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# Optimizing TRISHNA TIR channels configuration for improved land surface temperature and emissivity measurements

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# Abstract

In preparation of the Thermal infraRed Imaging Satellite for High-resolution Natural resource Assessment (TRISHNA) mission, we conducted a thorough analysis of sensitivity for the Temperature-Emissivity Separation (TES) method to the position of the four TRISHNA spectral channels, notably to find an optimal spectral configuration. To that purpose, we designed a fast-computing end-to-end simulator including several components, which we implemented to simulate both pixel-size TRISHNA measurements and land surface temperature (LST) retrievals. Firstly, simulations were conducted over a wide range of realistic scenarii, notably by including vegetation canopy-scale cavity effect. Secondly, the experimental design included the features of second generation Mercury-Cadmium-Telluride (MCT) cooled detectors with lower instrumental noises and finer channels. Thirdly, as opposed to previous studies that used predefined spectral configurations to determine the most suited one, we conducted an optimization of the spectral configuration by crossing, on a pair basis, several positions of the four TIR channels over a range of wavelengths. Fourthly, we quantified the TES sensitivity to atmospheric perturbations, by comparing LST retrievals with and without atmospheric noise.

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We observed an overall moderate sensitivity of TES LST retrievals to the spectral channel positions, with a maximum RMSE variation of 0.31 K within the atmospheric spectral windows. Furthermore, the TES method was sensitive to three main parameters, namely the instrumental noise, the atmospheric downwelling irradiance, and the transmittance due to ozone and water vapor, with RMSEs larger than 1 K for specific channel locations. Moreover, by considering possible superimposition of two channels, we noted that the TES method could achieve similar performance by considering three or four channels. Eventually, our study enabled us to recommend a new spectral configuration for the TRISHNA TIR instrument, that is more robust to atmospheric perturbations and to uncertainties on channel positions and bandwidths.

*Keywords:* Thermal infrared remote sensing, Satellite mission design, Spectral channels positioning, Temperature/emissivity separation, Vegetation canopy - scaled cavity effects, Mercury - Cadmium - Telluride cooled detectors, Atmospheric corrections, Sensitivity analysis

# 1 1. Introduction

The increasing need for efficient satellite retrieval of land surface temperature (LST) has become unanimous in the global scientific community (Murphy, 2006; Malenovský et al., 2012; Lagouarde et al., 2013; Weiss et al., 2020). Indeed, LST is a key parameter in environmental physics, since it drives numerous 5 land surface processes such as radiation budgets (Hulley and Hook, 2010; Mira et al., 2016; Ogawa et al., 2002, 2003; Yan et al., 2020), water depletion from soils through evapotranspiration (Bigeard et al., 2019; Carlson and Petropoulos, 2019; French et al., 2005; Galleguillos et al., 2011, 2010; Gómez et al., 2005; Jacob et al., 2002; Montes et al., 2014; Montes and Jacob, 2017; Pardo et al., 10 2014; Vinukollu et al., 2011), photosynthesis and soil respiration (Bayat et al., 11 2018; Inoue et al., 2004; Olioso et al., 2005), and it can be linked to paramount 12 environmental processes such as pollutant degradation (Louchart and Voltz, 13 2007) and pathogen dissemination (Courault et al., 2009; Schröder et al., 2006; 14

<sup>15</sup> Sobrino et al., 2016).

The consensus on the need for satellite-based LST retrievals is notably illus-16 trated by a recent soar of related satellite projects, with the launch in 2018 of the 17 NASA ECOsystem Spaceborne Thermal Radiometer Experiment onboard the 18 International Space Station (ECOSTRESS, Hulley et al., 2017), the ESA future 19 High Spatio-Temporal Resolution Land Surface Temperature Mission (HSTR 20 LSTM, Koetz et al., 2018), and the CNES (French spatial agency) / ISRO (In-21 dian Space Research Organisation) cooperation project on the Thermal infraRed 22 Imaging Satellite for High-resolution Natural resource Assessment mission (TR-23 ISHNA, Lagouarde et al., 2018). 24

The TRISHNA satellite will embark both a visible/short wave infrared and a 25 thermal infrared (TIR) sensor. With that payload, the TRISHNA mission will 26 focus on six monitoring scientific objectives, addressing 1/ energy and water 27 budgets of the continental biosphere in relation to ecosystem stress and water 28 use, 2/ coastal and inland waters, 3/ climatology and fluxes within urban areas 29 in relation to urban heating environments, 4/ geological phenomena such as 30 volcanoes or earthquakes in relation to solid Earth, 5/ cryospheric processes 31 such as snowmelt and related runoff, and 6/ atmospheric characterization (e.g., 32 water vapor). The two first objectives are the design drivers, with an emphasis 33 on ecosystem stress and water use monitoring. In a climate change context, the 34 latter is a critical aspect of the mission, notably for inter-tropical regions such 35 as India, for sub-humid southern France regions, or for semi-arid to arid regions 36 of the Mediterranean basin. 37

In order to efficiently reach the objectives of the TRISHNA mission, the 38 platform will be set to an approximately 760 km orbit with a 1,000 km swath 39  $\pm 34^{\circ}$  FOV), which allows a  $\sim 60$  m nadir spatial resolution and a 3-days revisit 40 time for global coverage (Lagouarde et al., 2019). These mission parameters 41 are critical since they will allow precise monitoring of, for instance, water stress 42 at crop field scale or urban heat islands. In order to guarantee such a precise 43 monitoring, another important mission parameter is the allocation of 4 channels 44 between 8 and 12  $\mu$ m for the thermal infrared sensor. These four channels are 45

<sup>46</sup> implemented by using second generation Mercury-Cadmium-Telluride (MCT)
<sup>47</sup> cooled detectors with better instrumental performance (lower noises and finer
<sup>48</sup> channels) than first generation microbolometers used on board previous satellite
<sup>49</sup> missions. In this configuration, high quality TRISHNA level 2 products are
<sup>50</sup> expected to be available to the community every 12 h.

In order to ensure precise measurements, suitable methods shall be used to retrieve LST from out-of-sensor TIR radiances. The retrieval issue lies in the fact that radiances in the TIR domain (7.5-13  $\mu$ m) depend on both temperature and emissivity of the observed surface. Hence, even with proper atmospheric corrections, the problem of retrieving these surface variables is ill-posed because of a larger number of unknowns ( $N_B$  emissivities + 1 temperature), than measurements ( $N_B$  radiances), with  $N_B$  the number of TIR spectral channels.

Among the existing methods for LST retrieval (see Dash et al., 2002; Jacob 58 et al., 2008; Li et al., 2013, for reviews on the matter), the best fitted for the 59 TRISHNA objectives is the Temperature-Emissivity Separation method (TES, 60 Gillespie et al., 1998), notably because it does not require ancillary information 61 on the observed scene and can directly produce LST estimates for each satel-62 lite overpass, provided that prior atmospheric corrections are conducted. This 63 method was developed specifically for TIR multispectral precursor sensors such 64 as the spaceborne Advanced Spaceborne Thermal Emission and Reflection Ra-65 diometer (ASTER) sensor or the airborne TIMS (Thermal Infrared Multispec-66 tral Scanner) sensor, and later on analysed or modified in order to stabilize and 67 enhance its performance (Mira et al., 2009; Coll et al., 2007; French et al., 2008; 68 Gillespie et al., 1999; Göttsche and Hulley, 2012; Grigsby et al., 2015; Gustafson 69 et al., 2006; Hulley et al., 2008, 2012; Jacob et al., 2004, 2017; Schmugge et al., 70 1998, 2002; Zheng et al., 2019). 71

The second alternative for LST estimation in the context of the TRISHNA mission is the Split Window (SW) method, notably because it permits to secure the operational delivery of LST retrievals in case of failure of at most two channels, and is more easily applicable than the TES method as it does not require prior atmospheric correction. It can even ensure better LST results for

grey-body observations, *i.e.*, scenes with low spectral contrast (see Li et al., 77 2013, for an overview of the existing SW methods). Eventually, the SW method 78 could be operationally used with retrievals of land surface emissivity and at-79 mospheric water vapor content, which should be indirectly obtained from the 80 TRISHNA VIS/NIR instrument via a NDVI-based method (Neinavaz et al., 81 2020; Vanhellemont, 2020; Jiménez-Muñoz and Sobrino, 2006; Sobrino et al., 82 2004; Olioso, 1995) and via a so-called 'water vapor channel' at 0.910  $\mu$ m, re-83 spectively. 84

As an important step of the TRISHNA mission preparation for high quality 85 observations, the TES and SW methods require prior optimization of the in-86 strument spectral configuration based on performance analyses. This consists in 87 establishing the locations and bandwidths of the spectral channels, namely two 88 channels between 10 and 12  $\mu$ m for SW, combined with two additional channels 89 between 8 and 9.5  $\mu$ m for TES, so that: 1/ atmospheric perturbations are en-90 hanced locally (for SW) or minimized overall (for TES), 2/ emissivity contrast 91 is spectrally minimized (for SW) or maximised (for TES), and 3/ instrumental 92 noises are reduced (for both TES and SW). 93

Pioneer work in that regard was conducted by Caselles et al. (1998), who 94 tested the performance of a SW method for several 2-channels configurations in 95 the context of the PRISM satellite development. Thanks to enhanced numeri-96 cal capabilities, further studies evaluated the performances of the TES and SW 97 methods for a larger number of spectral configurations (Sobrino et al., 2010; 98 Sobrino and Jiménez-Muñoz, 2014). In order to propose a spectral configu-99 ration for the MIcro Satellite for Thermal InfraRed GRound surface Imaging 100 (MISTIGRI) mission (Lagouarde et al., 2013), Jacob et al. (2021) examined 101 the performance of the TES method for six spectral configurations based on 102 microbolometer detectors. For this study, they used a representative emissivity 103 spectra dataset proposed by Jacob et al. (2017) to account for the cavity effect, 104 since the latter results in an upper (respectively lower) shift in emissivities (re-105 spectively radiometric temperature) as compared to the original TES settings. 106 Using the same emissivity dataset, and in order to propose a spectral config-107

<sup>108</sup> uration dedicated to the SW method for the TRISHNA mission, Vidal et al.
<sup>109</sup> (2021) developed a bi-dimensional optimization approach that simultaneously
<sup>110</sup> moves spectral channels rather than using predefined configurations.

In order to define an optimal spectral configuration, the current study aims 111 at conducting a sensitivity analysis of the TES method to the position of the 112 four TRISHNA spectral channels. It capitalizes on the contributions discussed 113 above, since it addresses the design of a TRISHNA-TES dedicated configura-114 tion 1/ by including the features of second generation MCT cooled detectors, 115 2/ by accounting for cavity effects in vegetation canopy, 3/ by conducting a sen-116 sitivity analysis that simultaneously moves spectral channels rather than using 117 predefined configurations, and 4/ by using an end-to-end simulator that includes 118 sensor measurements, atmospheric corrections and the TES retrieval procedure. 119 We first present a description of the TES method in § 2, since it is the main 120 object of our study. Then we describe the end-to-end simulator we specifically 121 developed for this work in  $\S$  3, as well as its implementation in the context of 122 TRISHNA studies in § 4. The experimental plan of our sensitivity analysis is 123 described in § 5, and § 6 details the various results obtained, which are then 124 discussed in  $\S$  7. We conclude in  $\S$  8 with limitations and opportunities of the 125 work to date. 126

# <sup>127</sup> 2. The Temperature/Emissivity Separation (TES) method

# 128 2.1. TES principle and overview

The TES method (Gillespie et al., 1998) was initially developed by the 129 ASTER Temperature Emissivity Working Group (hereafter TEWG) in order 130 to efficiently tackle the issue of surface temperature/emissivity separation ( $\S$  1). 131 After a thorough study of existing temperature and/or emissivity retrieval meth-132 ods, TES was proposed by combining key features of previous methods, espe-133 cially the Normalized Emissivity Method (NEM, Realmuto, 1990) and Minimum-134 Maximum Difference method (MMD, Matsunaga, 1994). More particularly, in 135 order to render the temperature/emissivity separation issue deterministic, the 136

TES method is based on the idea that over  $N_B \ge 3$  channels in the TIR domain, 137 the emissivity spectrum of a natural surface contains at least one value close 138 to unity. From this hypothesis, a first estimation of the surface temperature 139 and iterative sky irradiance correction is conducted by the NEM module. From 140 this correction a first emissivity spectrum is derived and ratioed to its mean 141 value, resulting in an relative emissivity spectrum that has the shape but not 142 the amplitude of the observed spectrum. This amplitude is then found within 143 the MMD module by relating the minimum emissivity to the spectral contrast, 144 which notably implies that the used spectral channels allow an efficient measure-145 ment of the latter. Eventually, a more correct final temperature is calculated 146 from the maximum emissivity value. 147

The ASTER TEWG and following studies have shown that TES can re-148 trieve surface emissivities within an 0.01 error margin (Coll et al., 2003) and 149 temperature within a 1 K error margin (Schmugge et al., 2003), assuming well 150 calibrated and accurate radiometric measurements. This error margin is satis-151 factory for the TRISHNA mission, which instrumental specification on absolute 152 calibration accuracy is about 0.7 K, for a required precision on LST of  $\sim 1.5$  K 153 (Lagouarde et al., 2019). Note that recently, Zheng et al. (2019) have suc-154 cessfully enhanced the TES method by proposing a SW/TES hybrid method 155 that reduces the TES precision to 0.87 K. However, their method is based on 156 a dry-atmosphere hypothesis, which is not relevant for our study, since most of 157 the regions of interest for the TRISHNA mission, such as inter-tropical regions, 158 sustain a humid climate. The TES processing flow is hereafter detailed. 159

# 160 2.1.1. NEM module

Using as input surface-leaving radiance  $L_i^{sur\uparrow}$ , which necessitates to apply atmospheric corrections to the out-of-sensor (OS) radiance before entering the TES method, the role of the NEM module is threefold:

1. to retrieve a first estimate for the surface temperature;

<sup>165</sup> 2. to estimate the shape for the emissivity spectrum;

a. to iteratively correct for the atmospheric downwelling irradiance impact
 on surface radiance.

