

Viscous transfer of momentum across a shallow laminar flow

O. Devauchelle¹†, P. Popović¹ and E. Lajeunesse¹

¹Institut de Physique du Globe de Paris, F-75238 Paris, France

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In a shallow channel, the flow transfers most of its momentum vertically. Based on this observation, one often neglects the momentum that is transferred across the stream—the core assumption of the shallow-water theory. In the context of viscous flows, this approximation is referred to as the “lubrication theory”, in which one assumes that the shear stress exerted by the fluid on the substrate over which it flows is proportional to its velocity. Here, we revise this theory to account for the momentum that viscosity transfers across a shallow laminar flow, while keeping the problem low-dimensional. We then test the revised lubrication theory against analytical and numerical solutions of the exact problem. We find that, at a low computational cost, the present theory represents the actual flow more accurately than the classical lubrication approximation. This theoretical improvement, devised with laboratory rivers in mind, should also apply to other geophysical contexts, such as ice flows or forming lava domes.

1. Introduction

Before the Navier-Stokes equations were fully established, hydraulic engineers developed empirical laws relating the turbulent flow of water in a canal to its slope and cross-section (Chézy 1775). One of the most successful among them is that of Darcy and Weisbach, which relates the average velocity U of the flow to the canal’s depth and downstream gradient (D and S , respectively; Brown 2002):

$$U = \sqrt{\frac{2gDS}{f_D}}, \quad (1.1)$$

where g is the acceleration of gravity, and f_D an empirical, dimensionless parameter.

Once an entirely empirical equation, the Darcy-Weisbach equation is now regarded as the result of the steady-state momentum balance, when the turbulent shear stress on the canal’s bottom, τ , is proportional to the squared flow velocity ($\tau = \rho f_D U^2 / 2$, where ρ is the density of water). In that light, equation (1.1) appears as an elementary solution of the shallow-water equations (which are often named after Saint-Venant 1871). As such, equation (1.1) becomes an approximate solution to the Navier-Stokes equations, and thus fits nicely into general fluid dynamics.

The shallow-water equations result from a twofold procedure (Stoker 2011): one first (i) integrates the Navier-Stokes equations vertically, and then (ii) approximates the resulting mass and momentum balances. Before step (ii) is carried

† Email address for correspondence: devauchelle@ipgp.fr

out, the integrated balances resulting from step (i) still contain contributions that cannot be expressed in terms of the vertically-averaged velocity U and the flow depth D . When the flow is almost horizontal, some of them can be neglected altogether. Others, such as the advection of momentum or the pressure gradient, must be approximated by functions of U and D . To approximate the basal friction, for instance, one still invokes the Darcy-Weisbach equation (1.1)—or some refined version of it. This step, although straightforward if one is content with the leading order of the approximation, becomes a delicate mathematical art when higher-order terms are needed (Ruyer-Quil & Manneville 2000).

Even at their simplest, however, the shallow-water equations prove wonderfully rich. They provide a first-order model for shallow water waves (Stoker 2011), tides (Gallagher & Munk 1971) and tidal bores (Chanson 2012), or the propagation of a tsunami (Popinet 2011). Still, the approximation they are based on limits their validity: A classic failure of theirs is that they cannot represent the propagation of a solitary wave (Korteweg & De Vries 1895). A variety of improvements now allow the shallow-water equations to better account for the streamwise diffusion of momentum (Ruyer-Quil & Manneville 2000; Audusse *et al.* 2011; De Vita *et al.* 2020), but its crosswise counterpart has attracted much less attention, probably because of its lesser importance for water waves (Marche 2007; Chauvet *et al.* 2014).

Yet, when all the other terms vanish, the flux of momentum across a shallow flow becomes obvious—even to the naked eye of an observer walking on a bridge. Indeed, in a straight rectangular channel, the downstream velocity of the flowing water decreases near the banks, and vanishes along them (Nezu *et al.* 1994; Chauvet *et al.* 2014). It is, of course, the turbulent transfer of momentum which propagates the influence of the banks across the flow, but the classical shallow-water equations cannot account for this mechanism. If one is only interested in the total discharge of the flow, an empirical law that takes into account the hydraulic radius of the channel, such as the celebrated formula of Manning (1891), can be substituted for equation (1.1). This, however, will not provide any detail about the flow profile, nor about the distribution of shear stress over the channel’s bed—a quantity crucial to the design of canals (Lacey 1930; Glover & Florey 1951), and to the formation of rivers (Parker 1978).

Alluvial rivers convey not only water, but also the sediment out of which they make their own bed—a natural fluid-structure interaction. The locus of this interaction is the bed surface, onto which the flow applies the stress that transports the sediment grains. To understand how an alluvial river selects its own size and shape, we thus need to understand how it distributes momentum over its bed (Parker 1978). Unfortunately, most rivers find themselves right at the threshold for sediment transport, which makes them dreadfully sensitive to the details of this distribution (Henderson 1961; Devauchelle *et al.* 2011; Métivier *et al.* 2017; Phillips & Jerolmack 2019). Near this critical point, indeed, a minute inaccuracy in the shear stress translates into a first-order error on the sediment flux. It is this demanding problem that motivates the present contribution, although the results presented here should apply to other systems as well.

The flow of water in canals and rivers is turbulent, and the shallow-water equations are classically devised for such flows. In the present paper, nonetheless, we focus on laminar flows, for two reasons. First, in the turbulent regime, the Navier-Stokes equations have no simple solution against which we could test our approximate theory. Second, the basic mechanism by which a river forms its

own channel does not rely on turbulence (although inertia drives the growth of some bedforms, Charru & Hinch 2006). One can indeed produce laminar rivers in the laboratory which, like their natural counterparts, prove highly sensitive to the distribution of the flow-induced stress (Seizilles *et al.* 2013; Abramian *et al.* 2019*b*, 2020).

When applied to thin, laminar films, the shallow-water approximation becomes the lubrication theory (Goodwin & Homsy 1991; Lister 1992). It applies to a variety of systems that ranges from the coating of solids (Levich & Landau 1942; Snoeijer *et al.* 2008) to fingering in free-falling viscous films (Huppert 1982; Craster & Matar 2009). It accounts for the Kapitza instability, which generates roll waves on window panes under heavy rain (Kapitza 1948; Benjamin 1957; Yih 1963), as well as for the slow, viscous intrusion of magma into the Earth’s crust and the formation of lava domes (Michaut 2011; Huppert *et al.* 1982; Stasiuk & Jaupart 1997). The ice sheets of Greenland and Antarctica creep over thousands of kilometers of land before reaching the ocean and, although their thickness can reach up to a few kilometers, their horizontal extension makes the “shallow-ice” approximation a suitable representation of their sluggish flow (Schoof & Hewitt 2013).

None of the above phenomena rely too heavily on the transfer of momentum across the extended (often horizontal) dimensions of the flow (at least to leading order); the classical lubrication theory therefore suits them well. Such is not the case, however, in a laminar laboratory river, where the crosswise transfer of momentum affects the distribution of shear stress over the sediment bed, and thus the equilibrium shape of the river (Abramian *et al.* 2019*a*). Likewise, ice sheets sometimes funnel their flow into narrow ice streams, in which the crosswise velocity gradient becomes strong enough to redistribute momentum horizontally (Suckale *et al.* 2014; Schoof & Mantelli 2021).

Marche (2007) proposed an improved version of the two-dimensional, shallow-water equations, which accounts for the horizontal flux of momentum in an inertial, non-stationary flow. To our knowledge, the horizontal flux of momentum due to viscosity, as represented in this model, has never been tested *per se*. In the present paper, we consider a simple configuration in which this can be done, and which is relevant for laminar rivers: a steady, viscous flow that is uniform downstream (figure 1). In this idealized configuration, the velocity aligns with the channel, and we can often derive an analytical expression for it (or at least a simple numerical solution). We take advantage of this simplicity to propose a new expression for the viscous flux of momentum across a shallow, laminar flow (section 2), and compare it to the exact, two-dimensional flow it should approximate (section 3). In this simplified context, we find that the new expression is more reliable than that of Marche (2007) (appendix A). Because it inherits the non-locality of the Stokes flow it approximates, the equation we propose accounts surprisingly well for the flow near the bank of a channel (section 4).

2. Momentum balance

Rivers are turbulent, and so are most of their laboratory analogues (Métivier *et al.* 2017). Some experimental flumes, however, are small enough, and the fluid they carry viscous enough, for the flow to remain laminar (Seizilles *et al.* 2013; Abramian *et al.* 2020). Provided their path does not bend too much, we are

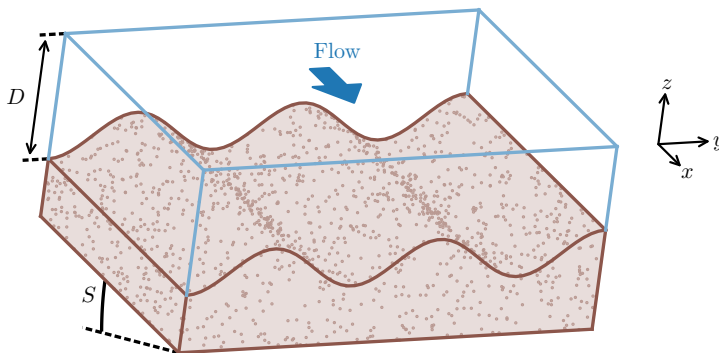


Figure 1: Typical flow configuration and notations. Flow is along the x direction, which is inclined with respect to horizontal. Free surface is flat ($z = 0$), and inclined downstream with slope S . Blue lines: fluid. Brown shading and dots: solid, impervious substrate.

left with a two-dimensional Poisson equation to account for the viscous diffusion of momentum across the flow (section 2.1). Integrated vertically, this equation becomes a one-dimensional equation (section 2.2) which, in its simplest form, yields the classical lubrication approximation (section 2.3). Finally, in section 2.4, we add the contribution of the crosswise stress to this approximation.

2.1. Continuity equation for a Poiseuille flow

We consider the laminar flow of figure 1, whose unidirectional velocity lies along the x axis. This configuration implies that the channel's bed, and thus the flow, do not vary in the x direction at all. Of course, such an idealized setting can only be an approximate representation of the actual flow, valid when the latter varies only weakly along the x direction.

To formalize this assumption, we call \mathcal{L} the crosswise extent of the channel (y axis), and \mathcal{L}_s the characteristic length over which the bed changes downstream (x axis). When the ratio $\mathcal{L}/\mathcal{L}_s$ is small enough, we may treat the flow as streamwise invariant, and represent it with its only component, $u(y, z)$. We use this approximation from now on. Under this assumption, the Stokes equation reduces to:

$$\nu \nabla^2 u + gS = 0, \quad (2.1)$$

where ν is the kinematic viscosity of the fluid, g the acceleration of gravity, and S the downstream slope of the channel (the sine of the angle it forms with the horizontal). The Laplacian operator, ∇^2 , applies only to directions orthogonal to the flow, that is, y and z .

A convenient consequence of our initial assumption that the flow is invariant along the x axis is that its free surface is also invariant. This assumption is very limiting: it precludes any free-surface wave and, in fact, any pressure gradient other than hydrostatic. In short, it limits the present theory to two-dimensional Poiseuille flows. This, we argue, is nonetheless a rich class of systems, to which the straight laminar rivers of Abramian *et al.* (2020) belong.

We further neglect surface tension, and any other stress on the free surface. Unlike the unidirectional-flow assumption, this hypothesis could well be relaxed in the present framework but, to keep our problem simple, we here assume that the free surface is horizontal (more exactly, it belongs to the (x, y) plane). We then place the origin on the free surface, which thus corresponds to $z = 0$. Accordingly,

$$\frac{\partial u}{\partial z} = 0 \quad \text{for } z = 0. \quad (2.2)$$

Finally, we treat the bed as an impervious wall, where the velocity of the fluid vanishes:

$$u = 0 \quad \text{for } z = -D(y), \quad (2.3)$$

where $D(y)$ is the flow depth. This assumption is only a rough approximation of reality when the bed is made of a granular sediment, into which the velocity profile can penetrate. One could account for this penetration by allowing some slip at the porous surface (Beavers & Joseph 1967). Similarly, in the context of flowing ice, a slippery bed surface would also translate into a Robin boundary condition (Schoof & Hewitt 2013). Here, for simplicity, we just keep equation (2.3) until Appendix C.2, where we touch upon the subject of slip.

