2	Analytical solution and Bayesian inference for interference pumping tests in fractal du-
3	al-porosity media
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## 22 Abstract

23 A new analytical solution is developed for interference hydraulic pumping tests in fractured 24 porous media using the dual-porosity concept. Heterogeneous fractured reservoirs are considered with hydrodynamic parameters assumed to follow power-law functions in radial dis-25 26 tance. The developed analytical solution is verified by comparison against a finite volume 27 numerical solution. The comparison shows that the numerical solution converges toward the 28 analytical one when the size of the time step decreases. The applicability of the fractal dual-29 porosity model is then assessed by investigating the identifiability of the parameters from a 30 synthetic interference pumping test with a set of noisy data using Bayesian parameter infer-31 ence. The results show that if the storage coefficient in the matrix is fixed, the rest of the pa-32 rameters can be appropriately inferred, otherwise, the identification of the parameters is faced 33 with convergence problems because of equifinality issues.

34 Keywords: Fractured porous media, interference pumping test, dual-porosity, fractal media,
35 Bayesian inversion.

#### 37 1 INTRODUCTION

38 Simulation of interference pumping tests in porous media allows testing different prospective 39 exploitation scenarios at low costs. These tests consist in pumping a well at a controlled flow 40 rate and measuring the water level response (drawdown) in one or more surrounding observa-41 tion wells. Pumping tests in fractured porous media have been discussed by several authors 42 (see [1-8], among others). Fractured media are often addressed using the dual-porosity model 43 suggested by Barenblatt et al. [9]. The main idea of this model is to consider two overlapping 44 continua: (i) the fractures which carry the major part of water to the pumping well and (ii) the 45 low permeability rock matrix which feeds fluid to the fractures. At each point of the domain, 46 water flow between the fractures and matrix are coupled via a linear exchange term propor-47 tional to the difference in hydraulic head between the two continua. As stated by De Smedt 48 [8], the simulation of interference pumping tests is often based on numerical models which 49 cannot insure full control of numerical accuracy [10]. Therefore, analytical solutions are high-50 ly sought-after because (i) they provide valuable insights into fundamental physics, as they 51 are free of numerical errors and (*ii*), although limited to simplified situations; they allow vali-52 dation of more complex numerical codes.

In the context of interference pumping tests in fractured porous media, an analytical solution 53 has been developed for homogenous aquifers by De Smedt [8]. Furthermore, Delay et al. [7] 54 55 developed a numerical solution using the dual-medium approach to describe radial convergent 56 flow in heterogeneous fractal media where power laws in space are used for the hydrodynam-57 ic parameters. As a consequence, parameter values decrease with the distance between the pumped well and the observed ones. This type of scaling laws inherits from theoretical works 58 on fractals [11] and has been used by several authors for single porous media ([12-15], among 59 60 others). Hence, the main objective of this paper is to develop, for the first time, an analytical 61 solution for the mathematical model describing interference pumping tests in heterogeneous 62 fractal dual media. Applicability of the model is then assessed by investigating identifiability

63 of the different parameters from a synthetic pumping test with a set of noisy data using Statis-

64 tical Calibration (SC) [16]. The latter means refining the prior distribution of uncertain pa-

<sup>65</sup> rameters based on matching simulation outputs with data [17]. In this work, the SC is per-

- 66 formed with the DREAM<sub>(ZS)</sub> software [16] based on the Markov Chain Monte Carlo process
- 67 (MCMC). DREAM<sub>(ZS)</sub> computes multiple sub-chains in parallel to thoroughly explore the
- 68 parameter space. Taking the last 25% of individuals of the MCMC (when the chains have
- 69 converged) yields multiple sets of parameters that adequately fit the model onto observations.
- 70 These sets can then be used to estimate the updated parameter distributions and therefore the
- 71 best estimates of the parameters and the associate confidence intervals.