It is first supposed that the anisotropic behavior of the surface is neglected, and that the emissivities are all equal to a given maximum emissivity close to unity  $(\varepsilon_i)_{i\in[1;N_B]} = \varepsilon_{max} = 0.99$ , typical of vegetated surfaces, snow and water. The ground emitted radiance  $L_i^{em}$  can then be estimated from the known surface-leaving radiance  $L_i^{sur\uparrow}$  and the atmospheric downwelling irradiance  $E_i^{atm\downarrow}$  via:

$$L_i^{em} = L_i^{sur\uparrow} - (1 - \varepsilon_{max}) \frac{E_i^{atm\downarrow}}{\pi}$$
(1)

The NEM temperature  $T_{NEM}$  is then calculated as the maximum temperature obtained from each ground emitted radiance  $L_i^{em}$  via inverse Planck's law:

$$T_i = B^{-1} \left( \frac{L_i^{em}}{\varepsilon_{max}}, \lambda_i \right)$$
(2)

$$T_{NEM} = \max(T_i) \tag{3}$$

$$B^{-1}(L_i, \lambda_i) = \frac{c_2}{\lambda_i \ln\left(\frac{c_1}{\pi \lambda_i^5 L_i} + 1\right)}$$
(4)

where  $c_1$  and  $c_2$  are the radiative constants respectively equal to  $3.74151 \times 10^{-16} \text{ W.m}^{-2}$  and  $0.0143879 \text{ m.K.} T_{NEM}$  is indeed most likely to be the best estimate of the actual surface temperature in presence of uncorrected effects of the atmospheric irradiance. Afterwards, the NEM emissivity spectrum is calculated as the ratio between the calculated ground emitted radiance  $L_i^{em}$ and that of a blackbody at the NEM temperature:

$$\varepsilon_{NEM,i} = \frac{L_i^{em}}{B(T_{NEM}, \lambda_i)} \tag{5}$$

where  $B(T_{NEM}, \lambda_i)$  is the Planck's law associated to the surface temperature  $T_{NEM}$  at the channel effective wavelength  $\lambda_i$ :

$$B(T,\lambda) = \frac{c_1}{\lambda^5 \pi \left[\exp\left(\frac{c_2}{T\lambda}\right) - 1\right]} \tag{6}$$

The obtained emissivity spectrum is then used to re-estimate  $L_i^{em}$  with Equa-185 tion 1, and the process from Equation 1 to 5 is repeated until convergence or 186 until the number of iteration reaches an arbitrary threshold  $N_{it}^{max} = 12$ . The 187 convergence condition is satisfied if the change in  $L_i^{em}$  between steps is lower 188 than a given threshold  $t_2$ , which is currently set to its ASTER value of 0.05 189  $W.m^{-2}.sr^{-1}.\mu m^{-1}$  and shall be refined for TRISHNA to its Noise Equivalent 190 Delta Radiance (Ne $\Delta L$ ) value. Divergence is also tested by calculating the sec-191 ond derivative of  $L_i^{em}$  versus the number of iterations  $N_{it}$ , *i.e.*, its variation 192 rate. If this rate exceeds a threshold  $t_1$ ,  $|\Delta^2 L_i^{em}/\Delta N_{it}^2| > t_1$ , TES is halted 193 and the last  $T_{NEM}$  and  $\varepsilon_{NEM,i}$  values are returned along with an error flag. 194 The threshold  $t_1$  is also set to its ASTER original value, which is the same 195 as  $t_2$ , and should be refined during the future TRISHNA level-2 study. Be-196 cause the TRISHNA mission will mainly focus on ecosystems, and therefore 197 surface emissivities greater than 0.5, TES can also be aborted in the case where 198 one of the NEM calculated emissivities exceeds reasonable limits, thus set to 199  $0.5 < \varepsilon_{NEM,i} < 1.0$ . Once convergence is reached, further  $\varepsilon_{max}$  refinement is 200 conducted, notably by evaluating the variance of the  $\varepsilon_{NEM,i}$  values obtained. 201 If the variance is larger or lower than a given threshold, the pixel is consid-202 ered to be composed mainly of rocks and soils, or of near graybody surfaces, 203 respectively. The  $\varepsilon_{max}$  refinement process used in this study is the same as de-204 scribed in Gillespie et al. (1998). Note that the aforementioned tests and flags 205 are paramount in terms of mission requirements, since they will allow to define 206 confidence levels on the results of the method for each pixel. 207

# 208 2.1.2. Ratio module

In the ratio module the relative emissivities  $\beta_i$  are calculated as the ratio between the NEM-obtained emissivities and their average value on the  $N_B$ channels:

$$\beta_i = \frac{\varepsilon_{NEM,i}}{\bar{\varepsilon}} = N_B \cdot \varepsilon_{NEM,i} \left( \sum_i \varepsilon_{NEM,i} \right)^{-1} \tag{7}$$

# 212 2.1.3. MMD module

The aim of the MMD module is to rescale the relative emissivity spectrum  $(\beta_i)_{i \in [1;N_B]}$  to an actual emissivity spectrum  $(\varepsilon_i^{MMD})_{i \in [1;N_B]}$  using the MMD  $(\equiv$  spectral contrast) of the  $(\beta_i)_{i \in [1;N_B]}$ . This is achieved by linking the MMD to the minimum emissivity  $\varepsilon_{min}$  of the original spectrum using an empiricallydetermined relationship (Matsunaga, 1994). The  $\varepsilon_{min}$  is afterwards used to scale the relative emissivity spectrum. For that purpose, the MMD is first calculated as:

$$MMD = \max_{i}(\beta_{i}) - \min_{i}(\beta_{i})$$
(8)

<sup>220</sup> and then linked to the minimum emissivity  $\varepsilon_{min}$  via:

$$\varepsilon_{min} = \mathbf{A} - \mathbf{B}.\mathbf{M}\mathbf{M}\mathbf{D}^C \tag{9}$$

where A, B and C are coefficients which depend upon the Instrumental Spectral Response Function (ISRF) of the TIR sensor to be used, which are obtained from a regression using a representative emissivity spectra database. Finally, the  $\beta_i$  values can be rescaled to the emissivity spectrum following:

$$\varepsilon_i^{TES} = \frac{\varepsilon_{min}}{\min_i(\beta_i)} \beta_i \tag{10}$$

The final surface temperature  $T_{sur}^{TES}$  is then computed by conducting a last NEM iteration (equations (1) to (3)), using the estimated surface emissivity spectrum  $(\varepsilon_i^{TES})_{i \in [1;N_B]}$ . A summary of the TES method described above can be found in Figures 2 and 3 of Gillespie et al. (1998).

# 229 3. PERSEUS description

In this section, we present the end-to-end simulator that we designed and implemented for our study, which aims to optimize the spectral configuration of the TRISHNA TIR instrument when used with the TES method. Here we provide an overview of the simulator, along with descriptions of each related module.

# 235 3.1. PERSEUS overview

For the purpose of this study, an end-to-end simulator named PERSEUS 236 (Python End-to-end Remote SEnsing instrument Simulator) has been devel-237 oped. Indeed, it was necessary to use a modular, versatile and fast-computing 238 simulator, since a large number of simulations were conducted. As such, PERSEUS 239 can simulate top-of-atmosphere (TOA) and out-of-sensor (OS) radiances from 240 any set of driving factors related to observation conditions, the latter including 241 surface temperature  $T_{sur}$  and emissivity spectrum ( $\varepsilon_{\lambda}$ ), atmospheric parameters 242 (downwelling irradiance, upwelling radiance, transmittance), and the spectral 243 configuration of the considered TIR sensor (response function for each of the 244  $N_B$  channels, so-called ISRF). Applying the inversion procedure to the obtained 245 OS radiances then permits to retrieve the surface properties of interest, namely 246  $T_{sur}$  and channel emissivities  $(\varepsilon_i)_{i \in [1;N_B]}$ . The retrieved values can afterwards 247 be compared to their input counterparts in order to conduct performance analy-248 ses, provided reference emissivity spectra are convolved with the ISRF. Figure 1 249 displays a schematic overview of PERSEUS and its three modules: the radia-250 tive transfer module, the instrument module and the inversion module. These 251 modules are presented in the following subsections. 252

# 253 3.2. The radiative transfer module

The PERSEUS radiative transfer module uses the IDL-encoded COMANCHE radiative transfer tool (Poutier et al., 2002) which has been adapted to use MODTRAN 5.2 (Berk and Anderson, 2008). It allows the calculation, within the TIR spectral domain, of atmospheric parameters and TOA radiance, from the scene surface variables  $T_{sur}$  and  $(\varepsilon_{\lambda})$ , for a given atmospheric profile. The TOA radiance is calculated via:

$$L_{\lambda}^{sur\uparrow} = \varepsilon_{\lambda} B(T_{sur}, \lambda) + (1 - \varepsilon_{\lambda}) \frac{E_{\lambda}^{atm\downarrow}}{\pi}$$
(11)

$$L_{\lambda}^{TOA} = \tau_{\lambda}^{atm} L_{\lambda}^{sur\uparrow} + L_{\lambda}^{atm\uparrow}$$
(12)

where  $L_{\lambda}^{sur\uparrow}$  and  $L_{\lambda}^{atm\uparrow}$  are the surface leaving radiance and atmospheric upwelling radiance, respectively,  $E_{\lambda}^{atm\downarrow}$  is the atmospheric downwelling irradiance,



Figure 1: Schematic overview of PERSEUS and its three modules : the radiative transfer module, the instrument module and the inversion module. The variables calculated by the simulator are highlighted in dotted lines.

 $L_{\lambda}^{TOA}$  is the TOA radiance, and  $\tau_{\lambda}^{atm}$  is the atmospheric transmittance between the sensor and the target.

# 264 3.3. The instrument module

The aim of the PERSEUS instrument module is to account for the signal modifications applied to the TOA radiance by the TIR sensor, which eventually permits to obtain the OS radiance. These within-sensor signal modifications consist in applying the ISRF and instrumental noises.

269 3.3.1. ISRF application

<sup>270</sup> Upon entering the instrument module, the total TOA radiance  $L^{TOA}(\lambda)$ , <sup>271</sup> with  $\lambda$  the wavelength, is convolved with the spectral response function  $S_i(\lambda)$ <sup>272</sup> of each of the  $N_B$  channels, in order to obtain the equivalent TOA radiance in 273 channel  $i, L_i^{TOA}$ :

$$L_i^{TOA} = \frac{\int_{\lambda_{min,i}}^{\lambda_{max,i}} L^{TOA}(\lambda) . S_i(\lambda) . d\lambda}{\int_{\lambda_{min,i}}^{\lambda_{max,i}} S_i(\lambda) . d\lambda}$$
(13)

where  $\lambda_{min,i}$  and  $\lambda_{max,i}$  are the limit wavelengths of channel *i*. It should be noted that the PERSEUS ISRF module can be used to apply the ISRF on any spectral variable, such as the prescribed emissivity spectra ( $\varepsilon_{\lambda}$ ).

277 3.3.2. Applying instrumental noises

In order to determine the radiance equivalent noise on each channel Ne $\Delta L_i$ , PERSEUS uses the following noise model:

$$Ne\Delta L_i = \sqrt{a_i + b_i L_i} \tag{14}$$

where the term *a* includes the quantification noise and the dark current noise, while the term  $bL_i$  refers to the shot noise modelled by a Poisson process. Once the Ne $\Delta L_i$  is calculated, the corresponding instrumental noises are applied to each channel *i* as a white Gaussian noise with a standard deviation of  $\sigma =$ Ne $\Delta L_i$ :

$$L_i^{OS} = L_i^{TOA} + \mathcal{N}(0, \sigma = \text{Ne}\Delta L_i)$$
(15)

where  $L_i^{OS}$  is the OS radiance in any of the  $N_B$  channels *i*.

