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Rate and State Friction as a Spatially Regularized Transient Viscous Flow Law

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Key Points:

- We reformulate the empirical rate and state friction law as a bulk viscous flow law in terms of anelastic shear strain rate.
- We show how mesh independence is achieved by including a gradient-like non-local anelastic shear strain rate equivalent.
- We show analytically and numerically that the proposed continuum model closely reproduces existing results of rate and state friction.

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Abstract

The theory of rate and state friction unifies field, laboratory, and theoretical analysis of the evolution of slip on natural faults. While the observational study of earthquakes and aseismic fault slip is hampered by its strong multi-scale character in space and time, numerical simulations are well-positioned to link the laboratory study of grain-scale processes to the scale at which rock masses move. However, challenges remain in accurately representing the complex and permanently evolving sub-surface fault networks that exist in nature. Additionally, the common representation of faults as interfaces may miss important physical aspects governing volumetric fault system behavior. In response, we propose a transient viscous rheology that produces shear bands that closely mimic the rate- and state-dependent sliding behavior of equivalent fault interfaces. Critically, we show that the expected tendency of the continuum rheology for runaway localization and mesh dependence can be halted by including an artificial diffusion-type regularization of anelastic strain rate in the softening law. We demonstrate analytically and numerically using a simplified fault transect that important aspects of the frictional behavior are not significantly affected by the introduced regularization. Any discrepancies with respect to the interfacial description of fault behavior are critically evaluated using 1D numerical velocity stepping and spring-slider experiments. Since no new physical parameters are introduced, our model may be straightforwardly used to extend the existing modeling techniques. The model predicts the emergence of complex patterns of shear localization and delocalization that may inform the interpretation of complex damage distributions observed around faults in nature.

Plain Language Summary

How, where, and when earthquakes nucleate is one of the great questions in science and society that, despite steady progress, has hardly been answered to any practical degree. Based on field observations, laboratory experiments, and theoretical work it is believed that a cocktail of escalating mechanical, chemical, and thermal grain-scale processes cause the sudden and rapid onset of earthquakes. The net effect of these processes are characterized by an immediate strengthening and a gradual weakening response to deformation and are unified in simplified form in the theory of ‘rate and state friction’. This theory is commonly used in computer simulations of earthquake sequences. We point out that rate and state friction, unlike some physical theories of earthquake

49 rupture, does not incorporate a diffusion process such as for example heat conduction.
 50 We show the introduction of an artificial diffusion process can prevent the mathemat-
 51 ical reduction of a fault zone to a two-dimensional interface while retaining the proper-
 52 ties of the original friction law. This in turn enables simulation techniques that rely on
 53 an interface-free description of the earth and promise to provide new insights into the
 54 spontaneous organization of seismic and aseismic phenomena in developing fault zones.

55 1 Introduction

56 Over the second half of the last century the study of the sliding behavior of fric-
 57 tional surfaces, such as those believed to occur in the earth's crust, has led to a general
 58 understanding that it is governed by competition of stabilizing viscous-like effects and
 59 potentially destabilizing processes that affect the texture of the interface in a time-dependent
 60 manner and self-organize into periods during which the interface is arbitrarily close to
 61 elastic stick punctuated with periods of anelastic slip (Rabinowicz, 1958; Brace & By-
 62 erlee, 1966; Bowden & Tabor, 1966; Scholz et al., 1972; Ida, 1972; Dieterich, 1978; Ru-
 63 ina, 1980, 1983; Dieterich & Kilgore, 1994).

64 1.1 Rate- and State-dependent Friction

65 This has led to the proposition of the phenomenological *rate- and state-dependent*
 66 *friction law* by Dieterich (1978, 1979a); Ruina (1980, 1983), which reads in its general
 67 form as

$$68 \quad f = f_0 + a \log V + b \log \Theta. \quad (1)$$

69
 70 In this law, the friction coefficient f is given by the linear combination of a reference fric-
 71 tion coefficient f_0 , a logarithmic contribution from dimensionless slip rate V multiplied
 72 by a coefficient a , and another logarithmic contribution from a dimensionless *state* vari-
 73 able Θ multiplied by a coefficient b . All quantities are taken to be positive.

74 The rate effect or *direct effect* is to a variable degree strengthening with increas-
 75 ing slip rate and thus exerts a stabilizing influence on the frictional interface for any strictly
 76 positive a . The state effect or *evolution effect* is to be governed by an evolution law that
 77 has the evolving steady state $\Theta_{ss} V(t) = 1$. Then, if $b > a$, this allows an externally
 78 loaded system with a frictional interface to move to a configuration with lower elastic
 79 strain energy and thus a frictional-mechanical instability may occur.

80 One of the most widely used state evolution laws is the *aging law* of Ruina (1980,
81 1983), given by

$$82 \dot{\Theta} = r_0 (1 - V\Theta), \quad (2)$$

84 with r_0 a rate constant that is commonly expressed by dividing the reference velocity
85 v_0 by a critical slip distance d_c . It should be noted that several other evolution laws that
86 govern the state variable have been proposed. The most commonly cited contender is
87 the *slip law* (Dieterich, 1979a; Ruina, 1980, 1983). In this work we do not consider it be-
88 cause its functional form ($\dot{\Theta} \propto \Theta V \log \Theta V$) is unfortunately not amenable to the type
89 of analytical treatment that is performed. We note that rate and state friction captures
90 only the first-order behavior of the stick-slip cycle and that laboratory experiments have
91 brought to light many secondary effects (e.g. Ruina, 1983; Marone, 1998; Mair & Marone,
92 1999; Chester, 1994; Passelègue et al., 2020).

93 Following Amonton’s law the friction coefficient is expressed as the ratio of shear
94 stress τ to normal stress σ , implying a cohesionless fault. It is further assumed that the
95 fault is always critically loaded, i.e. shear stress is equal to the shear strength and $V >$
96 0 . The assumption of criticality is necessary because (1) degenerates at $V = 0$, but has
97 also been suggested to be realistic for faults in nature (Bak & Tang, 1989).

98 The rate and state friction law is frequently applied in numerical studies of fault
99 slip, whether that be a study over the course of a single earthquake or a complex sequence
100 of slip transients (e.g. Dieterich, 1979b; P. G. Okubo, 1989; Rice, 1993; Rice & Ben-Zion,
101 1996; Ben-Zion & Rice, 1997; Lapusta et al., 2000; Lapusta & Liu, 2009; Rubin & Am-
102 puero, 2005; Gabriel et al., 2012; Jiang & Lapusta, 2016; D. Li & Liu, 2017; Erickson
103 et al., 2020). The use of interfacial friction laws in these models necessitates the treat-
104 ment of faults as mesh features on which internal boundary conditions can be applied.
105 It can be laborious to construct such meshes for non-trivial fault geometries, and the pro-
106 cedure does not scale well with increasing fault network complexity. There is currently
107 no established method to construct meshes with time-dependent geometry and topol-
108 ogy reflecting an evolving fault system, but we refer to K. Okubo et al. (2019) for work
109 in this direction. With some exceptions (e.g. D. Li & Liu, 2017; Sathiakumar et al., 2020;
110 Perez-Silva et al., 2021) many studies focus on the already complex behavior of long-term
111 sequences of slip transients on linear or planar faults in simple domains.

1.2 Continuum Fault Rheology

There is strong field and experimental evidence that faults in nature are not infinitely thin planes but consist of complex evolving networks of strongly localized shear zones within a wider region of damaged host rock (Tchalenko, 1970; Granier, 1985; Chester & Chester, 1998; Katz et al., 2004; Savage & Brodsky, 2011; Faulkner et al., 2011; Barth et al., 2013; Passelègue et al., 2016; Perrin et al., 2016; Locatelli et al., 2018, 2019; Pozzi et al., 2018, 2019, 2021; Ritter, Rosenau, & Oncken, 2018; Ritter, Santimano, et al., 2018; Ross et al., 2019). This knowledge has spurred a search for intermediate-scale continuum homogenizations of distributed micro-scale processes that obey the rate and state framework in a somewhat generalized sense. Currently proposed models focus on rheological feedback mechanisms involving temperature (Rice, 2006; Braeck & Podladchikov, 2007; John et al., 2009; Garagash, 2012; Roubek, 2014; Rice et al., 2014; Platt et al., 2014; Thielmann et al., 2015; Thielmann, 2018; Pozzi et al., 2021), grain size (Rozel et al., 2011; Thielmann et al., 2015; Thielmann, 2018; Barbot, 2019; Pozzi et al., 2021), porosity (Sleep, 1997; Niemeijer & Spiers, 2007; Chen & Spiers, 2016; Van den Ende et al., 2018), fluid pressure (Barbot & Fialko, 2010; Garagash, 2012; Rice et al., 2014; Platt et al., 2014; Poulet et al., 2014; Veveakis et al., 2014; Alevizos et al., 2014; Rattez, Stefanou, & Sulem, 2018; Rattez, Stefanou, Sulem, Veveakis, & Poulet, 2018), damage (Lyakhovskiy et al., 2011; Lyakhovskiy & Ben-Zion, 2014b, 2014a; Lyakhovskiy et al., 2016; Kurzson et al., 2019, 2020), granular physics (Daub et al., 2008; Daub & Carlson, 2008, 2009; Daub et al., 2010; Hermundstad et al., 2010; Elbanna & Carlson, 2014) or combinations thereof. Continuum models employing a rate and state formulation with the slip rate and tractions replaced by strain rate and stress invariants have been developed by Herrendörfer et al. (2018) and applied to self-organizing fault systems in Preuss et al. (2019, 2020); DalZilio et al. (2021); Behr et al. (2021). Since the state of the subsurface is typically difficult to accurately characterize on a regional basis, these models are important to develop a generic understanding of the long-term evolution of seismogenic fault systems.

A problem commonly associated with strain softening rheologies without internal length scale is unconstrained localization (Hobbs et al., 1990). As a shear zone localizes to the scale of a mesh element or grid cell they become poorly resolved and cause mesh dependence of the simulation. Models that possess an internal length scale perpendicular to the direction of shear may not suffer from this issue, provided that this length scale is resolved by the discretization. For example, as thermo-rheological feedback mech-

anism causes localization the heat diffusion term may gain prominence until localizing and delocalizing (e.g. diffusing) influences find a balance. Unfortunately some of the proposed fault rheologies don't possess an internal length scale, and some possess one that has physical meaning and can't be changed without changing the behavior of interest. If such a physical length scale is very small in nature, then that may place insurmountable constraints on the mesh resolution in two- or three-dimensional regional simulations (e.g. Rice et al., 2014; Platt et al., 2014). A few models *do* contain a controllable internal length scale; these are the unified rate and state friction theory of Sleep (1997) and the damage-breakage rheology of Lyakhovsky and Ben-Zion (2014b, 2014a); Lyakhovsky et al. (2016); Kurzon et al. (2019, 2020).

1.3 Objective

In summary, we state the need for a rigorous treatment of a continuum framework in which Dieterich–Ruina-type rate and state friction can be embedded based on the following observations:

1. Rate and state friction is a useful and powerful homogenization of the complex micro-scale processes that give rise to slip transients in nature.
2. Numerical modeling exploration of the complex and evolving relation that fault networks have with their tectonic environment is hampered by the long-standing challenges with mesh adaptive generation as well as lack of detailed knowledge of the structure of the subsurface.
3. The description of fault friction as a continuum process can be an elegant and practical way to avoid these problems provided they possess a controllable internal length scale in order to produce sensible results given reasonable computational resources.
4. A continuum faulting rheology may readily model important transient effects that might be missed in an equivalent interfacial description.
5. Much effort has been dedicated to understanding rate and state friction as an ingredient in numerical models – not all proposed continuum rheologies can equally benefit from this.

In response, we develop a Dieterich–Ruina-type continuum reformulation of rate and state friction that

- 175 1. possesses a controllable internal length scale λ_0 that enforces a constraint on lo-
176 calization,
- 177 2. yields mesh-independent results once λ_0 is sufficiently well resolved,
- 178 3. retains the parameters to the interfacial friction (1) and evolution (2) laws and
179 yields similar trajectories of V and f for the same choice of parameter values, but
180 that also
- 181 4. has interesting implications for fault zone behavior during and between earthquakes.

182 The model we propose here bears close relation to the earlier works of Sleep (1997) and
183 Herrendörfer et al. (2018). Contrary to Sleep (1997) we refrain from in-depth discussion
184 of fault physics but concentrate on the introduced spatial regularization, which is also
185 a necessary extension of Herrendörfer et al. (2018).

186 In Section 2 we will motivate our choice of constitutive model, focusing on the for-
187 mal links between interfacial and continuum models. In Section 3 we will present our
188 equivalent continuum rate and state rheology. In Section 4 we build our rheology into
189 a numerical model of a simplified one-dimensional fault transect to demonstrate both
190 its localization and delocalization behavior as well as its similarities and differences with
191 respect to the original laws of Dieterich and Ruina. We close with a discussion – in which
192 we will comment on possible links to existing theories of the physics of faulting – and
193 a summary in Sections 5 and 6.

194 2 Material Model and Continuum Mechanics

195 We will first describe how the existing interfacial description of rate and state fric-
196 tion can be embedded in a continuum mechanical framework with internal discontinu-
197 ities, and then generalize this to internal friction and distributed anelastic deformation
198 in an interface-free model.

199 2.1 A Fault Reference Frame

200 We consider the immediate neighborhood $\Omega \subset \mathbb{R}^3$ of a section of fault $\Gamma \subset \Omega$
201 (Figure 1) that is described as the set of points belonging to a differentiable 2D man-
202 ifold. We invoke the existence of a surjective function $\vec{x}_f(\vec{x})$ that maps a coordinate $\vec{x} \in$
203 Ω to its nearest Euclidian neighbor $\vec{x}_f \in \Gamma$. Next, we define the fault transect $T(\vec{x}_f) \subset$
204 Ω as the 1D curve that includes all the points $\vec{x} \in \Omega$ that map to a particular fault co-

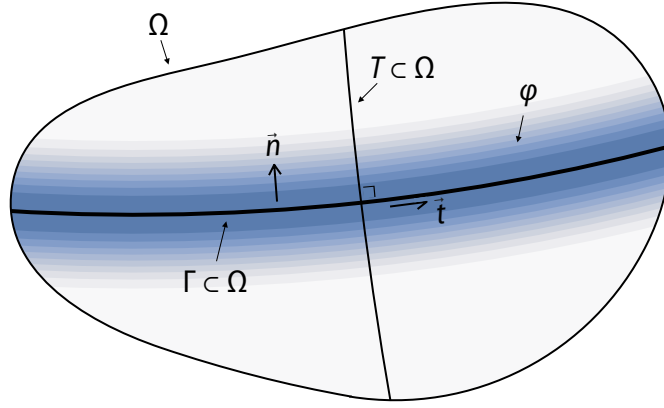


Figure 1. A region Ω in the enclosing a portion of a fault interface Γ , with normal and tangent vectors \vec{n} and \vec{t} indicated. A transect T is drawn perpendicular to the fault. The fault interface may be substituted by a shear band on which anelastic shear strain is distributed according to $\varphi(\vec{x})$, as indicated by the colored contours.

ordinate $\vec{x}_f \in \Gamma$, in short $T = \vec{x}_f^{-1}(\{\vec{x}\}) \subset \Omega$. Finally, we let $\hat{n}(\vec{x}) \in \mathbb{R}^3$, $\vec{x} \in \Omega$ be a unit vector field locally tangent to the transecting curve $T(\vec{x}_f(\vec{x}))$ and therefore normal to the fault.

The *state* of the body Ω is described by a symmetric Cauchy stress tensor field $\boldsymbol{\sigma}(t, \vec{x}) \in \mathbb{R}^3 \otimes \mathbb{R}^3$ and a velocity vector field $v_0 \vec{v}(t, \vec{x}) \in \mathbb{R}^3$. Their governing equations are stated after the introduction of some further scaffolding. Let t denote time.

We define the normal traction $\sigma = -\hat{n} \cdot \boldsymbol{\sigma} \cdot \hat{n}$ (positive in compression), the shear traction vector $\vec{\tau} = \boldsymbol{\sigma} \cdot \hat{n} + \sigma \hat{n}$, and the shear traction magnitude $\tau = \sqrt{\vec{\tau} \cdot \vec{\tau}}$. The rupture process is most efficient when the slip rate vector $\vec{V} \in \mathbb{R}^3$ is parallel to the in-plane shear traction vector $\vec{\tau}$, and so alongside the full velocity field \vec{v} we define a purely slip-induced velocity double couple field \vec{v}_{\equiv} around the coordinate \vec{x} as

$$\vec{v}_{\equiv}(t, \vec{x}) := \frac{1}{2} V(t, \vec{x}_f) \hat{t}(\vec{x}_f) \operatorname{sgn}(\hat{n}(\vec{x}_f) \cdot [\vec{x} - \vec{x}_f]), \quad (3)$$

with $\hat{t} \in \mathbb{R}^3 := \vec{\tau} / \tau$ the unit tangent vector and sgn the sign function, which has $\operatorname{sgn}(0) = 0$.

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2.2 Assumptions

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In the following, we will assume that continuum processes that occur around the fault are predominantly confined to the small neighborhood Ω , which itself includes only a small section of a whole fault. The neighborhood Ω is assumed to be sufficiently small that in-plane variations of fault properties such as state, slip rate, and curvature are negligible. This assumption comes at a loss of generality, especially around fault kinks, branches, and tips, but does not necessarily limit the applicability of our formulation. We will revisit this assumption in the Section 5.5 of the Discussion.

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We also assume that the material has uniform static properties throughout the domain. This means that the gouge layer is effectively assumed to be infinitely wide, or alternatively that the only quality that distinguishes the fault gouge from the host rock is its ‘state’. Undoubtedly there is a collection of state variables that govern the long-term structure of fault zones in nature that are not modeled here. We will revisit this assumption too in Section 5.5, but in the mean time note that its impact is largely confined to inferences made about the de-localization behavior of rate strengthening faults in Section 3

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2.3 Generalization of a Fault to a Shear Band

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We now generalize the fault interface to a distributed shear band by substituting a smoothed step function $\Phi = \Phi(t, x; \lambda_0) \in [-1, +1]$, $\lambda_0 > 0$ (e.g. $\Phi(t, x; \lambda_0) := \tanh(x/\lambda_0)$) for the sign function in (3). We require that the chosen function converges in a point-wise manner to the sign function as $\lambda_0 \rightarrow 0^+$, such that (3) can be seen as the result of taking the limit

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$$\begin{aligned} \vec{v}_{\rightleftharpoons}(t, \vec{x}; \lambda_0) &:= \frac{1}{2} V(t, \vec{x}_f) \hat{t}(\vec{x}) \Phi \left(t, \hat{n} \cdot \left[\frac{\vec{x} - \vec{x}_f}{\lambda_0} \right] \right) \\ \vec{v}_{\rightleftharpoons}(t, \vec{x}) &:= \lim_{\lambda_0 \rightarrow 0^+} \vec{v}_{\rightleftharpoons}(t, \vec{x}; \lambda_0). \end{aligned} \quad (4)$$

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Recognizing that fault slip is an anelastic process, the anelastic strain rate tensor $\dot{\epsilon}_{-e}$ is defined to be the symmetric gradient ∇^s of the (continuous) slip-induced velocity field $\vec{v}_{\rightleftharpoons}(t, \vec{x}; \lambda_0)$:

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$$\begin{aligned} \dot{\epsilon}_{-e}(t, \vec{x}) &:= v_0 \nabla^s \vec{v}_{\rightleftharpoons}(t, \vec{x}; \lambda_0) := \frac{1}{2} v_0 [(\nabla \vec{v}_{\rightleftharpoons})^T + (\nabla \vec{v}_{\rightleftharpoons})] (t, \vec{x}; \lambda_0) \\ &\approx \frac{1}{2} \frac{v_0}{\lambda_0} V(t, \vec{x}_f) \varphi \left(t, \hat{n} \cdot \left[\frac{\vec{x} - \vec{x}_f}{\lambda_0} \right] \right) [\hat{t} \otimes \hat{n} + \hat{n} \otimes \hat{t}] (\vec{x}) =: \gamma_0 \gamma(t, \vec{x}) \hat{s}(\vec{x}). \end{aligned} \quad (5)$$

251 Here the *equivalent anelastic shear strain rate* (a scalar value) has been introduced as

$$252 \quad \gamma_0 \gamma(t, x) = v_0 V(t) \varphi(t, x), \quad (6)$$

253
 254 with reference value $\gamma_0 = v_0/\lambda_0$ and *strain rate distribution* $\varphi(t, x; \lambda_0) = \partial_x \Phi(t, x; \lambda_0)$
 255 (illustrated in blue shading in Figure 1). The symbol $\hat{\mathbf{s}}$ denotes the Schmidt tensor $\hat{\mathbf{s}} =$
 256 $\frac{1}{2} [\hat{\mathbf{t}} \otimes \hat{\mathbf{n}} + \hat{\mathbf{n}} \otimes \hat{\mathbf{t}}]$, and its *Frobenius norm* $\|\hat{\mathbf{s}}\| = \sqrt{\text{tr} \hat{\mathbf{s}}^T \hat{\mathbf{s}}} = 1/\sqrt{2}$, which is an im-
 257 portant property to maintain when generalizing $\hat{\mathbf{s}}$ later on because it is compatible with
 258 the interpretation of $\gamma_0 \gamma$ as the fault-perpendicular derivative of fault-parallel velocity,
 259 i.e. as an accurate measure of simple shear. In the derivation of (5) the gradients of V
 260 and the orthonormal bases $\hat{\mathbf{t}}$ and $\hat{\mathbf{n}}$ have been dropped under the assumption set out in
 261 Section 2.2 that they are small compared to the gradient of Φ . For this assumption to
 262 be met, the across-fault length scale λ_0 must be small compared to the along-fault length
 263 scales that exists in the interfacial rate and state formulation.

