

Bore collapse and wave run-up on a sandy beach

Erwin W.J. Bergsma, Chris E. Blenkinsopp, Kévin Martins, Rafael Almar, Luis P. Melo de Almeida

▶ To cite this version:

Erwin W.J. Bergsma, Chris E. Blenkinsopp, Kévin Martins, Rafael Almar, Luis P. Melo de Almeida. Bore collapse and wave run-up on a sandy beach. Continental Shelf Research, 2019, 174, pp.132 - 139. 10.1016/j.csr.2019.01.009 . hal-03486880

HAL Id: hal-03486880 https://hal.science/hal-03486880

Submitted on 20 Dec 2021

HAL is a multi-disciplinary open access archive for the deposit and dissemination of scientific research documents, whether they are published or not. The documents may come from teaching and research institutions in France or abroad, or from public or private research centers. L'archive ouverte pluridisciplinaire **HAL**, est destinée au dépôt et à la diffusion de documents scientifiques de niveau recherche, publiés ou non, émanant des établissements d'enseignement et de recherche français ou étrangers, des laboratoires publics ou privés.



Distributed under a Creative Commons Attribution - NonCommercial 4.0 International License

Bore collapse and wave run-up on a sandy beach

Erwin W. J. Bergsma^{a,*}, Chris E. Blenkinsopp^b, Kévin Martins^{c,d}, Rafael Almar^a, Luis P. Melo de Almeida^e

^aCNES/IRD-LEGOS (CNES/CNRS/IRD/UPS) - UMR 5566, 14 Avenue Edouard Belin, 31400 Toulouse, France

^b University of Bath, Research Unit for Water, Environment and Infrastructure Resilience (WEIR), Bath, BA2 7YY, United Kingdom

^cLIENSs (CNRS-Université de La Rochelle) - UMR7266, 2 rue Olympe de Gouges, 17000 La Rochelle, France

^dEPOC (CNRS-Université de Bordeaux) - UMR5805, Allée Geoffroy Saint-Hilaire,, 33615 Pessac, France

^eUniversidade Federal do Rio Grande, Avenida Itália - Carreiros, 96203-900, Rio Grande, RS, Brasil

Abstract

Wave run-up on beaches and coastal structures is initiated and driven by collapsing incident bores, this process is often considered to define the seaward limit of the swash zone. It is hence a key feature in nearshore wave processes as extreme run-up can lead to structure overtopping and coastal inundation during storm conditions. In addition, the turbulent nature of incident bores and their collapse suspends and advects sediment, resulting in a highly morphologically dynamic swash zone. The cross shore bore collapse location varies from wave to wave and the process is very limited in both spatial and temporal extent, making direct measurement problematic. This paper presents high spatial-temporal resolution LiDAR field measurements of the evolving free-surface in the surf and swash zone which enable the bore collapse detection for 166 waves. These measurements are used to investigate the link between broken wave properties at bore collapse and wave run-up. Incident bores are identified at the seaward boundary of the LiDAR profiles and tracked through the inner surf and swash zones to the run-up limit. It is found that the vertical run-up height exceeds that which would be expected for a perfect conversion of potential to kinetic

Preprint submitted to Continental Shelf Research

December 19, 2018

^{*}Corresponding author

Email address: Erwin.Bergsma@legos.obs-mip.fr (Erwin W. J. Bergsma)

energy during bore collapse for 24% of the bores measured. By returning to an existing ballistic-type model to describe the run-up of individual waves, we show that wave run-up can be divided into three components: the bore collapse, terminal bore celerity and their non-linear interaction. For the present dataset, the contribution of the bore collapse and terminal bore celerity is 26% and 27% respectively, while non-linear interactions between the two dominates and account for 47% of the measured run-up. By including the terminal bore celerity, the ability to predict run-up is increased by 30% with the determination coefficient r increasing from 0.573 to 0.785. Likewise, the RMS-error for the wave run-up shows an approximately 10% reduction from 0.325 to 0.295 m. *Keywords:* Bore collapse, Swash zone, Run-up, Swash-swash interaction.