286 3.4. The inversion module

The PERSEUS inversion module includes the TES method with prior atmo-287 spheric corrections. As depicted in Figure 1, the atmospheric corrections can be 288 conducted considering reference or noised atmospheric parameters, where the 289 latter permit to account for uncertainties on atmospheric corrections. Noised 290 atmospheric parameters are obtained by adding Gaussian white noise at each 291 level of any atmospheric profile of temperature, relative humidity and ozone con-292 tent, before entering the radiative transfer module ( $\S$  3.2), by using the following 293 procedure: 294

• for the atmospheric temperature, a white Gaussian noise with a standard deviation of 0.8 K is added, as suggested by Barsi et al. (2003);

• for the relative humidity (RH) and the ozone concentration ( $[O_3]$ ), white Gaussian noises with respective standard deviations of  $0.1 \times \text{RH}$  and  $0.2 \times$  $[O_3]$  error are added, as suggested by Palluconi et al. (1999).

The outputs of the radiative transfer simulations conducted with these noised profiles are used to derive the atmospheric parameters of interest, namely atmospheric upwelling radiance  $L_i^{atm\uparrow}$ , atmospheric transmittance  $\tau_i^{atm}$  and atmospheric downwelling irradiance  $E_i^{atm\downarrow}$ . These atmospheric parameters are then used to conduct atmospheric corrections by computing the surface-leaving radiances  $L_i^{sur\uparrow}$  within each of the  $N_B$  channels *i*, following:

$$L_i^{sur\uparrow} = \frac{L_i^{OS} - L_i^{atm\uparrow}}{\tau_i^{atm}} \tag{16}$$

In the atmospheric noise-free case, atmospheric corrections rely on the outputs of radiative transfer simulations obtained from the original atmospheric profiles. Eventually, once the surface-leaving radiance  $L_i^{sur\uparrow}$  is obtained, the TES method is applied to the  $N_B$  surface-leaving radiances  $L_i^{sur\uparrow}$ , while using the atmospheric downwelling irradiance  $E_i^{atm\downarrow}$  (§ 2).

# 311 3.5. Performance analysis

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Once the surface temperatures  $T_{sur}^{TES}$  are estimated, they can be directly compared with their input counterparts, to assess the performance of the TES method for the retrieval of surface temperature and emissivity. The obtained emissivities  $(\varepsilon_i^{TES})_{i\in[1;N_B]}$  are compared against the ISRF-convolved surface emissivity spectra  $(\varepsilon_i)_{i\in[1;N_B]}$ . The latter are obtained by applying the ISRF of the instrument to the prescribed surface emissivity spectra  $(\varepsilon_{\lambda})$ .

It should be noted that the bandpass resampling of spectral values obtained by convolution with the ISRF is expected to add errors on the LST and LSE retrievals (Richter and Coll, 2002; Jiménez-Muñoz and Sobrino, 2006). However, as discussed in Jiménez-Muñoz and Sobrino (2006), this can not be avoided in end-to-end simulations study such as ours. Moreover, as we study relative
performances of various spectral configurations, and not absolute performances,
these effects should not significantly change our conclusions.

# 4. PERSEUS implementation in the context of the TRISHNA mis sion

For our study, we ran various simulations of TRISHNA measurements over a wide range of surface and atmospheric conditions. In the current section we present the datasets we used for characterizing the surface conditions (emissivity spectra, temperature) and the atmospheric conditions (variable profiles) of the pixel-sized scenes, along with the PERSEUS implementation used to obtain a representative dataset of TOA radiances.

# 333 4.1. The surface emissivity spectra dataset

We select the emissivity dataset simulated by Jacob et al. (2017) with the 334 SAIL-Thermique model (Olioso, 1995; Olioso et al., 2018), and previously used 335 to design the spectral configuration of the MISTIGRI TIR sensor (Jacob et al., 336 2021). This dataset, hereafter labelled SAIL271, comprises 271 simulated emis-337 sivity spectra of vegetation canopies over soils. It was implemented in order 338 to form a realistic and representative dataset of land surface emissivity spec-330 tra, dedicated to TIR studies. Indeed, the simulated emissivity spectra account 340 for radiance trapping within canopy and its subsequent cavity effect, and are 341 representative of a large range of soil and plant conditions, notably for Leaf 342 Area Index (LAI) and Average Leaf Angle (ALA). Among the 63,700 original 343 emissivity spectra simulated by considering a wide range of soil reflectance and 344 leaf reflectance / transmittance, the final 271 spectra were selected by using the 345 Spectral Angle Mapper method (SAM, Girouard et al., 2004). All the emissivity 346 spectra of the SAIL271 dataset were obtained considering a nadir viewing di-347 rection, because 1/ angular variation of canopy emissivity is low between nadir 348 and  $40^{\circ}$  (Olioso, 1995; Guillevic et al., 2003; Ren et al., 2015) although it can 349

Parameters	Values	
LAI	0, 0.25, 0.5, 1, 2, 4, 7	
ALA	$15^{\circ},  35^{\circ},  55^{\circ},  75^{\circ}$	
View zenith angle	Nadir	

Table 1: Summary of the soil and plant conditions taken into account in the SAIL271 dataset (from Jacob et al., 2017).

reach 0.01 over row-structured canopies (Sobrino and Caselles, 1990; Sobrino 350 et al., 2005), and 2/ the view zenith angle of the TRISHNA sensor is lower 351 than 34°. A summary of the range of these parameters is displayed in Table 1, 352 and a detailed description on the elaboration of the dataset can be found in 353 Jacob et al. (2017). The SAIL-Thermique model describes vegetation canopy 354 as a simple homogeneous volume of turbid medium, which induces errors when 355 considering most of continental surfaces with heterogeneous and discontinuous 356 canopies. According to the few comparison studies that are reported for emis-357 sivity models (Sobrino et al., 2005; Cao et al., 2018), resulting errors can reach 358 up to 0.005 over large ranges of LAI (0 to 7) and of view zenith angle (0 to  $50^{\circ}$ ). 359

# 360 4.2. The atmospheric profiles dataset

In order to work with a representative dataset of atmospheric profiles, we use the TIGR database that contains 2311 atmospheric profiles statistically selected over 80 000 radiosonde based records (Chevallier et al., 2000). We characterize these 2311 profiles by considering two main drivers of the atmospheric radiative transfer in the TIR spectral domain, namely the atmospheric water vapor content (AWVC) and the equivalent atmospheric temperature  $T_{eq}^{atm}$ . The AWVC drives the atmospheric transmittance, and is defined as (Jacob et al., 2003):

$$AWVC = \int_0^{z_{max}} \rho_v(z) dz$$
 (17)

where  $z_{max}$  is the highest altitude of the atmospheric profile, and  $\rho_v$  the water vapor density.  $T_{eq}^{atm}$  drives the atmospheric emission, and therefore  $E^{atm\downarrow}$  and  $L^{atm\uparrow}$ . It is defined as:

$$T_{eq}^{atm} = \frac{\int_{0}^{z_{max}} T(z) . \rho_{v}(z) . \mathrm{d}z}{\int_{0}^{z_{max}} \rho_{v}(z) . \mathrm{d}z}$$
(18)

In order to reduce the computational time, the TIGR dataset is filtered by comparing the 2311 profiles to each other, and by keeping those that meet the following conditions:

1. the difference between atmospheric water vapor content AWVC is larger than 0.15 g.cm<sup>-2</sup>;

2. the difference between equivalent atmospheric temperatures  $T_{eq}^{atm}$  is larger than 5 K.

These threshold values are empirically selected in order to obtain a number 378 of profiles consistent with our numerical capabilities, while conserving the enve-379 lope of the atmospheric profile distribution. The subsequent filtering results in a 380 selection of 24 atmospheric profiles. Figure 2 displays the original (Figure 2(a)) 381 and reduced (Figure 2(b)) atmospheric profile datasets in the  $T_{eq}^{atm}$  / AWVC 382 space. It shows that the reduced dataset preserves the envelope of the atmo-383 spheric profile distribution. However, the density of the atmospheric profiles is 384 homogenized in the reduced dataset as compared to the original one. Indeed, 385 the original dataset contains a larger number of dry atmospheres (AWVC<2.5 386  $g.cm^{-2}$ ) than wet ones (AWVC $\geq 2.5 g.cm^{-2}$ ), whereas the reduced dataset con-387 tains 13 wet and 11 dry atmospheres. Nevertheless, this new distribution fits 388 with the context of our study, since it accounts for wet atmospheres corre-389 sponding to tropical latitudes, the latter being of paramount interest for the 390 TRISHNA mission. The effect of this atmospheric profile selection on both the 391 TES performance and the representativeness of the results in the context of the 392 TRISHNA mission will be discussed in § 7. Note that the reduced dataset of 393 atmospheric profiles is hereafter labelled TIGR24. 394

395 4.3. TOA radiance simulations

Once the SAIL271 and TIGR24 datasets are obtained, and in order to take into account realistic surface/air temperature gradients, each surface emissivity



Figure 2: Representation of the atmospheric profiles in the  $T_{eq}^{atm}$ /AWVC plane for (a) the complete TIGR dataset, and (b) the TIGR24 one.

spectrum/atmospheric profile combination is associated to five surface temper-398 atures following  $T_{sur} = [T_{skin} - 5K; T_{skin}; T_{skin} + 5K; T_{skin} + 10K; T_{skin} + 15K],$ 399 where  $T_{skin}$  is the temperature at the lower level of the corresponding atmo-400 spheric profile. The  $271 \times 24 \times 5 = 32,520$  PERSEUS simulations are then 401 carried out in order to calculate the TOA radiances at nadir viewing direction, 402 for each possible combination of surface emissivity spectra/atmospheric pro-403 file/surface temperature. This results in a dataset of TOA radiances which is 404 structurally representative of a wide range of realistic situations. 405

# **5.** Setup for the sensitivity analysis

In this section we first present the reference TRISHNA instrumental configu-407 ration, including both the ISRF  $S_i(\lambda)$  and the Noise Equivalent Delta Radiance 408  $Ne\Delta L_i$  for each channel *i*. Next, we describe the sensitivity analysis based 409 on the changes in TES performance when moving the spectral channels. Fi-410 nally, we expose the methodological design for addressing the robustness of the 411 TRISHNA reference spectral configuration to uncertainties on ISRF due to in-412 strumental design, namely uncertainties on channel position and full width at 413 half maximum (FWHM). 414

# 415 5.1. TRISHNA reference instrumental configuration

The TRISHNA mission group set a baseline for the four spectral channels 416 of the TIR sensor, which theoretically allows a comfortable trade-off between 417 good performance of LST-retrieval methods and mission cost. Table 2 and Fig-418 ure 3 detail the central wavelength and FWHM for each of the four TRISHNA 419 spectral channels, labelled respectively TIR1, TIR2, TIR3 and TIR4. The four 420 TRISHNA TIR channels are pair-distributed on either side of the ozone ab-421 sorption feature spiking at 9.7  $\mu$ m, with TIR1 and TIR2 below and TIR3 and 422 TIR4 above. Moreover, water vapor absorption defines the atmospheric window 423 within which the four channels are located, with low atmospheric transmittance 424 below 8.25  $\mu$ m and above 12.5  $\mu$ m. The aforementioned ozone and water vapor 425

Table 2: Central wavelength  $\lambda_c$  and FWHM for the TRISHNA reference channels.

	TIR1	TIR2	TIR3	TIR4
$\lambda_c \; (\mu m)$	8.6	9.1	10.4	11.6
FWHM $(\mu m)$	0.35	0.35	0.7	1.0

Table 3: Instrumental noise parameters for the TRISHNA reference channels.

	TIR1	TIR2	TIR3	TIR4
$a_i \ (10^{-5} \ W^2.m^{-4}.sr^{-2}.\mu m^{-2})$	18.3	16.3	4.47	4.32
$b_i \; (10^{-8} \; W\!.m^{-2}.sr^{-1}.\mu m^{-1})$	411	547	8.13	175

<sup>426</sup> absorption features are expected to be the main constraining factors for channel<sup>427</sup> positioning.

The values of the a and b parameters used in our study for the calculation of the Ne $\Delta L_i$ , as described in § 3.3.2, are given in Table 3. These values were provided by CNES as first estimates of the TRISHNA instrumental noises. As expected with regards to their smaller FWHM, the noises on TIR1 and TIR2 are significantly larger than those on TIR3 and TIR4.

# 433 5.2. ISRF variations for the sensitivity analysis

The aim of this study is to define the best spectral configuration for the 434 TRISHNA TIR sensor. Therefore, the main task is to evaluate the effects of 435 changes in channel positions, *i.e.*, changes on the central wavelength  $\lambda_c$  of each 436 TIR channel. In order to do so, the TIR1/TIR2 and TIR3/TIR4 pairs are stud-437 ied separately. For each pair of channels, we conduct the sensitivity analysis by 438 considering simultaneous variations of both channels. This separation of chan-439 nels into pairs is motivated by the following three arguments, and consequences 440 on the results will be discussed in § 7. 441



Figure 3: Representation of the TRISHNA reference ISRF for spectral channels TIR1 (blue), TIR2 (red), TIR3 (green) and TIR4 (magenta). The mean atmospheric transmittance and downwelling irradiance over the profiles of the TIR24 dataset,  $\langle \tau_{atm} \rangle_{TIGR24}$  and  $\langle E_{atm}^{\downarrow} \rangle_{TIGR24}$ , are respectively displayed in black solid and dotted lines.

