arctic and Central Ocean. Methylmercury (MeHg) inflow to the
492	Arctic Ocean totalled 9±2 Mg/yr, most of which enters with deep and mid-depth waters of the Atlantic
493	Ocean (7.5±2 Mg/yr; Petrova et al., 2020). Similar to THg, more MeHg flows out of the Arctic Ocean
494	into the North Atlantic (16.5±3 Mg/yr, 14±2 Mg/yr via the Fram Strait, 2.5±1 Mg/yr via the Davis
495	Strait; Petrova et al. 2020) than flows in. The MeHg outflow via the Fram Strait conserves the shallow
496	MeHg peak found throughout the Arctic Ocean, and an average DMHg:MMHg ratio of 2.5:1 (Petrova
497	et al., 2020). In situ MeHg formation in the Arctic Ocean water column (Lehnherr et al., 2011; Wang
498	et al., 2012b, 2018; Heimbürger et al., 2015; Soerensen et al., 2016b; Petrova et al., 2020) results in a
499	net MeHg outflow of 7.5 ± 2 Mg/yr (Petrova et al. 2020).
500	A subsurface enrichment of MeHg in seawater has been observed in almost all major ocean basins
501	(AMAP/UNEP 2019), including the Atlantic Ocean (Bowman et al., 2015; Bratkič et al., 2016),
502	Pacific Ocean (Sunderland et al., 2009; Hammerschmidt and Bowman, 2012; Munson et al., 2015;
503	Bowman et al., 2016; Kim et al., 2017), Southern Ocean (Cossa et al., 2011; Gionfriddo et al., 2016);
504	Mediterranean Sea (Cossa et al., 2009, 2012, 2017; Heimbürger et al., 2010), Baltic Sea (Soerensen et
505	al., 2016b) and the Black Sea (Rosati et al., 2018) basins. However, the Arctic Ocean basin differs
506	from the others because the MeHg enrichment occurs at a much shallower depth (100-300 m depth)
507	where dissolved oxygen saturation is well above 75% (Wang et al., 2012b, 2018; Heimbürger et al.,
508	2015; see Figure 4). The Arctic subsurface enrichment of MMHg and DMHg is also detected in water
509	exiting the Arctic Ocean to the North Atlantic Ocean (Petrova et al. 2020). Heimbürger et al. (2015)
510	first hypothesized that the characteristic shallow MeHg peaks lead to enhanced biomagnification and
511	may be an alternative explanation for high arctic biota Hg levels, rather than anthropogenic Hg inputs.
512	As this subsurface MeHg maximum lies within the habitat of zooplankton and other lower trophic

level biota it has been suggested that biological uptake of the subsurface MeHg and subsequent biomagnification explain a long-standing mystery in the Canadian Arctic on why marine animals in the western Arctic have higher Hg levels than those in the east (see Figure 4; Wang et al., 2018). Increasing amounts of bioavailable Hg in surface waters could have an effect on Hg accumulation at the bottom of Arctic marine food webs. THg levels of up to 242 ng/g have been reported in zooplankton from Hudson Bay (Foster et al., 2012). As discussed further below, regional differences in food web length (and thus biomagnification) have also been provided as a contributing factor explaining the higher Hg levels marine animals in the western Arctic (St. Louis et al., 2011; Brown et al., 2016). The mechanisms of subsurface MeHg formation in the Arctic Ocean and other ocean basins in the world is under debate. Most of the studies have attributed subsurface MeHg to in situ Hg methylation in the water column associated with organic matter remineralization (Sunderland et al., 2009; Cossa et al., 2009; Wang et al., 2012b; Heimbürger et al., 2010, 2015), yet the biotic or abiotic processes that are responsible for methylation remain unidentified (Bowman et al., 2020a). Recent work in the Canadian Arctic (Wang et al., 2018) and the Siberian Shelf (Kim et al., 2020) suggest that advection of MeHg produced in shelf sediments also could contribute to the subsurface peak of MeHg found in the water column. This debate calls for further studies on MeHg sources and dynamics in the Arctic and global oceans.

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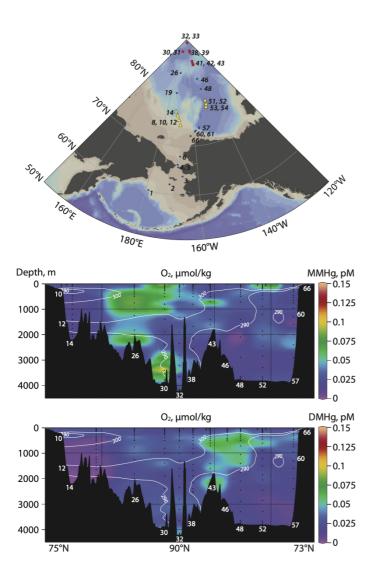
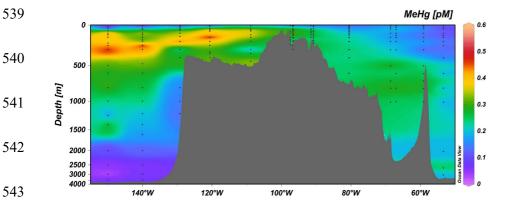


Figure 3. Monomethylmercury (MMHg) and dimethylmercury (DMHg) and dissolved oxygen (O_2) in filtered seawater along a looped latitudinal section (\sim 75°N to the North Pole and then back to \sim 73°N) in the western Arctic Ocean. The top panel shows sampling station locations along the section. Adapted from Agather et al., 2019.