1 1. Introduction

Incident waves dissipate their energy as they break and propagate in the surf zone as bores. As bores reach the shoreline, the wave form compresses as the wave decelerates and eventually collapses leading to wave run-up on beaches or coastal protection structures. The swash zone is recognized as a highly turbulent region with unsteady, non-uniform flows [14]. The turbulent flows suspend sediment into the water column leading to sediment transport [13] and relatively rapid morphological change on sandy [30, 25] and gravel beaches 8 [3]. Consequently, new insight into processes at the boundary between inner 9 and swash zones are valuable to enhance understanding of beach hydro and 10 morphodynamics. In this work we focus on the shallowest part of the inner-surf 11 zone, the bore collapse and associated vertical run-up. Furthermore, interaction 12 between consecutive swash events is discussed. 13

Incident bores in relation to run-up have been studied in scaled laboratories with single bores running up a slope, e.g. [9, 17]. Field measurements of the swash zone have tended to focus on maximum run-up and the statistical distribution of swash excursions using cameras or run-up wires e.g. [18, 20, 35]. More recent field studies [26, 10] have investigated more detailed swash hydro and morphodynamics using a variety of techniques in response to the recommendations of Puleo and Butt [31] and Puleo and Torres-Freyermuth [33] who
suggested that measurements of swash on a wave-by-wave basis was key to enhancing understanding.

Whitham [38] describes bore collapse and run-up on a wave-by-wave basis 23 by applying an analytical mathematical solution of a propagating bore in non-24 uniform water depth. This analytical solution shows that as bores propagate to 25 shore, they reach a maximum finite velocity (U_0) proportional to the local water 26 depth. This velocity is often applied as the starting point for run-up models of 27 individual bores such as the ballistic model of Shen and Meyer [34]. Yeh et al. 28 [40] calculated U_0 through a classical dam-break problem assuming a perfect 29 conversion of potential to kinetic energy. In the same work it is shown through 30 laboratory experiments that the theoretical value overestimates the measured 31 finite velocity for a single fully developed incident bore, while for undular bores 32 there seems to be a better match between theory and measurements. Baldock 33 and Holmes [8] recognized that the conversion efficiency varies with the type of 34 bore collapse (undular bores, uniform bores or waves breaking on the beach), 35 assuming that the theory of Whitham [38] is valid for the different types of bore 36 collapse or wave breaking. In order to take imperfect energy conversion into 37 account, Baldock and Holmes [8] introduced an energy conversion coefficient C. 38 The approaches above do not consider direct bore-bore interactions although 39 sediment transport in the swash zone is thought to be greatly affected by the 40 interaction between consecutive swash events [21, 26]. Capturing the complex-41 ity of such interactions is a major challenge to existing hydrodynamic modelling 42 approaches and introduces significant uncertainty into sediment transport pre-43 dictions [10]. Few existing studies have directly addressed this phenomenon, 44 though Hegge and Eliot [18] classified swash-swash interaction into 5 categories; 45 free, over-taking, over-riding, suppressed or composite. Over-taking represents 46 a bore that rides on top of the previous bore. Over-riding and suppressed 47 swash-swash interaction modes are linked through the strength of the backwash. 48 The composite mode consists of more than one of the other modes. Baldock 49

and Holmes [8] used their swash model to incorporate swash-swash interaction, 50 however in their study the effect of interactions was to translate the location of 51 bore collapse and the actual interaction between consecutive swashes was not 52 explicitly considered. More specific bore interaction-focussed laboratory exper-53 iments investigated the interaction of two successive bores [29, 14] in which the 54 former showed that the degree of swash-swash interaction relates to the soli-55 tary wave slope parameter [15]. Bore-bore interactions influence the cross shore 56 location of the bore collapse, bore height and advection of sediment into the 57 swash zone, which in turn affect swash-morphodynamics [5]. In line with this 58 observation, Alsina and Cáceres [4] showed that for saturated surf-zones the 59 amount of suspended sediment at the inner-surf, swash zone boundary is inde-60 pendent of offshore wave height, but caused by the combined action of incident 61 swell and swell related long-period water oscillations, for example modulation 62 of wave-wave interaction due to the wave group frequency [6]. 63