#### 72 2 MATHEMATICAL MODEL

We assume a horizontal, isotropic, confined and infinite aquifer with a uniform thickness. The pumping well only draws water from the fractures, has a constant rate and fully penetrates the aquifer. Using the dual-porosity concept, the flow in fractured porous media can be described by the following equations [18, 19]:

77 
$$S_{f} \frac{\partial h_{f}}{\partial t} = \nabla \cdot \left( K_{f} \nabla h_{f} \right) + \alpha \left( h_{m} - h_{f} \right)$$
(1)

78 
$$S_m \frac{\partial h_m}{\partial t} = \alpha \left( h_f - h_m \right)$$
(2)

where indexes *m* and *f* refer to the matrix and fracture continua, respectively,  $S[L^{-1}]$  is the specific storage, h[L] is the hydraulic head,  $K_f[LT^{-1}]$  is the hydraulic conductivity of the fractures,  $\alpha[L^{-1}T^{-1}]$  is the exchange rate coefficient between the fractures and the matrix. Following Warren and Root [18], equation (2) assumes pseudo-steady flow in the matrix continuum (term  $K_m$  is removed). Assuming radial symmetry, equations (1) and (2) can be formulated as:

85 
$$S_{f}(r)\frac{\partial h_{f}(r,t)}{\partial t} = \frac{1}{r}\frac{\partial}{\partial r}\left[rK_{f}(r)\frac{\partial h_{f}(r,t)}{\partial r}\right] + \alpha(r)\left[h_{m}(r,t) - h_{f}(r,t)\right]$$
(3)

86 
$$S_m(r)\frac{\partial h_m(r,t)}{\partial t} = \alpha(r) \Big[h_f(r,t) - h_m(r,t)\Big]$$
(4)

87 where r[L] is the radial coordinate with the center of the pumping well as origin.

88 The drawdown in an observation well intercepting the fractures can be expressed as follows

89 
$$S_{f}(r)\frac{\partial H_{f}(r,t)}{\partial t} = \frac{1}{r}\frac{\partial}{\partial r}\left(rK_{f}(r)\frac{\partial H_{f}(r,t)}{\partial r}\right) + \alpha(r)\left[H_{m}(r,t) - H_{f}(r,t)\right]$$
(5)

90 
$$S_m(r)\frac{\partial H_m(r,t)}{\partial t} = \alpha(r) \Big[ H_f(r,t) - H_m(r,t) \Big]$$
(6)

91 where  $H_f = h_0 - h_f$  and  $H_m = h_0 - h_m$  are respectively the drawdown in the fracture and in the 92 rock matrix. Initially, a uniform hydraulic head  $(h_0)$  is assumed in both the fracture and ma-93 trix continua. Therefore, the drawdowns satisfy the following initial conditions:

94 
$$H_f(r,0) = H_m(r,0) = 0$$
 (7)

95 Far away from the pumping well, the drawdowns are assumed to be zero

96 
$$\lim_{r \to \infty} H_f(r,t) = \lim_{r \to \infty} H_m(r,t) = 0$$
(8)

97 The limit where the radius approaches zero corresponds to the well withdrawal rate

98 
$$\lim_{r \to 0} \left[ -2\pi r B K_f(r) \frac{\partial H_f(r,t)}{\partial r} \right] = Q_0, \quad t > 0$$
(9)

99 where *B* is the thickness of the aquifer and  $Q_0$  is the pumping flow rate.

100 In the sequel, we limit our investigation for the semi-analytical solution of equations (5-9) to

101 the case of a confined reservoir of a fractal structure with a single well in an unbounded me-

- 102 dium as described in [7, 12]. Note that more complex fractal approaches have been developed
- 103 in the case of many wells with variable rates [20] and by considering transient flow from
- 104 block to fissure and skin between the fissure and the block [21].
- 105 In this work, following Delay et al. [7], we consider dual porosity media in which the hydro-
- 106 dynamic parameters are power-law functions in space (decreasing functions of the radial dis-
- 107 tance r between the pumping well and the observed one). Moreover, for sake of simplicity,
- 108 we assume the same functions for the exchange rate coefficient  $\alpha$  and for the storage coeffi-
- 109 cients  $S_f$  and  $S_m$ . Hence, the values of the hydraulic parameters vary with respect to radial
- 110 distance according to the following power-law functions:

111  

$$K_{f}(r) = K_{f_{0}}r^{-a}; \quad S_{f}(r) = S_{f_{0}}r^{-b}$$

$$S_{m}(r) = S_{m_{0}}r^{-b}; \quad \alpha(r) = \alpha_{0}r^{-b}$$
(10)

112 where the power-law exponents a and b are assumed to be positive.