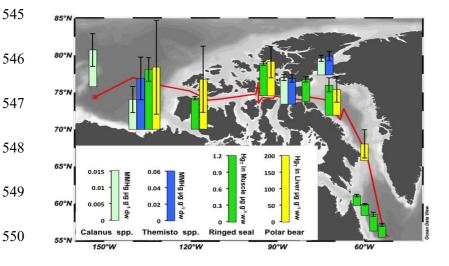


Figure 4. Methylmercury (MeHg) concentrations in unfiltered seawater (top panel) and the marine food web (bottom panel) along a longitudinal (west-to-east) section (the red line in the bottom panel) across the Canadian Arctic and Labrador Sea. Bar charts in the bottom panel show mean concentrations of monomethylmercury (MMHg) in *Calanus* spp. and *Themisto* spp. collected from 1998 to 2012 (±1SD), THg in muscle of adult ringed seals collected in 2007 and 2011 and THg in liver of polar bears (*Ursus maritimus*) collected from 2005 to 2008. Adapted from Wang et al., 2018.

4.2.2. Marine sediment

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Methylmercury concentration measurements in deep Arctic Ocean sediments are sparse. Based on a range of means, MeHg concentrations in Beaufort Sea and Chukchi Sea shelf sediments were estimated to range from 0.15 to 0.37 ng/g (dry weight, Fox et al., 2014). A recent study in the East Siberian Sea shows that MeHg concentrations range from 0.09 to 0.17 ng/g in shelf sediments, with much lower concentrations (~0.003 ng/g dry weight) in the slope sediments (Kim et al., 2020b). Liem-Nguyen et al (2022) reported MeHg concentrations ranging from 0.19 to 2.6 ng/g in Laptev Sea sediments along a 330 km long cross-shelf transect. The higher concentrations of MeHg (and MeHg/OC ratio) observed close to the Lena river delta than further out on the shelf was explained by trapping of terrestrial MeHg close to the river delta. Based on estimated MeHg concentrations of 0.2 ng/g (dry weight) for shelf sediments and 0.05 ng/g (dry weight) for the deep basin, Soerensen et al. (2016b) estimated a MeHg sedimentation flux of 1.3 Mg/yr and 0.1 Mg/yr to shelf and deep basin, respectively. Resuspension was not estimated due to an absence of data. Soerensen et al., (2016b) estimated MeHg diffusion to the overlying water column at 0.9 Mg/yr and 0.1 Mg/yr from shelf and deep basin sediments, respectively, for a total flux of 1.0 Mg/yr. However, a recent study in the East Siberian Sea suggests that the MeHg diffusion flux from sediments to the water column could be much larger, amounting to 4.6±0.2 Mg/yr for the Arctic shelf system alone (Kim et al., 2020b). There is a need to better constrain the benthic MeHg flux.

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4.2.3. Sea ice and snow

The sea ice environment, including first-year ice, multi-year ice and the overlying snow cover, provides a spatially and temporally dynamic interface between the ocean and the atmosphere in the Arctic Ocean where physical transport, chemical reactions, and biological uptake occur (Wang et al., 2017). While sea ice is semi-permeable (e.g., Loose et al. 2010), elevated Hg(0) concentrations under sea ice compared with ice-free waters suggest that ice is a barrier for gas exchange (Dimento et al. 2019). The sea ice environment is also an important habitat and feeding ground for many organisms,

including bacteria, algae, zooplankton and larger mammals such as polar bears and whales. Beattie et al. (2014) found elevated MeHg concentrations in multi-year Arctic sea ice of up to 0.5 ng/L (unfiltered) in the lower part of the ice (see Figure 5). Recent work has shown concentrations of both MMHg and DMHg in sea ice, especially within first-year ice (4.4% MMHg vs. 1.5% in multi-year ice) from across the central Arctic Ocean (Schartup et al., 2020). Snow, which covers large parts of the Arctic landscape as well as sea icescape has a large interstitial surface area for air-snow exchange of Hg and plays a major role in atmospheric mercury depletion events (Douglas et al., 2012, 2017; Steffen et al., 2014). However, Douglas et al. (2012) report that the small amounts of MeHg found in snowpack before spring snowmelt does not add significantly to MeHg levels in aquatic systems. As the high concentrations of MeHg in the bottom layer of sea ice (Beattie et al., 2014) cannot be explained by downward transport of MeHg from the overlying snow (Gionfriddo et al., 2016; Soerensen et al., 2016b), it has been postulated that Hg methylation could take place in sea ice. Discrete peaks in MMHg and DMHg concentrations within sea ice cores (Schartup et al., 2020) also suggest that they may be produced by microorganisms within the sea ice. This is supported by the fact that microbes present in snow and ice are metabolically active even under freezing temperatures (Douglas et al., 2012). Methylation of Hg has earlier been suggested to be strictly anaerobic and to occur in the anaerobic brine environment (droplets, channels or pockets of salt solution in sea ice), or in the lower part of the ice, due to the involvement of sulfate-reducing bacteria (Chaulk et al., 2011; Gilmour et al., 2013; Beattie et al., 2014). During spring melt, heterotrophic activity could lead to oxygen depletion in the bottom of the ice column and denitrification processes that favor anoxic environments could then be important in the lower parts of the ice (Rysgaard and Glud, 2004). A study in the Antarctic has suggested methylation of Hg in sea ice might be carried out by aerophilic bacteria, such as Nitrospina, or by sea ice phytoplankton (Gionfriddo et al., 2016), but this has yet to be confirmed in the Arctic. Gionfriddo et al. (2016) suggest that sea ice phytoplankton play a role in the stimulation of methylation of oxidized Hg that may occur within brine pockets and periphytic biofilms where organic matter is trapped and decaying. Due to the seasonality of primary production, MeHg formation could show seasonal variations. Chlorophyll a (Chla) showed a covariation with MeHg,