Equation (2.1), and the associated boundary conditions (2.2) and (2.3), represent a Nusselt flow down a channel of arbitrary section (Benjamin 1957). They sometimes have closed-form solutions, for instance when the channel is elliptic (section 3.1), but in general they do not, and one needs to approximate their solution with a series expansion, or some other numerical method (sections 3.2 and 3.4). Here, equations (2.1) to (2.3) serve as our reference, against which we can evaluate the new shallow-water approximation we introduce in section 2.4. In that sense, we refer to them as “the exact equations” although, of course, they are but an approximation of the actual flow.

2.2. Integrated momentum balance

Integrating the exact momentum balance (2.1) over a vertical section, and invoking the free-surface boundary condition (2.2), we find

$$\nu \int_{-D}^0 \frac{\partial^2 u}{\partial y^2} dz - \frac{\tau_z}{\rho} + gSD = 0, \quad (2.4)$$

where ρ is the density of the fluid and τ_z is the vertical component of the viscous stress on the bed’s surface. In a Newtonian fluid, the latter reads

$$\tau_z = \rho\nu \left. \frac{\partial u}{\partial z} \right|_{z=-D}. \quad (2.5)$$

A more usual notation for τ_z would be τ_{xz} , but in the present context the shorthand notation τ_z leaves no ambiguity.

To rewrite the integral in equation (2.4), we first remember that the free surface is flat and, accordingly, differentiate only the lower bound of the integral along the bed’s surface:

$$\int_{-D}^0 \frac{\partial^2 u}{\partial y^2} dz = \frac{\partial}{\partial y} \int_{-D}^0 \frac{\partial u}{\partial y} dz - \frac{\partial D}{\partial y} \left. \frac{\partial u}{\partial y} \right|_{z=-D}. \quad (2.6)$$

We then use the no-slip boundary condition (2.3) which, again, we differentiate

across the stream:

$$\left. \frac{\partial u}{\partial y} \right|_{z=-D} = \left. \frac{\partial D}{\partial y} \frac{\partial u}{\partial z} \right|_{z=-D}. \quad (2.7)$$

The above relation tells us that the vertical component of the shear stress, τ_z , its crosswise counterpart, τ_y , and the norm of the shear stress, τ , are all related to each other through

$$\tau \equiv \sqrt{\tau_y^2 + \tau_z^2} = \tau_z \sqrt{1 + \left(\frac{\partial D}{\partial y} \right)^2}. \quad (2.8)$$

Since the above relation is exact, and since the shape of the channel, $D(y)$, is a given of the problem, we can treat τ_z and τ as equivalent quantities. Hereafter, based on this observation, we will use τ_z or τ interchangeably, depending on which one is more convenient, or more telling.

The no-slip boundary condition (2.3) also allows us to write

$$\int_{-D}^0 \frac{\partial u}{\partial y} dz = \frac{\partial}{\partial y} \int_{-D}^0 u dz, \quad (2.9)$$

which, together with equation (2.7), we inject in the vertical integral of the Poisson equation, equation (2.4). We finally find

$$\rho\nu \frac{\partial^2}{\partial y^2}(UD) - \tau_z \left(1 + \left(\frac{\partial D}{\partial y} \right)^2 \right) + \rho gSD = 0, \quad (2.10)$$

where we have defined the vertically-averaged velocity U as

$$U = \frac{1}{D} \int_{-D}^0 u dz. \quad (2.11)$$

The one-dimensional momentum balance (2.10) follows from equations (2.1) to (2.3) without any additional assumption—in that sense, it is exact. The first term is the divergence of the cross-stream flux of momentum, the second one is the momentum lost to friction with the bottom, and the last one is the source of downstream momentum, powered by gravity.

A similar balance could be established without assuming that the free surface is flat, and the flow unidirectional; it would then explicitly couple the transverse flow to the deformations of the free surface. Assuming the flow to be unidirectional thus simplifies equation (2.10), but it makes the present theory unable to handle bedforms (unless they are invariant along the x direction, section 3.3).

Even in the simple form of equation (2.10), the integrated momentum balance is incomplete: it involves two unknown functions of the crosswise coordinate y (the shear stress τ_z and the average velocity U). This was to be expected, since we just integrated the Poisson equation vertically, without trying to solve it. At this point, unfortunately, we can postpone no further resorting to approximation.

2.3. Classical lubrication approximation

The lubrication theory relies on the flow being shallow. To quantify this statement, we define a small parameter, ϵ , as the aspect ratio of the channel's cross section:

$$\epsilon = \frac{\mathcal{H}}{\mathcal{L}} \quad (2.12)$$

where \mathcal{H} is the characteristic depth of the channel. Upon rescaling of the y and z coordinates with \mathcal{L} and \mathcal{H} respectively, the crosswise derivative in the exact Poisson equation (2.10) is multiplied by ϵ^2 (Appendix B.1). Assuming ϵ is small, one can thus neglect the flux of momentum that viscosity transfers across the flow. The classical lubrication approximation is based on this idea.

Neglecting the crosswise flux of momentum in the Poisson equation (2.10), we find the classical velocity profile of a Poiseuille flow,

$$u = \frac{3U}{2D^2} (D^2 - z^2). \quad (2.13)$$

where the vertically averaged velocity U is related to the vertical component of the shear stress on the bed, τ_z , through

$$\tau_z = \frac{3\rho\nu U}{D}. \quad (2.14)$$

Since the momentum that gravity delivers to the fluid does not propagate in the crosswise direction, the surface of the bed needs to absorb it all, and thus:

$$\tau_z = \rho g S D. \quad (2.15)$$

Equations (2.13) to (2.15) are exact when the channel is perfectly flat and infinitely wide ($\epsilon = 0$). They are also the leading-order term in the expansion of the flow as a power series in ϵ (Appendix B.1). The gist of the classical lubrication approximation is to assume that these leading-order equations provide a decent approximation of the flow when D varies slowly across the channel. In the case of a unidirectional flow in a straight channel, this simply translates into equations (2.13) to (2.15)—only with varying coefficients $U(y)$ and $D(y)$.

The main advantage of this approach is its wonderful simplicity. The cost of this simplicity, however, is that the resulting velocity $U(y)$ violates the momentum balance (2.10) at order ϵ^2 . In the next section, we improve the lubrication approximation by proposing a representation of the flow that satisfies it exactly.

2.4. Revised lubrication theory

We would like to improve the classical lubrication approximation so that it accounts for the momentum balance, while keeping the problem low-dimensional. Our objective is to predict the distribution of shear stress on the channel's bed better than the classical theory, at a small computational cost. This endeavor is in line with the work of Marche (2007), although here we limit ourselves to the configuration of a steady flow in a straight channel.

To refine an approximate theory based on a series expansion, it is customary to look for the next term in the expansion—here, the term of order ϵ^2 . Appendix B.1 is devoted to this standard method which, as expected, yields a correction to the classical lubrication approximation. Here, however, we propose a less conventional method based on the ansatz of the classical theory, in the hope of obtaining a more flexible representation of the flow. The expansion of Appendix B.1 will be our touchstone: to qualify as progress, the revised theory should be correct at least up to order ϵ^2 .

The reasoning we propose does not have the mathematical rigor of an order expansion, but it is extremely simple. It relies on the observation that it is only by enforcing equation (2.15) that we make equation (2.14) violate the momentum balance (2.10). In itself, equation (2.14) only links the average velocity,

U , with the shear stress on the bed, τ_z , when the vertical velocity profile is parabolic (2.13). In short, equation (2.14) just relates two unknowns, and thus spares one degree of freedom.

It is, we believe, in the spirit of the shallow-water approximation to make the most of this freedom, by dropping equation (2.15) while keeping equation (2.14). Specifically, we use the shape of the vertical profile (2.13) as an ansatz which relates the average velocity, U , to the shear stress, τ_z , through equation (2.14). In practice, this means substituting equation (2.14) into the momentum balance (2.10), which then reads

$$\frac{1}{3} \frac{d^2}{dy^2} (D^2 \tau_z) - \left(1 + \left(\frac{dD}{dy} \right)^2 \right) \tau_z + \rho g S D = 0. \quad (2.16)$$

This second-order, ordinary differential equation is the substitute we propose for equation (2.15). Its only unknown is τ_z . The flow depth, $D(y)$, although it appears at various orders of derivation, is the given boundary to which the flow adjusts.

Equation (2.14), which relates the friction on the bottom to the average velocity, is essentially a friction law. The approximate momentum balance (2.16) depends on this friction law, and it is therefore just as reliable as the latter (Appendix C).

Before we set off to investigate the validity of equation (2.16), we first acknowledge some of its encouraging features. First, it becomes equation (2.15) again over a flat bed—as it should. More interestingly, it is equivalent to the exact Poisson equation (2.1) up to order ϵ^2 , which ensures that it is an improvement upon the classical (zeroth order) theory (Appendix B.2). Finally, that it is a differential equation instead of an algebraic one is only slightly inconvenient, since equation (2.16) is linear in τ_z , and therefore amenable to classical resolution methods. We still need to check, however, that this equation is worth the effort of solving it.

The classical lubrication theory merely equates the vertical component of the shear stress, τ_z , with the momentum that gravity injects into a vertical slice of the flowing fluid, $\rho g S D$. Thus, not only does it neglect the transfer of momentum across the stream, but also approximates the flux of momentum into the bed with its vertical component, τ_z . The revised theory does neither. Indeed, integrating equation (2.16) between two points across the stream, say y_1 and y_2 , and invoking equation (2.7), we find:

$$\left[\frac{1}{3} \frac{d}{dy} (D^2 \tau_z) \right]_{y_1}^{y_2} - \int_{s(y_1)}^{s(y_2)} \tau ds + \rho g S \int_{y_1}^{y_2} D dy = 0. \quad (2.17)$$

where s denotes the arclength along the bed's surface (this change of variable is based on equation (2.8)). The first term in the above equation represents the crosswise flux of momentum, the second one is the total momentum transferred to the bed, and the last one is the momentum injected into the flow by gravity. The first term is an approximation, whereas the two others are exact. Regardless of the correctness of the first term, however, the three of them balance each other exactly—provided τ_z fulfills equation (2.16).

In other words, equation (2.16) is a proper continuity equation for momentum. The downstream momentum that gravity constantly supplies to the flow is either transmitted to the bed, or distributed sideways by viscosity. To close this balance, therefore, one needs to account for the momentum that is transferred across the stream. The value of the shear stress at a specific location on the bed thus depends

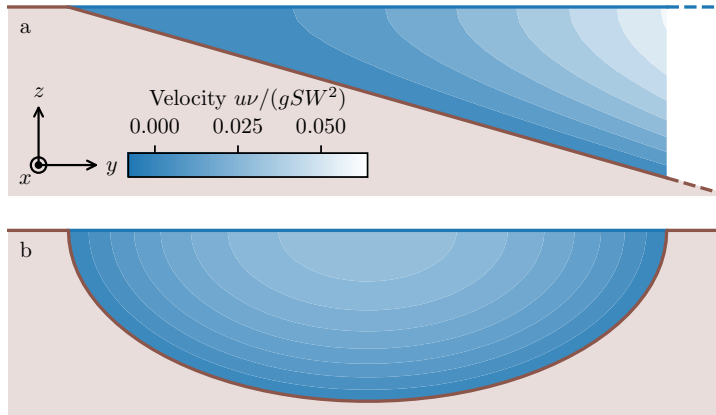


Figure 2: Exact solutions to both the two-dimensional Poisson equation (2.1), and the revised lubrication theory, equation (2.16). Aspect ratio is preserved. Blue shading show velocity contours. The width of the channel is W . (a) Wedge flow, equation (3.1) with $\mu = 1/3$. The flow extends to $y \rightarrow +\infty$ on the right-hand side. (b) Elliptic channel, with aspect ratio $R = 3.5$, equation (3.4).

on the shape of the entire channel. In mathematical terms, equation (2.16) is non-local—a property inherited from the Poisson equation (2.1).

