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From Lake to River: Documenting an Environmental Transition across the Jura/Knockfarril Hill Members Boundary in the Glen Torridon Region of Gale crater (Mars).

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Key Points:

- Sedimentary facies observed by *Curiosity* suggest lacustrine and fluvial environments preserved in the Glen Torridon stratigraphy.
- A paleoenvironmental transition is observed across the Jura/Knockfarril Hill members boundary, suggesting onset of fluvial conditions.
- This transition is progressive, and consistent within both the sedimentary and geochemical signals in the Glen Torridon region.

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Abstract

Between January 2019 and January 2021, the Mars Science Laboratory team explored the Glen Torridon region in Gale crater (Mars), known for its orbital detection of clay minerals. Mastcam, MAHLI and ChemCam data are used in an integrated sedimentological and geochemical study to characterize the Jura member of the upper Murray formation and the Knockfarril Hill member of the overlying Carolyn Shoemaker formation in northern Glen Torridon. The studied strata show a progressive transition represented by interfingering beds of fine-grained, recessive mudstones of the Jura member and coarser-grained, cross-stratified sandstones attributed to the Knockfarril Hill member. Whereas the former are interpreted as lacustrine deposits, the latter are interpreted as predominantly fluvial deposits. The geochemical composition seen by the ChemCam instrument show K₂O-rich mudstones (~1-2 wt.%) vs MgO-rich sandstones (>6 wt.%), relative to the average composition of the underlying Murray formation. We document consistent sedimentary and geochemical datasets showing that low-energy mudstones of the Jura member are associated with the K-rich endmember, and that high-energy cross-stratified sandstones of the Knockfarril Hill member are associated with the Mg-rich endmember, regardless of stratigraphic position. The Jura to Knockfarril Hill transition therefore marks a significant paleoenvironmental change, where a long-lived and comparatively quiescent lacustrine setting progressively changes into a more energetic fluvial setting, as a consequence of shoreline regression due to either increased sediment supply or lake-level drop.

Plain Language Summary

Between January 2019 and January 2021, the Mars Science Laboratory team explored the Glen Torridon region in Gale crater. This region is also known as the “clay-bearing unit” after the previous observation of strong spectral signatures of such minerals from orbit. As this indicates the presence of wetter conditions during the Martian past, this area was a main objective of the mission. There, sedimentary strata show a major change from fine-grained mudstones, characteristic of the underlying Murray formation (lake deposits), toward coarser-grained cross-stratified sandstones of the Knockfarril Hill member (river deposits). This change is also identified in the geochemical signal, by the ChemCam instrument. ChemCam data show higher magnesium and lower potassium in the coarse-grained fraction compared to the fine-grained sediment. We also observe a progressive transition with interbedding of both lithologies across the region. We interpret these strata to record a progressive transition from a quiet lake toward a more energetic lake shore to river plain, consequently to shoreline migration, due to either sedimentary fill of the basin and/or a drop in lake level.

1 Introduction

Gale crater, a ~3.7 Ga-old impact crater close to the Martian crustal dichotomy, was selected as the landing site of the Mars Science Laboratory (MSL) mission on the strength of well-expressed sedimentary successions visible from orbit, and spectral data (CRISM - Compact Reconnaissance Imaging Spectrometer for Mars) that indicated secular mineralogical changes in the strata exposed on the northern flank of Aeolis Mons (informally named Mt Sharp, Fig. 1a; Golombek *et al.*, 2012; Grotzinger *et al.*, 2012; Milliken *et al.*, 2010).

Since August 2012, the MSL rover *Curiosity* has explored the northern part of Gale crater to seek out geological evidence for ancient habitable environments (Grotzinger *et al.*, 2012). For ~9 years, the MSL team has documented a wide range of siliciclastic rocks that attest to the

presence of fluvial, lacustrine and aeolian depositional settings in the Mt Sharp sedimentary succession (e.g., Banham *et al.*, 2018; Bristow *et al.*, 2018; Edgar *et al.*, 2020; Grotzinger *et al.*, 2014, 2015; Le Deit *et al.*, 2016; Mangold *et al.*, 2015; Rice *et al.*, 2017; Rivera-Hernández *et al.*, 2019; Stack *et al.*, 2016; Schieber *et al.*, 2017; Stein *et al.*, 2018). Documenting past aqueous conditions is important with regard to past habitability, and particularly the identification of lacustrine sediments with their inherent tendency to concentrate and preserve organic matter within clay-bearing deposits (e.g., Farmer & DesMarais, 1999; Hays *et al.*, 2017; Summons *et al.*, 2011).

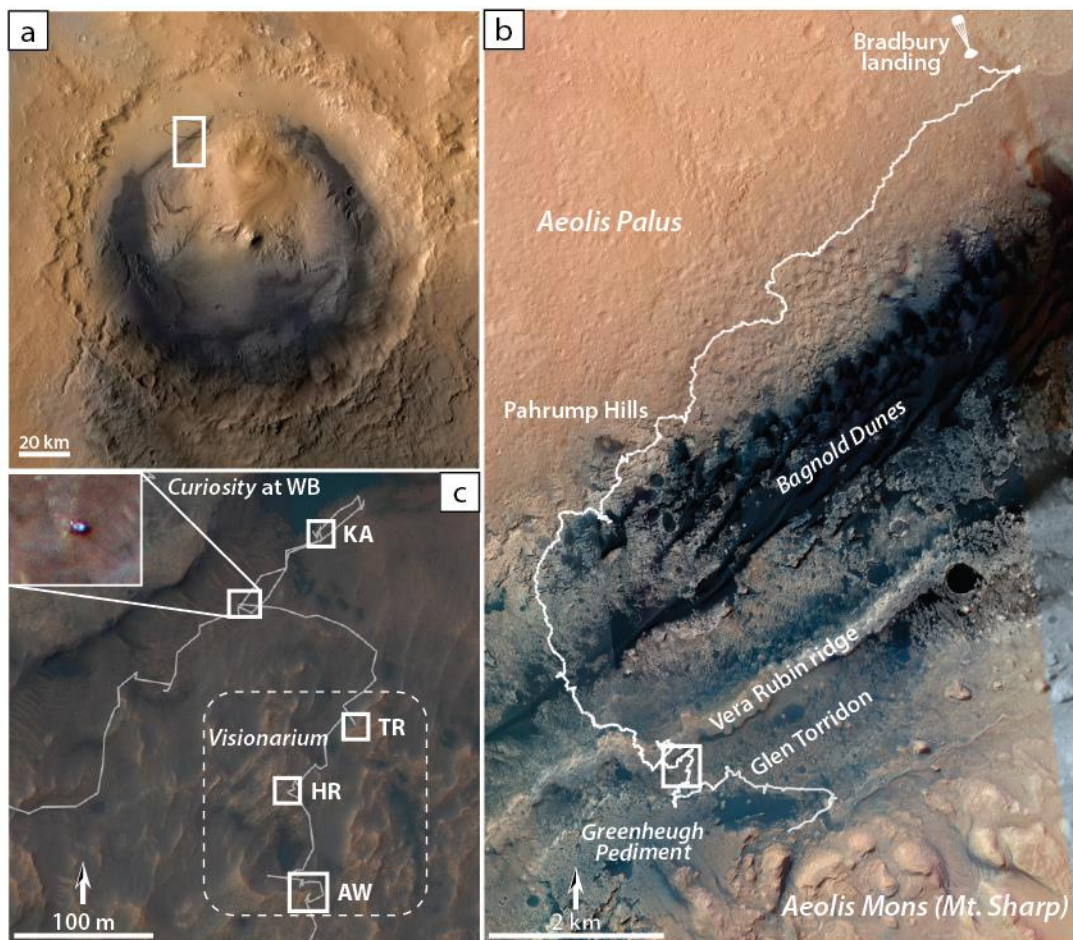


Figure 1: a) Location of the study area. The white box represents the area of operations of the Mars Science Laboratory rover *Curiosity* in northern Gale crater. b) The rover traverse (white line) from Bradbury landing and to Glen Torridon (as of June 2021). c) Study area and location of the selected outcrops along the traverse in northern Glen Torridon. The upper left inset shows the *Curiosity* rover at Woodland Bay from orbit (HiRISE image ESP_060207_1750). The dashed white line delimits the “Visionarium” area. WB: Woodland Bay, KA: Kilmarie/Aberlady drill sites, TR: Teal Ridge, HR: Harlaw rise, AW: Antonine Wall and Glen Etive drill site.

The Glen Torridon (GT) campaign (Bennett *et al.*, *this volume*) is focused on the area where the strongest clay signatures were detected from orbit (Milliken *et al.*, 2010). Specific

objectives of this campaign were lithologies and facies at GT because these strata are lateral equivalents to the Vera Rubin ridge (Fig. 1b), a prominent ridge of lacustrine strata that has undergone substantial diagenetic alteration (e.g., David *et al.*, 2020; Edgar *et al.*, 2020; Fraeman *et al.*, 2020; Frydenvang *et al.*, 2020; L'Haridon *et al.*, 2018, 2020). Because this alteration likely obliterated primary organic matter, it was of critical importance to study their potentially less modified equivalents at GT. In this study we use the instruments on *Curiosity* at several key locations in GT to document the sedimentary architecture, lithologies and facies of the exposed strata and the geochemical composition of this material as analyzed by the ChemCam instrument. Careful examination of facies and chemical changes allow us to identify marked lateral and vertical variations of depositional setting across the Jura to Knockfarril Hill transition in Glen Torridon.

2 Geological context and studied area:

Gale crater is ~155km wide (Fig. 1a) and formed ca. 3.7 ± 0.1 Ga (Le Deit *et al.*, 2013; Thomson *et al.*, 2011), close to the Noachian-Hesperian transition. It is located at 5.3°S, 137.8°E, near the Martian dichotomy boundary that juxtaposes the highly cratered southern highlands and the smoother northern lowlands. Gale crater belongs to a class of impact craters with thick sedimentary fills (e.g., Bennett & Bell, 2016; Day *et al.*, 2016; Grotzinger *et al.*, 2012; Malin & Edgett, 2000). It contains a central mound (Mt Sharp, Fig. 1a and 1b), that consists of a thick stack of sediments (~5km) that are thought to have been deposited shortly after the crater's formation and continuing into the Hesperian and possibly the Amazonian (Le Deit *et al.*, 2013; Palucis *et al.*, 2014; Thomson *et al.*, 2011). Initially investigated from orbit, the mound records a varied history of ancient sedimentary processes and conditions (e.g., Cabrol *et al.*, 1999; Malin & Edgett, 2000; Milliken *et al.*, 2010, 2014; Rossi *et al.*, 2008; Thomson *et al.*, 2011). Since the arrival of *Curiosity* on August 6, 2012, the MSL team has examined the crater floor along more than 27 km's of traverse from the northern crater floor plain (Aeolis Palus, Fig. 1b) to the lower reaches of Mt Sharp (Fig. 1b). A synthetic stratigraphic column based on vertical elevation has been prepared by the MSL Sedimentology and Stratigraphy Working Group (Fig. 2) to provide a representation of the vertical succession of sedimentary rocks encountered by the rover at specific elevations along the traverse. This column therefore does not indicate a true stratigraphic succession at a given specific location but proffers a representative log of the strata that form Mt. Sharp over more than 10 km of lateral distance, under the assumption of a nearly horizontal stacking pattern (cf. Fedo *et al.*, *this volume*).

The southward traverse from the Bradbury landing site to the foothills of Mt Sharp (sol 750; Fig. 1b) largely crossed fluvial conglomerates and sandstones (e.g., Edgar *et al.*, 2018; Mangold *et al.*, 2016; Rice *et al.*, 2017; Stack *et al.*, 2016; Vasavada *et al.*, 2014) as well as lacustrine deposits (e.g., Grotzinger *et al.*, 2014, 2015; Schieber *et al.*, 2017). Arriving at the Pahrump Hills region (Fig. 1b), the rover encountered the first deposits that are clearly part of the Mt Sharp strata. These rocks form the base of the Murray formation and are interpreted as lacustrine on the basis of fine-grained laminated mudstones, similar to those observed on Earth (e.g., Edgar *et al.*, 2020; Grotzinger *et al.*, 2015; Kah *et al.*, 2018; Rivera-Hernández *et al.*, 2019; Schieber *et al.*, 2015; Stack *et al.*, 2019). At present, interpreted sedimentary environments range from a succession dominated by hyperpycnal flows (Stack *et al.*, 2019) to one that is

characterized by saline and hypersaline waters and evaporitic conditions (Kah *et al.*, 2018; Schieber *et al.*, 2015).

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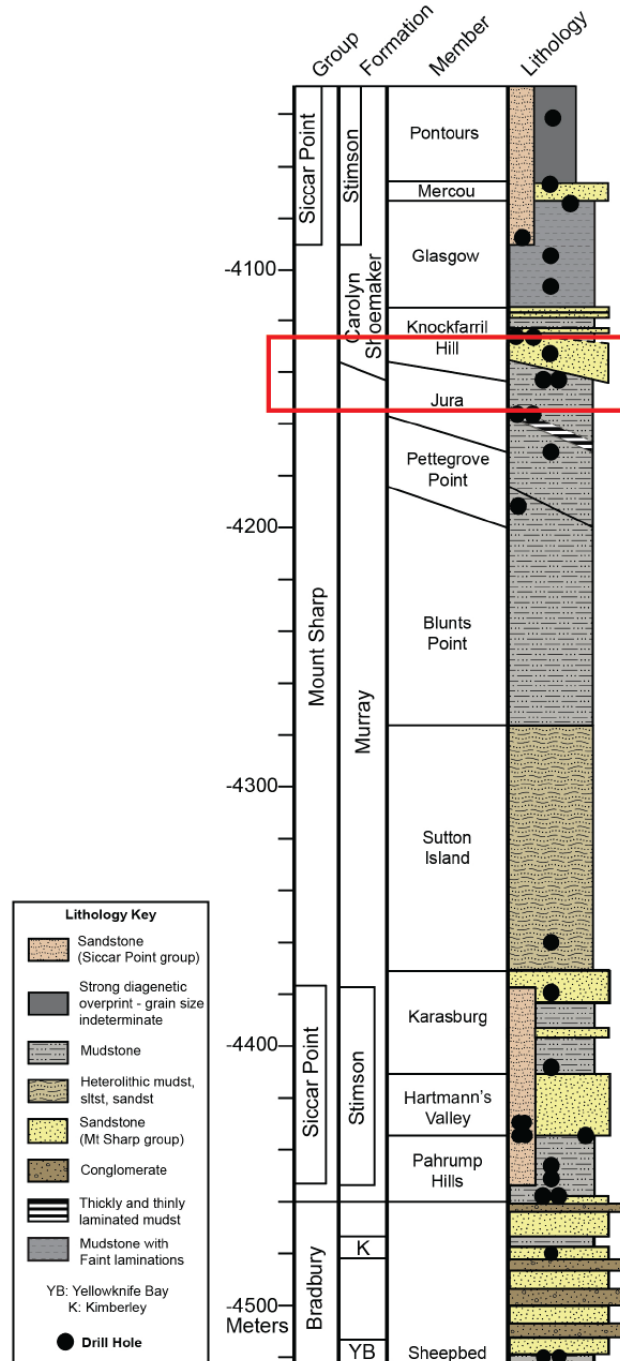


Figure 2: Synthetic stratigraphic column of the region explored by *Curiosity* (after Fedo *et al.*, *this volume*). The column illustrates the variety of deposits encountered by the rover at the elevation range they were observed, from coarse fluvio-deltaic sandstones to conglomerates in the Bradbury Group, to finer mudstones of lacustrine origin in the Murray formation and to the fluvial sandstones of the Carolyn Shoemaker formation. The aeolian sandstones of the Stimson

formation unconformably overlies the Mount Sharp group at several locations along the traverse. The red box marks the studied interval across the Jura - Knockfarril Hill members boundary. YB: Yellowknife Bay formation., K: Kimberley formation.