442	1.	The TES method is designed to perform best when the spectral configu-
443		ration captures the spectral variability of the observed scene (§ 1 and 2.1).
444		According to the subset of natural samples (rocks, soils, vegetation, wa-
445		ter and ice) within the ASTER database of emissivity spectra, the average
446		trend for emissivity variation is typified by minimum values between 8 $\mu{\rm m}$
447		and 9.5 $\mu\mathrm{m},$ and maximum values between 10 $\mu\mathrm{m}$ and 12 $\mu\mathrm{m}$ (see Figure
448		7 in Lagouarde et al., 2013). Therefore, the $\mathrm{TIR1}/\mathrm{TIR2}$ and $\mathrm{TIR3}/\mathrm{TIR4}$
449		pairs can be studied separately, since they permit to capture the minimum
450		and maximum emissivity values for typical natural targets.

<sup>451</sup> 2. This separation is consistent with the TRISHNA mission requirements, <sup>452</sup> which specify that two spectral channels shall be located within the [10  $\mu$ m -<sup>453</sup> 12  $\mu$ m] spectral domain for operational application of the SW method.

As compared to simultaneous variations of all four TRISHNA channels,
studying the TIR1/TIR2 and TIR3/TIR4 pairs separately significantly
reduces the number of channel configurations, and subsequently the computation load.

<sup>458</sup> For the TIR1/TIR2 pair study, the variations of position considered are:

- for  $\lambda_c^{TIR1}$ , from 8 to 9  $\mu$ m with a step of 0.1  $\mu$ m;
- for  $\lambda_c^{TIR2}$ , from 8.2 to 9.2  $\mu$ m with a step of 0.1  $\mu$ m;

In order to keep the same channel order, we set  $\lambda_c^{TIR1} \leq \lambda_c^{TIR2}$  with possible 461 channel superimposition. At the same time, TIR3 and TIR4 are set to their 462 reference positions, *i.e.*,  $\lambda_c^{TIR3} = 10.4 \ \mu \text{m}$  and  $\lambda_c^{TIR4} = 11.6 \ \mu \text{m}$  respectively 463  $(\S 5.1)$ . This experimental design leads to the generation of 85 ISRFs. When 464 convolved with the 32,520 TOA radiances described in § 4.3, we obtain 85  $\times$ 465 32,520 = 2,764,200 OS radiances for our sensitivity study. We note that  $\lambda_c^{TIR1}$ 466 can slide down to 8  $\mu$ m, where it crosses the water vapor absorption feature 467 below 8.25  $\mu$ m. This permits to highlight the sensitivity of the TES performance 468 to this absorption feature. 469



For the TIR3/TIR4 pair study, the variations of position considered are:

• for  $\lambda_c^{TIR3}$ , from 9.5 to 11  $\mu$ m with a step of 0.1  $\mu$ m;

# • for $\lambda_c^{TIR4}$ , from 10.5 to 12 $\mu$ m with a step of 0.1 $\mu$ m;

As for the TIR1/TIR2 pair, we set  $\lambda_c^{TIR3} \leq \lambda_c^{TIR4}$  with possible channel su-473 perimposition. At the same time, TIR1 and TIR2 are set to their reference 474 positions, *i.e.*,  $\lambda_c^{TIR1} = 8.6 \ \mu \text{m}$  and  $\lambda_c^{TIR2} = 9.1 \ \mu \text{m}$  respectively (§ 5.1). This 475 experimental design leads to the generation of 241 ISRFs, which results in a total 476 of  $241 \times 32,520 = 7,837,320$  OS radiances for our sensitivity study. Similarly to 477  $\lambda_c^{TIR1}$  in the TIR1/TIR2 case,  $\lambda_c^{TIR3}$  can slide down to 9.5  $\mu$ m, where it crosses 478 the ozone absorption feature at 9.7  $\mu$ m. This permits to highlight the sensitivity 479 of the TES performance to this absorption feature. This crossing with the ozone 480 absorption feature is done for TIR3 but not for TIR2, since similar results are 481 expected for both channels. Finally,  $\lambda_c^{TIR4}$  cannot slide above 12  $\mu$ m, which is 482 justified by a conceptual issue due to the complexity of implementing a detector 483 with such a high cutoff frequency. 484

# 485 5.3. Study on the robustness of the reference spectral configuration

Instrumental design makes mandatory the consideration of uncertainties on 486 the position and FWHM of each channel. For this purpose, we address the 487 robustness of the TRISHNA reference spectral configuration to these uncertain-488 ties. The goal is to anticipate how the TES performance would be impacted by 489 such uncertainties, and to evaluate if the observed impacts call for modifications 490 of the spectral configuration. For this, we define realistic uncertainty margins 491 for both the position and FWHM of each spectral channel, respectively labelled 492  $u(\lambda_i)$  and  $u(FWHM_i)$ , as presented in Table 4. Note that these uncertainty 493 margins are typical of those stated by instrument manufacturers. 494

Considering these uncertainties, we defined four worst-case spectral configurations with respect to atmospheric perturbations, where each of them corresponds to the displacement of one channel from the TRISHNA reference spectral configuration presented in § 5.1.

	TIR1	TIR2	TIR3	TIR4
$u(\lambda_i) \ (\mu \mathrm{m})$	$\pm 0.15$	$\pm 0.15$	$\pm 0.15$	$\pm 0.15$
$u(FWHM_i) \ (\mu m)$	$\pm 0.07$	$\pm 0.07$	$\pm 0.15$	$\pm 0.15$

Table 4: Uncertainties on the position of the central wavelength and on the FWHM of each TRISHNA spectral channel considered in the robustness study.

499	1.	A case where TIR1 is fully shifted to lower wavelengths, <i>i.e.</i> , consider-
500		ing $\lambda_c^{TIR1}$ = 8.6 $\mu$ m - 0.15 $\mu$ m = 8.45 $\mu$ m, with maximum bandwidth
501		$FWHM_{TIR1} = 0.35 \ \mu m + 0.07 \ \mu m = 0.42 \ \mu m$ . This corresponds to the
502		case where TIR1 would be most affected by the loss of transmission due
503		to atmospheric water vapor absorption below 8.25 $\mu \mathrm{m}.$
	0	A sees where TIDO is fully shifted to be used such as the issues iter

2. A case where TIR2 is fully shifted to larger wavelengths, *i.e.*, considering  $\lambda_c^{TIR2} = 9.1 \ \mu\text{m} + 0.15 \ \mu\text{m} = 9.25 \ \mu\text{m}$ , with maximum bandwidth  $FWHM_{TIR2} = 0.35 \ \mu\text{m} + 0.07 \ \mu\text{m} = 0.42 \ \mu\text{m}$ . This corresponds to the case where TIR2 would be most affected by the loss of transmission due to ozone absorption at 9.7  $\mu$ m.

509	3. A case where TIR3 is fully shifted to lower wavelengths, <i>i.e.</i> , consider-
510	ing $\lambda_c^{TIR3} = 10.4 \ \mu\text{m}$ - 0.15 $\mu\text{m} = 10.25 \ \mu\text{m}$ , with maximum bandwidth
511	$FWHM_{TIR3} = 0.70 \ \mu\text{m} + 0.15 \ \mu\text{m} = 0.85 \ \mu\text{m}$ . This corresponds to the
512	case where TIR3 would be most affected by the loss of transmission due
513	to ozone absorption at 9.7 $\mu \mathrm{m}.$

4. A case where TIR4 is fully shifted to larger wavelengths, *i.e.*, considering  $\lambda_c^{TIR4} = 11.6 \ \mu m + 0.15 \ \mu m = 11.75 \ \mu m$ , with maximum bandwidth  $FWHM_{TIR4} = 1.00 \ \mu m + 0.15 \ \mu m = 1.15 \ \mu m$ . This corresponds to the case where TIR4 would be most affected by the loss of transmission due to atmospheric water vapor absorption at 12.5 \ \mu m.

For each of these 4 worst-case scenarii, we calculated root mean square errors (RMSEs) on TES retrievals of surface temperature, by using the protocol presented in § 4.3 to simulate the dedicated 32,520 TOA radiances.

## 522 6. Results of the sensitivity analysis

We present in this section the results of our sensitivity analysis for both the 523 TIR1/TIR2 and TIR3/TIR4 variation cases. We first detail how the various 524 TES calibrations are conducted, to continue with the results of our sensitivity 525 analysis for both variation cases. Eventually, we present the results we obtained 526 when studying robustness of the TRISHNA reference spectral configuration to 527 uncertainties on channel position and FWHM. In the following, we mostly con-528 centrate on the results obtained for  $T_{sur}^{TES}$ . Indeed, as compared to land surface 529 emissivity, LST is our decision variable, since it is the first order parameter to 530 be used when defining the TRISHNA products. 531

# 532 6.1. Calibration of the TES MMD equation

For each of the 85 + 241 = 326 ISRFs used in this study, the coefficients A, B, and C of the  $\varepsilon_{min}$ -MMD relationship (equation 9) were calibrated over the SAIL271 dataset. The RMSE was calculated, for each spectral configuration, as the calibration residual error.

The obtained RMSE values for calibration and validation were all lower than 537 0.008. Table 5 displays averaged values of the coefficients A, B and C for each 538 of the TIR1/TIR2 and TIR3/TIR4 study cases, along with the corresponding 539 coefficients of variation. As observed in previous studies, the MMD calibration 540 values do not vary much from one spectral configuration to another for a given 541 number of channels (Sobrino and Jiménez-Muñoz, 2014; Jacob et al., 2017). 542 As compared to former studies dealing with 4 TIR bands, our coefficients are 543 larger, by 15% and 10% relative, for parameter B and C respectively. Apart 544 from maximum values of ISRF  $S_i(\lambda)$ , that were lower in our study (0.85 versus 545 (0.9), no other explanation could be found for these larger values of parameters 546 B and C. 547

# 548 6.2. TIR1/TIR2 case results

After performing atmospheric corrections, we computed  $T_{sur}^{TES}$  and  $(\varepsilon_i^{TES})_{i \in [1;N_B]}$ for each of the 2 764 200 OS radiances defined in § 5.2, and we calculated RM-

Table 5: Pairs of (mean value / coefficient of variation) obtained for the calibrated MMD coefficients A, B and C in both TIR1/TIR2 and TIR3/TIR4 study cases.

	А	В	С
TIR1/TIR2 case	0.984 / 0.17%	0.815 / 1.90%	0.912 / 1.80%
TIR3/TIR4 case	0.988~/~0.03%	0.834 / $2.80%$	0.926~/~0.10%

SEs for each of the 85 spectral configurations. Figure 4 displays the RMSEs obtained for surface temperature when considering (a) known and (b) noised atmospheric profiles, where known and noised atmospheric profiles correspond to accurate and inaccurate atmospheric corrections, respectively. What directly strikes in Figure 4 is that the TES retrievals of surface temperature are relatively in sensitive to the TIR1 and TIR2 channel positions, with respective maximum RMSE variations of 0.23 K and 0.15 K.

Figure 4(a) shows that for accurate atmospheric corrections (known atmo-558 spheric profiles), a large RMSE gradient is observed along the  $\lambda_c^{TIR1}$  axis. In 559 this case, RMSE increases as wavelength decreases, with maximum RMSEs ob-560 tained at  $\lambda_c^{TIR1} = 8.0 \ \mu \text{m}$  and a sharp increase from 8.1 to 8.0  $\mu \text{m}$ . Figure 4(b) 561 shows that for inaccurate atmospheric corrections (noised atmospheric profiles), 562 an additional RMSE gradient occurs along the  $\lambda_c^{TIR2}$  axis, although it remains 563 less significant than the one along the  $\lambda_c^{TIR1}$  axis. These RMSEs variations are 564 explained by the TES sensitivity to inaccurate atmospheric corrections (Gille-565 spie et al., 1998; Sobrino and Jiménez-Muñoz, 2014), where this sensitivity is 566 enhanced when both TIR1 and TIR2 channels come closer to the beginning of 567 the water vapor absorption feature at 8.25  $\mu$ m (decrease in atmospheric trans-568 mittance). 569

Interestingly, bringing the two channels closer seems to have little effect on the surface temperature RMSE, with minimum RMSE values obtained when both channels are centered on wavelengths larger than 8.7  $\mu$ m.



Figure 4: RMSEs in surface temperature obtained in the TIR1/TIR2 study, without (a) and with (b) atmospheric noise. The blue circle highlights in each case the reference spectral configuration: (a)  $RMSE(T_{sur}) = 0.31$ K, (b)  $RMSE(T_{sur}) = 0.85$ K.

#### 6.3. TIR3/TIR4 case results 573

Figure 5 is the same as Figure 4 but for the TIR3 and TIR4 positions defined 574 in § 5.2. When considering accurate atmospheric corrections (Figure 5(a)), and 575 similarly to the TIR1/TIR2 study case, the TES method is not very sensitive to 576 the position of the two channels, with a maximum RMSE variation of 0.19 K. 577 We note two main features from Figure 5(a): 578

1. increases in RMSEs are observed towards both limits of the  $\lambda_c^{TIR4}$  interval, 579 with maximum RMSEs at  $\lambda_c^{TIR4} = 10.4 \ \mu \text{m};$ 580

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2. a region with local minimum RMSEs is observed for  $\lambda_c^{TIR3} \in [9.5\mu m]$ ;  $9.6 \mu m$ ].