#### 113 3 ANALYTICAL SOLUTION

#### 114 **3.1** Analytical solution in the Laplace domain

In this section, we follow the strategy developed by De Smedt [8] to derive the analytical solution of the system formed by equations (5)-(10). The solution is derived using the Laplace transform technique with respect to time variable t. The Laplace transform of a timedependent function y is

119 
$$\overline{y}(p) = L[y(t); t \to p] = \int_{0}^{\infty} y(t)e^{-pt}dt$$
(11)

120 Applying the Laplace transform to (5)-(6) and using (10) leads to

121 
$$pS_{f_0}r^{-b}\overline{H}_f(r,p) = \frac{K_{f_0}}{r}\frac{\partial}{\partial r}\left[r^{1-a}\frac{\partial\overline{H}_f(r,p)}{\partial r}\right] + \alpha_0 r^{-b}\left[\overline{H}_m(r,p) - \overline{H}_f(r,p)\right]$$
(12)

122 
$$pS_{m_0}\overline{H}_m(r,p) = \alpha_0 \Big[\overline{H}_f(r,p) - \overline{H}_m(r,p)\Big]$$
(13)

123 Equation (13) can be rearranged

124 
$$\overline{H}_{m}(r,p) = \frac{\alpha_{0} / S_{m_{0}}}{p + \alpha_{0} / S_{m_{0}}} \overline{H}_{f}(r,p)$$
(14)

125 Substituting (14) into (12) yields

126 
$$r^{2} \frac{\partial^{2} \overline{H}_{f}}{\partial r^{2}} + (1-a)r \frac{\partial \overline{H}_{f}}{\partial r} - \frac{S_{f_{0}}}{K_{f_{0}}} \left[ p + \frac{\left(\alpha_{0} / S_{f_{0}}\right)p}{p + \alpha_{0} / S_{m_{0}}} \right] r^{a-b+2} \overline{H}_{f} = 0$$
(15)

127 Equation (15) is a second order ordinary differential equation; its general solution is of the

128 form (cf. Polyanin and Zaitsev [22], page 159, equation (127))

129 
$$\overline{H}_{f}(r,p) = r^{\frac{a}{2}} \left\{ C_{1}J_{\nu} \left[ \frac{2iS_{f_{0}}^{1/2}r^{n/2}}{nK_{f_{0}}^{1/2}} \sqrt{A(p)} \right] + C_{2}Y_{\nu} \left[ \frac{2iS_{f_{0}}^{1/2}r^{n/2}}{nK_{f_{0}}^{1/2}} \sqrt{A(p)} \right] \right\}$$
(16)

130 where 
$$n = a - b + 2$$
,  $v = \frac{a}{n}$ ,  $A(p) = p + \frac{(\alpha_0 / S_{f_0})p}{p + \alpha_0 / S_{m_0}}$ ,  $J_v(z)$  and  $Y_v(z)$  are the Bessel functions

131 of the first and second kinds, respectively. The parameters  $C_1$  and  $C_2$  are constant that can be

132 calculated using boundary conditions.

134 
$$\lim_{r \to \infty} \bar{H}_f(r, p) = 0 \tag{17}$$

135 and

136 
$$\lim_{r \to 0} \left[ -2\pi \mathbf{B} K_{f_0} r^{1-a} \frac{\partial \overline{H}_f(r, p)}{\partial r} \right] = \frac{Q_0}{p}$$
(18)

137 The solution of (15), taking into account the boundary conditions (17) and (18) is

138 
$$\bar{H}_{f}(r,p) = \frac{2r^{\frac{a}{2}}}{n^{1-\nu}\pi\Gamma(1-\nu)BK_{f_{0}}^{1-\nu/2}S_{f_{0}}^{\nu/2}} \frac{\bar{Q}(p)K_{\nu}\left[\frac{2S_{f_{0}}^{1/2}r^{n/2}}{nK_{f_{0}}^{1/2}}\sqrt{p+\frac{(\alpha_{0}/S_{f_{0}})p}{p+\alpha_{0}/S_{m_{0}}}}\right]}{\left[\sqrt{p+\frac{(\alpha_{0}/S_{f_{0}})p}{p+\alpha_{0}/S_{m_{0}}}}\right]^{\nu}}$$
(19)

139 where  $K_{\nu}$  is the modified Bessel function of the second kind of order  $\nu$ . More details about 140 the derivation of the solution (19) are given in the Appendix.

