



# **Dissolved inorganic nitrogen and phosphorus dynamics in seawater following an artificial Saharan dust deposition event**

Justine Louis, Matthieu Bressac, Maria L. Pedrotti, Cecile Guieu

## **► To cite this version:**

Justine Louis, Matthieu Bressac, Maria L. Pedrotti, Cecile Guieu. Dissolved inorganic nitrogen and phosphorus dynamics in seawater following an artificial Saharan dust deposition event. *Frontiers in Marine Science*, 2015, 2, <10.3389/fmars.2015.00027>. <hal-03502750>

**HAL Id: hal-03502750**

**<https://hal.science/hal-03502750v1>**

Submitted on 27 Dec 2021

**HAL** is a multi-disciplinary open access archive for the deposit and dissemination of scientific research documents, whether they are published or not. The documents may come from teaching and research institutions in France or abroad, or from public or private research centers.

L'archive ouverte pluridisciplinaire **HAL**, est destinée au dépôt et à la diffusion de documents scientifiques de niveau recherche, publiés ou non, émanant des établissements d'enseignement et de recherche français ou étrangers, des laboratoires publics ou privés.



Distributed under a Creative Commons CC BY 4.0 - Attribution - International License

# Dissolved inorganic nitrogen and phosphorus dynamics in seawater following an artificial Saharan dust deposition event

Justine Louis<sup>1,2</sup>, Matthieu Bressac<sup>1,2†</sup>, Maria L. Pedrotti<sup>1,2</sup> and Cécile Guieu<sup>1,2\*</sup>

## OPEN ACCESS

### Edited by:

Claire Mahaffey,  
University of Liverpool, UK

### Reviewed by:

Andrew King,  
Norwegian Institute for Water  
Research, Norway  
Arvind Singh,  
GEOMAR Helmholtz Centre for Ocean  
Research Kiel, Germany

### \*Correspondence:

Cécile Guieu,  
OOV-Centre National de la Recherche  
Scientifique, UMR 7093, Observatoire  
Océanologique de Villefranche,  
Laboratoire d'Océanographie de  
Villefranche/Mer, Bat. Jean Maetz,  
Chemin du Lazaret, 06234  
Villefranche/Mer, France  
guieu@obs-vlfr.fr

### † Present Address:

Matthieu Bressac,  
Institute for Marine and Antarctic  
Studies, University of Tasmania,  
Hobart, TAS, Australia

### Specialty section:

This article was submitted to  
Marine Biogeochemistry,  
a section of the journal  
Frontiers in Marine Science

**Received:** 06 January 2015

**Accepted:** 06 May 2015

**Published:** 26 May 2015

### Citation:

Louis J, Bressac M, Pedrotti ML and  
Guieu C (2015) Dissolved inorganic  
nitrogen and phosphorus dynamics in  
seawater following an artificial  
Saharan dust deposition event.  
Front. Mar. Sci. 2:27.  
doi: 10.3389/fmars.2015.00027

<sup>1</sup> Sorbonne Universités, Université Pierre et Marie Curie, UMR 7093, Observatoire Océanologique de Villefranche, Laboratoire d'Océanographie de Villefranche-sur-Mer, Villefranche-sur-Mer, France, <sup>2</sup> Centre National de la Recherche Scientifique, UMR 7093, Observatoire Océanologique de Villefranche-sur-Mer, Laboratoire d'Océanographie de Villefranche-sur-Mer, Villefranche-sur-Mer, France

The dynamics of dissolved inorganic nitrogen and phosphorus in seawater after a dust event were followed to better understand the impact of dust deposition in low nutrient waters of the Mediterranean Sea. Three independent abiotic experiments were performed over three seasons (winter, spring, end of summer) characterized by contrasted biogeochemical conditions. Experiments consisted of seeding evapocondensed Saharan dust at the surface of a polyethylene tank filled with filtered surface seawater. Phosphate ( $\text{PO}_4^{3-}$ ), nitrate ( $\text{NO}_3^-$ ), size and number of particles and transparent exopolymeric particles production (TEP) were measured over the course of 1 week following seeding. Dust deposition was followed by a transient increase in  $[\text{PO}_4^{3-}]$  during the first 3 h with a maximum input of 33, 9, and 39 nM, respectively in May, October and February. The removal of almost all the  $\text{PO}_4^{3-}$  initially released suggests a scavenging process of  $\text{PO}_4^{3-}$  back onto ferric oxide-rich particles leading to concentrations at the end of the experiment close to the initial values (7 nM in May and October, and 6 nM in February).  $\text{NO}_3^-$  released from dust was high especially in May and October (maximum input of 23 and 11  $\mu\text{M}$ , respectively) and was attributed to nitrogen dissolution from the large amount of small particles ( $<1 \mu\text{m}$ ) rich in nitrogen in the evapocondensed dust.  $[\text{NO}_3^-]$  remained high until the end of the experiment (16  $\mu\text{M}$  in May and 11  $\mu\text{M}$  in October), indicating that  $\text{NO}_3^-$  from dust is likely to be bioavailable for a longer period compared to  $\text{PO}_4^{3-}$  from dust. The release of  $\text{PO}_4^{3-}$  and  $\text{NO}_3^-$  was intrinsically linked to particle dynamics, governed by the quality/quantity of dissolved organic matter.

**Keywords:** atmospheric nutrients, dissolution, adsorption, particle size distribution, TEP, dissolved organic matter, Saharan dust, Mediterranean Sea

## Introduction

Studying the interactions between atmosphere and ocean is essential for better understanding biogeochemical cycles in the ocean, a critical step toward accurate modeling of ocean dynamics under climate and anthropogenic changes (e.g., Law et al., 2013; Liss and Johnson, 2014). With 500 Mt per year of dust deposited to the surface ocean (Law et al., 2013), the atmosphere is

a significant pathway for the transport of material from natural and anthropogenic sources to the coastal and open ocean. Atmospheric deposition is a source of external new nutrients to the ocean (i.e., Duce et al., 1991). Over the past decade, a number of studies have investigated post-depositional processes (dissolution, adsorption) of atmospheric nutrients in seawater and the impact of atmospheric deposition on biological activity (i.e., Guieu et al., 2014b). Most of the so-called “dissolution experiments” have been performed in bottles with a constant and homogeneous concentration of particles over the course of the experiment. As exchanges between dissolved and particulate phases occur while the atmospheric particles are sinking in the ocean, results from such “batch” approaches, that do not take into account the vertical dimension, could lead to incorrect quantifications of dissolution/scavenging processes involving atmospheric particles (de Leeuw et al., 2014). Additionally, most studies do not take into consideration the biogeochemical characteristics of the water used in experiments, such as the amount of dissolved organic matter.

Mineral aerosol particles are a dominant source of phosphorus (P) and iron (Fe) to the open ocean as they represent 83% of atmospheric total P and 95% of atmospheric total Fe globally (Jickells et al., 2005; Mahowald et al., 2008, 2009; Krishnamurthy et al., 2010; Liss and Johnson, 2014). Contrary to P and Fe, nitrogen (N) is dominated by anthropogenic emissions in gaseous form which represent 80% of the annual N deposition ( $67 \text{ Tg N yr}^{-1}$  in 2000) to the global ocean (Duce et al., 2008). Such input could support up to 3% of all oceanic annual new production (Duce et al., 2008) and 20% in regional seas subject to high anthropogenic N deposition such as the South China Sea (Kim et al., 2014). During the transport of dust, in particular in the marine atmosphere where the relative humidity is high, the atmospheric gas  $\text{HNO}_3$  reacts with mineral particles rich in carbonates and leads to coating of dust by nitrates (Krueger et al., 2004; Formenti et al., 2011). A strong mixing of dust in polluted air masses can allow an enrichment of anthropogenic N on mineral particles that may also become a significant source of N.

New nutrients from the atmosphere such as N, P, and Fe have the potential to impact marine biogeochemistry in Low Nutrient Low Chlorophyll waters (i.e., Guieu et al., 2014a); as is the case in tropical waters where  $\text{N}_2$  fixation can be co-limited by Fe and P (Mills et al., 2004) and in the Mediterranean Sea where the pelagic community can be limited by the bioavailability of P and N during the stratified period (Tanaka et al., 2011; Ridame et al., 2014).

The Mediterranean Sea, an oligotrophic basin characterized by an unusual high N/P ratio [about 21–23 in Western basin, (Bethoux et al., 1992)] is subjected to high Saharan dust deposition, which can serve as a significant source of new P and Fe to this semi-enclosed basin. Additionally, a strong internal mixing of dust and nitrate from polluted European and North African air masses (e.g., Putaud et al., 2004) by cloud processes implies that Saharan dust events are likely a source of anthropogenic N.

A previous study taking into account the vertical dimension necessary for the study of the fate of atmospheric nutrients in the ocean was carried out in the framework of the DUNE

project. Dust seeding over large, clean *in situ* mesocosms (Guieu et al., 2010, 2014b) showed that a wet deposition event resulted in a strong increase in the N/P ratio in surface waters due to the higher input of  $\text{NO}_3^-$  relative to  $\text{PO}_4^{3-}$  in the evapocondensed dust used for the experiment (Ridame et al., 2014). Such dust input can either be a source or a sink of dissolved iron (DFe) depending on the biogeochemical conditions of the seawater (Wagener et al., 2010; Ye et al., 2011; Wuttig et al., 2013). Phosphate turnover time ( $\text{TPO}_4^{3-}$ ), which represents the ratio between natural concentration and uptake by plankton (Thingstad et al., 1993), are very short in the surface waters of the Mediterranean Sea (e.g., Thingstad et al., 1996; Moutin et al., 2002) making it difficult to observe and follow the fate of new nutrients and to deconvolute biological processes from post-deposition abiotic processes onto dust (i.e., dissolution/adsorption). Therefore, this previously overlooked differentiation is crucial for future modeling of atmospheric nutrients inputs.

