

Handheld LIBS analysis for in situ quantification of Li and detection of the trace elements (Be, Rb and Cs)

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1	Handheld LIBS analysis for in situ quantification of Li and
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11 Abstract

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13 The world's lithium demand is currently increasing due to its use in batteries and in green 14 energy technologies. As a consequence, Li has recently been added to the list of critical materials in Europe. Finding new deposits is necessary in order to guarantee a mid- to long-15 16 term stable supply. Granites and associated pegmatites are the primary 'hard-rock' source for Li, where it is present in a variety of aluminosilicate minerals. Li is a light element that typically 17 18 occurs at low concentrations in rocks and minerals, which presents challenges for some 19 analytical techniques (e.g., portable X-ray fluorescence) for exploration. Laser induced 20 breakdown spectroscopy (LIBS) is a powerful analytical technique for the acquisition of 21 simultaneous almost all elements of the periodic table. LIBS has particular sensitivity to the light 22 elements (Z<13) and, therefore, handheld LIBS has significant potential for the exploration 23 of pegmatite- or granite-related Li mineralization. For this study, a suite of 16 pegmatite-hosted 24 Li ore minerals comprising the three main aluminosilicate species (spodumene, lepidolite and petalite) that contained between 1.36 to 3.54 wt.% Li was analyzed using a handheld 25 26 LIBS instrument. A univariate calibration curve was developed for LIBS quantification of 27 Li in these minerals using the strong Li emission lines at 610.36 nm and 670.79 nm ($R^2 > 0.9$, 28 RMSD = 0.2 wt. % Li). Simultaneous quantification of Li contents and measurement of 29 major element ratios (e.g., Si/Al) using portable LIBS permits rapid discrimination between 30 different Li-bearing silicates, such as spodumene and petalite, that can be difficult to identify 31 in the field on the basis of macroscopic criteria. Moreover, the detection of minor/trace elements such as Be, Rb and Cs is possible by handheld LIBS, which can provide useful 32 information for rare metal pegmatite and granite exploration. Thus, qualitative and quantitative 33 34 analysis by handheld LIBS represents a very promising approach for Li exploration.

36 Keywords: LIBS, lithium, Li minerals, handheld analysis

37

38 1. Introduction

39 The economic importance of Li is growing as it represents a key component in the 40 manufacturing of green energy storage-devices such as Li-ion batteries (Bibienne et al., 2020). As consequence, the global demand for Li has recently increased and Li was officially added 41 42 in 2020 to the list of critical raw materials published by the European Commission for the 43 European Union (COM-2020-474). Most Li resources around the world are hosted in brine 44 deposits from salars or geothermal fields, but other important sources are volcanic-sedimentary clays, granites, and granite-related pegmatites (Kesler et al., 2012; Bowell et al., 2020). 45 Although salar deposits contain the largest Li reserve, Li-rich mineral concentrates from 46 47 pegmatites are currently the main source of Li production due to its higher grade and lower 48 production cost (Bowell et al., 2020). As previously described by various authors (Černý, 1991; 49 Defnet et al., 2021; London, 2014, 2005a) the granite-associated pegmatites that are sources of 50 rare metals belong to two very distinct families - the Li-Cs-Ta (LCT) pegmatites and the Nb-51 Y-F (NYF) pegmatites. Quartz, potassium feldspar, and albite are the major constituents in 52 both the LCT and NYF pegmatites, with muscovite, biotite, garnet, tourmaline, and apatite 53 typically present as accessory phases. Pegmatites of the LCT family are characterized by 54 enrichment in Li, Rb, Cs, Be, Sn, Nb, Ta, B, P and F (Barros et al., 2020; London, 2014, 2005b). 55 Most LCT pegmatites occur worldwide in supracrustal rocks metamorphosed in the upper 56 greenschist to lower amphibolite facies that are are typically located in the vicinity of evolved, 57 S-type peraluminous granites and leucogranites from which they are inferred to be derived by 58 fractional crystallization . In chemically evolved LCT pegmatites, spodumene, petalite, and 59 lepidolite are the major Li-bearing minerals. Beryl is the primary host mineral for Be, pollucite 60 for Cs, the columbite group minerals for Nb and Ta, and cassiterite for Sn. Exploration and 61 assessment for LCT pegmatites are guided by a number of observations that include increasing 62 contents of Li in white mica and increasing Mn in garnet (Charoy et al., 1995). Thus, pegmatite-63 or granite-hosted Li mineralization is characterized by a wide geographic distribution and the 64 generation of valuable byproducts that include other strategic rare metals such as Ta, Nb, Sn, 65 Be, Cs, and Rb as well as raw materials such as quartz, feldspar, and micas (Defnet et al., 2021; Kesler et al., 2012; Linnen et al., 2012). As suggested by London (London, 1984), "... the 66 67 lithium aluminosilicate phase diagram constitutes a petrogenetic grid from which P-T

68 conditions of emplacement and crystallization can be ascertained for many lithium alumino69 silicate-bearing pegmatites".

