



# A Global-Scale Analysis of Hydrologic Extremes using Hidden Climate Indices

Benjamin Renard, David Mcinerney, Seth Westra, Michael Leonard, Dmitri Kavetski, Mark Thyer, Jean-Philippe Vidal

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# Hydrologic Extremes at the Global Scale

## 100-year Analysis and 180-year Reconstruction

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M. Leonard<sup>2</sup>   D. Kavetski<sup>2</sup>   M. Thyre<sup>2</sup>   J.-P. Vidal<sup>1</sup>

<sup>1</sup>INRAE, RiverLy Research Unit, Lyon, France

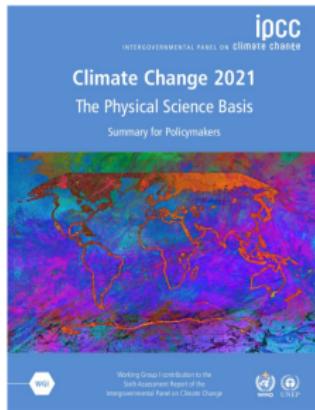
<sup>2</sup>School of Civil, Environmental and Mining Engineering, University of Adelaide, Australia

<sup>3</sup>INRAE, RECOVER Research Unit, Aix-en-Provence, France

EGU General Assembly, 27 May 2022

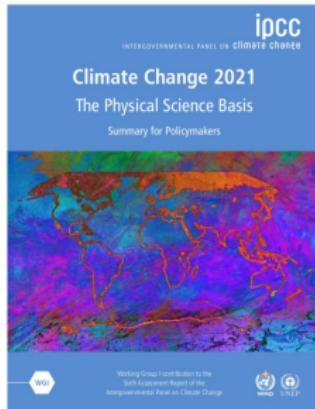


# Introduction



*"The frequency and intensity of heavy precipitation events have increased since the 1950s over most land area for which observational data are sufficient [...]"*

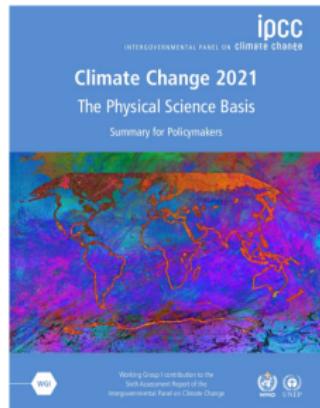
# Introduction



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*"Confidence about peak flow trends over past decades on the global scale is low [...]"*

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## Water Resources Research

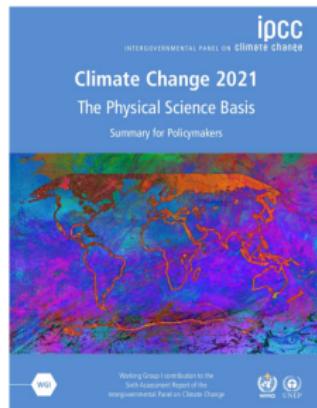
### COMMENTARY

10.1029/2018WR023749

### If Precipitation Extremes Are Increasing, Why Aren't Floods?

Ashish Sharma<sup>1</sup> , Conrad Wasko<sup>2</sup> , and Dennis P. Lettenmaier<sup>3</sup>

# Introduction



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*"Confidence about peak flow trends over past decades on the global scale is low [...]"*

## Water Resources Research

### COMMENTARY

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### If Precipitation Extremes Are Increasing, Why Aren't Floods?

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*"changes may be more complex than simple trends"*

## Water Resources Research

### RESEARCH ARTICLE

10.1029/2019WR026575

#### Key Points:

- A method from Scan statistics is used to detect periods of flood-rich and flood-poor years.

### Detecting Flood-Rich and Flood-Poor Periods in Annual Peak Discharges Across Europe

David Lun<sup>1</sup> , Svenja Fischer<sup>2</sup> , Alberto Viglione<sup>1</sup> , and Günter Blöschl<sup>1</sup>

# Objectives

Better understand the temporal variability of heavy precipitation (P) and flood (Q) at the global scale by means of an innovative probabilistic model

## 100-year analysis

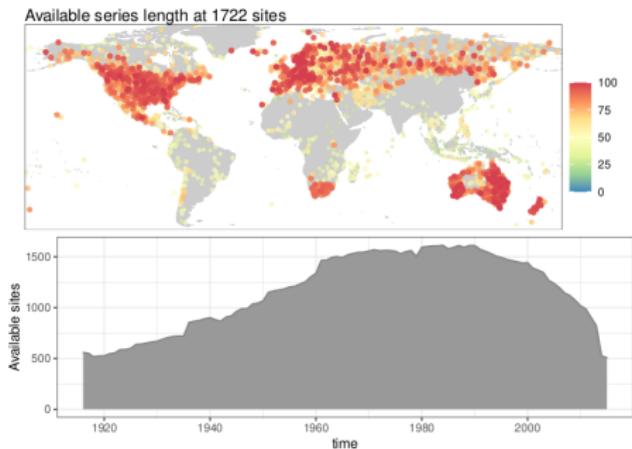
- Identify common (P+Q) vs. specific (P-only or Q-only) signals behind global extremes
- Look for trends and low-frequency variability in those signals

## 180-year reconstruction

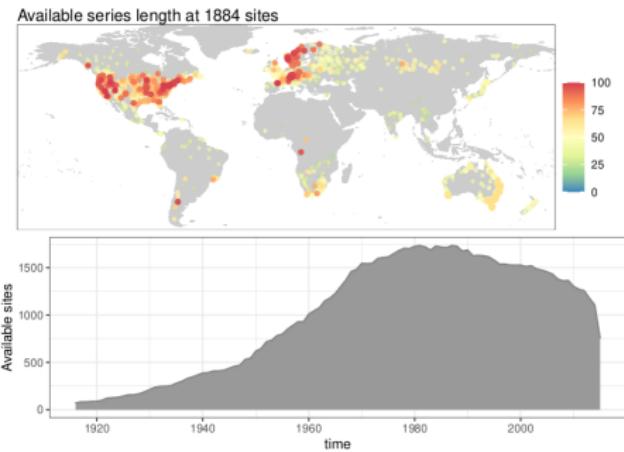
Using 20CRv3, reconstruct probabilities of extreme P/Q since 1836

# Global datasets

**P: a selection from Hadex 2+3**  
Donat et al. (2013); Dunn et al. (2020)



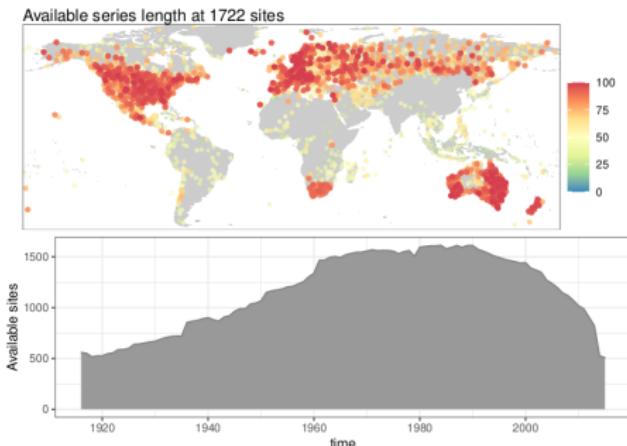
**Q: a selection from GSIM**  
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# Global datasets

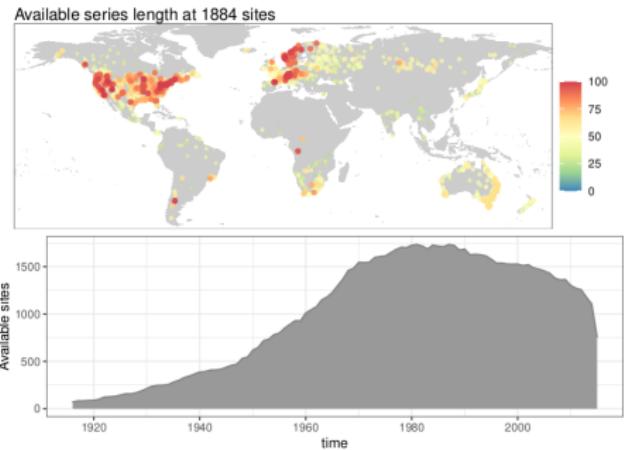
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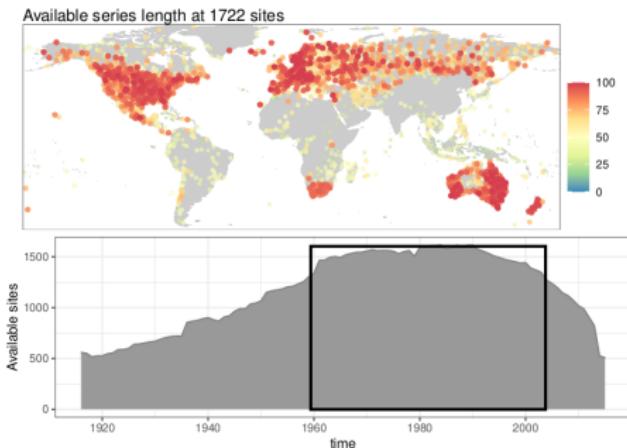


- Extract seasonal maxima at each site (SON, DJF, MAM, JJA)