A logical follow-up to the DUNE project was to study the fate of new nutrients after a dust deposition in abiotic conditions. As with the previous project, the strategy involved simulated dust depositions, but was performed on top of seawater volumes large enough to take into account the vertical dimension of particle sinking while also eliminating the processes involving living organisms. We performed three experiments in clean “minicosms” filled with 300 L filtered seawater ( $<0.2 \mu\text{m}$ ). The seawater was collected in the bay of Villefranche-sur-mer during 3 contrasting seasonal scenarios: (1) during winter mixing, (2) immediately after the spring phytoplankton bloom, and (3) at the end of the stratification period during the end of summer. The behavior of the DFe in these experiments, conducted inside a clean room, have been previously presented in Bressac and Guieu (2013). This study revealed contrasted post-deposition processes affecting the DFe concentration [DFe]: during the spring and summer experiments, the rapid formation of aggregates induced a negative feedback on [DFe] through scavenging, while a fraction of DFe was likely organically complexed. During the winter experiment, the absence of aggregation processes led to a large transient increase in [DFe] before being removed by adsorption onto settling particles. These contrasting behaviors were attributed to differences in the quantity and quality of the dissolved organic matter (DOM) in the seawater utilized. This study confirmed that atmospheric iron dissolution in seawater is not controlled by similar processes throughout the year, as it depends on seasonal *in situ* conditions such as DOM concentration, an outcome previously shown by Wagener et al. (2008). Indeed, during the spring bloom, DOM is more abundant than during winter. It accumulates in the euphotic zone from spring to late summer following the increase in primary and secondary production (Copin-Montégut and Avril, 1993). This fresh DOM is assumed to be labile and mainly composed of acid polysaccharides and sugars thought to serve as precursors to transparent exopolymeric particles (TEP). TEP acts as an organic glue and is required for the formation of sinking aggregates (Kepkay, 2000; Passow, 2000; Beauvais et al., 2003). During summer in the study area, the increase in TEP concentration is consistent with the accumulation of

carbon-rich DOM (Avril, 2002), with TEP-C concentrations corresponding to 25% of the total organic carbon (Mari et al., 2001).

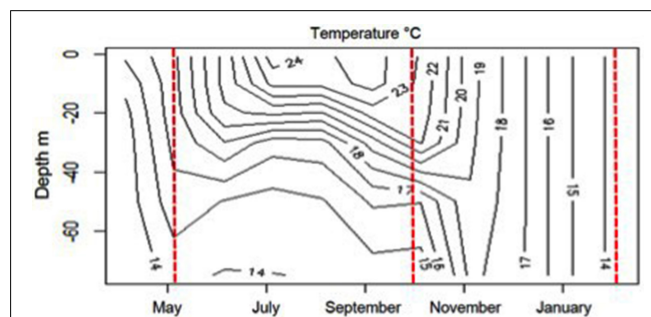
We measured the concentration of  $\text{NO}_3^-$  and  $\text{PO}_4^{3-}$  from samples collected during the same experiments as described in Bressac and Guieu (2013). The objectives of this study are (1) to present  $\text{PO}_4^{3-}$  and  $\text{NO}_3^-$  results using nanomolar sensitivity analytical techniques and (2) to link the dynamics of  $\text{PO}_4^{3-}$  and  $\text{NO}_3^-$  to the physical-chemical processes (dissolution, adsorption, and aggregation) (taking into consideration the DOM pool by measuring concentrations of TEP and absorption of the chromophoric DOM) occurring in low-nutrient waters after a dust event in contrasting biogeochemical conditions and, (3) to discuss the potential biogeochemical effect of such atmospheric input in LNLC waters.

## Material and Method

### Minicosm Experiment

The experimental setup is fully described in Bressac and Guieu (2013). A short summary is given here. A polyethylene (PE) tank (height = 1.09 m; diameter = 0.68 m; surface area = 0.36 m<sup>2</sup>; volume = 0.31 m<sup>3</sup>), hereafter called a “minicosm,” was installed in a clean room under controlled light and temperature settings (constant temperature of 20°C). A cover made of high density polyethylene (HDPE) was used to seal the minicosm tightly to avoid light penetration. The minicosm was conical at the bottom, and ended with a sediment trap from which exported material was collected at the end of each experiment. The minicosm was filled with filtered surface seawater collected at 5 m depth from the Bay of Villefranche (Mediterranean Sea) using three trace metal clean Teflon pumps connected to PE tubes. Inline filtration was performed using a 0.2 µm cartridge (Sartorius Sartobran-P-capsule with a 0.45 µm prefilter and a 0.2 µm final filter) and allowed to directly fill 25 L trace-metal clean HDPE containers that were kept in the dark until the seawater was transferred into the minicosm (less than 4 h for the entire operation). Seawater sampling and subsequent minicosm experiments were performed (1) in May 2011, at the beginning of the water column stratification while the phytoplankton bloom was ending (hereafter referred to as “May”), (2) in October 2011, at the end of the stratification period (“October”), and (3) in February 2012 during winter mixing (“February”) (Figure 1).

The same dust analog used during the DUNE experiments (see details in Guieu et al., 2010) was used to mimic a Saharan dust deposition at the surface of the minicosm. Briefly, the <20 µm fraction of soil collected in Tunisia was used after appropriate physico-chemical treatments in the laboratory, including dust processing with model cloud water to simulate processes during the atmospheric transport of aerosols. The particle number size distribution showed that 99% of particles had a size smaller than 1 µm. The dust analog, mostly made of quartz (40%), calcite (30%), and clay (25%), had the same P (and Fe) content with or without a chemical artificial aging, while concentration content of N was one order of magnitude higher after the dust was chemically processed due to the addition of  $\text{HNO}_3$  in the



**FIGURE 1 | Vertical distribution of temperature (°C) in the 0–70 m depth layer between April 2011 and February 2012 at Bay of Villefranche sur mer (data SOMLIT, <http://somlit.epoc.u-bordeaux1.fr/fr/>). The red lines correspond to the day when the water was collected in May, October, and February.**

simulated cloud water. Fe, N, and P represented 2.31, 1.15, and 0.05% (w/w) of the dust analog (Guieu et al., 2010).

To mimic a Saharan dust wet deposition flux of 10 g m<sup>-2</sup> at the minicosm surface (the same flux as during the DUNE experiment), 3.6 g of dust analog was diluted into 2 L of ultrapure water and sprayed at the surface of the minicosm. Sampling was done prior to the seeding, and 1, 3, 6, 12, 16, 24, 31, 48, 72, 96, 120, and 144 h after the seeding. At the end of each experiment, less than 15% of the initial volume of the minicosm (300 L) had been removed for sampling over the course of the experiment. At each sampling time, samples were taken at 0.1, 0.3, and 0.6 m-depths through three sampling tubes with inlets at the center of the minicosm. A system of “float and weight” was linked to the three tubes to keep the sampling depths consistent throughout the experiment (Bressac and Guieu, 2013). Aliquots for the measurements of  $\text{PO}_4^{3-}$  and  $\text{NO}_3^-$  concentrations ( $[\text{PO}_4^{3-}]$  and  $[\text{NO}_3^-]$ ) were immediately filtered with acid-cleaned 0.2 µm polycarbonate membrane filters (Nucleopore, Whatman). Filtrates were stored in 60 mL HDPE bottles and acidified to pH 2 (quartz-distilled HCl). The evolution of the particle size distributions, absorption of the chromophoric dissolved organic matter (<sup>a</sup>CDOM), concentration of TEP [TEP] and [DFe] were also measured in the minicosm over the course of 1 week following seeding; in addition particulate organic carbon (POC) was measured in the material collected from the sediment trap of the minicosms. All those data are available in Bressac and Guieu (2013).

### Nano-Molar Levels of Nutrients

$\text{PO}_4^{3-}$  and  $\text{NO}_3^-$  were analyzed using two different Liquid Waveguide Capillary Cells (LWCC) made of quartz capillary tubing.

### Dissolved Inorganic Phosphorus

We used the same protocol (and equipment), based on Murphy and Riley’s molybdenum blue method, as described by Pulido-Villena et al. (2010). Briefly LWCC was connected to a spectrophotometer and the measurements were performed in the visible spectrum at 710 nm. The 2-m length LWCC allowed for a detection limit of 1 nM and a relative standard deviation of



less than 10%. The reagents were prepared daily according to the same protocol as in Zhang and Chi (2002) and Zimmer and Cutter (2012).

### Dissolved Inorganic Nitrogen

Dissolved inorganic nitrogen (DIN) in seawater is the sum of nitrate ( $\text{NO}_3^-$ ) plus nitrite ( $\text{NO}_2^-$ ) and was measured using a 1-m long LWCC. Here, we focus solely on  $[\text{NO}_3^-]$  as  $\text{NO}_2^-$  concentrations ( $[\text{NO}_2^-]$ ) were close to the detection limit in May and October. The analytical setup used was based on several studies (Yao et al., 1998; Zhang, 2000; Adornato et al., 2007; Li et al., 2008). The analysis of  $\text{NO}_3^- + \text{NO}_2^-$  is based on the spectrophotometric Griess method (Chen et al., 2008).  $\text{NO}_2^-$  absorbance was measured at 540 nm. A further reduction of  $\text{NO}_3^-$  to  $\text{NO}_2^-$  by a copperized cadmium column (Gal et al., 2004) was necessary to quantify  $\text{NO}_3^-$  present in the sample.

The calibration curves were performed using standard solutions in ultrapure water diluting  $\text{NaNO}_2$  and  $\text{KNO}_3$  ("high purity" quality), which were previously dried at 60°C for 2 h. We obtained (1) a reduction efficiency between 90 and 96%, (2) a repeatability standard deviation (RSD) of 3% for the reduction process, (3) a detection limit for  $[\text{NO}_2^-]$  and  $[\text{NO}_2^- + \text{NO}_3^-]$  of 2 and 9 nmol  $\text{L}^{-1}$  after 10 blanks measurements, and (4) a  $p < 0.01$  for  $T_{\text{student}}$  and ANOVA tests to verify the accuracy and the linear regression of this analytical method. Samples with  $[\text{NO}_3^-] > 500 \text{ nM}$  (outside of the linear range of the calibration curve) were diluted with low  $[\text{NO}_3^-]$  seawater.