As encouraging as it may seem, the above features do not guarantee that equation (2.16) is a better approximation of the flow than the expansion of Appendix B.1, which is a local, algebraic equation: both are correct up to order ϵ^2 . We devote the next section to testing equation (2.16) in a variety of—hopefully instructive—configurations.

3. Test cases

Once the cross-section of the channel is fixed by choosing $D(y)$, we can solve equation (2.16), and therefore estimate the bottom shear stress, τ_z , according to the revised lubrication theory. To evaluate how accurate this approximation is, we need to solve the exact equation (2.1), and compute the associated shear stress. When the channel's shape is suitably simple, there are closed-form solutions of the Poisson equation (section 3.1). In a rectangular channel, we can still expand the exact solution into a series (section 3.2), but when the channel's bottom is corrugated, we need to resort to linearization (section 3.3) or finite elements (section 3.4) to solve the problem of reference.

3.1. Exact solutions

3.1.1. Wedge flow

Let us first consider a Poiseuille flow in a wedge (figure 2a). This configuration may represent a river bank, where the water surface ($z = 0$) intersects the bed ($z = -\mu y$). Mathematically, this problem is ill-posed, since it lacks a boundary condition on the open side of the wedge. We could fix such a condition, of course, but this would rule out any closed-form solution. Close enough to the corner, anyway, we expect that the flow velocity will be insensitive to the boundary

condition, and will thus behave like

$$u = \frac{gS}{\nu} \frac{(\mu y - z)(\mu y + z)}{2(1 - \mu^2)} = \frac{gS}{\nu} \frac{D^2 - z^2}{2(1 - \mu^2)}, \quad (3.1)$$

an expression that satisfies equation (2.1) exactly (except for $\mu = 1$, section 4). Based on this expression, the shear stress the flow exerts on the bed's surface is proportional to the water depth, μy :

$$\tau_z = \frac{\rho g S \mu y}{1 - \mu^2}. \quad (3.2)$$

We now need to evaluate the revised lubrication approximation against this result. This is straightforward: as it turns out, equation (3.2) is also a solution of equation (2.16). This is not true for the classical lubrication approximation since, in a wedge, equation (2.15) reads $\tau_z = \rho g S \mu y$. The revised theory thus proves an improvement in this case, at least when the bank is not too steep ($\mu < 1$; section 4 is devoted to $\mu \geq 1$).

3.1.2. Elliptic channel

We now turn our attention to a better-defined problem, and consider an elliptic channel of width W , and aspect ratio R (figure 2b), the depth of which reads

$$D = \frac{1}{R} \sqrt{W^2 - 4y^2}. \quad (3.3)$$

The corresponding solution of the Poisson equation (2.1) is (Boussinesq 1868, p. 388)

$$u = \frac{gS}{2\nu} \frac{(W^2 - 4y^2 - (Rz)^2)}{(R^2 + 4)}. \quad (3.4)$$

Differentiating the above equation, we find the expression of the vertical component of the shear stress on the bed's surface:

$$\tau_z = \frac{\rho g S R \sqrt{W^2 - 4y^2}}{R^2 + 4}. \quad (3.5)$$

Again, this expression happens to be an exact solution to equation (2.16)—the approximation could not be more accurate. In fact, the velocity field associated with the revised lubrication approximation, namely

$$u = \frac{D}{2\rho\nu} \left(R^2 - \left(\frac{z}{D} \right)^2 \right) \tau_z, \quad (3.6)$$

is the exact solution of the original Poisson equation. This result, which seems coincidental at first, appears inevitable after careful consideration. Indeed, in an elliptic channel, the vertical velocity profile of the exact solution is parabolic, making the ansatz of the revised theory a perfect representation of the flow. Conversely, the classical lubrication theory overestimates the shear stress by a factor of $(R^2 + 4)/R^2$.

Although limited to channels of specific shapes, the perfect match between the revised theory and the reference problem is encouraging. In particular, it is remarkable that the revised theory works so well in an elliptic channel of arbitrary aspect ratio, where the hypothesis of a shallow flow does not hold at all. In the next section, we test the revised theory under even worse conditions.

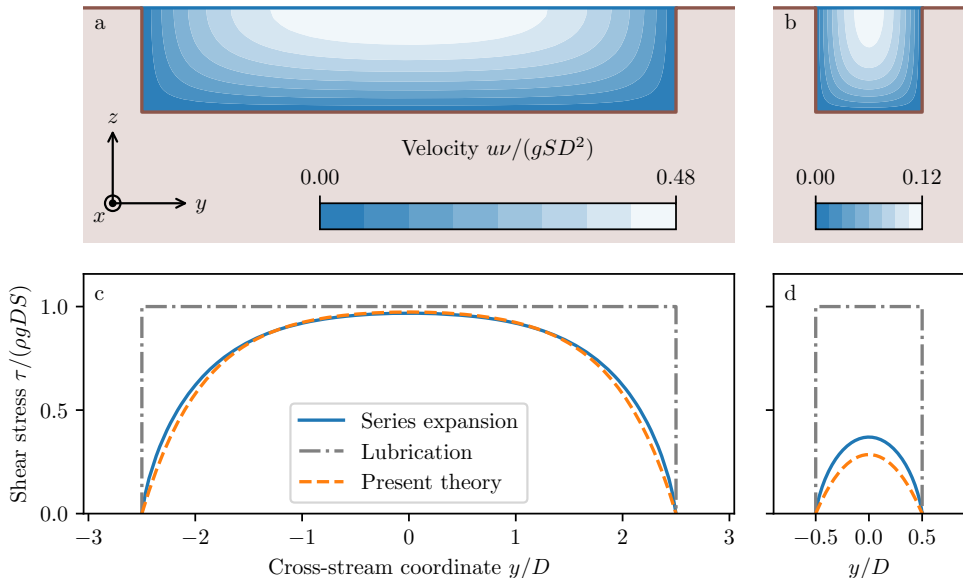


Figure 3: Laminar flow in a rectangular groove. (a,b) Streamwise velocity (blue shading), according to series expansion (3.7) truncated at $k = 99$. Aspect ratio is preserved. (c,d) Intensity of flow-induced shear stress on bottom (τ). Solid blue: series expansion (3.7) truncated at $k = 99$, for reference. Dash-dotted grey: classical lubrication approximation (2.14). Dashed orange: present theory (3.8). Channel aspect ratio is $W/D = 5$ (a,c) and $W/D = 1$ (b,d).

3.2. Rectangular groove

Let us, for a change of context, imagine a straight microfluidic groove engraved into a solid substrate, with a rectangular cross-section (figure 3a,b). The flat bottom of the channel, of course, would be no challenge for the lubrication theory, were it not for the two side walls, where the no-slip boundary condition applies. There is no closed-form solution of the Poisson equation in a channel of rectangular cross-section, but there exists a solution in the form of an infinite series (e.g. White 1991):

$$u = \frac{4gSW^2}{\nu\pi^3} \sum_{k=1,3,5,\dots}^{\infty} \frac{(-1)^{(k-1)/2}}{k^3} \left(1 - \frac{\cosh(k\pi z/W)}{\cosh(k\pi D/W)} \right) \cos\left(\frac{k\pi y}{W}\right) \quad (3.7)$$

where W is the width of the channel, and D its (uniform) depth. In this sum, the index k can be interpreted as a dimensionless wavenumber. Differentiating the above expression with respect to z yields the bottom shear stress τ_z with arbitrary precision (blue line in figure 3c,d, τ is exactly τ_z in this case). As expected, we find that viscosity conveys the influence of the walls across the flow, making the bottom shear stress reach a maximum in the middle of the channel.

The classical lubrication theory, in the form of equation (2.15), cannot account for the no-slip boundary condition at the side walls. In fact, according to this approximation, the shear stress just remains constant across the entire channel. As a consequence, although this constant is a reasonable estimate of the shear stress

far away from the walls, the classical theory fails entirely in their neighborhood (grey line in figure 3c,d).

Equation (2.16), conversely, requires two boundary conditions, which allows us to fix the velocity of the flow along the side walls: Equation (2.3) translates the no-slip condition into $\tau_z = 0$ in $y = \pm W/2$. The uniform depth of the channel makes this problem a textbook exercise, whose solution is

$$\tau_z = \rho g S D \left(1 - \frac{\cosh(\sqrt{3}y/D)}{\cosh(\sqrt{3}W/(2D))} \right). \quad (3.8)$$

We find that this expression is a much better approximation of the actual shear stress than the constant of the classical theory (orange line in figure 3c,d). In particular, the side walls affect the flow over a distance comparable to the channel's depth. In a channel of aspect ratio $W/D = 5$ (figure 3a,c), the largest error of the present theory is less than 10% of the average shear stress (the average error is about 2%). Combining equation (3.8) with the ansatz (2.13), and integrating the result across the channel, yields an estimate for the total discharge of a rectangular microfluidic groove; we find that this estimate lies within less than 3% of the actual value, whereas the classical theory is more than 30% off.

The revised lubrication theory relies on the friction law (2.14), which derives from the assumption that the flow profile is essentially that of a Nusselt film. Near the walls, however, this hypothesis breaks down, because the walls cause the profile to depart from its parabolic shape, thus affecting the friction law (Appendix C.1). It turns out, however, that the vicinity of a wall also lessens the contribution of bottom friction to the momentum balance. As a result, the friction term gets wrong where it does not matter much, which bolsters the validity of the revised lubrication theory near a wall.

In a narrower channel ($W/D = 1$, figure 3b,d), we expect both the classical and the revised theories to fail. Indeed they do, but not to the same extent. The revised theory underestimates the discharge by about 12%, whereas the classical theory overestimates it by a factor of almost 5. As visible on figure 3d, this difference results from the ability of the revised theory to account for friction along the sidewalls (Appendix B.1).

The non-locality of equation (2.16) comes at a cost but, as this case illustrates, it keeps the momentum balance under control. This, we argue, is why the revised theory is more reliable than the classical theory.

3.3. Linear perturbation

We now consider a channel of infinite width, the bottom of which is perturbed with a sinusoidal corrugation of infinitesimal amplitude (figure 1). In tune with our initial assumption, the flow remains invariant in the x direction; the crests and troughs of the corrugation, therefore, are aligned with the flow velocity. Mathematically,

$$D = D_0 + \text{Re} \left(D_1^* \exp \left(\frac{iky}{D_0} \right) \right) \quad (3.9)$$

where D_1^* is the (possibly complex) amplitude of the perturbation ($|D_1^*| \ll D_0$), D_0 is the unperturbed depth of the channel and k is the dimensionless wavenumber of the perturbation (the corresponding wavevector is aligned with

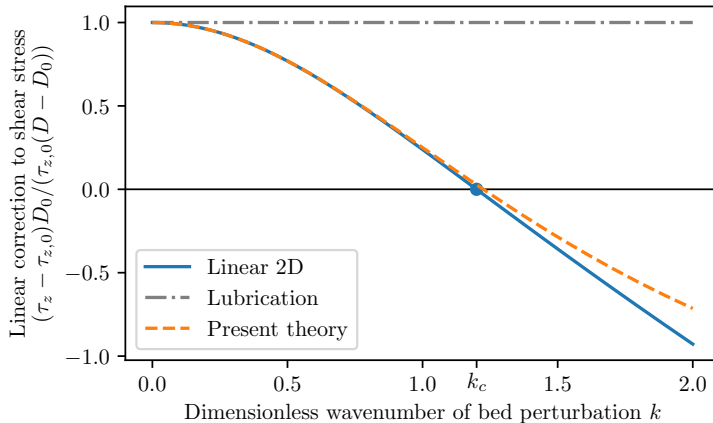


Figure 4: Spectral response of bottom shear stress to a bed perturbation of infinitesimal amplitude. Blue line: linearized two-dimensional solution (3.10) for reference (Abramian *et al.* 2019a). Blue dot: critical wavenumber $k_c \approx 1.20$. Dash-dotted grey line: classical lubrication approximation (3.11). Orange dashed line: present theory (3.13).

the y axis). When the perturbation vanishes, the bottom shear stress, $\tau_{z,0}$, is related to D_0 through equation (2.15).