# Global datasets

P: a selection from Hadex 2+3

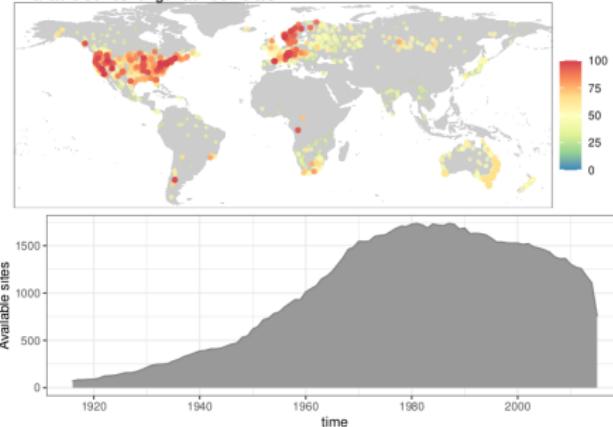
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Q: a selection from GSIM

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Available series length at 1884 sites

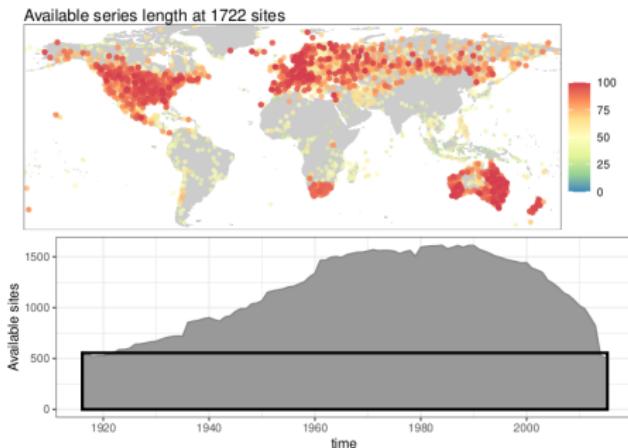


- Extract seasonal maxima at each site (SON, DJF, MAM, JJA)
- The rectangle dilemma...

# Global datasets

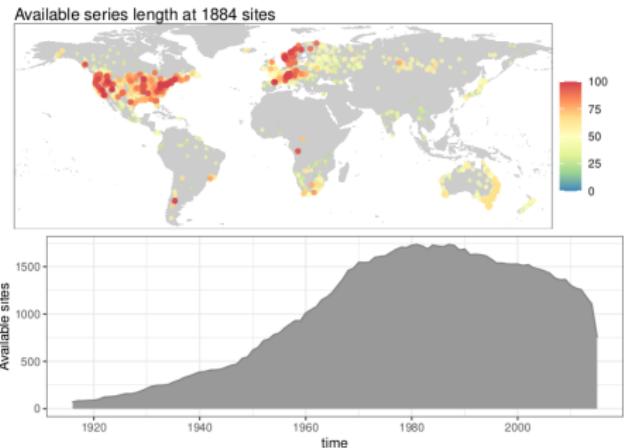
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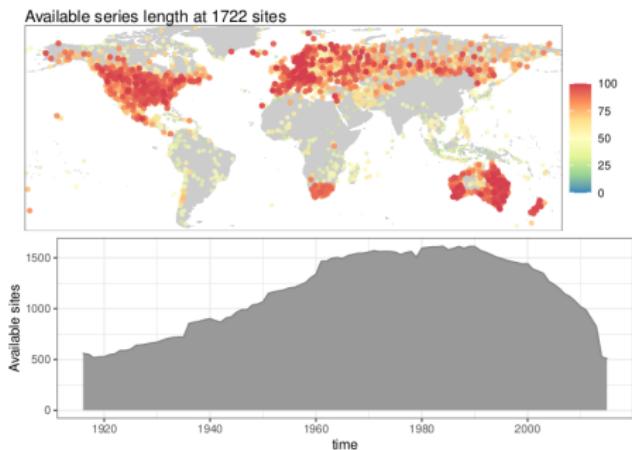
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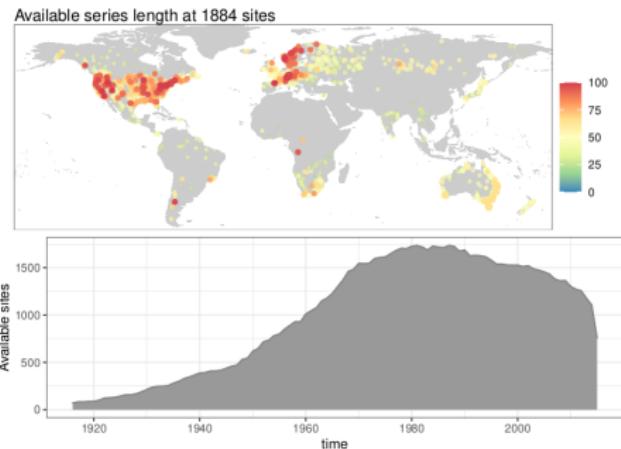
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# Global datasets

**P: a selection from Hadex 2+3**  
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**Q: a selection from GSIM**  
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- Extract seasonal maxima at each site (SON, DJF, MAM, JJA)
- The rectangle dilemma...
- In this work, we'll use all data available during 1916-2015

# Model

After suitable data transformation...

$$\begin{cases} \mathbb{E}[P(s, t)] = \\ \mathbb{E}[Q(s, t)] = \end{cases}$$

**Legend:** varies in space and time

# Model

After suitable data transformation...

$$\begin{cases} \mathbb{E}[P(s, t)] = c_P(s) \\ \mathbb{E}[Q(s, t)] = \underbrace{c_Q(s)}_{\text{constant (intercept)}} \end{cases}$$

**Legend:** varies in space and time ; varies in space

# Model

After suitable data transformation...

$$\begin{cases} \mathbb{E}[P(s, t)] = c_P(s) + \lambda_P(s)\tau(t) \\ \mathbb{E}[Q(s, t)] = c_Q(s) + \lambda_Q(s)\tau(t) \end{cases}$$

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$$\begin{cases} \mathbb{E}[P(s, t)] = c_P(s) + \lambda_P(s)\tau(t) + \pi(s)\delta(t) \\ \mathbb{E}[Q(s, t)] = c_Q(s) + \lambda_Q(s)\tau(t) + \theta(s)\omega(t) \end{cases}$$

**Legend:** varies in space and time ; varies in space ; varies in time

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After suitable data transformation...

$$\left\{ \begin{array}{l} \mathbb{E}[P(s, t)] = c_P(s) + \lambda_P(s)\tau(t) + \overbrace{\pi(s)\delta(t)}^{\text{P-specific covariate and its effects}} \\ \mathbb{E}[Q(s, t)] = c_Q(s) + \lambda_Q(s)\tau(t) + \overbrace{\theta(s)\omega(t)}^{\text{Q-specific covariate and its effects}} \end{array} \right.$$

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**Legend:** varies in space and time ; varies in space ; varies in time

- All covariates are considered unknown and are estimated  
→ *Hidden Climate Indices* (HCl)

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- $\tau(t) \sim \text{AR}(1) + \text{trend.}$

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- $\tau(t) \sim \text{AR}(1) + \text{trend}$ . Same for  $\delta(t)$  and  $\omega(t)$
- $\lambda(s) \sim \text{Spatial Gaussian Process}$ . Same for others

# Model

After suitable data transformation...

$$\left\{ \begin{array}{l} \mathbb{E}[P(s, t)] = c_P(s) + \lambda_P(s)\tau(t) + \pi(s)\delta(t) + \text{more components...} \\ \mathbb{E}[Q(s, t)] = c_Q(s) + \lambda_Q(s)\tau(t) + \theta(s)\omega(t) + \text{more components...} \end{array} \right.$$

  
one component

**Legend:** varies in space and time ; varies in space ; varies in time

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- One component not enough at the global scale → 5 used here

# Model

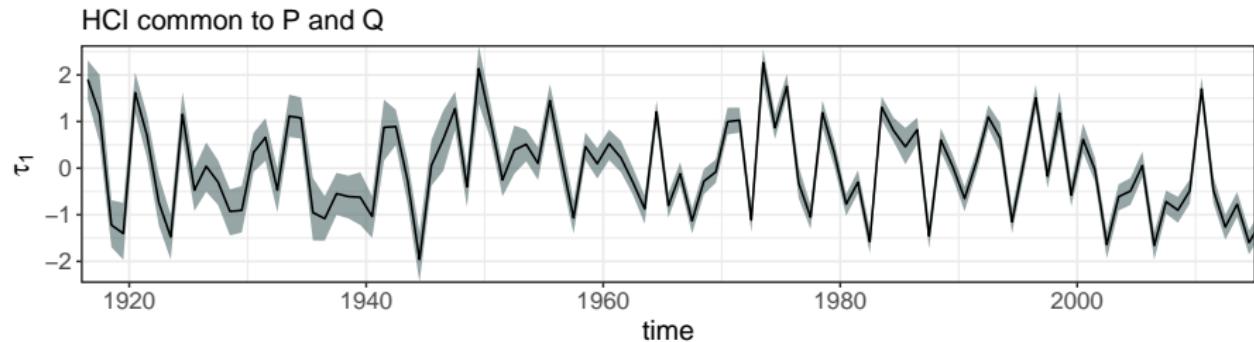
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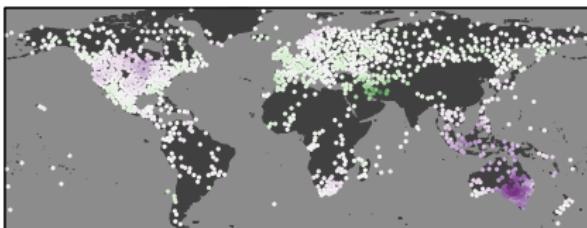
**Legend:** varies in space and time ; varies in space ; varies in time