264 Since the derivative of half a smoothed step function $\frac{1}{2} \Phi(x) \in [-\frac{1}{2}, +\frac{1}{2}]$ has the
 265 properties of a distribution – is a non-negative and importantly integrates to unity over
 266 its domain – we may express the integral relation

$$267 \quad \mathcal{V}(\gamma) := \gamma_0 \int_T \gamma(t, \vec{x}) dx = v_0 V(t, \vec{x}_f) \quad (7)$$

268
 269 over the fault transect T , which is again compatible with the notion that $\gamma_0 \gamma$ is a mea-
 270 sure of simple shear in the fault reference frame. We note that the above integral strictly
 271 requires that all anelastic strain occurs over the length of the fault transect T , imply-
 272 ing that the anelastic strain rate distribution has compact support in space. In practice
 273 we can relax that requirement if said distribution is narrow and thin-tailed.

274 Subtracting the anelastic shear strain rate tensor $\dot{\epsilon}_{-e}$ from the full strain rate ten-
 275 sor $\dot{\epsilon} = \nabla^s \vec{v}$ results in the elastic strain rate tensor. This additive decomposition of elas-
 276 tic and anelastic strains is known as the Maxwell model. Application of Hooke’s law of
 277 linear elasticity to the elastic strain rate tensor leads to

$$278 \quad \dot{\boldsymbol{\sigma}} = \mathbf{S} [\dot{\boldsymbol{\epsilon}} - \gamma_0 \gamma \hat{\mathbf{s}}] \quad (8)$$

279
 280 with fourth-order elastic stiffness tensor \mathbf{S} and the dot over a symbol denoting the time
 281 derivative. The relation above is standard in the context of elasto-plasticity (e.g. Mühlhaus
 282 & Aifantis, 1991), in which $\gamma_0 \gamma$ is known as the plastic multiplier and $\hat{\mathbf{s}}$ coincides with
 283 the derivative of the magnitude of shear traction τ with respect to the full stress ten-

284 sor $\boldsymbol{\sigma}$. In this case τ is to be interpreted as the *non-associated* plastic potential of a ma-
 285 terial that does not undergo anelastic volume change.

286 For completeness we list ordinary differential equation that governs \vec{v} , the *momentum balance equation*:
 287

$$288 \quad v_0 \dot{\vec{v}} - \vec{g} = \rho^{-1} \nabla \cdot \boldsymbol{\sigma}. \quad (9)$$

289 Here \vec{g} is the acceleration due to gravity and ρ the mass density. Gravity may be ignored
 290 in the remainder of this work at no loss of generality and the inertial term $\dot{\vec{v}}$ may be ig-
 291 nored at *some* loss of generality under the assumption of static momentum balance.
 292

293 2.4 Plasticity and Coordinate Invariance

294 For a spontaneously developing fault zone the shear and normal stress can not be
 295 defined in relation to a known plane. Plasticity models avoid this problem through the
 296 use of stress tensor invariants and scalar parameters. The Mohr-Coulomb plasticity model
 297 generates anelastic shear strain parallel to the Schmidt tensor $\hat{\mathbf{s}} = \frac{1}{2}(\hat{n} \otimes \hat{t} + \hat{t} \otimes \hat{n})$, in
 298 which \hat{t} makes an angle $\phi = \tan^{-1} f$ with respect to the largest principal stress axis $\vec{\sigma}_1$,
 299 and \hat{t} and \hat{n} form an orthogonal triad of vectors together with the intermediate princi-
 300 pal stress axis $\vec{\sigma}_2$. In other words, anelastic shear strain occurs in a plane perpendicu-
 301 lar to $\vec{\sigma}_2$. The corresponding cohesionless yield criterion is written as

$$302 \quad \frac{1}{2}(\sigma_1 - \sigma_3) = -\sin(\phi) \frac{1}{2}(\sigma_1 + \sigma_3), \quad (10)$$

303 with σ_1 and σ_3 the magnitudes of the largest and smallest principal stresses $\vec{\sigma}_1$ and $\vec{\sigma}_3$.
 304 The quantity $-\frac{1}{2}(\sigma_1 + \sigma_3) = \sigma_m$ is called the *mean stress* and $\frac{1}{2}(\sigma_1 - \sigma_3) = \sigma_d$ the
 305 *effective deviatoric stress*. We note that the angle of a shear band that is generated does
 306 not need to be parallel to the direction of shear strain (Vardoulakis, 1980; Vermeer, 1990;
 307 Marone et al., 1992; Kaus, 2010; Le Pourhiet, 2013), and moreover that a local change
 308 in friction is not expected to immediately alter the macroscopic fault angle (e.g. Preuss
 309 et al., 2019).
 310

311 In the Mohr-Coulomb model, the friction coefficient $f = \tan(\phi)$ of a yielding or
 312 even of a non-yielding material may be expressed from (10) as a function of the stress
 313 tensor $\boldsymbol{\sigma}$ as

$$314 \quad f(\boldsymbol{\sigma}) = \sigma_d(\boldsymbol{\sigma}) [\sigma_m(\boldsymbol{\sigma})^2 - \sigma_d(\boldsymbol{\sigma})^2]^{-\frac{1}{2}}. \quad (11)$$

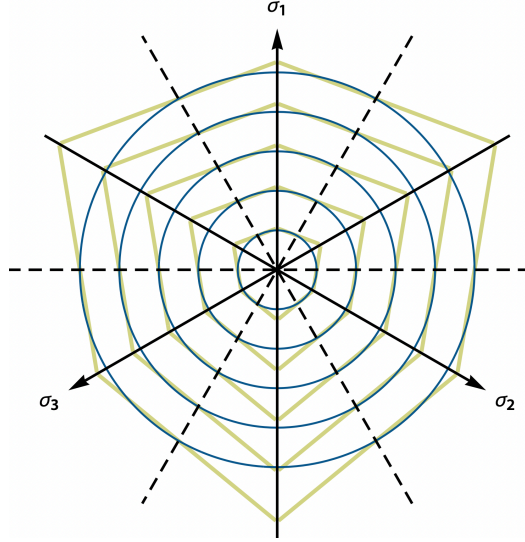


Figure 2. Comparison of Mohr-Coulomb (thick yellow lines) and Drucker-Prager yield envelopes (thin blue lines) at increments of pressure. The field of view is the octahedral plane of principal stress space. Projections of principal stress axes are indicated with arrows, and projections of planes of simple shear are indicated with dashed lines. A friction coefficient of 0.3 was used to generate this figure.

Herrendörfer et al. (2018) used the Drucker-Prager model (Drucker & Prager, 1952) as a simple and smooth approximation to Mohr-Coulomb plasticity. The model is defined in terms of the straightforwardly computable invariants

1. pressure $p = -\frac{1}{3} \text{tr } \boldsymbol{\sigma}$, and
2. effective shear stress τ_e proportional to the *Frobenius norm* of the deviatoric stress tensor $\boldsymbol{\tau} := \boldsymbol{\sigma} + p \boldsymbol{\delta}$, with $\boldsymbol{\delta}$ the Kronecker delta:

$$\tau_e := \frac{\|\boldsymbol{\tau}\|}{\sqrt{2}} = \sqrt{\frac{1}{2} \text{tr} (\boldsymbol{\tau}^T \cdot \boldsymbol{\tau})}. \quad (12)$$

In this model the existing definition of the Schmidt tensor $\hat{\boldsymbol{s}}$ is no longer applicable because the unit vectors \hat{t} and \hat{n} are not known *a priori*, and are not even uniquely constrained *a posteriori*. Keeping in line with plasticity theory, $\hat{\boldsymbol{s}}$ is redefined as the stress derivative of the non-associated plastic potential τ_e :

$$\hat{\boldsymbol{s}} := \frac{\partial \tau_e}{\partial \boldsymbol{\sigma}} = \frac{\boldsymbol{\tau}}{\tau_e}.$$

This definition still satisfies the criterion set out in the preceding section that $\|\hat{\boldsymbol{s}}\| = 1/\sqrt{2}$.

We calibrate the Drucker-Prager model to the Mohr-Coulomb model around a reference stress state $\boldsymbol{\sigma}_0$ of simple shear combined with isotropic compression, which encompasses all stress states possible in 2D plane strain and may be considered the most relevant stress state even in 3D tectonic settings. Let $\vec{\sigma}_0$ be given in ordered principal stress space by

$$\vec{\sigma}_0 := -p_0 \begin{bmatrix} 1 \\ 1 \\ 1 \end{bmatrix} + \tau_{e,0} \begin{bmatrix} +1 \\ 0 \\ -1 \end{bmatrix}.$$

At $\vec{\sigma} = \vec{\sigma}_0$, Mohr-Coulomb and Drucker-Prager measures of effective shear and normal stress coincide (up to a sign) and thus the Drucker-Prager model may be written as

$$\tau_e(\boldsymbol{\sigma}) = \sin(\phi) p(\boldsymbol{\sigma}). \quad (13)$$

Away from $\vec{\sigma}_0$ (and towards a uniaxial stress state) the Drucker-Prager yield surface becomes an increasingly worse approximation to the Mohr-Coulomb yield surface (Figure 2). Higher-order approximations like the Willam-Warnke yield envelope (Ulm et al., 1999) may be considered too. The yield surfaces discussed in this section serve as part of a general model proposition – in the one-dimensional numerical experiments considered in this work only the two effective stresses of Amonton’s friction law are defined.

Similar to the Mohr-Coulomb model (11), the friction coefficient $f = \tan \phi$ of the Drucker-Prager model may also be expressed from (13) as a function of the stress tensor $\boldsymbol{\sigma}$ as

$$f(\boldsymbol{\sigma}) = \tau_e(\boldsymbol{\sigma}) [p(\boldsymbol{\sigma})^2 - \tau_e(\boldsymbol{\sigma})^2]^{-\frac{1}{2}}. \quad (14)$$

In both the Mohr-Coulomb model and the Drucker-Prager model, the friction coefficient becomes ill-posed whenever the effective shear stresses σ_d or τ_e exceed the effective normal stresses σ_m or p . We have found a practical solution to be to add a constant value to the effective normal stresses, which for steady-state friction behaves as a cohesive strength.

Despite adopting a plasticity framework to enforce the frictional yield constraint we emphasise that the resulting model is still best regarded as ‘frictional viscous’ rather than frictional plastic because it lacks a distinct boundary between yielding and non-yielding states in space and time – the yield constraint (an equality, not an inequality) is enforced everywhere and anytime, and at any stress.

3 Continuum Evolution Law

Our goal is to find a continuum state evolution law that produces nearly the same history of loading and sliding as the original formulation of Dieterich and Ruina on a discrete fault. Like $V(t)$ is a ‘global’ measure of $\gamma(t, x)$ over a fictitious fault transect T , we introduce a local variable $\theta(t, x)$ of which $\Theta(t)$ is a global measure. We target continuum friction and evolution laws expressed in terms of γ and θ of the form

$$f = f_0 + a \log \gamma + b \log \theta, \quad (15)$$

$$\dot{\theta} = r_0 (c_2 - c_1 \mathcal{M}(\gamma) \theta), \quad (16)$$

with \mathcal{M} a non-local interaction operator that will be further elaborated in Section 3.3 and c_1 and c_2 additional coefficients that are required to calibrate (15) and (16) to their respective interfacial equivalent. The local friction and evolution laws (15) and (16) retain the structure of their interfacial counterparts (1) and (2) and generalize the continuum formulations of Sleep (1997) and Herrendörfer et al. (2018). We will revisit this connection in the Section 5.4 of the Discussion. The targeted form of the continuum equations does not come out of the blue. In coming sections we hope to make clear how it arises.

3.1 Analytical Framework

By

1. reorganizing the rate and state friction law (1) into a definition of $\Theta(V, f)$,
2. differentiating (1) with respect to time and reorganizing it into a definition of $\dot{V}(t)$,
and
3. substituting the definition of $\Theta(V, f)$ and the definition of $\dot{\Theta}$ that follows from the aging law (2) into the definition of $\dot{V}(t)$,

we arrive at the ordinary differential equation

$$a \dot{V} = b r_0 [V^2 - V \Theta(V, f)^{-1}] + V \dot{f}, \quad (17a)$$

$$\Theta(V, f)^{-1} := V^{\frac{a}{b}} \exp(-[f - f_0]/b) \quad (17b)$$

in which the friction coefficient $f = f(\boldsymbol{\sigma}; \hat{t}, \hat{n})$ is treated as a property of – rather than a constraint on – the stress tensor $\boldsymbol{\sigma}$.

By applying the same procedure to the proposed continuum friction and aging laws (15) and (16) we obtain the analogous ODE for γ :

$$a \dot{\gamma} = b r_0 [c_1 \gamma \mathcal{M}(\gamma) - c_2 \gamma \theta(\gamma, f)^{-1}] + \gamma \dot{f}, \quad (18a)$$

$$\theta(\gamma, f)^{-1} := \gamma^{\frac{a}{b}} \exp(-[f - f_0]/b), \quad (18b)$$

in which $f = f(\boldsymbol{\sigma})$ is given by the yield criterion (14) and $\dot{f} = \dot{f}(\boldsymbol{\sigma}, \dot{\boldsymbol{\sigma}})$ is expressible using the same yield criterion and the elastic constitutive equation (8).

Equations (17a,b) and (18a,b) are useful in mathematical and numerical analysis of the problem because they eliminate a variable and an algebraic constraint, and at the same time provide an ideal reference frame for establishing, and if necessary influencing, the ability of (18a,b) to reproduce the predictions made by (17a,b) in the sense of the integral relation (7).

In the following derivation we will assume that the field of effective friction $f(\vec{x})$, $\vec{x} \in \Omega$ in the neighborhood of a point \vec{x}_f on the fault is the same whether it is generated by a hard or a soft discontinuity. This assumption ties into the principal assumption set out in Section 2.2, which is revisited in Discussion section 5.5. We will also assume that $\mathcal{M}(\square(t) \diamond(x, \dots)) = \square(t) \mathcal{M}(\diamond(x, \dots))$. This assumption is satisfied in Section 3.3. Combining (17a,b), (18a,b), (6), and (7) leads to the following useful relation that describes the evolution of the anelastic shear strain rate distribution over time:

$$\dot{\varphi} \propto V [\lambda_0 c_1 \varphi \mathcal{M}(\varphi) - \varphi] - \Theta(V, f)^{-1} [\lambda_0^{\frac{a}{b}} c_2 \varphi^{1+\frac{a}{b}} - \varphi]. \quad (19)$$

This differential equation tells us several things. Most importantly, for γ to remain in accordance with V through the integral relation (7) over the transect T , the same integral of the left hand side of (19) must be zero. This requirement can only be met at arbitrary (V, f) if both the term multiplied by V and the term multiplied by $\Theta(V, f)^{-1}$ integrate to zero over T . This in turn can not in general be guaranteed unless

$$c_1 := \mathcal{C}_1(\varphi) := \lambda_0^{-1} \left[\int \varphi(x) \mathcal{M}(\varphi)(x) dx \right]^{-1}, \quad (20a)$$

$$c_2 := \mathcal{C}_2(\varphi) := \lambda_0^{-\frac{a}{b}} \left[\int \varphi(x)^{1+\frac{a}{b}} dx \right]^{-1}. \quad (20b)$$

We consider it undesirable for coefficients to depend in a time-dependent way on integrals of the modeled quantities and therefore will be restricting our attention to specific regimes of interest during which the values of the coefficients $c_{1,2}$ can be predicted analytically. We then employ those predicted values as model-specific constants in time and

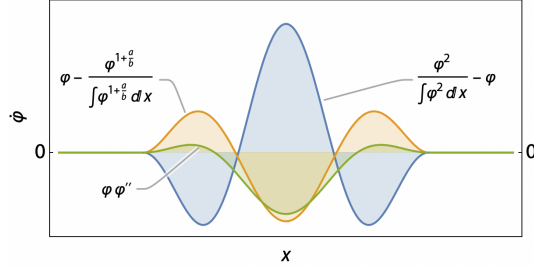


Figure 3. Influence of the different terms in Eq. (19) on the time derivative $\dot{\varphi}$ under the assumption that $\mathcal{M}(\varphi) = \varphi$. The first term (blue) can be seen to promote localization, while the second term (orange) favors the opposite. Also shown in this figure in green is the delocalizing influence of the nonlinear Laplacian term $\varphi(x)\varphi''(x)$, which is introduced in Sections 3.3 and 3.4. The input function φ follows a cosine-squared distribution.

426 space. We accept that this incurs a potential error whenever the state of the model is
 427 outside the selected regime, and will critically evaluate this error using numerical mod-
 428 els in Section 4.

429 Assuming the coefficients $c_{1,2}$ have been chosen appropriately and for simplicity
 430 that $\mathcal{M}(\varphi) = \varphi$, two additional observations can be made on the basis of (19):

- 431 1. The term that is multiplied by V promotes localization if $\mathcal{M}(\varphi) = \varphi$ because φ^2
 432 is a narrower distribution than φ (Figure 3).
- 433 2. The opposite (delocalization; Figure 3) holds true for the term that is multiplied
 434 by $\Theta(V, f)^{-1}$ on account of the corresponding minus sign and the requirement that
 435 $a, b > 0$.

436 This localizing and delocalizing behavior is of great interest and will be treated in de-
 437 tail in the following sections.

438 In the sections that follow we will assume to be operating on the fault-perpendicular
 439 transect T and will denote with the scalar coordinate x the distance with respect to the
 440 fault core, i.e. $x = \hat{n}(\vec{x}) \cdot [\vec{x} - \vec{x}_f(\vec{x})]$. This setting is in line with our assumption set
 441 out in Section 2.2 that across-fault variations in anelastic strain rate are more compact
 442 than its along-fault variations.

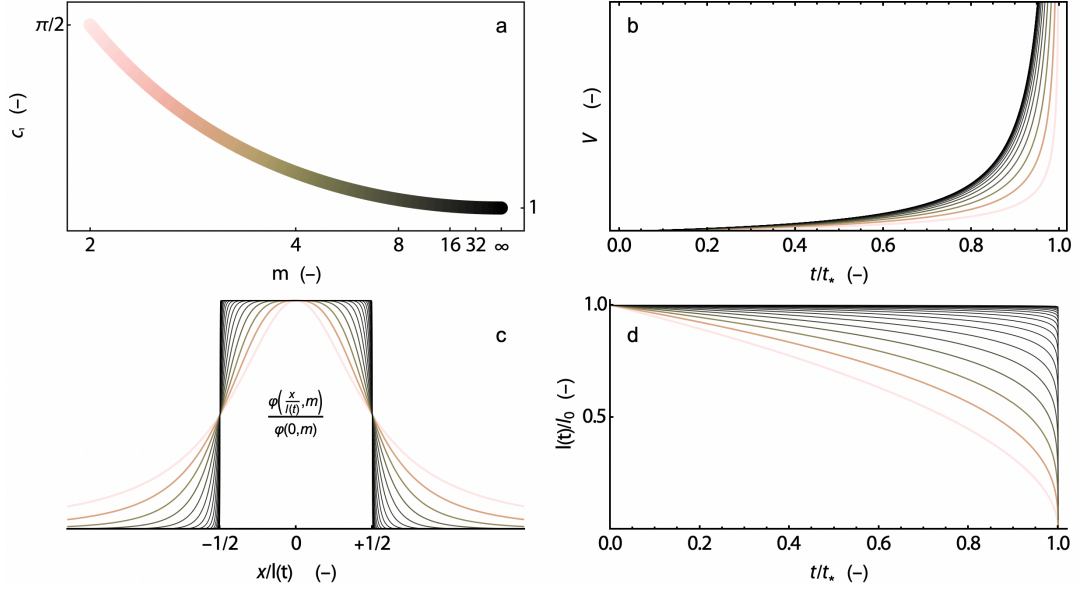


Figure 4. Asymptotic solutions during earthquake nucleation. a) Value of the dimensionless coefficient $c_1(m)$ versus distribution exponent m (see main text) represented on a reciprocal axis that ranges between 2 and ∞ . The thick line acts also as a color bar approximately indicating the value of m at which the individual lines in the remaining figures are plotted. b) Dimensionless slip rate V versus dimensionless time t/t_* to the slip rate asymptote for different values $m \in \{2^{1.0}, 2^{1.5}, 2^{2.0} \dots\}$. c) Nondimensionalized distributions $\varphi(x/l(t); m)/\varphi(0; m)$ versus dimensionless coordinate $x/l(t)$. d) Dynamic dimensionless length scale $l(t/t_*)/l_0$ versus dimensionless time t/t_* .

443

3.2 Runaway Sliding and Localization

444

445

Under rate-weakening conditions $a < b$, earthquakes occur as superexponential solutions to (17a,b). Prior to inertial damping the seismic slip rate V behaves as

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447

$$\dot{V} \propto V^2.$$

448

This ODE has the solution

449

450

$$V(t) = V(0) \left(1 - \frac{t}{t_*}\right)^{-1}, \quad (21)$$

451

in which $t_* \propto V(0)^{-1}$ is the time of the (hypothetical) singularity.

452

Assuming $\mathcal{M}(\square) = \square$, the same behavior occurs in (18a,b) in the limit

453

454

$$\dot{\gamma} \propto c_1 \gamma^2, \quad (22)$$

455

which is similarly solved by

456

457

$$\gamma(t, x) = \gamma(0, x) \left(1 - \frac{t}{t_*(x)}\right)^{-1},$$

458

in which $t_*(x) \propto (c_1 \gamma(0, x))^{-1}$.

459

We can make the following qualitative observations:

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1. small spatial variations in initial conditions $\gamma(0, x)$ can be amplified to infinity due to the fact that points with larger initial conditions on γ are closer to the strain rate asymptote at $t = t_*(x)$ than points with smaller initial conditions on γ , and
2. $\gamma(t, x)$ and $V(t)$ do not generally satisfy the integral relation (7) for all time $t \in [0, t_*)$.