This paper focuses on swash run-up observed in the field using a 2D LiDAR. A novel technique to capture and extract the bore collapse and incoming bore celerity from the data is presented and applied to investigate the nature of bore collapse and its importance to wave run-up.

68 2. Method

69 2.1. Study site and data collection

In-situ 2D LiDAR data was collected at Nha Trang beach, on the East coast 70 of Vietnam (Figure 1) during a 9-day field experiment from 26 November to 4 71 December 2015. The sandy beach of Nha Trang is situated in a semi-enclosed 72 bay, protected by a group of islands at the Southern part of the bay. The 5 km 73 long stretch of beach is therefore mostly exposed to North-Easterly swell. The 74 East-Vietnam coast experiences a wave climate that is primarily governed by two 75 monsoon seasons; the North East and South West monsoon. The former (latter) 76 is characterized by strong (mild) winds and energetic (moderate) waves. In addi-77 tion to monsoons, the region experiences occasional tropical storms (typhoons) 78



Figure 1: Left-hand map highlights Vietnam (darker area) and the red square indicates the zoomed area for the right-hand map. The right-hand map shows the location of Nha Trang.

leading to rapid erosion at Nha Trang bay [37, 1]. During the time-frame of 79 the experiment, the average significant wave height (H_s) was 1.07 m, with a 80 corresponding peak period of 11 seconds. The micro-tidal regime at Nha Trang 81 (maximum tidal range = 1.5 m) consists of a mix of diurnal and semi-diurnal 82 tides [27]. The upper beach slope was 0.1 while the inter-tidal terrace has a 83 slope of 0.01. The sediment sizes varies within Nha Trang bay from $D_{50} = 900$ 84 μm (coarse) in the North to 400 μm (medium-coarse) in the South [2]. During 85 the experiments, almost no wind was present, but this was not measured. 86

⁸⁷ During the field campaign a range of instruments were deployed, an off-⁸⁸ shore Acoustic Doppler Current Profiler (ADCP), near-shore ADV (Acoustic ⁸⁹ Doppler Velocimeter), shore mounted video cameras, a swash pole camera, pres-⁹⁰ sure transducers and a 2D LiDAR (for details see Almeida et al. [2]). This paper ⁹¹ will focus on data only from the 2D LiDAR which was deployed on a 4 m tall ⁹² tower above the high tide limit as shown in Figure 2. LiDAR data was collected ⁹³ at 25 Hz and was typically able to obtain beach profile and free surface data



Figure 2: The upper plot shows a photograph of the setup at low-tide at Nha Trang beach during the field experiment. The arrow in the top plot indicates the position of the 2D LiDAR. The bottom plot shows a snapshot of obtained and processed LiDAR data (beach and free surface - black line) at Nha Trang through a schematic representation of the LiDAR position and laser beams.

along a transect extending approximately 30 m seaward of the LiDAR position. 94 The obtained LiDAR data was post-processed using the methodology de-95 scribed by Almeida et al. [3] and Martins et al. [24] and interpolated onto a 1D 96 grid with $\Delta x = 10$ cm. The lower panel in Figure 2 shows a snapshot of the cap-97 tured water surface elevation and beach profile data. Here, a 40 minute subset 98 of the total collected dataset between 22:21 to 23:01 on the 27th of November 99 2015 containing 166 bores is analysed. During this time a significant wave height 100 of 1.2 m and peak period of 12 seconds was measured offshore. 101