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which also has been shown in other studies (Antarctica: Cossa et al., 2011; Arctic: Beattie et al., 2014), suggesting a microbial mercury methylation and the involvement of phytoplankton and phototrophs. Gionfriddo et al. (2016) also identify Proteobacteria, which possesses the *mer* operon, as a potentially important Hg reductive phylum. It should be noted that direct support for the production of MeHg in Arctic snow and ice (from e.g. Hg methylation assays) is still limited. From terrestrial sites, where Hg methylation and demethylation assays have been conducted on intact snow or melted snow, low or non-detectable potential Hg methylation rates were reported (Varty et al. 2021; Willis et al., 2018). In addition, both Varty et al. (2021) and Willis et al. (2018) found low methylation to demethylation ratios and concluded that *in situ* MeHg production likely did not contribute significantly to the pool of MeHg observed in the snowpack.

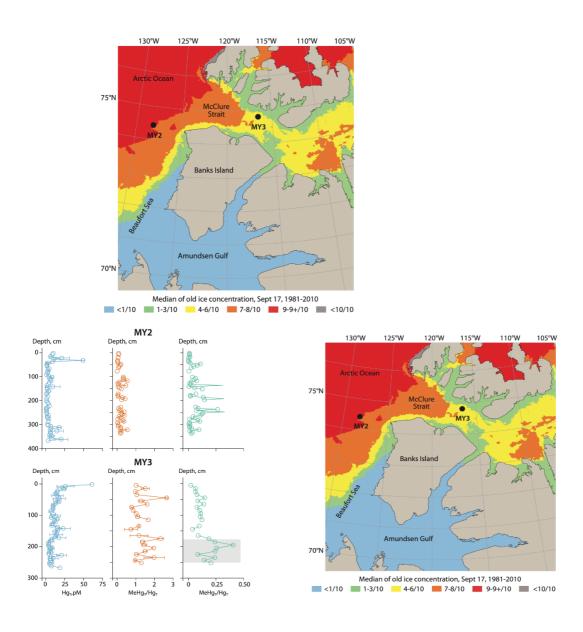


Figure 5 Profiles of total Hg (Hg_T) methylmercury (MeHg_T) concentrations and the ratio of MeHg to THg (MeHg:THg) in multi-year sea ice sampled at two sites (MY2 and MY3) in the eastern Beaufort Sea and the M'Clure Strait. Adapted from Beattie et al., 2014.

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The uptake and transport of Hg and subsequent formation of MeHg in sea ice is determined by its thickness and permeability, which varies dramatically with temperature and season. During winter, the ice is cold and less permeable, which slows the transport of Hg species through the ice. However, exchange continues with the seawater as the bottom of the ice is warmer (greater permeability) and with the atmosphere as brine exclusion expels Hg to the overlying snowpack. In spring, the ice warms up and the brine volume increases. Contaminants such as Hg then percolate from snow to ice and could lead to their flushing into the underlying seawater. Summer sea ice melt leads not only to dissolved Hg being released from the ice but also particle bound Hg and MeHg being incorporated in the bulk ice (Wang et al., 2017). The presence of sea ice also decreases photodegradation of MeHg in underlying water (Ebinghaus et al., 2002, Point et al. 2012). The dynamic motion of sea ice further makes it possible to transport Hg and MeHg over time and space as the ice can drift over large areas in the Arctic Ocean, depending on winds and underwater currents. The export of MeHg via drifting sea ice has been calculated by Wang et al. (2017) at less than 4% of the total Hg export from the Arctic Ocean. Petrova et al. (2020) estimated a net export of MeHg in sea ice of 0.4 Mg/yr, considering 10% of the total Hg export is in the form of MeHg. The total MeHg pool in sea ice in the Arctic Ocean was estimated by Schartup et al. (2020) to be 0.33 Mg, which is consistent with estimates of 0.3 to 2.6 Mg, based on studies performed in the M'Clure Strait and Beaufort Sea (Beattie et al., 2014; Soerensen et al., 2016b).

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5. Processes affecting mercury accumulation in Arctic biota

Methylmercury concentrations in fish and wildlife vary within and across geographic regions in the circumpolar Arctic, reflecting the variability in Hg loading to the ecosystem and environmental conditions that control *in situ* production of MeHg, both of which are discussed above. It also reflects on differences in biological uptake at the base of the food web, and ecological characteristics that control trophic transfer of MeHg within the food web (see Dietz et al., this issue). All of these conditions can change over time resulting in temporal variation of MeHg concentrations in animals (see Morris et al., this issue). The delineation of specific conditions promoting MeHg bioaccumulation is important for identifying Arctic ecosystems at greater risk of MeHg exposure (Munthe et al., 2007). Recent research has identified key environmental characteristics that enhance MeHg bioaccumulation in the Arctic.

Mercury speciation determines (a) how Hg is transported between different environmental

Mercury speciation determines (a) how Hg is transported between different environmental compartments, (b) organismal exposure, (c) the potential toxicity of the exposure to the organism, (d) the extent to which Hg is distributed between tissues within the organism, (e) the toxicity of the Hg compound (e.g., MMHg, DMHg, Hg(II)), (f) organism assimilation and excretion, and (g) its propensity to bioaccumulate and biomagnify through the food web. In this section, the uptake of Hg at the base of Arctic aquatic and terrestrial food webs is described, with an emphasis on the biomagnified chemical form MeHg.