In the absence of reference seawater at nanomolar concentrations, the reproducibility and the repeatability of the analysis of  $\text{PO}_4^{3-}$  and  $\text{NO}_3^-$  were controlled by an internal standard (the seawater collected during the experiment of May in the bay of Villefranche sur mer with  $[\text{PO}_4^{3-}] = 3 \text{ nM}$  and  $[\text{NO}_3^-] = 28 \text{ nM}$ ).

### Particle Size Distributions

Particle size distributions of the 0.7–18  $\mu\text{m}$  size range were determined by Coulter Counter (Beckman) using a 30  $\mu\text{m}$  aperture tube. The particle size distribution in number concentration  $N(D)$  (part  $\text{L}^{-1} \mu\text{m}^{-1}$ ) was obtained after 5 replicate measurements and normalized using the extent of each size class as described by Bressac and Guieu (2013). To follow the size evolution of the particles settling, the average particle diameter  $D_{\text{avg}}$  ( $\mu\text{m}$ ) was calculated as follows:

$$D_{\text{avg}} = \frac{\sum_i (N(D_i) \times D_i)}{n\text{Pc}}$$

where the symbol  $i$  corresponds to the different size classes over the 0.7–18  $\mu\text{m}$  size range,  $D_i$  is the mid-point diameter of a given size class, and  $n\text{Pc} = \sum_i N(D_i)$  is the total particle concentration over the 0.7–18  $\mu\text{m}$  size range (part  $\text{L}^{-1} \mu\text{m}^{-1}$ ).

### TEP Determination

Transparent exopolymeric particles (TEP) are acidic exopolysaccharides. This fraction can be quantified by spectrophotometry using Alcian blue stain. Samples (300 mL) were filtered through a 0.4  $\mu\text{m}$  polycarbonate filter (Poretic) at low and constant vacuum ( $<150 \text{ mmHg}$ ) to preserve cells.

Material retained on the filter was stained with 500  $\mu\text{L}$  of an aqueous solution of 0.06% acetic acid (pH 2.5) and 0.02% Alcian blue (Sigma, 8GX). After a few seconds, the filter was dried and rinsed with 1 mL milliQ water, and subsequently stored at  $-20^\circ\text{C}$  until analysis. TEP concentration was measured according to the dye-binding assay (Passow and Alldredge, 1995), 6 mL of 80% sulfuric acid was added to samples, followed by immediate vortexing. Following incubation at room temperature for 20 h, absorption was measured at 787 nm. TEP values are expressed as xanthan gum weight equivalent ( $X_{\text{equival}} \text{ L}^{-1}$ ) calculated by means of a calibration curve. The carbon content of TEP (TEP-C) was estimated from colorimetric determinations according to Engel and Passow (2001).

TEP-C was measured *in situ* for the 3 experiments. The evolution of the TEP-C after dust addition in the minicosm was followed in samples collected at 0.3 m depth in October and February. No minicosm data is available for the May experiment.

## Results

To better understand the relationship between the dynamics of the nutrients and the particles, we took the average values of the 3-depths for each parameter ( $n\text{Pc}$ ,  $D_{\text{avg}}$ ,  $[\text{NO}_3^-]$ ,  $[\text{PO}_4^{3-}]$ ,  $\Delta[\text{NO}_3^-]$ ,  $\Delta[\text{PO}_4^{3-}]$ ); these averages were used for clarity and consistency with Bressac and Guieu (2013).  $\Delta[\text{NO}_3^-]$  and  $\Delta[\text{PO}_4^{3-}]$  represent the release of  $\text{NO}_3^-$  ( $\mu\text{M}$ ) and  $\text{PO}_4^{3-}$  (nM) from dust. In effect, these correspond to the difference between  $[\text{NO}_3^-]$  and  $[\text{PO}_4^{3-}]$  after dust addition and before dust addition. The 3-depth averages  $[\text{NO}_3^-]$  and  $[\text{PO}_4^{3-}]$  are hereafter referred to as  $[\text{NO}_3^-]_{\text{avg}}$  and  $[\text{PO}_4^{3-}]_{\text{avg}}$ .

### Initial Conditions at the Three Different Seasons in the Bay of Villefranche

The evolution of the seawater temperature from April 2011 to February 2012 (Figure 1) clearly identifies the stratified period (May and October experiments) and winter mixing (February experiment). In May (Table 1), the seawater was typical of the post-bloom conditions with  $[\text{Chla}] = 0.28 \mu\text{g L}^{-1}$ , a nutrient depletion ( $[\text{NO}_3^-] = 28 \text{ nM}$ ;  $[\text{PO}_4^{3-}] = 3 \text{ nM}$ ) and a high value of  $[\text{TEP-C}]$  ( $27.1 \mu\text{mol C L}^{-1}$ ). In October,  $[\text{NO}_3^-]$  and  $[\text{PO}_4^{3-}]$  remained very low whereas  $[\text{Chla}]$  and  $[\text{TEP-C}]$  decreased ( $0.13 \mu\text{g L}^{-1}$  and  $15.4 \mu\text{mol C L}^{-1}$ , respectively). In February when the whole water column was well-mixed,  $[\text{NO}_3^-]$  was one order of magnitude higher compared to May and October (327 nM) whereas  $[\text{PO}_4^{3-}]$  was still low (7 nM). In February, TEP-C had the lowest concentration ( $2.7 \mu\text{mol C L}^{-1}$ ).

### Particle Size Distribution

The evolution of  $n\text{Pc}$  (Figures 2A,B,C) and  $D_{\text{avg}}$  (Figures 2D,E,F) are presented as the average of the values measured at the three sampling depths. Briefly, a maximum  $D_{\text{avg}}$  was observed during the first 3 h of the three experiments [ $1.9 \mu\text{m}$  at  $T_3$  (May),  $1.4 \mu\text{m}$  at  $T_1$  (October), and  $1.8 \mu\text{m}$  at  $T_3$  (February)], before decreasing and reaching the lowest values at  $T_{144}$  (0.98, 0.95, and  $0.91 \mu\text{m}$  in May, October, and February, respectively). The decreases in  $D_{\text{avg}}$  were concomitant with marked increases in  $n\text{Pc}$ , which were observed particularly in

**TABLE 1 | Initial conditions (unfiltered water: chlorophyll a, TEP-C; filtered water: nitrate, phosphate) for the three minicosm experiments in May, October and February (*in situ* values at 5-m depth).**

	May	October	February
Chla ( $\mu\text{g L}^{-1}$ )	0.28	0.13	0.17
$\text{NO}_3^-$ ( $\text{nmol L}^{-1}$ )	28	26	327
$\text{PO}_4^{3-}$ ( $\text{nmol L}^{-1}$ )	3	6	7
$a_{\text{CDOM}}(320)\text{m}^{-1}$	0.248	0.211	0.188
TEP-C ( $\mu\text{mol C L}^{-1}$ )	27.1	15.4	2.7

May and October (**Figures 2A,B,D,E**). In May between  $T_{12}$  and  $T_{31}$ , nPc increased from  $6.4 \times 10^8$  to  $1.9 \times 10^{10}$   $\text{part L}^{-1} \mu\text{m}^{-1}$ , and in October nPc increased from  $2.9 \times 10^9$  to  $3.7 \times 10^{10}$   $\text{part L}^{-1} \mu\text{m}^{-1}$  between  $T_6$  and  $T_{12}$ . After reaching a maximum, nPc decreased slightly but remained high until the end of both experiments ( $1.1 \times 10^{10}$  and  $1.5 \times 10^{10}$  in May and October, respectively at  $T_{144}$ ). In February, a one order of magnitude increase in nPc was observed at the end of the experiment between  $T_{96}$  and  $T_{144}$  ( $4.5 \times 10^9$  to  $3.5 \times 10^{10}$ ) whereas  $D_{\text{avg}}$  was  $\sim 1 \mu\text{m}$ . This pool of submicron particles corresponding to the lower size range of the dust population, are numerically dominant (cf. Section Minicosm Experiment).

### Evolution of Nitrate Concentrations

In May,  $[\text{NO}_3^-]_{\text{avg}}$  increased from 0.03 to  $0.12 \mu\text{M}$  between  $T_0$  and  $T_{0.5}$ , and the largest increase was observed at  $T_{31}$  (**Figure 2G**) with  $\Delta[\text{NO}_3^-]_{\text{max}}$  ( $[\text{NO}_3^-]_{\text{avg max}} - [\text{NO}_3^-]_{\text{avg } T_0}$ ) reaching  $23 \mu\text{M}$ . This increase was concomitant with the increase in nPc (**Figure 2A**).  $[\text{NO}_3^-]_{\text{avg}}$  released from dust was high with  $\Delta[\text{NO}_3^-]_{\text{end}}$  ( $[\text{NO}_3^-]_{\text{avg } T_{144}} - [\text{NO}_3^-]_{\text{avg } T_0}$ ) =  $16 \mu\text{M}$ . In October the release of  $\text{NO}_3^-$  was faster but lower than in May (**Figure 2H**), with a large increase in  $[\text{NO}_3^-]_{\text{avg}}$  that started 1 h after dust addition and a  $\Delta[\text{NO}_3^-]_{\text{max}}$  of  $11 \mu\text{M}$  measured at  $T_{16}$ . Simultaneously,  $D_{\text{avg}}$  decreased between  $T_1$  and  $T_{12}$  (**Figure 2E**). As in May,  $[\text{NO}_3^-]_{\text{avg}}$  remained stable until the end of the experiment ( $\Delta[\text{NO}_3^-]_{\text{end}} = 11 \mu\text{M}$ ). In February  $\Delta[\text{NO}_3^-]_{\text{max}}$  was the lowest among the three experiments ( $1.4 \mu\text{M}$ ) and occurred at the end of the experiment when nPc was maximum (**Figures 2C,I**).

The dust addition allowed a maximum change in  $\text{NO}_3^-$  stock of +19.2, +9.2, and  $+1.2 \text{ mmol m}^{-2}$  in May, October, and February, respectively (**Figure 3**).