465

Noting that this regime of runaway sliding corresponds in (19) to the limit

466

467

$$\dot{\varphi} \propto V [\lambda_0 c_1 \varphi \mathcal{M}(\varphi) - \varphi],$$

we can further illustrate the observations made above by choosing the class of solutions
(Appendix A)

$$V(t) \propto \left(1 - \frac{t}{t_*}\right)^{m^{-1}-1}, \quad (23a)$$

$$\varphi(t, x; m) := \left[l(t) c_1(m) \left(1 + \left| \frac{2x}{l(t)} \right|^m\right) \right]^{-1}, \quad m \in \mathbb{R} \geq 2 \quad (23b)$$

$$l(t) \propto V(t)^{(1-m)^{-1}}, \quad (23c)$$

$$\gamma_0 \gamma(t, x; m) := v_0 V(t) \varphi(t, x; m), \quad (23d)$$

$$c_1 = c_1(m) := \frac{\pi}{m} \operatorname{csc} \frac{\pi}{m}, \quad (23e)$$

in which the family of strain rate distributions $\varphi(t, x; m)$ generated by the exponent $m \in \mathbb{R} \geq 2$ (Figure 4c) are of generalized Cauchy type and evolve in accordance with a time-variable characteristic width $l(t)$ (Figure 4d). We emphasize that, though illustrative, this class of solutions is by no means unique because (18a,b) contains no constraint on the spatial distribution of strain rate. Nevertheless, we note the following:

1. In the limit $m \rightarrow \infty$, the distribution $\varphi(t, x; m \rightarrow \infty)$ reduces to a uniform distribution on the constant interval $x \in [-l(0)/2, +l(0)/2]$. The parameter $c_1(m \rightarrow \infty) = 1$ (Figure 4a) and the solution (23a) reduces to the interfacial solution (21). However this limit also requires infinite mesh resolution or alternatively the definition of two mesh discontinuities, which defeats our intent.
2. In the other end-member case where $m = 2$, the distribution $\varphi(t, x; m = 2)$ is known as the Cauchy distribution, and its characteristic width $l(t) \propto V(t)^{-1}$ for a given initial condition $l(0) = l_0$. The parameter c_1 is chosen in such a way that the distribution $\varphi(t, x; m)$ integrates to one and the time of the singularity matches the prediction of the interfacial model. However, the trajectory of V towards the asymptote is unavoidably different (Figure 4b).

We therefore come to the qualitative conclusion that for initially smooth and numerically resolvable strain rate distributions, orders-of-magnitude increases of slip rate readily correspond to orders-of-magnitude increase of strain rate locality, placing insurmountable demands on mesh resolution. If left unconstrained, this process culminates in finite

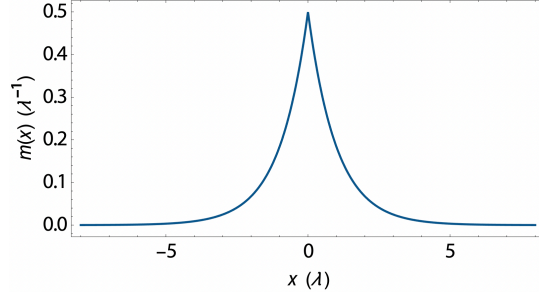


Figure 5. Plot of the mollifier $m(x) = \frac{1}{2\lambda} \exp\left(-\left|\frac{x}{\lambda}\right|\right)$.

496 time to an infinite slip rate concentrated in an infinitesimally thin region of space – a
497 plane.

498 3.3 Non-local Equivalent Strain Rate

499 In order to counter this tendency towards complete localization to a plane we in-
500 troduce the *non-local* interaction operator \mathcal{M} , which *mollifies* its operand by means of
501 convolution with the symmetric distribution $m(x)$:

$$502 \mathcal{M}(\gamma)(t, x) := \int_{\text{supp}(m)} m(\xi) \gamma(t, x - \xi) d\xi. \quad (24)$$

503 We note that $\mathcal{M}(\gamma)(t, x)$ still satisfies the relation (7) given that the distribution m in-
504 tegrates to unity over its domain.

505 Evaluation of $\mathcal{M}(\gamma)$ requires the evaluation of the integral of γ over the support
506 of $m(\xi)$ and therefore lacks the *mathematical* (as opposed to physical) locality that is
507 desirable for continuum models. We follow Peerlings et al. (1996) in constructing and
508 truncating a Taylor series linearization of (24). We find that a mollifier $m(x) = \frac{1}{2\lambda_0} \exp\left(-\left|\frac{x}{\lambda_0}\right|\right)$
509 (Figure 5) leads to series coefficients that are even powers of the length scale λ_0 :
510

$$511 \mathcal{M}(\gamma)(t, x) = \sum_{n=0}^{\infty} \lambda_0^{2n} \partial_x^{2n} \gamma$$

512 This infinite sum can be rolled up into a recursive or *implicit* definition of $\mathcal{M} = \mathcal{M}_{\text{im}}$

$$514 \mathcal{M}_{\text{im}}^{-1}(\bar{\gamma}) := \bar{\gamma} - \lambda_0^2 \partial_x^2 \bar{\gamma}, \quad (25)$$

515 which is to be solved for the auxiliary variable $\bar{\gamma} = \mathcal{M}_{\text{im}}(\gamma)$. Alternatively, the terms
516 beyond $n = 1$ can be dropped, leading to the *explicit* definition of $\mathcal{M} = \mathcal{M}_{\text{ex}}$
517

$$518 \mathcal{M}_{\text{ex}}(\gamma) := \gamma + \lambda_0^2 \partial_x^2 \gamma. \quad (26)$$

520 We note that any truncation of the Taylor series into a linear combination of even deriva-
 521 tives of γ adheres to the integral property expressed in (7), even if these truncations lose
 522 correspondence to a concrete mollifier. The implicit definition (25) has numerically ad-
 523 vantageous properties (Peerlings et al., 1996), but requires the potentially costly solu-
 524 tion of an elliptic PDE. For this reason we will use the explicit definition (26) in this work
 525 and discuss the implicit definition (25) as an option for future implementations.

526 We generalize the second partial derivatives ∂_x^2 with respect to the x coordinate
 527 that appear in (25) and (26) to Laplacians ∇^2 in 3D, again assuming that variations of
 528 strain rate across the fault are much bigger than variations along it.

529 A convolution over slip history $d(t)$ with a kernel $w(d) \propto \exp -d/d_c$, which is iden-
 530 tical to the kernel discussed above, was also employed by Ruina (1980, eq. 15, 16a, 17a)
 531 and Ruina (1983, eq. 13a,b) as a ‘generator’ of evolution laws. There, it represented the
 532 ‘fading memory’ of slip history on the fault. Here, we have consistently extended that
 533 notion with a non-local interaction of deformation that fades with distance away from
 534 a point in the shear band. Although not primarily intended to have physical meaning
 535 in this work, we note that the Laplacian of effective strain rate plays a role in granular
 536 physics (e.g. Bouzid et al., 2013, 2015). There, the length scale λ_0 is proportional to the
 537 grain size (Mühlhaus & Vardoulakis, 1987; Bouzid et al., 2013, 2015; Rice, 2006, and ref-
 538 erences in the latter).

539 3.4 Non-locality as a Localization Limiter

540 We repeat the analysis of Section 3.2, but now use the explicit non-local operator
 541 \mathcal{M}_{ex} in

$$542 \quad \dot{\gamma} \propto c_1 \gamma \mathcal{M}_{\text{ex}}(\gamma)$$

544 and

$$545 \quad \dot{\varphi} \propto V [\lambda_0 c_1 \varphi \mathcal{M}_{\text{ex}}(\varphi) - \varphi]. \quad (27)$$

547 As shown in Appendix C, we find that it is solved by

$$548 \quad V(t) \propto \left(1 - \frac{t}{t_*}\right)^{-1}, \quad (28a)$$

$$549 \quad \varphi(x) := \begin{cases} \frac{1}{\pi \lambda_0} \cos^2\left(\frac{1}{2} \frac{x}{\lambda_0}\right) & \forall x \in [-\pi \lambda_0, +\pi \lambda_0] \\ 0 & \forall x \notin [-\pi \lambda_0, +\pi \lambda_0] \end{cases}, \quad (28b)$$

$$550 \quad \gamma_0 \gamma(t, x) := v_0 V(t) \varphi(x), \quad (28c)$$

$$551 \quad c_1 := 2\pi, \quad (28d)$$

553 and observe that

- 554 1. The trajectory of V towards its asymptote can be made to exactly match the pre-
555 diction (21) of the interfacial rate and state friction formulation (17a,b) by choos-
556 ing $c_1 = 2\pi$.
- 557 2. The cosine-squared distribution with prescribed size λ_0 is an attractive steady state
558 during the process of earthquake nucleation. Further localization will not spon-
559 taneously occur. We interpret this distribution as striking a balance between the
560 localizing tendency of “ $\dot{\varphi}^2 - \varphi$ ” and the delocalizing tendency of “ $\varphi \nabla^2 \varphi - \varphi$ ”
561 (Equation (19) and Figure 3).

562 The implicit non-local operator \mathcal{M}_{im} in the same context does not have analyti-
563 cal solutions that simultaneously satisfy $\dot{V} \propto V^2$ and $\dot{\varphi} = 0$ with non-degenerate dis-
564 tribution $\varphi = \varphi(x/\lambda_0)$.

565 3.5 Steady-state Friction and Deformation

566 As mentioned in Section 1.1, an important property of the state evolution law is
567 the steady state $\Theta V = 1$. We wish to retain this steady state in the continuum equiv-
568 alent (16) of the state evolution law, ideally at a steady strain rate distribution $\varphi(x)$, cf.
569 (19). We are thus interested in a steady-state solution to

$$570 \quad \dot{\varphi} \propto \lambda_0 c_1 \varphi \mathcal{M}_{\text{ex}}(\varphi) - \lambda_0^{\frac{a}{b}} c_2 \varphi^{1+\frac{a}{b}}, \quad (29)$$

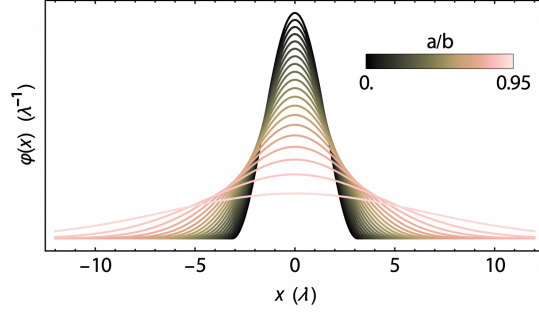


Figure 6. Steady-state distributions φ [cf. Eq. (30a)] of strain rate at various a/b versus distance x with respect to the fault core. The curve plotted for $a/b = 0$ also corresponds to the most extreme strain rate localization at any value of a/b (see main text).

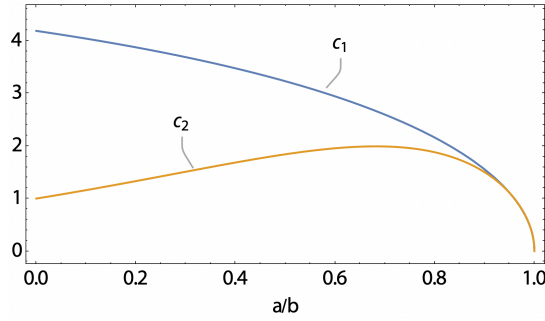


Figure 7. The coefficients $c_1(\frac{a}{b})$ and $c_2(\frac{a}{b})$, given by Eq. (31a) and (31b), are plotted versus the acceptable range of $\frac{a}{b}$.

572 which exists as long as $a, b > 0$ (standard requirements) and $a < b$ (a net weakening
573 fault). As detailed in Appendix B, this solution is given by

$$574 \quad \varphi_{\text{ss}}(x) = \begin{cases} \varphi_0 \left[\cos\left(\frac{1}{2} \frac{x}{\lambda_1}\right) \right]^{2[1-\frac{a}{b}]^{-1}} & \forall x \in [-\pi\lambda_1, +\pi\lambda_1] \\ 0 & \forall x \notin [-\pi\lambda_1, +\pi\lambda_1] \end{cases} \quad (30a)$$

$$575 \quad \lambda_1 := \lambda_0 \left[1 - \frac{a}{b} \right]^{-1}. \quad (30b)$$

577 Faults that are net strengthening or neutral ($a > b$) only possess a uniform steady state
578 strain rate distribution that is bounded by the (fictitious) walls of the gouge layer (Sec-
579 tions 2.2 and 5.5).

580 The coefficients c_1 and c_2 , and the newly introduced φ_0 , are given by

$$581 \quad c_1 = \frac{1}{2}\pi^{-\frac{1}{2}} \frac{\Gamma\left(1 + 2\left[1 - \frac{a}{b}\right]^{-1}\right)}{\Gamma\left(\frac{1}{2} + 2\left[1 - \frac{a}{b}\right]^{-1}\right)} [\lambda_1 \varphi_0]^{-2} \quad (31a)$$

$$582 \quad c_2 = \frac{1}{2}\pi^{-\frac{1}{2}} \frac{\Gamma\left(2\left[1 - \frac{a}{b}\right]^{-1}\right)}{\Gamma\left(\frac{1}{2}\left[3 + \frac{a}{b}\right]\left[1 - \frac{a}{b}\right]^{-1}\right)} [\lambda_1 \varphi_0]^{-1 - \frac{a}{b}} \quad (31b)$$

$$583 \quad \varphi_0 = \frac{1}{2}\pi^{-\frac{1}{2}} \frac{\Gamma\left(1 + \left[1 - \frac{a}{b}\right]^{-1}\right)}{\Gamma\left(\frac{1}{2} + \left[1 - \frac{a}{b}\right]^{-1}\right)} \lambda_1^{-1}, \quad (31c)$$

584
585 with Γ the gamma function that has the property $\Gamma(n) = (n-1)!\forall n \in \mathbb{N}^+$. The coef-
586 ficients c_1 and c_2 are plotted as functions of a/b in Figure 7. The solutions (30a) are plot-
587 ted in Figure 6, in which we can see that steady-state anelastic strain rate distributions
588 have finite width and are therefore numerically resolvable as long as the aforementioned
589 requirements on a , b , and a/b are met.

590 3.6 Damping

591 Finite time blow-up of the solution (Section 3.5) under quasi-static loading con-
592 ditions is understood to be the consequence of the model's inability to radiate or dissi-
593 pate liberated potential strain energy away from the fault. It appears impossible to di-
594 rectly counteract this problem by implementing inertia with absorbing boundary con-
595 ditions in the 1D models we present here, due to the resulting model's tendency to quickly
596 establish a global steady state at an unphysically high strain rate. Therefore, inspired
597 by the *radiation damping approximation* of Rice (1993) (see also Ben-Zion & Rice, 1995;
598 Rice & Ben-Zion, 1996; Lapusta et al., 2000; Thomas et al., 2014) and the Kelvin-Voigt
599 visco-plastic regularization of Needleman (1988) (see also F. Wu & Freund, 1984; Peirce
600 et al., 1983; Duretz et al., 2019, 2020, 2021; de Borst & Duretz, 2020; Stathas & Stefanou,
601 2022), we add a linear viscous damping term $\eta\gamma$ in a parallel configuration to the con-
602 tinuum friction law:

$$603 \quad f = f_0 + a \log \gamma + b \log \theta + \eta\gamma. \quad (32)$$

605 Informally, the action of this term may be understood to truncate the anelastic shear
606 strain rate, which then causes the distribution of said strain to widen under continued
607 loading. We formalize this claim by repeating the derivation in Section 3.1 of the anelas-
608 tic shear strain rate ODE (18a,b), now incorporating the damping term in (32). We ar-
609 rive at the following result, written in terms of the newly introduced auxiliary field ζ :

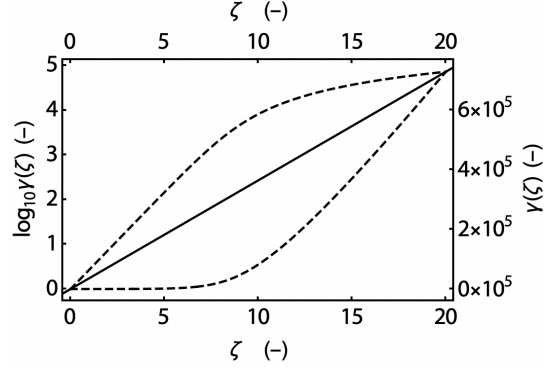


Figure 8. The function $\gamma(\zeta)$ defined in Eq. (33c) plotted on a logarithmic scale in the top left diagonal portion of the figure and on a linear scale in the opposing portion, highlighting the exponential behavior of $\gamma(\zeta)$ at low ζ and the linear behavior at high ζ .

610

611

$$a \dot{\zeta}(t, x) = b r_0 (c_1 \mathcal{M}_{\text{ex}} \circ \gamma \circ \zeta(t, x) - c_2 \theta(\zeta, f)^{-1}) + \dot{f}(t), \quad (33a)$$

612

$$\theta(\zeta, f)^{-1} := (\gamma \circ \zeta(t, x))^{\frac{a}{b}} \exp\left(-\frac{1}{b} [f - f_0 - \eta \gamma \circ \zeta(t, x)]\right), \quad (33b)$$

613

614

$$\gamma(\zeta) := \frac{a}{\eta} W_0\left(\frac{\eta}{a} \exp \zeta\right), \quad (33c)$$

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with the principal branch $W_0(\diamond)$ of the Lambert W function, which is an implicit transcendental function that is defined to be the solution to the equation $W_0 \exp W_0 = \diamond$. While we acknowledge that the use of implicit functions in physical descriptions is not ideal, from a practical standpoint there is little problem because reliable and fast algorithms for computing the common Lambert W function are included in many programming languages (Fritsch et al., 1973; Barry, Barry, & Culligan-Hensley, 1995; Barry, Culligan-Hensley, & Barry, 1995; Johansson, 2020).

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As indicated in Figure 8, the function $\gamma(\zeta)$ is approximately linear at high ζ (and correspondingly high γ). This means that in this limit, (33) simplifies to

624

$$\eta \dot{\gamma}(t, x) = b r_0 (c_1 \mathcal{M}_{\text{ex}} \circ \gamma(t, x) - c_2 \theta(\gamma, f)^{-1}) + \dot{f}(t), \quad (34a)$$

625

626

$$\theta(\gamma, f)^{-1} := \exp\left(-\frac{1}{b} [f - f_0 - \eta \gamma]\right), \quad (34b)$$

627

628

We can now compare reaction terms of (18) and (34). Ignoring spatial derivatives and the temporal derivative \dot{f} of friction in both, the reaction term of (18) that is active at

629 low strain rate looks like $\dot{\gamma} \propto \gamma^2 - \alpha(f)\gamma^{1+\frac{a}{b}}$ while the reaction term of (34) that is
 630 active at high strain rate looks like $\dot{\gamma} \propto \gamma - \alpha(f) \exp(\frac{\eta}{b}\gamma)$. The former is a convex func-
 631 tion of γ , the latter is a concave function of γ . We can thus see that the same conditions
 632 (large f , small $\alpha(f)$) that readily trigger runaway sliding in (18) also introduce a sta-
 633 ble steady state at high γ in (34) and therefore also in (33). We note that this quality
 634 of having a stable steady state at non-zero γ means that (33) can be classified as a reaction-
 635 diffusion equation of generalized Fisher or KolmogorovPetrovskyPiskunov (KPP) type
 636 (Fisher, 1937; Kolmogorov et al., 1937). This type of equation is known to exhibit out-
 637 ward propagating (strain rate) fronts, and consequently does not admit a steady strain
 638 rate distribution. This markedly limits the validity of the analytical values derived for
 639 $c_{1,2}$ during the coseismic phase. We will revisit this issue in detail throughout the results
 640 section.

641 In the interfacial representation of a fault in a homogeneous half-space, the shear
 642 traction $\tau(t, z)$ at a time t and location z along the fault may be decomposed into

$$643 \quad f(t, z)\sigma = \tau(t, z) = \tau_0(t, z) + \tau_d(t, z) - \eta_*V(t, z),$$

644
 645 where $\tau_0(t, z)$ is the loading stress, $\tau_d(t, z)$ is the convolution that expresses the dynamic
 646 stress transferred to the point z at time t by slip at all points within the causality cone
 647 around (t, z) (Rice, 1993; Cochard & Madariaga, 1994, 1996; Lapusta et al., 2000). The
 648 term η_*V exists to compensate for the removal of a singularity in aforementioned con-
 649 volution, and this sets the viscosity $\eta_* = \frac{1}{2} \frac{\mu}{c}$ equal to half the material's shear impedance,
 650 with $c_s = \sqrt{\mu/\rho}$ the shear wave speed and in turn μ the shear modulus and ρ the mass
 651 density. In the *quasi-dynamic* or *radiation damping* approximation (Rice, 1993), the dy-
 652 namic stress transfer term $\tau_d(t, z)$ is replaced by an easier to compute static stress trans-
 653 fer term $\tau_s(t, z)$, but the damping term η_*V is retained. Here we take the traction $\tau_0 +$
 654 τ_s to be the one produced by our static momentum balance equation (9) and elastic con-
 655 stitutive equation (8), and implicitly move over the damping term to the right-hand side
 656 of the equation. Calibration of (32) to the result gives our non-dimensional damping vis-
 657 cosity as

$$658 \quad \eta = c_3 \frac{1}{2} \frac{\mu v_0}{\sigma c_s},$$

659
 660 with $c_3 \sim 1$ an additional dimensionless calibration constant that is chosen somewhat
 661 arbitrarily to be equal to c_1 , which seems to give maximum macroscopic slip rates that
 662 are similar to those produced by the interfacial friction law with radiation damping.

4 Numerical Models

Analytical predictions made in preceding sections are complemented here with numerical simulation results that better illustrate the complex time-dependent behavior of the system of equations, and allow us to establish the consequences of approximations and other model choices made in the process of formulating a local equivalent continuum rate and state friction law. As before we restrict the scope to 1D models across the shear band.