- All covariates are considered unknown and are estimated  
→ *Hidden Climate Indices* (HCI)
- $\tau(t) \sim \text{AR}(1) + \text{trend}$ . Same for  $\delta(t)$  and  $\omega(t)$
- $\lambda(s) \sim \text{Spatial Gaussian Process}$ . Same for others
- One component not enough at the global scale → 5 used here
- (Bayesian + MCMC) estimation

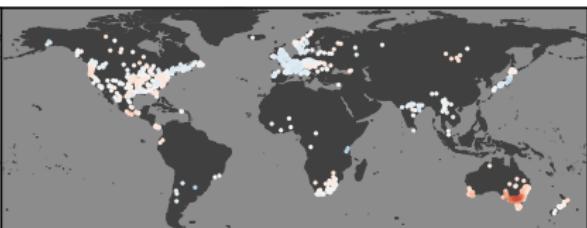
Ex.: SON season, 1st common HCI



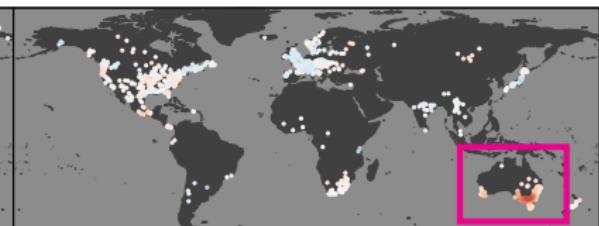
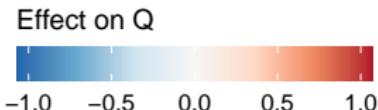
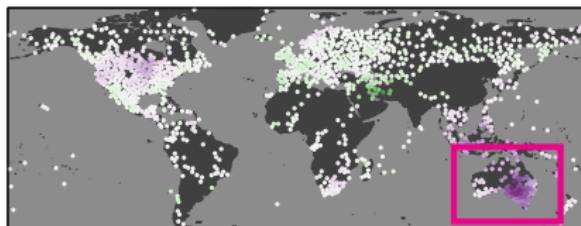
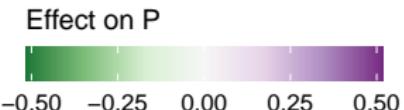
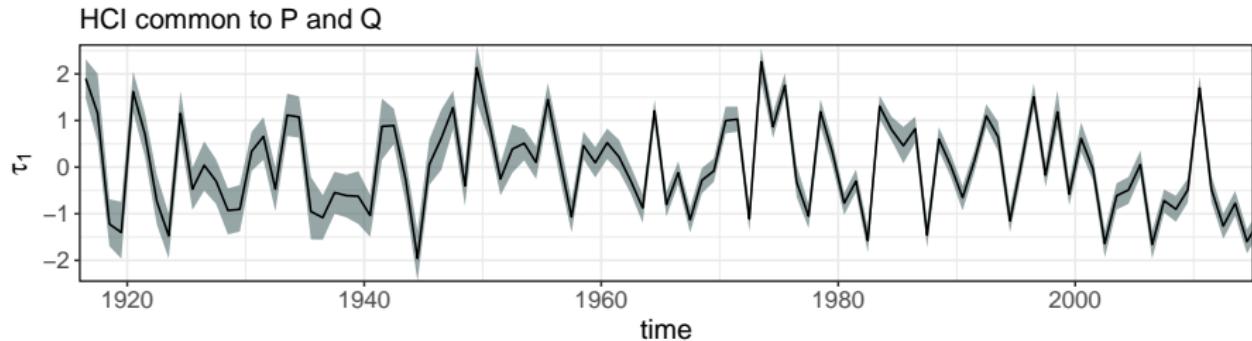
Effect on P