### Evolution of Phosphate Concentrations

For all seasons, the release of  $\text{PO}_4^{3-}$  occurred mainly during the first 3 h after dust addition following the  $D_{\text{avg}}$  pattern (**Figures 2D,E,F,J,K,L**).  $\Delta[\text{PO}_4^{3-}]_{\text{max}}$  ( $[\text{PO}_4^{3-}]_{\text{avg max}} - [\text{PO}_4^{3-}]_{\text{avg } T_0}$ ) were 33, 9, and  $39 \text{ nM}$  in May ( $T_{0.5}$ ), October ( $T_1$ ), and February ( $T_3$ ), respectively. In October and February,  $[\text{PO}_4^{3-}]_{\text{avg}}$  released from dust was totally removed since  $[\text{PO}_4^{3-}]_{\text{avg}}$  at the end of both experiments (7 and  $6 \text{ nM}$ , respectively) was close to the initial value ( $\pm 1 \text{ nM}$ ). In May, after a decrease in  $[\text{PO}_4^{3-}]_{\text{avg}}$  between  $T_{0.5}$  and  $T_{12}$  (from 36 to  $7 \text{ nM}$ ), a second increase in  $[\text{PO}_4^{3-}]_{\text{avg}}$  was observed at  $T_{31}$  ( $19 \text{ nM}$ ) and

then continually decreased until the end of the experiment when  $\Delta[\text{PO}_4^{3-}]_{\text{end}}$  ( $[\text{PO}_4^{3-}]_{\text{avg } T_{144}} - [\text{PO}_4^{3-}]_{\text{avg } T_0}$ ) was  $4 \text{ nM}$ .

The dust addition allowed a maximum change in  $\text{PO}_4^{3-}$  stock of +27.2, +7.5, and  $+32.5 \mu\text{mol m}^{-2}$  in May, October, and February, respectively (**Figure 3**). Six days after seeding, the  $\text{PO}_4^{3-}$  stock was  $+2.5 \mu\text{mol m}^{-2}$  (May),  $+0.8 \mu\text{mol m}^{-2}$  (October), and  $-0.8 \mu\text{mol m}^{-2}$  (February) compared with the initial  $\text{PO}_4^{3-}$  stock (**Figure 3**).

### Evolution of TEP-C Concentrations

In October, after the dust addition, the TEP-C concentration increased from 0.4 to  $5.0 \mu\text{mol C L}^{-1}$  between  $T_0$  and  $T_{0.5}$  and the largest increase ( $10.3 \mu\text{mol C L}^{-1}$ ) was observed 24 h after, with concentrations close to the level of those observed *in situ* ( $15.4 \mu\text{mol C L}^{-1}$ ). After TEP was completely removed by filtration, almost the same amount of TEP was formed by precursors 1 day after the dust addition. Afterward the TEP concentration decreased to  $\sim 3.2 \mu\text{mol C L}^{-1}$  until the end of the experiment (**Figure 4**). In February, the rate of TEP formation was slower than in October with an increase in production ( $6.9 \mu\text{mol C L}^{-1}$ ) 6 h after dust addition (**Figure 4**), then the TEP concentration decreased until the end of the experiment.

## Discussion

### Release of Phosphate and Nitrate Controlled by Particle Dynamics

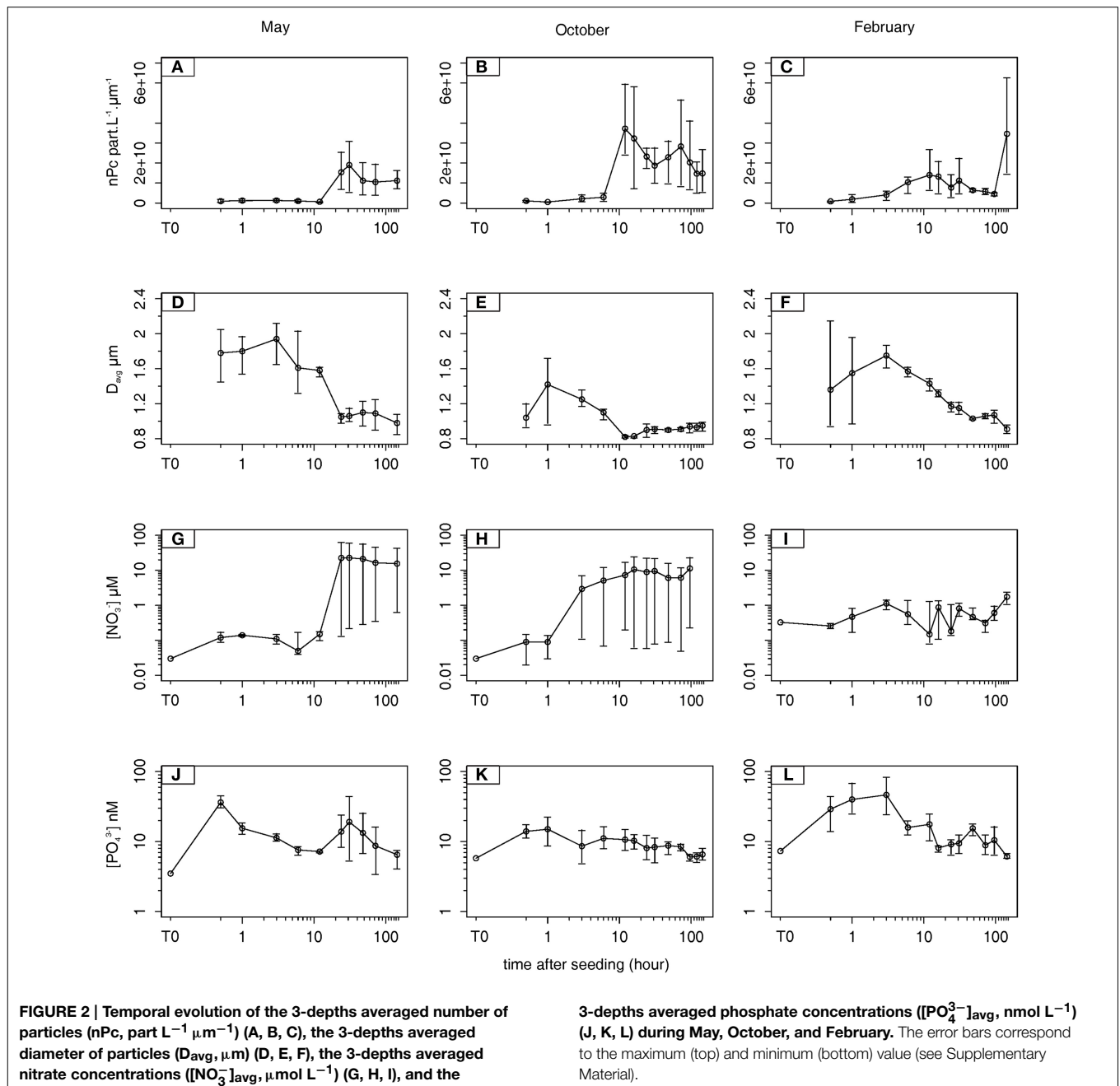
99% of the number of dust particles used to seed the minicosm were submicronic particles (Guieu et al., 2010). According to Stokes' law, the settling velocities of particles are a function of their diameter and we expected the settling of largest particles rapidly after dust addition, followed by a numerically dominant slowly settling particle population with a diameter smaller than  $1 \mu\text{m}$ . This is illustrated by **Figures 2A,B,C,D,E,F** as an increase in  $D_{\text{avg}}$  during the first 3 h after dust addition, followed by a decrease in  $D_{\text{avg}}$  until the end of the three experiments, while a concomitant increase in nPc was also observed.

### Nitrate vs. Particle Size Dynamics

For all seasons,  $\Delta[\text{NO}_3^-]_{\text{max}}$  was measured when the submicron mineral particles were abundant, i.e., high nPc and  $D_{\text{avg}}$  smaller than  $1 \mu\text{m}$ . This is in agreement with heterogeneous reactions on the surface of mineral particles in the atmosphere. Indeed, during the atmospheric transport of Saharan dust, heterogeneous reactions with compounds such as gaseous nitric acid occur on particle surfaces rich in calcium carbonate (Usher et al., 2003; de Leeuw et al., 2014). These reactions facilitate formation of nitrate coatings on particle surfaces (Laskin, 2005). The observed decrease in  $D_{\text{avg}}$  led to an increase in the overall surface area to volume ratio, and thus an increase in the relative amount of  $\text{NO}_3^-$  on the dust surface (Vlasenko et al., 2005).

### Phosphate vs. Particle Size Dynamics

An increase in  $[\text{PO}_4^{3-}]_{\text{avg}}$  occurred during the first 3 h after dust addition while the largest mineral particles were settling down. Such a fast release is in agreement with previous batch experiments indicating that 65% of the  $\text{PO}_4^{3-}$  release occurred



within the first 6 h (Ridame and Guieu, 2002). Despite a similar simulated dust flux of  $10 \text{ g m}^{-2}$  for each experiment, the maximum increase in  $PO_4^{3-}$  stock was lower in October than in May and February (Figure 3). The lowest  $D_{avg}$  in October (Figure 2E) suggests that the non-homogeneous distribution of mineral particles controlled the release of  $PO_4^{3-}$  from dust.