70 In western Europe, such 'hard-rock' Li deposits, which offer the best potential to supply the 71 continent for the immediate future (Gourcerol et al., 2019), are primarily associated 72 with rare-metal granites and pegmatites of the Variscan orogenic belt that extends for more 73 than 3,000 km from southwestern England, through Spain, France, and Germany to 74 the Czech Republic. Typical Li ore minerals include aluminosilicate phases, such as spodumene, petalite, micas (e.g., lepidolite), and phosphate minerals (e.g., montebrasite) that 75 76 are disseminated within granite and localized in associated pegmatite bodies and local 77 metasomatic greisens. The typically small size of these intrusions (generally <1 km³), together 78 with the absence of petrophysical characteristics of mineralization, are serious limitations for 79 exploration. Moreover, due to its small atomic weight, Li cannot be detected by X-ray 80 fluorescence (XRF) spectrometry so that the handheld XRF analyzers commonly used during field exploration campaigns for metal deposits cannot be used for identifying Li mineralization. 81 82

This work is part of LIGHTS (Lightweight Integrated Ground and Airborne Hyperspectraln 83 Topological Solution), a European research project aimed at improving techniques for Li 84 exploration, including the evaluation of portable tools. Our objective was to use handled laser-85 induced breakdown spectroscopy (LIBS) to identify and discriminate between the Li-bearing 86 phases spodumene, petalite, and lepitolite commonly encountered in metallogenic provinces 87 88 associated with rare-metal pegmatites and granites and to quantify their Li contents by developing a calibration curve for these minerals. Some minerals are quite difficult to 89 90 distinguish in the field due to their similar colors (see Supp. Figures S1a, S1b and S1c). We aimed to establish the LIBS calibration curve of Li content for different matrices. We also 91 92 assess the potential of using handheld LIBS to check the Al/Si emission line ratios 93 to discriminate between these Li minerals, as well as for the detection of other rare metals such 94 as Be, Cs and Rb that can be directly mined along with Li or be used as exploration vectors 95 for Li-mineralization.

96

97 2. Background

98 Details of the LIBS technique have been described in several reviews (Cremers and
99 Radziemski, 2006; Hahn and Omenetto, n.d.; Radziemski and Cremers, 2013; Senesi, 2014;
100 Fabre, 2020; Harmon and Senesi, 2021; Senesi et al., 2021). The components of a typical LIBS

101 system are basically the laser source (typically a Q-switched Nd:YAG at 1064 nm), a gated 102 detector attached to the spectrograph/spectrometer array, fiber optic cable that allows energy 103 transit from the laser source to the sample, and a detector that is sensitive to the plasma-related 104 emissions. LIBS has been attracting the interest of the scientific community over the past 50 105 years, which generated significant technical improvements especially during the last two 106 decades as a consequence of the miniaturization of the component lasers, spectrometers, 107 computers, and electronics along with an increase in the storage capacity of the batteries powering the devices. Multiple prototypes of field-portable LIBS systems were developed 108 109 during the 1990s that were (i) lightweight to provide maneuverability for field-use, (ii) solid and ruggedized for use under different field-conditions, (iii) simple to use, and (iv) optimized 110 111 for the detection of a large number of elements over a minimal acquisition time (Connors et 112 al., 2016; Cuñat et al., 2005; Palanco et al., 2003; Rakovský et al., 2014). This work led to the 113 more recent development and marketing of commercial handheld LIBS analyzers by multiple manufacturers. Several recent studies describe the use of handheld LIBS for different 114 115 geological applications that include mineral exploration (Connors et al., 2016; Defnet et al., 116 2021; Harmon et al., 2019, 2021; Lawley et al., 2021; Senesi, 2014; Senesi et al., 2021).

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118 Lithium is an important geochemical tracer for fluids or solids, but its abundance variation at the 119 micrometer scale is not readily determined through most analytical techniques and its bulk 120 concentration can only be measured using classical bulk analytical techniques. Although a 121 scanning electronic microprobe fitted with a soft X-ray emission spectrometer (SXES) can 122 analyze for Li at high spatial resolution (Takahashi et al., 2015), this is an expensive and time-123 consuming analytical procedure that can only be undertaken in the laboratory. Not only is 124 LIBS one of the few analytical techniques for rapid Li detection down to the ppm level (Fabre 125 et al., 2021, 2002) it is one that can be employed outside the laboratory in the field. Thus, the 126 analysis of Li using LIBS has been the subject of geological interest (Fabre et al., 2021, 2002, 127 1999; Sweetapple and Tassios, 2015; Verlaguet et al., 2016). Fabre et al. (Fabre et al., 2002) 128 used a univariate calibration model for Li quantification in synthetic glasses and natural Li-129 rich minerals using the characteristic Li line at 670.79 nm and then determined the Li contents 130 of geological materials that included spodumene, petalite, and fine-grained eucryptite from 131 granite pegmatite dikes in Portugal, melt inclusions from the Streltsovka uranium deposit in Siberia, and as a trace element in barren quartz veins from Spain. Estimated wt. % contents of 132 133 Li in spodumene and petalite Li compared well with abundances obtained from ion microprobe 134 analysis, with the possible exception of eucryptite because its grain size was smaller than the