After leaving the Pahrump Hills region (~sol 950 to 1800), the rover traversed a thick succession of Murray formation strata (350+ m elevation change, Fig. 2) up to the base of the Vera Rubin ridge (VRR). Being of considerable variability, these strata are currently considered a reflection of dominantly lacustrine deposition and highly variable lake levels (Edgar *et al.*, 2020; Fedo *et al.*, 2017; Schieber *et al.*, 2017; Stein *et al.*, 2018). The interval consists largely of mudstones of the Pahrump Hills, Karasburg, Sutton Island and Blunts Point members (Fig. 2). Coarser-grained strata (sandstones) have been described from the Hartmann's Valley and the Pettegrove Point members (Fig. 2; e.g., Edgar *et al.*, 2020), and may indicate episodic changes in local depositional settings, possibly due to aeolian or fluvial incursions in a lake margin setting (e.g., Edgar *et al.*, 2020; Gwizd *et al.*, 2018, 2019). Such episodic but important shifts in the local depositional conditions probably occurred at multiple points in Murray formation history, as indicated by lowstand deposits (Schieber *et al.*, 2017), desiccation cracks (Stein *et al.*, 2018), and hypersaline interludes (Rapin *et al.*, 2019).

Cross-stratified aeolian sandstones of the Stimson formation (Siccar Point group, Fig. 2; Banham *et al.*, 2018, 2021, *this volume*; Bedford *et al.*, 2020, 2022) have been observed to unconformably overlie the Murray and Carolyn Shoemaker formations in multiple places and elevations (as denoted by the two non-continuous placements illustrated on the column in Figure 2). They are remnants of an ancient aeolian sediment blanket, comparable to the current-day Bagnold active dune system (Fig. 1b). The Stimson formation is thought to be much younger than the Mount Sharp group as a whole (Banham *et al.*, 2018, Fedo *et al.*, 2020, *this volume*), significantly post-dating the wet and fluvio-lacustrine history of Gale, because it is also observed unconformably overlying the Carolyn Shoemaker formation strata in the Greenheugh Pediment area, which is located 300 meters higher in the stratigraphic succession than its first unconformable iteration encountered at Pahrump Hill (Figs. 1b and 2; Banham *et al.*, 2021, *this volume*, Bedford *et al.*, 2022). The Stimson formation does not crop out in the study area.

From sols ~1800 to ~2300, the rover traversed the Vera Rubin ridge, a prominent topographic feature associated with hematite detection in orbital spectra (e.g., Milliken *et al.*, 2010). VRR strata largely consist of mudstones with local intercalations of very fine- to fine-grained sandstones, and include the Blunts Point, Pettegrove Point and Jura members of the Murray formation. Considered to record lacustrine conditions (Edgar *et al.*, 2020), VRR strata show a complex diagenetic history that led to the formation of various minerals including hematite due to enhanced groundwater circulation and localized pH and/or redox fronts within partly to fully lithified sediments (e.g., David *et al.*, 2020; Fraeman *et al.*, 2020; Frydenvang *et al.*, 2020; L'Haridon *et al.*, 2018, 2020).

From Sol 2300 onward, *Curiosity* commenced detailed exploration of the Glen Torridon region (Fig. 1b). Whereas the initial traverse route (Mount Sharp Ascend Route, or MSAR) was planned along a roughly NW-SE transect, it was amended several times during the campaign to allow a thorough exploration of the broader region from VRR in the North, up to the Greenheugh pediment (GP) capping unit (Fig. 1b) to the South. Several lateral deviations provided a wealth of data to examine lateral variability within GT strata, particularly in the "Visionarium" area

(Fig. 1c; Bennett *et al.*, *this volume*, Fedo *et al.*, *this volume*), a region informally named due to the exceptional “vision” (*sensu* visibility) of strata allowed by the presence of well-developed outcrops. The GT campaign extended for more than 600 sols and resulted in 10 drill samples.

The GT sedimentary record encompasses the Jura member of the Murray formation and the Knockfarril Hill and Glasgow members of the Carolyn Shoemaker formation (Fig. 2). These units are composed of mudstones (Jura member; Edgar *et al.*, 2020; Fedo *et al.*, 2020) and very fine- to medium-grained sandstones (Knockfarril Hill and Glasgow members; Fedo *et al.*, 2020, *this volume*) that are laterally discontinuous. All these deposits contain significant amounts of clay minerals (e.g., Bristow *et al.*, 2021; Cousin *et al.*, 2021; Thorpe *et al.*, 2020), confirming prior detection with orbital spectrometry (e.g., Milliken *et al.*, 2010).

In this study we focus on the transition between the Jura and Knockfarril Hill members because of the major change in lithology observed in the northern part of GT along this boundary (Figs. 1c and 2). Five outcrops from that area have been selected for this study (Fig. 1c): Woodland Bay (sols 2414–2427), the Kilmarie/Aberlady drill site (sols 2359–2410), Teal ridge (sols 2440–2447), Harlaw rise (sols 2459–2463) and the Antonine Wall outcrop and Glen Etive drill site (sols 2476–2555). Those outcrops are identified by their informal names rather than the sol which they were observed for clarity. A detailed survey of sedimentary and geochemical features was performed for each outcrop, and the list of the individual ChemCam and MAHLI targets investigated at and between those five outcrops is tabulated in the supplementary materials in Table S1.

3 Data, Materials and Methods

3.1 Instruments and products

The *Curiosity* rover is equipped with multiple cameras, allowing for a wide range of observations: the Navcam (Navigation Camera) instrument consists of two pairs of greyscale wide-angle stereo cameras (only one pair is used at a time; Maki *et al.*, 2012). The Mastcam (Mast Camera) instrument consists of one pair of full-color high-resolution cameras with filter wheels for multispectral imaging. The two cameras have different focal lengths, 34mm and 100mm for Mastcam left (ML) and right (MR), respectively (Bell *et al.*, 2017; Malin *et al.*, 2010, 2017). The MAHLI (Mars Hand Lens Imager) instrument is a full-color high-resolution micro-imager positioned at the end of the robotic arm of the rover, able to resolve coarse silt grains (12 $\mu\text{m}/\text{px}$) from a distance of ~ 2 to ~ 4 cm (Edgett *et al.*, 2012). The RMI (Remote Micro-Imager) is a subsystem of the ChemCam instrument, designed to provide high-resolution greyscale context images of the targets investigated by the ChemCam instrument (Le Mouélic *et al.*, 2015; Maurice *et al.*, 2012; Wiens *et al.*, 2012). In this study, we use various image products to assess the outcrops explored by the rover, and to characterize the exposed rocks, their structures and texture (Mastcam, MAHLI and RMI); and to compute 3D Digital Outcrop Models (DOM) of key locations and structures (Navcam, Mastcam, MAHLI).

Geochemical data are investigated using the ChemCam instrument and its Laser-Induced Breakdown Spectrometer (LIBS). This instrument allows acquisition of elemental composition of individual rock targets up to ~ 7 meters of distance (effective range usually between 2–3 meters; Maurice *et al.*, 2012; Wiens *et al.*, 2012, 2021). LIBS measurements are performed by firing a short laser pulse at the rock target to create a plasma. When the excited ions cool, the emitted light is analyzed to determine the abundance of major rock-forming oxides (SiO_2 , TiO_2 ,

Al₂O₃, FeO, MgO, CaO, Na₂O and K₂O; Clegg *et al.*, 2017; Wiens *et al.*, 2013) and select minor elements (e.g., Li, Sr, Ru; Ollila *et al.*, 2014; Payré *et al.*, 2017), with all iron content reported as FeO. In this study we focus on analysis of 36 ChemCam targets from the Woodland Bay, Teal ridge, Harlaw rise and Antonine Wall sites mentioned above (Table S1). Only bedrock analyses are considered (soils, diagenetic veins, etc. have been excluded). For this work we focus on specific oxides (SiO₂, MgO, CaO, Na₂O and K₂O) to be used as geochemical markers characteristic of the sedimentary facies. In-depth analyses of the bulk bedrock geochemistry performed in the GT region are detailed in the companion work by Dehouck *et al.* (2022) and also in Gasda *et al.* (*this volume*).

3.2 Digital Outcrop Modelling

The various scales provided by the different cameras (from several μm per pixel and up to several cm per pixel) are helpful in observing multiple different levels of detail within the sedimentary architecture and allow multi-scale interpretation of the outcrops. In order to understand sedimentary architecture and spatial association of sedimentary structures, 3D reconstructions of outcrops and specific targets were done with Structure-from-Motion photogrammetry (Ullman, 1979). This technique is particularly suited to accurately recreate the 3D shape of geological objects of varying sizes using a set of overlapping images (e.g., Caravaca *et al.*, 2020; Le Mouélic *et al.*, 2020; Tavani *et al.*, 2014; Triantafyllou *et al.*, 2019; Verhoeven, 2011; Westoby *et al.*, 2012). Whereas the photogrammetric reconstruction process is complex due to varying optical parameters of the rover's instruments (as discussed by Caravaca *et al.*, 2020), we are nevertheless able to reconstruct both outcrops and individual targets to fit our needs for close inspection and reconstruction of architectural geometry.

The DOMs were computed using Agisoft Metashape Professional software (v.1.6; Agisoft LLC, 2020), whose embedded geospatial features also allow the scaling of the models and enable us to make accurate size measurements on the DOMs. We used the CloudCompare software (v.2.10; Girardeau-Montaut, 2020) for other 3D-related tasks (validation of the scale, visualization). Several of these DOMs were integrated within a game engine to enable true-scale visualization of the structures within a Virtual Reality system (e.g., Caravaca *et al.*, 2020; Le Mouélic *et al.*, 2018, 2020). As a complement, we also used the NASA-developed Augmented Reality application OnSight (Abercrombie *et al.*, 2017). This allows users to observe positions of the scientific targets within a terrain reconstruction generated from rover data, which is helpful for contextualizing targets.

3.3 Determination of key sedimentary characteristics

Determination of the size and shape of the sedimentary structures was performed using 2D images from Mastcam (ML or MR), or DOMs produced using Navcam and Mastcam images. Scaling on the 2D Mastcam imagery varies according to target distance, but usually allows observation of sub-cm-scale details on MR images (Malin *et al.*, 2017). Accurate measurement of feature size, and allowing for their spatial orientation, was performed using scaled DOMs (cf. Caravaca *et al.*, 2020). This allowed sub-cm measurement of objects on the 3D models.

Grain-size, and fine-scale structures (lamination, diagenetic veins, etc.) were visually investigated and characterized using MAHLI and RMI high-resolution images (e.g., Mangold *et al.*, 2016; Stack *et al.*, 2019). MAHLI images were scaled from the instrument standoff distance, resulting in a pixel size varying from ~ 17 to ~ 100 $\mu\text{m}/\text{pixel}$ (for 1 to 25 cm standoff distance).

respectively; Edgett *et al.*, 2020a, 2020b). A similar scaling is performed for the RMI images, resulting in pixel sizes ranging from ~40 to ~150 $\mu\text{m}/\text{pixels}$ (Le Mouélic *et al.*, 2015). These image datasets are useful in resolving grain-size down to coarse silt. Grain-size is classified according to Wentworth (1922) and Lazar *et al.* (2015). In this study, a “mudstone” is a rock composed of more than 50% particles with a grainsize of $<62.5 \mu\text{m}$ (Lazar *et al.*, 2015). Lamina thickness was measured using MAHLI 2D images or 3D micro-DOMs derived from MAHLI images. These products allow measurement of sub-mm-scale grains and structures. The laminae thickness was preferentially measured on 3D models to take advantage of the ability to measure orthogonally to the bedding plans. An arithmetic mean of these measurements was calculated for each outcrop and target observed. A complete set of measurements available in supplementary Table S2.

4 Outcrop characterizations

4.1 Woodland Bay

The Woodland Bay outcrop was first observed during the initial eastward traverse (cf. Fig. 1c) on Sol 2357. After the drill campaign at Kilmarie/Aberlady and before entering the Visionarium area to the South (Fig. 1c), the rover stopped at this outcrop during Sols 2414-2427 to examine the meter-scale outcrop patches with alternating massive and laminated cm-scale beds (Fig. 3a). These heterolithic strata show low-angle truncation surfaces (red lines in Fig. 3b), that differentiate them from the surrounding, more massive rocks. Based on examination of close range MAHLI images, the grain size of these beds is inferred to be in the mudstone/siltstone range (locally up to very fine/fine sandstone). At Woodland Bay several bedding plane exposures show cm-scale polygonal cracks (Fig. 3a) that are not observed as distinctly elsewhere in the region (e.g., Fig. 4a).