On the other hand, a clear increase in RMSEs appears around 9.7  $\mu m$  when 583 considering inaccurate atmospheric corrections (Figure 5(b)), which clearly indi-584 cates a sensitivity of the TES method to ozone and water vapor absorption. This 585 sensitivity increases the maximum RMSE variation by more than a factor 2, as 586 compared to the case of accurate atmospheric corrections, namely 0.42 K. Thus, 587 a maximum RMSE of 1.16 K is obtained for channel locations corresponding to 588 lower atmospheric transmittance, namely  $\lambda_c^{TIR3} = 9.6 \ \mu m$  that corresponds to 589 ozone absorption and  $\lambda_c^{TIR4} = 12.0 \ \mu m$  that corresponds to water vapor absorp-590 tion. For the same reason, an RMSE gradient is observed along the  $\lambda_c^{TIR4}$  axis, 591 with RMSEs increasing as  $\lambda_c^{TIR4}$  increases, since atmospheric transmittance de-592 creases because of water vapor absorption (see Figure 3). When restricting the 593 TIR3 location between 10.0  $\mu$ m and 11.0  $\mu$ m, the sensitivity of the TES method 594 to TIR3 and TIR4 positions is lowered, with a maximum RMSE variation of 595 0.305 K. Overall, and conversely to the TIR1/TIR2 case, the sensitivity of the 596 TES method is increased when considering noise in the TIR3/TIR4 case, which 597 could be explained by larger atmospheric perturbations on TIR measurements 598 within the  $[11 \ \mu\text{m} - 12 \ \mu\text{m}]$  spectral interval, as compared to the  $[8.5 \ \mu\text{m} - 9.5 \ \mu\text{m}]$ 599 one (see Figure 3). 600

Another result displayed by Figure 5 is that the proximity and even su-601 perimposition of the two channels does not significantly deteriorate the TES 602



Figure 5: RMSEs in surface temperature obtained in the TIR3/TIR4 study, without (a) and with (b) atmospheric noise. The blue circle highlights in each case the reference spectral configuration: (a)  $RMSE(T_{sur}) = 0.31$ K, (b)  $RMSE(T_{sur}) = 0.85$ K.

performance. This trend, observed for both the TIR1/TIR2 and TIR3/TIR4
cases, suggests that the TES method may have comparable performance when
considering 3 or 4 spectral channels. Its consequence on the design of the TRISHNA instrument, as well as other future similar ones, will be discussed in
§ 7.

# 608 6.4. Results on emissivity

The results we obtain for land surface emissivity are similar to those obtained 609 for land surface temperature. In the case of inaccurate atmospheric corrections, 610 the sensitivity of the TES method to channel positions is similar for emissivity 611 and temperature retrievals. Indeed, we observe similar RMSE gradients, espe-612 cially for  $\varepsilon_{TIR2}^{TES}$  and  $\varepsilon_{TIR4}^{TES}$  (Figure 6 top right versus Figure 4(b) and Figure 6 613 bottom right versus Figure 5(b)). Furthermore,  $\varepsilon_{TIR2}^{TES}$  is less sensitive to the de-614 crease in  $\lambda_c^{TIR1}$  than  $\varepsilon_{TIR1}^{TES}$  (Figure 6 top left versus top right), with a difference 615 in RMSE variations = 0.013. Similarly,  $\varepsilon_{TIR3}^{TES}$  is less sensitive to the increase in 616  $\lambda_c^{TIR4}$  than  $\varepsilon_{TIR4}^{TES}$  (Figure 6 bottom left versus bottom right). Conversely, the 617 RMSEs obtained on  $\varepsilon_{TIR4}^{TES}$  are less impacted by the proximity of  $\lambda_c^{TIR3}$  with 618 the ozone absorption feature at 9.7  $\mu \mathrm{m},$  as compared to the RMSEs obtained on 619  $\varepsilon_{TIB3}^{TES}$ . Eventually, these results highlight that the TES retrievals of emissivities 620 in the case of inaccurate atmospheric corrections behave in the same way as 621 their temperature counterparts, with a sensitivity that seems directly linked to 622 the atmospheric transmittance within their respective channels. 623

#### 6.5. Results on the robustness of the spectral configuration

For each of the 4 worst-case scenarii for the TRISHNA spectral configuration related to uncertainties on channel positions and FWHM (§ 5.3), we calculated RMSEs on TES retrievals of surface temperature. The results, summarized in Table 6, confirm the overall low sensitivity of the TES method to channel position and FWHM, even when considering inaccurate atmospheric corrections. Indeed, the maximum difference with the reference case is 0.07 K, when TIR4 is most affected by water vapor absorption. The results also show that all



Figure 6: RMSEs in emissivity as obtained in the TIR1/TIR2 study for  $\varepsilon_{TIR1}^{TES}$  and  $\varepsilon_{TIR2}^{TES}$  (top panels), and in the TIR3/TIR4 study for  $\varepsilon_{TIR3}^{TES}$  and  $\varepsilon_{TIR4}^{TES}$  (bottom panels), all considering atmospheric noise. The blue circle highlights in each case the reference spectral configuration:  $RMSE(\varepsilon_{TIR1}^{TES}) = 0.026, RMSE(\varepsilon_{TIR2}^{TES}) = 0.025, RMSE(\varepsilon_{TIR3}^{TES}) = 0.026, RMSE(\varepsilon_{TIR4}^{TES}) = 0.026$ .

Table 6: RMSEs on  $T_{sur}^{TES}$  obtained in the worst-case scenarii related to uncertainties on channels positions and FWHM defined in § 5.3

Channel ( $\lambda_c$ ,FWHM)	Case	RMSEs on $T_{sur}^{TES}$ (in K)
TIR1 $(8.45\mu m, 0.42\mu m)$	H2O absorption	0.856
TIR2 $(9.25\mu m, 0.42\mu m)$	$O_3$ absorption	0.863
TIR3 (10.25 $\mu\mathrm{m},\!0.85\mu\mathrm{m})$	$O_3$ absorption	0.883
TIR4 (11.75 $\mu$ m,1.15 $\mu$ m)	H2O absorption	0.923
Reference spectral configuration	-	0.850

four channels should be located away from their respective nearby absorption feature, in order to maintain low RMSEs ( $\leq 1$  K) when considering worst-cases, and therefore to obtain a TRISHNA spectral configuration that is more robust to uncertainties on channel positions and FWHM.

# 636 6.6. Recommended TRISHNA spectral configuration

The results of the previous section 6.5 permit to define criteria for a possible 637 improvement of the TRISHNA spectral configuration with regards to uncer-638 tainties on channel positions and FWHM. Despite the apparent equivalence we 639 report for the TES performance between the three and four channels configura-640 tions, we recommend a four-channels spectral configuration for the TRISHNA 641 mission. This choice is motivated by two main operational arguments, namely 642 (1) to secure the mission continuity in case of channel failures, as encountered 643 with ASTER, LANDSAT and ECOSTRESS, and (2) to allow a better character-644 ization of land surface emissivity spectra for downstream scientific applications. 645 In this 4-channel configuration, the channel-wise criteria are the following. 646

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• For TIR1, a shift towards larger wavelengths would attenuate the impact of water vapor absorption below 8.25  $\mu$ m. We therefore recommend to slightly shift  $\lambda_c^{TIR1}$  from 8.6  $\mu$ m to 8.65  $\mu$ m. For TIR2, a shift towards lower wavelengths would mitigate the impact of ozone absorption at 9.7 μm. We therefore recommend to shift λ<sup>TIR2</sup><sub>c</sub> from 9.1 μm to 9.0 μm. Then, the overlap between TIR1 and TIR2 channels would be small, given FWHM values for both channels.

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• For TIR3, a shift towards larger wavelengths would limit the impact of ozone absorption at 9.7  $\mu$ m. Vidal et al. (2021) reported an optimal position at 10.6  $\mu$ m when dealing with the performance of the Split Window method in the context of the TRISHNA mission, and showed that this positioning of TIR3 was robust to the uncertainties on the parameters used in the SW formulations, such as emissivity and AWVC. We therefore recommend to shift  $\lambda_c^{TIR3}$  at 10.6  $\mu$ m.

For TIR4, a shift towards lower wavelengths would attenuate the impact
 of water vapor absorption above 12 μm. Nonetheless, a limited overlap
 between TIR3 and TIR4 channel is required to ensure satisfactory per formance of the Split Window method in the context of the TRISHNA
 mission (see Vidal et al., 2021). We therefore recommend to keep the
 original TIR4 position at 11.6 μm.

The recommendations listed above lead to the definition of a new TRISHNA 667 spectral configuration, detailed in Table 7 and Figure 7. In order to ensure that 668 this spectral configuration is more robust to uncertainties on channel positions 669 and FWHM than the reference configuration, we compare the corresponding 670 RMSE values on LST to those obtained with the reference spectral configu-671 ration. For the four worst-case scenarii we consider, apart from TIR4 channel 672 located close to 12  $\mu$ m, the new configuration slightly improves the performance 673 of the TES inversion, with differences in RMSEs of about 0.02K. 674

The RMSE differences between the two spectral configurations is even lower when comparing the overall TES performance. Indeed, for each case we obtain a RMSE over our 32,520 simulations of 0.850 K and 0.849 K respectively, for a global gain of 0.001 K with the recommended configuration. This apparent low

	TIR1	TIR2	TIR3	TIR4
$\lambda_c \; (\mu m)$	$8.65{\pm}0.10$	$9.0{\pm}0.10$	$10.6 {\pm} 0.15$	$11.6 {\pm} 0.15$
FWHM $(\mu m)$	$0.35{\pm}0.07$	$0.35{\pm}0.07$	$0.7 {\pm} 0.15$	$1.0 {\pm} 0.15$

Table 7: Central wavelength  $\lambda_c$  and FWHM for the TRISHNA recommended channels.



Figure 7: Representation of the TRISHNA recommended ISRF with spectral channels TIR1 (blue), TIR2 (red), TIR3 (green) and TIR4 (magenta). The mean atmospheric transmittance and downwelling irradiance over the profiles of the TIR24 dataset,  $\langle \tau_{atm} \rangle_{TIGR24}$  and  $\langle E_{atm}^{\downarrow} \rangle_{TIGR24}$ , are respectively displayed in black solid and dotted lines.

<sup>679</sup> gain achieved by the spectral channel positions optimization will be discussed<sup>680</sup> in § 7.

# 681 7. Discussion

# <sup>682</sup> 7.1. On the selection of atmospheric profiles

In  $\S$  4.2, we detail the TIGR24 subset of atmospheric profiles selected for 683 our study, and highlight that this subset does not conserve the initial profile 684 density of the TIGR dataset in the  $T_{atm}^{eq}$ /AWVC plane. Indeed, the comparison 685 between Figure 2(a) and 2(b) shows that the TIGR24 subset drastically reduces 686 the density of dry atmosphere profiles (AWVC  $< 2.5 \text{ g.cm}^{-2}$ ) as compared to 687 the original TIGR dataset. As such, the TIGR24 subset includes fewer dry 688 atmosphere profiles (11) than wet atmosphere profiles (13). However, the cur-689 rent study is conducted in the context of the TRISHNA mission, which regions 690 of interest include tropical regions such as India. Therefore, it makes sense to 691 assess the TES performance by weighting the occurrence of wet atmospheres. 692 First, atmospheric corrections are less accurate with such wet atmosphere pro-693 files. Second, accuracy on atmospheric corrections is a first order driver for the 694 performance of the TES method (Gillespie et al., 1998; Jacob et al., 2004, 2017; 695 Schmugge et al., 1998, 2002; Sobrino and Jiménez-Muñoz, 2014). Therefore, 696 accounting for such wet atmosphere profiles to analyze TES performance is rel-697 evant for the preparation of the TRISHNA mission. Indeed, it is expected that 698 considering more dry atmospheres than in the 'TIGR24' dataset would reduce 699 the RMSEs and result in better TES performance, while mitigating the impact 700 of the more problematic wet atmospheres. In order to provide a realistic evalu-701 ation of the TES performance in the context of the TRISHNA mission, it would 702 be valuable to quantify the occurrence of dry and wet atmospheric profiles on a 703 global scale. Such a study should be conducted in the context of this mission. 704

# 705 7.2. On the TIR1/TIR2 and TIR3/TIR4 pairs separation

It is explained in § 5.2 that our sensitivity study is based on sliding the positions of two TRISHNA channels at the same time, by considering separately the TIR1/TIR2 and TIR3/TIR4 pairs of channels. As opposed to moving the positions of all four channels at the same time, this trade-off is motivated by the considerations discussed below.

1. For the use of the TES method, these two pairs of channels have distinct 711 tasks, namely capturing the minimum (for TIR1/TIR2) and maximum 712 (for TIR3/TIR4) emissivity values across the emissivity spectra of natural 713 targets. In order to do so, we locate these pairs within spectral intervals 714 for which atmospheric and emissivity-related properties are known to be 715 optimal for TES application, namely  $[8 \ \mu m - 9.5 \ \mu m]$  and  $[10 \ \mu m - 12 \ \mu m]$ , 716 respectively (Lagouarde et al., 2013). Therefore, sliding in pairs permits 717 to study the positioning of channels in the context of their respective roles 718 in these spectral intervals. 719

2. The consideration of four sliding channels would result in a number of simulations that would exceed our computational capability. Indeed, even considering a reduced interval for TIR3 outside of atmospheric ozone absorption between 10  $\mu$ m and 11  $\mu$ m would increase the number of simulations by a factor 50, for over 500 millions of simulations in that case.