135 laser spot size. Since then, numerous studies described the use of LIBS for analysis of silicate 136 gem minerals or other applications in geology (Harmon et al., 2009, 2006; Kochelek et al., 137 2015; McMillan et al., 2018). For example, Rossi et al. (Rossi et al., 2014) reported Be and B 138 concentrations within gem-quality spodumene, topaz, and tournaline that were in excellent 139 agreement with LA-ICP-MS and EMPA analyses for the same samples. A LIBS study of 140 spodumene by Sweetapple and Tassios (Sweetapple and Tassios, 2015) showed that it was possible to measure Li, Be, and B in a hydrothermally altered pegmatite, with microscale 141 mapping of Li used to discriminate between spodumene, its alteration mineralogy, and matrix 142 143 silicates. More recently, new studies have been done combining handheld LIBS tool and 144 geochemical exploration for elemental mapping in Li-rich rocks (Harmon et al., 2021; Lawley 145 et al., 2021).

146 This study of handheld LIBS analysis had two objectives, to obtain rapid discrimination of Li-147 rich minerals that are often difficult to distinguish visually in the field and to perform Li 148 quantification beforehand in the laboratory in order to subsequently apply the Li calibration to 149 pegmatite outcrops in the field for the LIGHTS project (Fabre et al., 2021). One of the 150 objectives of this study is then the access to fast and *in situ* use of internal calibration and then 151 reduce at maximum any other use of external treatment of the data. All of the LIBS spectra (each individual LIBS spectra and mean spectra) and more information about the sample 152 153 analyses or location used in this paper are archived for open access at 154 https://doi.org/10.24396/ORDAR-65.

155

156 **3.** Samples and methods

157 *3.1 Li-bearing samples selection and bulk analyses*

158 There are more than a hundred lithium-bearing minerals, most of which are silicates (Grew, 159 2020). The most common and economically significant Li bearing minerals are the 160 aluminosilicates present in rare-metal pegmatites and granites, namely spodumene - LiAlSi₂O₆, 161 petalite - LiAlSi₄O₁₀, and lepidolite - KLi₂AlSi₄O₁₀F₂ (Table 1). Other Li-bearing minerals are 162 phosphates, such as montebrasite, amblygonite, and the rarely encountered lithiophosphates. Li contents are the higher in the phosphate minerals, but these are rarely of economic importance 163 164 (Gourcerol et al., 2019; Kesler et al., 2012). Therefore, this study focused on the Li-bearing silicate minerals, although geological information such as location, mineralogy and Li contents, 165 166 continuity and size and relation to enclosing rocks, is relatively sparse and scattered (Gourcerol et al., 2019; Kesler et al., 2012; Tadesse et al., 2019). 167

- 169 Table 1
- 170 Characteristic of the most common Li-bearing minerals found in economic granite- or
- 171 pegmatite-related Li deposits. Li wt. % is based on stoichiometry.

Mineral Name	Mineral Group	Chemical Formula	Theoretical Li content (wt. %)
Spodumene	Inosilicate (pyroxene)	LiAlSi2O6	3-7
Lepidolite	Phyllosilicate (mica)	K ₂ (Li,Al) ₅₋₆ (Si ₆₋₇ Al ₂₋₁ O ₂₀)(OH,F) ₄	1.4 to 3.6
Petalite	Phyllosilicate	LiAlSi ₄ O ₁₀	1.6 to 2.3
Amblygonite/ Montebrasite	Phosphate	LiAl(PO ₄)(F,OH)	3.4 to 4.7
Hectorite	Phyllosilicate (clay mineral)	Na ₀₋₃ (Mg,Li) ₃ Si ₄ O ₁₀ (OH) ₂	0.54
Jadarite	Nesosilicate	LiNaSiB ₃ O ₇ (OH)	7.3

173 Sixteen large (>2 cm) and visually homogeneous pegmatite-hosted samples of spodumene (n 174 = 9), petalite (n = 4) and lepidolite (n = 3) were selected for our study. The samples were cut 175 in half using a diamond saw. One half with a flat surface was retained for LIBS analyses and 176 the other half was crushed using an agate mortar to obtain a powder fraction for independent 177 chemical analysis. These powders were analyzed at the SARM laboratory (CRPG-CNRS, Nancy, France) using inductively coupled plasma atomic emission spectroscopy (ICP-AES) for 178 179 major-element analysis, ICP-MS for trace-element analysis, and flame atomic absorption 180 spectrometry (AAS) for Li analysis. Bulk mineral geochemical compositions are provided in 181 Table S1, and the Li contents, which may differ from the theoretical values, are reported in 182 Table 2, and location information of the samples and their structural formulae. Measured bulk 183 concentrations were used as reference concentrations to be compared to the LIBS analyses 184 undertaken on the second sample piece.