5.1 Mercury uptake at the base of Arctic food chains

The bioconcentration of aqueous MeHg in bacteria and algae, either suspended in the water column (plankton) or growing in biofilms on surfaces such as sediment, ice and rocks, is understood to be a critical step in the transfer of Hg to top predators in freshwater and marine environments. Aquatic food webs with greater MeHg uptake by microorganisms also have higher concentrations in subsequent trophic levels (Wu et al., 2019), and greater uptake at the base of terrestrial food webs likely has the same consequence (Tsz-Ki Tsui et al., 2019).

Despite the critical role of phytoplankton as the dominant energy source for food webs in the ocean, few measurements of MeHg have been made of Arctic phytoplankton. In the Subarctic estuary of Lake Melville, phytoplankton contained 0.03 ng/g dw of MeHg (Schartup et al., 2015a), which is low relative to concentrations in marine phytoplankton at more temperate latitudes (Schartup et al., 2018). Phytoplankton of the Chukchi Sea had MeHg concentrations below the study analytical detection limit of 1.5 ng/g dw (Fox et al., 2017). Cellular uptake of MeHg by phytoplankton may occur by passive or active transport across the cell membrane (Lee and Fisher, 2016; Skrobonja et al., 2019). Smaller phytoplankton cells have larger surface to volume ratios and, consequently, bioconcentrate more MeHg than larger cells (Lee and Fisher, 2016; Schartup et al., 2018). Smaller phytoplankton are also known to stimulate MeHg production, because they have a higher surface to volume ratio and are more effectively remineralized (Heimbürger et al., 2010). In the Arctic Ocean, water column MeHg concentrations are often highest at subsurface depths of ~100–200 m, where the proximity to phytoplankton may enhance uptake into pelagic food webs (Heimbürger et al., 2015; Wang et al., 2018). Arctic estuaries may also be important zones for MeHg uptake by phytoplankton because of salinity-driven water-column stratification and the stimulation of MeHg production by inputs of terrestrial organic matter (Schartup et al., 2015b). Seasonal changes also affect Hg uptake by marine phytoplankton, such as in the Baltic Sea, where reduced sea ice cover and eutrophication enhanced bioaccumulation during warmers months (Bełdowska and Kobos, 2016; Soerensen et al., 2016a). Sea ice is another habitat where microbial uptake of MeHg may be critical for basal uptake into food webs. As stated before, MeHg production can occur in polar sea ice (Beattie et al., 2014; Gionfriddo et al., 2016; Shartup et al., 2020) and in the marginal sea ice zone (Heimbürger et al., 2015). Algal growth on the under-ice surface is a significant energy source for pelagic food webs in the Arctic (Kohlbach et al., 2016). Burt et al. (2013) characterized uptake of inorganic Hg by ice algae during a spring bloom associated with first-year ice and suggested that this could represent an important entry point for marine invertebrates grazing either directly on the ice or on settling particulate matter released from melting ice.

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Methylmercury within the sea ice can be taken up by ice-associated biota, algae, under-ice amphipods and ice-dependent animals such as fishes, polar bears and birds and can bioaccumulate in marine food chains throughout the year. When sea ice melts, Hg and MeHg in the ice are released to the surface ocean. Fast settling of sea ice algae upon melt (Boetius et al. 2013) may also be an important transport route for MeHg into the deep ocean, but this needs further research.

With greater light penetration into the water column, the melting period is also a period of high primary production in surface water, and the concurrent release of contaminants during this period could lead to additional uptake of contaminants within the marine food web. Mercury attached to particles is released from the ice during complete ice melt. Sinking particles may be remineralized in the water column, releasing Hg for uptake by pelagic microbes, or reach the sediment and become a source of Hg available for uptake by benthic organisms. Hg uptake by sea ice algae is only limited by the amount of Hg available for uptake when spring blooms commence (Burt et al., 2013). This is important since ice algae represent 10 to 60% of annual primary production in the Arctic, and they may become even more abundant under mild climate conditions. Burt et al. (2013) estimate that with the replacement of multi-year sea ice with first-year ice due to climate change, an extra 48 kg/y of particle-bound Hg could enter the Beaufort Sea in the coming years.

Benthic algae and biofilms are a dominant primary energy source for food webs in many Arctic lakes (Vadeboncoeur et al., 2003; Chételat et al., 2010; Rautio et al., 2011). In the eastern Canadian Arctic, MeHg concentrations ranged from 2 to 12 ng/g dw for rocky shoreline biofilms (Chételat et al., 2018), 0.3 to 1.3 ng/g dw for microbial mats (Camacho et al., 2015), and 2 to 56 ng/g dw for periphyton and filamentous algae (van der Velden et al., 2013a). In comparison, MeHg concentrations of water-column seston measured in some of those same lakes ranged from less than 0.4 to 30 ng/g dw, and bioconcentration factors of MeHg in biofilms and seston were higher in lakes with low aqueous DOC (Chételat et al., 2018).

Invertebrates play a key role in the trophic transfer of MeHg entering food webs via microorganisms.