141

#### 142 **3.2** Analytical solution in the time domain

143 The solution in the time domain  $H_f(r,t)$  is sought using the convolution theorem of the La-144 place transform [23]

145 
$$L\left[\int_{0}^{t} y(\tau, t-\tau) d\tau; t \to p\right] = L\left\{L\left[y(t_{1}, t_{2}); t_{2} \to p\right]; t_{1} \to p\right\}$$
(20)

146 Hence, equation (19) is written as

147 
$$\bar{H}_{f}(r,p) = \bar{H}_{f}(r;p_{1},p_{2}) = \frac{2Q_{0}r^{\frac{a}{2}}}{n^{1-\nu}\pi\Gamma(1-\nu)BK_{f_{0}}^{1-\nu/2}S_{f_{0}}^{\nu/2}} \frac{K_{\nu}\left[\frac{2S_{f_{0}}^{1/2}r^{n/2}}{nK_{f_{0}}^{1/2}}\sqrt{p_{1}+\frac{(\alpha_{0}/S_{f_{0}})p_{2}}{p_{2}+\alpha_{0}/S_{m_{0}}}}\right]^{\nu}}{p_{2}\left[\sqrt{p_{1}+\frac{(\alpha_{0}/S_{f_{0}})p_{2}}{p_{2}+\alpha_{0}/S_{m_{0}}}}\right]^{\nu}}$$
(21)

- 148 where  $p_1 = p_2 = p$  but different indexes are used to indicate that the inverse transformation is
- 149 performed in two steps.
- 150 The different indexes used in (21) indicate that the inverse Laplace transform is developed in
- 151 two steps. In a first step, the inverse Laplace transform with respect to the variable  $p_1$  is ex-
- 152 pressed as follows

153 
$$L^{-1}\left[\bar{H}_{f}\left(r;p_{1},p_{2}\right);p_{1}\rightarrow t_{1}\right] = \frac{n^{\nu}Q_{0}e^{-\frac{S_{f_{0}}r^{n}}{2n^{2}K_{f_{0}}t_{1}}} \cdot W_{\frac{1-\nu}{2},\frac{\nu}{2}}\left(\frac{S_{f_{0}}r^{n}}{n^{2}K_{f_{0}}t_{1}}\right)}{\pi\Gamma(1-\nu)BK_{f_{0}}^{\frac{1-\nu}{2}}S_{f_{0}}^{\frac{1+\nu}{2}}r^{\frac{2-b}{2}}} \cdot \frac{e^{-\frac{(\alpha_{0}/S_{f_{0}})t_{1}p_{2}}{p_{2}+\alpha_{0}/S_{m_{0}}}}}{p_{2}}$$
(22)

154 where  $t_1$  is the time variable obtained by the inverse Laplace transform with respect to the

155 variable  $p_1$  and  $W_{\kappa,\mu}(z)$  is the Whittaker function (Abramowitz and Stegun [24] p. 505).

156 In the second step, the inverse Laplace transform with respect to  $p_2$  is obtained using the in-

157 verse transform of the *J*-function [25] based on the formula  $L^{-1}\left[\frac{1}{p}e^{-\frac{\alpha p}{p+\beta}};p \to t\right] = J(\alpha,\beta t)$ 

158 which yields

159 
$$\bar{H}_{f}(r;p_{1},p_{2}) = \frac{n^{\nu}Q_{0}e^{-\frac{S_{f_{0}}r^{n}}{2n^{2}K_{f_{0}}t_{1}}} \cdot t_{1}^{\frac{\nu-1}{2}} \cdot W_{\frac{1-\nu}{2},\frac{\nu}{2}}\left(\frac{S_{f_{0}}r^{n}}{n^{2}K_{f_{0}}t_{1}}\right) \cdot J\left(\frac{\alpha_{0}t_{1}}{S_{f_{0}}},\frac{\alpha_{0}t_{2}}{S_{m_{0}}}\right)}{2\pi\Gamma(1-\nu)BK_{f_{0}}^{\frac{1-\nu}{2}}S_{f_{0}}^{\frac{1+\nu}{2}}r^{\frac{2-b}{2}}}$$
(23)

160 The Bessel integral *J*-function is defined as follows [8, 26]

161 
$$J(a,b) = 1 - e^{-b} \int_{0}^{a} e^{-\lambda} I_0(2\sqrt{b\lambda}) d\lambda = e^{-a-b} \sum_{n=0}^{\infty} \frac{b^n}{n!} \sum_{m=0}^{n} \frac{a^m}{m!}$$
(24)

162 In view of (20) and (23), the analytical expression of  $H_{f}(r,t)$  is

$$163 \qquad H_{f}(r,t) = \frac{n^{\nu}Q_{0}}{2\pi\Gamma(1-\nu)BK_{f_{0}}^{\frac{1-\nu}{2}}S_{f_{0}}^{\frac{1+\nu}{2}}r^{\frac{2-b}{2}}\int_{0}^{t}\tau^{\frac{\nu-1}{2}} \cdot e^{-\frac{S_{f_{0}}r^{n}}{2n^{2}K_{f_{0}}\tau}} \cdot W_{\frac{1-\nu}{2},\frac{\nu}{2}}\left(\frac{S_{f_{0}}r^{n}}{n^{2}K_{f_{0}}\tau}\right) \cdot J\left(\frac{\alpha_{0}\tau}{S_{f_{0}}},\frac{\alpha_{0}(t-\tau)}{S_{m_{0}}}\right) d\tau$$

164

165 A similar expression can be obtained for  $H_m(r,t)$  using the convolution theorem of the La-

(25)

166 place transform for equation (14) and using equation (19).