185

186 Table 2

187 Mineralogy, provenance and measured Li contents of the samples (AAS). The error obtained

188 for the Li content is 5% (relative percentage)

Sample ID	Mineral	Country of Origin	Li content (wt. %)
Kun-Afg	Kunzite (Spodumene)	Afghanistan	3.54
Spo-Kok	Spo-Kok Spodumene South Afr Namaqualand- Kok		3.31
Spo-Lac	Spodumene	Canada	2.83
Hid-Afg	Hiddenite (Spodumene)	Afghanistan	3.36
Spo-Nor	Spodumene	South Africa Namaqualand-Norabees I	3.25
Spo-Big	Spodumene	Canada/Big Lake	3.17
Spo-Boa	Spodumene	Boa Vista, Portugal	3.46
Spo-Pak	Spodumene	Pakistan	3.36
Spo-Zai	Spodumene	Zaire	1.33
Pet-Bik2	Petalite	Zimbabwe/Bikita	2.04
Pet-Bik	Petalite	Zimbabwe/Bikita	2.16
Pet-Big	Petalite	Canada/Big Lake	2.01
Pet-Tan	Petalite	Canada/Tanco	2.18
Lep-Moz	Lepidolite	Mozambique	2.57
Lep-Cun	Lepidolite	Mauritania	3.02
Lep-Pak	Lepidolite	Pakistan	2.47

As illustrated by the pictures of typical Li minerals shown in a Figure S1, the appearance of these minerals cannot be used for discrimination, as color varies varied from translucent to white, with some brown to pink variants. In order to avoid any bias in the analysis and to obtain the best estimate of Li content, the selection of samples was done on the basis of their visual homogeneity, which may be an indication of uniform composition. As the lepidolite micas have a sheet-like structure, it was not necessary to cut these samples to obtain a flat surface for the LIBS analysis.

197

198 *3.2 LIBS technique and protocol of analysis*

The handheld LIBS analyzer used exclusively for this study was a SciAps model Z-300 analyzer that utilizes a 1064 nm laser source featuring a 5-6 mJ/pulse and 10Hz/50 Hz repetition rate. This instrument has three charge coupled diode (CCD) spectrometers that cover a broad spectral range from 190-950 nm, which permits the detection of one or more emission lines for elements of interest from the ultraviolet (UV) to visible and near-infrared (VNIR)

portions of the electromagnetic spectrum. The LIBS analysis has been also done under a 204 205 constant flux of Ar gas, with the canister, in order to confine the plasma emission and 206 increase its intensity. The analyzer is configured to permit rastering of the laser beam across a 2 207 mm segment of sample and the pattern of the raster grid can be configured to optimize the 208 number of laser points on the sample surface. The number of laser shots can be adjusted to 209 include a few "cleaning" shots prior to data acquisition if dust or an environmental stain is present on the sample. In the case of a surface coating, a LIBS-depth profile can be acquired by 210 211 taking a number of additional laser shots at any location. These specific characteristics of LIBS analysis are of particular relevance for in situ geologic analysis and exploration 212 213 application when a presence of coating is suspected. The data acquisition processes described above have be modified for different minerals depending on a sample's physical character and 214 size. For this study, the LIBS raster was acquired generally along a 3x5 point grid with four 215 216 laser shots for data acquisition made at each point following one cleaning shot. This rater grid 217 was duplicated on each sample and the 30 averaged LIBS spectra from the four laser shots 218 were aggregated to emulate a bulk analysis (Fig. 1). 219 Here, concerning the temporal conditions of the plasma analysis, the delay time was fixed as

following: the beginning of the acquisition of the LIBS emission is 400 ns and the acquisitiontime window is fixed to 1 ms.



Fig. 1 Schematic representation of the LIBS analysis for this study illustrating the two 3x5
raster grids that comprised of 15 points of laser spot size of 50 µm in diameter at a 12.5µ
spacing.

227

229

228 *3.3 Data processing*

230 As noted above, LIBS spectra were acquired over the broad spectral range from 190-950 nm. 231 The maximum emission intensity observed was typically around 30,000 a.u. (arbitrary units) 232 due to the strong Li emission. Emission line selection is the first step in LIBS spectral data 233 processing. In the case of Li, few electronic transitions are possible, which limits the number of 234 emission lines present in a LIBS spectrum. The strongest lines are those at 610.36 nm, 670.79 235 nm and 812.64 nm. For this study, only the lines at 610.36 nm and the 670.79 nm were used, 236 as these lines exhibited the best signal to noise ratio, displayed simple Voigt profiles, and had 237 no interferences with other peaks (Fig. 2). As the minerals in our study are not Ca-rich 238 (Guezenoc et al., 2019b), the strong interference by the Ca line at 610.23 nm was not observed. 239 Spectra can be exported from the instrument either as individual raw spectral or as averaged 240 spectra, with both options explored in this study. Even if the spectra, which are calibrated in wavelength, are minimally processed after acquisition by the internal software of the analyzer,
these spectra can be used for elemental detection (such as the identification of Be and Cs) or
for developing a calibration curve. Examples of mean LIBS spectra for the petalite,
spodumene and lepidote minerals, in the UV, VIS and VNIR ranges, are reported in the
supplementary figures (S2, S3 and S4).