4.1 Equations

On the one-dimensional line $x \in [-L/2, +L/2]$ that deforms under simple shear, analogous to the transect T defined in Figure 1, we model the compact equation (33a,b) together with the one-dimensional static momentum balance equation $\partial\tau/\partial x = 0$ cf. (9), elastic constitutive equation $\dot{\tau}(t, x) = \mu [\partial v(t, x)/\partial x - 2\gamma_0\gamma(t, x)]$ cf. (8) (with μ the shear modulus as before), yield equality $\tau(t) = f(t)\sigma$, and finally the boundary conditions $v(t, x = \mp L/2) = \mp V_p/2$ combined into the single ODE

$$\dot{f}(t) = \frac{\mu}{\sigma} \left[\frac{V_p}{L} - \frac{\mathcal{V}(\gamma \circ \zeta)}{L} \right], \quad (33d)$$

with V_p the ‘plate’ driving velocity that is applied antisymmetrically on both ends of the domain. This equation uses (7) that relates slip velocity to the integral (from $-L/2$ to $+L/2$) of anelastic shear strain rate γ , as well as (33c) for the relation between γ and the auxiliary field ζ .

The system is closed with natural boundary conditions $\partial\gamma/\partial x = 0$ on $\gamma(t, x = \mp L/2)$, and initial conditions $f(t = 0) = f_0 + \Delta f$ and $\gamma(t = 0, x) \propto \varphi_{ss}(x) + \epsilon$, with $\varphi_{ss}(x)$ deriving from the analytical prediction (30a) and ϵ a necessary but small homogeneous background value of anelastic shear strain rate. As long as this value is small enough ($\epsilon \ll V_p/L \ll v_0/L$) and the ‘tails’ of the strain rate distribution can be ignored, the required computational domain size may be based on the length scale λ_0 and therefore cover only a potentially small fraction of the line $[-L/2, +L/2]$, greatly reducing the computational cost.

The solution procedure is outlined in Appendix D.

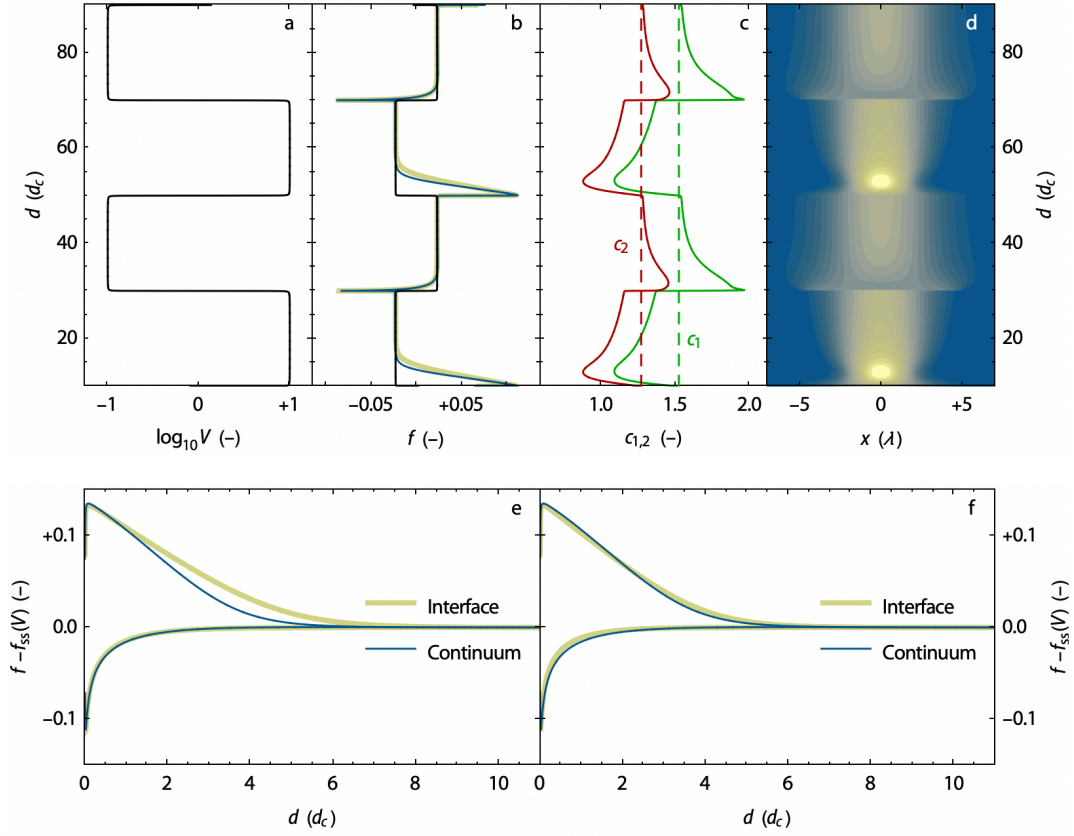


Figure 9. Results of the numerical velocity stepping experiments: a) dimensionless sliding velocity V closely tracking a smooth square wave signal (Eq. (36)); b) frictional response of the interfacial reference model (thick yellow line) and continuum model (dark blue line) with for reference the theoretical steady-state friction $f_{ss}(V)$ (black line); c) measured (solid lines; cf. Eq. (20a,b)) vs. used (dashed lines; cf. Eq. (31a,b)) values of the dimensionless coefficients $c_{1,2}$; d) distribution $\varphi = \gamma/V$ of anelastic shear strain rate with darker blue colors reflecting very low values and bright yellow colors reflecting high values, and contours distributed evenly on a linear scale. Reflecting the ultimate slip dependence (rather than direct time dependence) of the rate and state friction laws, curves in panels a)–d) are plotted against sliding distance d (measured in critical slip distances d_c) on the vertical axis. Panel e) shows an enhanced view of the evolution of the instantaneous friction coefficient towards the steady-state value with slip distance d , simultaneously showing a positive and negative step for both the interface and continuum models. Panel f) displays the same, but artificially increases the critical slip weakening distance of the interfacial friction law by 20%.

Table 1. Parameter values used in the velocity stepping experiment

parameter	value	unit
a	$2 \cdot 10^{-2}$	1
b	$3 \cdot 10^{-2}$	1
d_c	10^{-5}	m
v_0	10^{-6}	m s^{-1}
ρ	10^{+3}	kg m^{-3}
μ	10^{+10}	Pa
c_s	$3.162... \cdot 10^{+3}$	m s^{-1}
η	$1.581... \cdot 10^{-6}$	1
σ	10^{+6}	Pa
d_c	$5.00 \cdot 10^{-4}$	m
λ_0	$\sim 1.4 \cdot 10^{-4}$	m
c_1	$1.531...^a$	1
c_2	$1.276...^a$	1
r_0	v_0/d_c	s^{-1}
ν	10^{-3}	1
n	20	1

Some rate and state parameter values are roughly based on Erickson et al. (2020).

^aUsing (31a) and (31b) with stated values of a and b .

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4.2 Continuum Velocity Stepping Friction Experiments

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Velocity stepping laboratory experiments form the foundation on which rate and state friction was proposed (Dieterich, 1978). We compare numerical velocity stepping experiments of a traditional rate and state friction-governed interface with our proposed continuum equivalent under the same conditions. We impose a driving velocity that follows a smooth square wave function with respect to sliding distance d :

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$$\log_{10}(V_p(d)/v_0) = \frac{2}{\pi} \tan^{-1} \left(-\frac{1}{\nu} \cos \left(\frac{\pi d}{n d_c} \right) \right), \quad (36)$$

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with $n = 20$ the amount of slip weakening distances d_c between velocity steps, and $\nu = 10^{-3}$ a parameter that controls the smoothness of the smooth square wave, with small numbers giving the squarest result (see Figure 9a). Some smoothness is important to maintain numerically stable results. The low and high driving velocities are $10^{-1}v_0$ and $10^{+1}v_0$ respectively. Other model parameters are given in Table 4.2. Inertial effects may be ignored (both in the original lab experiments and in our simulations) due to the low driving velocity. Thus, the quasi-radiation damping viscosity η is set to zero and its effects not discussed here. In the limit $\eta \rightarrow 0^+$, (33c) simply reduces to $\gamma(\zeta) = \exp \zeta$ or equivalently $\zeta(\gamma) = \ln \gamma$, which reduces (33a,b) back to (18a,b).

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Because the domain size L is very small and the system (33a–33d) therefore stiff, the friction tends to a stable steady state after being perturbed by a velocity step rather than developing a limit cycle. This also means that $\mathcal{V}(\gamma) \approx V_p(d)$. In fact, $\mathcal{V}(\gamma)$ and $V_p(d)$ are so close that their difference would not register in Figure 9a.

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In the following, we use the terms time dependence and slip dependence interchangeably but note that the rate of change of either is not constant from the perspective of the other. A slip-centric presentation of velocity-stepping results is commonplace in the earlier cited literature.

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In Figure 9b (and its detailed view 9e) we explore the similarities and differences in slip dependence of the friction coefficient f between the interfacial and continuum velocity stepping experiments. Based on (19), we expect that a velocity step perturbs not just the magnitude of anelastic shear strain, but also its distribution (Figure 9d) and therefore the theoretical value of the dimensionless coefficients $c_{1,2}$ (Section 3.3; Equations (20a) and (20b); Figure 9c, solid lines) away from steady state. Since these coefficients are actually set to their constant steady-state predictions (Section 3.5; Equations (31a)

724 and (31b); dashed lines in Figure 9c), we expect some differences in time or slip depen-
725 dence to occur. Notably, a positive velocity step leads to a temporary localization of anelas-
726 tic deformation and a corresponding decrease in the theoretical values of $c_{1,2}$ with re-
727 spect to steady state. This leads to an overestimation of corresponding terms in the right-
728 hand-side of (33a) and a faster decay towards the steady-state friction coefficient (Fig-
729 ure 9b,e) of the continuum model with respect to the interfacial model. A negative ve-
730 locity step causes temporary delocalization (Figure 9d) and under-estimation of terms
731 involving $c_{1,2}$ in (33a), but does not lead to any noticeable difference in friction coeffi-
732 cient evolution towards the steady state between the two models (Figure 9b,e). We con-
733 clude that those terms involving $c_{1,2}$ are negligible in this situation.

734 In Figure 9c and 9d it can also be seen that the *distribution* of anelastic shear strain
735 rate tends more slowly to a steady state than the friction coefficient itself, and moreover
736 that this decay is slower after a positive velocity step than after a negative one. In fact,
737 the 40 d_c wavelength we use for the input signal (36) is too small to enable a somewhat
738 complete evolution towards steady-state of the anelastic strain rate distribution during
739 the high-velocity regime, but we have taken care that this does not meaningfully impact
740 the results during the subsequent low-velocity regime.

741 Most of the difference between the interfacial and continuum velocity stepping re-
742 sults disappears if the critical slip weakening distance d_c is seen as an additional cali-
743 bration parameter. Continuing the analogy to the classical experimental setup, when we
744 regard the continuum formulation proposed here to be the ‘generator’ of experimental
745 data, and the result of the existing interfacial rate and state friction and aging laws as
746 a data fitting curve, we observe a critical slip weakening distance d_c that is about 20%
747 larger than the value that was used to generate the data. The result is plotted in Fig-
748 ure 9f. There it can be seen that the discrepancies remaining after the calibration of d_c
749 are small compared to the expected noise level of experimental results.

750 4.3 Continuum Spring-slider Experiments

751 Whereas velocity stepping experiments are useful for studying the way in which
752 friction can be attracted to a steady state, permanently out-of-equilibrium ‘limit cycle’
753 behavior is more interesting in the study of earthquake dynamics. We achieve this be-
754 havior by increasing the size of the domain and the distance of the boundary conditions

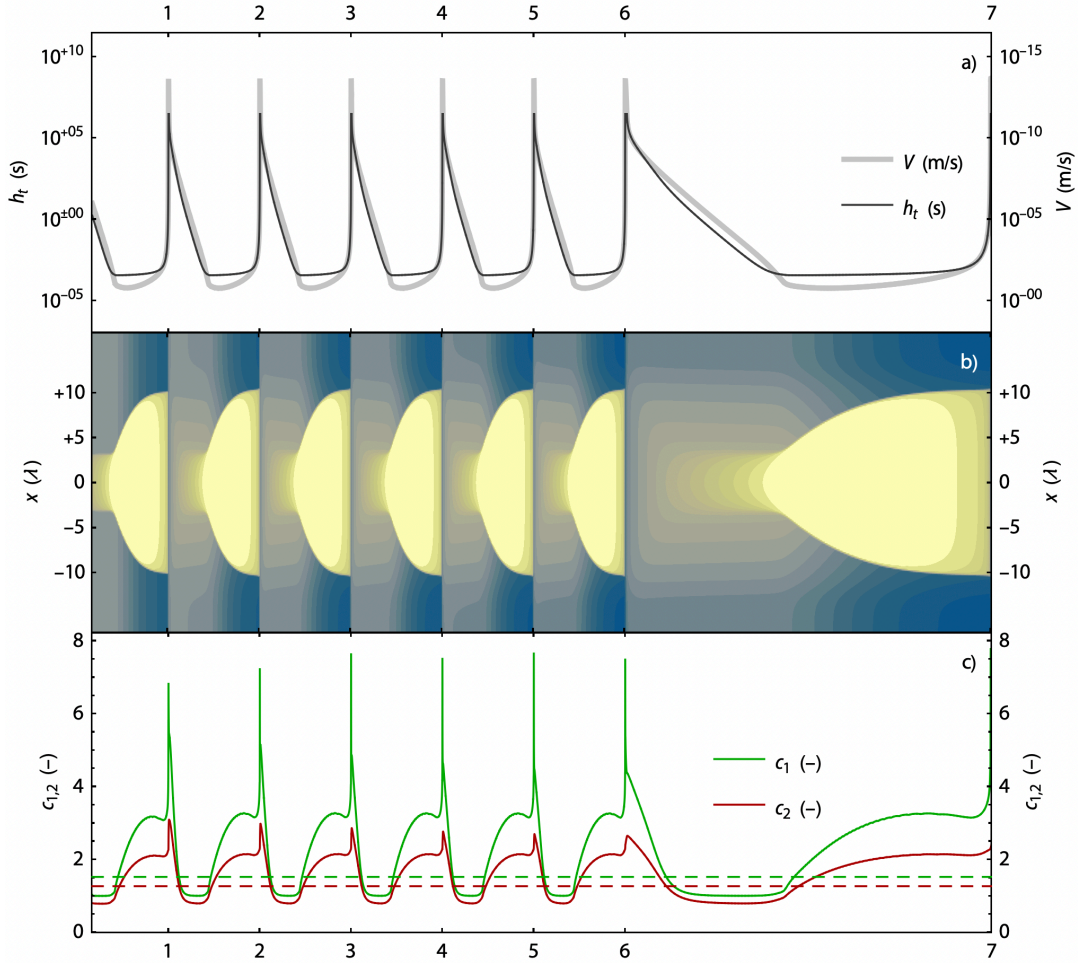


Figure 10. Evolution of the continuum seismic cycle simulation demonstrated by a succession of six events plotted against time step count on the horizontal axis, with a seventh event stretched by a factor four in order to better show some features of an individual event. The end of each event is punctuated by a slip velocity minimum and numbered 1 to 7 on the horizontal axes. Each revolution around the limit cycle takes 83.00 years and 107,500 time steps to complete. Panel a: time step size h_t (thin dark line; left axis; logarithmic scale) and slip velocity V (thick bright line, right axis; logarithmic scale). The former serves to aid the interpretation of the time-dependent progression of each event while the latter gives a sense of magnitude to the color scale of panel b. Axes are chosen so that together these curves express the imperfect reciprocal relation between time step and slip velocity. Panel b: logarithm of dimensionless anelastic shear strain rate versus signed distance x measured in number of length scales λ_0 away from the shear zone center. Bright yellow colors indicate high strain rate; dark blue colors indicate low strain rate. Contours are drawn at equal intervals in log space. No explicit color scale is given because the magnitude of anelastic shear strain rate γ is tied to the value of λ_0 in order to produce a length-scale independent history of sliding velocity V (panel a). Panel c: measured (solid lines; cf. Eq. (20a,b)) vs. used (dashed lines; cf. Eq. (31a,b)) values of the dimensionless parameters $c_{1,2}$.

$c_{1,2}$.

755 by many orders of magnitude ($L = 80$ km) with respect to the velocity stepping ex-
 756 periments. Thereby, we greatly reduce the effective stiffness of the medium as can be seen
 757 in (33d). This is analogous to performing the so-called spring-slider experiment in which
 758 a mass is dragged over a surface by a spring that is tensed at a constant rate to give an
 759 educational example of the apparent stick-slip behavior of seismogenic faults. Our goal
 760 in performing this experiment is to determine the degree to which the here proposed con-
 761 tinuum friction laws match the predictions made by Dieterich and Ruina's interfacial laws
 762 and to describe the transient behavior of anelastic strain rate in the added dimension.
 763 Damping plays an essential role in these experiments to close the limit cycle that oth-
 764 erwise extends to $V \rightarrow \infty$ (Section 3.6). As discussed in the same section, our imple-
 765 mentation of damping as a linear bulk viscosity precludes a steady-state anelastic strain
 766 rate distribution at high slip velocity, a fact that directly contradicts one of our precepts
 767 (Section 3.1). Therefore we anticipate a larger discrepancy between interface and con-
 768 tinuum models here than was observed during the velocity stepping experiments.

769 Figure 10 depicts the evolution of quantities of interest over the course of multi-
 770 ple orbits of the solution as it converges to the limit cycle. The first important obser-
 771 vation is that the continuum model still appears to exhibit spontaneous limit cycle be-
 772 havior. Due to the appropriately chosen initial condition, the wind-up period is short
 773 and there is little variability between successive revolutions. As predicted (Section 3.5),
 774 the nucleation phase is marked by runaway amplitude increase of a quasi-steady strain
 775 rate distribution (panel b). As before, we measure the dynamic value of the coefficients
 776 $c_{1,2}$ from the strain rate distribution following (20a,b), and compare to the analytically
 777 derived steady-state values (31a,b). In line with analytical predictions this phase is also
 778 accompanied by measured values of $c_{1,2}$ (panel c) that are slightly lower than the con-
 779 stant values that are prescribed (dashed lines in panel c). The opposite situation occurs
 780 throughout the remainder of each cycle where anelastic strain is more widely distributed.
 781 We can attribute this directly to the effect of damping at high strain rate, which exerts
 782 a delocalizing influence. Figure 11 provides a more intuitive view of the short-lived de-
 783 localizing behavior of a strain pulse using linear rather than logarithmic scales. We em-
 784 phasize that the ultimate extent of coseismic delocalization compared to the length scale
 785 λ_0 is dependent on model and material parameters, among which the effective stiffness
 786 of the medium.

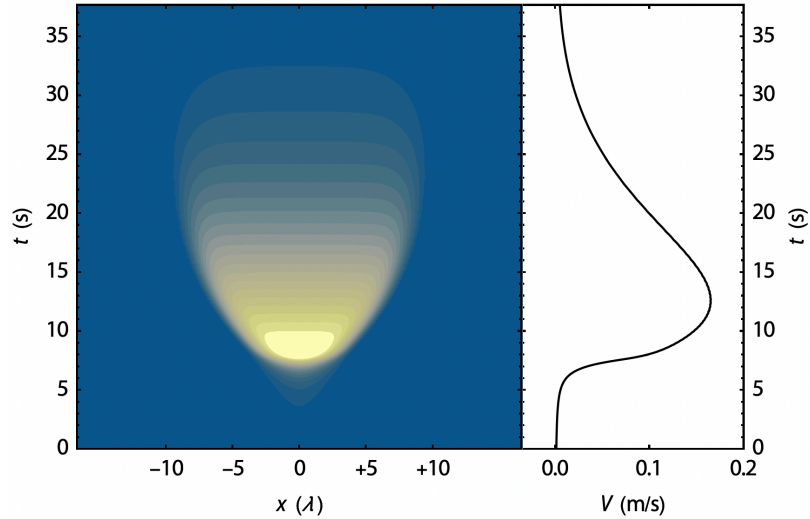


Figure 11. Focus on the last event of Figure 10, here visualized in panel a by strain rate on a linear color and contour scale and in panel b by slip rate V , both versus a linear time scale in seconds.

787 Large peaks in the measured values of $c_{1,2}$ shown in panel c of Figure 10 occur dur-
 788 ing the post- and interseismic phase, where anelastic shear strain rate is broadly and quite
 789 uniformly distributed. The values that are actually used (dashed lines) are way too small
 790 here, however, terms that do not involve these coefficients dominate the time-dependent
 791 behavior in this regime and the discrepancy's consequences are limited. This is largely
 792 confirmed in Figure 12, which tabulates the evolving scalar properties of a limit cycle
 793 of the interfacial and continuum models of rate and state friction. In particular, the du-
 794 ration of the limit cycle (first column, panels a–c, time in years on the horizontal axis)
 795 is very similar even though the timing of the event in relation to its enclosing slip ve-
 796 locity minima is somewhat different.

797 Markedly different are the coseismic aspects of the limit cycle. Panel e of the mid-
 798 dle column and panels h–j of the right column of Figure 12 clearly show this. The con-
 799 tinuum model has a faster stress drop (panel h, $U' \propto f$) and a more symmetric slip rate
 800 response with time compared to the interfacial model. The limit cycle depicted in lin-
 801 ear phase space (U', V) in Figure 12e shows a skewed triangular trajectory of the inter-
 802 facial model and a more parabolic trajectory for the continuum model. It appears that
 803 the amount of stress drop that happens *before* peak slip rate in the interfacial model is
 804 small – about 10% of the total stress drop. At the same point in the continuum model

805 the stress drop is already about one third of the total. The same limit cycle shown in
 806 Figure 12f in log-linear phase space ($U', \log_{10}V$) allows us to see that there are subtle
 807 differences between the limit cycles of the interfacial and continuum models over the whole
 808 range of slip velocities.

809 A further result of practical interest is the difference in adaptively chosen time step
 810 between the interfacial and continuum models. Figure 10a shows an approximate inverse
 811 relation between the time step of continuum model to the aggregate slip velocity v_0V ,
 812 which is unsurprising seeing how $d_c/[v_0V(t)]$ gives a sensible local time scale for both
 813 interfacial and continuum rate and state friction models. The approximate inverse re-
 814 lation breaks down at large time scales where the corresponding time steps are truncated
 815 to a fixed value of 0.1 yr, and at large slip velocity and small time step, where the time
 816 scale of the continuum model becomes dominated by large spatial gradients and is gen-
 817 erally much smaller than that of the interfacial model (Figures 12d,g,k). From these same
 818 plots it can be seen that, just like stress begins to drop in significant advance of an event,
 819 so too does the time step of the continuum model.

