









High temperature in-situ emissivity study of basaltic dry magmas

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Context

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Aufaristama et al. (2018)



- Volcanic eruptions are regularly monitored by remote sensing (RS).
- The accuracy of these RS techniques depends on crucial parameters such as spectral emissivity (ε).
- In litterature, ε is mostly considered constant for a given rock family or is estimated indirectly from Reflectance data (Harris 2013); for example $\varepsilon = 1$ or 0.95 in the case of basalts.
- To date, only a few studies have measured laboratory in-situ ε at relevant magmatic T: Lee et al. (2013), Rogic et al. (2019a, b), Ramsey et al. (2019), and Thompson and Ramsey (2020)
 - ε is measured in limited spectral range: 5-16 μm (2000-625 cm⁻¹).
 - Most data are acquired in limited temperature range: i.e. at T≤ glass transition. Only in Lee et al. (2013), T reaches 1400 °C.
 - Results show that ε decreases with T in TIR range (8-14 μ m)

→ Emissivity of melt (ε_{hot}) < Emissivity of crystallised lava (ε_{crust}).



• These determined ε values do not allow to obtain reliable modeling results \rightarrow overestimation of radiant flux measured by RS and modelled distance-to-run of lava.

To solve this issue: high temperature and broad spectral range ε data are necessary!

This work is aimed at determining the in-situ Emissivity response of three basalts over a range of magmatic temperatures (up to 1300 K) in the TIR-MIR-SWIR spectral regions (500-8 000 cm⁻¹).

In-situ emissivity measurements (De Sousa Meneses et al. 2015)

Apparatus developped at CEMHTI (Orléans, France):

- two FTIR spectrometers (Bruker 8ov and 7o) + CO2 Coherent K500 laser + sample chamber
 - Sample chamber: turntable with sample holder + Black Body furnace (Pyrox PY8)
- Wide spectral range 20-18 000 cm⁻¹, NIR-MIR-TIR
- High Temperature range from ambient to 2000 °C
- Solid cylindrical glass samples of 1.5 x 10 ϕ mm
- Emissivity is determined using direct method (see Eq. enclosed), method validated for phonolites in Li PhD thesis (2018).
- Temperature is calculated via the Christiansen wavenumber (i.e. where ε =1) with uncertainty ~1 %).
 Bruker Vertex 70 (FT-IR) Bruker Vertex 80v (FT-IR)



	Spectrometer	Spectral range (cm ⁻¹)	Beam splitter	Detector
•	Vertex 80v	[20, 600]	Multilayer	Bolometer
-	Vertex 80v	[380, 5000]	Ge/KBr	DLaTGS/KBr
	Vertex 70	[380, 5000]	Ge/KBr	DLaTGS/KBr
	Vertex 70	[3800, 10,000]	Vis/CaF ₂	InGaAs
	Vertex 70	[9000, 18,000]	Vis/CaF ₂	Si-diode

$$E(\sigma, T, \theta) = \frac{L(\sigma, T, \theta)}{L^{0}(\sigma, T)}$$

Glass Samples

- **Basalt: MORB** (Juan de Fuca, provided by Prouteau et al. 2001)
 - 48.8 wt. % SiO₂ + 3.0 wt. % Alkali
- **Basalt: 2014-2015 Holuhraun** (Bardarbunga, provided by U. Iceland-A. Hoskuldson)
 - 46.9 wt. % SiO2 + 2.6 wt. % Alkali
- Basanite: 1704-1705 Fasnia (Tenerife Canary Island)
 - 43.2 wt. % SiO2 + 5.1 wt. % Alkali

Textural, chemical and structural characterisation before and after IR measurement: SEM, EMPA and µRaman spectroscopy.



Holuhraun

2014-2015,

volcano

Bardarbunga





16

14

10

8

2

0

% 12

wt.

 $Na_2O + K_2O$



• ε increases in TIR range from 0.80 to 0.85 at 10 μm.

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In MIR and SWIR; ε shows very high, constant values
 (~0.95), close to a black body behavior.

- Absorption feature (AF) \rightarrow vibrations of silicate network.
- The Christiansen feature (CF, around 1230 cm⁻¹) shifts towards lower wavenumber.

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Holuhraun 2014-2015, Bardarbunga volcano (Iceland)





 ϵ increases in TIR range from 0.80 to 0.84 at 10 μ m.

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- ε almost constant (~0.95) in MIR and SWIR until 1173 K. With further T increase, ε slightly decreases to 0.90. ٢
- Modifications of Absorption feature (AF) at T> 1173 K and shift of Christiansen feature (CF, around 1215 cm⁻¹) to lower wavenumber. \rightarrow structural reorganization of glass network. 6



 ϵ increases in TIR range from 0.80 to 0.84 at 10 μ m.

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- ε almost constant (~0.95) in MIR and SWIR until 1069 K, With further T increase, ε slightly decreases to 0.91.
- Modifications of Absorption feature (AF) at T> 1069 K and shift of Christiansen feature (CF, around 1210 cm⁻¹) to lower wavenumber → structural reorganization of glass network.





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Litterature

Our study

Overall ε increase with T in TIR in all basalt samples.	Overall ε decrease with T in TIR in all studies (Lee et al. (2013), Rogic et al. (2019a, b), Ramsey et al. (2019), and Thompson and Ramsey (2020)).
ε shows almost a blackbody behavior in MIR and SWIR.	No data in MIR and SWIR.
Absorption feature (8.5-12 μ m): vibration variations of Si/Al/Fe silicate tetrahedra in Restsrahlen bands = Rearrangement in the disordered glass structure at highest T.	Absorption feature between 8.5-12 μm.
Overall ε values remains above 0.80 within all conditions applied. In TIR: Emissivity of melt (ε_{hot}) > Emissivity of crystallised lava (ε_{crust}).	 ε values are variable: Ramsey et al. (2019): ε_{hot}(0.60)< ε_{crust}(0.95) Lee et al. (2013): ε from 0.50 to anomalous values >1 between 6-16 μm. Rogic et al. (2019a, b): ε above 0.90
Basalts and basanite present similar ϵ behavior.	Rogic et al. (2019a, b): ɛ behavior of Etna basalts differs from the basalts presented here.
	Lee et al. (2013): large ε variations at T>1200 K.





Conclusion

- Emissivity is dependent of:
 - Composition
 - Wavelength
 - Temperature
- Emissivity increases with temperature in TIR, and remains constant, close to blackbody behavior, in MIR and SWIR until 1069 K for Fasnia and 1091 K for Bardarbunga. Then decreases in T range where structural modifications were detected.
- Basanite and basalts have similar emission behaviours, similar to Fe-rich phonolites (Li PhD (2018), Li et al. in prep).
- In-situ Laboratory ε must be compared to field data.

Improvements on E uncertainty will enhance our capacity to better constrain kinetic T and surface radiance from RS and accurately model lava flows with appropriate rheological evolution; information that are crucial to improve hazard assessment in volcanic systems.



Thank you for reading me...



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25/10/19 Piton

de la Fournaise

(credit: R. Bouhet)