A micro-DOM has been computed on the MAHLI target Morningside (Sol 2424; Fig. 3b). On this model, we can observe intercalation of bundles of mm-scale mudstone laminae with cm-scale erosion-resistant mudstone/siltstone beds. Localized occurrences of (very) fine-grained sandstones (e.g., yellow arrow in Fig. 3b) are observed, which can be difficult to differentiate from coarse mudstones/siltstones. In the resistant beds, layering is obscured, probably due to cementation, we do observe in the recessive laminated part that layering is wavy-non-parallel (blue lines in Fig. 3b), laminae diverge and converge forming small-sized lenticular features (green lines in Fig. 3b, green arrows in Fig. 3c), and in places show low-angle truncations (red lines in Fig. 3b).

The resistant layers show vertical jointing that does not continue into the intervening softer intervals. These are probably the side views of the polygonal fracture networks that these harder layers show on bedding planes.

The thickness of individual laminae has been measured to be $\sim 0.35 \pm 0.06 \text{ mm}$ for the resistant beds, and $\sim 0.63 \pm 0.13 \text{ mm}$ in the recessive intervals.

Collectively, these elements point at two depositional settings at Woodland Bay differing by the energy involved for the deposition. On one hand, massive mudstones beds denote a (very) low energy setting, and on the other hand, the laminated beds would represent more dynamic settings, as on Earth, this lamina geometry typically is associated with ripple lamination. This intricate interbedding of massive and laminated beds suggest that the energy of the setting varied rapidly on a local scale, with alternating lower and seemingly higher energetic conditions. This

variation in deposition could also be at the origin of the differential weathering behavior observed here.

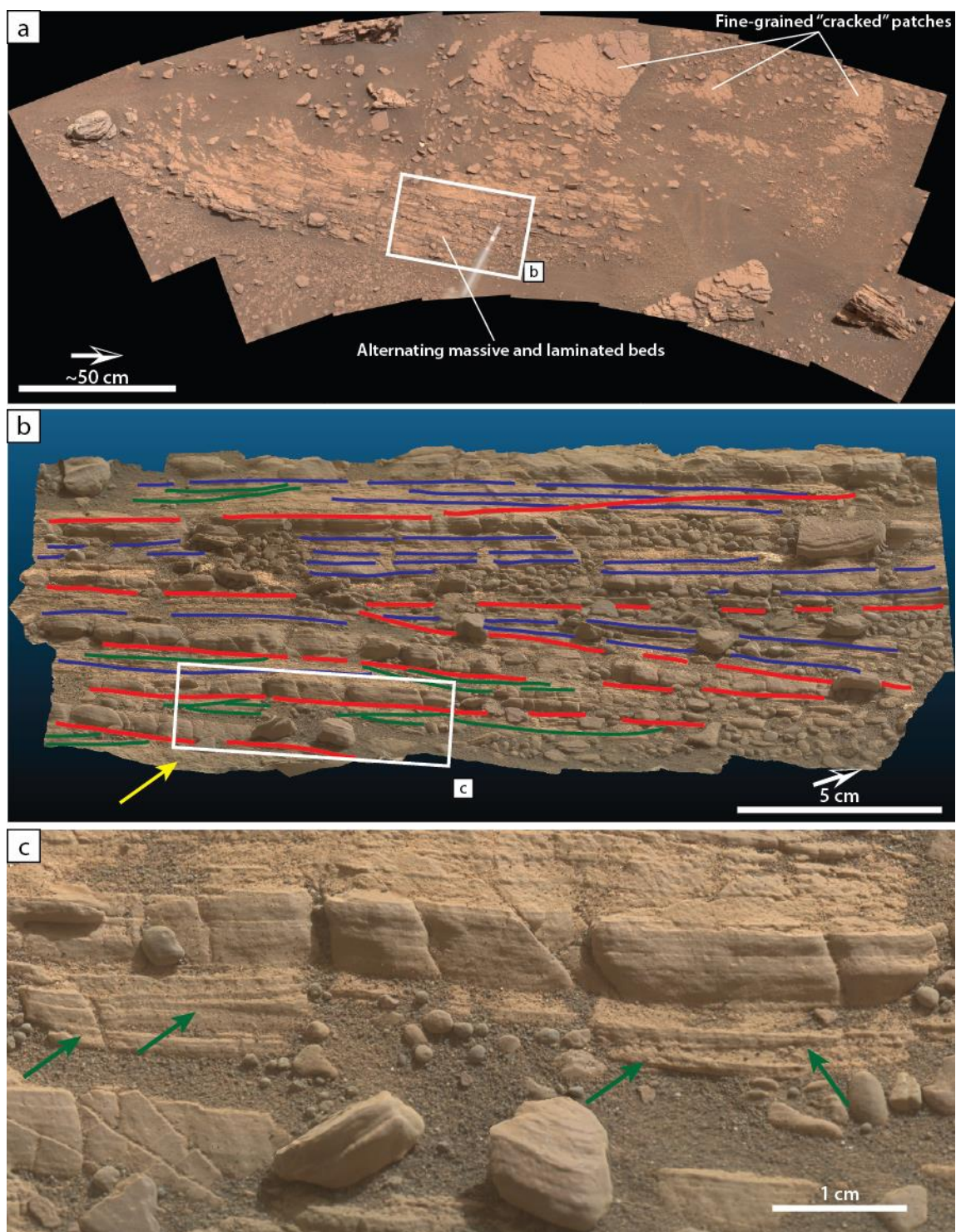


Figure 3: a) General view of the Woodland Bay outcrop, showing occurrence of alternating massive and laminated cm-scale coarse mudstones (locally up to fine sandstones), and several exposures of cm-scale polygonal cracks. Mastcam mosaic acquired on Sol 2427 (sequence mcam012854). b) View of the micro-DOM of the Morningside target, showing interstratified, thicker, resistant beds with vertical jointing, and more thinly bedded softer intervals (blue lines), as well as low-angle truncation surfaces (red lines). Green lines highlight small-sized lenticular features. The yellow arrow points to a coarse-grained part, possibly up to fine sandstone. micro-DOM was computed from 16 MAHLI images taken on Sol 2424; mesh can be viewed on Sketchfab at: <https://skfb.ly/6SE7y>). c) Detail view on the converging and diverging laminae (green arrows) forming small-sized lenticular shapes.

4.2 Kilmarie/Aberlady drill site

The first drill samples of the GT campaign (Kilmarie and Aberlady) were collected between Sols 2359 to 2410 at an outcrop situated in the northeastern part of the GT area, just below the VRR (cf. Figs. 1b and 1c, Fig. 4). This outcrop is characterized by the presence of several dm- to meter-scale slabs of erosion-resistant fine- to medium-grained sandstones (Fig. 4a). These slabs outcrop up to ~5 cm (Fig. 4b) above a rubbly- to smooth-weathering mudstone bedrock which is fractured or covered by regolith.

A DOM has been computed for the drill operations area, with specific focus on the drilled slabs (Fig. 4b). In this 3D model, we can observe that the rock slabs occur within a restricted part of the exposure that covers only ~6m² and is surrounded by smooth weathering mudstone bedrock. The slabs are likely from the same stratum, and in side view they show cm-scale cross-stratification in places (Fig. 4c). The dip direction of foresets suggests sediment transport (and current flow) in a northerly direction (Fig. 4a). Laminae have a mean thickness of $\sim 0.75 \pm 0.11$ mm and $\sim 1.09 \pm 0.14$ mm on the Kilmarie and Aberlady slabs, respectively. These coarser-grained cross-stratified slabs appear more resistant to erosion than the surrounding rubbly bedrock and host a well-developed network of light-toned veins. There are particularly well developed on the Kilmarie drill slab (Fig. 4d).

Collectively, these elements point at a depositional setting dominated by deposition of coarser-grained, cross-strata forming sandstones within condition allowing for a moderately energetic but sustained unidirectional flow.

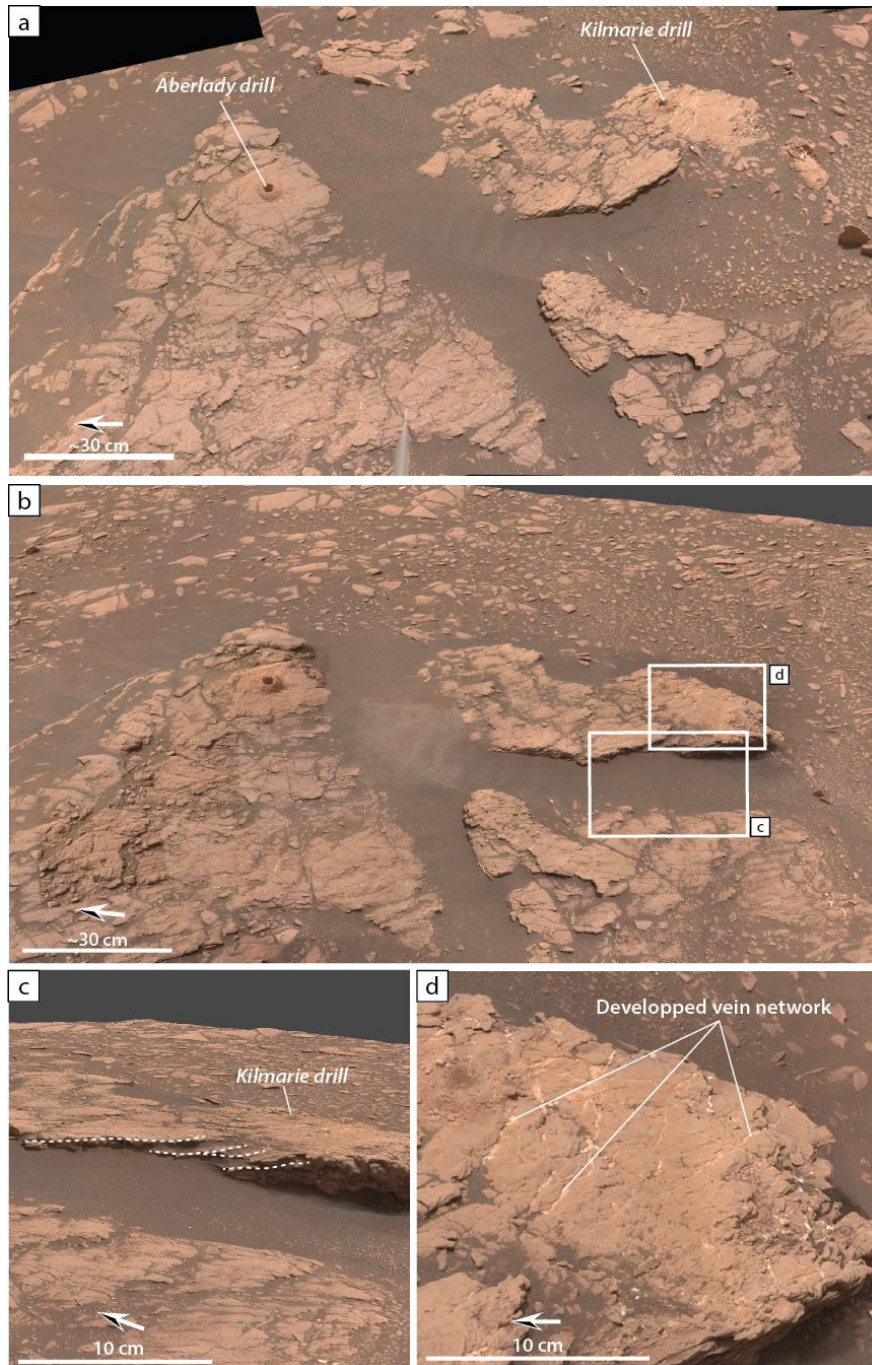


Figure 4: a) General view of the Kilmarie/Aberlady drills site, showing the position of each drillhole on their respective resistant slab, cropping out of the surrounding rubbly and smoother terrain. Composite Mastcam mosaic with images acquired on Sol 2371 (sequence mcam012571) and on Sol 2405 (sequence 12753). b) General view of the DOM of the locality computed from 178 Navcam and 732 Mastcam images taken during Sols 2365 to 2406; mesh visible on Sketchfab at: <https://skfb.ly/ooNz9>). c) Detail view of the side of the Kilmarie slab, highlighting the presence of cm-scale cross-stratifications visible in section. d) Detail of the top of the same slab with visible the laminae terminations on the flat surface. A light-toned vein network can also be observed on this surface.

4.3 Teal ridge

The Teal ridge locality was one of the first outcrops encountered in the Visionarium area, between Sols 2440 to 2447. This outcrop is one of the best exposures of the Visionarium area, and forms a mesa ~8 meters-long and ~1-meter-high (Fig. 5a). While the mesa is composed of fine- to medium-grained sandstones, the surrounding area is composed of a smoother mudstone bedrock, mostly covered by the regolith. We therefore infer a probable sharp contact between those two lithologies at the base of the Teal Ridge mesa (dashed line in Fig. 5a). The rocks composing the mesa have been sculpted out of the surrounding rocks due to their erosion resistance. The central and western part of the mesa (Fig. 5a) present a well-preserved cliff-face that allows examination and description of dm- to meter-scale cross-stratification within the sandstones. Some of the structures on the westernmost end of the mesa show a concave base characteristic of a meter-scale channel form (Fig. 5b).

For further study of sedimentary features, a DOM of the Teal Ridge outcrop has been computed (Fig. 5c). On this model we see a variety of sedimentary structures within the fine- to medium-grained sandstones, and more particularly the presence of compound crossbedding (Fig. 5d). Cross-stratification is seen in dm- to meter-scale bedsets (blue lines in Fig. 5d), as well as at the mm- to cm-scale (green lines in Fig. 5d). Whereas cross-stratification is already well visible at the macro-scale, it can also be observed in close-up images.

For the latter observation, a micro-DOM has been computed on the MAHLI target Stack_of_Glencoul (Sol 2444), situated on the eastern part of the Teal Ridge outcrop (Figs. 5c and 6). This DOM (Fig. 6b) helps to confirm the original character of small-scale cross-stratification, showing them to be an integral part of the larger scale cross-stratification (Figs. 5d and 6a). We observe mm-scale laminations with a mean thickness of 0.61 ± 0.22 mm (blue lines, Fig. 6c), organized into at least four bedsets of low angle ($<15^\circ$), cm-scale cross-strata, separated by sharp truncation surfaces (red lines, Fig. 6c). The DOM shows apparently opposing transport directions (left to right, from bottom to top in Fig. 6c), probably indicating that we observe a plane approximately perpendicular to the transport direction of 3D bedforms. On other parts of the Teal ridge outcrop some layering shows an undulating character, for example on the Beaulieu MAHLI target (Sol 2443, Fig. S1). On this target the mean lamina thickness is 0.60 ± 0.14 mm.