820 Finally, we remark that spatial resolution tests indicate that the quality of the so-
 821 lution is only influenced by the ratio of cell size h_x to λ_0 , with acceptable results achieved
 822 when $h_x/\lambda_0 \geq 10$. For the results presented here we have used $h_x/\lambda_0 = 20$. Chang-
 823 ing λ_0 in proportion to h_x does nothing except to cause a wider or narrower but equally
 824 well resolved strain distribution and a virtually identical limit cycle.

825 5 Discussion

826 So far we have argued for a visco-elastic continuum rheology that resembles both
 827 the mathematical presentation as well as the resulting behavior of traditional interfa-
 828 cial rate and state friction as described by Dieterich and Ruina in their respective sem-
 829 inal publications. Before addressing this work in the context of a much broader body of
 830 existing research and reflecting on critical assumptions made in this work and its exten-
 831 sion towards the future, we briefly touch upon some alternate branches of our proposed
 832 theory that have been left out so far for the sake of clarity.

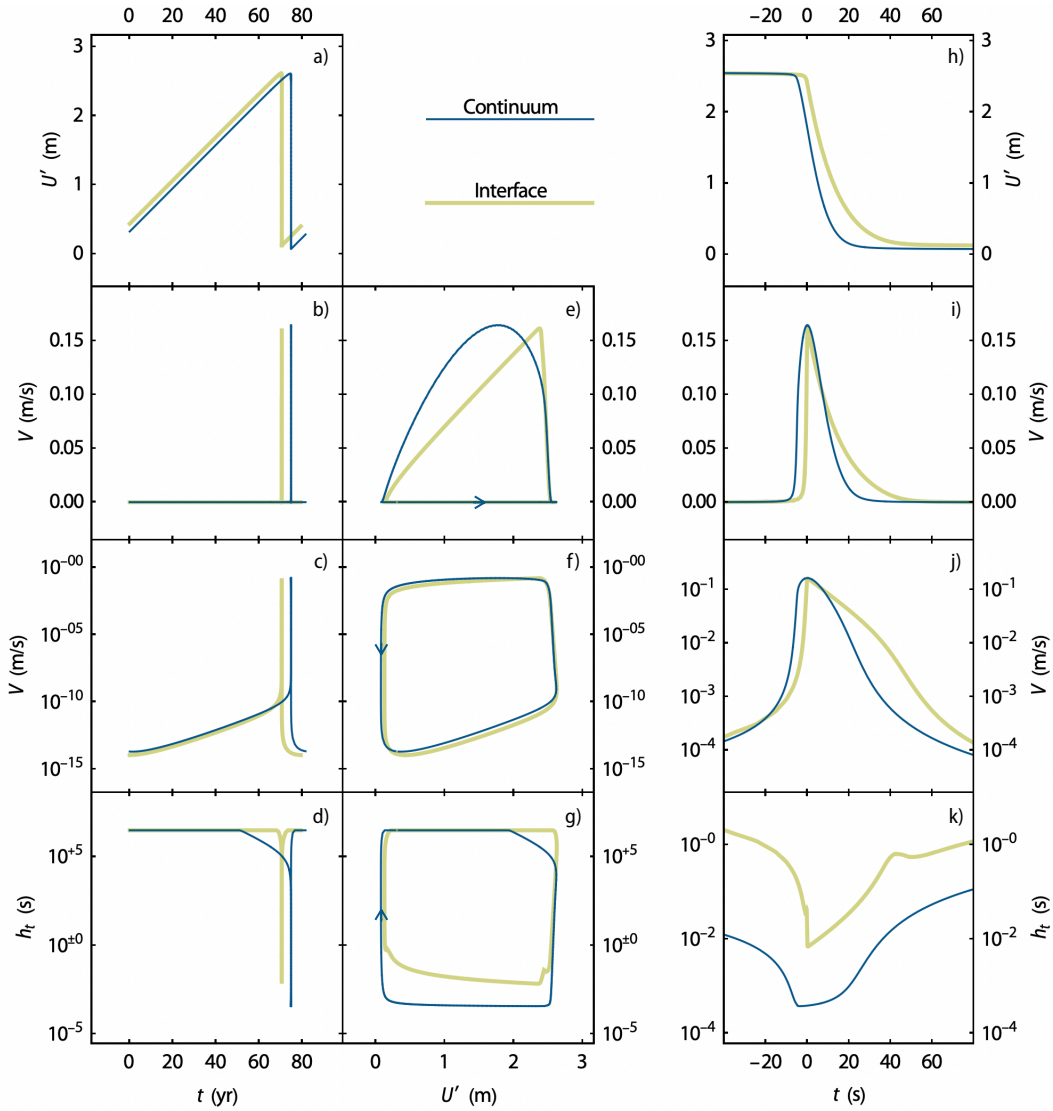


Figure 12. Panels a–g: behavior of the interfacial (bright yellow lines) and continuum (dark blue lines) models for a full limit cycle. Panels h–k: focus on the 100 seconds around an event. From left to right, horizontal axes are: time t in years, slip deficit U' in meters (linearly related to shear stress τ and friction coefficient f), and time t in seconds. Vertical axes from top to bottom are slip deficit U' in meters, slip velocity V in meters per second on a linear scale, the same on a logarithmic scale, and finally the time step size h_t in seconds on a logarithmic scale. Where limit cycles are shown (middle column, panels e–g), the cycle sense is indicated with small arrows.

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5.1 A Primitive Reformulation

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In contrast to traditional models of plasticity, the continuum rate and state friction laws presented here possess a continuous time dependence that makes the anelastic shear strain rate γ a predictable quantity rather than one that is to be solved by a constrained optimization algorithm (e.g. Simo & Taylor, 1985; Duretz et al., 2018). This is apparent from the ODE (18a,b) for γ , but is hidden in the Dieterich-Ruina form (15) and (16) of the continuum rate and state friction rheology. However, (18a,b) lacks the simplicity and elegance of the Dieterich-Ruina system and is difficult to interpret as the resultant of physical phenomena. By defining a set of primitive variables κ and ψ as

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$$\dot{\kappa} = r_{\kappa}\gamma \quad (37)$$

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$$\dot{\psi} = r_{\psi}\gamma^{a/b} \exp(-[f(\boldsymbol{\sigma}) - f_0]/b),$$

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with the respective reference rates $r_{\kappa} = r_0 c_1$ and $r_{\psi} = r_0 c_2$, we are able to write a more expressive formulation of (18a,b) as

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$$\dot{\kappa} = r_{\kappa}\gamma = r_{\kappa} \left(\dot{\psi}/r_{\psi} \right)^{\frac{b}{a}} \exp([f(\boldsymbol{\sigma}) - f_0]/a) \quad (38)$$

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$$\dot{\psi} = r_{\psi} \exp(\mathcal{M}_{\text{ex}}(\kappa) - \psi), \quad (39)$$

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that still closes the momentum balance equation (9) and elasto-plastic constitutive equation (8) without algebraic constraints and thus without requiring a constrained optimization algorithm. Equation (39) can optionally accommodate damping using the Lambert W function analogous to the way it was discussed in Section 3.6.

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We recognize $\kappa(t, x)$ as a dimensionless measure of accumulated anelastic shear strain and therefore assign to ψ the meaning of a ‘virtual’ or ‘unrealized’ strain. In spite of the conceptual appeal of (38) and (39), it can be difficult to choose appropriate initial conditions for ψ .

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5.2 Quadratic Diffusion

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Instead of the mollified term $\gamma \mathcal{M}(\gamma)$ that was introduced in (16) and used throughout the remainder of this work, we could have chosen to use $\mathcal{M}(\gamma^2)$:

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$$a \dot{\gamma} = b r_0 [c_1 \mathcal{M}(\gamma^2) - c_2 \gamma \theta(\gamma, f)^{-1}] + \gamma \dot{f}. \quad (40)$$

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and the definition of c_1 is adapted to the new way of mollifying according to the same principles as laid out in Section 3.

865 This choice has advantages and disadvantages. A major disadvantage is that it does
 866 not seem to allow a return from (40) to the more elegant Dieterich-Ruina form (15) and
 867 (16), nor to the compact form (38)–(39) proposed in the preceding section. A clear ad-
 868 vantage is that, taken together with the explicit non-local operator \mathcal{M}_{ex} , (40) produces
 869 a *degenerate* reaction-diffusion equation of generalized Fisher or KolmogorovPetrovskyPiskunov
 870 (KPP) type (Fisher, 1937; Kolmogorov et al., 1937) that remains well-posed at and around
 871 $\gamma = 0$. This leads to shear zone solutions that can propagate into perfectly intact rock
 872 without taking recourse to an arbitrary small initial condition on γ . Furthermore, un-
 873 der quadratic diffusion, the function $\gamma(\zeta)$ used in the ODE form (33a–d) in Section 3.6
 874 becomes algebraic instead of transcendently implicit.

875 We have reproduced all results presented in Section 4 with this quadratic diffusion
 876 term and zero background anelastic strain rate too, and observe numerically smoother
 877 and better resolved shear fronts during the coseismic delocalization phase. These fronts
 878 also travel faster and further than in the linear diffusion case leading to a larger discrep-
 879 ancy between measured and used values of the coefficients $c_{1,2}$ and consequently larger
 880 deviations of the limit cycle with respect to that of the interfacial rate and state friction
 881 model. The generalized Fisher–KPP type equation has merited a large body of theoret-
 882 ical research (e.g. Sánchez-Garduño & Maini, 1994, 1995, 1997; Sherratt & Marchant,
 883 1996; Petrovskii & Li, 2003, 2006; Gilding & Kersner, 2005; Y. Wu et al., 2006; Y. Li &
 884 Wu, 2008; Malaguti & Ruggerini, 2010; Broadbridge & Bradshaw-Hajek, 2016) and there-
 885 fore some of the above mentioned empirical findings may be given a theoretical under-
 886 pinning in future work.

887 5.3 Relation to Regularized Damage or Plasticity Models

888 As shown in Section 4.3 there are circumstances under which the transient viscous
 889 rheology proposed in this work promotes a spontaneous organization of periods and re-
 890 gions of negligible anelastic strain rate and those of significant anelastic strain rate. Even
 891 though a critical yield stress seems to be an emergent rather than an inherent property
 892 of the system of equations and initial and boundary conditions, it is tempting to think
 893 of this rheology as a smooth plasticity model.

894 We have shown (Section 3.5) that our model requires regularization by a non-local
 895 strain rate measure to avoid spurious mesh dependence. Non-locality in one form or an-

896 other has been applied to combat ill-posedness in e.g. Bažant et al. (1984); Triantafyl-
897 lidis and Aifantis (1986); Schreyer and Chen (1986); Pijaudier-Cabot and Bažant (1987);
898 de Borst and Mhlhaus (1992); Peerlings et al. (1996); Jirásek (1998); Bažant and Jirásek
899 (2002); Engelen et al. (2003); Jirásek and Rolshoven (2009b, 2009a); Burghardt et al.
900 (2012); Lyakhovsky et al. (2011); Lyakhovsky and Ben-Zion (2014b, 2014a); Lyakhovsky
901 et al. (2016); Kurzon et al. (2019, 2020); Kiefer et al. (2018); Abdallah et al. (2020).

902 The comprehensive non-local damage-breakage rheology of Lyakhovsky et al. (2011);
903 Lyakhovsky and Ben-Zion (2014b, 2014a); Lyakhovsky et al. (2016), firmly rooted in ther-
904 modynamic theory and well-calibrated to match observations, is one of few damage the-
905 ories developed to cover the full process of earthquake generation and healing (Lyakhovsky
906 et al., 2016; Kurzon et al., 2019, 2020). As in our model, the nonlocality in theirs is of
907 Gradient type, but differs subtly in that it acts on the damage parameter rather than
908 on the anelastic shear strain rate. Another example is the Godunov-Peshkov-Romenski
909 model (S. Godunov & Romenskii, 1972; S. K. Godunov & Romenskii, 2003; Resnyan-
910 sky et al., 2003; Romenskii, 2007; Romenski et al., 2020), which was first used in Gabriel
911 et al. (2021) to simulate dynamic rupture and off-fault damage generation. This model
912 differs from our model and the aforementioned damage-breakage rheology in a funda-
913 mental way in that shear bands are produced by damage waves described by hyperbolic
914 equations rather than parabolic equations of reaction-diffusion type.

915 Our model differs from phase-field models of fracture in some respects. We have
916 used the distribution φ of anelastic shear strain rate as an analytical tool during the deriva-
917 tion of the continuum rheology proposed in this work. While this distribution could be
918 renormalized to form a phase field $d \in (0, 1)$, the necessary inclusion of a damping vis-
919 cosity in Section 3.6 causes temporal variations in the anelastic shear strain rate distri-
920 bution that transcend the modeled distributions, and make renormalization to a phase
921 field possible at best as a post-processing step in a simulation. We recognize that the
922 inclusion of a damping viscosity constitutes the use of a ‘double-well potential’ in phase
923 field terminology, which multiple authors have commented on critically in this context
924 (e.g. Kuhn et al., 2015; J.-Y. Wu, 2017).

925 Our linear viscous damping appears as a regularization technique in some plastic-
926 ity models (Needleman, 1988; Peirce et al., 1983; F. Wu & Freund, 1984; Duretz et al.,
927 2019, 2020, 2021; de Borst & Duretz, 2020; Stathas & Stefanou, 2022). In these mod-

928 els the Kelvin–Voigt arrangement of yield strength and Newtonian viscosity truncates
929 the steady-state anelastic strain rate that may be achieved, forcing a shear zone to have
930 a finite width in order to slide at a certain macroscopic rate. In our model, the same vis-
931 cosity also introduces an effective upper bound on anelastic strain rate attained during
932 the limit cycle and causes subsequent delocalization, but we must still rely on the gra-
933 dient regularization to combat the unbounded localization that would otherwise happen
934 even at low strain rate. A detailed analysis of the efficacy of this viscous regularization
935 in dynamical problems is made in Stathas and Stefanou (2022).

936 5.4 Relation to Other Transient Continuum Rheologies

937 The framework here proposed can be seen as a generalization of the work of Herrendörfer
938 et al. (2018); Preuss et al. (2019, 2020), who made the purely local substitution $v_0 V \rightarrow$
939 $h_x \gamma_0 \gamma$ and used a Drucker-Prager elastoplastic model similar to the one set out in Sec-
940 tion 2.4. Setting $\lambda_0, \eta = 0$ and assuming that strain rate fully localizes into a discrete
941 Dirac function sampled every h_x , we find that the coefficients c_1 and c_2 become h_x/d_c
942 and 1 respectively, substitution of which into (16) yields Herrendörfer’s version of the
943 aging law. Their model was first applied along a predefined staggered grid line in Herrendörfer
944 et al. (2018), obeying the discrete Dirac distribution of anelastic shear strain automat-
945 ically and effectively yielding a numerical method analogous to the stress glut method
946 of Andrews (1999). Herrendörfer’s model was subsequently applied in an unconstrained
947 evolving continuum model in Preuss et al. (2019, 2020), but notwithstanding measures
948 put in place that acknowledge the changing distribution of shear strain rate within a shear
949 zone, their model ultimately lacks regularizations that remove mesh dependence. In the
950 chapters titled ‘Localization of Deformation’ and ‘Relationship of Localization to Insta-
951 bility’ of his PhD thesis, Ruina (1980) gives a thoughtful take on aspects of the local-
952 ization behavior of a strain rate formulation of rate and state friction (without spatial
953 regularization), which is in some aspects in line with findings reported in this work, and
954 complementary in others.

955 As noted in Section 1.2 of the Introduction, a variety of continuum theories have
956 been developed to explain the general rate- and state-like behavior of deformation in faults
957 and shear zones (e.g. Sleep, 1997; Braeck & Podladchikov, 2007; Daub et al., 2008; Daub
958 & Carlson, 2008, 2009; Daub et al., 2010; Elbanna & Carlson, 2014; John et al., 2009;
959 Thielmann et al., 2015; Thielmann, 2018; Rozel et al., 2011; Barbot, 2019; Niemeijer &

960 Spiers, 2007; Chen & Spiers, 2016; Van den Ende et al., 2018; Roubek, 2014; Lyakhovsky
961 et al., 2011; Lyakhovsky & Ben-Zion, 2014b, 2014a; Lyakhovsky et al., 2016; Kurzon et
962 al., 2019, 2020; Poulet et al., 2014; Veveakis et al., 2014; Alevizos et al., 2014; Rattetz,
963 Stefanou, & Sulem, 2018; Rattetz, Stefanou, Sulem, Veveakis, & Poulet, 2018; Pozzi et
964 al., 2021). We note again that an internal length scale in these models typically arises
965 from the inclusion of a diffusion process (of e.g. temperature, pore pressure) but that
966 may not always be adjusted to meet the constraints imposed by scale and computational
967 power without changing the outcome of the model. In our proposed formulation an arti-
968 ficial diffusion process acts directly on the anelastic shear strain rate, resulting in a ro-
969 bust and controllable internal length scale.

970 In this last respect our work is preceded by nearly 25 years by Sleep (1997). In this
971 work Sleep combined and extended earlier works (Sleep, 1995; Segall & Rice, 1995; Linker
972 & Dieterich, 1992; Chester, 1994, 1995) in which rate and state friction was interpreted
973 as the product of crack generation and healing, associated rheological weakening, and
974 dissipative heating. This physical reasoning resulted in a model that contains only quan-
975 tities that are either directly measurable or can be modeled by independent methods.
976 This contrasts with our purely mathematical argumentation that serves to retain close
977 correspondence to the original phenomenological description of rate and state dependent
978 sliding on a frictional interface. Sleep (1997) neglects the fluxes and associated spatial
979 gradients of the pore fluid but does include heat diffusion, which they note does how-
980 ever not play a significant role at the scale of their numerical experiments. Instead, re-
981 sembling our approach, they impose an artificial length scale and forced strain distribu-
982 tion by explicit mollification of the anelastic shear strain rate with a Gaussian kernel.
983 This leads to an aging law (Equation 53 in Sleep (1997)) that is structurally identical
984 to our result (16). Like us, they find that strain localization can only occur when $a <$
985 b , and that a rate-strengthening effect that activates at high strain rate leads to strain
986 delocalization. With respect to Sleep's valuable contribution, in this work we provide
987 a more complete argumentation for this type of spatial regularization and analysis of the
988 resulting patterns of strain localization and delocalization over the seismic cycle.

989 5.5 Assumptions and Future Work

990 We proposed our model in a very general three-dimensional continuum mechan-
991 ics framework, but for simplicity have considered only a small fault neighborhood in which

992 in-plane variations of fault properties can be neglected (Section 2.2) so that the model
993 becomes effectively one-dimensional. This same assumption was also taken in the nu-
994 merical models that we have used in our analysis. The assumption is clearly violated around
995 fault branches, at fault kinks or on rough faults, and near the fault or rupture tip. We
996 note however that rate and state friction was proposed based on laboratory studies that
997 also neglect these geometrical complexities. Although rate and state faults with branches
998 and kinks are still largely non-standard in present-day numerical modeling studies, plenty
999 of attention has been given to the critical nucleation patch and the structure of the co-
1000 hesive zone near the rupture tip (e.g. Rice, 1993; Cocco & Bizzarri, 2002; Cocco et al.,
1001 2004; Rubin & Ampuero, 2005; Day et al., 2005; Lapusta & Liu, 2009; Viesca, 2016b,
1002 2016a; Putelat et al., 2017). Given our main assumption, these features can only be ac-
1003 curately reproduced with our continuum formulation in higher-dimensional numerical
1004 models if the regularization length scale λ_0 is significantly smaller than the length scales
1005 associated with the critical nucleation patch and the cohesive zone. In turn the cell size
1006 must be sufficient to resolve λ_0 , and so we expect to need a grid resolution that is sig-
1007 nificantly higher than that of existing methods to simulate rate and state frictional in-
1008 terfaces. We have also seen that the diffusion process is associated with a smaller time
1009 scale, and thus, stricter time step constraints than the interfacial model. We note that
1010 both spatial and temporal resolution requirements already place challenging constraints
1011 on simulations of seismic and aseismic slip sequences (Erickson et al., 2020; Jiang et al.,
1012 2022).

1013 It seems prudent to first make a more detailed assessment of the computational de-
1014 mands and the techniques that may be required to meet the resolution requirements (e.g.
1015 adaptive mesh refinement, local adaptive time stepping). In the process of constructing
1016 higher-dimensional models one may first concentrate on the friendliest regions of param-
1017 eter space, for example those that promote stable sliding, or undamped runaway local-
1018 ization as in Viesca (2016b, 2016a, 2020).

1019 The temporal patterns of localization and delocalization that occur in our model
1020 yield testable predictions that tie in to a recent surge in interest in similar patterns ob-
1021 served in the lab and in nature (e.g. Ben-Zion & Zaliapin, 2020; McBeck et al., 2018; McBeck,
1022 Aiken, et al., 2020; McBeck, Ben-Zion, & Renard, 2020; McBeck et al., 2021). It could
1023 be the scope of future research to reinterpret anelastic strain rate in our model as a mea-
1024 sure of the activity of a statistical distribution of cracks of various properties and com-

1025 pare to aforementioned lab and field observations. In this context, the delocalization that
1026 is in our models induced by a rate-limiting Kelvin viscosity is reminiscent of the growth
1027 of off-fault fracture networks during dynamic rupture (e.g. Templeton & Rice, 2008; K. Okubo
1028 et al., 2019; Gabriel et al., 2021).