Abramian *et al.* (2019a) investigated the stability of a similar perturbation when the bed over which the fluid flows is made of mobile sediment. Since it is the flow-induced shear stress that drives sediment transport in this case, one essential step of their analysis was to linearize the exact Poisson equation (our equation (2.1)), and to analytically solve the resulting linear problem. Their equation (3.7) will serve as our reference; in the present notations, it reads:

$$\tau_z = \tau_{z,0} \left(1 + (1 - k \tanh k) \operatorname{Re} \left(\frac{D_1^*}{D_0} \exp \left(\frac{iky}{D_0} \right) \right) \right). \quad (3.10)$$

The above equation indicates that the shear stress is in phase with the sinusoidal perturbation; its extrema lie where the perturbation's are. More remarkably, the sign of the shear-stress perturbation depends on the wavelength of the bed perturbation (blue line in figure 4). When this wavelength is large enough ($k < k_c$, where $k_c \approx 1.20$), the shear stress is at its highest in the troughs, where the flow is deeper—in tune with the classical lubrication theory. On the contrary, when the perturbation's wavelength is short enough ($k > k_c$), the crests of the corrugation find themselves more exposed to the bulk of the flow, and are therefore subjected to a stronger stress. In other words, the crests then shield the troughs, which collect a smaller share of the flow's momentum. This mechanism, the signature of the Laplacian in equation (2.1), is essentially the same as the shadowing that sets the shape of diffusion-limited aggregates (Witten Jr & Sander 1981) and causes the Saffman-Taylor instability (Saffman 1986). It is crucial to the stability analysis of Abramian *et al.* (2019a), as it prevents the perturbations with the shortest wavelengths to grow infinitely fast, thus selecting the wavelength most likely to appear in an experiment.

The classical lubrication theory cannot account for this phenomenon. Indeed, it directly relates the shear stress to the amplitude of the perturbation, regardless

of its wavelength:

$$\frac{\tau_z - \tau_{z,0}}{\tau_{z,0}} = \frac{D - D_0}{D_0} = \operatorname{Re} \left(\frac{D_1^*}{D_0} \exp \left(\frac{iky}{D_0} \right) \right). \quad (3.11)$$

In the Fourier space, this translates into a flat response (grey line in figure 4). The present theory, on the other hand, is designed to account for the crosswise flux of momentum. To compare it to equation (3.10), we first linearize equation (2.16) about a horizontal bed. We find

$$\frac{D_0^2}{3} \frac{d^2}{dy^2} \frac{\tau_{z,1}}{\tau_{z,0}} + \frac{2D_0}{3} \frac{d^2 D_1}{dy^2} - \frac{\tau_{z,1}}{\tau_{z,0}} + \frac{D_1}{D_0} = 0, \quad (3.12)$$

where D_1 and $\tau_{z,1}$ are the linear corrections to the base state of the flow depth and the shear stress, respectively. Applying a Fourier transform to the above expression, we get

$$\frac{\tau_{z,1}^*}{\tau_{z,0}} = \frac{1 - 2k^2/3}{1 + k^2/3} \frac{D_1^*}{D_0}, \quad (3.13)$$

where the symbol $*$ denotes the Fourier amplitude of the quantity it decorates. For long wavelengths ($k \ll 1$), this expression correctly approximates the exact linear result (3.10), but departs from it as the wavelength of the perturbation approaches the flow depth ($k \sim 1$) (orange line in figure 4). Before this happens, however, the present theory accounts for the sign change of the shear stress near k_c , which it estimates at $\sqrt{3/2} \approx 1.22$ —within 2% of the true value.

The quality of the present approximation turns out to be surprisingly good. In a fashion quite typical of shallow-water theories, it works better than it should. Indeed, if we expand equations (3.12) and (3.13) into Taylor series, they match up to order k^4 only (odd orders all vanish); reasonably good, perhaps, but insufficient to account for the sign change in figure 4, since this well-controlled, but truncated, expansion fails to ever change sign, unlike the original equation (3.13). In short, it seems we can push the revised lubrication theory beyond where a well-controlled order expansion endorses it—at our own risks, of course.

Taking such risks, however, would be of little use if the revised lubrication theory were to fail beyond the linear regime, in which the two-dimensional equation it approximates can be solved relatively easily. The next section addresses this point.

3.4. Finite-amplitude corrugation

The revised lubrication theory fairly reproduces the results of the linearized, two-dimensional Poisson equation (section 3.3). We now perturb a flat bed with a corrugation of finite amplitude, thus precluding any linearization of the problem. To test the present theory, we need a reference, which analytical calculations can no longer provide. Instead, we solve equation (2.1) in a corrugated channel with finite elements (FreeFem++, Hecht 2012, blue shading in figure 5(a,b)). Upon differentiation with respect to y and z , these numerical simulations yield an approximation of the flow-induced shear stress on the bed (blue line in figure 5(c,d)).

We consider two configurations, in which the corrugations differ in amplitude and wavelength. In the first case, the wavelength of the corrugation is shorter than the flow depth ($k = 6$), while its amplitude remains small, albeit not vanishing ($D_1^* = 0.03D_0$, figure 5(a,c)). We expect both the classical lubrication theory

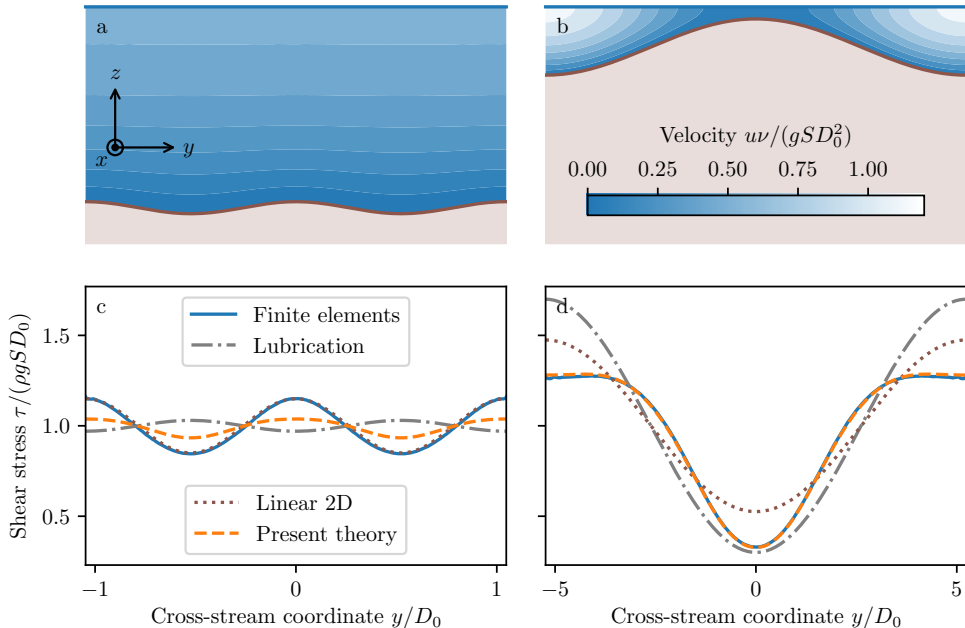


Figure 5: Laminar flow above a corrugated bottom (a,b), and associated shear stress on the bed (c,d). Blue shading in (a,b) shows two-dimensional velocity (2.1) (finite-elements simulations, arbitrary units). Flow is towards viewer. Aspect ratio is preserved. Solid blue lines in (c,d): shear stress from finite-elements simulations. Dotted brown line: linearized two-dimensional flow (3.10) (Abramian *et al.* 2019a). Dashed orange lines: present theory (2.16) (numerical solution). (a,c) $k = 6$ and $D_1^* = 0.03D_0$. (b,d) $k = 0.6$ and $D_1^* = 0.7D_0$.

and the present theory to fail in this case, due to the short wavelength of the perturbation. In the second case, the wavelength is significantly larger than the flow depth ($k = 0.6$), while the amplitude is comparable to the depth ($D_1^* = 0.7D_0$, figure 5(b,d)). It is in such a flow configuration that the revised theory could prove useful.

In the first case, the amplitude of the perturbation is small enough (with respect to depth) for the linearized theory of section 3.3 to provide a good approximation of the shear stress (dotted brown line in figure 5(c)). This is not true, however, for the classical lubrication theory, which wrongly locates the maxima of the shear stress in the troughs (dashed grey line). This was to be expected, since our choice of wavenumber is far beyond k_c , the wavenumber above which the crests shield the troughs (section 3.3). The revised lubrication theory, in the form of a numerical solution of equation (2.16), does a little better, since it locates the maxima correctly, but the amplitude of its prediction is off. For very short wavelengths, therefore, the only alternative to the exact Poisson equation remains the linear, two-dimensional theory.

Conversely, in the second case, the amplitude of the perturbation is too large (with respect to depth) for the linear, two-dimensional theory (dotted brown line in figure 5(d)). Moreover, its wavelength is too short for the classical lubrication

approximation (dash-dotted grey line). The revised lubrication theory, on the other hand, handles this case gracefully: the numerical solution of equation (2.16) matches the finite element simulation of the full two-dimensional problem everywhere within 2.3% of the average stress (dashed orange line). Equation (2.16) thus represents accurately the diffusion of momentum across the flow, which makes it a convenient alternative to the Poisson equation in this case.

This encouraging result would not survive a significant shortening of the perturbation’s wavelength, that is, a reduction of the aspect ratio of the flow. As it is, however, the periodic configuration of figure 5(b) evokes a series of channels of aspect ratio $\pi D_0/(kD_1^*) \approx 7.5$ —a value comparable to that of laboratory rivers (Seizilles *et al.* 2013). In the next section, inspired by these experiments, we consider a configuration in which the corrugation and the free surface intersect, thus forming the bank of a river.

4. Flow along a corner

In section 3.1.1, we noted that the analytical solution of the Poisson equation in a wedge diverged at a special value of the wedge angle, namely $\mu = 1$ —and we quickly moved on. We now consider this problem in details. In fact, this is a typical difficulty with any shallow-water or lubrication theory: when the bottom of the flow reaches the free surface, their intersection creates a singularity. This singular point, or line, often becomes an issue when one numerically simulates a sheet flow over a rough bed that can protrude through the fluid’s surface (Delestre *et al.* 2014).

In this section, true to the form of the paper, we consider a simpler version of this problem: the laminar flow of a viscous fluid in a channel with a sharp corner (figure 6). At the bank, the flow forms a wedge bounded by the bed and the free surface, which intersect each other at an angle β (the slope of the bank is $\mu = \tan \beta$). Away from the corner, the bottom gradually returns to the horizontal, until it reaches an axis of symmetry (dashed blue line on figure 6). The assumption, here, is that the flat part of the bed, and the exact location of the axis of symmetry, do not really matter to the flow near the corner.

We begin with our usual reference, the Poisson equation (2.1), for which we find a variety of asymptotic behaviors near the corner (section 4.1). We then try to identify the same regimes in the revised lubrication theory (section 4.2). The classical lubrication theory, of course, is irrelevant for this problem, since it assumes that the fluid’s velocity is proportional to the local flow depth, and therefore undergoes no transition when $\mu = 1$.

4.1. Poisson equation

The Poisson equation (2.1) is linear and non-homogeneous. It is thus convenient to decompose its solutions into the sum of a homogeneous solution and a special solution. Dropping the source term of the Poisson equation, we are left with the Laplace equation,

$$\nabla^2 u = 0, \tag{4.1}$$

to which we can find solutions in the form of analytical functions of the complex variable $\omega = y + iz$ (the bank is located at $\omega = 0$). Then, we will need to find a special solution to the original Poisson equation. Unless $\mu = 1$, this will simply be equation (3.1), the solution we encountered in section 3.1.1.