No mudstone interval was observed in the Teal ridge section, and most outcropping features seem to be relatively resistant to erosion. Multiple light-toned veins occur throughout the outcrop at the macro-scale (Figs. 5d and 6a) and at the micro-scale (Figs. 6b and 6c). Some of these veins crosscut the sediment layers, whereas others seem to follow them (Figs. 6b and 6c).

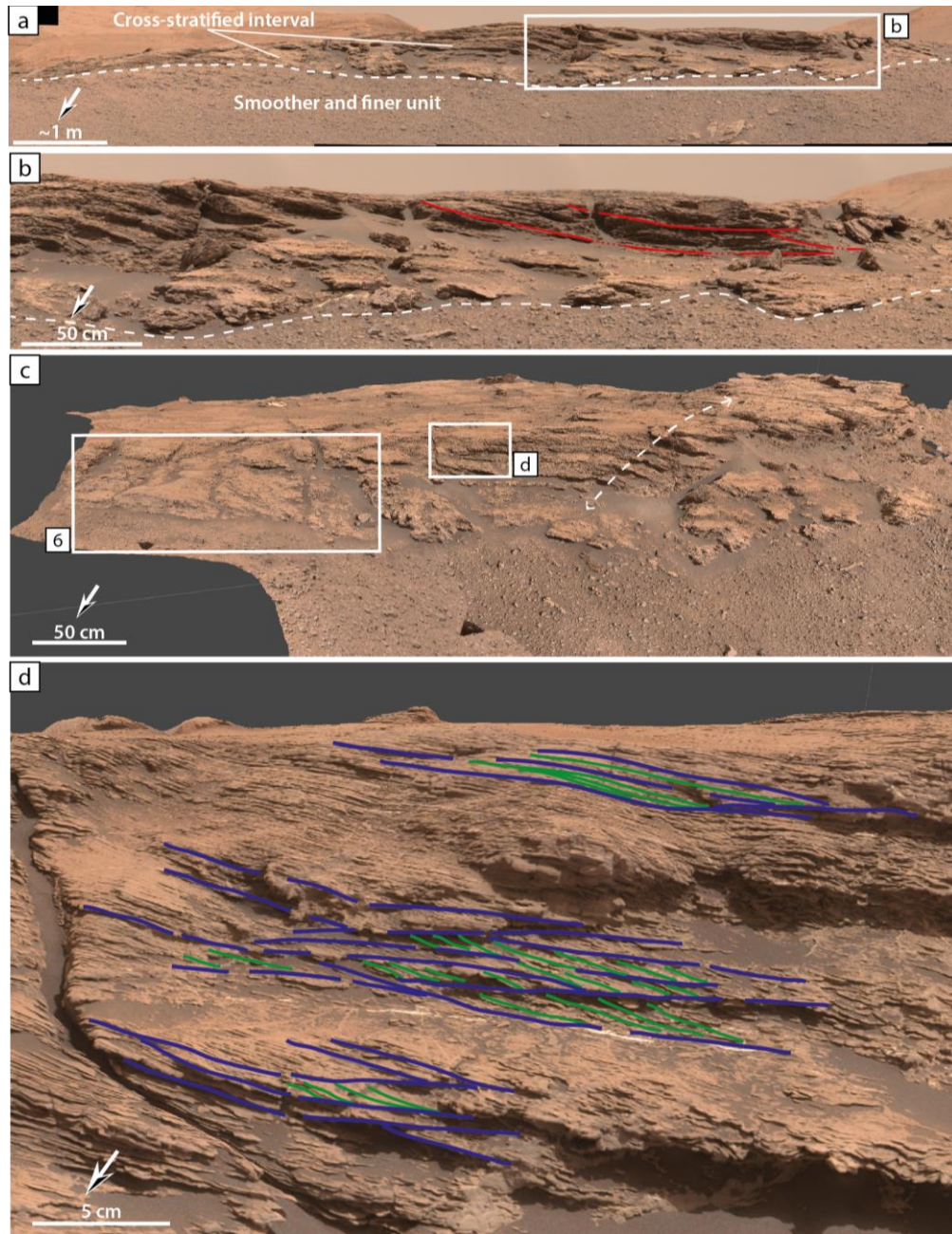


Figure 5: a) General view of the Teal ridge outcrop in the Visionarium. Mastcam mosaic acquired on Sol 2435 (sequence mcam12925). The mosaic highlights a continuous exposure of cross-stratified sandstones. The dashed line indicates a probable sharp contact of the sandstones overlying mudstone bedrock. b) Detail view of the western part of the outcrop, showing possible meter-scale channel-shapes with concave bases (red lines). c) General view of the DOM of the locality, showing the mesa structure of the 8 m long outcrop. Dashed line indicates inferred deformation axis. The DOM was computed from 214 Navcam and 540 Mastcam images taken between Sols 2435 and 2450; mesh visible on Sketchfab at: <https://skfb.ly/opv8r>. d) Detail view of the central part of the section, showing compound cross-stratification (blue lines: dm- to meter-scale stratification; green lines: mm- to cm-scale cross-strata within these larger beds).

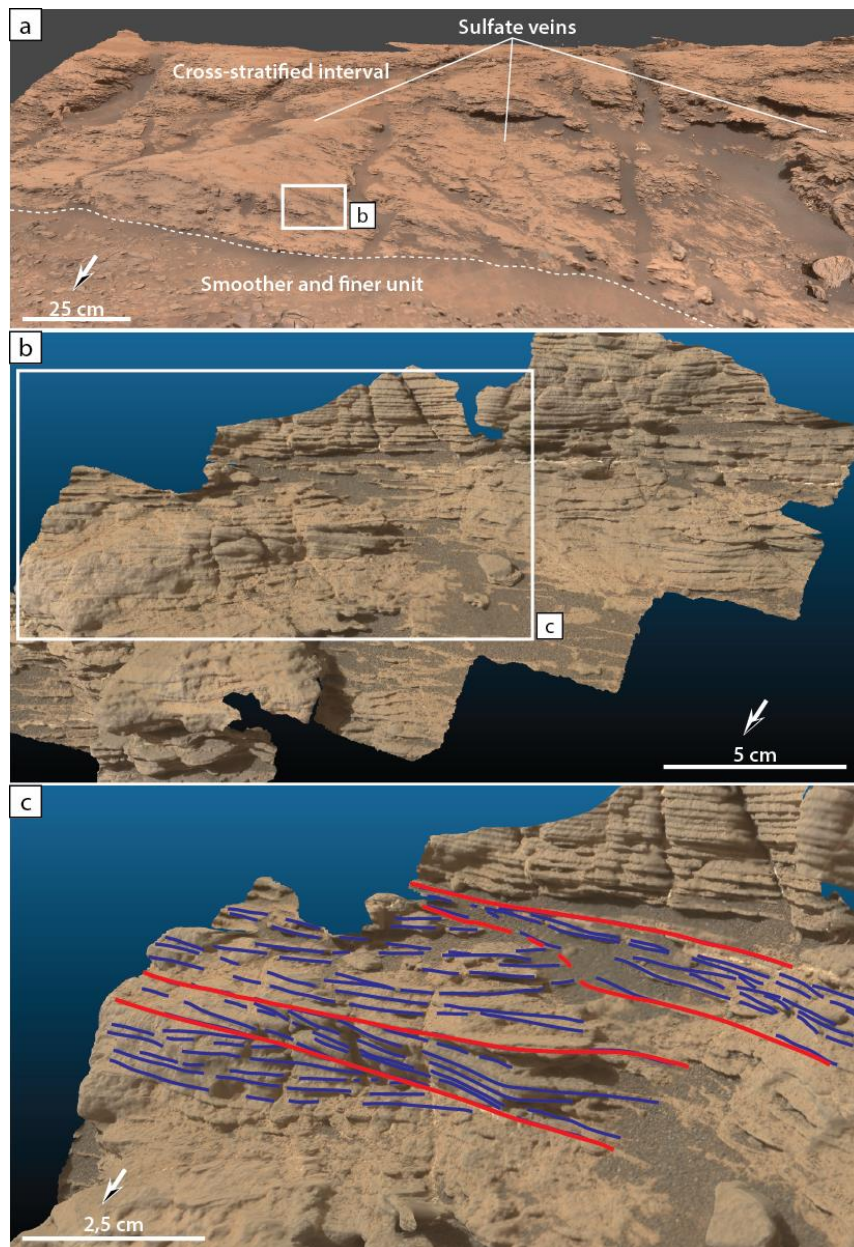


Figure 6: a) Detailed view of the eastern part of the Teal ridge outcrop on the DOM. The dashed line indicates the position of the inferred sharp contact between the mudstone bedrock and the cross-stratified sandstones that compose the mesa. b) View of the micro-DOM of the Stack_of_Glencoul target, computed from 15 MAHLI images taken on Sol 2444; mesh visible on Sketchfab at: <https://skfb.ly/6SE7z>. c) Detail view of the micro-DOM showing several bedsets of cm-scale cross-strata with varying direction of apparent sediment transport (blue lines) separated by sharp truncation surfaces (red lines).

The DOM shows that most strata outcropping at Teal ridge are gently dipping southward regardless of cross-stratification. Strata in the center of the outcrop (Figs 5a to 5c) seem to

“bend” or “roll-over” without significant change of cross-stratification, suggesting a syn- to post-depositional deformation of the structure. The cross-stratification observed at Teal ridge highlights a direction of transport mainly oriented along a SW-NE apparent axis, with a preferential direction toward the SW (e.g., cross-stratification in Fig. 5d). However, the concave-based channel-shapes (Fig. 5b), a higher order architectural element, are oriented NW-SE, suggesting that channel-scale transport occurred along a NW-SE axis. On Earth, preferential transport direction being perpendicular to the overall direction of the channels is reminiscent of the formation of point bars.

Collectively, these elements point at a depositional setting within a higher energy setting. The coarse-grained cross-stratified sandstones were probably deposited within meter-scale channelized bodies carved in the local finer-grained bedrock. This and the presence of dm- to meter-scale compound cross-stratifications highlights sustained flow at the time of deposition.

4.4 Harlaw rise

During the southward traverse of the Visionarium area, *Curiosity* stopped at the Harlaw rise locality from Sol 2459 to 2463. *Curiosity* observed several dm- to meter-scale blocks made of laminated and erosion-resistant sandstones. One of these (Fig. 7a), consisting of fine- to medium- sandstones, shows a conspicuous pattern of undulating laminations throughout most of its thickness. Several light-toned veins occur either conformable to or cross-cutting the laminations (Figs. 7a and 7b).

To better observe the sedimentary features of this block, a micro-DOM has been computed on the MAHLI target Strathdon (Sol 2462; Fig. 7b). This model presents two different parts with specific characters, highlighted by differential erosion of the outcrop (Fig. 7b). The upper part shows sub-cm-scale undulating laminations, highlighted by differential erosion of the outcrop (green lines on Fig. 7c), and the lower part, a recessive bed with faint planar laminae (Fig. 7c) whose thickness (when visible) has been measured to be $\sim 0.34 \pm 0.09$ mm. Whereas the undulations in the upper part could be due to soft sediment deformation, the fact that the undulations do not propagate into the lower layer suggests that they may actually represent climbing ripple lamination, a sedimentary feature associated with rapid sedimentation from sediment-rich currents (Allen, 1982). These laminations show a mean lamination thickness of $\sim 0.62 \pm 0.15$ mm. Finally, cross-stratifications observed on the Strathdon target (Figs. 7b and 7c) indicate transport along a roughly SW-NE apparent axis.

Collectively, these elements point at two different settings recorded at Harlaw rise. On the one hand, the fine massive mudstones represent a (very) low energy setting; on the other hand, the undulated, cross-stratified sandstones evidence a high energy unidirectional flow. Also, the sharp contact between both finer- and coarser-grained beds indicate fluctuating energy levels.