1029 We have assumed an infinite gauge layer and observed that materials that are rate-
1030 strengthening at steady state ($a > b$) can only feature delocalization of strain. This ap-
1031 pears inconsistent with the widespread observation of faults and fault gauges that are
1032 more or less stable (e.g. Ikari et al., 2011; Carpenter et al., 2012; Coble et al., 2014; Bed-
1033 ford et al., 2021). A possible explanation is that a gauge layer is initially formed by a
1034 set of strain-softening processes, subsequently chemically, petrologically, texturally or ge-
1035 ometrically matured over time to become rate-strengthening, but generally weaker than
1036 host rock. The gauge layer walls may then act as a barrier to further delocalization. While
1037 this situation could be simulated in our approach by varying the frictional properties ac-
1038 cordingly, the consistent way to model the long term evolution of faults likely requires
1039 at least two internal state variables. The damage-breakage rheology of Lyakhovsky and
1040 Ben-Zion (2014b, 2014a) has this feature.

1041 In this work, we have restricted ourselves to classical Dieterich-Ruina rate and state
1042 friction with aging law. In Section 1.2, we have hinted at a sizable number of physical
1043 interpretations of the phenomenological rate and state friction problem. This concerns
1044 notably the behavior of weakening and the interpretation given to the ‘state’ variable.
1045 At the same time, a number of alternative phenomenological evolution laws have been
1046 proposed, among which the slip law (Dieterich, 1979a). It would be worthwhile to in-
1047 vestigate the possibility of applying the mathematical techniques developed here to a wider
1048 range of friction and evolution laws, especially those involving multiple states, such as
1049 for example temperature, pore pressure, or grain size. However, this might not be straight-
1050 forward because we have relied heavily on analytical solutions to calibrate the interfa-
1051 cial and continuum descriptions of friction, and finding these solutions is frequently a
1052 time-consuming task with uncertain prospects for success.

1053 6 Conclusions

1054 In this work we have carefully constructed a coordinate-invariant and mesh-independent
1055 transient visco-elastic continuum rheology that behaves in a way that is consistent with

1056 rate and state friction on an interface. We have shown that inclusion of a diffusion-like
1057 spatial regularization ensures a limit to strain localization and thus guarantees mesh con-
1058 vergence. In a simplified 1D fault transect, important metrics of the seismic cycle – such
1059 as slip rate and friction – are independent of the diffusion length scale associated with
1060 the regularization. However, throughout this work we have assumed the regularization
1061 length scale to be small compared to the length scales associated with other features of
1062 interest, such as fault curvature or along-strike variations of slip rate and stress. There-
1063 fore, going forward, high-resolution 2D or 3D numerical models are required to apply this
1064 model to the study of the seismogenic behavior of emerging and evolving fault zone net-
1065 works. Our continuum rheology resembles a reaction-diffusion equation for anelastic strain
1066 rate. Processes described by such equations are ubiquitous in nature, and it is tempt-
1067 ing to compare temporal patterns of localization and delocalization produced by our model
1068 with natural observations.

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1086 ‘Scientific Colour Maps’ package (Cramer, 2021) for distortion-free representation of the
1087 model results – also for readers with color vision deficiencies (Cramer et al., 2020).

1088 **Data Availability Statement**

1089 All codes and algorithms to generate and visualize the results discussed in this work
 1090 can be found as Wolfram Mathematica (Wolfram Research, Inc., 2017) notebooks in the
 1091 supplementary information.

1092 **Appendix A Strain Localization in the Runaway Slip Regime**

1093 In this appendix we derive a generic family of solutions to the runaway slip regime
 1094 discussed in Section 3.2. We consider the ODE

1095
$$\dot{\gamma}(t, x) = \gamma(t, x)^2, \quad (\text{A1})$$

 1096

1097 in which t and γ are taken to be non-dimensionalized in such a way that the rate con-
 1098 stant equals one and is thus dropped. We emphasize that since no derivatives with re-
 1099 spect to x are included, the ultimate collective behavior of $\gamma(t, x)$ is not uniquely defined.
 1100 Even so, our Ansatz is a product-wise decomposition of the solution into a purely time-
 1101 dependent term Γ and a self-similar contribution g :

1102
$$\gamma(t, x) =: \Gamma(t)g\left(\frac{x}{l(t)}\right), \quad g(0) = 1. \quad (\text{A2})$$

 1103

1104 We then take the time derivative of the Ansatz (A2) and equate the result to (A1), yield-
 1105 ing

1106
$$\dot{\gamma}(t, x) = \gamma(t, x)^2 = \dot{\Gamma}(t)g\left(\frac{x}{l(t)}\right) - \Gamma(t)\frac{\dot{l}(t)}{l(t)}\frac{x}{l(t)}g'\left(\frac{x}{l(t)}\right) = \Gamma(t)^2g\left(\frac{x}{l(t)}\right)^2. \quad (\text{A3})$$

 1107

1108 At $x = 0$, we have:

1109
$$\dot{\Gamma}(t) = \Gamma(t)^2. \quad (\text{A4})$$

 1110

1111 The solution to (A4) is

1112
$$\Gamma(t) = \Gamma_0 (1 - \Gamma_0 t)^{-1}, \quad \Gamma_0 = \Gamma(t = 0). \quad (\text{A5})$$

 1113

1114 When we substitute this solution into (A3), we get

1115
$$\frac{\dot{l}(t)}{l(t)}\frac{x}{l(t)}g'\left(\frac{x}{l(t)}\right) = -\Gamma(t)\left[g\left(\frac{x}{l(t)}\right)^2 - g\left(\frac{x}{l(t)}\right)\right]. \quad (\text{A6})$$

 1116

1117 If we now take the additional Ansatz that

1118
$$\frac{x}{l(t)}g'\left(\frac{x}{l(t)}\right) = m\left[g\left(\frac{x}{l(t)}\right)^2 - g\left(\frac{x}{l(t)}\right)\right], \quad m \in \mathbb{R} \geq 2 \quad (\text{A7})$$

 1119

1120 then we obtain

$$1121 \quad g(t, x) = \left(1 + 2 \left| \frac{x}{l(t)} \right|^m\right)^{-1}, \quad (A8)$$

1122 and

$$1124 \quad l(t) = l_0 (1 - l_0 t)^{\frac{1}{m}}, \quad l_0 = l(t = 0), \quad (A9)$$

1126 The function $g(t, x/l(t))$ integrates to

$$1127 \quad \Lambda(t) := \int_{-\infty}^{+\infty} g\left(t, \frac{\xi}{l(t)}\right) d\xi = l(t) \frac{\pi}{m} \csc \frac{\pi}{m}, \quad (A10)$$

1129 and thus the *distribution* $\varphi(t, x)$ corresponding to $g(t, x/l(t))$ is given by

$$1130 \quad \varphi(t, x) = \Lambda(t)^{-1} \left(1 + 2 \left| \frac{x}{l(t)} \right|^m\right)^{-1}, \quad (A11)$$

1132 Finally, we use (7) and (A11) to write (A5), (A8), and (A9) in terms of $V(t)$, and
1133 obtain:

$$1134 \quad V(t) \propto \left(1 - \frac{t}{t_*}\right)^{m^{-1}-1},$$

$$1135 \quad l(t) \propto V(t)^{(1-m)^{-1}}. \quad (A12a)$$

1137 Appendix B Derivation of the Steady-state Strain Rate Distribution

1138 Here we derive the steady-state solution presented in Section 3.4 to the time-dependent
1139 PDE

$$1140 \quad \dot{\varphi} \propto \lambda_0 c_1 \varphi \mathcal{M}_{\text{ex}}(\varphi) - \lambda_0^{\frac{a}{b}} c_2 \varphi^{1+\frac{a}{b}}, \quad (B1)$$

1142 subject to $a, b > 0$ and $a < b$, with \mathcal{M}_{ex} given by (26) and the coefficients c_1 and c_2
1143 measured from the sought solution $\varphi(\infty, x)$ by means of (20a) and (20b). We summa-
1144 rize the PDE that is to be solved as

$$1145 \quad \varphi^2 + \varphi \lambda_0^2 \partial_x^2 \varphi - c \varphi^{1+\frac{a}{b}} = 0, \quad (B2)$$

1147 with c a parameter to be determined later. Dividing the whole equation by φ , and mak-
1148 ing the substitution $\lambda_0 \rightarrow [1 - \frac{a}{b}] \lambda_1$, (B2) becomes

$$1149 \quad \varphi + \left[1 - \frac{a}{b}\right]^2 \lambda_1^2 \partial_x^2 \varphi - \xi \varphi^{\frac{a}{b}} = 0. \quad (B3)$$

1151 We now introduce $f(\varphi) = [1 - \frac{a}{b}] \lambda_1 \partial_x \varphi(x)$, which allows us to write

$$1152 \quad \left[1 - \frac{a}{b}\right]^2 \lambda_1^2 \partial_x^2 \varphi = [1 - \frac{a}{b}] \lambda_1 \partial_x f(\varphi) = [1 - \frac{a}{b}] \lambda_1 f'(\varphi) \partial_x \varphi = f'(\varphi) f(\varphi). \quad (B4)$$

1154 Applying this reasoning to (B3) and reorganizing slightly gives

$$1155 \quad f'(\varphi)f(\varphi) = c\varphi^{\frac{a}{b}} - \varphi. \quad (B5)$$

1157 We now integrate (B5) by parts as follows:

$$1158 \quad \int_0^\varphi f'(\varepsilon)f(\varepsilon)d\varepsilon = \frac{1}{2} (f(\varphi)^2 - f(0)^2) = c \int_0^\varphi \varepsilon^{\frac{a}{b}} d\varepsilon - \int_0^\varphi \varepsilon d\varepsilon = c \left[1 + \frac{a}{b}\right]^{-1} \varphi^{1+\frac{a}{b}} - 2^{-1}\varphi^2. \quad (B6)$$

1159 We arbitrarily take the distribution $\varphi(\infty, x)$ to be symmetric around $x = 0$, where it
 1160 reaches its maximum value $\varphi_0 = \varphi(\infty, 0)$. At this point, as well as at the extremities
 1161 of the distribution where $\varphi = 0$, we set the gradient, i.e. $f(\varphi)$, to zero. This eliminates
 1162 the term $f(0)^2$ in (B6) and constrains the free parameter c to

$$1164 \quad c = \frac{1}{2} \left[1 + \frac{a}{b}\right] \varphi_0^{1-\frac{a}{b}}. \quad (B7)$$

1166 We now multiply both sides of (B6) by 2 and take the square root to obtain

$$1167 \quad f(\varphi) = \sqrt{\varphi_0^{1-\frac{a}{b}} \varphi^{1+\frac{a}{b}} - \varphi^2}. \quad (B8)$$

1169 Applying the inverse function theorem to $f(\varphi)^{-1} = \left(\left[1 - \frac{a}{b}\right] \lambda_1^{-1} \varphi'(x)\right)^{-1} = \left(\left[1 - \frac{a}{b}\right] \lambda_1\right)^{-1} x'(\varphi)$,
 1170 we obtain from (B8)

$$1171 \quad \frac{x'(\varphi)}{\lambda_1} = \left[1 - \frac{a}{b}\right] \left(\varphi_0^{1-\frac{a}{b}} \varphi^{1+\frac{a}{b}} - \varphi^2\right)^{-\frac{1}{2}}. \quad (B9)$$

1173 We seek to integrate this relation once more over the region $\varepsilon \in [0, \varphi]$ to obtain a so-
 1174 lution for $x(\varphi)$. The integral of the right hand side of (B8) can be reverse-engineered
 1175 from the known derivatives

$$1176 \quad \cos^{-1}(f) = -(1 - f^2)^{-\frac{1}{2}}. \quad (B10)$$

1178 and $d_f f^\alpha = \alpha f^{\alpha-1}$, noting also that $\varphi/\varphi_0 \leq 1$, to give

$$1179 \quad \left[1 - \frac{a}{b}\right] \left(\varphi_0^{1-\frac{a}{b}} \varphi^{1+\frac{a}{b}} - \varphi^2\right)^{-\frac{1}{2}} = \partial_\varphi \left[-2 \cos^{-1} \left(\left[\frac{\varphi}{\varphi_0} \right]^{\frac{1}{2} \left[1 - \frac{a}{b}\right]} \right) \right]. \quad (B10)$$

1181 This yields

$$1182 \quad \frac{x(\varphi)}{\lambda_1} - \frac{x(0)}{\lambda_1} = -2 \cos^{-1} \left(\left[\frac{\varphi}{\varphi_0} \right]^{\frac{1}{2} \left[1 - \frac{a}{b}\right]} \right) + \pi. \quad (B11)$$

1184 Setting $\frac{x(0)}{\lambda_1} = -\pi$ and inverting for $\varphi(\infty, x)$ gives

$$1185 \quad \varphi(\infty, x) = \varphi_0 \cos \left(\frac{1}{2} \frac{x}{\lambda_1} \right)^{2 \left[1 - \frac{a}{b}\right]^{-1}}. \quad (B12)$$

1187 It is possible and in fact desirable to isolate a single strain pulse in the range $x \in [-\pi\lambda_1, +\pi\lambda_1]$:

$$1188 \quad \varphi(\infty, x) = \begin{cases} \varphi_0 \cos\left(\frac{1}{2}\frac{x}{\lambda_1}\right)^2 [1 - \frac{a}{b}]^{-1} & \forall x \in [-\pi\lambda_1, +\pi\lambda_1] \\ 0 & \forall x \notin [-\pi\lambda_1, +\pi\lambda_1]. \end{cases} \quad (B13)$$

1189 Finally, we determine the coefficient φ_0 that ensures that the distribution φ integrates
1190 to one over its domain, and the coefficients c_1 and c_2 following the expressions (20a) and
1191 (20b). Definite integrals of (B13) are evaluated using Wolfram Mathematica (Wolfram
1192 Research, Inc., 2017), and we find

$$1194 \quad \varphi_0 = \frac{1}{2}\pi^{-\frac{1}{2}} \frac{\Gamma\left(1 + [1 - \frac{a}{b}]^{-1}\right)}{\Gamma\left(\frac{1}{2} + [1 - \frac{a}{b}]^{-1}\right)} \lambda_1^{-1} \quad (B14a)$$

$$1195 \quad c_1 = \frac{1}{2}\pi^{-\frac{1}{2}} \frac{\Gamma\left(1 + 2[1 - \frac{a}{b}]^{-1}\right)}{\Gamma\left(\frac{1}{2} + 2[1 - \frac{a}{b}]^{-1}\right)} [\lambda_1\varphi_0]^{-2} \quad (B14b)$$

$$1196 \quad c_2 = \frac{1}{2}\pi^{-\frac{1}{2}} \frac{\Gamma\left(2[1 - \frac{a}{b}]^{-1}\right)}{\Gamma\left(\frac{1}{2}[3 + \frac{a}{b}][1 - \frac{a}{b}]^{-1}\right)} [\lambda_1\varphi_0]^{-1 - \frac{a}{b}}. \quad (B14c)$$

1197 We have verified that (B13) is a solution to (B2) given (B7) using automated symbolic
1198 manipulation in Wolfram Mathematica.
1199

1200 Appendix C Derivation of the Runaway Strain Rate Distribution

1201 We simply note that the time-dependent PDE

$$1202 \quad \dot{\varphi} \propto V[\lambda_0 c_1 \varphi \mathcal{M}_{\text{ex}}(\varphi) - \varphi] \quad (C1)$$

1203 to which a steady-state solution is sought in Section 3.5, closely relates to (B1) if in that
1204 equation the substitution $\frac{a}{b} \rightarrow 0$ is made. Making the same substitution in the result
1205 of Appendix B, we obtain:
1206

$$1207 \quad \varphi(\infty, x) = \begin{cases} \frac{1}{\pi\lambda} \cos\left(\frac{1}{2}\frac{x}{\lambda_0}\right)^2 & \forall x \in [-\pi\lambda_0, +\pi\lambda_0] \\ 0 & \forall x \notin [-\pi\lambda_0, +\pi\lambda_0]. \end{cases} \quad (C2)$$

1208 Appendix D Numerical Solution Procedure

1209 We discretize the Laplacian using a standard second-order accurate central differ-
1210 ence stencil, with natural boundary conditions of the same accuracy implemented by stag-
1211 gering the fields with respect to the physical domain walls. The problem size is halved
1212

1213 by exploiting symmetry across the shear zone. Integrals are evaluated using a midpoint
 1214 rule. This gives the discrete system

$$1215 \quad \dot{\tilde{\zeta}} = \mathbf{g}(\tilde{\zeta}, \tilde{f})$$

$$1216 \quad \dot{\tilde{f}} = \mathbf{h}(\tilde{\zeta}, \tilde{f}),$$

1218 where tildes indicate approximate space-discrete quantities and numerical arrays are in-
 1219 dicated in bold face. We form a symbolic rules for computing the Jacobian matrix \mathbf{J} of
 1220 this system of non-linear space-discrete equations,

$$1221 \quad \mathbf{J}(\tilde{\zeta}, \tilde{f}) = \begin{bmatrix} \nabla_{\tilde{\zeta}} \mathbf{g}(\tilde{\zeta}, \tilde{f}) & \nabla_{\tilde{f}} \mathbf{g}(\tilde{\zeta}, \tilde{f}) \\ \nabla_{\tilde{\zeta}} \mathbf{h}(\tilde{\zeta}, \tilde{f}) & \nabla_{\tilde{f}} \mathbf{h}(\tilde{\zeta}, \tilde{f}) \end{bmatrix},$$

1223 with the upper left block a dense $N \times N$ matrix with a dominant sparse band structure
 1224 (N being the problem size), the lower right block a 1×1 empty matrix, and the off-diagonal
 1225 blocks densely populated vectors of compatible shape. The system is then linearized as

$$1226 \quad \mathbf{d} \begin{bmatrix} \dot{\tilde{\zeta}} \\ \dot{\tilde{f}} \end{bmatrix} = \mathbf{J}(\tilde{\zeta}, \tilde{f}) \mathbf{d} \begin{bmatrix} \tilde{\zeta} \\ \tilde{f} \end{bmatrix}$$

1228 and the smallest time scale $S_t(\tilde{\zeta}, \tilde{f})$ to be resolved is computed as the reciprocal of the
 1229 largest eigenvalue of the Jacobian matrix \mathbf{J} . The actual time step Δ_t is computed as some
 1230 fixed fraction of S_t , optionally bounded by a maximum value and/or maximum growth
 1231 rate to prevent time step overestimation as the time scale increases. The equations are
 1232 discretized in time with forward and backward Euler schemes, both first-order accurate,
 1233 respectively as

$$1234 \quad \begin{bmatrix} \tilde{\zeta}^{k+1} \\ \tilde{f}^{k+1} \end{bmatrix} = \begin{bmatrix} \tilde{\zeta}^k \\ \tilde{f}^k \end{bmatrix} + \Delta_t \begin{bmatrix} \mathbf{g}(\tilde{\zeta}^k, \tilde{f}^k) \\ \mathbf{h}(\tilde{\zeta}^k, \tilde{f}^k) \end{bmatrix} \quad (\text{D1})$$

$$1235 \quad \mathbf{f} = \begin{bmatrix} \tilde{\zeta}^{k+1} \\ \tilde{f}^{k+1} \end{bmatrix} - \begin{bmatrix} \tilde{\zeta}^k \\ \tilde{f}^k \end{bmatrix} - \Delta_t \begin{bmatrix} \mathbf{g}(\tilde{\zeta}^{k+1}, \tilde{f}^{k+1}) \\ \mathbf{h}(\tilde{\zeta}^{k+1}, \tilde{f}^{k+1}) \end{bmatrix} = \mathbf{0}. \quad (\text{D2})$$

1237 Our algorithm makes an explicit-in-time prediction using (D1) and evaluates the l_2 norm
 1238 of the implicit-in-time residual \mathbf{f} . It includes the possibility to perform Newton-Raphson
 1239 iterations using the Jacobian \mathbf{J} to keep the residual bounded, although we find it to be
 1240 more efficient to experimentally set the dimensionless time step Δ_t/S_t sufficiently small
 1241 (e.g. 0.1) to never cause tolerances to be violated. This value is then taken to correspond
 1242 to a stable time step size of the explicit problem.

1243 We make our algorithms and scripts to produce figures available as Wolfram Math-
 1244 ematica (Wolfram Research, Inc., 2017) notebooks in the Supplementary Materials.