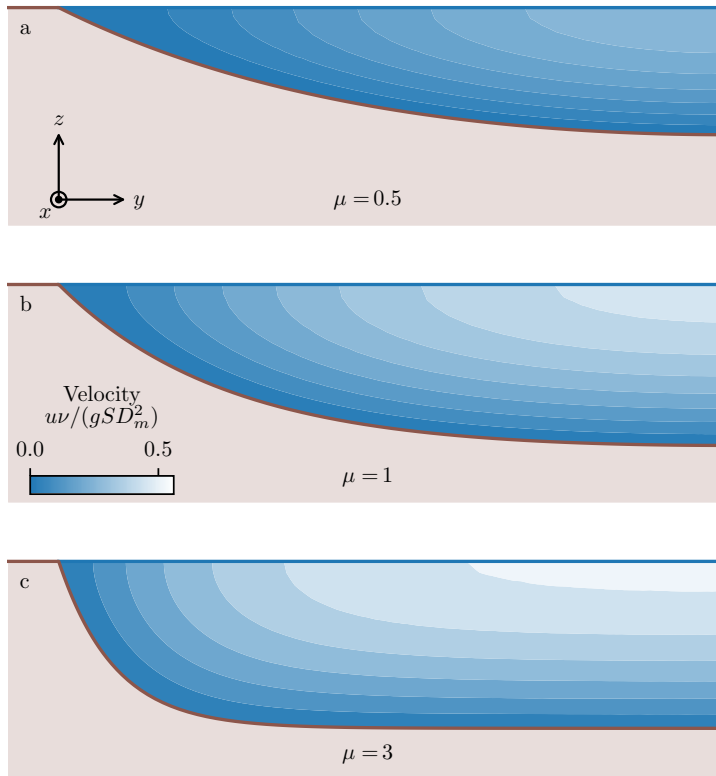


Figure 6: Laminar flow in a channel with a sharp corner (finite-element simulation). Channel shape corresponds to equation (4.4). The tangent of the corner angle is μ . Blue shading: flow velocity.

4.1.1. Homogeneous solution

In a corner of angle β , there exists a power-law solution to the Laplace equation (4.1) that satisfies boundary conditions (2.2) and (2.3). Namely,

$$u^{(h)} = C \operatorname{Re} \left[\omega^{\pi/(2\beta)} \right] \quad (4.2)$$

where $\omega = y + iz$ is the complex coordinate, and C is a positive constant with physical dimensions (Polubarinova-Kochina 1962). Differentiating the above expression, we get the vertical component of the shear stress, τ_z :

$$\tau_z^{(h)} = C' y^{\pi/(2\beta)-1} \quad (4.3)$$

where C' is another positive constant. The exponent of the above expression is already an indication that a slope of $\mu = 1$ might be a special value. Indeed, when $\beta = \pi/4$, the exponent of the above expression is exactly 1—the exponent of the special solution (3.2). Motivated by this observation, we now compare the homogeneous solution (4.3) with the special solution.

4.1.2. *Shallow bank* ($\mu < 1$)

When the bank is shallow enough ($\mu < 1$ or equivalently $\beta < \pi/4$), the special solution (3.1) remains finite and positive, and thus physically acceptable. In a given channel, such as those of figure 6, the actual flow is then the sum of this special solution and a homogeneous solution (4.2), the prefactor of which, C , adjusts to the far field, that is, to the overall shape of the channel.

Provided the corner's angle, β , is less than $\pi/4$, the exponent of the homogeneous solution (4.3) is larger than one, and therefore larger than that of the special solution. As a consequence, the special solution dominates the flow in the bank's neighborhood ($\omega \rightarrow 0$), and the shear stress is well approximated by equation (3.2), in which there is no parameter that adjusts to the far field.

To compare this asymptotic behavior to numerical simulations, we first devise a concrete example of a channel with a bank slope of μ . We expect that, apart from the bank, the general shape of the channel will not alter the asymptotic behavior of the flow in the corner. To fix the channel's shape, we arbitrarily define its cross section as

$$D = D_m \frac{\cosh(\mu(y/D_m - R/2)) - \cosh(\mu R/2)}{\sinh(\mu R/2)} \quad (4.4)$$

where D_m sets the depth of the channel, and R sets its aspect ratio (figure 6). In practice, we choose $R = 8$, whereas D_m does not matter once the problem is made dimensionless. We then run finite-element simulations in such a channel to numerically approximate the velocity field u , which is a solution to the Poisson equation (2.1) (blue shadings in figure 6). By differentiating this velocity, we find the shear stress on the bed, τ_z . For a shallow bank ($\mu < 1$), we find good agreement between the numerical solution (solid blue line in figure 7) and the special solution (3.2) (dashed blue line in figure 7), without adjusting any parameter. This shows that, near a shallow bank, the flow is indeed dominated by the special solution, the asymptotic behavior of which is independent of the rest of the flow (the far field).

However, as the bank's angle approaches 45° ($\mu = 1$), equation (3.1) breaks down, and the flow near the bank takes another form.

4.1.3. *Steep bank* ($\mu > 1$)

When the bank is steep ($\mu > 1$), the special solution (3.2) yields a negative shear stress in the corner—quite unrealistic. This is an indication that, in this case, we cannot neglect the homogeneous solution anymore.

In fact, the homogeneous solution (4.2) takes over the flow near the bank. Mathematically, the exponent of the homogeneous solution (4.2) becomes less than one and, therefore, the homogeneous solution overshadows the special solution (3.2) near the bank ($\omega \rightarrow 0$). As a consequence, the leading-order term in the velocity field is now a power law with a varying exponent (namely $\pi/(2\beta)$), which depends on the bank's slope—and so does the shear stress through equation (4.3). The dependence of an exponent on the shape of the boundary is not unheard of; it is in fact quite typical of the Laplace equation (Polubarinova-Kochina 1962). That it happens in a viscous flow is reminiscent of the recirculation loops identified by Moffatt (1964), although the transition we find here is mathematically simpler.

Again, the finite element simulations match this asymptotic regime near the bank although, this time, we need to fit the constant C' to the numerical data (dashed and solid brown lines in figure 7). The value of this prefactor, indeed,

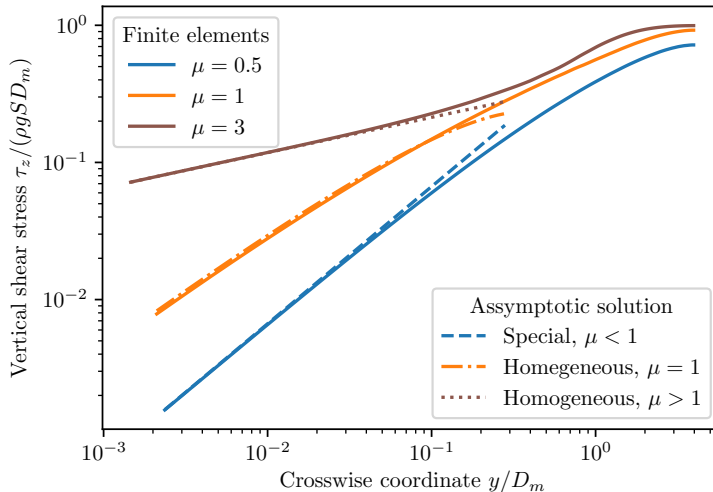


Figure 7: Shear stress along the bed of a channel with sharp corner. Solid lines correspond to the finite-element simulations of figure 6. Dashed blue line is equation (3.2), without any fitted parameter. Dotted brown line is equation (4.3) with coefficient C' fitted to finite-element simulation. Dash-dotted orange line: equation (4.6), with C_b arbitrarily set to D_m .

depends on the shape of the entire channel—the far field matters for the flow along a steep corner, even at leading order.

4.1.4. Intermediate bank ($\mu = 1$)

When the slope of the bank is exactly one ($\beta = \pi/4$), finding a special solution of the Poisson equation (2.1) requires more work. One method is to write the solution as the sum of a radially symmetric term that does not match the boundary conditions, with an analytical term that corrects this mismatch. A simple expression for the first is $-gS\omega\bar{\omega}/(4\nu)$ (the overbar denotes complex conjugation), which naturally satisfies the free-surface boundary condition, but not the no-slip condition on the channel's bed. We now need to find an analytical function that compensates for this shortcoming. There might be a principled method to do so, but we used trial and error to identify, among the usual suspects of complex analysis (power laws and logarithm), the one that suits our problem, namely:

$$u = -\frac{gS}{\nu} \operatorname{Re} \left(\frac{\omega\bar{\omega}}{4} + \frac{\omega^2}{\pi} \log \left(\frac{\omega}{C_b} \right) \right), \quad (4.5)$$

where C_b is an arbitrary positive constant. This constant relates the flow near the corner to the far field. Unlike for a steep bank (section 4.1.3), however, this constant plays only a minor role in the neighborhood of the bank, as $\log \omega$ overcomes $\log C_b$. This was to be expected, perhaps, in the special case that marks the transition between the shallow bank, for which the far field does not matter, and the steep bank, for which the far field dominates the flow.

Differentiating equation (4.5) along the z direction, we finally find an expression

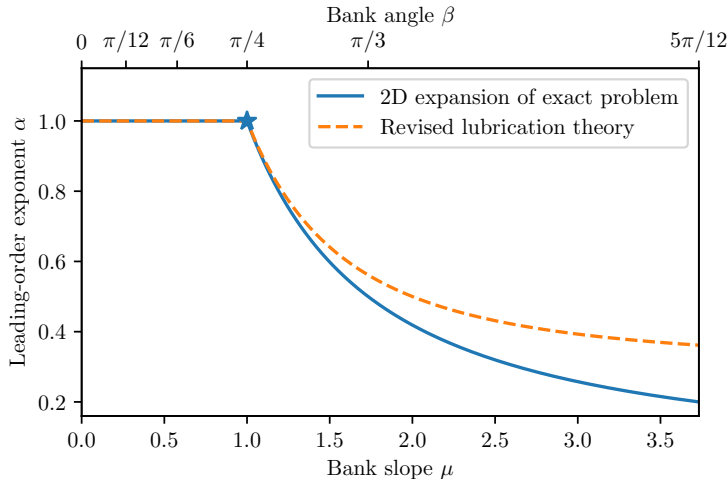


Figure 8: Exponent of the leading order term in the expansion of the shear stress, τ_z , near a bank of angle β . Solid blue line: two-dimensional expansion of the exact problem for reference, equations (3.2) and (4.3). Dashed orange line: revised shallow-water theory, equations (3.2) and (4.13). Blue star: intermediate bank, for which the leading-order term is not a power law, equation (4.6).

for the shear stress near the bank:

$$\tau_z \approx -\frac{2\rho g S}{\pi} y \log\left(\frac{y}{C_b}\right). \quad (4.6)$$

The finite-element simulations match this peculiar scaling, even when the constant C_b is arbitrarily set to D_m (orange lines in figure 7).

To represent the transition from the shallow-bank asymptotic regime to steep-bank regime, we plot the exponent of the leading-order term in the shear-stress expansion, α , as a function of the bank slope, μ (figure 8). After a plateau at $\alpha = 1$ that corresponds to a shallow bank ($\mu < 1$), the exponent suddenly switches to $\pi/(2\beta) - 1$ when the bank gets steeper than 45° (blue line on figure 8). Right at the transition, when the bank's slope is exactly one, the leading-order term is not a power-law and α is thus ill-defined.

Now that we have established the asymptotic behavior of the flow near a bank, we can ask how accurately the revised lubrication theory accounts for this behavior. This is the purpose of the next section.

4.2. Revised lubrication approximation

We now consider the corner flow of section 4.1 in the framework of the revised lubrication theory. The findings of the previous section will serve as a reference, against which we will test the theory introduced in this paper. Near a bank of slope μ , equation (2.16) becomes

$$\frac{1}{3} \frac{d^2}{dy^2} (y^2 \tau_z) - \left(1 + \frac{1}{\mu^2}\right) \tau_z + \frac{\rho g S}{\mu} y = 0. \quad (4.7)$$

Once again, to solve this linear, non-homogeneous equation, we need to find a special solution and a family of homogeneous solutions. In doing so, we hope to find the counterparts of the special and homogeneous solutions of section 4.1.