In the last decade, large amounts of data were compiled on MeHg concentrations in marine
zooplankton from across the circumpolar Arctic (Foster et al., 2012; Pućko et al., 2014; Pomerleau et

al., 2016). Zooplankton MeHg concentrations ranged from 8 to 96 ng/g dw and were highest in the southern Beaufort Sea region compared to the Chukchi Sea, Canadian Arctic Archipelago, Hudson Bay and northern Baffin Bay (Pomerleau et al., 2016; Wang et al., 2018). Larger taxa of marine zooplankton such as *Themisto* and *Paraeuchaeta* predate on smaller zooplankton, and biomagnification of MeHg was found between omnivorous and herbivorous taxa (Foster et al., 2012; Pomerleau et al., 2016). Recent measurements of MeHg concentrations in zooplankton from Arctic lakes (Lescord et al., 2015; Chételat et al., 2018, Varty et al., 2021) add to data from earlier studies (Chételat and Amyot, 2009; Gantner et al., 2010), which showed that zooplankton in lakes can have higher concentrations (>100 ng/g dw) than those in the marine environment. In a freshwater system on Ellesmere Island (Canada), zooplankton MeHg concentrations were lowest under ice, highest at the onset of ice-free conditions and then subsequently declined, indicating considerable seasonality in bioaccumulation (Varty et al., 2021). While copepods tend to dominate in lacustrine zooplankton communities in the Arctic, large-bodied grazers (Cladocera, Daphnia) are also present in some systems and can have elevated MeHg concentrations (Chételat and Amyot, 2009; Lehnherr et al., 2012a; Lescord et al. 2015; MacMillan, 2018). Modelling shows a potential for increase distribution of daphnia across Arctic lakes due to climate-induced changes in lake productivity, a trend that could enhance MeHg bioaccumulation patterns in pelagic food chains (Chételat and Amyot 2009). The developmental stage can also influence the MeHg concentration of Arctic freshwater copepods, with dilution occurring as adults accumulate lipid stores (Chételat et al., 2012). A considerable number of studies of Arctic food webs have also characterized Hg concentrations in a variety of benthic invertebrate taxa in marine environments (Bidleman et al., 2013; van der Velden et al., 2013a; Fox et al., 2014; Jörundsdóttir et al., 2014; Clayden et al., 2015; McMeans et al., 2015; Góngora et al., 2018; Pedro et al., 2019) as well as in lakes (van der Velden et al., 2013a; Chételat et al., 2015, 2018; Kahilainen et al., 2016, 2017; Keva et al., 2017; Rohonczy et al., 2020). These studies showed variation in Hg bioaccumulation among benthic invertebrate taxa, which was partly related to trophic position.

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5.2. Processes affecting mercury bioaccumulation at higher trophic levels

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Mercury concentrations in fish and wildlife reflect ecological processes that influence dietary exposure as well as animal physiology that affects internal accumulation of Hg (Bradley et al., 2017, Chételat et al., 2020). Trophic position, shifts in diet, habitat-specific feeding, cross-ecosystem subsidies and movement or migration can be important determinants of dietary Hg consumed by fish and wildlife. Differences in the extent of assimilation and elimination of MeHg related to sex, age, body condition and species also affect animal Hg burdens. Recent advances have identified Arctic-specific processes affecting Hg bioaccumulation in fish and wildlife. Arctic marine ecosystems have long food chains (~4–5 levels), and top predators (such as polar bear, ringed seal, beluga whale, and Greenland shark) have elevated Hg concentrations due to feeding at a higher trophic position (St. Louis et al., 2011; McMeans et al., 2015; Brown et al., 2016, 2018). These slow-growing, long-lived top predators are prone to long-term Hg bioaccumulation with age (Masbou et al., 2018; Ewald et al., 2019). Recent quantification of food-chain length using nitrogen stable isotopes demonstrated that regional differences in Hg concentrations of top predators, such as between western and eastern regions of the Canadian Arctic (St. Louis et al., 2011; Brown et al., 2016) is, in part, due to food-chain length. There can also be spatial variation in the carbon sources which support Arctic marine food webs, with potential consequences for food web bioaccumulation of Hg (Cyr et al., 2019). Recent variation of Hg concentrations in beluga whales reflect complex influences of diet, foraging habitat (e.g., benthic versus pelagic feeding), and climate drivers (Desforges et al., 2022; Loseto et al., 2015, Loseto et al., 2008). Considerable research effort has also focused on measuring the rate that Hg biomagnifies through food webs in Arctic ecosystems. A global meta-analysis showed that biomagnification rates were higher at northern latitudes compared with the tropics (Lavoie et al., 2013). Within the Arctic, Hg biomagnification rates can vary among marine and freshwater food webs (Clayden et al., 2015; Lescord et al., 2015; McMeans et al., 2015; Ruus et al., 2015; Fox et al., 2017); although in a study of co-located marine and lacustrine food webs of Arctic char, no difference was found in biomagnification rates (van der Velden et al., 2013b). While some evidence suggests that