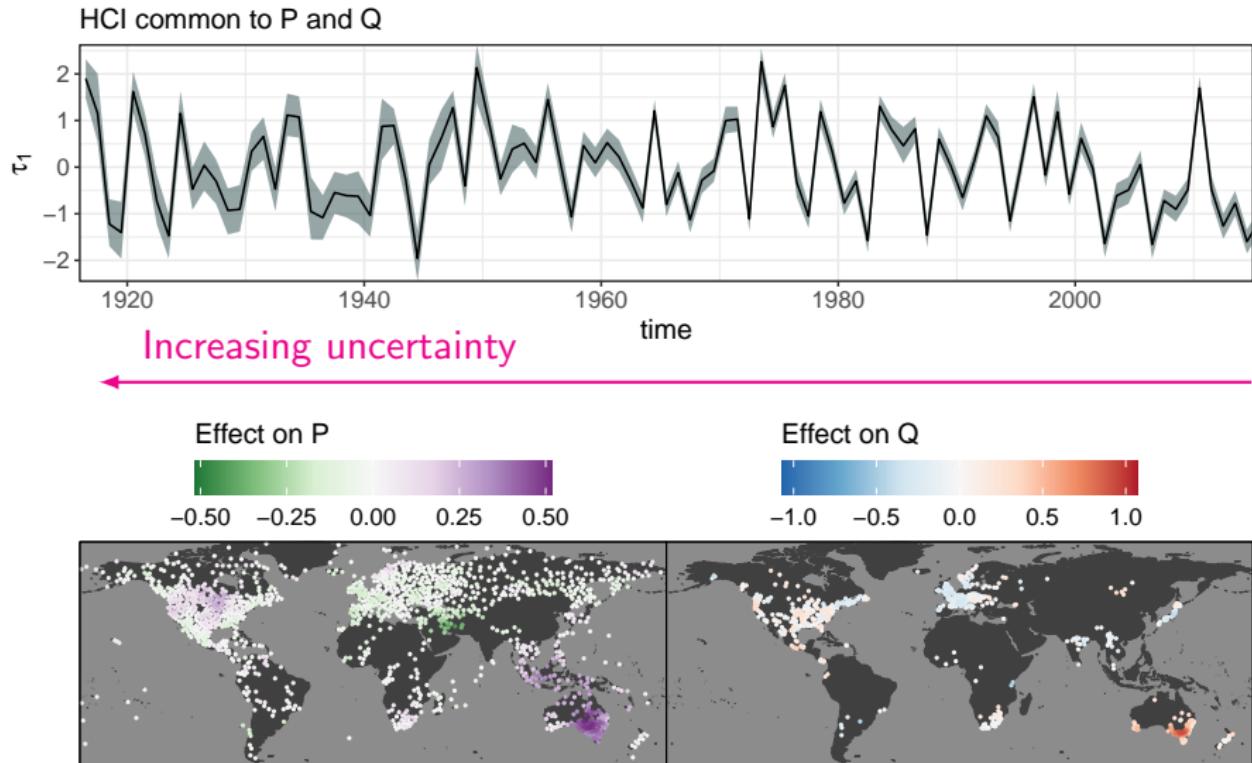
Effect on Q



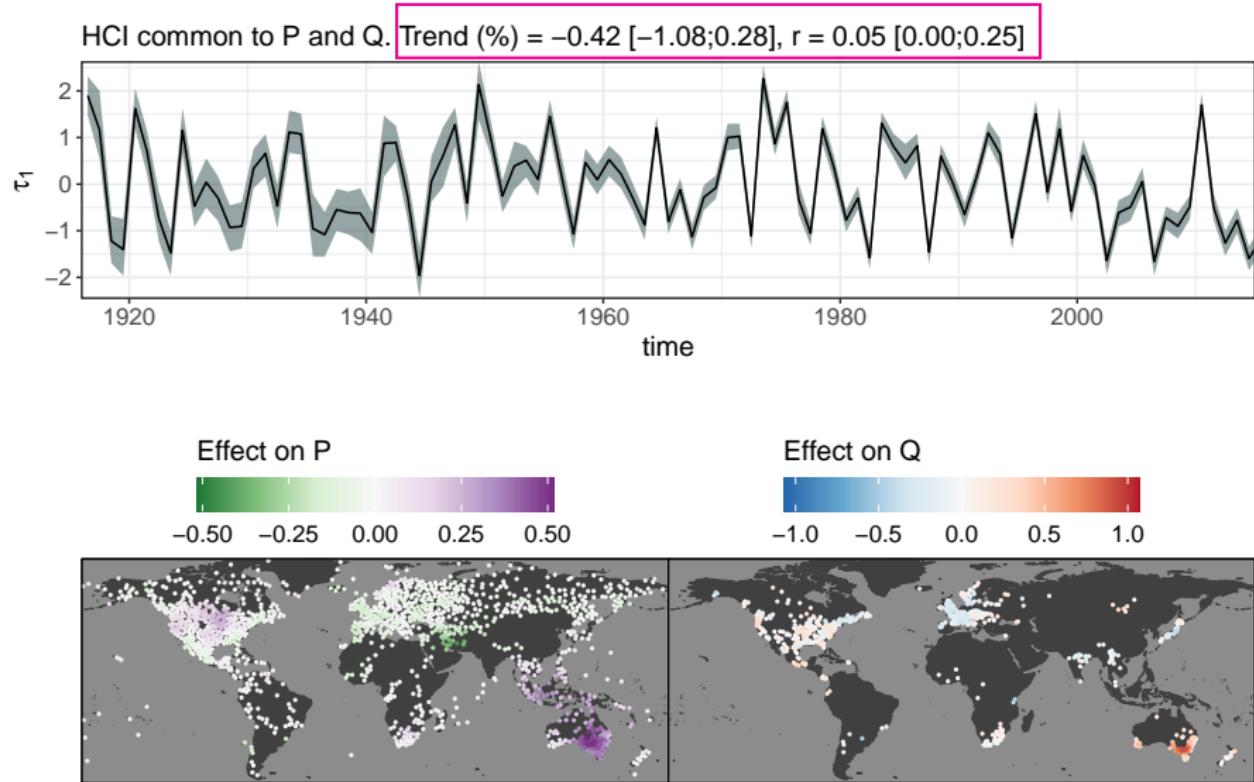
Ex.: SON season, 1st common HCI



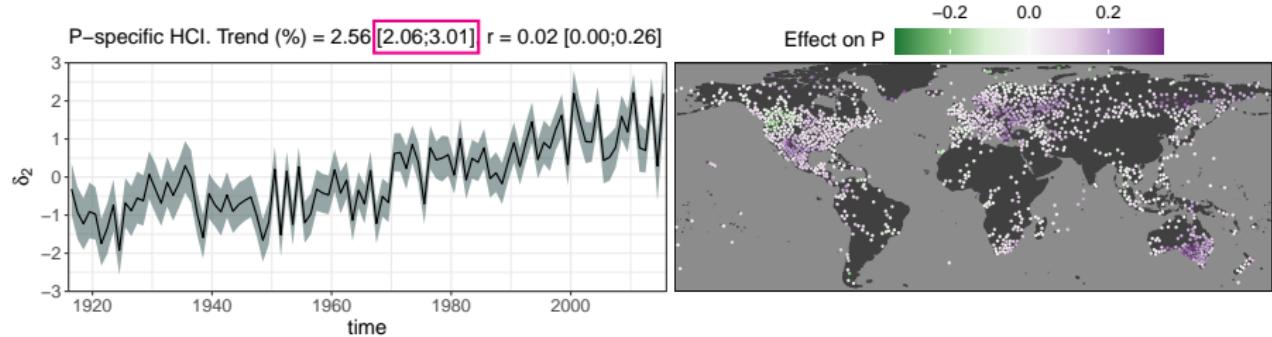
## Ex.: SON season, 1st common HCI



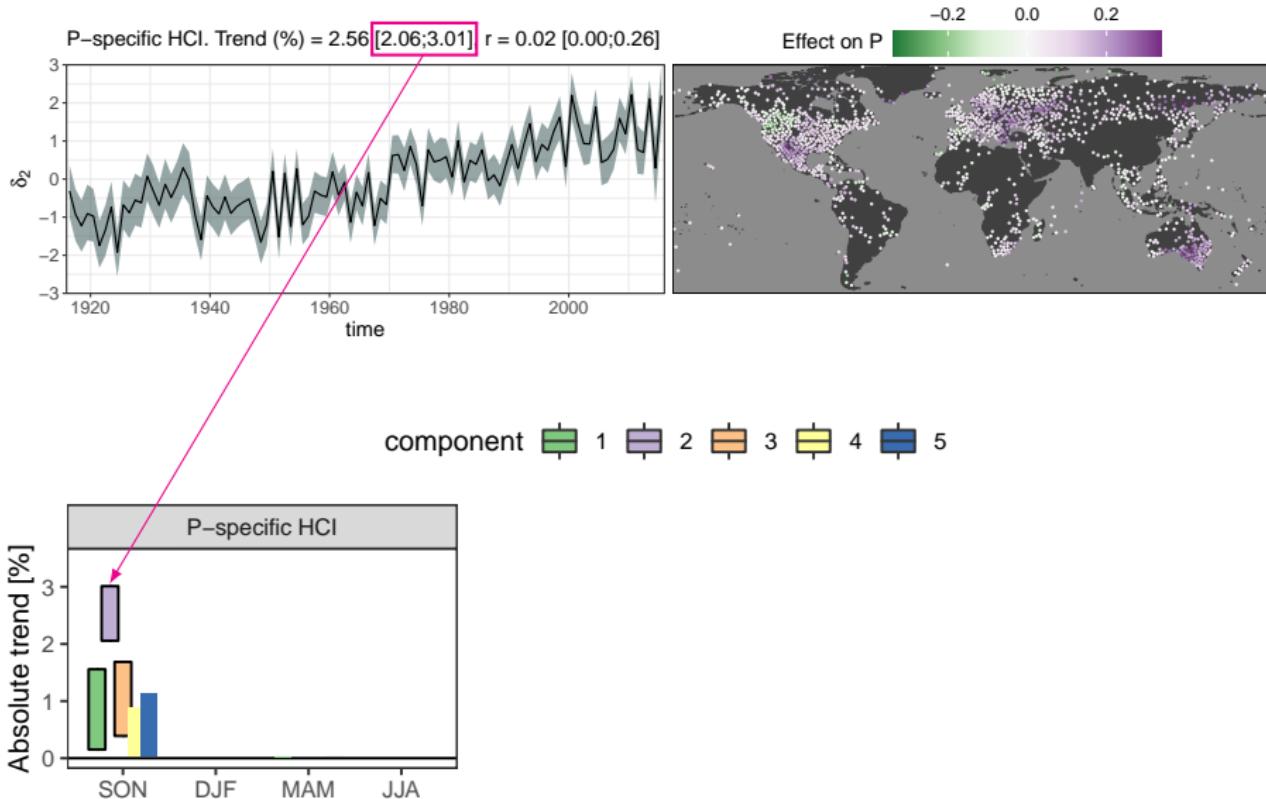
# Ex.: SON season, 1st common HCI



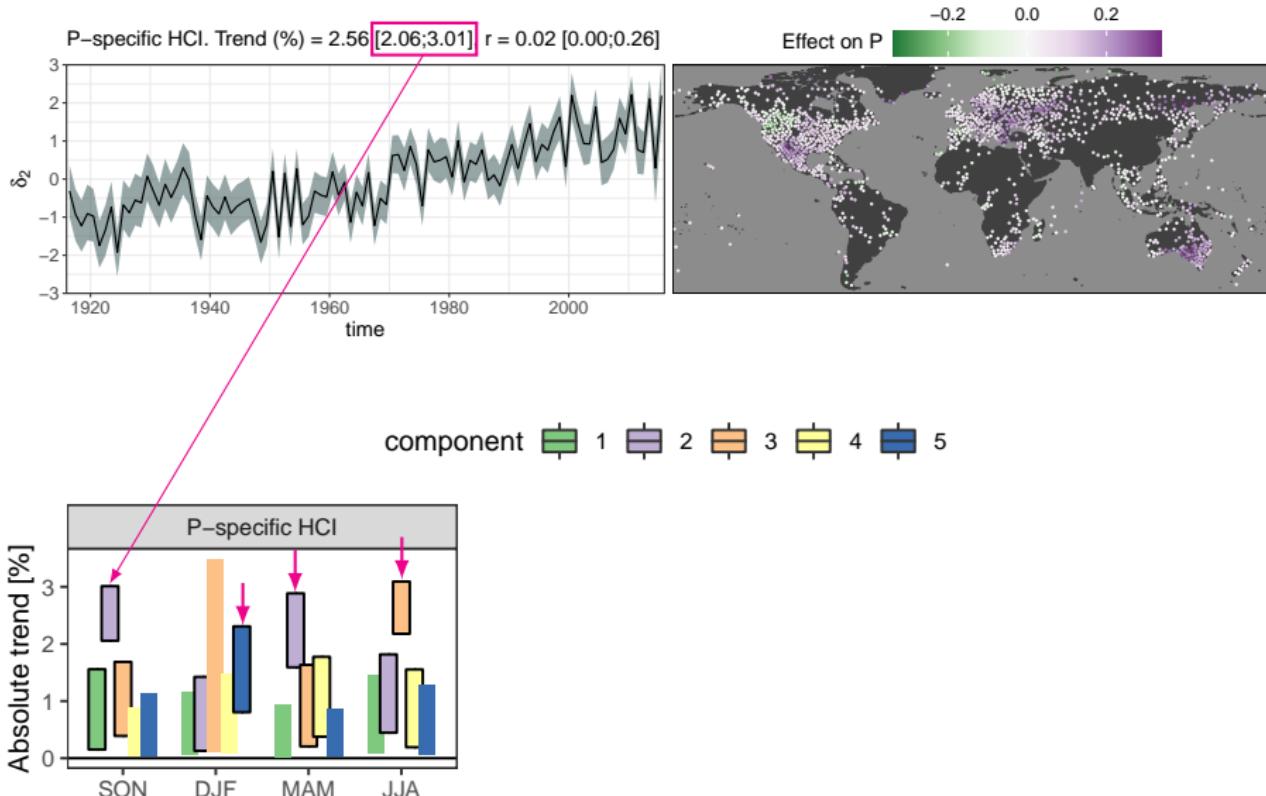