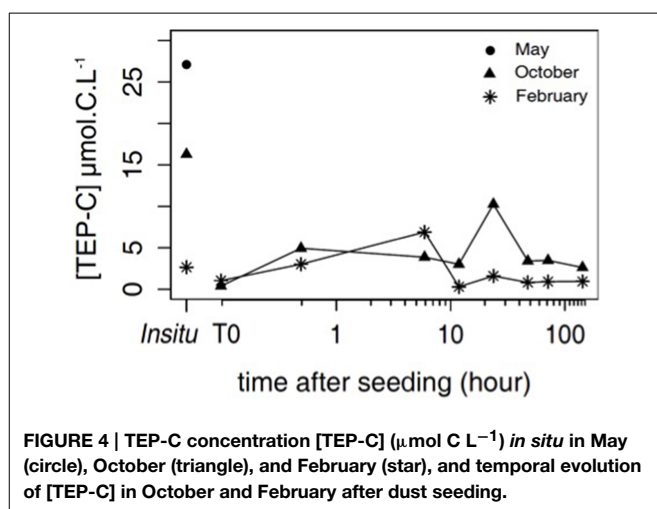
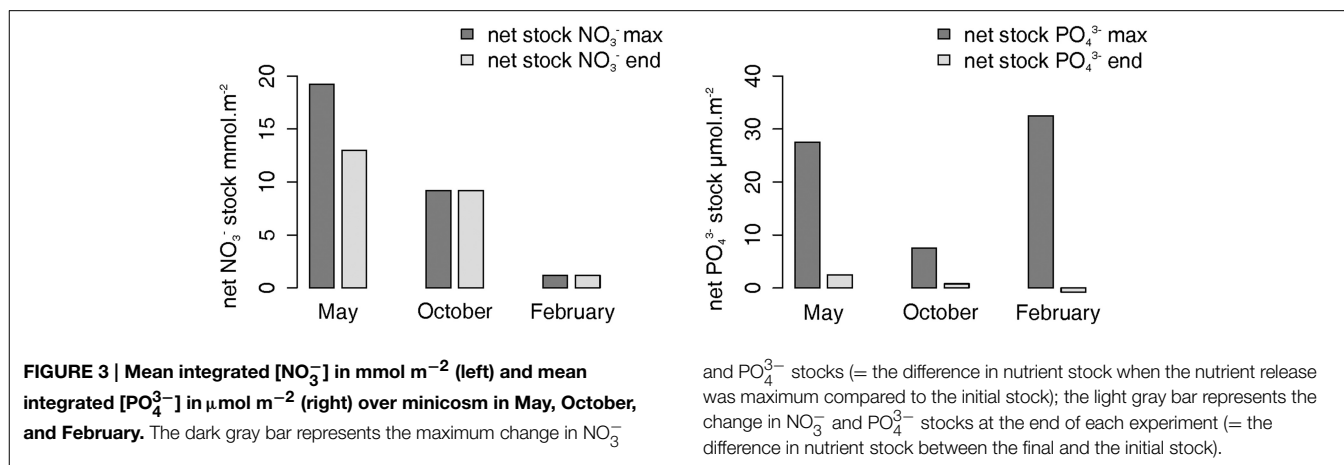
### Adsorption of Phosphate onto Ferric Oxide-Rich Particles and Link between DFe and $PO_4^{3-}$

After the dust addition, there was a sudden transient increase in  $[PO_4^{3-}]_{avg}$  attributed to a release from the largest particles, followed by a decrease in  $[PO_4^{3-}]_{avg}$  for the three experiments

corresponding to the decrease in  $D_{avg}$  (Figures 2D,E,F,J,K,L). In abiotic conditions, such a decrease can be attributed to the removal process of adsorption of freshly released  $PO_4^{3-}$  onto the submicron sinking particles. Almost all  $[PO_4^{3-}]$  released from dust was scavenged-back onto sinking particles within few days.

### Correlation between DFe and $PO_4^{3-}$

When comparing the respective evolutions of [DFe] (Bressac and Guieu, 2013) and  $[PO_4^{3-}]$  (Figure 5), a good correlation ( $R^2 = 0.91$ ,  $p < 0.001$ ) was observed in February indicating that the post-depositional processes (both dissolution and removal processes) could be similar for the two elements. In February,



DFe dissolution was very fast (occurring during the first 3 h after seeding) and was followed by DFe adsorption onto the submicron mineral particles (Bressac and Guieu, 2013). Strong correlation between DFe and  $\text{PO}_4^{3-}$  in February indicates that both DFe and  $\text{PO}_4^{3-}$  removal from seawater was due to their adsorption back onto the submicron mineral particles.

In May, a good linear relationship ( $R^2 = 0.84$ ) between  $[\text{DFe}]$  and  $[\text{PO}_4^{3-}]$  was only observed at 0.3 and 0.6 m but not at 0.1 m-depth. At 0.1 m, DFe release in seawater was found to be controlled by a dust-DOM aggregation process (Bressac and Guieu, 2013). The aggregation process likely resulted in a decrease in accessibility to surface adsorption sites. While DFe, for which some dissolved organic compounds have a strong affinity (Van den Berg, 1995; Barbeau, 2006), was removed from the dissolved phase,  $\text{PO}_4^{3-}$  was likely not affected by the aggregation process. At 0.3 and 0.6 m, the positive linear correlation seems to indicate that the removal of both  $\text{PO}_4^{3-}$  and DFe from seawater was independent of dust-DOM aggregation but rather controlled by adsorption processes onto the submicron mineral particles, similar to what was hypothesized for February.

In October, the correlation between the two nutrients (at the 3-depths) was weaker than in February ( $R^2 = 0.70$ ), but still

significant ( $p < 0.001$ ). This could be explained by the dust-DOM aggregation process still present but weaker than in May (Bressac and Guieu, 2013).

### Adsorption of $\text{PO}_4^{3-}$ onto the Submicron Particles Rich in Iron Oxide

Based on previous studies of the interaction between iron oxides and  $\text{PO}_4^{3-}$ , the observed removal of  $\text{PO}_4^{3-}$  was likely due to the strong  $\text{PO}_4^{3-}$  adsorption capacity at the surface of the mineral particles rich in iron oxides. Berner (1973) focused on the  $\text{PO}_4^{3-}$  removal from seawater by the adsorption capacity of the volcanogenic iron oxides and suggested  $\text{PO}_4^{3-}$  can react with poorly crystallized hydrous iron oxides. Sei et al. (2002) showed the presence of iron oxides in kaolinite with a large specific surface area enhanced  $\text{PO}_4^{3-}$  adsorption, while Parfitt and Atkinson (1976) confirmed the complex involved in  $\text{PO}_4^{3-}$  adsorption on goethite was of the type Fe-O-P-O-Fe.

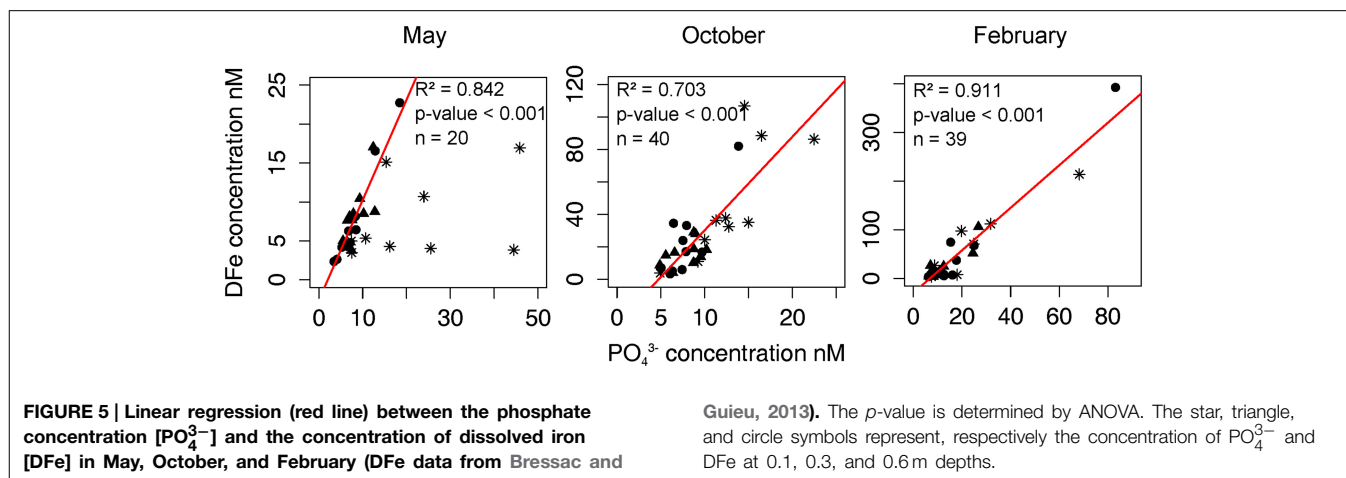
The dissolved-particulate exchanges of phosphorus from dust are controlled by iron oxide rich particles. We assume that the dust-DOM aggregation process, occurring particularly in May, could prevent the interaction between  $\text{PO}_4^{3-}$  and iron oxide at the surface of mineral particles, and could explain the second transient increase in  $[\text{PO}_4^{3-}]$  (Figure 2J).

### The Seasonal Variability of Nitrate Input from Saharan Dust: Does the DOM Matter?

By taking into account the vertical dimension and thus the temporal variability of the distribution of mineral particles, our minicosm approach showed that the release of  $\text{NO}_3^-$  was not instantaneous or total after dust addition. This differs from results presented by Ridame et al. (2014) who conducted a batch dissolution experiment using the same dust analog but in artificial DOM-free seawater. This study showed a complete dissolution of  $\text{NO}_x$  occurred within 3 h after dust addition (Ridame et al., 2014).

Here we found a strong variability of  $\text{NO}_3^-$  dissolution from dust at the 3 explored seasons (the difference between nutrient stock when  $\text{NO}_3^-$  release was maximum and initial stock were 19.2, 9.2, and 1.2  $\text{mmol m}^{-2}$  in May, October, and February, respectively) (Figure 3). One possibility for the





observed variability could be a link between  $\text{NO}_3^-$  dissolution and the quantity and quality of DOM described by Bressac and Guieu (2013). Indeed, it was quite variable according to the season: in May (post-bloom condition), the seawater was characterized by an abundant and newly produced DOM, whereas in February (winter mixing), the DOM was less abundant and likely older (Table 1). Although to our knowledge such link has not been proposed in the ocean, some interesting hypotheses can be made. Bressac and Guieu (2013) showed that [DFe] resulted from contrasted post-deposition processes linked to the DOM pool. In May and October, with abundant and newly produced DOM conditions, the rapid formation of aggregates inducing a negative feedback on the [DFe] through scavenging, while a fraction of the DFe was likely organically complexed. In February, the low DOM content strongly limited aggregation allowing a strong iron dissolution that was followed by the (re)adsorption of DFe onto the submicron dust particles. This suggests that in February the submicron mineral particles would be richer in iron oxides than in May in the first few hours after seeding. The same particles would release 10–20 times less  $\text{NO}_3^-$  compared to any other season (see above Section). One possible explanation could be that some reactions involving iron and DOM could transform the chemical form of N following “the ferrous wheel hypothesis” as defined Davidson et al. (2003) to explain abiotic immobilization of N in mineral soils rich in iron. This would involve a reduction of  $\text{NO}_3^-$  to  $\text{NO}_2^-$  by a redox reaction with Fe(III) and a nitration reaction with DOM to transform  $\text{NO}_2^-$  to dissolved organic nitrogen (DON). This is speculative, however, because DON was not measured in our study. Nevertheless, the link between  $\text{NO}_3^-$  and DOM in the ocean still need further investigations to explore these potentially important dynamics.