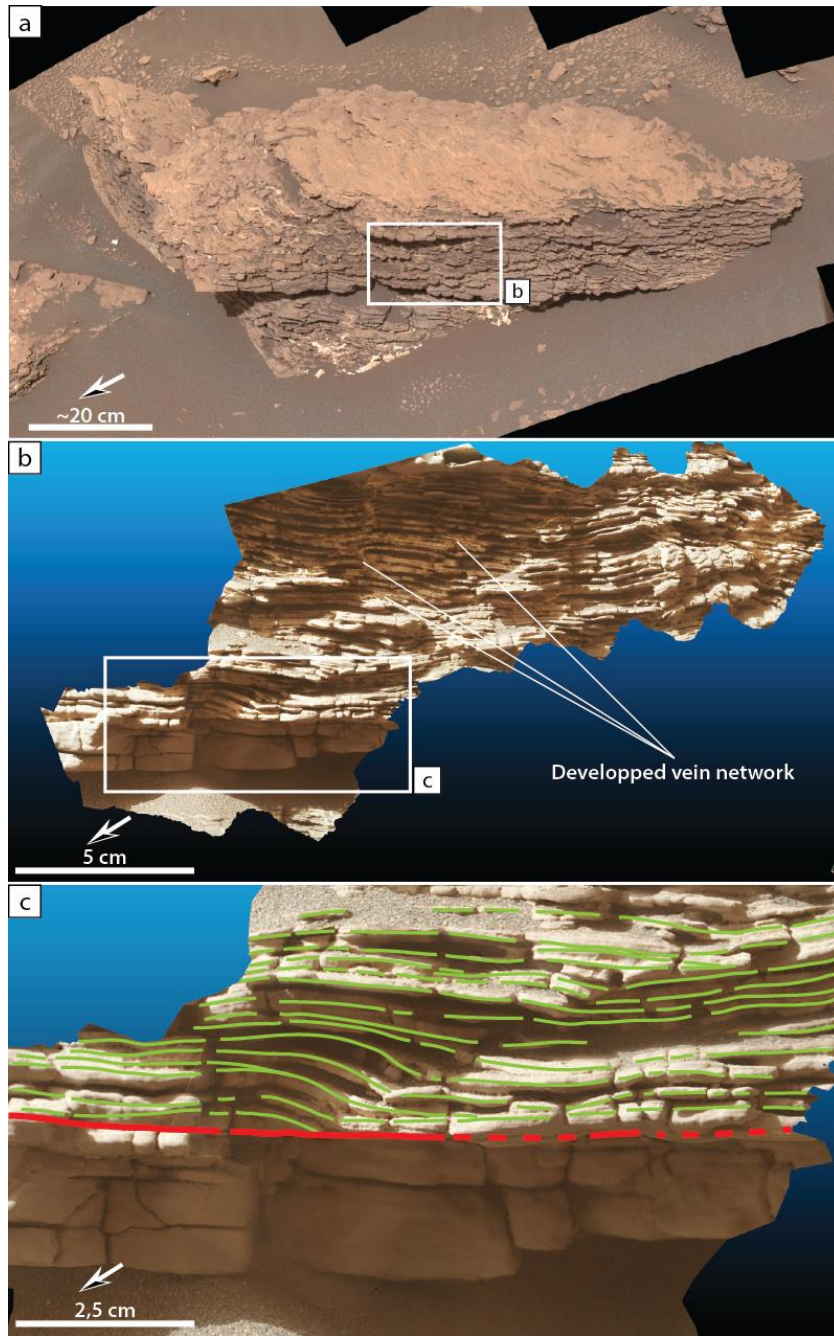


Figure 7: a) General view of a metric-scale block outcropping at the Harlaw rise locality in the Visionarium. Mastcam mosaic acquired on Sol 2461 (sequence mcam013066). The mosaic shows a prominent erosion-resistant meter-scale block cropping out of smoother material. b) View of the micro-DEM of the Strathdon target, computed from 18 MAHLI images taken on Sol 2462; mesh visible on Sketchfab at: <https://skfb.ly/6UEwZ>. A developed vein network is present, with veins in-between and cutting through the laminations. c) Detail view of the micro-DEM highlighting a sharp (possibly erosive, red line) contact between a lower (faintly laminated) fine-grained and recessive bed and an upper coarser-grained interval of erosion-resistant undulating and cross-stratified (green lines, climbing ripples?) sandstones.

4.5 Antonine Wall

The Antonine Wall outcrop is situated at the southern end of the Visionarium area and was traversed between Sols 2476 and 2555. The observations there were made in two phases; first from the North and observation of the cliff-face, and later a climb to the top of the mesa to the Glen Etive drill site. Antonine Wall is one of the biggest mesas of the entire Visionarium area, extending about 20 meters in continuous cliff-face exposure, and a height of up to 2 meters (Fig. 8a). It resembles Teal ridge, with a prominent cliff-face that shows mostly fine- to medium-grained sandstones with compound cross-stratification (Fig. 8a). In the westernmost part of the outcrop, several dm- thick cross-stratified beds also show a concave base, suggestive of possible channel-shapes, in a similar fashion to that of Teal ridge (cf. Fig. S2). The mesa is elevated above smooth weathering mudstones that surround this exposure. A sharp contact between those two lithologies is inferred at the base of the mesa (dashed line in Figs. 8a and 8b).

For closer inspection of this mesa, a DOM of the Antonine Wall outcrop has been computed (Fig. 8b). On this model, we observe cm- to dm-scale cross-stratification in the cliff-face sandstones (Fig. 8c). Compared to Teal ridge, cross-stratification is generally smaller scale and shows more variability of sedimentary transport direction (Fig. 8c, green lines toward the East, and blue line toward the West). Cross-stratification also appears in planar view at the top of the mesa, for example around the Glen Etive drill site (Fig. 8d). There, several sets of mm- to cm-scale cross- lamination is clearly visible.

This exposure also highlights the resistance to erosion of these cross-stratified sandstones, compared to the mudstone bedrock which decomposes into a multitude of clasts. However, the sandstones seem much more fractured at Glen Etive than at any other outcrop of the Visionarium area (Figs. 8b and 8c). Veins are also present on this outcrop within the cross-bedded sandstones interval and can be observed in-between and/or cross-cutting the stratification. The DOM also allows us to observe that most strata of the outcrop are dipping southward, similarly to Teal ridge's. This suggests that the overall transport (and channel orientation) is to the south, and that the observed cross stratification (Fig. 8c) reflects 3D bedforms that migrated within these channels.

Collectively, these elements point at a depositional setting similar to that of Teal ridge, albeit being slightly less energetic as hinted by the presence of smaller-scale (cm- to dm-scale) cross-strata. Here too, the coarse-grained cross-stratified and mesa-forming sandstones are likely to have been deposited within meter-scale channelized bodies carved in the local finer-grained bedrock. This and the presence of dm- to meter-scale compound cross-stratifications is highlighting sustained flow at the time of deposition.

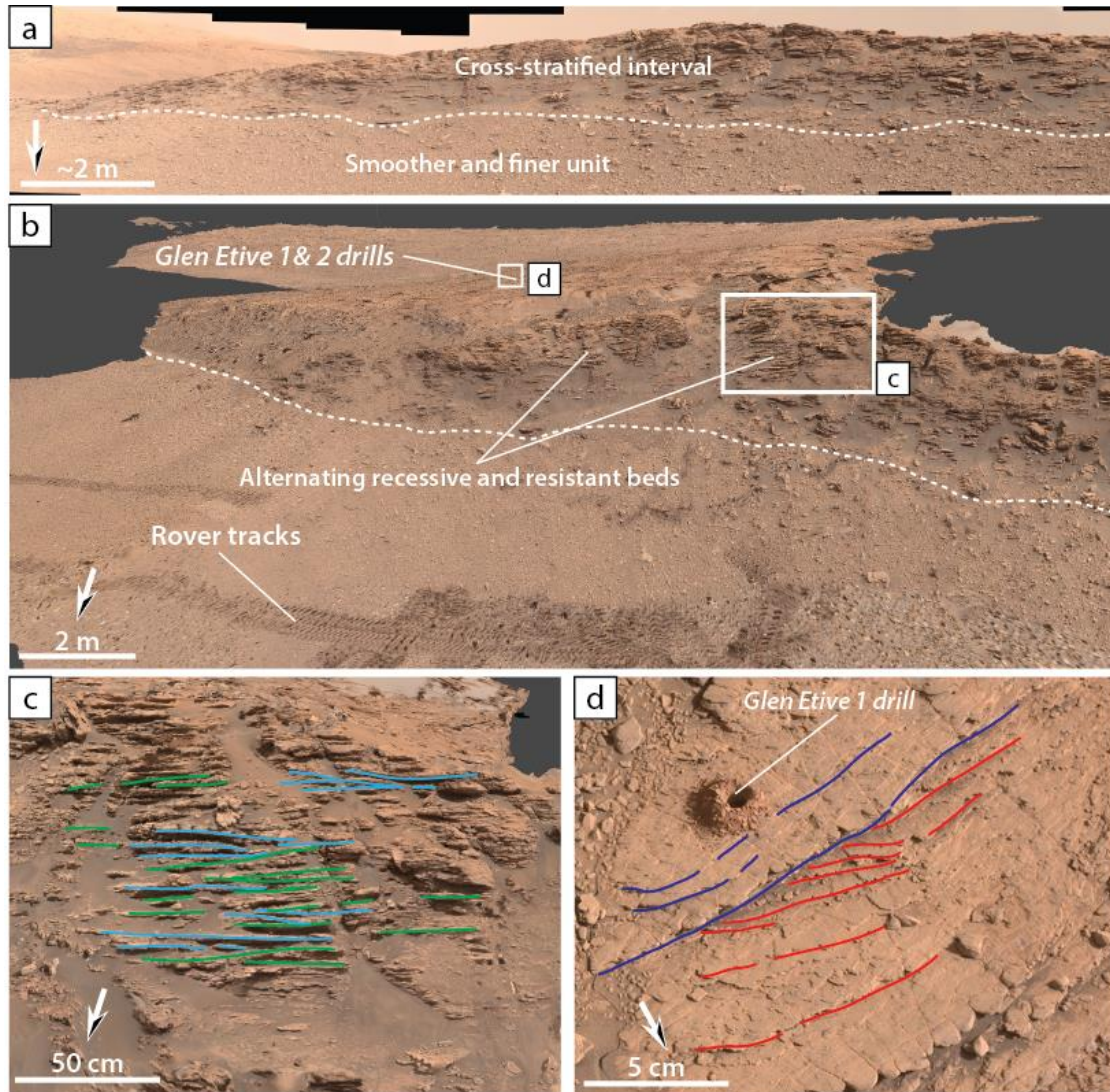


Figure 8: a) General view of Antonine Wall outcrop in the Visionarium. Mastcam mosaic acquired on Sol 2471 (sequence mcam013112). The mosaic shows a continuous exposure of cross-stratified sandstones. The dashed line evidences a probable sharp contact of the sandstones over material smoother and finer than sand-sized grains. b) General view of the DOM of the locality, showing the mesa structure of the ~15 m long outcrop. The sharp contact between smooth fine-grained material and the overlying coarser-cross-stratified interval (dashed line) is clearly visible. Rover tracks have a width of about 40 cm. The DOM was computed from 593 Navcam and 1197 Mastcam images taken between Sols 2471 and 2553; mesh visible on Sketchfab at: <https://skfb.ly/6Z7VL>. c) Detail view of the central part of Antonine Wall cliff-face, showing cross-stratification with varying directions of apparent sedimentary transport (green lines toward the East, and blue line toward the West). d) Close view on the Glen Etive 1 drill site on top of the outcrop, that shows an expression of cross-stratification in plan view. Mastcam image 2490ML0132270010904925C00_DRCL

5 Interpretation of the texture and facies

5.1 Facies and depositional settings

Differentiation of facies and sub-facies is based on grain size, multi-scale sedimentary structures and bedforms, and post-depositional features. We defined 2 distinct facies, each facies being subdivided into 2 sub-facies that are detailed in Table 1 and illustrated in Figure 9. These facies represent two different depositional settings and were consistently encountered throughout the studied area, in both the Jura and the Knockfarril Hill members.

5.1.1 Facies A: Massive to finely laminated rubbly mudstones/siltstones.

Facies A is primarily characterized by its fine-grained texture, usually mudstones and siltstones locally up to very fine-grained sandstone (Table 1, Figs. 9a and 9b). Most of the exposures of facies A show massive to faintly laminated cm-scale beds (Fig. 9a) which characterize the sub-facies A1, whereas sub-facies A2 can locally show mm-scale planar stratification (Table 1, Fig. 9c). When laminations are visible, we observe a lamina thickness of 0.35 ± 0.07 mm (Table S2).

Facies A, and more particularly sub-facies A1 is often associated with the smooth and fine-grained bedrock that is mostly encountered in the central and northern parts of the studied area, and within the few massive recessive intercalated beds present at Strathdon or at Glen Etive, in-between more resistant and coarser-grained sandstones of facies B. In some places, such as at Woodland Bay, the sub-facies A1 exposures show the presence of polygonal cracks linked to present day weathering (e.g., Fig. 9a, see §5.3). Overall, facies A rocks are not very resistant to erosion, and in most places are recessive relative to sandy lithologies (Fig. 7c) or form a cover of cm-scale rubbly pieces of bedrock (cf. Figs. 3a, 4b, 9b). Due to its mostly featureless character and the presence of only faint, planar mm-scale laminations, sub-facies A1 is considered to record low energy deposition.

Sub-facies A2 shows an overall similar appearance to sub-facies A1, characterized by fine-grained mudstones and siltstones. The grain size is observed to coarsen locally up to very fine-grained sandstones (e.g., Fig. 3c). Sub-facies A2 also features mm-scale planar parallel lamination (cf. Figs. 3b, 3c, 9c) with the local presence of low-angle truncation surfaces and current produced lamination, suggesting that in places at least facies A was deposited in shallow water with frequent current or wave activity.

Facies A is therefore interpreted as representing deposition in a lacustrine environment that that was in places shallow enough to preserve a record of current and possibly wave reworking.

Facies	Grain-size	Micro-structures	Macro-structures	Post-depositional features	Dynamism	Depositional setting	
A	A1	Mudstones	Massive to sub-mm-scale planar laminations	N/A	Polygonal cracks locally	Very low to low energy	Quiet lacustrine environment
	A2	Mudstones and siltstones (locally up to very fine sandstones)	Fine (sub-)mm-scale planar laminations	Mm-scale planar parallel if present		Low (+) to moderate energy	Shallow lacustrine to lake shore environment, with local current or wave influence
B	B1	Very fine and fine (up to medium) sandstones	Mm- to cm-scale cross-stratified (locally undulated) laminations	Dm- to meter-scale crossbedding	Developed light-toned vein networks	Moderate energy in a sustained flow	Distal fluvial environment to shallow fluvially-influenced lake margin
	B2	Fine-medium (locally up to coarse) sandstones	Cm-scale cross-stratified laminations	Dm- to meter-scale crossbedding and metric-scale channelized bodies	Developed light-toned vein networks	Moderate to high energy in a sustained flow	Dynamic fluvial plain (divagating river channels?)

Table 1: Sedimentary facies observed in northern Glen Torridon region.