Concentration estimation by calibration curves is one of the most common aspects of LIBS spectral analysis. This approach to quantification is quite sensitive and depends strongly on the analysis protocol and data processing (Babos et al., 2019; Bennett et al., 2018; Fabre et al., 2014, 2011). The information required for calibration curve development is the total area under a peak, also referred as the LIBS area intensity, which is expressed by the analyzer in arbitrary units (a.u.).





Fig. 2 Averaged LIBS spectrum obtained from analysis of 30 points on a spodumene sample with blowups presented for the two main Li lines (610.36 and 610.79 nm) that were used for the development of the calibration curve for estimation of Li concentrations. The grey area corresponds to the value used for the calibration.

257

258 **4. Results**

259 *4.1 Univariate calibration curve*

260 In order to evaluate the accuracy of our calibration model for Li, figures of merit were 261 calculated according to Equations 1 and 2. These are the root mean squared error of 262 prediction (RMSE-P) and the squared multiple correlation coefficient (R²). RMSE is a 263 commonly used statistical parameter to compare different models with similar analytical 264 conditions and for a particular variable, as it is scale dependent. In case of linear regression, 265 RMSE provides an indication of accuracy in the model, translating the standard deviation of prediction errors to evaluate the degree of data concentration around the line of best fit. R² 266 is also called the "coefficient of determination" and is an indicator of the linear relationship 267 268 between the X and Y variables, an R² value close to 1 indicates a good degree of fit.

Equation 1

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- 271
- 272
- 273

 $R^{2} = \frac{\sum_{i} (y_{pred} - y_{m})^{2}}{\sum_{i} (y_{red} - y_{m})^{2}}$ Equation 2 274 275

 $\text{RMSE} - \text{P} = \sqrt{\frac{1}{n} \sum_{i=1}^{n} (y_{pred} - y_{ref})^2}$

- The term y_{pred} is the predicted value from the calibration model, y_{ref} is the reference value from 276 277 bulk analysis, and y_m is the mean value of the predictors. As recently recalled (Duponchel et 278 al., 2020), a univariate calibration curve is the graphical display of a regression model in which 279 two variables of the same statistical weight are plotted against each other to evaluate the 280 relationship (i.e. correlation) between them. The situation where more than one X variable is 281 simultaneously used to forecast a single dependent Y variable (or vice versa) is a multivariate 282 model. The calibration curve approach used for this study was the univariate model, where X 283 corresponds to the emission area of a selected Li peak in the LIBS spectrum and Y to the independently determined wt. % Li concentration in a sample. 284
- 285

286 4.2 Lithium calibration on Li minerals

Calibration curves can be developed internally on the Z-300 using its proprietary Profile 287 288 Builder software based on the selection of different element emission peaks, which may be 289 used as collected or after normalization of a spectrum to its entire intensity or through the 290 standard normal variation (SNV) method. This is common practice when analyzing different 291 types of minerals by LIBS to ameliorate potential matrix effects resulting from difference in 292 physical character such as surface texture and hardness (El Haddad et al., 2014; Guezenoc et 293 al., 2019a; Zorov et al., 2010). For this study, total peak areas for both the 610.36 nm and 670.79

294 nm Li spectral lines were extracted and normalized to the entire LIBS spectra. Spectral 295 normalization is a part of the onboard software of the analyzer that can be selected by the 296 analyst. This may not have been necessary for the hard minerals spodumene and petalite, but 297 it should improve the results as the mica lepidolite displays a very different surface texture and 298 hardness. The Profile Builder software first associates with each standard or sample its actual 299 bulk Li content. The total area under a selected Li peak was determined from the 30 averaged 300 analyses, with each sample analyzed across the two 3x5 shot grids. The calibration curve was 301 obtained by taking the area under the curve for each Li peak in the LIBS spectrum and dividing 302 this by the area under the entire LIBS spectrum. Other elements can be added by developing 303 multiple calibration models. This calibration curve is automatically saved in memory and is 304 made available for subsequent analyses on the field. Fig. 3 presents the lithium calibration curve 305 obtained for our suite of 16 Li minerals using the two major Li emission lines. The Y error bars 306 in the figure correspond to the error value derived from bulk chemical analysis of the second 307 piece of each sample, calculated from the flame AAS analyses. This value is a specific function 308 of the Li content of the sample (around 5% relative percentage). The X error bars correspond to 309 the standard error obtained for the 30 averaged spectra for each sample.

310



311

Fig. 3 Calibration curve obtained for the 16 Li-rich minerals. Error bars are shown for both the
LIBS analyses and the Li bulk contents (AAS). Symbols: diamond for spodumene, triangle for
lepidolite and square for petalite.