cold, low productivity ecosystems have higher Hg biomagnification (Lavoie et al., 2013), factors 780 controlling the biomagnification rate of Hg through Arctic food webs remain unclear. 781 Marine subsidies to the diet of terrestrial carnivores and freshwater fish can strongly affect Hg 782 bioaccumulation in the Arctic. Wolves and Arctic foxes that scavenged on carcasses of marine 783 mammals had higher Hg concentrations than those that had an exclusively terrestrial diet (Bocharova et al., 2013; McGrew et al., 2014; Hallanger et al., 2019). In the case of Arctic fox at Svalbard, 784 785 interannual variation in their Hg concentration was positively related to the extent of sea ice cover, 786 which provides a platform to access marine carcasses (Hallanger et al., 2019). Sea-run Arctic char tend 787 to be larger and have faster growth rates than their lacustrine conspecifics, and therefore they have lower muscle Hg concentrations (Swanson et al., 2011; van der Velden et al., 2013b). The lower Hg 788 789 bioaccumulation in sea-run char was also attributed to lower Hg exposure at the base of marine food 790 webs relative to lakes in the Arctic (van der Velden et al., 2013a). 791 Migration by wildlife such as seabirds and marine mammals can influence their dietary Hg exposure. 792 Little auks (Alle alle) that breed in Greenland during summer and overwinter off the coast of eastern 793 Canada were found to accumulate more Hg during the wintering period (Fort et al., 2014). Similarly, 794 yellow-billed loons (Gavia adamsii), which breed in lakes of Arctic North America and overwinter in 795 marine areas of eastern Asia, showed evidence of higher winter exposure (Evers et al., 2014). An 796 evaluation of seasonal variation in Hg contamination of nine species of Arctic seabirds showed, on 797 average, three times higher Hg concentrations in feathers grown at lower latitudes during the non-798 breeding period than those grown while breeding in the Arctic (Albert et al., 2021). Greater Hg 799 bioaccumulation was found for seabirds overwintering in the western Atlantic relative to the eastern 800 Atlantic or the Pacific region (Albert et al., 2021). Thus, consideration of overwinter exposure at lower 801 latitudes is necessary for interpreting Hg bioaccumulation in migratory Arctic seabirds. Bowhead 802 whales (Balaena mysticetus) undergo long seasonal migrations within the Arctic in response to sea ice 803 conditions and prey availability, and examination of their baleen plates provided a record of temporal variation in Hg exposure (Pomerleau et al., 2018). Annual cycles of Hg and δ^{13} C values (an isotopic 804

805 indicator of diet) in the baleen plates showed seasonal shifts in diet and Hg accumulation by bowhead 806 whales in both the eastern and western Arctic. 807 Biological transport is also a pathway for the mobility of contaminants in the Arctic environment such 808 as from migratory spawning fish and bird colonies (Michelutti et al. 2010, Choy et al. 2010, 809 Kristiansen et al., 2019). Migrating salmonids that return to remote freshwater ecosystems to spawn 810 act as important transport (bio)vectors of Hg from ocean to rivers (Blais et al., 2007). Similarly, high-811 density nesting of seabirds and their deposition of guano can lead to inputs of marine-derived nutrients 812 and Hg to nearby ponds and tundra (Michelluti et al., 2010; Kristiansen et al., 2019, Choy et al., 2010). 813 The body condition of fish and wildlife is sometimes correlated with their Hg concentration, although 814 reported observations have been both positive and negative. Trudel and Rasmussen (2006) showed 815 that Hg concentration may be either positively or negatively correlated to changes in growth and 816 growth efficiency, depending on how fish allocate energy between activity costs and growth rates. 817 Mercury concentrations of northern pike were negatively correlated with growth rate though also 818 influenced by water chemistry and characteristics of the subarctic catchments (Moslemi-Aqdam et al., 819 2022). Tissue Hg concentrations were negatively correlated with body condition in sea-run char 820 (Evans et al., 2015), caribou (Rangifer tarandus groenlandicus; Gamberg et al., 2016) and little auks 821 (Amélineau et al., 2019), while positive correlations were observed for eider (Provencher et al., 822 2016). For polar bear, both negative and positive correlations were found between tissue Hg 823 concentrations and body condition (Bechshoft et al., 2016; McKinney et al., 2017). Those contrasting 824 observations reflect several potential underlying physiological processes driving the effect of body 825 condition. Fasting or lower food availability and the subsequent loss of body mass can concentrate Hg 826 already in tissues (Keva et al., 2017; Seewagen et al., 2016; Peterson et al., 2018; Martyniuk et al. 827 2020). The reverse has also been suggested, that poorer body condition may be a negative effect of a 828 higher Hg burden (Provencher et al., 2016; Amélineau et al., 2019), although this has been discounted 829 for northern pike (Esox lucius; Sandheinrich and Drevnick, 2016). A positive association between 830 body condition and Hg concentration could result from greater Hg intake associated with increased 831 feeding (Bechshoft et al., 2016). Given these bioenergetic influences, seasonal and interannual

changes in food resource availability in the Arctic can have important consequences for Hg concentrations in animals (Amélineau et al., 2019; Eckbo et al., 2019; Hallanger et al., 2019).

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5.3. Organic matter influences on mercury bioaccumulation

While numerous water chemistry parameters have been shown to alter MeHg bioaccumulation in Arctic systems, organic matter has complex linkages with the Hg cycle that can both dampen or promote MeHg bioaccumulation. The complexity of these interactions is reflected in the varying relationships reported between organic matter and Hg in the literature (Lavoie et al., 2019). A recent global review found these relationships to be broadly influenced by geographic location and temporal scale, as well by as human activities (Lavoie et al., 2019). The chemical composition of DOC also plays a critical role in its effects on Hg biogeochemical cycling (Schartup et al., 2015b; Lescord et al., 2018; Jiang et al., 2020), including bioaccumulation (see Section 4.4). For example, OC of terrestrial origin is involved in the transport of MeHg to downstream freshwater and marine environments. In the Subarctic Baltic Sea, interannual variation of seawater MeHg concentrations were related to DOC inputs and composition (Soerensen et al., 2017). In Arctic fresh waters, DOC can promote or inhibit the photodegradation of MeHg depending on the amount of DOC present (Girard et al., 2016) and alter Hg(II) bioavailability to microbes (Chiasson-Gould et al., 2014) because of the multiple roles that organic matter plays in those processes. Dissolved organic carbon can play a key role in promoting MeHg bioaccumulation, although those interactions are complex, and consideration of site-specific conditions are likely important. Numerous studies on the effects of organic matter on Hg accumulation in Arctic biota have been conducted in lakes and rivers across the Arctic over the past ten years. For example, Hudelson et al. (2019) recently reported strong negative correlations between DOC and POC concentrations versus total Hg (THg) levels in landlocked Arctic char (Salvelinus alpinus) in lakes around Resolute Bay in Nunavut, Canada. In the same set of lakes, Lescord et al. (2015) reported negative correlations

between DOC (in combination with total nutrient levels) and MeHg concentrations in zooplankton. In