## Biogeochemical Effects

### Bioavailability of Atmospheric Nutrients

At initial conditions, the  $\text{NO}_3^-/\text{PO}_4^{3-}$  ratio suggests P-limitation conditions for phytoplankton activity in February ( $\text{NO}_3^-/\text{PO}_4^{3-} \gg 16$ ) and a NP co-limitation in May and October ( $\text{NO}_3^-/\text{PO}_4^{3-} \ll 16$  with low  $[\text{PO}_4^{3-}]$ ) (Table 1). These results are in agreement with the results presented by Tanaka

et al. (2011) for summer conditions at “station A” ( $39^\circ 5.96' \text{N}$ ,  $5^\circ 21.00' \text{E}$ ) in the North Western Mediterranean basin.

A N or NP co-limitation with initial  $\text{NO}_3^-/\text{PO}_4^{3-} < 6$  and low  $[\text{PO}_4^{3-}]$  was also found in June 2010 [DUNE 2 (R1), (Ridame et al., 2014)] and is comparable to our initial conditions in May and October. At the end of the DUNE 2 (R1) mesocosms experiment, the  $\text{NO}_3^-/\text{PO}_4^{3-}$  ratio was  $> 150$  which led to P limitation conditions. In the present study, we also observed a strong increase in the  $\text{NO}_3^-/\text{PO}_4^{3-}$  ratio along the course of the experiments in May and October due to the higher input of  $\text{NO}_3^-$  relative to  $\text{PO}_4^{3-}$  in the evapocondensed dust used for the experiment.

This is in good agreement with higher  $\text{NO}_3^-/\text{PO}_4^{3-}$  in atmospheric deposition compared to the classical Redfield Ratio (i.e., Herut et al., 1999, 2002; Baker et al., 2010; Markaki et al., 2010). In the Mediterranean Sea ratio up to 87 were found in seawater leachable dust samples (Herut et al., 2002) and ratio in total atmospheric deposition ranging 46–161 were found over the whole Mediterranean Sea (Markaki et al., 2010). Much higher ratio have even been found in the NW Mediterranean Sea for very intense anthropogenic episodes (soluble inorganic N/soluble inorganic P  $> 1000$ ; Sandroni et al., 2007). Such unbalanced  $\text{NO}_3^-/\text{PO}_4^{3-}$  in the atmospheric deposition, including dust deposition, likely contribute to the unusual  $\text{NO}_3^-/\text{PO}_4^{3-}$  found in the Mediterranean Sea as hypothesized by many authors (i.e., Herut et al., 1999; Markaki et al., 2010).

The results obtained in this abiotic study along with those obtained from mesocosm experiments representing the natural environmental assemblage can facilitate a better understanding of the fate of atmospheric nutrients in LNLC waters. For example, our May experiment can be compared to *in situ* results obtained during the DUNE-1-P and DUNE-2 (R1) mesocosm experiments (Pulido-Villena et al., 2014; Ridame et al., 2014). This mesocosm experiment was conducted in a nutrient-depleted system during the stratification period.  $\text{NO}_3^-$  and  $\text{PO}_4^{3-}$  released from dust after the simulation of a wet deposition, was followed by an increase in bacteria respiration (Pulido-Villena et al., 2014) and primary production (Ridame et al., 2014). The phytoplankton  $\text{PO}_4^{3-}$  requirements estimated in DUNE-1-P and DUNE-2 (R1)

were +25 and +18  $\mu\text{mol m}^{-2}$ , respectively (Ridame et al., 2014). The increase in  $\text{PO}_4^{3-}$  stock in May (+27  $\mu\text{mol m}^{-2}$ , **Figure 3**) was of the same order of magnitude as for the transient increase observed in DUNE-1-P (+21  $\mu\text{mol m}^{-2}$ ) and DUNE-2 (R1) (+24  $\mu\text{mol m}^{-2}$ ) experiments (Ridame et al., 2014). This indicates first that  $\text{PO}_4^{3-}$  release from dust can fulfill the phytoplankton requirement. Second, the fast disappearance of the  $\text{PO}_4^{3-}$  following its release from dust in the mesocosm experiment was due to the rapid uptake ( $\text{PO}_4^{3-}$  turnover times were estimated to be  $\sim 1$  h using a  $^{33}\text{P}$  uptake technique on the upper stratified water of the Mediterranean sea, Moutin et al. 2002) resulting in a fast increase (within 2 days) in bacteria respiration and primary production by heterotroph/autotroph communities that characterize this oligotrophic system. This suggests that biological uptake was faster than the adsorption of  $\text{PO}_4^{3-}$  back onto the submicron particles rich in iron oxide observed in abiotic conditions (Section Adsorption of Phosphate onto Ferric Oxide-Rich Particles and Link between DFe and  $\text{PO}_4^{3-}$ ). The large increase in  $\text{NO}_3^-$  stock after dust addition [+19.2 and +10.2  $\text{mmol m}^{-2}$ , respectively in May and DUNE-2 (R1) experiments] was much higher than the phytoplankton  $\text{NO}_3^-$  requirement [0.7 and 0.8  $\mu\text{mol m}^{-2}$  in DUNE-1-P and DUNE-2 (R1), respectively, (Ridame et al., 2014)] and this explains that a large amount of  $\text{NO}_3^-$  released from dust remained bioavailable until the end of the DUNE experiments.

In February, phytoplankton growth is primarily limited by the light availability (Bonilla-Findji et al., 2010), and  $\text{PO}_4^{3-}$  released from a dust event would rapidly be removed by the adsorption process onto sinking mineral particles. Nevertheless, with favorable weather conditions the rapid release of  $\text{PO}_4^{3-}$  and DFe (Bressac and Guieu, 2013) from dust could stimulate biological activity during months characterized by low irradiance.

### Abiotic Formation of TEP

Owing to their surface-reactive nature, the abiotic formation of TEP is an important pathway to convert dissolved into POC (**Figure 4**) increasing the possibility to aggregate DOM. The organic aggregates can interact with the mineral particles and lead to “abiotic POC export” that influences carbon export to the deep ocean (Ternon et al., 2010). The highest flux of abiotic POC was found in May [20.0  $\text{mg m}^{-2} \cdot \text{d}^{-1}$ , compared to 9.6 and 6.1  $\text{mg m}^{-2} \cdot \text{d}^{-1}$  in October and February, respectively (Bressac and Guieu, 2013)] and suggests that this month has the largest abiotic formation of TEP. The consequences of such strong aggregation and export, in particular on the microbial food web, are unknown.

### Conclusion

Minicosm experiments mimicking artificial dust deposition in abiotic conditions were performed during three seasons representing contrasted biogeochemical conditions. Monitoring  $[\text{NO}_3^-]$  and  $[\text{PO}_4^{3-}]$  during particle settling highlighted the link between dissolved-particulate exchanges of P and N and particle dynamics. A transient increase in  $[\text{PO}_4^{3-}]$  occurred during the first 3 h after dust addition while the larger particles were still settling. The maximum 3-depths averaged  $\text{PO}_4^{3-}$  release was 33, 9, and 39 nM in May, October, and February, respectively.

The maximum 3-depths averaged  $\text{NO}_3^-$  release was 23, 11, and 1.4  $\mu\text{M}$  in May, October, and February experiments, respectively.  $[\text{NO}_3^-]$  released from dust remained relatively stable and will be likely bioavailable over a 1 week scale at minimum in surface waters. This could be important because in May and October this represents 10–30 times the maximum  $[\text{NO}_3^-]$  concentrations observed in the North Mediterranean Sea (2–3  $\mu\text{mol L}^{-1}$  in winter) (Marty et al., 2002). The seasonal variability of  $[\text{NO}_3^-]$  released from dust suggests that freshly produced DOM would enhance the dissolution process of  $\text{NO}_3^-$  (as is the case for the May experiment after the spring bloom). This link between atmospheric  $\text{NO}_3^-$  and DOM will require further exploration.

The adsorption back onto ferric oxide-rich particles of  $\text{PO}_4^{3-}$  released from dust leads to short-term bioavailability. The dust-DOM aggregation process could hinder the interaction between  $\text{PO}_4^{3-}$  and iron oxide at the surface of mineral particles. The input of new nutrients and the possible shift toward higher  $\text{NO}_3^-/\text{PO}_4^{3-}$  ratios induced by a Saharan dust event mixed with anthropogenic nitrogen have a potential impact on the autotrophic and heterotrophic communities in Mediterranean Sea and other oligotrophic areas receiving such inputs.

This work demonstrates the importance of taking into account the settling of particles while the dissolved/particulate phase reactions occur, as these new nutrient dynamics are intrinsically linked to particle dynamics. Additionally, particle dynamics are conditioned by the quality/quantity of DOM, suggesting that the post-deposition processes of  $\text{NO}_3^-$  and  $\text{PO}_4^{3-}$  are likely indirectly linked to the quality/quantity of DOM. Consequently, the seasonality of the quality/quantity of DOM implies a non-constant potential impact of atmospheric new nutrients in the surface waters over the course of the year. The next step of this work will be to consider how on-going environmental changes (such as temperature increase and pH decrease) affect the release of atmospheric nutrients, including organic forms.

### Acknowledgments

This work was funded by the MISTRALS/MERMEX project and the CHIPIE-EMERGENCE-UPMC project. A grant provided by the University of Paris 6 supported JL. A grant provided by the ACRI-ST company and the French National Association for Research and Technology (ANRT) supported MB. The authors thank K. Desboeufs and F. Dulac for providing the Saharan dust analog. J.-Y. Carval from the Observatoire d’Océanologie de Villefranche sur mer is acknowledged for his helpful work at sea. We thank our colleagues from SOMLIT (<http://somlit.epoc.u-bordeaux1.fr/fr/>) for providing the *in situ* temperature and Chla data. The authors thank C. Ridame for her constructive comments on a previous version of the manuscript and K. Frischkorn for kindly editing the manuscript.