4.2.1. Homogeneous equation

We begin with the homogeneous equation associated with equation (4.7):

$$\frac{d^2}{dy^2} (y^2 \tau_z) - A \tau_z = 0, \quad (4.8)$$

where, for simplicity, we have introduced $A = 3(1 + 1/\mu^2)$. The solutions of the above equation are the sum of two independent solutions:

$$\tau_{z,h} = C_+ y^{\alpha_+} + C_- y^{\alpha_-} \quad (4.9)$$

where C_{\pm} denote integration constants, and the two exponents read

$$\alpha_{\pm} = \frac{-3 \pm \sqrt{1 + 4A}}{2}. \quad (4.10)$$

Among the above exponents, one is positive ($\alpha_+ > 0$), and the other negative ($\alpha_- < 0$). We discard the latter to ensure that the shear stress, τ_z , remains finite at the bank. We then rewrite the homogeneous solution as

$$\tau_{z,h} = C_+ y^{\alpha_+}. \quad (4.11)$$

At this point, it is tempting to identify α_+ with the exponent of equation (4.3). This comparison, however, makes sense only if equation (4.11) represents the leading-order term of the flow. To find when this is the case, we need to look for a special solution of equation (4.7).

4.2.2. Special solution

From section 3.1.1 we know that, unless the slope of the bank is $\mu = 1$, a special solution of equation (4.7) is equation (3.2)—the very equation that directed our attention to the bank problem in the first place. Again, this solution disappears when $\mu = 1$, just like the special solution of the Poisson equation (section 4.1.1). In the light of section 4.1.4, we expect that we will need to pay some dedicated attention to this singular case.

Now that we possess a special solution to equation (4.7) and a family of homogeneous solutions (4.11), we are well equipped to evaluate their relative importance near the bank.

4.2.3. Shallow bank ($\mu < 1$)

When the bank is shallow enough ($\mu < 1$), equation (4.10) tells us that the exponent of the homogeneous solution, α_+ , is larger than 1. As a result, near the bank ($y \rightarrow 0$), the homogeneous solution becomes negligible with respect to the special solution (3.2). Therefore,

$$\tau_z \sim \frac{\rho g \mu S}{1 - \mu^2} y, \quad (4.12)$$

which is just equation (3.2). At leading order, the shear stress thus grows linearly with the distance from the bank, which translates into a plateau on figure 8 ($\alpha = 1$). This plateau matches exactly the one we found in section 4.1.2 and,

in retrospect, supports the rough reasoning of section 3.1.1. Encouraged by this agreement, we now consider the case of a steep bank.

4.2.4. Steep bank ($\mu > 1$)

When the bank is steep ($\mu > 1$), the shear stress associated with the special solution (3.2) becomes negative. This would be physically unacceptable, if the homogeneous solution did not overcome the special solution. According to equation (4.10), however, we find that the exponent of the special solution, α_+ , is now less than 1. This means that, near a steep bank, the homogeneous solution indeed dominates the flow. Therefore,

$$\tau_z \sim C'_+ y^{\alpha_+}, \quad (4.13)$$

where C'_+ is a constant set by the far field. This expression, together with equation (4.10), approximates well the exact expansion derived in section 4.1.3 when μ is close to one (figure 8). This agreement, however, becomes less accurate as the bank's slope steepens. Beyond the critical slope, the revised lubrication theory gradually becomes less reliable, and can only provide qualitative estimates of the exponent α .

In the limit of a vertical corner ($\mu \rightarrow \infty$), the bed becomes a wall that extends down to infinity. The revised lubrication approximation then fails because the vertical integral of the momentum balance, which yields equation (2.4), diverges. Adding a bottom at some finite depth, however, provides a lower bound to this integral, and allows the approximation to recover (section 3.2). In fact, near the walls of a rectangular channel, the revised lubrication approximation gives a decent estimate of the stress on the bottom (appendix C.1). The quality of the approximation thus depends on the shape of the bottom, rather than the behavior of the flow near the corner—another manifestation of non-locality.

4.2.5. Intermediate bank ($\mu = 1$)

When the bank is at an angle of exactly 45° (that is, $\mu = 1$), the special solution (3.2), breaks down. Instead, we find the following special solution:

$$\tau_z = -\frac{3\rho g S}{5} y \log\left(\frac{y}{C'_b}\right), \quad (4.14)$$

where C'_b is a positive constant which, like C_b , is set by the far field. The exponent of the homogeneous solution is $\alpha_+ = 1$, and equation (4.14) is thus the leading-order term of the shear stress near the bank. This expression matches its two-dimensional counterpart, equation (4.6), but for the value of the prefactor, which we find to be less than 6% off.

Just like in section 4.1.4, the flow near a bank of slope 1 features a logarithmic correction, with an integration constant that adjusts to the far field. That, in the revised lubrication theory, the prefactor of this regime is a bit off is typical of a shallow-water approximation: it is qualitatively correct, but not exact. The error, in this case, vanishes suddenly when the bank's angle becomes less than 45° .

5. Conclusion

For the cost of solving a linear, ordinary differential equation, we can account for the crosswise transfer of momentum across a shallow laminar flow. In all the configurations we have tested, doing so proved an improvement over the classical

lubrication theory, which discards this transfer altogether. The revised theory, however, is only an approximation of the original two-dimensional problem. It fails when the flow varies too abruptly across the stream and, in its present form, cannot account for any variation along the stream. In that sense, one gets what one pays for: the classical lubrication theory requires little effort, but it applies only to very shallow flows. At the other end of the spectrum, one can numerically solve the full, two-dimensional problem to reach an arbitrary accuracy. The revised theory can be a fair compromise.

As is often the case with shallow-water equations, we found that this compromise is a good representation of the flow, while remaining a tractable approximation. The accuracy of this approximation depends entirely on the vertical profile of the flow: the revised theory is reliable when the latter is almost parabolic. When is this true? This is a difficult question in general. To answer it, one could turn the parabolic profile into the first term of a projection onto a suitably chosen basis, and then evaluate the next term (Ruyer-Quil & Manneville 2000)—a task far beyond the scope of the present paper.

The order expansion detailed in Appendix B, nonetheless, shows that the revised theory is accurate up to order ϵ^2 at least, where ϵ is the aspect ratio of the flow (the classical theory is of zeroth order). Based on a series of comparisons with the exact Poisson equation, we argue that the revised theory is yet better than an order expansion, because it keeps track of the non-locality of the original equation. In contrast with an expansion, indeed, one needs to know the entire shape of the channel in which the revised lubrication equation is to be solved. This feature is what allows the revised theory to account for the side walls of a channel (section 3.2), or the matching of a local asymptotic regime with the bulk of the flow (section 4).

In principle, the method we propose might be extended to three-dimensional flows, provided they remain laminar and shallow, and that their free surface is fixed. Such would be the case, for instance, in a Hele-Shaw cell, where the viscous transfer of momentum in the cell's plane is usually neglected. By accounting for this transfer, the present theory could provide a better representation of the Saffman-Taylor instability, and especially of its regularization, which is classically attributed to surface tension only (Saffman 1986).

Outside the laboratory, the present theory could apply to the viscous flow of ice shelves. Ice streams, in particular, flow over a layer of glacial till which reduces their basal friction; as a consequence, although they are often shallow, a large proportion of the shear stress that opposes their sliding comes from their margins, and propagates crosswise toward their center (Suckale *et al.* 2014; Schoof & Mantelli 2021). The classical lubrication approximation is useless in this case, but the revised theory could help us understand the formation of these fast-flowing channels, without resorting to the complexity of a three-dimensional model. This theory, however, would apply only to an ice stream that does not vary along its course—a rather limiting requirement. Such an abstract configuration, nonetheless, could illustrate the basic mechanism by which a nascent ice stream appears, and thus prove illuminating.

A simplified, low-dimensional configuration often becomes the archetype of a linear instability. One such instability, for instance, can be found below the flowing ice of continental shelves, where meltwater carves channels in the underlying sediment layer (Kasmalkar *et al.* 2019). The stability analysis that explains the initiation of this pattern might well be simplified by the use of the revised

lubrication approximation, as long as the bedforms are aligned with the flow. In fact, these meltwater canals might result from the same instability as the streamwise streaks of Abramian *et al.* (2019a), the stability of which is controlled by the transfer of momentum across the flow—just what the revised theory is designed for.

Finally, it is in temperate latitudes that we find the most ordinary example of natural streams where the present theory applies: alluvial rivers. Indeed, to maintain its bed just above the threshold for sediment transport, a river needs to hand out some of its momentum to the banks (Parker 1978; Abramian *et al.* 2020). To derive the equilibrium shape of an alluvial river, one thus needs to understand how the flow distributes stress over a boundary whose location is unknown beforehand. The present theory might simplify this free-surface problem, and thus help us understand how rivers build their own channels. That would be well worth the effort of revising the lubrication theory.

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Author ORCID. O. Devauchelle, <https://orcid.org/0000-0002-7295-4896>; P. Popović, <https://orcid.org/0000-0001-7937-9153>; E. Lajeunesse, <https://orcid.org/0000-0002-0950-6054>.

Appendix A. Alternative model

Using the notations and hypotheses of the present paper, equation (5.20) in Marche (2007, p. 59) reduces to

$$\frac{gSD}{3\nu} - \frac{U}{D} + \frac{1}{3} \frac{d}{dy} \left(D \frac{dU}{dy} \right) = 0. \quad (\text{A } 1)$$

For comparison with the present model, we rewrite equation (2.16) in terms of the depth-averaged velocity U , using equation (2.14):

$$\frac{gSD}{3\nu} - \frac{U}{D} \left(1 + \left(\frac{dD}{dy} \right)^2 \right) + \frac{1}{3} \frac{d^2}{dy^2} (DU) = 0. \quad (\text{A } 2)$$

Equations (A 1) and (A 2) are obviously equivalent when the water depth is uniform (D constant, section 3.2). Over a corrugated bed, however, they yield distinct results. To see this, we consider a flat bed perturbed with a sinusoidal corrugation of wavenumber k (section 3.3). Up to fourth order in k , we find that the amplitude of the perturbation reads

$$\frac{U_1^*}{U_0} = -2 + \frac{2k^2}{3} - \frac{2k^4}{9} + \mathcal{O}(k^6) \quad (\text{A } 3)$$

for equation (A 1), whereas it reads

$$\frac{U_1^*}{U_0} = -2 + k^2 - \frac{k^4}{3} + \mathcal{O}(k^6) \quad (\text{A } 4)$$

for equation (A 2).

We now wish to compare the above expansions to the linear, two-dimensional theory they should approach. Following Abramian *et al.* (2019a), we linearize the two-dimensional flow equations above a sinusoidal bed perturbation of small amplitude D_1^* :

$$u = \frac{3U_0}{2D_0^2} (z^2 - D_0^2) + \frac{3U_0 D_1^*}{D_0 \cosh k} \cos\left(\frac{ky}{D_0}\right) \cosh\left(\frac{kz}{D_0}\right) + \mathcal{O}(D_1^{*2}). \quad (\text{A } 5)$$

Integrating vertically the above expression, and expanding the result up to fourth order in k , we find

$$\frac{U_1^*}{U_0} = -2 + k^2 - \frac{2k^4}{5} + \mathcal{O}(k^6). \quad (\text{A } 6)$$

The theory we propose here (equation (A 2)) accords with the above expression up to second order in k —an improvement with respect to the classical lubrication theory, which truncates the expansion after the zeroth-order term. The model of Marche (2007) (equation (A 1)) does not yield the right second-order term, although it correctly approximates its trend. Both one-dimensional models are wrong at the fourth order.

Appendix B. Series expansion of the velocity field

In section 2, we have established the revised lubrication equation (2.16) based on an ansatz, that we hoped would be a fair representation of a shallow flow. This method is straightforward, but offers no simple control on the quality of the approximation (such control might be possible, but would involve elaborate mathematics, Ruyer-Quil & Manneville 2000). Here we tread a more familiar path; it does not reach as far as the ansatz method, but it leads to a conservative estimate for the validity of the revised theory.