316 Part of the information that can be extracted from the LIBS handheld instrument and used for development of the calibration curve is reported in Table 3. The relative standard deviation 317 318 (RSD) values obtained on the averaged Li emission line intensities are around 13%, can be down to 9%, but generally do not exceed 18%. These variations have two possible causes. 319 320 The first is compositional heterogeneity within a sample, in that minerals may be 321 compositionally zoned in ways not observed by visual inspection but would be detected by the LIBS analysis as the laser beam spot size on the sample is 50 µm. The second is a matrix effect 322 related to likely differences in the volume of sample ablated by each laser shot at the different 323 324 points along the two raster grids. Both of these potential complications have been addressed by 325 making 30 LIBS analyses at two places on each sample in order to obtain a pseudo-bulk analysis 326 that somewhat representative of the entire sample. As shown in Fig. 3, there is a high degree of 327 correlation, with $R^2 > 0.9$ and RMSE-P = 0.19 wt. % for this Li mineral calibration. This means 328 that for each quantification of Li on new mineral, the error will be ± 0.19 wt. % Li, thus allowing 329 the discrimination of the mineral end-members and delivering very good Li quantification when 330 used in the field.

331

332 Table 3

Information from the Z-300 handheld LIBS analyzer used for the establishment of the Li
calibration curve. Li emission areas are reported as baseline corrected emission intensities in
arbitrary units (a.u.). RSD values are calculated for the 30 LIBS analyses of each mineral.
Reference Li contents (in wt. %) are from the independent analysis by flame AAS (5%). The
calculated Li contents are the values obtained from the LIBS calibration curve (Fig. 3).

Sample ID	Li Emission Intensity (a.u.)	RSD LIBS (%)	Reference Li Content (wt. %)	Error content (Li wt. %)	Calculated Li content (wt. %)
Spo-Zai	8952	17.6	1.33	0.06	1.64
Spo-Pak	17991	10.8	3.36	0.17	3.52
Spo-Nor	15648	11.3	3.25	0.16	3.03
Spo-Lac	14017	9.4	2.83	0.14	2.70
Spo-Kok	15908	10.8	3.31	0.16	3.09
Spo-Boa	16461	10.8	3.46	0.17	3.20
Spo-Big	17198	19.7	3.17	0.16	3.36
Pet-Tan	10751	11.6	2.18	0.11	2.02
Pet-Bik2	10337	15.4	2.04	0.10	1.93
Pet-Bik	11597	11.6	2.16	0.11	2.19

Pet-Big	10225	9.1	2.01	0.10	1.91
Lep-Pak	13923	13.8	2.47	0.12	2.68
Lep-Moz	14022	16.1	2.57	0.13	2.70
Lep-Cun	16959	18.1	3.02	0.15	3.31
Kun-Afg	17173	12.7	3.54	0.18	3.35
Hid-Afg	17603	14.3	3.36	0.17	3.44

340 *4.3 Discrimination of Li-bearing minerals*

341 Since LIBS has the potential of detecting most elements of the periodic table, it is common to 342 use LIBS analysis to discriminate mineral phases of similar composition (McMillan et al., 2006; Harmon et al., 2009; Farnsworth-Pinkerton et al., 2018). This capability would be 343 344 especially useful during fieldwork, where it might be used to rapidly distinguish between Li 345 minerals of similar appearance but different major element composition, as is the case for 346 spodumene and petalite. Although these minerals are generally difficult to distinguish visually, 347 they are characterized by distinct Si and Al proportions and, therefore, have different Si/Al 348 ratios that can be observed in LIBS spectra. The diagrams of Fig. 4 show LIBS spectra between 349 287-311 nm for petalite, spodumene, and lepitolite, displaying the emission line for Si at 288.16 350 nm and the two emission lines for Al at 308.21 and 309.27 nm. The relative difference in peak 351 intensities between Al and Si results in a high Si/Al ratio for petalite and low ratios for 352 spodumene and lepidolite. Indeed, a rapid look at such spectra on the LIBS analyzer in the field 353 could help to quickly discriminate a mineral like petalite with a relatively high Si/Al ratio from 354 spodumene with a low Si/Al ratio. Similarly, it should be possible to distinguish between petalite 355 and spodumene according to the intensity K emission lines at 766.49 nm and 769.89 nm (not 356 visible here). Lepidolites can be easily identified in the field due to their purple color. This 357 capability to rapidly identify elements and element ratios for minerals in the field represents one 358 of the primary utilities of handheld LIBS instrument.





Fig. 4 Representative LIBS spectra between 287–311 nm, with intensities expressed in
arbitrary units (a.u.), a) petalite, b) spodumene and c) lepidolite.