167 In the case of homogeneous media (a=b=0), we have  $W_{\frac{1}{2},0}(z) = z^{\frac{1}{2}}e^{-\frac{z}{2}}$  and (25) reduces to

168 the De Smedt [8] solution:

169 
$$H_{f}(r,t) = \frac{Q_{0}}{4\pi B K_{f_{0}}} \int_{0}^{t} \tau^{-1} e^{-\frac{S_{f_{0}}r^{2}}{4K_{f_{0}}\tau}} \cdot J\left(\frac{\alpha_{0}\tau}{S_{f_{0}}}, \frac{\alpha_{0}(t-\tau)}{S_{m_{0}}}\right) d\tau$$
(26)

170

171

#### 173 4 COMPARISON TO NUMERICAL SOLUTIONS

174

175 dual media. For each test, the numerical solution is compared against the corresponding ana-176 lytical solution using (25). The numerical solution of the system (5)-(10) is developed using 177 the standard Finite Volume (FV) method. The domain of size L=10,000 m is discretized using a uniform spatial discretization of size  $\Delta r = 5$  m. A similar solution is obtained using a 178 179 finer spatial discretization of size  $\Delta r = 1$  m which shows the weak sensitivity of the numerical 180 solution to the spatial discretization. The duration of the simulation is  $2 \times 10^6$  s. Three time step sizes are investigated with  $\Delta t = 10^5$  s,  $\Delta t = 5 \times 10^4$  s and  $\Delta t = 10^4$  s. The pumping flow rate 181 is  $Q_0 = 5 \text{ m}^3/\text{h}$ . The initial head in the aquifer is  $h_0 = 100 \text{ m}$  and the aquifer thickness is 182 B = 100 m. The specific storage of the fractures and rock matrix are respectively  $S_{f_0} = 5 \times 10^{-6}$ 183 m<sup>-1</sup> and  $S_{m_0} = 10^{-4}$  m<sup>-1</sup>. The hydraulic conductivity of the fractures is  $K_{f_0} = 10^{-4}$  m/s and the 184 rate of exchange between the matrix and the fractures is  $\alpha_0 = 5 \times 10^{-11} \text{ m}^{-1} \text{ s}^{-1}$ . 185 186 Figure 1 depicts the drawdown in both the fracture and matrix at 10m and 100m. The draw-

We numerically solve two interference pumping tests dealing with homogeneous and fractal

187 down in the fissure starts at early times because storage of the fissures is immediately effec-

188 tive at the start of the pumping test, while storage of the matrix needs much more time to be-

189 come noticeable. A flattening of the drawdown curve in the fissure can be observed at inter-

190 mediate pumping time which is typical for double-porosity behavior because of the transition

191 from fracture to matrix water storage release.

192 The numerical FV solution is compared to the analytical solution (25) in the case of homoge-

193 neous (a = b = 0) and fractal (a = 0.6 and b = 0.4) dual media. The comparison is performed

194 at two observation wells located respectively at distances  $r_1 = 10$  m and  $r_2 = 100$  m from the

195 pumping well. Significant discrepancies can be observed between analytical and numerical

196 solutions when the latter is evaluated using a large time step size (Figure 2). The drawdown in 197 the case of fractal dual media is larger than in homogeneous dual media because of the reduc-198 tion of the hydrodynamic parameters with the distance r from the pumping well. In fact, in the case of fractal dual media, the permeability of the fractures is proportional to  $r^{-0.6}$  where-199 200 as the exchange rate and the specific storage of both the fractures and the matrix are proportional to  $r^{-0.4}$ . Thereby, the discrepancy between analytical and numerical solutions (Figure 201 202 2) is more pronounced with dual fractal media than with homogeneous dual media. For both 203 homogeneous and fractal dual media, the numerical solution converges toward the analytical 204 solution (25) when the size of the time step decreases (Figure 2). Note that the reduction of 205 the permeability of the fractures with respect to the distance r is also responsible, via the 206 Whittaker function in equation (25), of the delay for the stabilisation of the solution in the 207 fractal case (Figure 2 b) compared to the homogeneous case (Figure 2a).