# Trends and autocorrelations



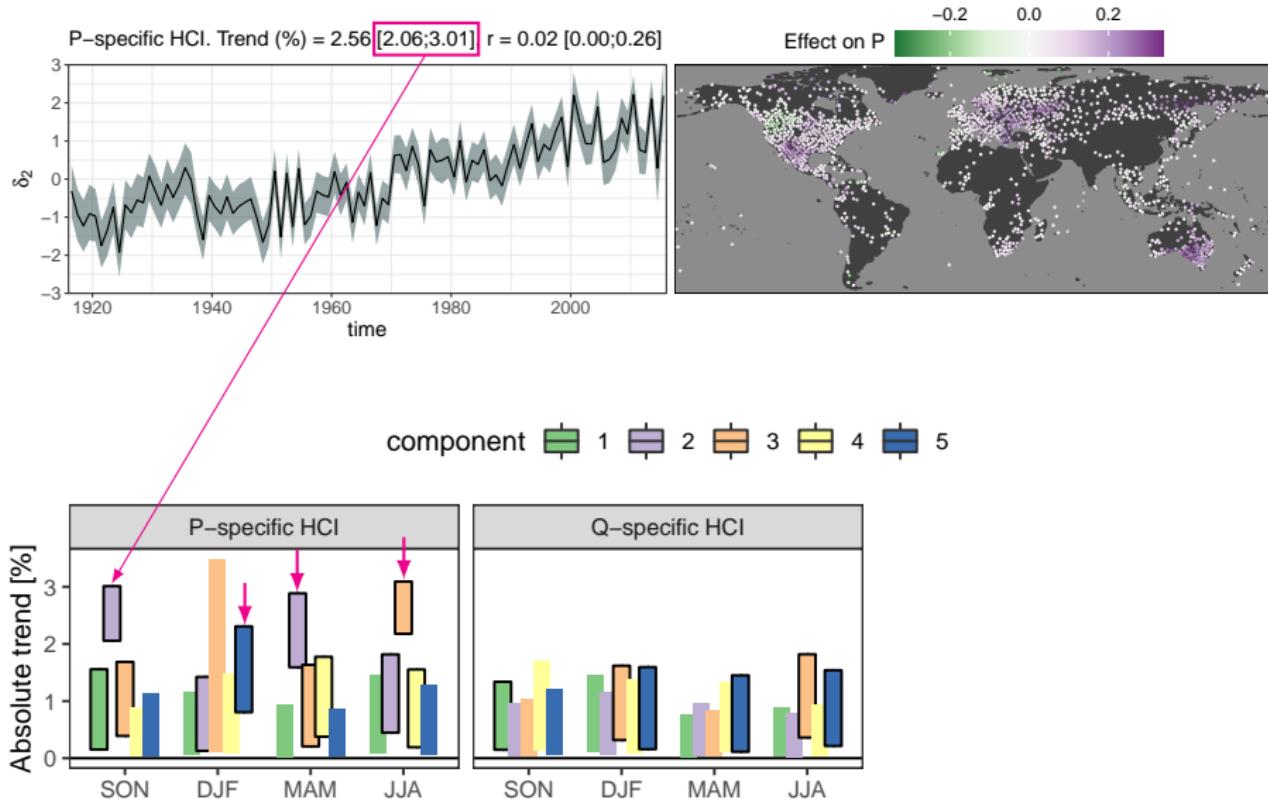
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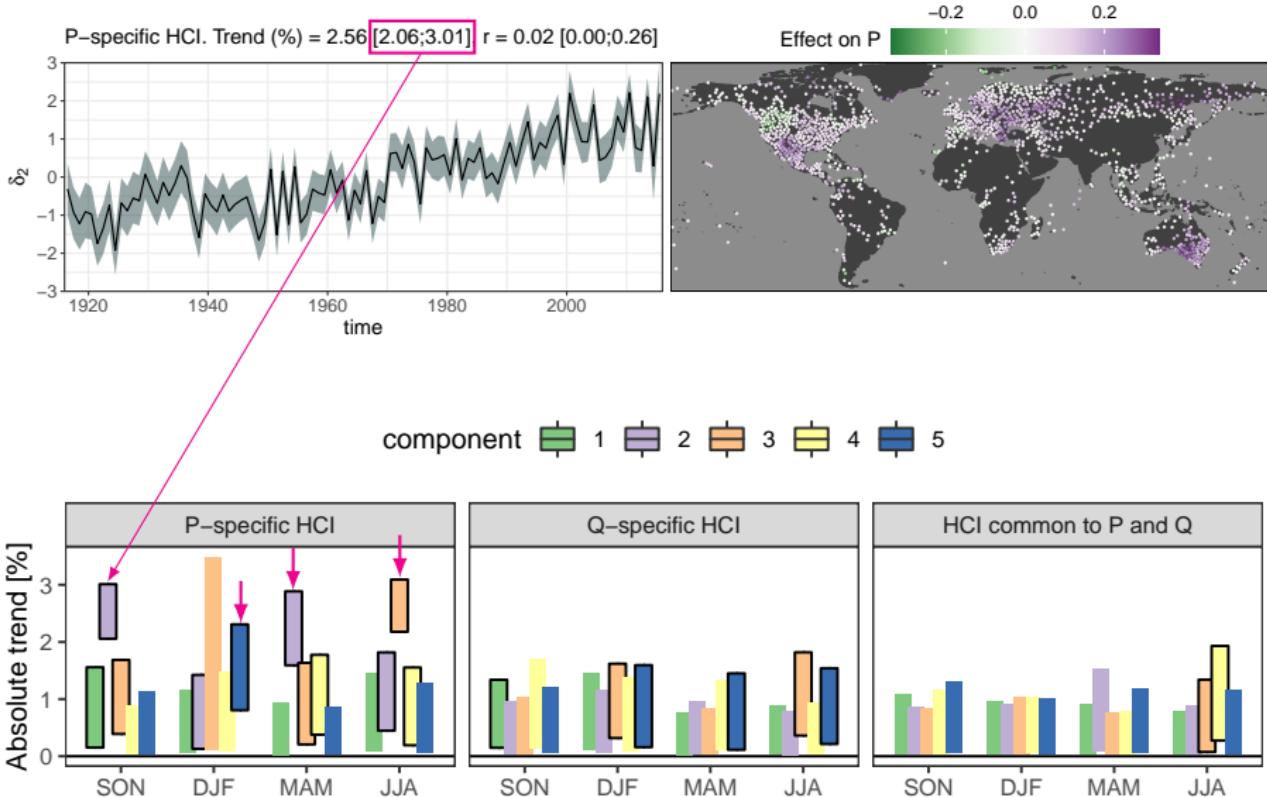
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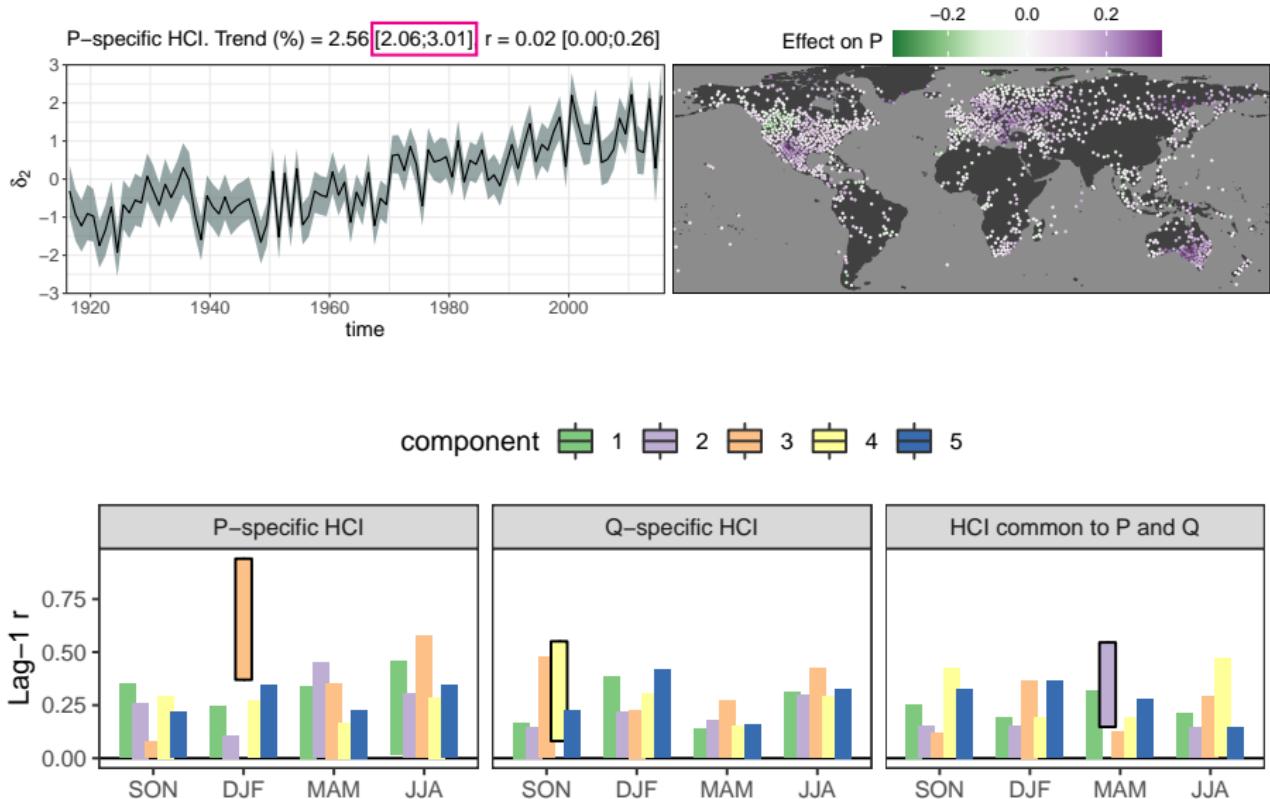
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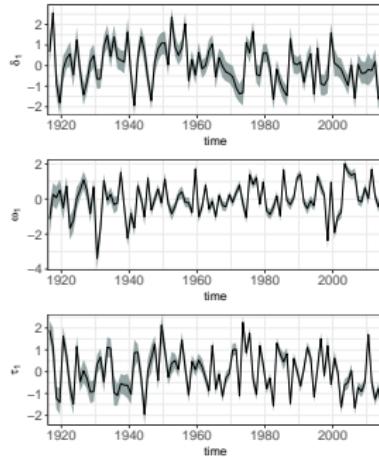