1245 References

- 1246 Abdallah, Y., Sulem, J., & Stefanou, I. (2020). Compaction banding in high-porosity
 1247 carbonate rocks: 2. a gradient-dependent plasticity model. *Journal of Geophys-*
 1248 *ical Research: Solid Earth*, *125*(12), e2020JB020610.
- 1249 Alevizos, S., Poulet, T., & Veveakis, E. (2014). Thermo-poro-mechanics of chemi-
 1250 cally active creeping faults. 1: Theory and steady state considerations. *Journal*
 1251 *of Geophysical Research: Solid Earth*, *119*(6), 4558–4582.
- 1252 Andrews, D. (1999). Test of two methods for faulting in finite-difference calculations.
 1253 *Bulletin of the Seismological Society of America*, *89*(4), 931–937.
- 1254 Bak, P., & Tang, C. (1989). Earthquakes as a self-organized critical phenomenon.
 1255 *Journal of Geophysical Research: Solid Earth*, *94*(B11), 15635–15637. doi: 10
 1256 .1029/JB094iB11p15635
- 1257 Barbot, S. (2019). Modulation of fault strength during the seismic cycle by grain-
 1258 size evolution around contact junctions. *Tectonophysics*, *765*, 129 - 145. doi:
 1259 10.1016/j.tecto.2019.05.004
- 1260 Barbot, S., & Fialko, Y. (2010). A unified continuum representation of post-seismic
 1261 relaxation mechanisms: semi-analytic models of afterslip, poroelastic rebound
 1262 and viscoelastic flow. *Geophysical Journal International*, *182*, 1124–1140.
- 1263 Barry, D., Barry, S., & Culligan-Hensley, P. (1995). Algorithm 743: Wapr—a for-
 1264 tran routine for calculating real values of the w-function. *ACM Transactions*
 1265 *on Mathematical Software (TOMS)*, *21*(2), 172–181.
- 1266 Barry, D., Culligan-Hensley, P., & Barry, S. (1995). Real values of the w-function.
 1267 *ACM Transactions on Mathematical Software (TOMS)*, *21*(2), 161–171.
- 1268 Barth, N. C., Boulton, C., Carpenter, B. M., Batt, G. E., & Toy, V. G. (2013). Slip
 1269 localization on the southern alpine fault, new zealand. *Tectonics*, *32*(3), 620-
 1270 640. doi: 10.1002/tect.20041
- 1271 Bažant, Z. P., Belytschko, T. B., & Chang, T.-P. (1984). Continuum theory for
 1272 strain-softening. *Journal of Engineering Mechanics*, *110*(12), 1666–1692.
- 1273 Bažant, Z. P., & Jirásek, M. (2002). Nonlocal integral formulations of plasticity and
 1274 damage: survey of progress. *Journal of engineering mechanics*, *128*(11), 1119–

- 1275 1149.
- 1276 Bedford, J. D., Faulkner, D. R., Allen, M. J., & Hirose, T. (2021). The stabilizing
1277 effect of high pore-fluid pressure along subduction megathrust faults: Evidence
1278 from friction experiments on accretionary sediments from the Nankai Trough.
1279 *Earth and Planetary Science Letters*, *574*, 117161.
- 1280 Behr, W. M., Gerya, T. V., Cannizzaro, C., & Blass, R. (2021). Transient slow slip
1281 characteristics of frictional-viscous subduction megathrust shear zones. *AGU*
1282 *Advances*, *2*(3), e2021AV000416.
- 1283 Ben-Zion, Y., & Rice, J. R. (1995). Slip patterns and earthquake populations along
1284 different classes of faults in elastic solids. *Journal of Geophysical Research:*
1285 *Solid Earth*, *100*(B7), 12959–12983.
- 1286 Ben-Zion, Y., & Rice, J. R. (1997). Dynamic simulations of slip on a smooth fault in
1287 an elastic solid. *Journal of Geophysical Research: Solid Earth*, *102*(B8), 17771-
1288 17784. doi: 10.1029/97JB01341
- 1289 Ben-Zion, Y., & Zaliapin, I. (2020). Localization and coalescence of seismicity before
1290 large earthquakes. *Geophysical Journal International*, *223*(1), 561-583. doi: 10
1291 .1093/gji/ggaa315
- 1292 Bouzid, M., Trulsson, M., Claudin, P., Clément, E., & Andreotti, B. (2013). Non-
1293 local rheology of granular flows across yield conditions. *Physical review letters*,
1294 *111*(23), 238301.
- 1295 Bouzid, M., Trulsson, M., Claudin, P., Clément, E., & Andreotti, B. (2015). Mi-
1296 cro-rheology to probe non-local effects in dense granular flows. *EPL (Euro-*
1297 *physics Letters)*, *109*(2), 24002.
- 1298 Bowden, F. P., & Tabor, D. (1966). Friction, lubrication and wear: a survey of work
1299 during the last decade. *British Journal of Applied Physics*, *17*(12), 1521–1544.
1300 doi: 10.1088/0508-3443/17/12/301
- 1301 Brace, W., & Byerlee, J. (1966). Stick-slip as a mechanism for earthquakes. *Science*,
1302 *153*(3739), 990–992.
- 1303 Braeck, S., & Podladchikov, Y. Y. (2007). Spontaneous thermal runaway as an ul-
1304 timate failure mechanism of materials. *Phys. Rev. Lett.*, *98*, 095504. doi: 10
1305 .1103/PhysRevLett.98.095504
- 1306 Broadbridge, P., & Bradshaw-Hajek, B. H. (2016). Exact solutions for logistic
1307 reaction–diffusion equations in biology. *Zeitschrift für angewandte Mathematik*

- 1308 *und Physik*, 67(4), 93. doi: 10.1007/s00033-016-0686-3
- 1309 Burghardt, J., Brannon, R., & Guilkey, J. (2012). A nonlocal plasticity formu-
1310 lation for the material point method. *Computer methods in applied mechanics*
1311 *and engineering*, 225, 55–64.
- 1312 Carpenter, B., Saffer, D., & Marone, C. (2012). Frictional properties and sliding sta-
1313 bility of the san andreas fault from deep drill core. *Geology*, 40(8), 759–762.
- 1314 Chen, J., & Spiers, C. J. (2016). Rate and state frictional and healing behavior
1315 of carbonate fault gouge explained using microphysical model. *Journal of Geo-*
1316 *physical Research: Solid Earth*, 121(12), 8642–8665.
- 1317 Chester, F. M. (1994). Effects of temperature on friction: Constitutive equations
1318 and experiments with quartz gouge. *Journal of Geophysical Research: Solid*
1319 *Earth*, 99(B4), 7247–7261.
- 1320 Chester, F. M. (1995). A rheologic model for wet crust applied to strike-slip faults.
1321 *Journal of Geophysical Research: Solid Earth*, 100(B7), 13033–13044.
- 1322 Chester, F. M., & Chester, J. S. (1998). Ultracataclasite structure and friction pro-
1323 cesses of the punchbowl fault, san andreas system, california. *Tectonophysics*,
1324 295(1-2), 199–221.
- 1325 Coble, C. G., French, M. E., Chester, F. M., Chester, J. S., & Kitajima, H. (2014).
1326 In situ frictional properties of San Andreas Fault gouge at SAFOD. *Geophysi-*
1327 *cal Journal International*, 199(2), 956-967. Retrieved from [https://doi.org/](https://doi.org/10.1093/gji/ggu306)
1328 [10.1093/gji/ggu306](https://doi.org/10.1093/gji/ggu306) doi: 10.1093/gji/ggu306
- 1329 Cocco, M., & Bizzarri, A. (2002). On the slip-weakening behavior of rate- and state
1330 dependent constitutive laws. *Geophysical Research Letters*, 29(11), 11-1-11-4.
1331 doi: 10.1029/2001GL013999
- 1332 Cocco, M., Bizzarri, A., & Tinti, E. (2004). Physical interpretation of the breakdown
1333 process using a rate- and state-dependent friction law. *Tectonophysics*, 378(3),
1334 241 - 262. (Physics of Active Faults - Theory, Observation and Experiments)
1335 doi: 10.1016/j.tecto.2003.09.015
- 1336 Cochard, A., & Madariaga, R. (1994). Dynamic faulting under rate-dependent fric-
1337 tion. *Pure and Applied Geophysics*, 142, 419-445. doi: 10.1007/BF00876049
- 1338 Cochard, A., & Madariaga, R. (1996). Complexity of seismicity due to highly rate-
1339 dependent friction. *Journal of Geophysical Research: Solid Earth*, 101(B11),
1340 25321–25336.

- 1341 Cramer, F. (2021). *Scientific colour maps*. Zenodo. Retrieved from [https://doi](https://doi.org/10.5281/zenodo.5501399)
1342 [.org/10.5281/zenodo.5501399](https://doi.org/10.5281/zenodo.5501399) doi: 10.5281/zenodo.5501399
- 1343 Cramer, F., Shephard, G. E., & Heron, P. J. (2020). The misuse of colour in science
1344 communication. *Nature communications*, *11*(1), 1–10.
- 1345 DalZilio, L., Lapusta, N., Avouac, J.-P., & Gerya, T. (2021). Subduction earthquake
1346 sequences in a non-linear visco-elasto-plastic megathrust. *Geophysical Journal*
1347 *International*, *229*(2), 1098–1121. Retrieved from [https://doi.org/10.1093/](https://doi.org/10.1093/gji/ggab521)
1348 [gji/ggab521](https://doi.org/10.1093/gji/ggab521) doi: 10.1093/gji/ggab521
- 1349 Daub, E. G., & Carlson, J. M. (2008). A constitutive model for fault gouge deforma-
1350 tion in dynamic rupture simulations. *Journal of Geophysical Research: Solid*
1351 *Earth*, *113*(B12).
- 1352 Daub, E. G., & Carlson, J. M. (2009). Stick-slip instabilities and shear strain local-
1353 ization in amorphous materials. *Physical Review E*, *80*(6), 066113.
- 1354 Daub, E. G., Manning, M. L., & Carlson, J. M. (2008). Shear strain localization in
1355 elastodynamic rupture simulations. *Geophysical research letters*, *35*(12).
- 1356 Daub, E. G., Manning, M. L., & Carlson, J. M. (2010). Pulse-like, crack-like, and
1357 supershear earthquake ruptures with shear strain localization. *Journal of Geo-*
1358 *physical Research: Solid Earth*, *115*(B5).
- 1359 Day, S. M., Dalguer, L. A., Lapusta, N., & Liu, Y. (2005). Comparison of fi-
1360 nite difference and boundary integral solutions to three-dimensional sponta-
1361 neous rupture. *Journal of Geophysical Research: Solid Earth*, *110*(B12). doi:
1362 [10.1029/2005JB003813](https://doi.org/10.1029/2005JB003813)
- 1363 de Borst, R., & Duretz, T. (2020). On viscoplastic regularisation of strain-softening
1364 rocks and soils. *International Journal for Numerical and Analytical Methods in*
1365 *Geomechanics*, *44*(6), 890–903. doi: 10.1002/nag.3046
- 1366 de Borst, R., & Mhlhaus, H.-B. (1992). Gradient-dependent plasticity: Formulation
1367 and algorithmic aspects. *International Journal for Numerical Methods in Engi-*
1368 *neering*, *35*(3), 521–539. doi: 10.1002/nme.1620350307
- 1369 Dieterich, J. H. (1978). Time-dependent friction and the mechanics of stick-slip.
1370 *pure and applied geophysics*, *116*(4), 790–806. doi: 10.1007/BF00876539
- 1371 Dieterich, J. H. (1979a). Modeling of rock friction: 1. experimental results and
1372 constitutive equations. *Journal of Geophysical Research: Solid Earth*, *84*(B5),
1373 2161–2168. doi: 10.1029/JB084iB05p02161

- 1374 Dieterich, J. H. (1979b). Modeling of rock friction: 2. simulation of preseismic
 1375 slip. *Journal of Geophysical Research: Solid Earth*, *84*(B5), 2169–2175. doi:
 1376 10.1029/JB084iB05p02169
- 1377 Dieterich, J. H., & Kilgore, B. D. (1994). Direct observation of frictional contacts:
 1378 New insights for state-dependent properties. *pure and applied geophysics*,
 1379 *143*(1), 283–302. doi: 10.1007/BF00874332
- 1380 Drucker, D. C., & Prager, W. (1952). Soil mechanics and plastic analysis or limit de-
 1381 sign. *Quarterly of applied mathematics*, *10*(2), 157–165.
- 1382 Duretz, T., de Borst, R., & Le Pourhiet, L. (2019). Finite thickness of shear bands
 1383 in frictional viscoplasticity and implications for lithosphere dynamics. *Geo-*
 1384 *chemistry, Geophysics, Geosystems*. doi: 10.1029/2019GC008531
- 1385 Duretz, T., de Borst, R., & Yamato, P. (2021). Modeling lithospheric deforma-
 1386 tion using a compressible visco-elasto-viscoplastic rheology and the effective
 1387 viscosity approach. *Geochemistry, Geophysics, Geosystems*, *22*(8). doi:
 1388 10.1029/2021GC009675
- 1389 Duretz, T., de Borst, R., Yamato, P., & Le Pourhiet, L. (2020). Toward ro-
 1390 bust and predictive geodynamic modeling: The way forward in frictional
 1391 plasticity. *Geophysical Research Letters*, *47*(5), e2019GL086027. doi:
 1392 10.1029/2019GL086027
- 1393 Duretz, T., Souche, A., de Borst, R., & Le Pourhiet, L. (2018). The benefits of using
 1394 a consistent tangent operator for viscoelastoplastic computations in geodynam-
 1395 ics. *Geochemistry, Geophysics, Geosystems*, *19*(12), 4904–4924.
- 1396 Elbanna, A. E., & Carlson, J. M. (2014). A two-scale model for sheared fault gouge:
 1397 Competition between macroscopic disorder and local viscoplasticity. *Journal of*
 1398 *Geophysical Research: Solid Earth*, *119*(6), 4841–4859.
- 1399 Engelen, R. A., Geers, M. G., & Baaijens, F. P. (2003). Nonlocal implicit
 1400 gradient-enhanced elasto-plasticity for the modelling of softening be-
 1401 haviour. *International Journal of Plasticity*, *19*(4), 403 - 433. doi:
 1402 10.1016/S0749-6419(01)00042-0
- 1403 Erickson, B. A., Jiang, J., Barall, M., Lapusta, N., Dunham, E. M., Harris, R., . . .
 1404 Wei, M. (2020). The Community Code Verification Exercise for Simulating
 1405 Sequences of Earthquakes and Aseismic Slip (SEAS). *Seismological Research*
 1406 *Letters*. doi: 10.1785/0220190248

- 1407 Faulkner, D., Mitchell, T., Jensen, E., & Cembrano, J. (2011). Scaling of fault dam-
1408 age zones with displacement and the implications for fault growth processes.
1409 *Journal of Geophysical Research: Solid Earth*, 116(B5).
- 1410 Fisher, R. A. (1937). The wave of advance of advantageous genes. *Annals of Eugen-*
1411 *ics*, 7(4), 355-369. doi: 10.1111/j.1469-1809.1937.tb02153.x
- 1412 Fritsch, F. N., Shafer, R., & Crowley, W. (1973). Algorithm 443: Solution of the
1413 transcendental equation $we^w = x$. *Communications of the ACM*, 16(2), 123–
1414 124.
- 1415 Gabriel, A.-A., Ampuero, J.-P., Dalguer, L. A., & Mai, P. M. (2012). The transition
1416 of dynamic rupture styles in elastic media under velocity-weakening friction.
1417 *Journal of Geophysical Research: Solid Earth*, 117(B9).
- 1418 Gabriel, A.-A., Li, D., Chiochetti, S., Tavelli, M., Peshkov, I., Romenski, E., &
1419 Dumbser, M. (2021). A unified first-order hyperbolic model for nonlinear
1420 dynamic rupture processes in diffuse fracture zones. *Philosophical Transactions*
1421 *of the Royal Society A*, 379(2196), 20200130.
- 1422 Garagash, D. I. (2012). Seismic and aseismic slip pulses driven by thermal pressur-
1423 ization of pore fluid. *Journal of Geophysical Research: Solid Earth*, 117(B4).
- 1424 Gilding, B., & Kersner, R. (2005). A fisher/kpp-type equation with density-
1425 dependent diffusion and convection: travelling-wave solutions. *Journal of*
1426 *Physics A: Mathematical and General*, 38, 3367. doi: 10.1088/0305-4470/38/
1427 15/009
- 1428 Godunov, S., & Romenskii, E. (1972). Nonstationary equations of nonlinear elastic-
1429 ity theory in eulerian coordinates. *Journal of Applied Mechanics and Technical*
1430 *Physics*, 13(6), 868–884.
- 1431 Godunov, S. K., & Romenskii, E. (2003). *Elements of continuum mechanics and*
1432 *conservation laws*. Springer Science & Business Media.
- 1433 Granier, T. (1985). Origin, damping, and pattern of development of faults in gran-
1434 ite. *Tectonics*, 4(7), 721-737. doi: 10.1029/TC004i007p00721
- 1435 Hermundstad, A., Daub, E., & Carlson, J. (2010). Energetics of strain localiza-
1436 tion in a model of seismic slip. *Journal of Geophysical Research: Solid Earth*,
1437 115(B6).
- 1438 Herrendörfer, R., Gerya, T., & Van Dinther, Y. (2018). An invariant rate-and
1439 state-dependent friction formulation for viscoelastoplastic earthquake cycle

- 1440 simulations. *Journal of Geophysical Research: Solid Earth*, *123*(6), 5018–5051.
- 1441 Hobbs, B. E., Mühlhaus, H.-B., & Ord, A. (1990). Instability, softening and local-
1442 ization of deformation. *Geological Society, London, Special Publications*, *54*(1),
1443 143–165. doi: 10.1144/GSL.SP.1990.054.01.15
- 1444 Ida, Y. (1972). Cohesive force across the tip of a longitudinal-shear crack and grif-
1445 fith’s specific surface energy. *Journal of Geophysical Research*, *77*(20), 3796–
1446 3805.
- 1447 Ikari, M. J., Marone, C., & Saffer, D. M. (2011). On the relation between fault
1448 strength and frictional stability. *Geology*, *39*(1), 83–86.
- 1449 Jiang, J., Erickson, B. A., Lambert, V. R., Ampuero, J.-P., Ando, R., Barbot, S. D.,
1450 ... van Dinther, Y. (2022). Community-driven code comparisons for three-
1451 dimensional dynamic modeling of sequences of earthquakes and aseismic slip.
1452 *Journal of Geophysical Research: Solid Earth*, *127*(3), e2021JB023519. doi:
1453 10.1029/2021JB023519
- 1454 Jiang, J., & Lapusta, N. (2016). Deeper penetration of large earthquakes on seismi-
1455 cally quiescent faults. *Science*, *352*(6291), 1293–1297.
- 1456 Jirásek, M. (1998). Nonlocal models for damage and fracture: comparison of
1457 approaches. *International Journal of Solids and Structures*, *35*(31-32), 4133–
1458 4145.
- 1459 Jirásek, M., & Rolshoven, S. (2009a). Localization properties of strain-softening gra-
1460 dient plasticity models. part ii: Theories with gradients of internal variables.
1461 *International Journal of Solids and Structures*, *46*(11-12), 2239–2254.
- 1462 Jirásek, M., & Rolshoven, S. (2009b). Localization properties of strain-softening gra-
1463 dient plasticity models. part i: Strain-gradient theories. *International Journal*
1464 *of Solids and Structures*, *46*(11-12), 2225–2238.
- 1465 Johansson, F. (2020). Computing the lambert w function in arbitrary-precision com-
1466 plex interval arithmetic. *Numerical Algorithms*, *83*(1), 221–242.
- 1467 John, T., Medvedev, S., Rüpke, L. H., Andersen, T. B., Podladchikov, Y. Y., &
1468 Austrheim, H. (2009). Generation of intermediate-depth earthquakes by self-
1469 localizing thermal runaway. *Nature Geoscience*, *2*, 137–140.
- 1470 Katz, Y., Weinberger, R., & Aydin, A. (2004). Geometry and kinematic evolution of
1471 riedel shear structures, capitol reef national park, utah. *Journal of Structural*
1472 *Geology*, *26*(3), 491-501. doi: 10.1016/j.jsg.2003.08.003

- 1473 Kaus, B. J. (2010). Factors that control the angle of shear bands in geodynamic nu-
1474 merical models of brittle deformation. *Tectonophysics*, *484*(1), 36 - 47. (Quan-
1475 titative modelling of geological processes) doi: 10.1016/j.tecto.2009.08.042
- 1476 Kiefer, B., Waffenschmidt, T., Sprave, L., & Menzel, A. (2018). A gradient-enhanced
1477 damage model coupled to plasticity multi-surface formulation and algorithmic
1478 concepts. *International Journal of Damage Mechanics*, *27*(2), 253-295. doi:
1479 10.1177/1056789516676306
- 1480 Kolmogorov, A. N., Petrovskii, I. G., & Piskunov, N. S. (1937). A study of the
1481 diffusion equation with increase in the amount of substance, and its applica-
1482 tion to a biological problem. In V. Tikhomirov (Ed.), *Selected works of a.n.*
1483 *kolmogorov 1 (1991)* (pp. 242–270). Kluwer.
- 1484 Kuhn, C., Schlüter, A., & Müller, R. (2015). On degradation functions in phase field
1485 fracture models. *Computational Materials Science*, *108*, 374–384.
- 1486 Kurzon, I., Lyakhovsky, V., & Ben-Zion, Y. (2019). Dynamic rupture and seismic
1487 radiation in a damage–breakage rheology model. *Pure and Applied Geophysics*,
1488 *176*(3), 1003–1020.
- 1489 Kurzon, I., Lyakhovsky, V., & Ben-Zion, Y. (2020). Earthquake source properties
1490 from analysis of dynamic ruptures and far-field seismic waves in a damage-
1491 breakage model. *Geophysical Journal International*, *224*(3), 1793-1810. doi:
1492 10.1093/gji/ggaa509
- 1493 Lapusta, N., & Liu, Y. (2009). Three-dimensional boundary integral modeling of
1494 spontaneous earthquake sequences and aseismic slip. *Journal of Geophysical*
1495 *Research: Solid Earth*, *114*(B9). doi: 10.1029/2008JB005934
- 1496 Lapusta, N., Rice, J. R., Ben-Zion, Y., & Zheng, G. (2000). Elastodynamic analysis
1497 for slow tectonic loading with spontaneous rupture episodes on faults with
1498 rate- and state-dependent friction. *Journal of Geophysical Research: Solid*
1499 *Earth*, *105*(B10), 23765-23789. doi: 10.1029/2000JB900250
- 1500 Le Pourhiet, L. (2013). Strain localization due to structural softening during pres-
1501 sure sensitive rate independent yielding. *Bulletin de la Société géologique de*
1502 *France*, *184*(4-5), 357–371.
- 1503 Li, D., & Liu, Y. (2017). Modeling slow-slip segmentation in cascadia subduction
1504 zone constrained by tremor locations and gravity anomalies. *Journal of Geo-*
1505 *physical Research: Solid Earth*, *122*(4), 3138–3157.