#### 208 5 BAYESIAN INVERSION OF A PUMPING TEST IN FRACTAL DUAL MEDIA

209 In practical applications, most of the hydraulic parameters cannot be measured directly and 210 should be estimated using an inversion procedure. In the case of fractal dual media, the inver-211 sion can be challenging because of model equifinality, implying that different parameter com-212 binations can fit the model responses to data. To assess the applicability of the fractal dual 213 porosity model, the identifiability of the parameters of the analytical solution (25) is assessed 214 using a synthetic interference pumping test with a set of noisy data. Parameter identification is 215 performed in a Bayesian framework where the prior knowledge about the parameters and the 216 observed data are merged to define the joint posterior probability distribution function (pdf) of 217 the parameters. In this work, the pdf is performed using the  $DREAM_{(ZS)}$  software [16] based 218 on the Markov Chain Monte Carlo (MCMC) sampler. DREAM<sub>(ZS)</sub> generates random se-219 quences of parameter sets that asymptotically converge toward the target joint posterior dis-

- tribution [27]. Thus, if the number of runs is sufficiently high, the generated samples allows
- 221 exploring the entire parameter space of the posterior distribution of the parameters and pro-
- 222 vides the pairwise parameter correlations and the uncertainty of model predictions.
- 223 The Bayes theorem states that the probability density function of the model parameters condi-
- 224 tioned onto data can be expressed as:

225 
$$p(\boldsymbol{\xi} | \boldsymbol{y}_{mes}) \propto p(\boldsymbol{y}_{mes} | \boldsymbol{\xi}) p(\boldsymbol{\xi})$$
(27)

where  $p(\boldsymbol{\xi} | \boldsymbol{y}_{mes})$  is the likelihood function measuring how well the model fits the observations  $\boldsymbol{y}_{mes}$ , and  $p(\boldsymbol{\xi})$  is the prior information about the parameter before the observations are made. In this work, a Gaussian distribution defines the likelihood function because the *observations* are simulated and corrupted with Gaussian errors. Hence, the parameter posterior distribution is expressed as:

231 
$$p(\boldsymbol{\xi}/\boldsymbol{y}_{mes}) \propto \sigma_h^{-N_h} \exp\left(-\frac{SS_h(\boldsymbol{\xi})}{2\sigma_h^2}\right)$$
(28)

- where  $SS_h(\xi)$ , are the sums of the squared differences between the observed and modeled drawdowns. For instance,  $SS_h(\xi) = \sum_{k=1}^{Nh} \left( h_{mes}^{(k)} - h_{mod}^{(k)}(\xi) \right)^2$ , which includes the observed  $h_{mes}^{(k)}$ and predicted  $h_{mod}^{(k)}$  drawdowns at time  $t_k$  for the number of observations  $N_h$ .
- 235 A synthetic pumping test is generated by running the analytical solution using the reference
- parameter values given in Table 1 for  $K_{f_0}$ ,  $S_{f_0}$ ,  $\alpha_0$ , a and b. As in Fahs et al. [28], the stor-
- age coefficient in the matrix  $S_{m_0}$  is fixed in this study to be  $10^{-4}$  m<sup>-1</sup>.
- Delay et al [7] showed that results of calibration improved when several drawdown curves at several locations are inverted at once. In the sequel, the analytical solution is run for a long period of  $2 \times 10^7$  s to calculate the drawdown at four observation wells located respectively at  $r_1 = 1 \text{ m}, r_2 = 10 \text{ m}, r_3 = 100 \text{ m}$  and  $r_4 = 500 \text{ m}$  from the pumping well. Four datasets of 2000

values each (corresponding to the drawdown at the observation wells measured each  $10^4$  s) are used as target responses for model inversion.

Gaussian noise with a standard deviation of 0.01m and a mean of zero was added to the target responses. We assume that all hydraulic parameters have uniform prior distributions over the ranges given in Table 1. As recommended in Vrugt *et al.* [29], we consider that the posterior distribution is stationary if the Gelman and Ruban [30] criterion is less than 1.2 and that the chains are not autocorrelated.