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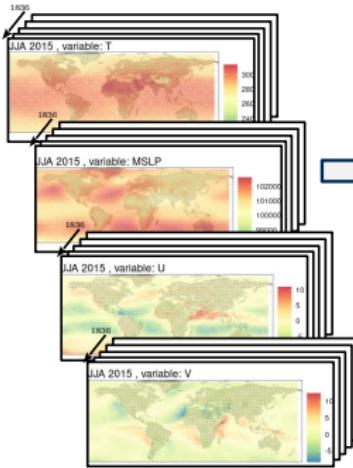
# 180-year reconstruction

## Hidden Climate Indices

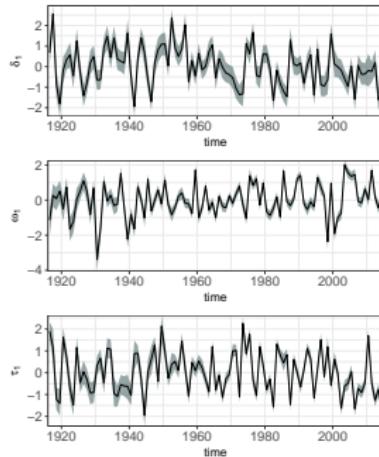


# 180-year reconstruction

## Climate predictors

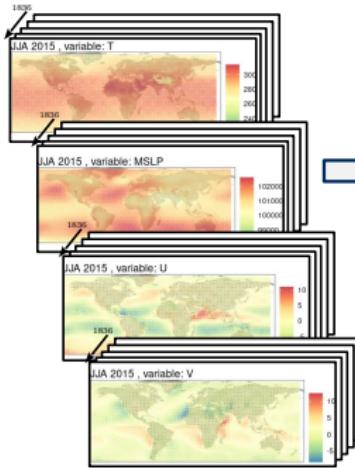


## Hidden Climate Indices

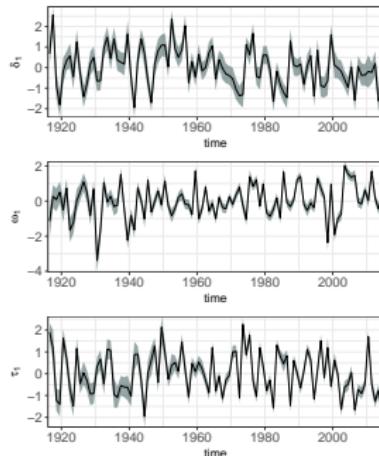


# 180-year reconstruction

## Climate predictors



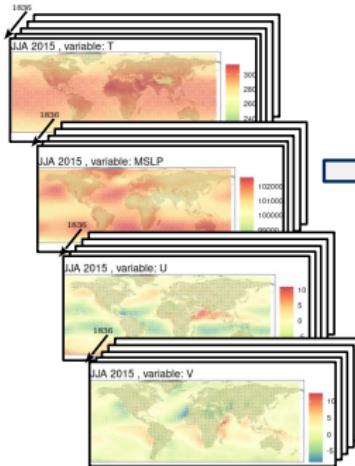
## Hidden Climate Indices



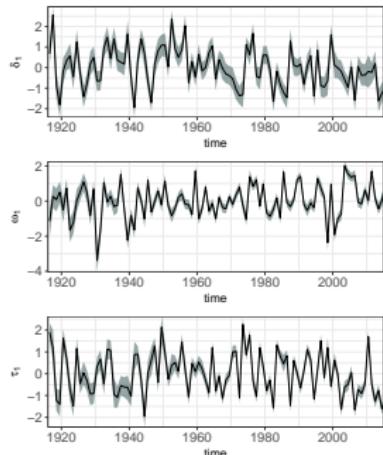
- Using 20CRv3, reconstruction of HCIs from 1836
- Hydro-extreme probability maps from 1836

# 180-year reconstruction

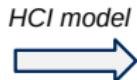
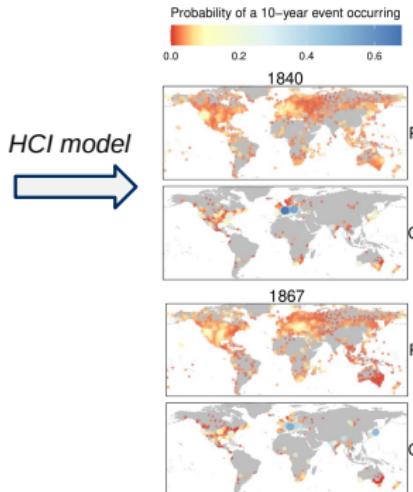
Climate predictors



Hidden Climate Indices



$\text{Pr}[\text{hydro-xtreme}]$



- Using 20CRv3, reconstruction of HCIs from 1836
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# Thank you!

Renard & Thyer (2019). Revealing Hidden Climate Indices from the Occurrence of Hydrologic Extremes. *Water Resources Research*.

Renard et al. (2021). A Hidden Climate Indices Modeling Framework for Multi-Variable Space-Time Data. *Water Resources Research*.



<https://globxblog.inrae.fr/>



<https://github.com/STooDs-tools>



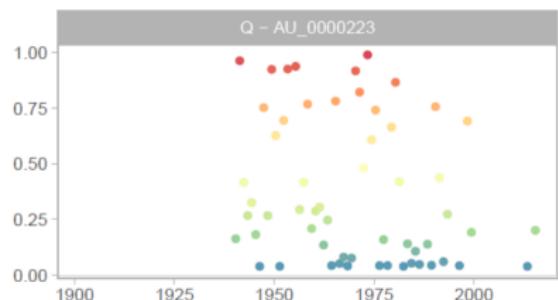
This project has received funding from the European Union's Horizon 2020 research and innovation programme under the Marie Skłodowska-Curie grant agreement No 835496

# Analyzed variables

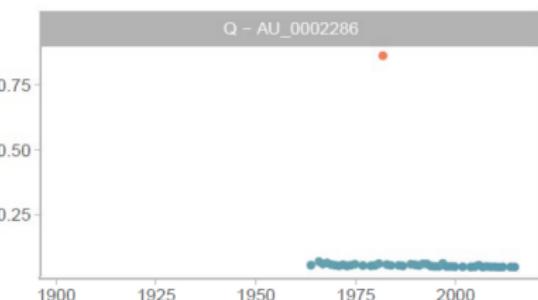
Non-exceedance probability ( $\Leftrightarrow$  return period) of the largest event of the season

**Example:** Maximum streamflow in December-January-February for 2 Australian stations

Barker Creek at Brooklands (QLD)



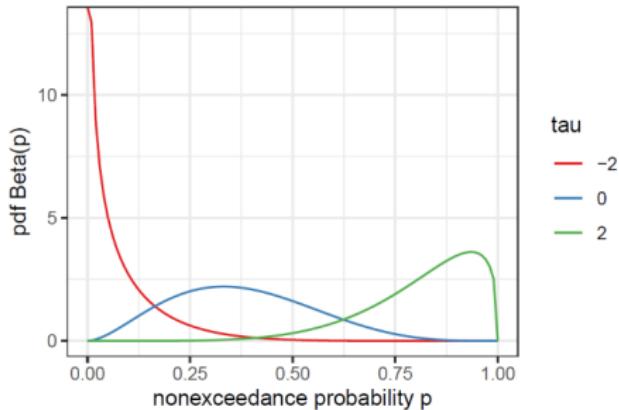
Clarke Brooke at Hillview Farm (WA)