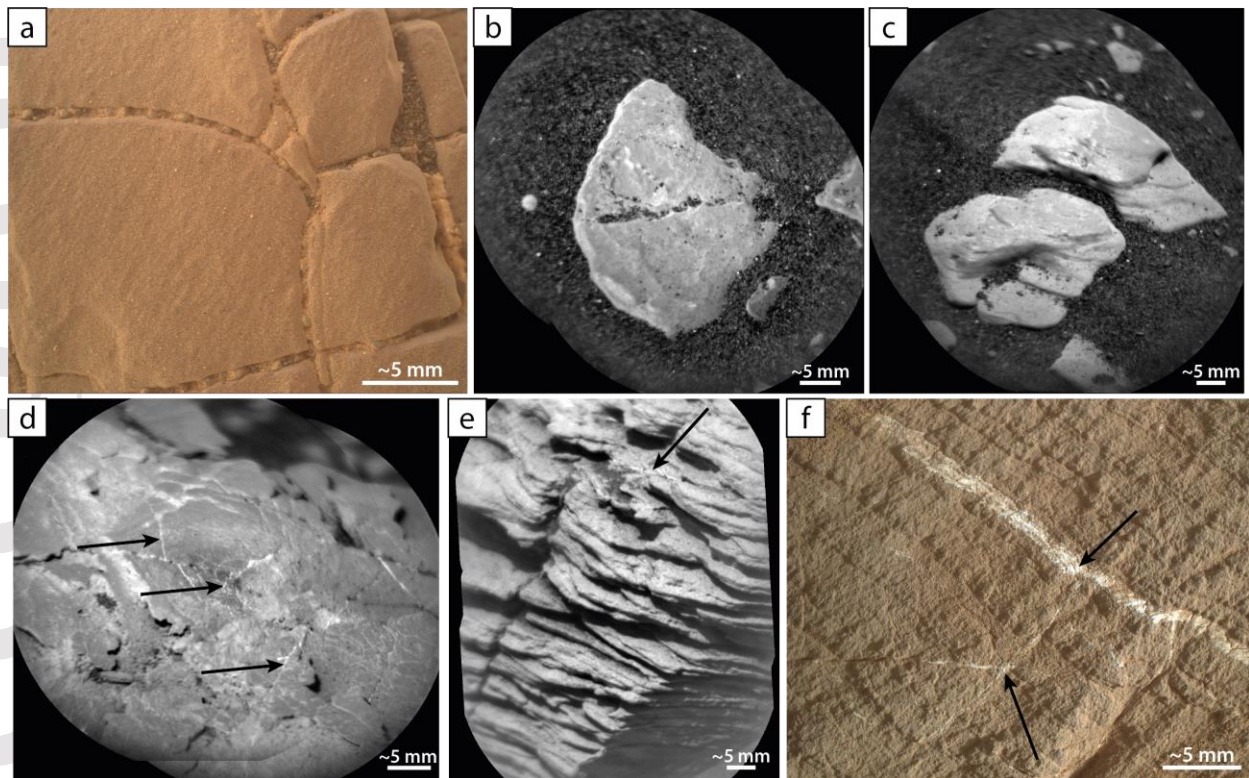


Figure 9: Sedimentary facies observed across the Jura/KHm boundary; Arrows point to calcium sulfate veins. a) Facies A1, smooth and rubbly massive mudstone, locally exhibiting cracks. MAHLI image of target “Crakaig” acquired on sol 2422 (2422MH0007150010900895C00). b) Facies A1, smooth rubbly mudstone without crack. ChemCam RMI image of target

“Hillhead_ccam” acquired on sol 2419 (ChemCam sequence ccam01419). c) Facies A2, smooth and rubbly mudstone/siltstone, exhibiting fine (sub-)mm-scale planar laminae. ChemCam RMI image of target “aegis_post_2418a” acquired on sol 2418 (ChemCam sequence ccam05416). d) Facies B1, erosion-resistant cross-stratified fine to medium sandstone. ChemCam RMI image of target “Feshie_ccam” acquired on sol 2461 (ChemCam sequence ccam02461) with LIBS raster hitting a sulfate vein. e) Facies B2, coarser erosion-resistant cross-stratified medium (to coarse) sandstone. ChemCam RMI image of target “Yesnaby_Stacks” acquired on sol 2444 (ChemCam sequence ccam01444). f) Facies B2, laminated medium sandstone. MAHLI image of target “Beaully” acquired on sol 2443 (2443MH0007230010901502C00).

5.1.2 Facies B: Cross-stratified fine to medium sandstones with light-toned veins.

Facies B is primarily characterized by its coarser-grained texture, from fine (Table 1, Fig. 9d) to medium-grained sandstone, locally up to coarse-grained sandstone (Table 1, Figs. 9e and 9f). This coarser facies is also characterized by the presence of compound cross-stratification, with bedset thickness ranging from mm-scale to dm scale and even m-scale (Table 1, Figs. 5d, 6c and 8c). Lamination thickness in facies B is typically between 0.5 and 1 mm (Figs. 9e and 9f, Table S2), but can be up to cm-scale locally. Sub-facies B1 and B2 are mainly differentiated by their grain-size, with B2 showing a larger proportion of medium, locally up to coarse, sand grains with cm-scale cross-stratification (Table 1). Most exposures of facies B are in the form of dm- to meter-scale blocks or slabs that protrude from the finer grained and rubbly bedrock in the northern part of GT (e.g., Kilmarie/Aberlady drill site and Harlaw rise, Figs. 4 and 7). Facies B is also the predominant facies in the meter-scale mesas observed in the Visionarium (e.g., Teal ridge and Antonine Wall outcrops, Fig. 5, 6 and 8). It becomes the dominant observed facies from Teal ridge southwards (cf. Fig. 1c). This facies is also characterized by the presence of well-developed light-toned vein networks (Figs. 4d, 6 and 7b and arrows in Fig. 9d to 9e). Some of these veins have been hit by ChemCam’s LIBS, as seen in Fig. 9d. Facies B rocks are overall resistant to erosion, as observed through their consistent exposures as mesas or individual blocks and slabs cropping out of the finer bedrock. In some places, the coarser sandstones are also fractured, such as at Antonine Wall (Fig. 8b).

The compound cross-stratifications observed in facies B sandstones (e.g., Fig. 5) suggests sustained flow over extended time periods (Allen, 1982), plausibly associated with fluvial processes (channels, and sandy braids) or a fluvially-influenced lake margin setting (distal deltaic distributaries). Comparable cross-stratification has been observed elsewhere in the Murray formation (Schieber *et al.*, 2017; Stack *et al.*, 2019) but does not appear to occur in the same depositional context.

At GT, the observation of compound cross-stratification in likely channel-fill sandstone bodies (Figs. 5b, 5c and S2) argues against deposition by short-lived flow events. Instead, bedform organization at the observed scale implies relatively steady uniform (and unidirectional) flow of at least many hours duration (Bohacs, 1981), an interpretation that is consistent with sustained flow in fluvial channels. The variable orientation of cross-beds seen in facies B outcrops (e.g., Teal ridge, Harlaw rise, Antonine Wall) is also consistent with migration of 3D bedforms (dunes or braids) through fluvial channels. A general southward dip of facies B strata and SE-NW oriented channel forms also argue for the former presence of a well-developed and temporally continuous fluvial system.

Sub-facies B1 may represent deposition under moderate energetic conditions that can be observed in both a distal fluvial system or within a fluvially-influenced proximal shallow lake margin. Sub-facies B2 probably represents more energetic conditions encountered in fluvial channels or deltaic distributaries. These observations and deductions suggest that the irregularly distributed resistant outcrops and mesas seen throughout the Visionarium are remnants of fluvial channels.

5.2 Geometric relations between facies A and B

In the studied area, facies A and B denote differing environmental and energetic conditions that are encountered in close lateral proximity without any conspicuous unconformity separating them (e.g., at Kilmarie/Aberlady drill site), and even appear interstratified in places (e.g., at Woodland Bay or at Antonine Wall). It is common to observe sandstone bodies with affinity to facies B within areas dominated by facies A, as well as occurrences of facies A type lithologies within larger bodies of facies B.

We observe two types of stratigraphic relationships between facies A and B throughout the northern GT area and within the Visionarium. First, a progressive transition can be documented (left on Fig. 10) showing dm- to meter-scale alternations and interfingering between both facies. This progressive transition is evidenced within alternating fine- and coarser-grained beds at Woodland Bay (Fig. 3). A similar relationship is also seen in some intervals of the Antonine Wall outcrop with occurrences of recessive finer-grained layer alternating with resistant and cross-stratified sandstones (Fig. 8c). This alternating/interfingering architecture is observed in GT area and was called “Flodigarry facies”, notably at Woodland Bay (e.g., Bristow *et al.*, 2021; Edgar *et al.*, 2020; Fedo *et al.*, *this volume*; Rudolph *et al.*, 2020). We propose that the “Flodigarry” outcrop is not a “facies” *per se* but rather reflects the progressive transition between facies A and B. Second, a sharp (possibly erosive) meter-scale discontinuity between facies A and B is observed in places (Fig. 10), most particularly at the base of Teal ridge and Antonine Wall mesas (Figs. 6 and 8). There, coarse-grained cross-stratified deposits belonging to facies B are seen deposited above facies A fine-grained deposits. Some intermediate stages can be observed at a smaller scale, for example at Harlaw rise (Fig. 7) where a sharp truncation occurs at the base of the facies B interval that overlies finer-grained and recessive facies A bed.

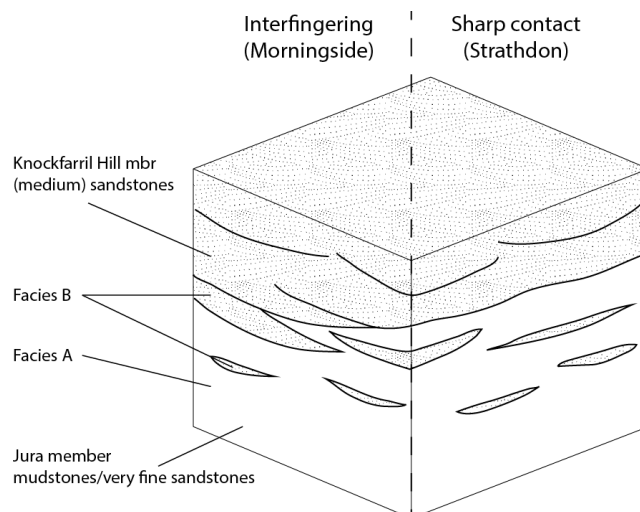


Figure 10: Schematic representation of the two types of geometric relations observed at the outcrop between facies A and facies B. Left: progressive transition marked by the presence of alternating/interfingering beds of both facies (e.g., Woodland Bay). Right: Sharp, likely erosive and laterally continuous contact between facies A and B, with B replacing A (e.g., mesas of Teal ridge and Antonine Wall).

The presence of these two types of spatially distributed transitions between the same two facies suggests differing levels of depositional and erosive energy across the area. First, we observe facies interfingering in the northern part of the study area (cf. Figs. 1 and 10) which also represents the lowest elevations in the GT area (at ~-4154m). In that case the transition involves facies A and B1, representing respectively a presumed low to moderate energy (shore) lacustrine environment, and moderate but sustained energy settings related to a fluvial influence on the lake-margin, probably reflecting shoreline regression as the lake margin was infilled with sediment. Upsection and farther South, these gradual transitions are “replaced” with channel sand bodies that appear to fill erosive incisions into the underlying lacustrine mudstones (cf. Figs. 1 and 10). These erosive contact relationships suggest deposition on a fluvial plain with channel incision forced by either high flow events or lowering of lake level. Assuming a prevalence of southward sediment transport, the upwards progression from facies A to dominant facies B with channel incisions can be considered a plausible outcome of southward sediment transport that first filled up (exhausted) accommodation along the lake margin, and thus, over time, pushed the shoreline and lake margin fluvial plains further south.

5.3 The origin of polygonal cracks at Woodland Bay

The Woodland Bay outcrop (Fig. 1c) contains polygonal crack patterns in multiple resistant sediment layers that haven't been observed elsewhere in the studied area (Figs. 3a and 11a). These cm-scale cracks are present on several low-relief exposures rising from the regolith and surrounding rubble as dm- to meter-scale patches, most likely representing the same bedding plane (Fig. 11a).

The observed polygonal-cracked layers are interstratified with softer thinly layered intervals, parting planes are well developed, and in sideview the cracks are parallel-sided and may terminate upon encountering the softer intervals. In contrast, desiccation cracks would be expected to taper downwards. In outcrop the cracked layers tend to “peel off” from the underlying softer strata. These observations suggest mechanical decoupling between polygon layers and soft interbeds. Although these cracks resemble geometrically those at “Old Soaker” (lower in the section; Schieber *et al.*, 2017; Stein *et al.*, 2018), polygonal geometries such as these form whenever a coherent layer undergoes shrinkage, no matter what the underlying cause (Chan *et al.*, 2008; Cosgrove & Hudson, 2016; Loope & Burberry, 2018; Williams & Robinson, 1989). The observation that we see both rectilinear and hexagonal fracture patterns on the same surface (Fig. 11b) suggests that the formation of the observed crack patterns involved many opening (contraction) cycles (Goehring, 2013). The grain size of polygon layers observed with MAHLI (Fig. 11c, Fig. S3) suggests that they likely are coarse mudstones (aka siltstones). Because of framework grain support, siltstones do not contract upon drying and thus are unlikely to form desiccation cracks. Also, had these cracks formed at the time of sediment deposition, one should expect to see evidence of lithified crack fills (e.g., Stein *et al.*, 2018). The absence of such

fill lends further credence to a weathering related origin of these cracks because we only observe modern-day, wind-blow sediment filling those cracks (Fig. 11c).

Collectively, the observations made on the Woodland Bay polygonal-cracked layers are consistent with and of the same scale and characteristics as polygonal networks observed during outcrop weathering of mudstones on Earth, and likely caused by thermal cycling. Given the large day-night temperature differences on Mars, thermal expansion-contraction of surface layers is a plausible process for formation of the polygonal-cracked layers, and the cracks are most likely unrelated to the original depositional environment of these strata.