### Supplementary Material

The Supplementary Material for this article can be found online at: <http://journal.frontiersin.org/article/10.3389/fmars.2015.00027/abstract>

## References

- Adornato, L.R., Kaltenbacher, E. A., Greenhow, D. R., and Byrne, R. H. (2007). High-resolution *in situ* analysis of nitrate and phosphate in the oligotrophic ocean. *Environ. Sci. Technol.* 41, 4045–4052. doi: 10.1021/es0700855
- Avril, B. (2002). DOC dynamics in the northwestern Mediterranean Sea (DYFAMED site). *Deep-sea Res. II Top. Stud. Oceanogr.* 49, 2163–2182. doi: 10.1016/S0967-0645(02)00033-4
- Baker, A. R., Lesworth, T., Adams, C., Jickells, T. D., and Ganzeveld, L. (2010). Estimation of atmospheric nutrient inputs to the Atlantic Ocean from 50°N to 50°S based on large-scale field sampling: fixed nitrogen and dry deposition of phosphorus. *Global Biogeochem. Cycles* 24, GB3006. doi: 10.1029/2009GB003634
- Barbeau, K. (2006). Photochemistry of organic iron (III) complexing ligands in oceanic systems. *Photochem. Photobiol.* 82, 1505–1516. doi: 10.1111/j.1751-1097.2006.tb09806.x
- Beauvais, S., Pedrotti, M. L., Villa, E., and Lemée, R. (2003). Transparent exopolymer particle (TEP) dynamics in relation to trophic and hydrological conditions in the NW Mediterranean Sea. *Mar. Ecol. Prog. Ser.* 262, 97–109. doi: 10.3354/meps262097
- Berner, R. A. (1973). Phosphate removal from sea water by adsorption on volcanogenic ferric oxides. *Earth Planet. Sci. Lett.* 18, 77–86. doi: 10.1016/0012-821X(73)90037-X
- Bethoux, J. P., Morin, P., Madec, C., and Gentili, B. (1992). Phosphorus and nitrogen behaviour in the Mediterranean Sea. *Deep Sea Res. A Oceanogr. Res. Pap.* 39, 1641–1654. doi: 10.1016/0198-0149(92)90053-V
- Bonilla-Findji, O., Gattuso, J.P., Pizay, M.D., and Weinbauer, M. G. (2010). Autotrophic and heterotrophic metabolism of microbial planktonic communities in an oligotrophic coastal marine ecosystem: seasonal dynamics and episodic events. *Biogeosciences* 7, 3491–3503. doi: 10.5194/bg-7-3491-2010
- Bressan, M., and Guieu, C. (2013). Post-depositional processes: what really happens to new atmospheric iron in the ocean's surface? *Global Biogeochem. Cycles* 27, 859–70. doi: 10.1002/gbc.20076
- Chen, G., Yuan, D., Huang, Y., Zhang, M., and Bergman, M. (2008). In-field determination of nanomolar nitrite in seawater using a sequential injection technique combined with solid phase enrichment and colorimetric detection. *Anal. Chim. Acta* 620, 82–88. doi: 10.1016/j.aca.2008.05.019
- Copin-Montégut, G., and Avril, B. (1993). Vertical distribution and temporal variation of dissolved organic carbon in the North-Western Mediterranean Sea. *Deep Sea Res. Part I Oceanogr. Res. Pap.* 40, 1963–1972. doi: 10.1016/0967-0637(93)90041-Z
- Davidson, E.A., Chorover, J., and Dail, D. B. (2003). A mechanism of abiotic immobilization of nitrate in forest ecosystems: the ferrous wheel hypothesis. *Glob. Change Biol.* 9, 228–236. doi: 10.1046/j.1365-2486.2003.00592.x
- de Leeuw, G., Guieu, C., Arneth, A., Bellouin, N., Bopp, L., Boyd, P. W., et al. (2014). “Ocean-atmosphere interactions of particles,” in *Ocean-Atmosphere Interactions of Gases and Particles*, eds S. Peter Liss and T. Martin Johnson (Berlin; Heidelberg: Springer Earth System Sciences Springer), 171–246. Available online at: [http://link.springer.com/chapter/10.1007/978-3-642-25643-1\\_4](http://link.springer.com/chapter/10.1007/978-3-642-25643-1_4)
- Duce, R. A., LaRoche, J., Altieri, K., Arrigo, K. R., Baker, A. R., Capone, D. G., et al. (2008). Impacts of atmospheric anthropogenic nitrogen on the open ocean. *Science* 320, 893–897. doi: 10.1126/science.1150369
- Duce, R. A., Liss, P. S., Merrill, J. T., Atlas, E. L., and Buat-Menard, P., Hicks, B. B., et al. (1991). The Atmospheric input of trace species to the World Ocean. *Global Biogeochem. Cycles* 5, 193–259. doi: 10.1029/91GB01778
- Engel, A., and Passow, U. (2001). Carbon and Nitrogen Content of Transparent Exopolymer Particles (TEP) in relation to their alcian blue adsorption. *Mar. Ecol. Prog. Ser.* 219, 1–10. doi: 10.3354/meps219001
- Formenti, P., Schütz, L., Balkanski, Y., Desboeufs, K., Ebert, M., Kandler, K. A., et al. (2011). Recent progress in understanding physical and chemical properties of African and Asian mineral dust. *Atmos. Chem. Phys.* 11, 8231–8256. doi: 10.5194/acp-11-8231-2011
- Gal, C., Frenzel, W., and Miller, J. (2004). Re-examination of the cadmium reduction method and optimisation of conditions for the determination of nitrate by flow injection analysis. *Microchimica Acta* 146, 155–164. doi: 10.1007/s00604-004-0193-7
- Guieu, C., Aumont, O., Paytan, A., Bopp, L., Law, C. S., Mahowald, N., et al. (2014a). The significance of the episodic nature of atmospheric deposition to low nutrient low chlorophyll regions. *Global Biogeochem. Cycles* 28, 1179–1198. doi: 10.1002/2014GB004852
- Guieu, C., Dulac, F., Desboeufs, K., Wagener, T., Pulido-Villena, E., Grisoni, J.-M., et al. (2010). Large clean mesocosms and simulated dust deposition: a new methodology to investigate responses of marine Oligotrophic Ecosystems to atmospheric inputs. *Biogeosciences* 7, 2765–2784. doi: 10.5194/bg-7-2765-2010
- Guieu, C., Dulac, F., Ridame, C., and Pondaven, P. (2014b). Introduction to Project DUNE, a DUST experiment in a low nutrient, Low Chlorophyll Ecosystem. *Biogeosciences* 11, 425–442. doi: 10.5194/bg-11-425-2014
- Herut, B., Collier, R., and Krom, M. D. (2002). The role of dust in supplying nitrogen and phosphorus to the Southeast Mediterranean. *Limnol. Oceanogr.* 47, 870–878. doi: 10.4319/lo.2002.47.3.0870
- Herut, B., Krom, M. D., Pan, G., and Mortimer, R. (1999). Atmospheric input of nitrogen and phosphorus to the Southeast Mediterranean: sources, fluxes, and possible impact. *Limnol. Oceanogr.* 44, 1683–1692. doi: 10.4319/lo.1999.44.7.1683
- Jickells, T. D., An, Z. S., Andersen, K. K., Baker, A. R., Bergametti, G., Brooks, N., et al. (2005). Global iron connections between desert dust, ocean biogeochemistry, and climate. *Science* 308, 67–71. doi: 10.1126/science.1105959
- Kepkay, P. E. (2000). “Colloids and the Ocean carbon cycle,” in *Marine Chemistry. The Handbook of Environmental Chemistry 5D*, ed P. J. Wangersky (Berlin Heidelberg: Springer) 35–56. Available online at: [http://link.springer.com/chapter/10.1007/10683826\\_2](http://link.springer.com/chapter/10.1007/10683826_2)
- Kim, T.-W., Lee, K., Duce, R., and Liss, P. (2014). Impact of atmospheric nitrogen deposition on phytoplankton productivity in the South China Sea. *Geophys. Res. Lett.* 41, 3156–3162. doi: 10.1002/2014GL059665
- Krishnamurthy, A., Moore, J. K., Mahowald, N., Luo, C., and Zender, C. S. (2010). Impacts of atmospheric nutrient inputs on Marine biogeochemistry. *J. Geophys. Res.* 115, G01006. doi: 10.1029/2009JG001115
- Krueger, B. J., Grassian, V. H., Cowin, J. P., and Laskin, A. (2004). Heterogeneous chemistry of individual mineral dust particles from different dust source regions: the importance of particle mineralogy. *Atmos. Environ.* 38, 6253–6261. doi: 10.1016/j.atmosenv.2004.07.010
- Laskin, A. (2005). Heterogeneous Chemistry of individual mineral dust particles with Nitric Acid: a combined CCSEM/EDX, ESEM, and ICP-MS Study. *J. Geophys. Res.* 110, D10208. doi: 10.1029/2004JD005206
- Law, C. S., Brévière, E., de Leeuw, G., Garçon, V., Guieu, C., and Kieber, D. J., et al. (2013). Evolving Research Directions in Surface Ocean - Lower Atmosphere (SOLAS) Science. *Environ. Chem.* 10, 1. doi: 10.1071/EN12159
- Li, Q. P., Hansell, D. A., and Zhang, J.-Z. (2008). Underway monitoring of nanomolar nitrate plus nitrite and phosphate in oligotrophic seawater. *Limnol. Oceanogr. Methods* 6, 319–326. doi: 10.4319/lom.2008.6.319
- Liss, P. S., and Johnson, M. T. (eds.). (2014). *Ocean-Atmosphere Interactions of Gases and Particles*. (Berlin, Heidelberg: Springer Berlin Heidelberg). Available online at: <http://link.springer.com/10.1007/978-3-642-25643-1>
- Mahowald, N., Jickells, T. D., Baker, A. R., Artaxo, P., and Benitez-Nelson, C. R., Bergametti, G. et al. (2008). Global distribution of atmospheric phosphorus sources, concentrations and deposition rates, and anthropogenic impacts. *Global Biogeochem. Cycles* 22, GB4026. doi: 10.1029/2008GB003240
- Mahowald, N. M., Engelstaedter, S., Luo, C., Sealy, A., Artaxo, P., Benitez-Nelson, C., et al. (2009). Atmospheric iron deposition: global distribution, variability, and human perturbations. *Ann. Rev. Mar. Sci.* 1, 245–278. doi: 10.1146/annurev.marine.010908.163727
- Mari, X., Beauvais, S., Lemée, R., and Pedrotti, M. L. (2001). Non-Redfield C: N ratio of transparent exopolymeric particles in the northwestern Mediterranean Sea. *Limnol. Oceanogr.* 46, 1831–1836. doi: 10.4319/lo.2001.46.7.1831
- Markaki, Z., Loÿe-Pilot, M. D., Violaki, K., Benyahya, L., and Mihalopoulos, N. (2010). Variability of Atmospheric deposition of dissolved Nitrogen and Phosphorus in the Mediterranean and Possible Link to the Anomalous Seawater N/P Ratio. *Mar. Chem.* 120, 187–194. doi: 10.1016/j.marchem.2008.10.005
- Marty, J.-C., Chiavérini, J., Pizay, M.-D., and Avril, B. (2002). Seasonal and interannual dynamics of nutrients and phytoplankton pigments in the western Mediterranean Sea at the DYFAMED time-series station (1991–1999). *Deep*