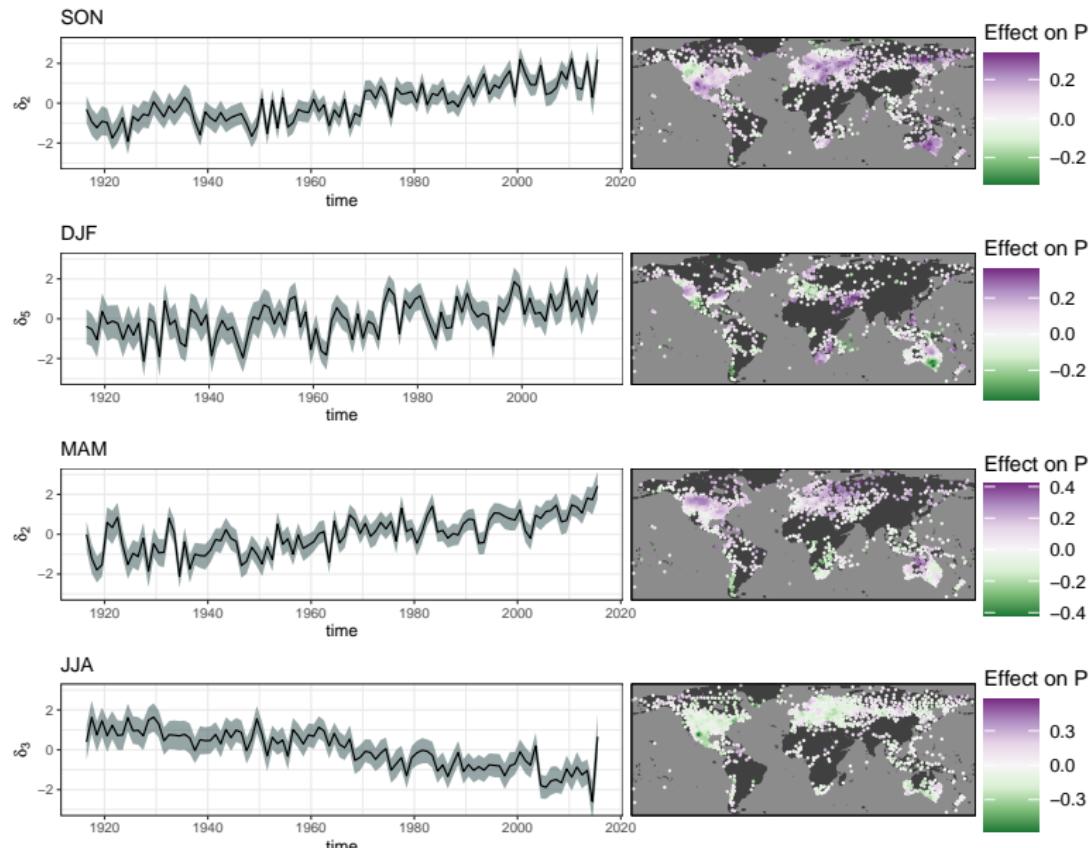
# Model

Beta distribution reparameterized in terms of mean  $\mu$  and precision  $\gamma$

$$\begin{cases} P(s, t) \sim \text{Beta}(\mu_P(s, t), \gamma_P(s)) ; Q(s, t) \sim \text{Beta}(\mu_Q(s, t), \gamma_Q(s)) \\ \text{logit}(\mu_P(s, t)) = \lambda_{P,0}(s) + \sum_{k=1}^K \lambda_{P,k}(s) \tau_k(t) + \sum_{k=1}^K \pi_k(s) \delta_k(t) \\ \text{logit}(\mu_Q(s, t)) = \lambda_{Q,0}(s) + \sum_{k=1}^K \lambda_{Q,k}(s) \tau_k(t) + \sum_{k=1}^K \theta_k(s) \omega_k(t) \end{cases}$$

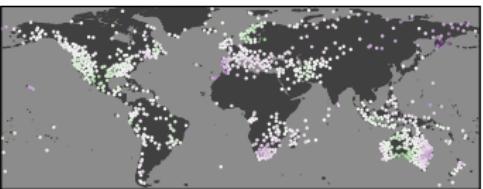
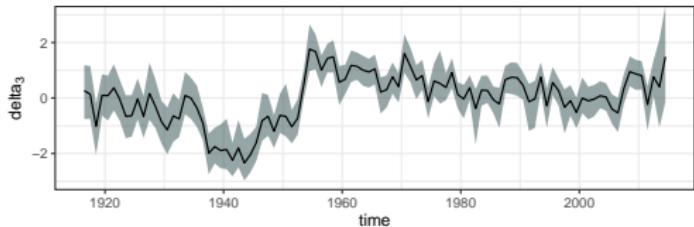


# P-specific HCIs with large trends



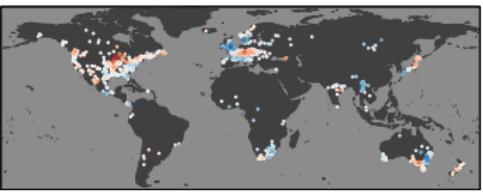
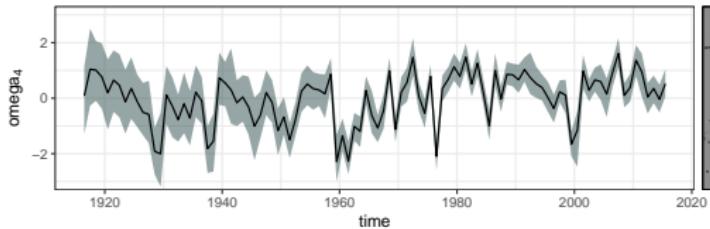
# HCLs with notable autocorrelation

DJF



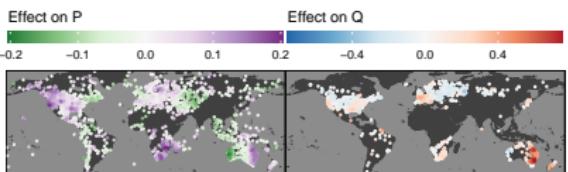
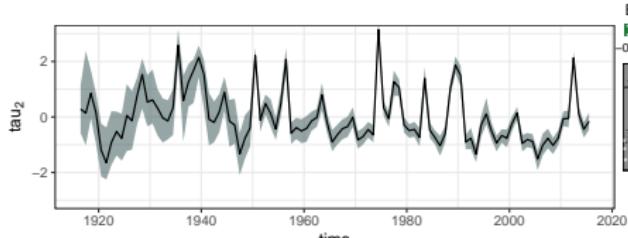
Effect on P  
0.3  
0.0  
-0.3

SON



Effect on Q  
0.25  
0.00  
-0.25

MAM



Effect on P  
-0.2 0.0 0.2  
Effect on Q  
-0.4 0.0 0.4

# Downscaling approach

## Method: inverted regression

**Step 1:**  $w(s, t)$ : climate field at time  $t$  and location  $s$

$\hat{\tau}_k(t)$ : estimated HCI's (from previous analysis)

Goal: estimate  $\psi_k(s)$ 's in:

$$w(s, t) = \psi_0(s) + \psi_1(s)\hat{\tau}_1(t) + \dots + \psi_K(s)\hat{\tau}_K(t) + \varepsilon(s, t)$$

**Step 2:**  $w(s, t^*)$ : climate field at time  $t^*$  and location  $s$

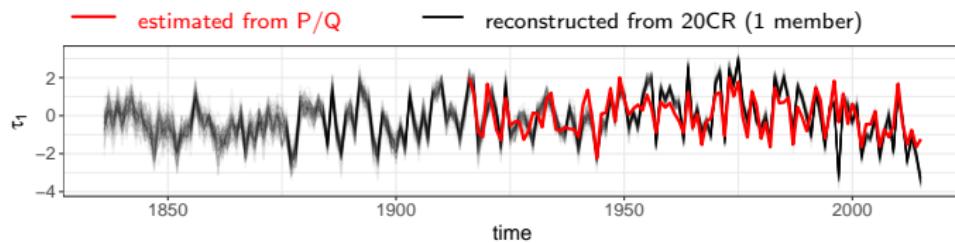
$\hat{\psi}_k(s)$ : estimated from previous step

Goal: estimate  $\tau_k(t^*)$ 's in:

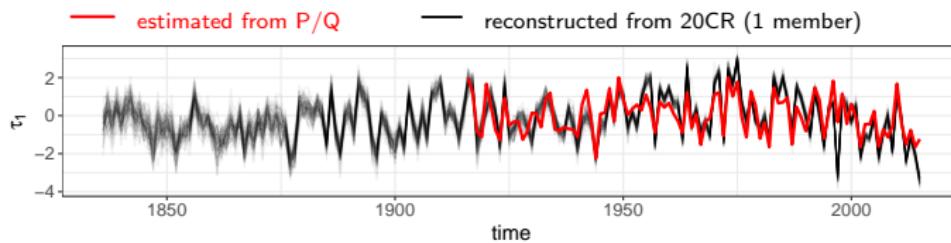
$$w(s, t^*) = \psi_0(s) + \hat{\psi}_1(s)\tau_1(t^*) + \dots + \hat{\psi}_K(s)\tau_K(t^*) + \varepsilon(s, t^*)$$

**Alternatives:** LASSO, RIDGE and other form of penalised regression, but first attempts inconclusive

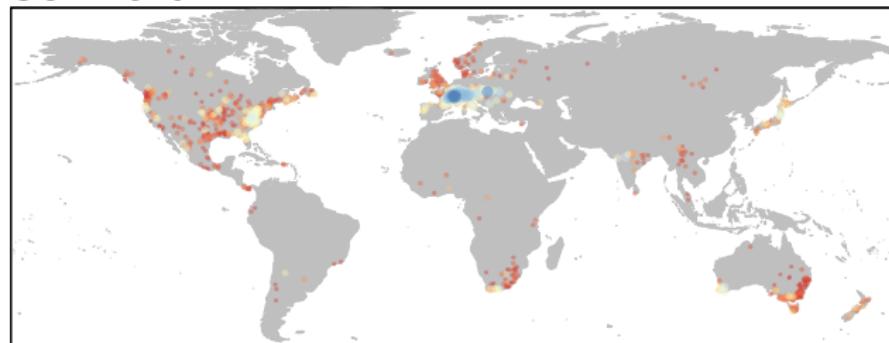
# Reconstructions from 20CRv3 (1836-2015)