Given the operational aspect of the TRISHNA mission, the conclusions 725 drawn in this study are based on the results obtained in the operational case 726 (inaccurate atmospheric corrections). In that case, the factors that can affect 727 the TES performance are supposedly (1) the instrumental noise, (2) inaccurate 728 atmospheric corrections, and (3) variation of the observed emissivity which con-729 trast has to be correctly measured for the TES method. On the one hand, we 730 do not observe any sensitivity to the instrumental noise in the operational case. 731 Moreover, we run for the sake of completeness the TIR1/TIR2 and TIR3/TIR4 732 cases with inaccurate atmospheric corrections but no instrumental noise. The 733 results obtained show similar RMSE variations than in the operational case 734 (Figure 4(b) and 5(b)), and the RMSE differences between both cases are of an 735 order of magnitude of a few  $10^{-3}$  K. This suggests that TES is faintly sensitive 736 to instrumental noise in the operational case. On the other hand, results of 737

 $\S$  6.2 and 6.3 show a moderate sensitivity of TES to channel positions, which 738 appears to be due to variations of atmospheric perturbations in the operational 739 case, notably in the case of the TIR1/TIR2 pair. Indeed, this case was expected 740 to result in a higher TES sensitivity to channel position than the TIR3/TIR4, 741 given the variation of emissivity observed in the  $[8 \ \mu\text{m} - 9.5 \ \mu\text{m}]$  spectral interval 742 as compared to the  $[10 \ \mu\text{m} - 12 \ \mu\text{m}]$  interval (Lagouarde et al., 2013). This re-743 sult supports the hypothesis that in the operational case, the TES performance 744 depends mainly on atmospheric perturbations rather than emissivity variations. 745 This main dependency to atmospheric conditions, which implies that the TES 746 sensitivity to channel positions within atmospheric windows is moderate, hints 747 that sliding at the same time the four TRISHNA TIR channels is likely to 748 provide similar results on TES performance as obtained in this study, as well 749 as similar recommendations for an optimized TRISHNA spectral configuration, 750 provided that TRISHNA channels are not located nearby the limits of the at-751 mospheric window. This statement is also supported by the similarity between 752 the results of the current study and those reported for the MISTIGRI mission 753 (Jacob et al., 2021) with different spectral configurations including more than 754 two channels. 755

# 756 7.3. On the TES performance with known atmospheres

On the one hand, result analysis in § 6.2 and 6.3, including interpretation of Figures 4(b) and 5(b), underline the sensitivity of the TES method to atmospheric perturbations related to water vapor absorption (TIR1, TIR2 and TIR4) and ozone absorption (TIR3). This is consistent with former studies that reported similar sensitivities (Gillespie et al., 1998; Jacob et al., 2004; Schmugge et al., 1998, 2002; Sobrino and Jiménez-Muñoz, 2014).

<sup>763</sup> On the other hand, Figures 4(a) and 5(a) indicate a large sensitivity of <sup>764</sup> the TES method to the TIR1 and TIR4 channel positions when considering <sup>765</sup> accurate atmospheric corrections. For the TIR1/TIR2 case, RMSE values on <sup>766</sup> TES retrievals of surface temperature increase as TIR1 slides toward 8.0  $\mu$ m. For <sup>767</sup> the TIR3/TIR4 case, RMSE values increase when TIR4 comes closer to either <sup>768</sup> 10.5  $\mu$ m or 12  $\mu$ m, which induces a region with lower RMSE values between <sup>769</sup> both limits. Additionally, the TES sensitivity to TIR3 channel position relative <sup>770</sup> to ozone absorption at 9.7  $\mu$ m is very low. Considering the underlying physical <sup>771</sup> processes at play, possible explanations involve the impact of both atmospheric <sup>772</sup> downwelling irradiance and instrumental noise.

1. Jacob et al. (2017) reported a large sensitivity of the TES method to 773 the atmospheric downwelling irradiance below 8.5  $\mu$ m. This is consistent 774 with the gradient of RMSE values along the TIR1 axis in Figure 4(a), 775 and in accordance with the change of atmospheric downwelling irradiance 776 within this spectral domain due to water vapor absorption (Figures 3) 777 and 7). The same physical process, namely water vapor absorption, also 778 explains the increase of RMSE values when TIR4 comes closer to  $12.0 \ \mu m$ 779 in Figure 5(a). Overall, the impact of atmospheric downwelling irradiance 780 can be explained by Equation 11, where any error on emissivity retrieval 781 is enhanced by the weighting of atmospheric downwelling irradiance. 782

2. In case of accurate atmospheric corrections, a possible perturbation of TES 783 performance is the instrumental noise. In the current study, the latter is 784 expressed as a white noise that accounts for dark current and quantifi-785 cation (§ 3.3.2). Figure 8 displays the quadratic value of Ne $\Delta$ L obtained 786 with the PERSEUS instrument module for the TIR3 (left panel) and TIR4 787 (right panel) channels, in the TIR3/TIR4 study case. On the one hand, 788 a clear RMS gradient along the  $\lambda_c^{TIR3}$  axis is observed in the Ne $\Delta L_{TIR3}$ 789 case, with RMS decreasing as wavelength decreases down to 9.6  $\mu$ m near 790 the ozone atmospheric feature. On the other hand, an RMS gradient is 791 observed along the  $\lambda_c^{TIR4}$  axis for the Ne $\Delta L_{TIR4}$  case, with RMS increas-792 ing as wavelength decreases. The combination of both quadratic noises 793  $Ne\Delta L_{TIR3}$  and  $Ne\Delta L_{TIR4}$  can explain the RMSE variations observed in 794 Figure 5(a) for TES retrievals of surface temperature. More particularly, 795 it can explain the lowest RMSE values obtained across the ozone atmo-796 spheric feature at  $9.7\mu m$ , and the largest RMSE values obtained when 797

 $\lambda_c^{TIR4}$  decreases down to 10.5  $\mu$ m. Eventually, these results are confirmed by the fact that the magnitude of the Ne $\Delta$ L variations observed in Figure 8 corresponds to the magnitude of their corresponding surface temperature RMSE variations in Figure 5(a), where both magnitudes are compared using Planck's law (Equation 6).

The aforementioned considerations also suggest that the TES method is more sensitive to instrumental noise than to uncertainties on atmospheric downwelling irradiance, when used without atmospheric perturbations. Indeed, the lowest RMSE values for surface temperature in Figure 5(a) are obtained across the ozone emission feature centered at 9.7  $\mu$ m. However, further investigations on the sensitivity of the TES method to the combination of these two driving factors are needed to conclude on this issue.

For sake of completeness, we also address in Figure 9 the variations of calibra-810 tion RMSEs for the  $\varepsilon_{min} - MMD$  relationship (Equation 9). The obtaining of 811 these calibration RMSE values for both TIR1/TIR2 and TIR3/TIR4 study cases 812 are described in § 6.1. Both Figure 9(a) and (b) display a low sensitivity of the 813 MMD calibration to the channel positions, with maximum RMSE variations of 814 0.0007 and 0.0021 for the TIR1/TIR2 and TIR3/TIR4 case study, respectively. 815 These low variations cannot explain the surface temperature RMSE variations 816 of 0.233 K and 0.185 K observed in Figure 4(a) and 5(a), respectively. Moreover, 817 we observe different patterns in Figure 9(a) and (b), as compared to those dis-818 played in Figure 4(a) and 5(a), except for the increase in RMSE with decreasing 819  $\lambda_c^{TIR4}$  in Figure 9(b). However, as mentioned above, this decrease in Figure 5(a) 820 is likely to be due to instrumental noise, notably given the low magnitude of 821 the calibration RMSE values. These results support the idea that with accurate 822 atmospheric corrections, TES is primarily sensitive to instrumental noise and 823 atmospheric downwelling irradiance. 824

# 825 7.4. On the TES performance with 3 or 4 channels

The results presented in this paper hint at the fact that when superimposing two of the four spectral channels considered, the TES performance does not



Figure 8: Quadratic values (labelled RMS for Root mean square) of Ne $\Delta$ L obtained from the instrument module when considering the TIR3 (left panel) and TIR4 (right panel) channels. The blue circle highlights in each case the reference spectral configuration:  $RMS(Ne\Delta L_{TIR3}) = 6.725 \times 10^{-3} \text{ W.m}^{-2} \text{.sr}^{-1} . \mu \text{m}^{-1} \text{ and } RMS(Ne\Delta L_{TIR4}) = 7.526 \times 10^{-3} \text{ W.m}^{-2} \text{.sr}^{-1} . \mu \text{m}^{-1}.$ 



Figure 9: Calibration RMSEs of the MMD equation (9) as computed in the TIR1/TIR2 case (a), and in the TIR3/TIR4 case (b). The blue circle highlights in each case the reference spectral configuration with a calibration RMSE of 0.0052.

drastically change as compared to other effective spectral configurations. This could reveal that the TES method has similar performance when considering 3 or 4 spectral channels. This observation is reinforced by results of previous studies that obtained similar, and sometimes even slightly better, TES performance with three-channels configurations than four-channels ones (Hulley et al., 2010; Sobrino and Jiménez-Muñoz, 2014; Jiménez-Muñoz et al., 2014; Jacob et al., 2017, 2021).

From a general point of view, this issue is paramount since a 3-channels con-835 figuration could sensibly reduce the production cost for any TIR instrument as 836 compared to a 4-channels configuration. A study specifically dedicated to com-837 pare the performance of these configurations should be conducted in the context 838 of future missions. However, in the context of the TRISHNA mission, as the 839 aforementioned results do not specifically favor the 3 or 4 channels configurations 840 in term of TES performances, we recommend a 4-channels configuration for two 841 main operational reasons, namely (1) to secure the mission continuity in case of 842 channel failures, as encountered with ASTER, LANDSAT and ECOSTRESS, 843 and (2) to allow a better characterization of land surface emissivity spectra for 844 downstream scientific applications. 845

## <sup>846</sup> 7.5. On the improvements of the recommended TRISHNA spectral configuration

As compared to the reference TRISHNA spectral configuration, we showed 847 in § 6.6 that the gain on TES performance with the recommended configuration 848 is not significant ( $\approx 0.001$  K). On the one hand, this result can be explained 849 by the overall low sensitivity of TES to channel positions observed in  $\S$  6.2 850 and 6.3, especially when considering channels within the atmospheric window. 851 On the other hand, and as reported by Vidal et al. (2021), the performance 852 of the Split Window method is improved with the recommended configuration, 853 showing an overall gain in RMSE of 0.2 K as compared to the reference config-854 uration. Moreover in the SW case, this configuration is shown to be robust to 855 uncertainties on emissivity, AWVC, and channel positioning. Therefore, even if 856 our recommended spectral configuration does not significantly improve the TES 857

performance for the retrieval of land surface temperature, it improves the performance of the Split Window method, while being more robust to uncertainties
on channel position and FWHM.

# <sup>861</sup> 7.6. On the future optimization of the TES method for TRISHNA applications

Although the objective of this study was to evaluate the sensitivity of the 862 TES method to the TRISHNA spectral configuration, the TES method was 863 not numerically optimized for the TRISHNA instrument. We expect that such 864 an optimization would increase the observed performance, therefore improving 865 the quality of the TRISHNA LST and land surface emissivity (LSE) retrievals. 866 Indeed, the several thresholds used in the NEM module (§ 2.1.1) were set in 867 the current study to their ASTER mission values, due to a lack of information 868 for other sensors. On the one hand, considering that both the convergence and 869 divergence thresholds (respectively  $t_2$  and  $t_1$ ) are supposed to be fixed at the 870 corresponding instrument Ne $\Delta L$  value, and given the fact that the TRISHNA 871  $Ne\Delta L$  value is an order of magnitude lower than the ASTER one, the TES 872 surface temperature estimation is expected to be improved. On the other hand, 873 the remaining thresholds for emissivity refinement could as well be improved 874 using more comprehensive emissivity databases than the ones used to set their 875 ASTER value. This update of the thresholds is also believed to possibly improve 876 the TES estimation of surface temperature and emissivity, especially regarding 877 the distinction between reflective and grey-body materials. 878

# 879 8. Conclusion

From a process perspective, the current study evaluates the sensitivity of the TES method to channel positions and bandwidths in the context of the TR-ISHNA mission with a four-channel configuration. We highlight the sensitivity of the TES method to atmospheric downwelling irradiance and instrumental noise, under conditions of accurate atmospheric corrections. In the case of inaccurate atmospheric corrections, the TES methods is sensitive to channel locations and bandwidths, more particularly when filter response functions extend over spectral intervals with atmospheric perturbations, namely absorption
by water vapor and ozone. These results are strengthened by the two following
considerations.