High Arctic lakes, waters with very low concentrations of DOC were associated with greater bioaccumulation of MeHg at the base of the food web (e.g., chironomid larvae and zooplankton) even when concentrations of MeHg in the water were low (Chételat et al., 2018). Overall, these results suggest that Arctic aquatic systems with lower DOC concentrations, in combination with other physicochemical variables, may have higher rates of Hg bioaccumulation in biota. However, variation in the influence of DOC likely occurs. Furthermore, climate change is rapidly altering both carbon and Hg cycling, resulting in a general increase in concentrations of organic matter in aquatic environments across the Arctic (MacDonald et al., 2005; Isles et al., 2020; Chételat et al. 2022). Some research in Arctic and Subarctic systems has also suggested that the relationship between biotic Hg and organic matter concentrations may not be linear and may, in fact, be dependent on DOC concentrations. More specifically, several studies have reported a curvilinear relationship between DOC concentrations and THg or MeHg concentrations in aquatic biota, with an inflection point of ~8-11 mg/L of DOC (French et al., 2014; Braaten et al., 2018). This relationship was first suggested based on data from the Adirondack lakes and is dependent on system pH (Driscoll et al., 1994). However, given that concentrations of DOC are often strongly correlated with aqueous concentrations of both THg and MeHg, the curvilinear relationship between DOC and biotic Hg concentrations warrants further investigation with consideration of other factors such as water chemistry and trophic position (Chételat et al., 2018; Broadley et al., 2019). For example, Chételat et al. (2018) found that the ratio of aqueous MeHg concentrations to DOC concentrations was a strong and linear predictor of invertebrate MeHg concentrations across an array of Arctic lakes. Recent studies, some related to this curvilinear Hg-DOC relationship, also consider the effect of various measures of DOM quality rather than quantity (i.e., DOC concentrations) have on Hg bioaccumulation. French et al. (2014) suggest that changes in DOM compounds and bioaccessibility may account for changes in Hg bioavailability and bioaccumulation in Arctic freshwater invertebrates; at lower DOC concentrations, Hg was found to be bound with smaller fulvic acids, while at higher DOC concentrations Hg was bound with larger and less bioaccessible humic acids. In the Subarctic, biotic Hg concentrations were similarly higher in systems with higher proportions of more

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bioaccessible DOM compounds and lower proportions of large and aromatic humic substances (Lescord et al., 2018). Conversely, other studies have reported that humic lakes are associated with greater MeHg bioaccumulation in Subarctic boreal ecoregions (Braaten et al., 2018; Poste et al., 2019). Future studies on Arctic Hg and DOM interactions may benefit from incorporating simple structural measurements, such as SUVA254, which provides an estimate of DOM aromaticity (e.g., Skyllberg et al., 2009; Burns et al., 2013). While several studies have shown changes in aqueous Hg levels in relation to permafrost degradation and changing organic matter dynamics, none to our knowledge have examined their combined effect on biotic Hg levels, Hg bioaccumulation potentials, or rates of biomagnification. In the retrogressive thaw slump in the western Canadian Arctic, over 95% of the Hg was released in the particulate phase (St. Pierre et al. 2018) and may not be readily bioavailable. Furthermore, a recent review by Tank et al. (2020) found considerable variation in DOC concentrations and DOM quality among studies examining the impact of permafrost thawing on biogeochemical cycling across aquatic networks. Therefore, future studies assessing the impact of thawing permafrost on downstream bioaccumulation patterns would benefit from considering the effect of DOM, including the structural or quality measurements discussed above.

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6. Conclusions and recommendations

Conclusions (in numbered bullets) are provided below under section headings followed by recommendations in italics where appropriate.

How much methylmercury is circulating in the Arctic environment?

- While no estimates of MeHg pools in the environment reservoirs, nor fluxes, were available in the previous AMAP 2011 Mercury Assessment, measurements of MeHg in the Arctic environment over the last decade have allowed for a model budget, as presented in Table 1, Table 2 and Figure 1.
- 2. Seawater *in situ* Hg methylation is an important feature leading to the high MeHg portion (32%) and a net export of MeHg (~5–11 Mg/y) from the Arctic Ocean to the North Atlantic.
- 3. Arctic sediments are a potential source of MeHg to the Arctic marine water column.

Estimated and modeled MeHg budgets and fluxes are based on limited in situ measurements. Hence, more field measurements are needed to fill important knowledge gaps and complete the total Arctic MeHg budget. Geographically, the Siberian shelf seas and the central Arctic Ocean are important regions where there are important data gaps to address. In addition, future models would benefit from a better understanding of how the cycling of Hg varies over seasons and between years. Further work is also warranted to better understand and improve the comparability of data between studies (e.g. through inter laboratory comparisons and development of new certified reference materials) as future synthesis and modelling work will continue to rely on data from multiple labs.