# Reconstructions from 20CRv3 (1836-2015)



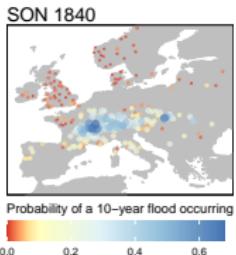
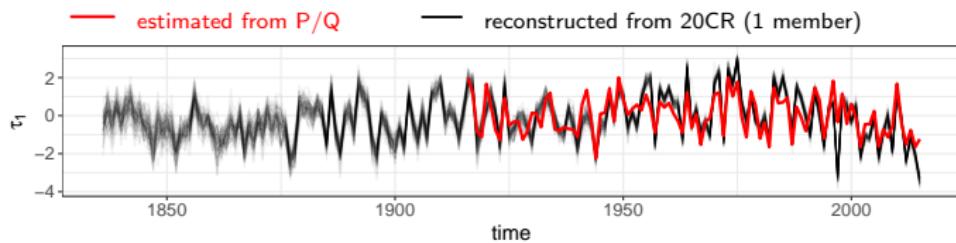
SON 1840



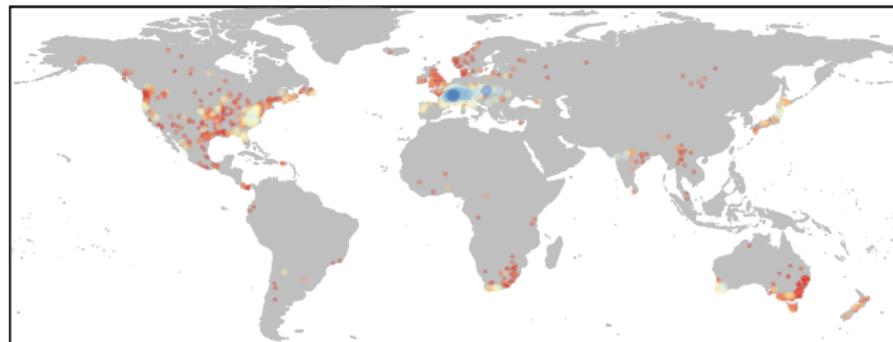
Probability of a 10-year flood occurring



# Reconstructions from 20CRv3 (1836-2015)



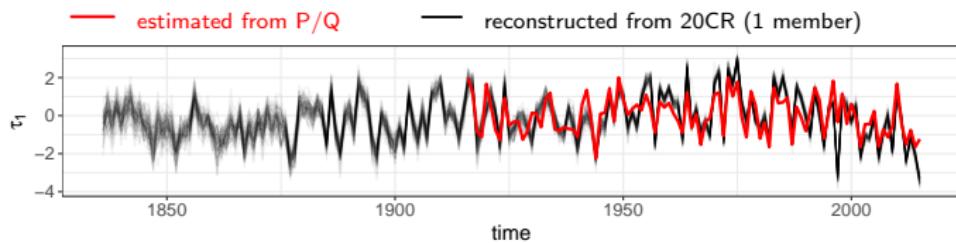
SON 1840



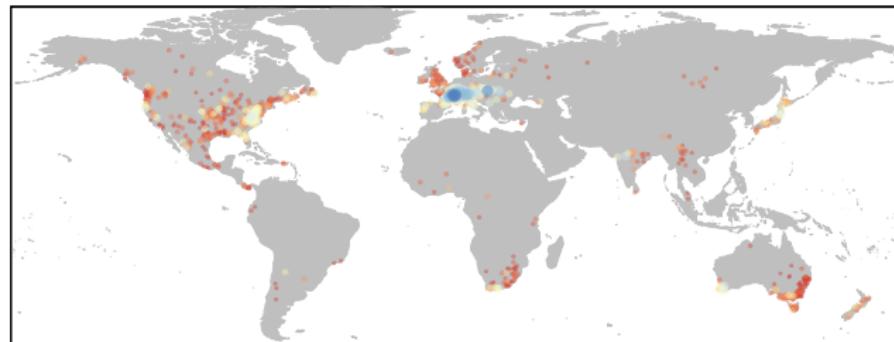
Probability of a 10-year flood occurring



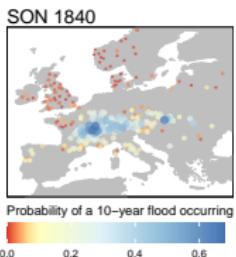
# Reconstructions from 20CRv3 (1836-2015)



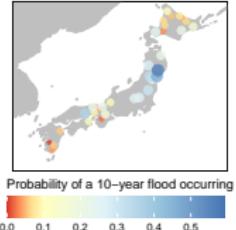
SON 1840



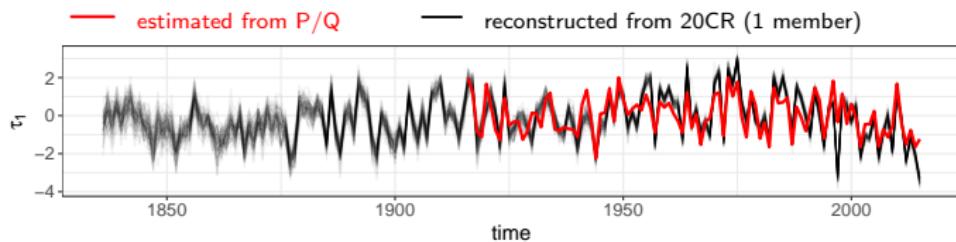
Probability of a 10-year flood occurring



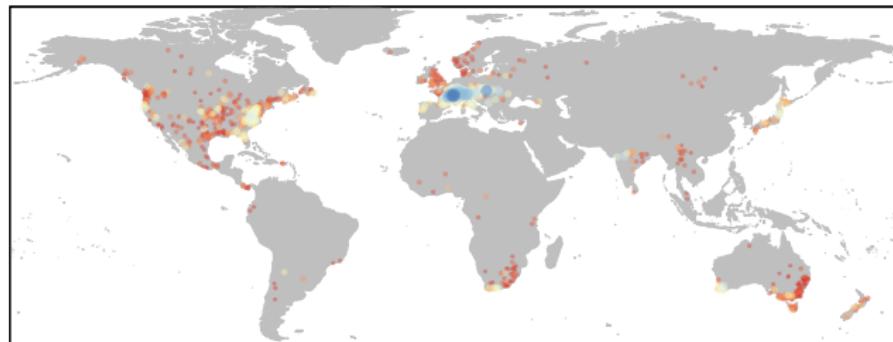
SON 1867



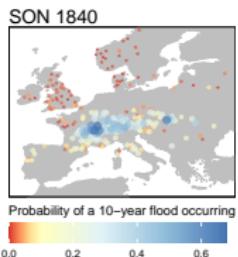
# Reconstructions from 20CRv3 (1836-2015)



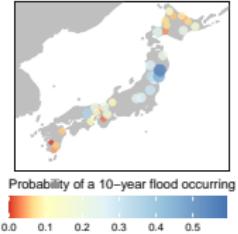
SON 1840



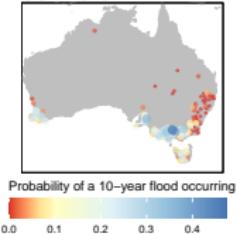
Probability of a 10-year flood occurring



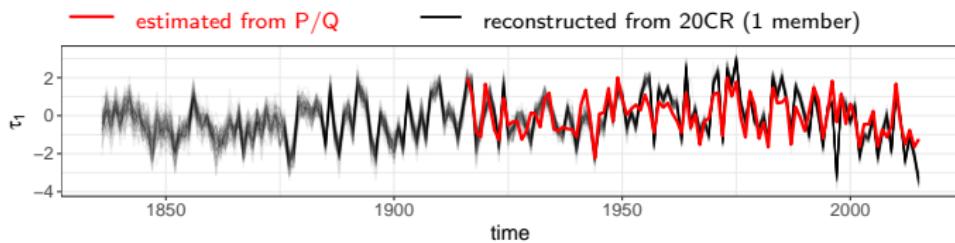
SON 1867



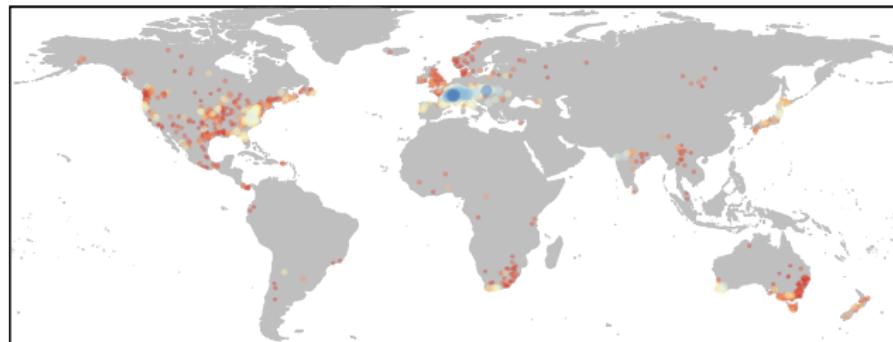
SON 1916



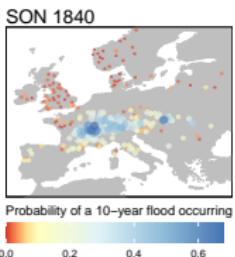
# Reconstructions from 20CRv3 (1836-2015)



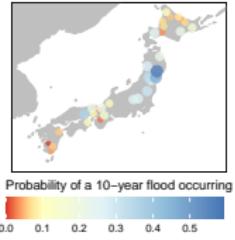
SON 1840



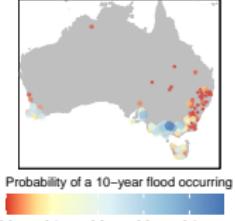
Probability of a 10-year flood occurring



SON 1867



SON 1916



→ **Reliability:** good (cross-validation); **Sharpness:** poor (P) to good (Q)