LIBS also has the potential to detect other light elements, such as Be, Rb and Cs, other rare metals that commonly occur in pegmatites along with Li. The detection of trace concentrations of such elements has important implications for Li exploration. Indeed, high Be, Rb, and Cs contents in muscovite micas can be used as indicator of Li-fertility, which permits discrimination between barren and Li-mineralized granites and pegmatites (Ballouard et al., 2020; Černý, 1989; Gordiyenko, 1971; Selway et al., 2005). Similarly, secondary or fluid-modified minerals, which result from the pervasive circulation of hydrothermal fluids that produces country-rock 372 aureoles of Li mineralization in metasomatized pegmatites and granites, can be strongly enriched in Li, Be, Rb, and Cs (Ballouard et al., 2020; Barros et al., 2020; Selway et al., 2005). 373 Typically, such enrichment is the most pronounced in the first 10s of meters away from the 374 country rocks surrounding mineralized intrusions. Li, Cs, and Rb concentrations in metasomatic 375 376 biotite and muscovite at the intrusion-country rock contact can commonly be between 1000 ppm and 1 wt. % (Ballouard et al., 2020) with such anomalies sometimes extending outward 377 for a distance of >1 km (Selway et al., 2005). Therefore, the direct measurement of trace and 378 379 minor elements in micas using handled LIBS in the field can be a powerful tool for rare metal 380 deposit exploration.



381

Fig. 5 LIBS spectra over selected wavelength intervals displaying Rb, Sr, Cs and Be most intensive emission lines. The fine dotted line denotes lepidolite, the dashed line denotes petalite and the solid line denotes spodumene. The LIBS spectra are averaged spectra obtained on the 30 points for each mineral and they are not normalized. The legend for each plot indicates the elemental concentrations for Rb, Sr, Cs and Be. a) spectra between 775-795 nm for Rb, b) spectra between 406-409 nm for Sr, c) spectra between 845-860 nm for Cs, d) spectra between 310-316 nm for Be.

389

The LIBS spectra shown in Fig. 5 are averaged spectra for each of the three Li mineralsanalyzed, i.e. the lepidolite profile is the average spectrum for the three lepidolite samples. Their

392 contents, obtained from the independent bulk analysis, testify that the limit of detection are in 393 agreement with such low concentration. As these spectra are not normalized, it is not possible 394 to directly compare emission intensities of the trace elements among the different minerals. 395 According to these results, Be and Sr can be detected down to the ppm level, as previously 396 observed for Li (0.1 ppm estimated), Rb can be detected down to 20 ppm, and the detection 397 limit for Cs is estimated at around 15 ppm. These limits of detection for such low-abundance trace elements are in agreement to the concentrations commonly reported for such Li-minerals 398 399 (Selway et al., 2005; Van Lichtervelde et al., 2008).

400

401 **5. Discussion**

The analytical approach followed for this work was developed around a number of 402 403 interrogations, regarding the quite recent portable LIBS technology and its application to 404 geological material. Qualitative LIBS analysis has been the focus of most previous studies due 405 to high emissivity of atoms and ions in plasma conditions in reaction to the laser beams, and 406 most importantly for the wide range of detection, from a few ppms to percentage concentration 407 levels during the same analysis. These features are widely exploited, especially in the 408 microchemical mapping of mineral surfaces and various materials. (Sweetapple and Tassios, 409 2015a; Kuhn et al., 2016; Fabre et al., 2018; Jolivet et al., 2019; Lawley et al., 2021). 410 However, the quantitative aspect is still challenging because of both chemical and physical 411 effects related to the type of material analyzed and the type of LIBS used (Rosenwasser et al., 412 2001; Sobron, 2019; Harmon and Senesi, 2021). Similarly to conventional analytical methods, 413 standards with known physical and chemical properties have to be used to build predictive 414 models and provide concentration estimates.

415

416 This study started by testing different analytical parameters in order to optimize the LIBS 417 sample acquisition for the analyte and the samples to be tested. AS previously described, the 418 optimized number of laser shots is dependent on the chemical homogeneity of the sample and 419 the grain size of the sample in relation to the laser spot size. Regarding the size of the minerals, 420 one should multiply the number of the laser shots to obtain a bulk composition. As the collection 421 of spectra is fast, around 30s for each raster grid, the LIS analysis seems to be faster than for 422 handheld XRF analyzers, with the option to directly control the LIBS spectra are attractive 423 aspects for exploration purpose. Similarly, minimal sample preparation is also an advantage avoiding any use of powders or pressed pellets, that are sometimes required for quantitative 424 425 analysis for other portable instrument such as XRF (Gazley and Fisher, 2014; Lemière and 426 Harmon, 2021). In addition to being able to use a priori compositional knowledge of the 427 standards used for calibration, the portable LIBS instrument allows the analyst to define and 428 compare different calibration tests while varying the spectra acquisition parameters (e.g. 429 number of analysis points, number of laser shots, laser repetition rate, ...etc.). The statistical 430 parameters of RMSE-P (as displayed by error bars in the calibration curve of Fig. 3) and RSD, 431 which are calculated by the internal software, are strong indicators regarding the accuracy of the analytical results and, consequently, the calibration model. For exploration context, this 432 value of 0.19 wt% is sufficient to discriminate the different minerals and to provide a good 433 434 estimate of the Li content. The use of handheld LIBS is of first interest to limit the sampling 435 with the good identification of Li-bearing minerals and to limit the future chemical analyses.