102 2.2. Bore collapse and vertical run-up

The seaward boundary of the swash zone is characterized by a rapid steepening of the incoming bore and ultimately, as the water depth in front of the bore approaches zero, the bore collapses, driving swash up-rush [39, 20]. As briefly

discussed above, previous analytical work by Whitham [38] suggests that a fi-106 nite velocity U_0 proportional to the local water depth $\propto \sqrt{h}$ is reached. It is 107 noted that the empirical approach presented here takes no account of several 108 processes known to influence swash flows, including friction swash-groundwater 109 interactions and porosity e.g. [32, 22, 12]. Yeh et al. [40] used the classic bore 110 collapse theory to calculate the finite velocity at collapse, assuming a perfect 111 conversion of potential to kinetic energy during the bore collapse process e.g. 112 [8, 11, 36]. This initial shoreline velocity can then be used in to estimate swash 113 trajectory via a ballistic-type model e.g. [34]. Hence vertical run-up (R) can be 114 approximated as a function of the initial shoreline velocity U_0 : 115

$$R = \frac{U_0^2}{2g} \tag{1}$$

In which g is the acceleration due to gravity. The maximum velocity in the case where no energy is lost during the transformation of potential to kinetic energy is approximated as a function of the bore height at collapse following $U_0 = 2\sqrt{gH_{b,c}}$ [40]. Baldock and Holmes [8] replaced the factor of 2 (perfect conversion) by an empirical bore collapse efficiency coefficient C to approximate the initial velocity which, in theory, is in the range 1 to 2:

$$U_0 = C\sqrt{gH_{b,c}} \tag{2}$$

where $H_{b,c}$ is the bore height at collapse in which according to Shen and Meyer [34], the bore height must be taken in slope-normal direction. Following (2), the vertical run-up from the bore collapse location can now be calculated as:

$$R = \frac{C^2 H_{b,c}}{2} \tag{3}$$

126 2.3. Detection of bore collapse

The high spatial and temporal resolution of the LiDAR data collected enables individual incident bores to be tracked through the near shore. Tracking of individual bores allows evolving bore characteristics such as bore shape, height, period and the bore collapse process to be captured. The bore tracking methodology is similar to that used in the surf zone by Martins et al. [24], though instead of tracking peaks in the surface elevation, peaks in its spatial derivative are tracked instead. The tracking was initiated at the cross-shore position x = -18 m. Then, individual bores are subsequently tracked inshore for every Δx by identifying the maximum surface gradient around the previously detected peak. An example of this process is shown in Figure 3.

Inshore of the breakpoint, the bore front gradient varies considerably in the 137 surf zone with the breaking intensity [23]. As the bores approach the boundary 138 between inner surf and swash, the front steepens, reaching a local maximum 139 gradient just before the bore collapses (red dot in the bottom plot of Figure 3). 140 The LiDAR is able to detect this and here we define bore collapse as the point 141 at which the local maximum bore front steepness occurs, before the bore front 142 suddenly and rapidly flattens as observed in Figure 3. Thus the location of bore 143 collapse is defined as the location of the local maximum bore front gradient in 144 time: 145

Bore Collapse = max
$$\left(\left| \frac{\Delta \eta_{bf}}{\Delta x} \right| \right)_t$$
 (4)

In which η_{bf} is the free surface elevation of the bore front which is defined between a seaward and landward limit taken here as:

Fitting Limits
$$\eta_{bf} = \left|\frac{\Delta\eta}{\Delta x}\right| < \frac{1}{8} \max\left(\left|\frac{\Delta\eta}{\Delta x}\right|\right)_t$$
 (5)

where η is the free surface elevation. In Figure 4, a dot-dashed line is drawn 148 connecting the points defined by (5) to indicate the bore front slope at collapse. 149 A second line is fitted at the moment of bore collapse to the free surface elevation 150 0.5 metres seaward of the seaward-limit of (5) as illustrated in Figure 4. The 151 fitting limits are illustrated by the grey vertical dashed lines in Figure 4. The 152 intersection of the two lines is taken as the bore head point; indicated by the blue 153 circle in Figure 4. The bore height is ultimately determined at the bore head 154 point as the vertical distance from the bore head to the bed. The fitting limits 155