1. On the one hand, the TES method presents a low sensitivity to channel positions within both [8.25  $\mu$ m - 9.5  $\mu$ m] and [10  $\mu$ m - 12  $\mu$ m] atmospheric windows. On the other hand, the TES performance is slightly degraded when locating channels on spectral interval with atmospheric perturbations. Therefore, it does not seem necessary to investigate a finer spectral sampling.

Our results are similar to those reported for the MISTIGRI mission with
 spectral configurations that differ from the TRISHNA reference (Jacob
 et al., 2021). Therefore, scanning the spectral space for each of the four
 channel is likely to provide similar results than scanning the spectral space
 using separated pairs of channels.

From an operational viewpoint, the results of the current study do not con-901 tradict those reported by Vidal et al. (2021) who addressed the sensitivity of 902 the Split Window method to channel positions and bandwidths in the context 903 of the TRISHNA mission. Overall, both studies converge toward a unique rec-904 ommendation for the spectral configuration of the TRISHNA mission, with an 905 overall error on surface temperature retrievals of 0.85 K. This is satisfactory in 906 term of mission requirements ( $\leq 1.5$  K), and is of the same order of magnitude 907 as the precision requirements for further applications (0.8 K, cf Seguin et al., 908 1999; Hernandez-Baquero, 2000). 909

Finally, the precision of the TES method in the context of the TRISHNA mission is likely to be improved in the near future thanks to forthcoming studies, as discussed below.

The prior use of the water vapor scaling method (WVS method, Tonooka,
 2005) before atmospheric corrections is expected to considerably reduce
 the RMSE values on TES retrievals of surface temperatures, especially

for wet atmospheric conditions (AWVC  $\geq 2.5 \text{ g.cm}^{-2}$ ). Indeed, Tonooka 916 (2005) reported that the expected RMSE on TES LST retrievals is about 917 0.6 K when using the WVS method. This is sensibly lower than the 918 RMSE we obtain in the current study for inaccurate atmospheric correc-919 tions (0.85 K), as well as twice lower than the mission requirement (1.5 K). 920 2. Research should be conducted on how the TES method could be strength-921 ened by integrating a priori visible and shortwave infrared data, as for 922 instance a first estimate on the emissivity of the considered scene. As 923 mentioned in  $\S$  2.1, recent research has shown that such an enhancement 924 could increase the accuracy on TES retrievals for dry atmospheres (Zheng 925 et al., 2019). However, such improvements need to be carefully imple-926 mented, since it could introduce spatial discontinuities in the temperature 927 retrievals, for instance between vegetated and arid area, as observed with 928 the original TES threshold classifier (Gustafson et al., 2006). 929

From a methodological perspective, we propose an optimization method 930 based on the sliding of channels within respective spectral ranges. This method 931 allows the finding of minimum values of RMSEs in surface temperature and 932 emissivity, within the space defined by the aforementioned spectral ranges. Pro-933 vided adequate computational capabilities are available, it can be used as such 934 for future multispectral and hyperspectral sensors to find optimal channel posi-935 tions with respect to any inversion method, either within a given spectral domain 936 among solar and thermal infrared ones, or across both simultaneously. Indeed, 937 due to computational limitations in the current study, we consider the sliding 938 of TRISHNA channels on a pair-basis rather than considering the simultaneous 939 sliding of all four channels. We expect that future improvements of computa-940 tional capacities will allow more comprehensive studies, such as conducted by 941 Vidal et al. (2021) for the SW channels of the TRISHNA instrument. 942

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# 946 References

- Barsi, J.A., Barker, J.L., Schott, J.R., 2003. An atmospheric correction parameter calculator for a single thermal band earth-sensing instrument, in: IGARSS
  2003. 2003 IEEE International Geoscience and Remote Sensing Symposium.
- <sup>949</sup> 2003. 2003 IEEE International Geoscience and Remote Sensing Symposiu
- Proceedings (IEEE Cat. No. 03CH37477), IEEE. pp. 3014–3016.
- Bayat, B., van der Tol, C., Verhoef, W., 2018. Integrating satellite optical
  and thermal infrared observations for improving daily ecosystem functioning
  estimations during a drought episode. Remote sensing of environment 209,
  375–394.
- Berk, A., Anderson, G.P., 2008. Impact of modtran® 5.1 on atmospheric
  compensation, in: IGARSS 2008-2008 IEEE International Geoscience and
  Remote Sensing Symposium, IEEE. pp. III–127.
- Bigeard, G., Coudert, B., Chirouze, J., Er-Raki, S., Boulet, G., Ceschia, E.,
  Jarlan, L., 2019. Ability of a soil-vegetation-atmosphere transfer model and
  a two-source energy balance model to predict evapotranspiration for several
  crops and climate conditions. Hydrology and Earth System Sciences 23, 5033–
  5058.
- <sup>963</sup> Cao, B., Guo, M., Fan, W., Xu, X., Peng, J., Ren, H., Du, Y., Li, H., Bian,
  <sup>964</sup> Z., Hu, T., et al., 2018. A new directional canopy emissivity model based on
  <sup>965</sup> spectral invariants. IEEE Transactions on Geoscience and Remote Sensing
  <sup>966</sup> 56, 6911–6926.
- <sup>967</sup> Carlson, T.N., Petropoulos, G.P., 2019. A new method for estimating of evapo-
- <sup>968</sup> transpiration and surface soil moisture from optical and thermal infrared mea-

- <sup>969</sup> surements: The simplified triangle. International Journal of Remote Sensing
  <sup>970</sup> 40, 7716–7729.
- Caselles, V., Rubio, E., Coll, C., Valor, E., 1998. Thermal band selection for
  the prism instrument: 3. optimal band configurations. Journal of Geophysical
  Research: Atmospheres 103, 17057–17067.
- <sup>974</sup> Chevallier, F., Chédin, A., Chéruy, F., Morcrette, J.J., 2000. Tigr-like
  <sup>975</sup> atmospheric-profile databases for accurate radiative-flux computation. Quar<sup>976</sup> terly Journal of the Royal Meteorological Society 126, 777–785.
- <sup>977</sup> Coll, C., Caselles, V., Valor, E., Niclòs, R., Sánchez, J.M., Galve, J.M., Mira,
  <sup>978</sup> M., 2007. Temperature and emissivity separation from aster data for low
  <sup>979</sup> spectral contrast surfaces. Remote sensing of environment 110, 162–175.
- <sup>980</sup> Coll, C., Caselles, V., Valor, E., Rubio, E., 2003. Validation of temperature<sup>981</sup> emissivity separation and split-window methods from tims data and ground
  <sup>982</sup> measurements. Remote Sensing of Environment 85, 232–242.
- <sup>983</sup> Courault, D., Jacob, F., Benoit, V., Weiss, M., Marloie, O., Hanocq, J.F., Fillol,
  <sup>984</sup> E., Olioso, A., Dedieu, G., Gouaux, P., et al., 2009. Influence of agricultural
  <sup>985</sup> practices on micrometerological spatial variations at local and regional scales.
  <sup>986</sup> International Journal of Remote Sensing 30, 1183–1205.
- Dash, P., Göttsche, F.M., Olesen, F.S., Fischer, H., 2002. Land surface temperature and emissivity estimation from passive sensor data: Theory and
  practice-current trends. International Journal of remote sensing 23, 2563–
  2594.
- French, A., Schmugge, T., Ritchie, J., Hsu, A., Jacob, F., Ogawa, K., 2008.
  Detecting land cover change at the jornada experimental range, new mexico
  with aster emissivities. Remote Sensing of Environment 112, 1730–1748.
- French, A.N., Jacob, F., Anderson, M.C., Kustas, W.P., Timmermans, W.,
  Gieske, A., Su, Z., Su, H., McCabe, M.F., Li, F., et al., 2005. Surface energy

- <sup>996</sup> fluxes with the advanced spaceborne thermal emission and reflection radiome-
- ter (aster) at the iowa 2002 smacex site (usa). Remote sensing of environment
  998 99, 55–65.
- Galleguillos, M., Jacob, F., Prévot, L., French, A., Lagacherie, P., 2011. Comparison of two temperature differencing methods to estimate daily evapotranspiration over a mediterranean vineyard watershed from aster data. Remote
  Sensing of Environment 115, 1326–1340.
- Galleguillos, M., Jacob, F., Prévot, L., Lagacherie, P., Liang, S., 2010. Mapping
  daily evapotranspiration over a mediterranean vineyard watershed. IEEE
  Geoscience and Remote Sensing Letters 8, 168–172.
- Gillespie, A., Rokugawa, S., Matsunaga, T., Cothern, J.S., Hook, S., Kahle,
  A.B., 1998. A temperature and emissivity separation algorithm for advanced
  spaceborne thermal emission and reflection radiometer (aster) images. IEEE
  Transactions on Geoscience And Remote Sensing 36, 1113–1126.
- Gillespie, A.R., Rokugawa, S., Hook, S.J., Matsunaga, T., Kahle, A.B., 1999.
  Temperature/emissivity separation algorithm theoretical basis document, version 2.4. ATBD contract NAS5-31372, NASA .
- Girouard, G., Bannari, A., El Harti, A., Desrochers, A., 2004. Validated spectral
  angle mapper algorithm for geological mapping: comparative study between
  quickbird and landsat-tm, in: XXth ISPRS congress, geo-imagery bridging
  continents, Istanbul, Turkey, pp. 12–23.
- Gómez, M., Olioso, A., Sobrino, J., Jacob, F., 2005. Retrieval of evapotranspiration over the alpilles/reseda experimental site using airborne polder sensor and a thermal camera. Remote Sensing of Environment 96, 399–408.
- <sup>1020</sup> Göttsche, F.M., Hulley, G.C., 2012. Validation of six satellite-retrieved land
- <sup>1021</sup> surface emissivity products over two land cover types in a hyper-arid region.
- <sup>1022</sup> Remote Sensing of Environment 124, 149–158.

<sup>1023</sup> Grigsby, S.P., Hulley, G.C., Roberts, D.A., Scheele, C., Ustin, S.L., Alsina,

<sup>1024</sup> M.M., 2015. Improved surface temperature estimates with master/aviris sen-

- <sup>1025</sup> sor fusion. Remote Sensing of Environment 167, 53–63.
- Guillevic, P., Gastellu-Etchegorry, J., Demarty, J., Prévot, L., 2003. Thermal
   infrared radiative transfer within three-dimensional vegetation covers. Journal
   of Geophysical Research: Atmospheres 108.
- Gustafson, W.T., Gillespie, A.R., Yamada, G.J., 2006. Revisions to the aster
  temperature/emissivity separation algorithm, in: 2nd International Symposium on Recent Advances in Quantitative Remote Sensing, Global Change
  Unit, University of Valencia Torrent, Spain. pp. –.
- Hernandez-Baquero, E.D., 2000. Characterization of the Earth's surface and
  atmosphere from multispectral and hyperspectral thermal imagery. Technical
  Report. Air Force Institute of Technology Wright-Patterson, AFB OH, School
  Of Engineering.
- Hulley, G., Hook, S., Baldridge, A., 2008. Aster land surface emissivity database
  of california and nevada. Geophysical Research Letters 35.
- Hulley, G., Hook, S., Fisher, J., Lee, C., 2017. Ecostress, a nasa earth-ventures
  instrument for studying links between the water cycle and plant health over
- $_{1041}$   $\,$  the diurnal cycle, in: IGARSS 2017-2017 IEEE International Geoscience and
- <sup>1042</sup> Remote Sensing Symposium, IEEE. pp. 5494–5496.
- Hulley, G.C., Hook, S.J., 2010. Generating consistent land surface temperature
  and emissivity products between aster and modis data for earth science research. IEEE Transactions on Geoscience and Remote Sensing 49, 1304–1315.
- Hulley, G.C., Hook, S.J., Baldridge, A.M., 2010. Investigating the effects of
  soil moisture on thermal infrared land surface temperature and emissivity
  using satellite retrievals and laboratory measurements. Remote Sensing of
  Environment 114, 1480–1493.

- Hulley, G.C., Hughes, C.G., Hook, S.J., 2012. Quantifying uncertainties in land
   surface temperature and emissivity retrievals from aster and modis thermal
   informal data. Learning of Geometrical Research. Atmospheres 117
- <sup>1052</sup> infrared data. Journal of Geophysical Research: Atmospheres 117.
- Inoue, Y., Olioso, A., Choi, W., 2004. Dynamic change of co2 flux over bare
   soil field and its relationship with remotely sensed surface temperature. In ternational Journal of Remote Sensing 25, 1881–1892.
- Jacob, F., Gu, X., Hanocq, J.F., Tallet, N., Baret, F., 2003. Atmospheric
  corrections of single broadband channel and multidirectional airborne thermal
  infrared data: application to the reseda experiment. International Journal of
  Remote Sensing 24, 3269–3290.
- Jacob, F., Lesaignoux, A., Olioso, A., Weiss, M., Caillault, K., Jacquemoud, S.,
  Nerry, F., French, A., Schmugge, T., Briottet, X., et al., 2017. Reassessment
  of the temperature-emissivity separation from multispectral thermal infrared
  data: Introducing the impact of vegetation canopy by simulating the cavity
  effect with the sail-thermique model. Remote Sensing of Environment 198,
  160–172.
- Jacob, F., Olioso, A., Gu, X.F., Su, Z., Seguin, B., 2002. Mapping surface
  fluxes using airborne visible, near infrared, thermal infrared remote sensing
  data and a spatialized surface energy balance model. Agronomie 22, 669–680.
- Jacob, F., Petitcolin, F., Schmugge, T., Vermote, E., French, A., Ogawa, K.,
  2004. Comparison of land surface emissivity and radiometric temperature
  derived from modis and aster sensors. Remote Sensing of Environment 90,
  137–152.
- Jacob, F., Schmugge, T., Olioso, A., French, A., Courault, D., Ogawa, K.,
  Petitcolin, F., Chehbouni, G., Pinheiro, A., Privette, J., 2008. Modeling and
  inversion in thermal infrared remote sensing over vegetated land surfaces, in:
  Advances in Land Remote Sensing. Springer, pp. 245–291.