- Sea Res. Part II Top. Stud. Oceanogr.* 49, 1965–1985. doi: 10.1016/S0967-0645(02)00022-X
- Mills, M. M., Ridame, C., Davey, M., La Roche, J., and Geider, R. J. (2004). Iron and Phosphorus Co-Limit Nitrogen fixation in the Eastern Tropical North Atlantic. *Nature* 429, 292–294. doi: 10.1038/nature02550
- Moutin, T., Thingstad, T. F., Van Wambeke, F., Marie, D., Slawyk, G., et al. and Claustre, H. (2002). Does competition for nanomolar phosphate supply explain the predominance of the cyanobacterium *Synechococcus*? *Limnol. Oceanogr.* 47, 1562–1567. doi: 10.4319/lo.2002.47.5.1562
- Parfitt, R. L., and Atkinson, R. J. (1976). Phosphate Adsorption on Goethite ( $\alpha$ -FeOOH). *Nature* 264, 740–742. doi: 10.1038/264740a0
- Passow, U. (2000). Formation of transparent exopolymer particles, TEP, from dissolved precursor material. *Mar. Ecol. Prog. Ser.* 192, 1–11. doi: 10.3354/meps192001
- Passow, U., and Alldredge, A. L. (1995). A dye-binding assay for the spectrophotometric measurement of transparent exopolymer particles (TEP). *Limnol. Oceanogr.* 40, 1326–1335. doi: 10.4319/lo.1995.40.7.1326
- Pulido-Villena, E., Baudoux, A.-C., Obernosterer, I., Landa, M., Caparros, J., Catala, P., et al. (2014). Microbial food web dynamics in response to a Saharan Dust event: results from a mesocosm study in the Oligotrophic Mediterranean Sea. *Biogeosciences* 11, 5607–5619. doi: 10.5194/bg-11-5607-2014
- Pulido-Villena, E., Rérolle, V., and Guieu, C. (2010). Transient fertilizing effect of dust in P-deficient LNLC surface ocean. *Geophys. Res. Lett.* 37, L01603. doi: 10.1029/2009GL041415
- Putaud, J.-P., van Dingenen, R., Dell'Acqua, A., Raes, F., Matta, E., Decesari, S., et al. (2004). Size-segregated aerosol mass closure and chemical composition in Monte Cimone (I) during MINATROC. *Atmos. Chem. Phys.* 4, 889–902. doi: 10.5194/acp-4-889-2004
- Ridame, C., Dekazemack, J., Guieu, C., Bonnet, S., L'Helguen, S., and Malien, F. (2014). Contrasted saharan dust events in LNLC environments: impact on nutrient dynamics and primary production. *Biogeosciences* 11, 4783–4800. doi: 10.5194/bg-11-4783-2014
- Ridame, C., and Guieu, C. (2002). Saharan input of phosphate to the oligotrophic water of the open western Mediterranean Sea. *Limnol. Oceanogr.* 47, 856–869. doi: 10.4319/lo.2002.47.3.0856
- Sandroni, V., Raimbault, P., Migon, C., Garcia, N., and Gouze, E. (2007). Dry atmospheric deposition and diazotrophy as sources of new nitrogen to northwestern Mediterranean oligotrophic surface waters. *Deep Sea Res. Part I Oceanogr. Res. Pap.* 54, 1859–1870. doi: 10.1016/j.dsr.2007.08.004
- Sei, J., Jumas, J. C., Olivier-Fourcade, J., Quiquampoix, H., and Staunton, S. (2002). Role of iron oxides in the phosphate adsorption properties of kaolinites from the Ivory Coast. *Clays Clay Miner.* 50, 217–222. doi: 10.1346/000986002760832810
- Tanaka, T., Thingstad, T. F., Christaki, U., Colombet, J., and Cornet-Barthaux, V., Courties, C., et al. (2011). Lack of P-limitation of phytoplankton and heterotrophic prokaryotes in surface waters of three anticyclonic eddies in the stratified Mediterranean Sea. *Biogeosciences* 8, 525–538. doi: 10.5194/bg-8-525-2011
- Ternon, E., Guieu, C., Loje-Pilot, M.-D., Leblond, N., Bosc, E., Gasser, B., et al. (2010). The impact of Saharan dust on the particulate export in the water column of the North Western Mediterranean Sea. *Biogeosciences* 7, 809–826. doi: 10.5194/bg-7-809-2010
- Thingstad, T. F., Dolan, J. R., and Fuhrman, J. A. (1996). Loss rate of an oligotrophic bacterial assemblage as measured by H-thymidine and PO<sub>4</sub>: good agreement and near-balance with production. *Aquat. Microb. Ecol.* 10, 29–36. doi: 10.3354/ame010029
- Thingstad, T. F., Skjoldal, E. F., Böhne, R. A. (1993). Phosphorus cycling and algal-bacterial competition in Sandsfjord, western Norway. *Mar. Ecol. Prog. Ser.* 99, 239–239. doi: 10.3354/meps099239
- Usher, C. R., Michel, A. E., and Grassian, V. H. (2003). Reactions on mineral dust. *Chem. Rev.* 103, 4883–4940. doi: 10.1021/cr020657y
- Van den Berg, C. M. G. (1995). Evidence for organic complexation of iron in seawater. The chemistry of iron in seawater and its interaction with phytoplankton. *Mar. Chem.* 50, 139–157. doi: 10.1016/0304-4203(95)00032-M
- Vlasenko, A., Sjogren, S., Weingartner, E., Stemmler, K., Gäggeler, H. W., and Ammann, M. (2005). Effect of humidity on nitric acid uptake to mineral dust aerosol particles. *Atmos. Chem. Phys. Discuss.* 5, 11821–11860. doi: 10.5194/acpd-5-11821-2005
- Wagener, T., Guieu, C., and Leblond, N. (2010). Effects of dust deposition on iron cycle in the surface mediterranean sea: results from a mesocosm seeding experiment. *Biogeosciences* 7, 3769–3781. doi: 10.5194/bg-7-3769-2010
- Wagener, T., Pulido-Villena, E., and Guieu, C. (2008). Dust iron dissolution in seawater: results from a one-year time-series in the Mediterranean Sea. *Geophys. Res. Lett.* 35, L16601. doi: 10.1029/2008GL034581
- Wuttig, K., Wagener, T., Bressac, M., Dammshäuser, A., Streu, P., Guieu, C., et al. (2013). Impacts of dust deposition on dissolved trace metal concentrations (Mn, Al and Fe) during a mesocosm experiment. *Biogeosciences* 10, 2583–2600. doi: 10.5194/bg-10-2583-2013
- Yao, W., Byrne, R. H., and Waterbury, R. D. (1998). Determination of nanomolar concentrations of nitrite and nitrate in natural waters using long path length absorbance spectroscopy. *Environ. Sci. Technol.* 32, 2646–2649. doi: 10.1021/es9709583
- Ye, Y., Wagener, T., Völker, C., Guieu, C., and Wolf-Gladrow, D. A. (2011). Dust Deposition: iron source or sink? a case study. *Biogeosciences* 8, 2107–2124. doi: 10.5194/bg-8-2107-2011
- Zhang, J.-Z. (2000). Shipboard automated determination of trace concentrations of nitrite and nitrate in oligotrophic water by gas-segmented continuous flow analysis with a liquid waveguide capillary flow cell. *Deep Sea Res. Part I Oceanogr. Res. Pap.* 47, 1157–1171. doi: 10.1016/S0967-0637(99)00085-0
- Zhang, J. Z., and Chi, J. I. E. (2002). Automated analysis of nanomolar concentrations of phosphate in natural waters with liquid waveguide. *Environ. Sci. Technol.* 36, 1048–1053.
- Zimmer, L. A., and Cutter, G. A. (2012). High resolution determination of nanomolar concentrations of dissolved reactive phosphate in ocean surface waters using long path liquid waveguide capillary cells (LWCC) and spectrometric detection. *Limnol. Oceanogr. Methods* 10, 568–580. doi: 10.4319/lom.2012.10.568

**Conflict of Interest Statement:** The authors declare that the research was conducted in the absence of any commercial or financial relationships that could be construed as a potential conflict of interest.

Copyright © 2015 Louis, Bressac, Pedrotti and Guieu. This is an open-access article distributed under the terms of the Creative Commons Attribution License (CC BY). The use, distribution or reproduction in other forums is permitted, provided the original author(s) or licensor are credited and that the original publication in this journal is cited, in accordance with accepted academic practice. No use, distribution or reproduction is permitted which does not comply with these terms.