Figure 3: A schematic representation of bore tracking using the LiDAR data. In both plots, the time between every line is 3 time steps ($\Delta t = 25$ Hz). The upper plot shows the gradient of the measured free surface. The dots represent the tracked bore positions. The dots are the local maximum gradient determined for every gridded cross shore point in time. The lower plot shows the measured free surface elevation and the grey dots are the positions as derived from the upper plot. The red-lines represent the bore front gradient and the red-dot indicates the bore collapse (maximum bore front gradient in time).



Figure 4: Schematic representation of the bore collapse detection. The solid red line represents the measured free surface elevation, the dash-dotted (-.) line is the slope of the bore front, the dashed green line represents the slope of the free surface elevation on the seaward side of the bore head. In the upper part of the plot, the black solid line represents the absolute derivative of the surface elevation and the dashed red line indicates the fitting limit. The grey dashed lines give the used fitting boundaries.

¹⁵⁶ used here are calibrated for the current Nha Trang dataset and are therefore ¹⁵⁷ site-specific. Nonetheless, considering bore self-similarity, it is expected that ¹⁵⁸ similar thresholds are likely to be valid for other datasets with differing site and ¹⁵⁹ wave conditions. The same method was used successfully to define surf zone ¹⁶⁰ bores by Martins et al. [23].

The bore height is estimated at every cross-shore location using the surface 161 elevation data. Tracking the incident bores through the surf zone allows for an 162 estimation of the bore celerity. Until bore collapse, changes in the bore shape 163 are minimal, leading to a robust celerity estimate. After collapse, the earlier 164 bore features such as the steep front appear less distinct and as such the celerity 165 estimate is equivalent to the shoreline velocity. At bore collapse, a local bore 166 related Froude number (similar to that presented by Yeh et al. [40], Zhang and 167 Liu [41]) based on the bore celerity $(c_{b,c})$ and bore height $(H_{b,c})$ is defined: 168

$$Fr_{b,c} = \frac{c_{b,c}}{\sqrt{gH_{b,c}}} \tag{6}$$

¹⁶⁹ 2.4. Determination of run-up (R) from LiDAR data

The run-up of every wave is defined as the distance (horizontal and vertical) between the toe of the bore at collapse and the most landward shoreline position. In order to calculate the vertical run-up (R), the shoreline is extracted from the LiDAR data using the 3 cm water depth contours which is tracked throughout the up-rush/backwash cycle.

175 3. Results

176 3.1. Bore collapse signature from LiDAR

To date, the bore collapse process has predominantly been observed in lab-177 oratory experiments and modelled with numerical models (e.g. [40, 28]). The 178 LiDAR data provides the opportunity to observe the nature of the bore collapse 179 process on a wave-by-wave basis in the field. An example observation of the 180 most commonly occurring bore collapse sequence is shown in Figure 5. Figure 181 5a shows a bore approaching the shoreline which reaches a maximum steepness 182 at the point of collapse (Figure 5b). The bore collapse process initiates the 183 swash up-rush (Figure 5c-d), following flow reversal (Figure 5e) the backwash 184 flow is then observed to interact with the proceeding bore (Figure 5f). 185

¹⁸⁶ By tracking the bore properties, the spatial development of the bore front ¹⁸⁷ gradient can be investigated. Figure 6 shows the spatial variations of the bore ¹⁸⁸ front gradient in the vicinity of bore collapse. At x = 0, the gradient of the ¹⁸⁹ bore front is at its maximum which indicates bore collapse, following (4). The ¹⁹⁰ grey lines represent a subset of individual bores from the collected dataset and ¹⁹¹ show the variation in pre/post collapse bore front gradient.