436

437 The LIBS analyses were obtained on a single LIBS analyzer after internal wavelength 438 calibration, with spectral pre- processing and processing performed with its onboard software. 439 It is worth noting that LIBS spectra acquired can be used as raw data, i.e. without any treatment 440 or correction, by exporting the data for external processing that gives the analyst more 441 options to pre-process and process data according to the particular objectives of a study. Pre-442 processing of LIBS spectra was performed via the integrated software, which is an important 443 capability for field use. For classical LIBS analysis, baseline and noise corrections are optional 444 depending on evaluation by the analyst of the LIBS spectra and emission lines present for 445 specific elements of interest. Normalization is also commonly used to attenuate the shot-to-446 shot fluctuations in laser energy and laser-sample interactions under the same acquisition 447 protocol. Total area normalization was used for this study, as different types of Li minerals were 448 analyzed. For our application, the normalization process only takes into account a potential 449 chemical effect, considering that all samples analyzed are naturally-occurring mineral 450 specimens (no grains or powders) and were subject to the same preparation.

451

452 The potential of LIBS for the detection and quantification of elements lighter than Mg is very 453 important for geological purposes. No other technique is able to rapidly analyze these elements 454 on unprepared samples. Hence, having such a capability in a handheld device is quite 455 revolutionary for field geologists. The use of LIBS spectra for material discrimination is also a 456 very useful field analysis capability, as it exploits abundance differences in detected elements 457 to gain insight into a sample's mineralogy, which can be especially useful during the early 458 stages of a mineral exploration and sampling campaign when the conventional macroscopic 459 tests for mineral identification (e.g. luster, color, streak, hardness, cleavage) are difficult to

460 perform due to sample size, shape, or degree of alteration degree. Due to the small size of the 461 laser spot, handheld LIBS can precisely analyze a small area of a sample of ~50µm diameter 462 and detect a wide range of major, minor, and trace elements, which allows the analyst to 463 correctly identify the sample's mineralogy through comparison against a pre-assembled library of 464 spectra and/or element ratio. LIBS analysis is rapid and the internal software ranks detected elements according to their "percentage of likelihood", which is a parameter that the software 465 466 determines by matching the emission lines detected lines in a spectrum of an element to an 467 internal database derived from the NIST Atomic Spectra Database (Kramida et al., 2018) and 468 then compares the intensity levels recorded for each line. This option was very useful for trace 469 elements as shown in Fig. 5. The detection of high-Li with a high K-content can help the 470 geologist to identify the lepidolite minerals in place. Elements such as Be, Rb, and Cs were 471 detected in almost every analyzed mineral but usually with higher intensity levels in 472 lepidolite than in petalite and spodumene. The bulk analyses not only confirmed the presence 473 of these trace elements in the sample suite and documented that lepidolite was more enriched 474 in Rb especially and then Be and Cs. Importantly, this highlights that samples with only a few 475 10s of ppm abundance of these trace elements could be detected by the handheld LIBS 476 instrument.

477

478 **6.** Conclusions

479 Handheld LIBS is a powerful tool with a strong potential for both qualitative and quantitative 480 analysis. More such case studies are needed to realize its advantages for field use. This study began 481 with a focus on setting up the most appropriate LIBS analysis (multiplying the number of 482 points) and data processing protocol (internal normalization) for a dataset obtained for a suite 483 of 16 Li minerals comprised of three major mineral species (lepidolite, petalite and 484 spodumene). Thus, the resulting calibration curve showed that LIBS data compared well with 485 Li contents determined from conventional techniques with an RMSE of 0.19 wt. % Li. Although 486 very encouraging, the results of this study for Li quantification should be validated through analysis 487 of a larger collection of Li minerals. The discrimination aspect of LIBS is quite promising given 488 its capability for major, minor, and trace element analysis, which allows the analyst more 489 freedom in terms of element selection for discrimination depending on the particular 490 characteristics of different sample types. For this study, the major elements Si, Al, and K were detected and their measured relative intensities can be used for mineral discrimination. The ratio 491 492 between the Si and Al intensities can be observed to successfully discriminated petalite from 493 spodumene and lepidolite (Fig. 4). It was proposed that a similar comparison of K intensity

- 494 comparison could be used to identify lepidolite. As is the situation with quantitative analysis,
 495 use of LIBS for mineral discrimination should also be designed for the specific mineralogy of
 496 the sample type being investigated and the particular elements associated with target ore minerals.
 497 Considering that trace elements occur in various amounts depending on geological setting,
 498 handheld LIBS analysis could be used to discriminate mineral species of interest. Calibration
 499 curves can be developed for LIBS data from a portable device rapidly and efficiently, with its
- 500 use in an exploration program able to provide a considerable time and cost saving.
- 501

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