- Jacob, F., Vidal, T.H., Lesaignoux, A., Olioso, A., Weiss, M., Nerry, F., Jacquemoud, S., Gamet, P., Caillault, K., Labarre, L., et al., 2021. A simulationbased error budget of the tes method for the design of the spectral configuration of the micro-bolometer-based mistigri thermal infrared sensor. IEEE
  Transactions on Geoscience and Remote Sensing .
- Jiménez-Muñoz, J., Sobrino, J., 2006. Error sources on the land surface tem perature retrieved from thermal infrared single channel remote sensing data.
   International Journal of Remote Sensing 27, 999–1014.
- Jiménez-Muñoz, J.C., Sobrino, J.A., Mattar, C., Hulley, G., Göttsche, F.M.,
   2014. Temperature and emissivity separation from msg/seviri data. IEEE
   Transactions on Geoscience and Remote Sensing 52, 5937–5951.
- Koetz, B., Bastiaanssen, W., Berger, M., Defourney, P., Del Bello, U., Drusch,
  M., Drinkwater, M., Duca, R., Fernandez, V., Ghent, D., et al., 2018. High
  spatio-temporal resolution land surface temperature mission-a copernicus candidate mission in support of agricultural monitoring, in: IGARSS 2018-2018
  IEEE International Geoscience and Remote Sensing Symposium, IEEE. pp.
  8160-8162.
- Lagouarde, J.P., Bach, M., Sobrino, J.A., Boulet, G., Briottet, X., Cherchali,
  S., Coudert, B., Dadou, I., Dedieu, G., Gamet, P., et al., 2013. The mistigri thermal infrared project: scientific objectives and mission specifications.
  International Journal of Remote Sensing 34, 3437–3466.
- Lagouarde, J.P., Bhattacharya, B., Crébassol, P., Gamet, P., Adlakha, D.,
  Murthy, C., Singh, S., Mishra, M., Nigam, R., Raju, P., et al., 2019.
  Indo-french high-resolution thermal infrared space mission for earth natural resources assessment and monitoring-concept and definition of trishna, in:
  ISPRS-GEOGLAM-ISRS Joint International Workshop on "Earth Observa-
- tions for Agricultural Monitoring", p. 403.
- 1104 Lagouarde, J.P., Bhattacharya, B., Crébassol, P., Gamet, P., Babu, S., Boulet,
- G., Briottet, X., Buddhiraju, K.M., Cherchali, S., Dadou, I., et al., 2018. The

- <sup>1106</sup> indian-french trishna mission: Earth observation in the thermal infrared with
- high spatio-temporal resolution, in: IGARSS 2018-2018 IEEE International

- 1109 Li, Z.L., Tang, B.H., Wu, H., Ren, H., Yan, G., Wan, Z., Trigo, I.F., Sobrino,
- J.A., 2013. Satellite-derived land surface temperature: Current status and perspectives. Remote Sensing of Environment 131, 14–37.
- Louchart, X., Voltz, M., 2007. Aging effects on the availability of herbicides to runoff transfer. Environmental science & technology 41, 1137–1144.
- Malenovský, Z., Rott, H., Cihlar, J., Schaepman, M.E., García-Santos, G., Fernandes, R., Berger, M., 2012. Sentinels for science: Potential of sentinel-1,-2, and-3 missions for scientific observations of ocean, cryosphere, and land. Remote Sensing of Environment 120, 91–101.
- Matsunaga, T., 1994. A temperature-emissivity separation method using an empirical relationship between the mean, the maximum, and the minimum of the thermal infrared emissivity spectrum. Journal of the Remote Sensing Society of Japan 14, 230–241.
- Mira, M., Olioso, A., Gallego-Elvira, B., Courault, D., Garrigues, S., Marloie,
  O., Hagolle, O., Guillevic, P., Boulet, G., 2016. Uncertainty assessment of
  surface net radiation derived from landsat images. Remote sensing of Environment 175, 251–270.
- Mira, M., Schmugge, T.J., Valor, E., Caselles, V., Coll, C., 2009. Comparison
  of thermal infrared emissivities retrieved with the two-lid box and the tes
  methods with laboratory spectra. IEEE transactions on geoscience and remote
  sensing 47, 1012–1021.
- Montes, C., Jacob, F., 2017. Comparing landsat-7 etm+ and aster imageries to estimate daily evapotranspiration within a mediterranean vineyard watershed. IEEE Geoscience and Remote Sensing Letters 14, 459–463.

<sup>1108</sup> Geoscience and Remote Sensing Symposium, IEEE. pp. 4078–4081.

- 1133 Montes, C., Lhomme, J.P., Demarty, J., Prévot, L., Jacob, F., 2014. A three-
- source svat modeling of evaporation: Application to the seasonal dynamics of
- a grassed vineyard. Agricultural and forest meteorology 191, 64–80.
- Murphy, R.E., 2006. The npoess preparatory project, in: Earth Science Satellite
  Remote Sensing. Springer, pp. 182–198.
- Neinavaz, E., Skidmore, A.K., Darvishzadeh, R., 2020. Effects of prediction
  accuracy of the proportion of vegetation cover on land surface emissivity
  and temperature using the ndvi threshold method. International Journal of
  Applied Earth Observation and Geoinformation 85, 101984.
- Ogawa, K., Schmugge, T., Jacob, F., French, A., 2002. Estimation of broadband
  land surface emissivity from multi-spectral thermal infrared remote sensing.
  Agronomie 22, 695–696.
- Ogawa, K., Schmugge, T., Jacob, F., French, A., 2003. Estimation of land surface window (8–12  $\mu$ m) emissivity from multi-spectral thermal infrared remote sensing—a case study in a part of sahara desert. Geophysical Research Letters 30.
- Olioso, A., 1995. Simulating the relationship between thermal emissivity and
  the normalized difference vegetation index. International Journal of Remote
  Sensing 16, 3211–3216.
- Olioso, A., Inoue, Y., Ortega-Farias, S., Demarty, J., Wigneron, J.P., Braud,
  I., Jacob, F., Lecharpentier, P., Ottlé, C., Calvet, J.C., et al., 2005. Future
  directions for advanced evapotranspiration modeling: Assimilation of remote
  sensing data into crop simulation models and svat models. Irrigation and
  Drainage Systems 19, 377–412.
- <sup>1157</sup> Olioso, A., Jacob, F., Weiss, M., 2018. First evaluation of land surface emissiv-<sup>1158</sup> ity spectra simulated with the sail-thermique model, in: IGARSS 2018-2018
- IEEE International Geoscience and Remote Sensing Symposium, IEEE. pp.3951–3954.

- Palluconi, F., Hoover, G., Alley, R., Jentoft-Nilsen, M., Thompson, T., 1999.
- An atmospheric correction method for aster thermal radiometry over land.
  Algorithm Theoretical Basis Document .
- Pardo, N., Sánchez, M.L., Timmermans, J., Su, Z., Pérez, I.A., García, M.A.,
  2014. Sebs validation in a spanish rotating crop. Agricultural and Forest
  Meteorology 195, 132–142.
- Poutier, L., Miesch, C., Lenot, X., Achard, V., Boucher, Y., 2002. Comanche
  and cochise: two reciprocal atmospheric codes for hyperspectral remote sensing, in: 2002 AVIRIS Earth Science and Applications Workshop Proceedings,
  pp. 1059–0889.
- Realmuto, V., 1990. Separating the effects of temperature and emissivity: Emissivity spectrum normalization, in: Proc. of the Second Thermal Infrared Multispectral Scanner (TIMS) Workshop, Jet Propulsion Lab., 1990, pp. 31–35.
- Ren, H., Liu, R., Yan, G., Li, Z.L., Qin, Q., Liu, Q., Nerry, F., 2015. Performance evaluation of four directional emissivity analytical models with thermal sail model and airborne images. Optics express 23, A346–A360.
- Richter, R., Coll, C., 2002. Bandpass-resampling effects for the retrieval of
  surface emissivity. Applied Optics 41, 3523–3529.
- Schmugge, T., French, A., Ritchie, J.C., Rango, A., Pelgrum, H., 2002. Temperature and emissivity separation from multispectral thermal infrared observations. Remote Sensing of Environment 79, 189–198.
- Schmugge, T., Hook, S., Coll, C., 1998. Recovering surface temperature and
  emissivity from thermal infrared multispectral data. Remote Sensing of Environment 65, 121–131.
- 1185 Schmugge, T., Ogawa, K., Jacob, F., French, A., Hsu, A., Ritchie, J., Rango,
- A., 2003. Validation of emissivity estimates from aster data, in: International
- <sup>1187</sup> Geoscience and Remote Sensing Symposium, pp. III–1873.

- Schröder, W., Schmidt, G., Hasenclever, J., 2006. Geostatistical analysis of
  data on air temperature and plant phenology from baden-württemberg (germany) as a basis for regional scaled models of climate change. Environmental
  Monitoring and Assessment 120, 27–43.
- Seguin, B., Becker, F., Phulpin, T., Gu, X., Guyot, G., Kerr, Y., King, C.,
  Lagouarde, J., Ottlé, C., Stoll, M., Tabbagh, A., Vidal, A., 1999. Irsute: a
  minisatellite project for land surface heat flux estimation from field to regional
  scale. Remote Sensing of Environment 68, 357–369.
- Sobrino, J., Caselles, V., 1990. Thermal infrared radiance model for interpreting the directional radiometric temperature of a vegetative surface. Remote
  Sensing of Environment 33, 193–199.
- Sobrino, J., Lagouarde, J., Boulet, G., Briottet, X., Cherchali, S., Coudert, B.,
  Dadou, I., Dedieu, G., Gillespie, A., Hagolle, O., et al., 2010. Overview of the
  thermal infrared explorer (tirex) mission. Proc. Abstract Book 3rd Recent
  Adv. Quantitative Remote Sens , 213–214.
- <sup>1203</sup> Sobrino, J.A., Del Frate, F., Drusch, M., Jiménez-Muñoz, J.C., Manunta, P.,

Regan, A., 2016. Review of thermal infrared applications and requirements for

<sup>1205</sup> future high-resolution sensors. IEEE Transactions on Geoscience and Remote <sup>1206</sup> Sensing 54, 2963–2972.

1204

- Sobrino, J.A., Jiménez-Muñoz, J.C., 2014. Minimum configuration of thermal
  infrared bands for land surface temperature and emissivity estimation in the
  context of potential future missions. Remote Sensing of Environment 148,
  158–167.
- Sobrino, J.A., Jimenez-Munoz, J.C., Paolini, L., 2004. Land surface temperature
  retrieval from landsat tm 5. Remote Sensing of environment 90, 434–440.
- Sobrino, J.A., Jiménez-Muñoz, J.C., Verhoef, W., 2005. Canopy directional
  emissivity: Comparison between models. Remote Sensing of Environment 99,
  304–314.

- Tonooka, H., 2005. Accurate atmospheric correction of aster thermal infrared
  imagery using the wvs method. IEEE Transactions on Geoscience and Remote
  Sensing 43, 2778–2792.
- Vanhellemont, Q., 2020. Combined land surface emissivity and temperature
  estimation from landsat 8 oli and tirs. ISPRS Journal of Photogrammetry
  and Remote Sensing 166, 390–402.
- Vidal, T.H., Jacob, F., Olioso, A., Gamet, P., 2021. Optimization of instrumental spectral configurations for the split-window method in the context of the
  trishna mission. IEEE Transactions on Geoscience and Remote Sensing .
- Vinukollu, R.K., Wood, E.F., Ferguson, C.R., Fisher, J.B., 2011. Global estimates of evapotranspiration for climate studies using multi-sensor remote
  sensing data: Evaluation of three process-based approaches. Remote Sensing
  of Environment 115, 801–823.
- Weiss, M., Jacob, F., Duveiller, G., 2020. Remote sensing for agricultural applications: A meta-review. Remote Sensing of Environment 236, 111402.
- Yan, G., Jiao, Z.H., Wang, T., Mu, X., 2020. Modeling surface longwave radiation over high-relief terrain. Remote Sensing of Environment 237, 111556.
- <sup>1233</sup> Zheng, X., Li, Z.L., Nerry, F., Zhang, X., 2019. A new thermal infrared channel
  <sup>1234</sup> configuration for accurate land surface temperature retrieval from satellite
  <sup>1235</sup> data. Remote Sensing of Environment 231, 111216.