In Figure 6, the solid blue line shows a relatively stable incoming bore with a bore height of 0.63 m, a terminal bore celerity 1.41 m/s and a Froude number, $Fr_{b,c} = 0.57$. The bore front slope at x = 0 m is approximately 17 degrees, 70% of the maximum steepness at bore collapse. Steepening of the bore front occurs until the bore collapses at x = 0 when the maximum slope (~ 25 degrees) is reached. Here, the terminal bore front slope angle is in the range of 12 to 35



Figure 5: Example of a swash event highlighting the bore collapse process/sequence observed at Nha Trang beach. The sequence shows: a) the incoming bore and the retreating toe of the previous bore 1 second prior to collapse, b) the bore at collapse, c) the initiation of swash motion following bore collapse 1 second post collapse, d) 4 seconds after the collapse flow divergence between the upper and lower parts of the swash flow, e) latter stages of backwash and f) the arrival of the subsequent bore, 10.5 seconds after a). The black and green dots indicate the shoreline tracking and the lines the calculated slopes during the collapsing process. Our detection here, mainly focusses on the b)-d) in which the green dots are used to determine the run-down limit at collapse in b) and maximum run-up as the upper green dot in d).



Cross shore distance relative to the collapse location [m]

Figure 6: Cross-shore variation of normalised bore front gradient. The grey lines represent the bore front gradient for an arbitrary set of bores. The blue line shows a representative bore discussed in the text. x = 0 is the point of bore collapse. The dashed black line is the mean gradient in space of all observed bores within this dataset and the error bars indicate the associated standard deviation.

degrees which is much flatter than previously observed in the laboratory [40] 198 or in a numerical test case [28] which indicated a near-vertical bore front at 199 collapse. After the moment the bore collapses, the collapsing bore slope reduces 200 at a higher rate compared to the steepening observed prior to collapsing. The 201 mean bore front gradient shows that the steepening typically occurs within 202 the last half metre before the bore collapse. This rapid process highlights the 203 need for high spatio-temporal resolution measurements to fully capture bore 204 collapse. The individual bore collapse signatures shown in Figure 6 highlight 205 the variability of this process. 206

²⁰⁷ 3.2. Observed bore celerity

Bore celerity through the surf and swash zone can be estimated through 208 the bore tracking methodology. Detection and magnitude of the bore celerity 209 is influenced by changes in bore shape, front slope changes and instabilities. 210 Tracking the bore-head typically over-estimates the celerity as the front steep-211 ens. Likewise, tracking the bore toe leads to an underestimate. It was found 212 that the most stable results were obtained by tracking the bore's maximum 213 gradient. Figure 7 represents the celerity corresponding to the same bores as 214 shown in Figure 6. 215



Cross shore distance relative to the collapse location [m]

Figure 7: Cross-shore variation of normalised bore celerity (before collapse) and normalised shoreline velocity (after bore collapse). The grey lines represent bore celerity/shoreline velocity for the same bores as in Figure 6. The blue line shows a representative bore. The dashed black line is the mean bore celerity in space of all observed bores for this dataset and the error bars indicate the associated standard deviation

The bore celerity in Figure 7 is normalized by the minimum bore celerity for 216 each detected wave. Minimum bore celerities are therefore indicated by a value 217 of 1. The blue solid line indicates the estimated celerity for the same bore as 218 highlighted in Figure 6. Prior to bore collapse, a reduction of the bore celerity 219 can be observed as the bore approaches zero depth. Minimum bore celerity is 220 reached at the point of bore collapse (x = 0), where the absolute bore celerity is 221 1.41 m/s for the highlighted bore. Immediately following bore collapse, a rapid 222 acceleration occurs as the swash flow is initiated [19]. The average bore celerity 223 (black dashed line) shows a very similar behaviour with a deceleration prior 224 to the minimum value at bore collapse and subsequent acceleration following 225 the collapse process. The significantly larger error bars after collapse can be 226 explained by the fact that the swash tip is significantly harder to detect due to 227 the flattening of the front (swash-tip) slope and small flow depths. 228