- 1506 Li, Y., & Wu, Y. (2008). Stability of travelling waves with noncritical speeds for
1507 double degenerate fisher-type equations. *Discrete and Continuous Dynamical*
1508 *Systems - Series B*, *10*, 149-170. doi: 10.3934/dcdsb.2008.10.149
- 1509 Linker, M., & Dieterich, J. H. (1992). Effects of variable normal stress on rock
1510 friction: Observations and constitutive equations. *Journal of Geophysical Re-*
1511 *search: Solid Earth*, *97*(B4), 4923-4940.
- 1512 Locatelli, M., Verlaguet, A., Agard, P., Federico, L., & Angiboust, S. (2018).
1513 Intermediate-depth brecciation along the subduction plate interface (monviso
1514 eclogite, w. alps). *Lithos*, *320*, 378-402.
- 1515 Locatelli, M., Verlaguet, A., Agard, P., Pettke, T., & Federico, L. (2019). Fluid
1516 pulses during stepwise brecciation at intermediate subduction depths (monviso
1517 eclogites, w. alps): First internally then externally sourced. *Geochemistry,*
1518 *geophysics, geosystems*, *20*(11), 5285-5318.
- 1519 Lyakhovsky, V., & Ben-Zion, Y. (2014a). A continuum damage-breakage faulting
1520 model and solid-granular transitions. *Pure and Applied Geophysics*, *171*(11),
1521 3099-3123.
- 1522 Lyakhovsky, V., & Ben-Zion, Y. (2014b). Damage-breakage rheology model and
1523 solid-granular transition near brittle instability. *Journal of the Mechanics and*
1524 *Physics of Solids*, *64*, 184-197.
- 1525 Lyakhovsky, V., Ben-Zion, Y., Ilchev, A., & Mendecki, A. (2016). Dynamic rup-
1526 ture in a damage-breakage rheology model. *Geophysical Journal International*,
1527 *206*(2), 1126-1143.
- 1528 Lyakhovsky, V., Hamiel, Y., & Ben-Zion, Y. (2011). A non-local visco-elastic dam-
1529 age model and dynamic fracturing. *Journal of the Mechanics and Physics of*
1530 *Solids*, *59*(9), 1752-1776.
- 1531 Mair, K., & Marone, C. (1999). Friction of simulated fault gouge for a wide range of
1532 velocities and normal stresses. *Journal of Geophysical Research: Solid Earth*,
1533 *104*(B12), 28899-28914. doi: 10.1029/1999JB900279
- 1534 Malaguti, L., & Ruggerini, S. (2010). Asymptotic speed of propagation for fisher-
1535 type degenerate reaction-diffusion-convection equations. *Advanced Nonlinear*
1536 *Studies*, *10*. doi: 10.1515/ans-2010-0306
- 1537 Marone, C. (1998). The effect of loading rate on static friction and the rate of fault
1538 healing during the earthquake cycle. *Nature*, *391*(6662), 69-72. doi: 10.1038/

1539

34157

1540

Marone, C., Hobbs, B. E., & Ord, A. (1992). Coulomb constitutive laws for friction: Contrasts in frictional behavior for distributed and localized shear. *pure and applied geophysics*, *139*(2), 195–214. doi: 10.1007/BF00876327

1541

1542

1543

McBeck, J. A., Aiken, J. M., Mathiesen, J., Ben-Zion, Y., & Renard, F. (2020). Deformation precursors to catastrophic failure in rocks. *Geophysical Research Letters*, *47*(24), e2020GL090255.

1544

1545

1546

McBeck, J. A., Ben-Zion, Y., & Renard, F. (2020). The mixology of precursory strain partitioning approaching brittle failure in rocks. *Geophysical Journal International*, *221*(3), 1856–1872.

1547

1548

1549

McBeck, J. A., Kobchenko, M., Hall, S. A., Tudisco, E., Cordonnier, B., Meakin, P., & Renard, F. (2018). Investigating the onset of strain localization within anisotropic shale using digital volume correlation of time-resolved x-ray microtomography images. *Journal of Geophysical Research: Solid Earth*, *123*(9), 7509–7528.

1550

1551

1552

1553

1554

McBeck, J. A., Zhu, W., & Renard, F. (2021). The competition between fracture nucleation, propagation, and coalescence in dry and water-saturated crystalline rock. *Solid Earth*, *12*(2), 375–387.

1555

1556

1557

Mühlhaus, H.-B., & Aifantis, E. (1991). A variational principle for gradient plasticity. *International Journal of Solids and Structures*, *28*(7), 845–857.

1558

1559

Mühlhaus, H.-B., & Vardoulakis, I. (1987). The thickness of shear bands in granular materials. *Geotechnique*, *37*(3), 271–283.

1560

1561

Needleman, A. (1988). Material rate dependence and mesh sensitivity in localization problems. *Computer Methods in Applied Mechanics and Engineering*, *67*(1), 69–85. doi: 10.1016/0045-7825(88)90069-2

1562

1563

1564

Niemeijer, A., & Spiers, C. (2007). A microphysical model for strong velocity weakening in phyllosilicate-bearing fault gouges. *Journal of Geophysical Research: Solid Earth*, *112*(B10).

1565

1566

1567

Okubo, K., Bhat, H. S., Rougier, E., Marty, S., Schubnel, A., Lei, Z., . . . Klinger, Y. (2019). Dynamics, radiation, and overall energy budget of earthquake rupture with coseismic off-fault damage. *Journal of Geophysical Research: Solid Earth*, *124*(11), 11771–11801. doi: 10.1029/2019JB017304

1568

1569

1570

1571

Okubo, P. G. (1989). Dynamic rupture modeling with laboratory-derived consti-

- 1572 tutive relations. *Journal of Geophysical Research: Solid Earth*, *94*(B9), 12321–
1573 12335.
- 1574 Passelègue, F. X., Almakari, M., Dublanchet, P., Barras, F., Fortin, J., & Violay,
1575 M. (2020). Initial effective stress controls the nature of earthquakes. *Nature*
1576 *communications*, *11*(1), 1–8.
- 1577 Passelègue, F. X., Spagnuolo, E., Violay, M., Nielsen, S., Di Toro, G., & Schubnel,
1578 A. (2016). Frictional evolution, acoustic emissions activity, and off-fault dam-
1579 age in simulated faults sheared at seismic slip rates. *Journal of Geophysical*
1580 *Research: Solid Earth*, *121*(10), 7490–7513.
- 1581 Peerlings, R. H. J., de Borst, R., Brekelmans, W. A. M., & De Vree, J. H. P. (1996).
1582 Gradient enhanced damage for quasi-brittle materials. *International Journal*
1583 *for Numerical Methods in Engineering*, *39*(19), 3391–3403. doi: 10.1002/(SICI)
1584 1097-0207(19961015)39:19<3391::AID-NME7>3.0.CO;2-D
- 1585 Peirce, D., Asaro, R. J., & Needleman, A. (1983). Material rate dependence and
1586 localized deformation in crystalline solids. *Acta metallurgica*, *31*(12), 1951–
1587 1976.
- 1588 Perez-Silva, A., Li, D., Gabriel, A.-A., & Kaneko, Y. (2021). 3d modeling of long-
1589 term slow slip events along the flat-slab segment in the guerrero seismic gap,
1590 mexico. *Geophysical Research Letters*, *48*(13), e2021GL092968.
- 1591 Perrin, C., Manighetti, I., & Gaudemer, Y. (2016). Off-fault tip splay networks:
1592 A genetic and generic property of faults indicative of their long-term propa-
1593 gation. *Comptes Rendus Geoscience*, *348*(1), 52–60. (From past to current
1594 tectonics) doi: 10.1016/j.crte.2015.05.002
- 1595 Petrovskii, S. V., & Li, B.-L. (2003). An exactly solvable model of population dy-
1596 namics with density-dependent migrations and the allee effect. *Mathematical*
1597 *Biosciences*, *186*(1), 79 - 91. doi: 10.1016/S0025-5564(03)00098-1
- 1598 Petrovskii, S. V., & Li, B.-L. (2006). *Exactly solvable models of biological invasion*.
1599 Chapman and Hall/CRC.
- 1600 Pijaudier-Cabot, G., & Bažant, Z. P. (1987). Nonlocal damage theory. *Journal of*
1601 *engineering mechanics*, *113*(10), 1512–1533.
- 1602 Platt, J. D., Rudnicki, J. W., & Rice, J. R. (2014). Stability and localization of
1603 rapid shear in fluid-saturated fault gouge: 2. localized zone width and strength
1604 evolution. *Journal of Geophysical Research: Solid Earth*, *119*(5), 4334–4359.

- 1605 Poulet, T., Veveakis, E., Regenauer-Lieb, K., & Yuen, D. A. (2014). Thermo-poro-
1606 mechanics of chemically active creeping faults: 3. the role of serpentinite in
1607 episodic tremor and slip sequences, and transition to chaos. *Journal of Geo-*
1608 *physical Research: Solid Earth*, *119*(6), 4606–4625.
- 1609 Pozzi, G., De Paola, N., Nielsen, S. B., Holdsworth, R. E., & Bowen, L. (2018). A
1610 new interpretation for the nature and significance of mirror-like surfaces in
1611 experimental carbonate-hosted seismic faults. *Geology*, *46*(7), 583-586. doi:
1612 10.1130/G40197.1
- 1613 Pozzi, G., De Paola, N., Nielsen, S. B., Holdsworth, R. E., Tesei, T., Thieme, M.,
1614 & Demouchy, S. (2021). Coseismic fault lubrication by viscous deformation.
1615 *Nature Geoscience*, *14*(6), 437–442.
- 1616 Pozzi, G., Paola, N. D., Holdsworth, R. E., Bowen, L., Nielsen, S. B., & Dempsey,
1617 E. D. (2019). Coseismic ultramylonites: An investigation of nanoscale viscous
1618 flow and fault weakening during seismic slip. *Earth and Planetary Science*
1619 *Letters*, *516*, 164 - 175. doi: 10.1016/j.epsl.2019.03.042
- 1620 Preuss, S., Ampuero, J. P., Gerya, T., & van Dinther, Y. (2020). Characteristics of
1621 earthquake ruptures and dynamic off-fault deformation on propagating faults.
1622 *Solid Earth*, *11*(4), 1333–1360.
- 1623 Preuss, S., Herrendrfer, R., Gerya, T., Ampuero, J.-P., & van Dinther, Y. (2019).
1624 Seismic and aseismic fault growth lead to different fault orientations.
1625 *Journal of Geophysical Research: Solid Earth*, *124*(8), 8867-8889. doi:
1626 10.1029/2019JB017324
- 1627 Putelat, T., Dawes, J., & Champneys, A. (2017). A phase-plane analysis of local-
1628 ized frictional waves. *Proceedings of the Royal Society A: Mathematical, Physi-*
1629 *cal and Engineering Sciences*, *473*(2203), 20160606.
- 1630 Rabinowicz, E. (1958). The intrinsic variables affecting the stick-slip process. *Pro-*
1631 *ceedings of the Physical Society (1958-1967)*, *71*(4), 668.
- 1632 Rattez, H., Stefanou, I., & Sulem, J. (2018). The importance of thermo-hydro-
1633 mechanical couplings and microstructure to strain localization in 3d continua
1634 with application to seismic faults. part i: Theory and linear stability analysis.
1635 *Journal of the Mechanics and Physics of Solids*, *115*, 54 - 76.
- 1636 Rattez, H., Stefanou, I., Sulem, J., Veveakis, M., & Poulet, T. (2018). The im-
1637 portance of thermo-hydro-mechanical couplings and microstructure to strain

- 1638 localization in 3d continua with application to seismic faults. part ii: Numerical
 1639 implementation and post-bifurcation analysis. *Journal of the Mechanics
 1640 and Physics of Solids*, 115, 1–29.
- 1641 Resnyansky, A., Romensky, E., & Bourne, N. (2003). Constitutive modeling of frac-
 1642 ture waves. *Journal of applied physics*, 93(3), 1537–1545.
- 1643 Rice, J. R. (1993). Spatio-temporal complexity of slip on a fault. *Journal of Geo-
 1644 physical Research: Solid Earth*, 98(B6), 9885–9907. doi: 10.1029/93JB00191
- 1645 Rice, J. R. (2006). Heating and weakening of faults during earthquake slip. *Journal
 1646 of Geophysical Research: Solid Earth*, 111(B5).
- 1647 Rice, J. R., & Ben-Zion, Y. (1996). Slip complexity in earthquake fault models. *Pro-
 1648 ceedings of the National Academy of Sciences*, 93(9), 3811–3818. doi: 10.1073/
 1649 pnas.93.9.3811
- 1650 Rice, J. R., Rudnicki, J. W., & Platt, J. D. (2014). Stability and localization of
 1651 rapid shear in fluid-saturated fault gouge: 1. linearized stability analysis. *Jour-
 1652 nal of Geophysical Research: Solid Earth*, 119(5), 4311–4333.
- 1653 Ritter, M. C., Rosenau, M., & Oncken, O. (2018). Growing faults in the lab: In-
 1654 sights into the scale dependence of the fault zone evolution process. *Tectonics*,
 1655 37(1), 140–153. doi: 10.1002/2017TC004787
- 1656 Ritter, M. C., Santimano, T., Rosenau, M., Leever, K., & Oncken, O. (2018). Sand-
 1657 box rheometry: Co-evolution of stress and strain in riedel and critical wedgeex-
 1658 periments. *Tectonophysics*, 722, 400–409. doi: 10.1016/j.tecto.2017.11.018
- 1659 Romenski, E., Reshetova, G., Peshkov, I., & Dumbser, M. (2020). Modeling wave-
 1660 fields in saturated elastic porous media based on thermodynamically compat-
 1661 ible system theory for two-phase solid-fluid mixtures. *Computers & Fluids*,
 1662 206, 104587. doi: 10.1016/j.compfluid.2020.104587
- 1663 Romenskii, E. (2007). Deformation model for brittle materials and the structure of
 1664 failure waves. *Journal of applied mechanics and technical physics*, 48(3), 437–
 1665 444.
- 1666 Ross, Z. E., Idini, B., Jia, Z., Stephenson, O. L., Zhong, M., Wang, X., ... others
 1667 (2019). Hierarchical interlocked orthogonal faulting in the 2019 ridgecrest
 1668 earthquake sequence. *Science*, 366(6463), 346–351.
- 1669 Roubek, T. (2014). A note about the rate-and-state-dependent friction model in a
 1670 thermodynamic framework of the Biot-type equation. *Geophysical Journal In-*

- 1671 *ternational*, 199(1), 286-295. doi: 10.1093/gji/ggu248
- 1672 Rozel, A., Ricard, Y., & Bercovici, D. (2011). A thermodynamically self-consistent
1673 damage equation for grain size evolution during dynamic recrystallization.
1674 *Geophysical Journal International*, 184(2), 719–728.
- 1675 Rubin, A. M., & Ampuero, J.-P. (2005). Earthquake nucleation on (aging) rate and
1676 state faults. *Journal of Geophysical Research: Solid Earth*, 110(B11).
- 1677 Ruina, A. L. (1980). *Friction laws and instabilities: A quasistatic analysis of some*
1678 *dry friction behaviour* (Unpublished doctoral dissertation). Ph.D. Thesis, Divi-
1679 sion of Engineering, Brown University, Providence RI.
- 1680 Ruina, A. L. (1983). Slip instability and state variable friction laws. *Jour-*
1681 *nal of Geophysical Research: Solid Earth*, 88(B12), 10359-10370. doi:
1682 10.1029/JB088iB12p10359
- 1683 Sánchez-Garduño, F., & Maini, P. (1994). An approximation to a sharp type so-
1684 lution of a density-dependent reaction-diffusion equation. *Applied Mathematics*
1685 *Letters*, 7(1), 47 - 51. doi: 10.1016/0893-9659(94)90051-5
- 1686 Sánchez-Garduño, F., & Maini, P. (1995). Traveling wave phenomena in some degen-
1687 erate reaction-diffusion equations. *Journal of Differential Equations*, 117(2),
1688 281 - 319. doi: 10.1006/jdeq.1995.1055
- 1689 Sánchez-Garduño, F., & Maini, P. (1997). Travelling wave phenomena in non-linear
1690 diffusion degenerate nagumo equations. *Journal of Mathematical Biology*,
1691 35(6), 713–728. doi: 10.1007/s002850050073
- 1692 Sathiakumar, S., Barbot, S., & Hubbard, J. (2020). Earthquake cycles in fault-bend
1693 folds. *Journal of Geophysical Research: Solid Earth*, 125(8), e2019JB018557.
- 1694 Savage, H. M., & Brodsky, E. E. (2011). Collateral damage: Evolution with dis-
1695 placement of fracture distribution and secondary fault strands in fault damage
1696 zones. *Journal of Geophysical Research: Solid Earth*, 116(B3).
- 1697 Scholz, C., Molnar, P., & Johnson, T. (1972). Detailed studies of frictional sliding of
1698 granite and implications for the earthquake mechanism. *Journal of geophysical*
1699 *research*, 77(32), 6392–6406.
- 1700 Schreyer, H. L., & Chen, Z. (1986). One-Dimensional Softening With Localization.
1701 *Journal of Applied Mechanics*, 53(4), 791-797. doi: 10.1115/1.3171860
- 1702 Segall, P., & Rice, J. R. (1995). Dilatancy, compaction, and slip instability of a
1703 fluid-infiltrated fault. *Journal of Geophysical Research: Solid Earth*, 100(B11),

- 1704 22155-22171. doi: 10.1029/95JB02403
- 1705 Sherratt, J., & Marchant, B. (1996). Nonsharp travelling wave fronts in the fisher
1706 equation with degenerate nonlinear diffusion. *Applied Mathematics Letters*,
1707 9(5), 33 - 38. doi: 10.1016/0893-9659(96)00069-9
- 1708 Simo, J., & Taylor, R. (1985). Consistent tangent operators for rate-independent
1709 elastoplasticity. *Computer Methods in Applied Mechanics and Engineering*,
1710 48(1), 101 - 118. doi: 10.1016/0045-7825(85)90070-2
- 1711 Sleep, N. H. (1995). Ductile creep, compaction, and rate and state dependent fric-
1712 tion within major fault zones. *Journal of Geophysical Research: Solid Earth*,
1713 100(B7), 13065–13080.
- 1714 Sleep, N. H. (1997). Application of a unified rate and state friction theory to the
1715 mechanics of fault zones with strain localization. *Journal of Geophysical Re-*
1716 *search: Solid Earth*, 102(B2), 2875-2895. doi: 10.1029/96JB03410
- 1717 Stathas, A., & Stefanou, I. (2022). The role of viscous regularization in
1718 dynamical problems, strain localization and mesh dependency. *Com-*
1719 *puter Methods in Applied Mechanics and Engineering*, 388, 114185. doi:
1720 10.1016/j.cma.2021.114185
- 1721 Tchalenko, J. (1970). Similarities between shear zones of different magnitudes. *Geo-*
1722 *logical Society of America Bulletin*, 81(6), 1625–1640.
- 1723 Templeton, E. L., & Rice, J. R. (2008). Off-fault plasticity and earthquake rup-
1724 ture dynamics: 1. dry materials or neglect of fluid pressure changes. *Journal of*
1725 *Geophysical Research: Solid Earth*, 113(B9).
- 1726 Thielmann, M. (2018). Grain size assisted thermal runaway as a nucleation mech-
1727 anism for continental mantle earthquakes: Impact of complex rheologies.
1728 *Tectonophysics*, 746, 611–623. doi: 10.1016/j.tecto.2017.08.038
- 1729 Thielmann, M., Rozel, A., Kaus, B., & Ricard, Y. (2015). Intermediate-depth earth-
1730 quake generation and shear zone formation caused by grain size reduction and
1731 shear heating. *Geology*, 43(9), 791. doi: 10.1130/G36864.1
- 1732 Thomas, M. Y., Lapusta, N., Noda, H., & Avouac, J.-P. (2014). Quasi-dynamic
1733 versus fully dynamic simulations of earthquakes and aseismic slip with and
1734 without enhanced coseismic weakening. *Journal of Geophysical Research: Solid*
1735 *Earth*, 119(3), 1986-2004. doi: 10.1002/2013JB010615
- 1736 Triantafyllidis, N., & Aifantis, E. C. (1986). A gradient approach to localization of

- 1737 deformation. i. hyperelastic materials. *Journal of Elasticity*, 16(3), 225–237.
- 1738 Ulm, F.-J., Coussy, O., & Bažant, Z. P. (1999). The “chunnel” fire. i: Chemoplas-
1739 tic softening in rapidly heated concrete. *Journal of engineering mechanics*,
1740 125(3), 272–282.
- 1741 Van den Ende, M., Chen, J., Ampuero, J.-P., & Niemeijer, A. (2018). A comparison
1742 between rate-and-state friction and microphysical models, based on numerical
1743 simulations of fault slip. *Tectonophysics*, 733, 273–295.
- 1744 Vardoulakis, I. (1980). Shear band inclination and shear modulus of sand in biax-
1745 ial tests. *International journal for numerical and analytical methods in geome-
1746 chanics*, 4(2), 103–119.
- 1747 Vermeer, P. (1990). The orientation of shear bands in biaxial tests. *Geotechnique*,
1748 40(2), 223–236.
- 1749 Veveakis, E., Poulet, T., & Alevizos, S. (2014). Thermo-poro-mechanics of chemi-
1750 cally active creeping faults: 2. transient considerations. *Journal of Geophysical
1751 Research: Solid Earth*, 119(6), 4583–4605.
- 1752 Viesca, R. C. (2016a). Self-similar slip instability on interfaces with rate-and state-
1753 dependent friction. *Proceedings of the Royal Society A: Mathematical, Physical
1754 and Engineering Sciences*, 472(2192), 20160254.
- 1755 Viesca, R. C. (2016b). Stable and unstable development of an interfacial sliding in-
1756 stability. *Physical Review E*, 93(6), 060202.
- 1757 Viesca, R. C. (2020). On the existence of a nucleation length for dynamic shear rup-
1758 ture. *arXiv preprint arXiv:2008.11854*.
- 1759 Wolfram Research, Inc. (2017). *Mathematica, Version 11.1.1.0*. (Champaign, IL,
1760 2017)
- 1761 Wu, F., & Freund, L. (1984). Deformation trapping due to thermoplastic instability
1762 in one-dimensional wave propagation. *Journal of the Mechanics and Physics of
1763 Solids*, 32(2), 119–132. doi: 10.1016/0022-5096(84)90014-0
- 1764 Wu, J.-Y. (2017). A unified phase-field theory for the mechanics of damage and
1765 quasi-brittle failure. *Journal of the Mechanics and Physics of Solids*, 103, 72–
1766 99.
- 1767 Wu, Y., Xing, X., & Ye, Q. (2006). Stability of travelling waves with algebraic
1768 decay for n-degree fisher-type equations. *Discrete and Continuous Dynamical
1769 Systems*, 16. doi: 10.3934/dcds.2006.16.47