Given the biogeochemical cycling of Hg in the Arctic environment is sensitive to climate change, its effects on MeHg production and breakdown in the Arctic, including the connection between climate

change and changes in DOM discharges, needs further study. Further studies should thus cover both mercury speciation (i.e. quantification of methylated Hg forms) as well as ancillary parameters to link Hg speciation with changes in water chemistry and biological factors (such as the composition of the microbial community).

What controls the pool of mercury available for bioaccumulation?

- 4. Methylation and demethylation are key processes controlling the pool of MeHg available for bioaccumulation in Arctic food webs. Methylation of inorganic Hg in the environment is primarily a microbial process while demethylation of MeHg occurs through both biotic and abiotic pathways.
- Organic matter quantity and quality vary greatly among Arctic environments and plays a key role in both the availability of Hg for microbial methylation as well as the uptake of MeHg into organisms.
- 6. Microbial Hg methylation occurs via organisms that carry the *hgcAB* gene pair. Microorganisms carrying the *hgcAB* gene pair have been identified in Arctic soils and thawing permafrost, and those carrying HgcA or hgcAB-like genes have recently been identified in global ocean seawater, though not in the Arctic Ocean. Organisms carrying genes that encode the *merA* and *merB* operons (a biotic pathway for Hg reduction and demethylation) have also been observed in the Arctic environment.
- 7. Evasion of DMHg from the marine boundary layer, and subsequent demethylation of DMHg in the atmosphere, may be a source of MeHg deposition to Arctic snow, surface ocean water, and coastal terrestrial ecosystems. In addition, DMHg may degrade to MeHg through photochemical or dark abiotic pathways while still being dissolved in surface waters.

Future mapping of microbial communities carrying the genes involved in methylation and demethylation of Hg should be carried out as such mapping will provide critical information on these processes across the Arctic, as well as reveal key linkages to the biogeochemical

cycle of other elements such as carbon (e.g. by understanding the involvement of methanogens in Hg methylation).

The thawing of permafrost represents a potential source of both DOM and Hg to Arctic and Subarctic aquatic environments; future work should examine the effects of permafrost-derived organic matter on Hg bioaccumulation in Arctic biota, which may add to our understanding of the complex Hg-DOM relationship as permafrost thawing increases in a changing climate. Future research should also be directed towards a better understanding of the formation and environmental fate of DMHg in the Arctic Ocean, and its role in controlling the pool of Hg available for bioaccumulation in marine and coastal food webs.

Where are the hotspots of methylmercury in the Arctic?

- 8. We previously knew that the production of MeHg in natural systems was controlled by substrate availability (e.g., inorganic Hg and organic matter), the presence and abundance of microbial Hg methylators, and MeHg degradation processes. Thus, anoxic systems, such as organic-rich sediments and flooded soils, are understood to be typical hotspots for MeHg production in the Arctic as elsewhere.
- 9. In the terrestrial environment, foliar uptake of Hg from the atmosphere, a process that could be enhanced in recently or soon-to-be deglaciated foreland, is a key driver for sequestering inorganic Hg that may be available for methylation.
- 10. Permafrost thaw in low-relief regions with ice-rich soils and sediments may result in the formation of thermokarst wetlands, ponds and lakes that have been repeatedly identified as Hg methylation hotspots across the Arctic due to the concurrent mobilization of carbon and nutrients, creating conditions suitable for methylation.
- 11. Locally significant MeHg production may occur in newly-deglaciated forelands (especially where wetlands develop), in cryoconite holes on the margins of glacier surfaces or margins,

979	and at the head of glacially-fed fjords and in the sediments within these fjords, though limited
980	evidence is available on the extent of occurrence.
981	12. In the Arctic Ocean., MeHg is found to be enriched in subsurface, oxic seawater at shallower
982	depths (100-300 m) than in the Global Ocean.
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984	Future work is needed to better identify sources and pathways of MeHg production and
985	storage in the Arctic terrestrial environment and how melting glaciers (and glacier sediments)
986	and thawing permafrost would contribute to MeHg loading of downstream environments.
987	The processes that are responsible for the enrichment of MeHg in Arctic subsurface seawater
988	remain poorly understood. Methylation and demethylation rates of Hg in Arctic seawater need
989	to be better quantified.
990	Future work should also aim to confirm the occurrence, or lack thereof, of Hg methylation in
991	Arctic sea ice and/or snow.
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993	What processes are important for biological uptake of methylmercury?
994	13. As we understood previously, bioconcentration of aqueous MeHg in bacteria and algae is a
995	critical step in the transfer of Hg to top predators in freshwater and marine environments.
996	Methylmercury is biomagnified through the food web. Thus, the concentration of Hg Arctic
997	animals at the top of the food web depends on the amounts of MeHg entering the food web,
998	the length of the food web as well as seasonal dietary changes, migration, and body condition.
999	14. Sea ice algae may be an important entry point for MeHg in marine ecosystems.
1000	15. Zooplankton in Arctic lakes can have higher concentrations than those in the marine
1001	environment.
1002	16. Recent nitrogen stable isotope measurements specific to Arctic biota have highlighted the
1003	importance of variable trophic position in controlling MeHg concentrations among species and
1004	among populations of the same species distributed across the circumpolar Arctic.

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1006	While some evidence suggests cold, low-productivity ecosystems have higher Hg
1007	biomagnification, factors controlling the biomagnification rate of Hg through Arctic food
1008	webs needs to be understood more clearly.
1009	The large magnification at the base of the food chain and its connection to DOM is a critical
1010	process that needs further study.
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