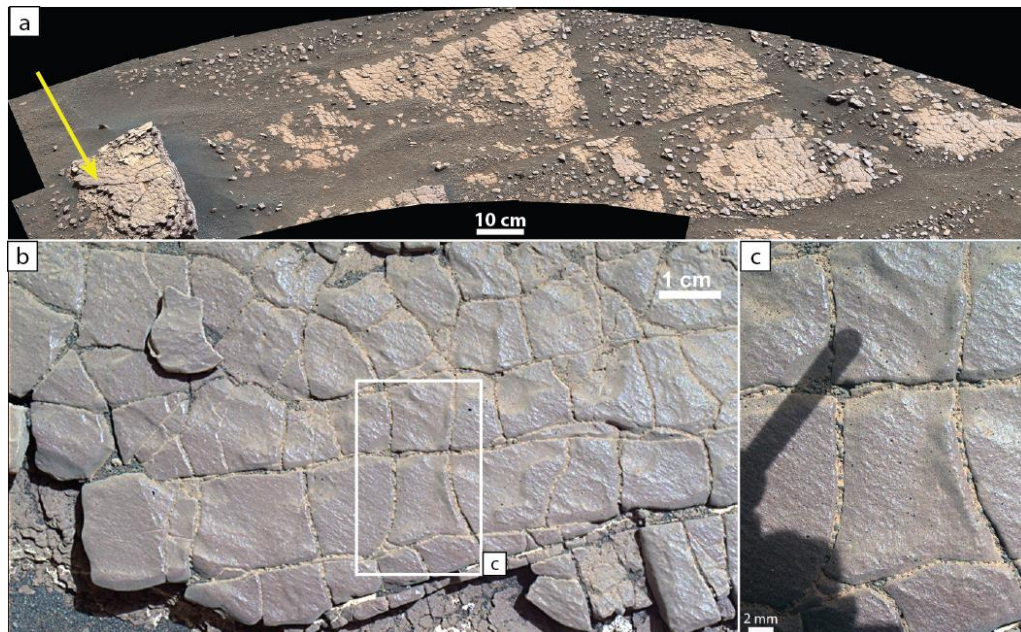


Figure 11: a) General view of the Woodland Bay outcrop toward the polygonal-cracked patches. Mastcam mosaic acquired on Sol 2424 (sequence mcam012841). With exception of the far left block (yellow arrow), these may represent a single bedding plane. b) Detail view of the target Crakaig yellow arrow on a) showing well-developed polygonal crack patterns (MAHLI image 2422MH0007060010900870C00 acquired on Sol 2422). c) Closer detail view showing that the cracks are filled with modern surface sediment, eolian sand, silt, and dust, with no indication of a lithified crack fill.

6 Geochemical compositions and trends

6.1 Major Oxide Composition

Here, we investigate the Major Oxide Composition (MOC) measured by the ChemCam instrument on our subset of 36 targets within or near the main outcrops described above (detailed list in Table S1, complete data in Table S3). We want to test the existence of a possible relationship between the sedimentary facies A and B described above and the geochemical variability observed in GT. To that purpose, we focus on oxides SiO_2 , CaO , MgO , Na_2O and K_2O . Other compositional trends observed within bedrock geochemistry (e.g., TiO_2 , Al_2O_3 , FeO)

and their in-depth analyses are beyond the scope of this study and are addressed in detail in the companion work by Dehouck *et al.* (2022). To test the existence of a possible relationship between the sedimentary and geochemical characters, individual points of measurements are colored in Figure 12 according to the facies determined for each corresponding target (Table S3).

Figure 12 shows the abundances for CaO, MgO, K₂O and SiO₂. Most of the CaO values (including all facies A samples) are concentrated in an area of the plot close to GT mean values (indicated by the whiskers on Fig. 12a, <2 wt.%). Facies B targets show higher CaO values (>2 wt.%), especially for those belonging to sub-facies B2 (> 2wt.% and up to ~8 wt.%). This enrichment is likely related to the light-toned veins frequently observed in targets of this sub-facies (e.g., Figs. 10d and 10f), which are composed of calcium sulfate, as previously observed along *Curiosity*'s traverse (e.g., Nachon *et al.*, 2014, 2017; Rapin *et al.*, 2016; 2019). The lower SiO₂ content of the B2 targets (< 55 wt.%) is consistent with this interpretation. Overall, facies B shows values closer to the Martian average crust (Taylor & McLennan, 2009) and Gale conglomerates (Mangold *et al.*, 2016; Fig. 12a) for the two oxides considered here.

A clear compositional difference between samples from facies A and B is also observed in the MgO vs K₂O plot (Fig. 12b). LIBS points identified as facies A show a lower mean MgO content (5.7 ± 0.8 wt.%) and a higher mean K₂O content (1.9 ± 0.4 wt.%). Conversely, LIBS points belonging to facies B show a higher mean MgO content (7.1 ± 1.0 wt.%), and a lower mean K₂O content (1.3 ± 0.7 wt.%). Here again, facies B shows the closest values to average Martian crust. On the other hand, when comparing to the Gale conglomerates, we observe that facies A1 is closest to the Kimberley conglomerates, a K-rich deposit previously identified on Aeolis Palus (Mangold *et al.*, 2016). Despite the observed difference in K₂O content, facies A and B show similar Na₂O values, close to both the average Martian crust and the Kimberley conglomerates (Fig. S4).

Thus, facies A and B display significant differences in major element abundances that may be a consequence of their distinct texture and depositional mode.

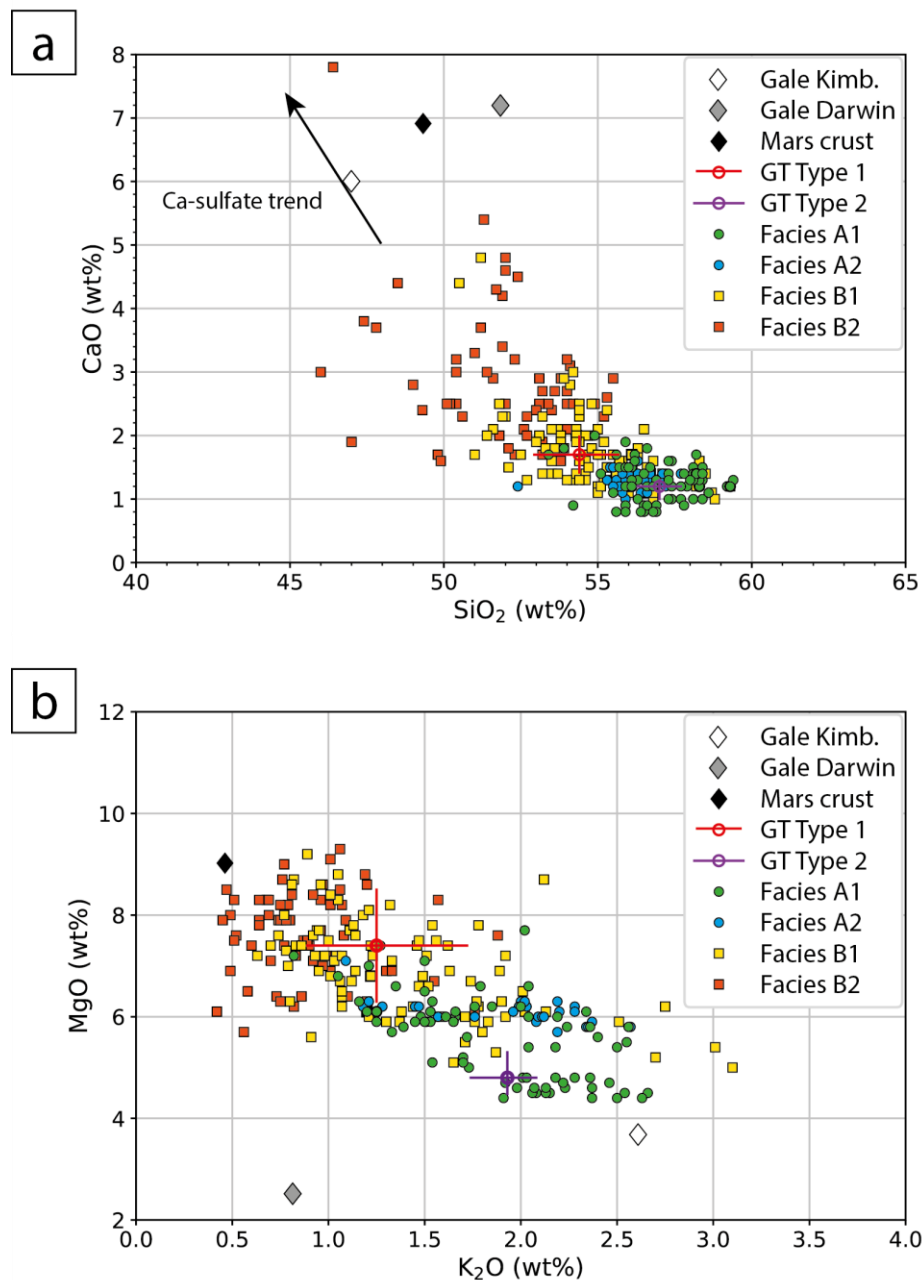


Figure 12: Major-oxide compositions of representative GT targets in CaO, SiO₂, MgO and K₂O (in wt.%). Each dot represents a single analysis point, colored as a function of its sedimentary (sub-)facies A1, A2, B1 and B2. Types 1 and 2 are the “coherent” and “rubbly” endmembers described by Dehouck *et al.*, (2022). a) CaO vs SiO₂ plot. Arrow represents the trend to Ca-sulfate compositions. b) MgO vs K₂O plot. Average Martian crust values from Taylor & McLennan (2009); average Gale basin conglomerates values from Mangold *et al.*, (2016) (Kimb.: Kimberley).

6.2 Relation between the studied facies and regional geochemical endmembers

The GT region shows a geochemical composition of the major elements that is broadly in line with the underlying Murray formation, but with some notable variability (e.g., Dehouck *et al.*, 2021, *this volume*). In the Jura member, two endmembers have been identified from both the APXS and ChemCam measurements (e.g., Dehouck *et al.*, 2021, 2022; O’Connell-Cooper *et al.*, 2022). These two endmembers were referred to as “rubbly” and “coherent” from the relative resistance to erosion of the corresponding rocks (Dehouck *et al.*, 2022), with the “rubbly” consistently outcropping as small pieces of fragmented bedrock (e.g., Fig. 9a), whereas the “coherent” is represented by larger blocks and slabs (as observed at e.g., Kilmarie/Aberlady, Teal ridge).

These two endmembers are mainly discriminated based on their MgO and K₂O contents (Dehouck *et al.*, 2021, 2022; O’Connell-Cooper *et al.*, 2022). The “rubbly” endmember shows a K₂O-rich (> 1.5 wt.%) and MgO-poor (usually <6 wt.%) composition, defined as the regional Type 2 composition by Dehouck *et al.*, (2022). The “coherent” endmember shows higher MgO content (~6 to 11wt.%) and a lower K₂O content (< 1.5 wt.%), defined as the regional Type 1 composition by Dehouck *et al.*, (2022).

Figure 12b is thus most important in seeking a potential relation between the geomorphic expression at the outcrop, sedimentary facies, and geochemical signals in GT. We observe a clear distinction between facies A and B samples in the Mg and K contents. The fine-grained and rubblier facies A is in line with high-K Type 2; and the coarse-grained, slab-forming facies B is mostly in line with high-Mg/low-K Type 1 compositions. However, the sub-facies B1 shows a more important scattering of the K₂O values than the other sub-facies. This could be the result of a local mixing of sedimentary material at the interface between facies A and B, which is compatible with the alternating and interfingering depositional mechanism that we have highlighted on several outcrops (see section 5.2).

Overall, targets from facies B have a composition closer to the average Martian crust compared to targets from facies A (Fig. 12). This difference could reflect a less pronounced aqueous alteration of facies B material. However, the values of the Chemical Index of Alteration are similar for the Types 1 and 2 endmembers (Dehouck *et al.*, 2022). Alternatively, the difference in composition could be due a difference in sediment provenance, with facies A material being derived from a more K-rich source. The existence of a K-rich source in the Gale watershed was already inferred earlier in the mission based on the observation of potassic sandstones and conglomerates at the Kimberley location (Le Deit *et al.*, 2016; Mangold *et al.*, 2016; Treiman *et al.*, 2016).

7 Discussion

7.1 Significance of the facies and stratigraphic evolution from Jura member to Knockfarril Hill member

As we elucidate the stratigraphic and paleoenvironmental evolution of the studied area and broader regional context, it is important to recognize the relations between the facies characterized in this work and the members of the Murray and Carolyn Shoemaker formations defined by the MSL team. These members are defined by the team from both the characters as

seen from orbital (spectro-) images, and the textures deciphered using *in situ* observations made using *Curiosity*. Yet, from their sedimentary characters (textures, laminae thicknesses, structures) and their overall spatial distribution, both facies A and B fall in line with previously described members present in the GT area: the Jura and Knockfarril Hill members, respectively (e.g., Bennett *et al.*, *this volume*; Edgar *et al.*, 2020; Fedo *et al.*, 2020, *this volume*). However, from a textural and sedimentary point of view, the Jura/KHm boundary (Fig. 13a) has been drawn from orbital and *in situ* observations based on the difference in resistance to erosion, an approach that does not capture the smaller-scale local variability in facies. Thus, observed occurrences of facies A can be referred to as “Jura-type deposits” whereas facies B beds can be referred to as “KHm-type deposits”, to avoid any ambiguity with the mapped units.

We have shown above that the chemical composition of the bedrock varies between two regionally observed endmembers that are consistent with sedimentary defined Jura-type and KHm-type deposits, implying a link between the facies and the geochemical signal. Specifically, on the one hand, the Type 2 endmember (as defined by Dehouck *et al.*, 2022) is associated to rubbly outcrop expressions and morphologies, therefore corresponding to fine-grained Jura member facies. On the other hand, the Type 1 endmember corresponds to coherent and slab/ridge-forming morphologies, therefore associated with coarser-grained KHm facies.

Outcrops with a Jura-type signature within the mapped KHm member can be interpreted as occurrences of Jura member material within the coarse-grained KHm member, and *vice versa*, where KHm-type signatures within the mapped Jura member can be interpreted as occurrences of KH member. Consequently, the two samples acquired within the mapped Jura member during the GT campaign (Aberlady and Kilmarie) have sedimentological and geochemical characteristics that link them to KHm-type deposits. The “true” Jura-type material, on the other hand, was too fragmented to be sampled with the drill.

Whereas a clear topographic and textural change is observed between Jura and KHm members in the regional orbital imagery from HiRISE, allowing for mapping those two units at the regional scale (Fig. 13a, Bennett *et al.*, *this volume*; Fedo *et al.*, *this volume*), this distinction is not as clear and obvious from rover observations along the traverse. Indeed, Figure 13 shows that the spatial distribution of Jura-type and KHm-type material exposures across the region are scattered laterally (Fig. 13a) as well as vertically (Fig. 13b). The cross-section in Figure 13b demonstrates that KHm-type deposits are observed as low as -4154m in elevation, which is 10 meters below the mapped limit of the KHm *sensu stricto*. Conversely, Figure 13a also shows the recurring presence of fine-grained Jura-type deposits within the KHm map area (blue squares in Fig. 13), indicating coexistence of energetic fluvial and less energetic lake margin depositional settings. This geographic distribution means that strata typical of the KHm occur somewhat lower, and therefore earlier in the history of GT than suggested by the regional mapping. It also indicates a more complex history, with laterally and vertically alternating deposition of Jura-type and KHm-type materials during the time interval represented by the ~10m-thick succession observed in the study area.