The classical way to simplify an equation based on the smallness of some parameter is by order expansion. Let \mathcal{L} be the characteristic length over which the bed changes, and \mathcal{H} the typical depth of the flow; when the flow is shallow, its aspect ratio $\epsilon = \mathcal{H}/\mathcal{L}$ is the small parameter upon which we can base our series expansion. We first expand the exact Poisson equation (2.1) into a power series around $\epsilon = 0$, and then do the same with the revised lubrication equation (2.16). This will tell us how the latter converges towards the former as ϵ goes to zero.

B.1. Poisson equation

Rescaling horizontal and vertical coordinates with \mathcal{L} and \mathcal{H} respectively, and the flow velocity with $gS\mathcal{H}^2/\nu$, the Poisson equation (2.1) becomes

$$\epsilon^2 \frac{\partial^2 \tilde{u}}{\partial \tilde{y}^2} + \frac{\partial^2 \tilde{u}}{\partial \tilde{z}^2} = -1 \quad (\text{B } 1)$$

where a tilde indicates a dimensionless quantity ($y = \tilde{y}\mathcal{L}$, $z = \tilde{z}\mathcal{H}$ and $u = \tilde{u}gS\mathcal{H}^2/\nu$). The characteristic aspect ratio of the flow, ϵ , appears naturally in the above equation, thus prompting us to expand the velocity field in terms of ϵ^2 . Accordingly, we define the first two terms of such an expansion as,

$$\tilde{u} = \tilde{u}_0 + \epsilon^2 \tilde{u}_1 + \mathcal{O}(\epsilon^4), \quad (\text{B } 2)$$

and inject this expression into equation (B 1) and the associated boundary conditions, equations (2.2) and (2.3). At zeroth order, we find:

$$\tilde{u}_0 = \frac{1}{2} \left(\tilde{D}^2(\tilde{y}) - \tilde{z}^2 \right), \quad (\text{B 3})$$

where $\tilde{D} = D/\mathcal{H}$ is the dimensionless flow depth. At this order, the diffusion of momentum across the stream does not play any role; therefore, this expression is just the Poiseuille profile of the classical lubrication approximation, that is, the dimensionless form of equation (2.13). More surprisingly, the correction to the leading order, \tilde{u}_1 , also takes the form of a parabolic vertical profile, although with a different prefactor:

$$\tilde{u}_1 = \frac{1}{4} \left(\tilde{D}^2 \right)' \left(\tilde{D}^2(\tilde{y}) - \tilde{z}^2 \right), \quad (\text{B 4})$$

where a prime denotes derivation with respect to \tilde{y} . The combination of equations (B 2), (B 3) and (B 4) yields the expression of the velocity field up to second order in ϵ :

$$\tilde{u} = \frac{1}{2} \left(\tilde{D}^2(\tilde{y}) - \tilde{z}^2 \right) \left(1 + \frac{\epsilon^2}{2} \left(\tilde{D}^2 \right)'' \right) + \mathcal{O}(\epsilon^4). \quad (\text{B 5})$$

The above expression is an improvement upon the lubrication approximation, which might prove useful in itself. Here we simply note that, up to order ϵ^2 the vertical velocity profile remains parabolic (this is not true at the next order, ϵ^4). The vertical profile can thus be factorized in equation (B 5)—an indication that the Poiseuille profile is a reasonable ansatz. In that sense, the above expansion supports the revised lubrication approximation. In the next section, we check that the latter indeed matches equation (B 5) up to order ϵ^2 .

B.2. Revised lubrication equation

The revised lubrication equation (2.16) is supposed to approximate the flow when the aspect ratio ϵ is small enough. Its expansion around $\epsilon = 0$ should, therefore, be consistent with that of the exact Poisson equation (2.1). To confirm this, we first need to make equation (2.16) dimensionless. Defining the vertical component of the dimensionless shear stress as $\tilde{\tau}_z = \tau_z/(\rho g \mathcal{H} S)$, we find

$$\frac{\epsilon^2}{3} \left(\tilde{D}^2 \tilde{\tau}_z \right)'' - \left(1 + \epsilon^2 \left(\tilde{D}' \right)^2 \right) \tilde{\tau}_z + \tilde{D} = 0, \quad (\text{B 6})$$

where, like in section B.1, the aspect ratio of the flow naturally appears as ϵ^2 . Expanding the shear stress $\tilde{\tau}_z$ around $\epsilon = 0$ again yields, at leading order, the term that corresponds to the classical lubrication equation, $\tilde{\tau}_{z,0} = \tilde{D}$. At the next order (ϵ^2), we find

$$\tilde{\tau}_{z,1} = \frac{\tilde{D}}{2} \left(\tilde{D}^2 \right)''. \quad (\text{B 7})$$

Combining the zeroth- and first-order terms into a single expression, we get

$$\tilde{\tau}_z = \tilde{D} \left(1 + \frac{\epsilon^2}{2} \left(\tilde{D}^2 \right)'' \right) + \mathcal{O}(\epsilon^4). \quad (\text{B 8})$$

To compare this expression to the expansion of the Poisson equation (section B.1), we only need to differentiate equation (B 8) with respect to the vertical coordinate

\tilde{z} ; it then becomes equation (B 5). The revised lubrication approximation is thus equivalent to the original Poisson equation up to order ϵ^2 , whereas the validity of the classical theory ends at zeroth order.

The formal expansion of this appendix demonstrates that the revised approximation we propose improves the classical lubrication equation, at least up to order ϵ^2 . However, why would anyone want to solve the ordinary differential equation (2.16), when equation (B 8) provides an explicit solution that is also valid up to order ϵ^2 ? A straightforward answer is to note that, in an elliptic channel, equation (B 8) does not yield the exact solution, whereas the revised lubrication equation does (section 3.1). The revised theory, therefore, must be a bit better than the classical expansion.

This answer is conveniently simple, but we believe it misses a deeper point. The main difference between the two approaches is that the revised lubrication equation (2.16) is non-local, whereas expansion (B 8) depends only on the flow depth and its second derivative. Whether non-locality, which requires us to solve a differential equation with boundary conditions, is an advantage or an inconvenience depends on the context. It is computationally more expensive than the explicit expansion (B 8), but it is more flexible. A good example of this versatility is the rectangular groove of section 3.2: since the channel's bed is flat, the second derivative of depth vanishes everywhere, and expansion (B 8) yields the same result as the classical lubrication approximation. The revised theory, conversely, accounts for the presence of the two walls with decent accuracy. Likewise, the non-local theory represents qualitatively the transition of the flow along a corner (section 4), a feat out of any local theory's reach.

Appendix C. Friction coefficient

The present theory, like the lubrication approximation, is based on an assumption about the vertical velocity profile. The details of this profile, however, do not really matter; only the relation it implies between the average flow velocity U and the shear stress does (sections 2.3 and 2.4). We can express this relation in terms of a friction coefficient, C_f , which we define as follows:

$$C_f = \frac{D\tau_z}{\rho\nu U}. \quad (\text{C } 1)$$

In the present paper, this coefficient is assumed to be $C_f = 3$, the value it takes when the velocity profile is exactly that of a Nusselt film (2.13). It is the expression of C_f , as opposed to the entire shape of the profile, that allows us to write equation (2.16) in closed form. As a consequence, the revised lubrication theory is exactly as reliable as the friction law it is based on—but this law needs not be $C_f = 3$.

C.1. Approximation

In the rectangular groove of section 3.2, the revised lubrication approximation fails near the side walls. This is because the velocity profile, perturbed by the presence of the wall, departs from the ansatz (2.13).

Figure 9(a) shows how the vertical velocity profile changes near the side wall of a rectangular groove of aspect ratio 5. At most, the velocity profile differs from the Nusselt profile by about 12% (when $y + W/2 = 0.01W$), but this small difference

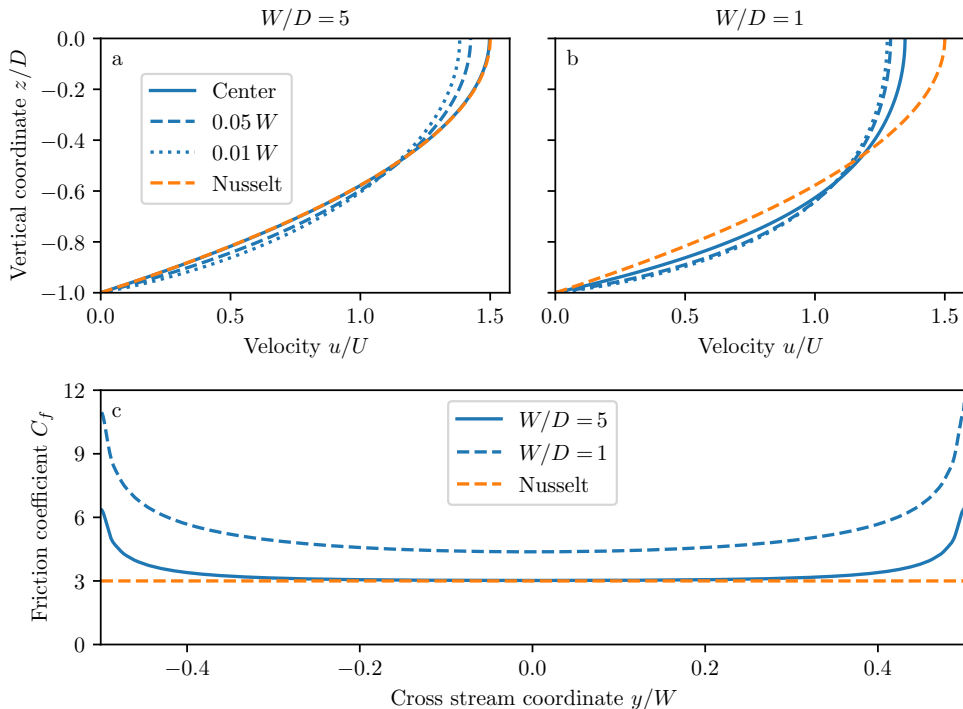


Figure 9: Velocity profile and friction coefficient in the rectangular grooves of Figure 5. (a) and (b): Vertical velocity profile at different locations across the channel, according to equation (3.7) (blue lines), and according to the ansatz (2.13) (dashed orange line). Legend indicates distance to the left wall. (c): Friction coefficient C_f across the channel according to equation (3.7) (blue lines). Dashed orange line: ansatz friction coefficient ($C_f = 3$).

is enough to increase the friction coefficient by a factor of about 2 (Figure 9(c)). The departure from the Nusselt profile is even more pronounced in a narrower channel (Figure 9(a)).

Near the wall, however, the horizontal transfer of momentum is large, whereas neither the friction on the bottom nor gravity contribute much to the momentum balance. In other words, the first term in equation (2.16) needs to vanish—hence the linearity of the cross-stream profile in Figure 3(c). This is fortunate, because it means that the shape of the profile does not depend on the value of the friction coefficient. This mitigates the error introduced by the ansatz, thus explaining why the lubrication approximation still performs well near the walls (Figure 3(c)).

There is no reason to believe that this mitigation is inevitable, and one might devise an exotic wall condition that would strongly affect the vertical profile of the flow, without absorbing much momentum. This would more methodically break the revised lubrication theory. Instead of this mischievous endeavor, we now propose a simple extension of the theory which illustrates the role of the friction coefficient.

C.2. Slip condition

So far, we have assumed that the fluid cannot slip along the bed, but this needs not be always the case. Although this is seldom significant at the macroscopic scale, a liquid can slip along a solid (Lauga & Stone 2003). This phenomenon, often referred to as “Navier slip”, is of prime importance in microfluidic devices, or near a moving contact line (Snoeijer *et al.* 2006). Similarly, the velocity of a fluid can remain finite along the surface of a porous solid (Beavers & Joseph 1967) and, yet at a larger scale, glaciers and ice sheets are known to slide over their rock bed (Schoof & Hewitt 2013).

The simplest way to represent this sliding is to replace the no-slip boundary condition (2.3) with a more general (but still linear) Robin condition:

$$u = \lambda_s \frac{\partial u}{\partial z} \quad \text{for } z = -D(y), \quad (\text{C } 2)$$

where λ_s is a slip length which depends on the fluid, and on the solid over which it flows. According to this new boundary condition, the slip velocity is simply proportional to the shear stress the flow applies on the surface of the solid.