249 The MCMC method was terminated after 15000 model runs. The convergence was reached at

250 around 10000 model runs. Figure 3 shows the results of the identification based on the last 251 25% of individuals of the MCMC sampler (when the chains have converged). The "on-252 diagonal" plots in these figures display the inferred parameter distributions, whereas the "off-253 diagonal" plots represent the pairwise correlations in the MCMC sample. If the drawings are 254 independent, non-sloping scatterplots should be observed. However, if a good value of a giv-255 en parameter is conditioned by the value of another parameter, then their pairwise scatterplot 256 should show a narrow sloping stripe. The results of Figure 3 show that the five parameters  $K_{f_0}$ ,  $S_{f_0}$ ,  $\alpha_0$ , a and b are appropriately estimated; they have almost symmetric bell-shaped 257 258 posterior distributions and have strongly narrowed their prior intervals. The scatter plots in 259 this figure show a moderate correlation between the parameters a and b (a correlation factor of 0.61), a high correlation between the parameters  $S_{f_0}$  and b (a correlation factor of 0.79) 260 and a very high correlation between the parameters  $K_{f_0}$  and a (a correlation factor of 0.98). 261 262 The mean and 95% confidence intervals of the samples that adequately fit the model onto observations are reported in Table 2. From this table, we can see that the parameters  $K_{f_0}$ , a 263 and b are well identified with respectively a mean value of  $0.99 \times 10^{-5}$  ms<sup>-1</sup>, 0.299 and 0.2. 264 265 These values are almost identical to the reference values of Table 1. The uncertainty on these parameters is small since the 95% confidence intervals are respectively  $[0.9-1.1]\times10^{-5}$  ms<sup>-1</sup>, [0.28-0.32] and [0.17-0.23]. The parameters  $S_{f_0}$  and  $\alpha_0$  are less well identified. Indeed, although their mean values (respectively  $5.03\times10^{-5}$  m<sup>-1</sup> and  $3.06\times10^{-11}$  m<sup>-1</sup>s<sup>-1</sup>) are very close to the reference values, their posterior uncertainty remains quite large since their confidence intervals are respectively  $[3.2-7.6]\times10^{-5}$  m<sup>-1</sup> and  $[2.1-4.5]\times10^{-11}$  m<sup>-1</sup>s<sup>-1</sup>.

Finally note that the identification procedure does not converge if the storage coefficient in the matrix  $S_{m_0}$  is involved in the inversion procedure (not fixed). In that case, the identification of the six parameters  $K_{f_0}$ ,  $S_{f_0}$ ,  $S_{m_0}$ ,  $\alpha_0$ , *a* and *b* was not possible. Indeed, in this case, the MCMC algorithm was faced to convergence problem because of equifinality issues.

275

#### 276 6 CONCLUSION

277 A new analytical solution for interference hydraulic pumping tests in fractal fractured porous 278 media is developed using the dual-porosity concept and power laws in space for all hydrody-279 namic parameters. The developed analytical solution is compared to a numerical one obtained 280 with the FV method using different time step sizes. Discrepancies are observed between ana-281 lytical and numerical solutions when the latter is used with large time steps. These discrepan-282 cies are more pronounced with dual fractal media than with homogeneous dual media. When 283 the size of the time step decreases, the numerical solution converges toward the analytical one 284 for both homogeneous and fractal dual media.

The applicability of the fractal dual-porosity model was then assessed by investigating identifiability of the hydraulic parameters from a synthetic interference pumping test with a set of noisy data using the MCMC sampler. The results show that if the storage coefficient in the

288	matrix is fixed, the rest of the parameters can be appropriately identified, whereas conver-
289	gence problems can be encountered if all the hydraulic parameters are to be estimated.
290	
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## 297 Appendix: Derivation of the Analytical Solution in the Laplace Domain

298

#### 299 Using the boundary condition at infinity (17) and the following asymptotic expansions for

300 large arguments  $(|z| \rightarrow \infty)$  of  $J_{\nu}$  and  $Y_{\nu}$  [22],  $J_{\nu}(z) = \sqrt{2/\pi z} \{P(\nu, z) \cos \chi - Q(\nu, z) \sin \chi\}$ 

301 and 
$$Y_{\nu}(z) = \sqrt{2/\pi z} \{ P(\nu, z) \cos \chi + Q(\nu, z) \sin \chi \}$$
 for  $|\arg z| < \pi$  with  $\chi = z - \left(\frac{1}{2}\nu + \frac{1}{4}\right)\pi$ ,

302 
$$P(v,z) = \sum_{k=0}^{\infty} (-1)^k \frac{(v,2k)}{(2z)^{2k}}$$
 and  $Q(v,z) = \sum_{k=0}^{\infty} (-1)^k \frac{(v,2k+1)}{(2z)^{2k+1}}$ , equation (16) gives