In summary, the fine-grained facies of Jura member represents quieter settings dominated by lacustrine processes, whereas coarser-grained and cross-stratified facies of KHm represents more energetic, fluvially-influenced lake margin and fluvial channel settings. The spatial distribution of both facies, their gradual as well as sharp contact relationships, together with their specific geochemical signature (Types 2 and 1 endmembers, respectively), collectively are the expression of a dynamic depositional setting and lateral facies migration. The widespread

dominance of fluvial deposition at the onset of the KHm *sensu stricto* (as mapped) may reflect accelerated southward facies migration due to increased sediment supply, or a forced regression of the lake shoreline due to either a lake level drop and/or a change of localization of the regional depocenter.

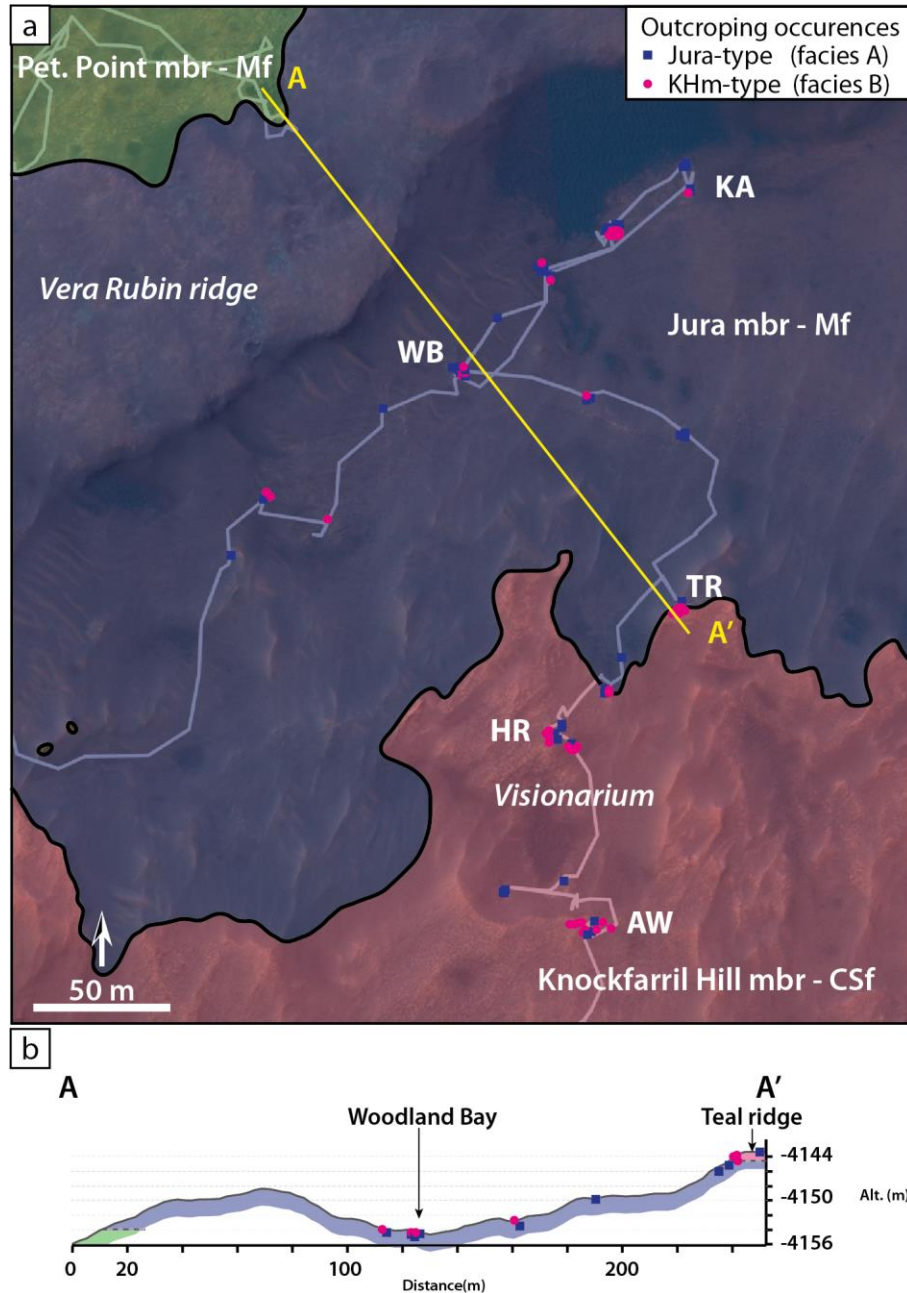


Figure 13: a) Spatial distribution of the studied targets along the traverse after their characterized sedimentary Jura- or KHm-type facies. This map shows that both facies are observed occurring within both members mapped area, demonstrating a progressive transition between them rather than a sharp geographic boundary. Regional geological basemap after Fedo *et al.* (this volume). WB: Woodland Bay, KA: Kilmarie/Aberlady drill sites, TR: Teal Ridge, HR: Harlaw rise, AW:

Antonine Wall and Glen Etive drill site; Mf: Murray formation, CSf: Carolyn Shoemaker formation (cf. Fig. 2). b) Regional cross-section of the northern GT area from VRR and down to the Visionarium area.

7.2 A progressive but marked change in the regional depositional settings in Glen Torridon across the Jura/Knockfarril Hill members boundary

Up to the VRR, *Curiosity* has been exploring 300+ meters of mostly fine-grained lacustrine strata of the Murray formation. Arriving at GT, however, we observe a change in the stratigraphy representing an important shift in the regional depositional setting toward more energetic conditions over a comparatively short time period (~10 m thick interval, likely representing multiples of 1000's to 10,000's of years). This transition is marked by the onset and increasing occurrence of the KHm-type sandstones, characterized by fluvial processes. We therefore propose that this interval records a significant change in the regional organization of the Gale crater depositional system. This change could have been caused by migration of the lake margin, perhaps due to substantial drop in the regional base level. Such a drop would force basinward migration of the shoreline and favor the observed lateral and vertical shift from lacustrine to fluvial settings in the study area. We can hypothesize that changes in climate could have caused a decrease of water supply, lead to lake level drop and migration of the regional depocenter, and thus a general expansion of fluvial plains. Indeed, the KHm sandstones interval continues up section for more than 10 additional meters (cf. Fig. 2; Fedo *et al.*, *this volume*), indicating that fluvial conditions were sustained for a significant time interval in the area.

This trend toward a subaerial evolution of the basin also has potential to be investigated at a larger scale. As the rover continues to climb up, it will be important to test whether the paleoenvironmental change recorded at the Jura/KHm boundary is restricted to the GT area, or if it is a part of a broader evolution of the depositional settings in Gale crater, as hinted at by decameter-scale cross-stratified deposits observed in the overlying so-called “sulfate unit” (corresponding to the Pontours member of Carolyn Shoemaker formation; cf. Fig 2 and Fedo *et al.*, *this volume*), that are thought of to be of aeolian origin with local fluvial inputs (Rapin *et al.*, 2021).

7.3 Geochemical processes associated with sedimentation

A lacustrine setting often produces a “homogeneous” geochemical signal because sediments from the various sources and inputs of the basin are mixed; conversely, the composition of a fluvial deposit reflects more directly the composition of the source rocks encountered in the watershed (e.g., Garzanti *et al.*, 2011; Jin *et al.*, 2006; Sawyer, 1986). As we observe variations in the GT paleoenvironmental conditions across the Jura/KHm boundary, the correlation between sedimentary and geochemical parameters is interpreted to reflect this regional change in depositional settings.

We have demonstrated that in GT, a strong relationship exists between the sedimentary and the geochemical records, with each facies displaying a characteristic geochemical signature, which is Si- and K-rich for the Jura-type material on one hand, and Si-poor, K-poor and Mg-rich for the KHm-type material on the other hand. These variations in abundances could be associated with several processes such as (i) the impact of fluid circulation during cementation and/or diagenesis, (ii) a mechanical segregation related to the changing energy of the depositional setting and/or proximity to primary sources, or (iii) the influence of the accumulation and/or (iv)

in situ formation of clay minerals, which is of particular interest in the GT area, aka the “clay-bearing unit”.

The KHm-type material is characterized by its coarser-grained texture (fine- to medium-grained sandstones) due to deposition under energetic fluvial conditions, and good outcropping conditions as erosion-resistant pluri-dm-scale slabs to pluri-meter-scale mesas (particularly in the Visionarium area). This resistance to erosion can be linked to a higher degree of cementation. The larger grain size of these sandstones is a favoring factor because it implies a primary porosity that increases the potential for fluid circulation and subsequent cementation during burial (e.g., Bjørlykke, 1988; Hiatt & Kyser, 2000). Cementation and fluid circulation are plausible processes for enhanced accumulation of Mg (notably in cements), and for the depletion of K due to destruction of labile minerals (e.g., Argast & Donnelly, 1987; Jin *et al.*, 2006).

The KHm-type material is also characterized by the presence of well-developed light-toned vein networks, that are both observed at the outcrop, and detected through the geochemical signal with higher values of CaO and lower values of SiO₂ in the coarsest targets of the facies B2, which are commonly associated in Gale crater with secondary sulfate vein formation (e.g., Nachon *et al.*, 2014, 2017; Rapin *et al.*, 2016). Those veins are absent from the more massive and fine-grained Jura-type material, highlighting the mechanical contrast between KHm and Jura rocks. The latter, being poorly cemented and mechanically weak, more likely deformed instead of fracturing, whereas the former are likely to have behaved rigidly and prone to fracture formation. Thus, secondary porosity due to fracture formation, as well as later stage fluid circulation and mineralization should largely be restricted to the sandy KHm rocks.

Mechanical segregation of minerals may be a cause for K-enrichment within finer-grained Jura-type material. This mechanism implies that throughout their transport, minerals can be sorted according to their density, and therefore, K bearing minerals (such as K-feldspar) may have been preferentially transported and deposited under lower energy conditions due to its lower density compared to mafic mineral grains (e.g., Argast & Donnelly, 1987; Jin *et al.*, 2006; Sawyer, 1986). As the Jura-type material is representative of lower energy depositional settings, such a mechanism could explain the observed local enrichment in K. A similar relationship was already observed earlier in the mission by Siebach *et al.* (2017) with K concentrated in the fine fraction resulting from the alteration of feldspar-rich protolith. This hypothesis of a role of the transport is also supported by comparing both materials' signature with the average Martian crust (Taylor & McLennan, 2009) and the Gale basin average conglomerates (Mangold *et al.*, 2016). We can see that the coarser-grained fraction of the KHm-type material is overall closer to the average crustal composition, indicative of a sediment possibly closer to its primary source. On the contrary, K-rich Jura-type material shows a conspicuously differing composition and systematically more distant from the average crustal values, indicative of a reworked and/or more evolved sediment. This latter's origin might be related to a different sedimentary source, similar to previously identified K-rich bedrock sources within the Gale crater watershed indicated by the K-rich sandstones and conglomerates observed at Kimberley (Le Deit *et al.*, 2016; Mangold *et al.*, 2016). As Jura-type material exhibits K₂O values close to that of the Kimberley conglomerates (Fig. 12b), a common origin with Kimberley's sediments cannot be ruled out.

Finally, whereas this hypothesis is compatible with a model in which K is leached from coarser-grained sediments with higher porosity such as from the KHm-type material, an alternative hypothesis, supported by observations on minor elements (Cousin *et al.*, 2021; Dehouck *et al.*, 2022), suggests that enrichment in K could also be due to the presence of illite, a K-rich clay mineral that would have been segregated mechanically during transport from the coarser sediment fraction (KHm). Post-depositional incorporation of potassium in clay minerals (as suggested by Thorpe *et al.*, 2020) is yet another process that needs to be considered.

8 Conclusions

As the *Curiosity* rover entered the Glen Torridon area, it encountered strata that differed notably from the underlying ~300m of Murray formation. The sedimentary succession of the area consists of two contemporaneous facies: a mudstone-dominated “facies A” that accumulated in a low-energy lacustrine setting and shows evidence of current reworking in presumed lake margin locations and a sandstone-dominated cross-stratified “facies B” associated with sustained energy, channel fill deposits in alluvial plains and fluvial-influenced lake margin settings. Facies A and B are typical of rocks from the Jura and Knockfarril Hill members, respectively, with significant exceptions as these facies are observed to alternate and/or interfinger with each other across the mapped Jura/KHm boundary. This suggests that these facies co-existed across the area, and through time shifted laterally in response to sediment supply and/or base level changes.

Geochemically, facies A and B can be considered compositional endmembers, corresponding to regionally recognized Types 2 and 1 compositions, respectively (2 being K-rich; 1 being K-poor and Mg-rich). Data points along the traverse reflect the intermingling of facies and the upwards shift from facies A to facies B.

The paleoenvironmental shift across the Jura/KHm boundary has been documented through sedimentological analysis of outcrops and is mirrored by the changes in the chemical compositions collected from these same locations. This transition, first progressive as illustrated by the interfingering of the two facies, and then expressed up section as sharper erosive contacts between the two, reflects the onset and then steady expansion of fluvial dominated depositional environments in the study area. The underlying causes for this change could be local or regional drop of base level, possibly in relation with changing global climatic conditions at the Noachian/Hesperian boundary.

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All of the Mastcam (Malin, 2013), Navcam (Maki, 2013), and MAHLI (Edgett, 2013) images used in this manuscript are freely available through the Planetary Data System Cartography and Imaging Sciences. All of the ChemCam RMI images (Wiens, 2013b) and LIBS data (Wiens, 2013a) used in this manuscript are freely available through the Planetary Data System Geosciences node. All 3D DOMs presented in this manuscript are freely accessible for visualization, including in Virtual Reality, on the Sketchfab web-platform (<https://sketchfab.com/LPG-3D>). Complete data tables presented in the supplementary online material of this work are available in the Zenodo open-source online repository (Caravaca *et al.*, 2022).

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