Over a flat bed, the velocity profile that satisfies the momentum balance (2.1) and the above boundary condition is again a parabola, but one that allows for some slip along the bed:

$$u = \frac{gSD^2}{2\nu} \left(1 - \left(\frac{z}{D} \right)^2 + 2 \left(\frac{\lambda_s}{D} \right)^2 \right). \quad (\text{C } 3)$$

Replacing the Nusselt profile (2.13) with the above parabola, we can finally calculate the friction coefficient:

$$C_f = \frac{3}{1 + 3(\lambda_s/D)^2}. \quad (\text{C } 4)$$

This expression is all we need to account for slip in the framework of the revised lubrication theory. In practice, we would just inject it into the integrated momentum balance (2.10). We would then find an equation similar to equation (2.16), but for the first term, which would depend on the slip length λ_s .

As the slip length λ_s increases, sliding becomes easier—until the bottom friction becomes negligible. In the integrated momentum balance (2.10), this translates into the first term (cross-stream diffusion) dominating the second (exchange with the bed). In short, the revised lubrication theory transitions smoothly to a “sliding membrane” model in which momentum is exchanged only laterally (Schoof & Hewitt 2013).

REFERENCES

- ABRAMIAN, ANAÏS, DEVAUCHELLE, OLIVIER & LAJEUNESSE, ERIC 2019a Streamwise streaks induced by bedload diffusion. *Journal of Fluid Mechanics* **863**, 601–619.
- ABRAMIAN, A., DEVAUCHELLE, O. & LAJEUNESSE, E. 2020 Laboratory rivers adjust their shape to sediment transport. *Phys. Rev. E* **102**, 053101.
- ABRAMIAN, A., DEVAUCHELLE, O., SEIZILLES, G & LAJEUNESSE, E 2019b Boltzmann distribution of sediment transport. *Physical review letters* **123** (1), 014501.
- AUDUSSE, EMMANUEL, BRISTEAU, MARIE-ODILE, PERTHAME, BENOÎT & SAINTE-MARIE, JACQUES 2011 A multilayer saint-venant system with mass exchanges for shallow water flows. derivation and numerical validation. *ESAIM: Mathematical Modelling and Numerical Analysis* **45** (1), 169–200.

- BEAVERS, GORDON S & JOSEPH, DANIEL D 1967 Boundary conditions at a naturally permeable wall. *Journal of fluid mechanics* **30** (1), 197–207.
- BENJAMIN, T. BROOKE 1957 Wave formation in laminar flow down an inclined plane. *Journal of Fluid Mechanics* **2** (6), 554–573.
- BOUSSINESQ, JOSEPH 1868 Mémoire sur l'influence des frottements dans les mouvements réguliers des fluides. *Journal de mathématiques pures et appliquées* **13** (2), 377–424.
- BROWN, G.O. 2002 *The History of the Darcy-Weisbach Equation for Pipe Flow Resistance*, pp. 34–43. ASCE.
- CHANSON, HUBERT 2012 *Tidal bores, aegir, eagre, mascaret, pororoca: Theory and observations*. World Scientific.
- CHARRU, F. & HINCH, E.J. 2006 Ripple formation on a particle bed sheared by a viscous liquid. part 1. steady flow. *Journal of Fluid Mechanics* **550**, 111–121.
- CHAUVET, HUGO, DEVAUCHELLE, OLIVIER, MÉTIVIER, FRANÇOIS, LAJEUNESSE, ERIC & LIMARE, ANGELA 2014 Recirculation cells in a wide channel. *Physics of Fluids* **26** (1), 016604.
- CHÉZY, A 1775 Mémoire sur la vitesse de l'eau conduite dans une rigole donnée. *Dossier* **847**, 363–368.
- CRASTER, RV & MATAR, OK 2009 Dynamics and stability of thin liquid films. *Reviews of modern physics* **81** (3), 1131.
- DE VITA, FRANCESCO, LAGRÉE, PIERRE-YVES, CHIBBARO, SERGIO & POPINET, STÉPHANE 2020 Beyond shallow water: appraisal of a numerical approach to hydraulic jumps based upon the boundary layer theory. *European Journal of Mechanics-B/Fluids* **79**, 233–246.
- DELESTRE, OLIVIER, CORDIER, STÉPHANE, DARBOUX, FRÉDÉRIC, DU, MINGXUAN, JAMES, FRANCOIS, LAGUERRE, CHRISTIAN, LUCAS, CARINE & PLANCHON, OLIVIER 2014 Fullswof: A software for overland flow simulation. In *Advances in hydroinformatics*, pp. 221–231. Springer.
- DEVAUCHELLE, O., PETROFF, A. P., LOBKOVSKY, A. E. & ROTHMAN, D. H. 2011 Longitudinal profile of channels cut by springs. *Journal of Fluid Mechanics* **667**, 38–47.
- GALLAGHER, BRENT S & MUNK, WALTER H 1971 Tides in shallow water: Spectroscopy. *Tellus* **23** (4-5), 346–363.
- GLOVER, R. E. & FLOREY, Q. L. 1951 Stable channel profiles. *Tech. Rep.*. U.S. Bur. Reclamation, Denver, CO, USA.
- GOODWIN, RALPH & HOMSY, GM 1991 Viscous flow down a slope in the vicinity of a contact line. *Physics of Fluids A: Fluid Dynamics* **3** (4), 515–528.
- HECHT, F. 2012 New development in freefem++. *J. Numer. Math.* **20** (3-4), 251–265.
- HENDERSON, FRANCIS M 1961 Stability of alluvial channels. *Journal of the Hydraulics Division* **87** (6), 109–138.
- HUPPERT, HERBERT E 1982 Flow and instability of a viscous current down a slope. *Nature* **300** (5891), 427–429.
- HUPPERT, HERBERT E, SHEPHERD, JOHN B, SIGURDSSON, R HARALDUR & SPARKS, STEPHEN J 1982 On lava dome growth, with application to the 1979 lava extrusion of the soufriere of st. vincent. *Journal of Volcanology and Geothermal Research* **14** (3-4), 199–222.
- KAPITZA, PL 1948 Wave flow of thin layer of viscous fluid (in russian). *Zh. Eksp. Teor. Fiz.* **18**, 3–28.
- KASMALKAR, INDRANEEL, MANTELLI, ELISA & SUKALE, JENNY 2019 Spatial heterogeneity in subglacial drainage driven by till erosion. *Proceedings of the Royal Society A* **475** (2228), 20190259.
- KORTEWEG, DIEDERIK JOHANNES & DE VRIES, GUSTAV 1895 Xli. on the change of form of long waves advancing in a rectangular canal, and on a new type of long stationary waves. *The London, Edinburgh, and Dublin Philosophical Magazine and Journal of Science* **39** (240), 422–443.
- LACEY, GERALD 1930 Stable channels in alluvium. In *Minutes of the Proceedings of the Institution of Civil Engineers*, , vol. 229, pp. 259–292. Thomas Telford-ICE Virtual Library.
- LAUGA, E. & STONE, H.A. 2003 Effective slip in pressure-driven stokes flow. *Journal of Fluid Mechanics* **489**, 55–77.

- LEVICH, B & LANDAU, L 1942 Dragging of a liquid by a moving plate. *Acta Physicochim. URSS* **17**, 42.
- LISTER, JOHN R 1992 Viscous flows down an inclined plane from point and line sources. *Journal of Fluid Mechanics* **242**, 631–653.
- MANNING, R. 1891 On the flow of water in open channels and pipes. *Trans. Inst. Civil Eng. Ireland* **20**, 161–207.
- MARCHE, FABIEN 2007 Derivation of a new two-dimensional viscous shallow water model with varying topography, bottom friction and capillary effects. *European Journal of Mechanics-B/Fluids* **26** (1), 49–63.
- MÉTIVIER, F., LAJEUNESSE, E. & DEVAUCHELLE, O. 2017 Laboratory rivers: Lacey’s law, threshold theory, and channel stability. *Earth Surface Dynamics* **5** (1), 187–198.
- MICHAUT, CHLOÉ 2011 Dynamics of magmatic intrusions in the upper crust: Theory and applications to laccoliths on earth and the moon. *Journal of Geophysical Research: Solid Earth* **116** (B5).
- MOFFATT, H KEITH 1964 Viscous and resistive eddies near a sharp corner. *Journal of Fluid Mechanics* **18** (1), 1–18.
- NEZU, IEHISA, NAKAGAWA, HIROJI & JIRKA, GERHARD H 1994 Turbulence in open-channel flows. *Journal of Hydraulic Engineering* **120** (10), 1235–1237.
- PARKER, GARY 1978 Self-formed straight rivers with equilibrium banks and mobile bed. part 2. the gravel river. *Journal of Fluid mechanics* **89** (1), 127–146.
- PHILLIPS, CB & JEROLMACK, DJ 2019 Bankfull transport capacity and the threshold of motion in coarse-grained rivers. *Water Resources Research* **55** (12), 11316–11330.
- POLUBARINOVA-KOCHINA, P.IA. 1962 *Theory of ground water movement*. Princeton University Press.
- POPINET, STÉPHANE 2011 Quadtree-adaptive tsunami modelling. *Ocean Dynamics* **61** (9), 1261–1285.
- RUYER-QUIL, CHRISTIAN & MANNEVILLE, PAUL 2000 Improved modeling of flows down inclined planes. *The European Physical Journal B-Condensed Matter and Complex Systems* **15** (2), 357–369.
- SAFFMAN, PG 1986 Viscous fingering in hele-shaw cells. *Journal of Fluid Mechanics* **173**, 73–94.
- SAINT-VENANT, AJC DE 1871 Théorie du mouvement non permanent des eaux, avec application aux crues des rivières et à l’introduction de marées dans leurs lits. *Comptes rendus des seances de l’Academie des Sciences* **36**, 174–154.
- SCHOOF, CHRISTIAN & HEWITT, IAN 2013 Ice-sheet dynamics. *Annual Review of Fluid Mechanics* **45**, 217–239.
- SCHOOF, CHRISTIAN & MANTELLI, ELISA 2021 The role of sliding in ice stream formation. *Proceedings of the Royal Society of London Series A* **477** (2248), 20200870.
- SEIZILLES, G, DEVAUCHELLE, O, LAJEUNESSE, E & MÉTIVIER, F 2013 Width of laminar laboratory rivers. *Physical Review E* **87** (5), 052204.
- SNOEIJER, JH, ZIEGLER, J, ANDREOTTI, B, FERMIGIER, M & EGGERS, J 2008 Thick films of viscous fluid coating a plate withdrawn from a liquid reservoir. *Physical review letters* **100** (24), 244502.
- SNOEIJER, JACCO H, DELON, GILES, FERMIGIER, MARC & ANDREOTTI, BRUNO 2006 Avoided critical behavior in dynamically forced wetting. *Physical review letters* **96** (17), 174504–174504.
- STASIUK, MARK V & JAUPART, CLAUDE 1997 Lava flow shapes and dimensions as reflections of magma system conditions. *Journal of Volcanology and Geothermal Research* **78** (1-2), 31–50.
- STOKER, JAMES JOHNSTON 2011 *Water waves: The mathematical theory with applications*, , vol. 36. John Wiley & Sons.
- SUCKALE, JENNY, PLATT, JOHN D, PEROL, THIBAUT & RICE, JAMES R 2014 Deformation-induced melting in the margins of the west antarctic ice streams. *Journal of Geophysical Research: Earth Surface* **119** (5), 1004–1025.
- WHITE, FRANK M. 1991 *Viscous fluid flow*, 2nd edn. *Mechanical Engineering* . McGraw-Hill, Inc.
- WITTEN JR, TA & SANDER, LEONARD M 1981 Diffusion-limited aggregation, a kinetic critical phenomenon. *Physical review letters* **47** (19), 1400.

YIH, CHIA-SHUN 1963 Stability of liquid flow down an inclined plane. *The physics of Fluids*
6 (3), 321–334.