303  $C_1 + iC_2 = 0$ . Therefore, this latter can be written as

304 
$$\bar{H}_{f}(r,p) = C_{1} r^{\frac{a}{2}} H_{\nu}^{(1)} \left[ \frac{2i S_{f_{0}}^{1/2} r^{n/2}}{n K_{f_{0}}^{1/2}} \sqrt{A(p)} \right]$$
(A.1)

305 where  $H_{\nu}^{(1)}$  is the Hankel function of the first kind defined by  $H_{\nu}^{(1)}(z) = J_{\nu}(z) + iY_{\nu}(z)$  for 306 any complex number z.

307 The boundary condition at the well (18) is then used to find  $C_1$ . Note that for  $\nu > 0$  the

308 Hankel function  $H_{\nu}^{(1)} \sim -\frac{i}{\pi} \Gamma(\nu) \left(\frac{1}{2}z\right)^{-\nu}$  as z approaches zero [22], where  $\Gamma$  is the Gamma

309 function. Using this expansion we can write

$$310 \qquad r^{1-a} \frac{\partial H_f(r,t)}{\partial r} \sim \frac{i^{1-\nu} S_{f_0}^{\nu/2} C_1}{2\pi K_{f_0}^{\nu/2}} \left\{ -r^{-a} n^{\nu} \left[ \sqrt{A(p)} \right]^{-\nu} \left\{ a \Gamma(\nu) - n \Gamma(1+\nu) \right\} + n^{1-\nu} \left[ \sqrt{A(p)} \right]^{\nu} \Gamma(1-\nu) \right\}$$

(A.2)

312 as *r* appraches zero. Using the property of the Gamma function  $\Gamma(1+\nu) = \nu \Gamma(\nu)$  for  $\nu > 0$ ,

313 the first term of right hand side of equation (A.2) vanishes since nv = a. Substituting (A.2)

314 into (18) we get 
$$C_1 = -\frac{\bar{Q}(p)}{i^{1-\nu}n^{1-\nu}\Gamma(1-\nu)BK_{f_0}^{1-\nu/2}S_{f_0}^{\nu/2}\left[\sqrt{A(p)}\right]^{\nu}}$$

315 Finally, inserting the value of  $C_1$  into (A.1) and using the relationship 316  $K_{\nu}(z) = \frac{1}{2}\pi i e^{\frac{1}{2}\nu\pi i} H_{\nu}^{(1)}\left(ze^{\frac{1}{2}\pi i}\right)$  for  $-\pi < \arg z \le \frac{1}{2}\pi$ , where  $K_{\nu}$  is the modified Bessel func-

317 tion of the second kind of order v, we get the solution in the Laplace domain as defined by 318 (19).

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410 neous and (*ii*) fractal dual media.

411 Figure 3: MCMC solutions of the interference pumping test in fractal dual media. The diago-

412 nal plots represent the inferred posterior probability distribution of model parameters. The off-

- 413 diagonal scatterplots represent the value of a parameter  $\alpha$  versus that of a parameter  $\beta$  in the
- 414 best solutions inferred by MCMC.

415

416

# **Table 1**

## 

Parameters	Lower bounds	Upper bounds	Reference values
$K_{f_0}  [{\rm ms}^{-1}]$	5×10 <sup>-7</sup>	5×10 <sup>-4</sup>	1×10 <sup>-5</sup>
$S_{f_0} [{ m m}^{-1}]$	10 <sup>-6</sup>	10-3	5×10 <sup>-5</sup>
$\alpha_0  [\text{m}^{-1}\text{s}^{-1}]$	10-12	10-9	3×10 <sup>-11</sup>
a [-]	0.1	0.5	0.3
b [-]	0.1	0.5	0.2

## **Table 2**

Parameters	Mean	95% Conf. Int.
$K_{f_0}  [{ m ms}^{-1}]$	0.99×10 <sup>-5</sup>	[0.9-1.1]×10 <sup>-5</sup>
$S_{f_0}$ [m <sup>-1</sup> ]	5.03×10 <sup>-5</sup>	[3.2-7.6]×10 <sup>-5</sup>
$\alpha_0  [\text{m}^{-1}\text{s}^{-1}]$	3.06×10 <sup>-11</sup>	[2.1-4.5]×10 <sup>-11</sup>
a [-]	0.299	[0.28-0.32]
b [-]	0.2	[0.17-0.23]







**Figure 2** 