229 3.3. Wave run-up

As discussed in the Methods section, previous authors have suggested that the vertical run-up can be considered a function of the bore height at collapse using (3). Figure 8 shows the vertical run-up as a function of the terminal bore height where the lines indicate constant values of the coefficient C. It can be



Figure 8: Bore height versus vertical run-up per detected bore collapse. The colour indicates the local bore related Froude number following (6). The lines represent C thresholds, the red dashed line is C = 2 (perfect conversion), 95% threshold is represented by blue line and the solid green line indicates the 99% interval of the scatter and the theoretical minimum C(C = 1) is represented by the dark black line.

observed that the variability of C is much greater in the current field dataset 234 compared to the large flume experiments described by Blenkinsopp et al. [11] 235 where values of C were between 1.95 and 2.25 for monochromatic waves. The 236 average value of C for the current data is 1.79 with a standard deviation of 0.265, 237 compared to a mean \overline{C} of 2.09 and standard deviation of 0.08 in Blenkinsopp 238 et al. [11]. Here, we find the majority (75.2%) of the bores have a C-value 239 indicating an imperfect conversion form potential to kinetic energy (C < 2)240 while a significant portion of the bores experience greater run-up than predicted 241 by equation (3) assuming a perfect conversion (C = 2). It is suggested that while 242 the bore collapse process is the primary factor in determining initial swash 243 velocity, other processes including swash-swash interaction and terminal bore 244 celerity seem to contribute. Note that C-values greater than 2 have previously 245 been observed in laboratory experiments with fully developed [7] and solitary 246 bores [16]. 247

248

The dots in Figure 8 are coloured according to the Froude number at bore

collapse defined in (6). It is observed that relatively small values of $Fr_{b,c}$ tend 249 to correspond to lower values of C implying that bores arriving with a rela-250 tively low celerity tend to generate smaller than expected run-up. Observations 251 suggest that such events typically feature strong interaction between the pre-252 ceding backwash flow and the incoming bore which acts to retard wave run-up. 253 Conversely, relatively large values of $Fr_{b,c}$ tend to correspond with the higher C-254 values, indicating larger than expected run-up for a given bore collapse height. 255 In this case, observations indicate that such events correspond to overtaking 256 swash events according to the definitions of Hegge and Eliot [18] which act 257 to enhance wave run-up. Over the total dataset 50.3% of the bores are free 258 swashes without bore-bore interaction. 19.3% of the bores are overtaken by 259 the subsequent incident bore (6.2%) of the dataset consists of the subsequent 260 bores). Overriding and suppressed bores collectively comprise 16.8% of the 261 total dataset and the remaining 7.4% corresponds to composite swash-swash 262 interactions. While there is clear scatter in measured values of the coefficient 263 C, if the average measured value (1.79 as found above) is taken to calculate the 264 run-up for each swash event in the dataset using (3), the RMS error is 0.325 m. 265 The LiDAR data and tracking routines allow individual incident bores and 266 bore-pairs to be tracked, which enables an analysis of bore-bore interactions. 267 Figure 9 shows 3 example cases with increasing swash duration (hence reducing 268 swash saturation as described in (6)) from left to right: overtaking (a), (par-269 tially) suppressing (b) and a free bore (c). In all of the presented cases, the bore 270 heights of two consecutive bores are of similar order at -17.5 m cross shore (sea-271 ward of all collapse locations for the presented bores). The maximum difference 272 is 6 cm, which corresponds to 5% of the maximum height of the two consecutive 273 bores. Figure 9d-f shows the variation in the time between the two consecutive 274 bores as they progress shoreward, and these demonstrate a characteristic be-275 haviour for the different types of bore-bore interaction. For the overtaking case 276 (Figure 9 a and d), the second bore propagates before the flow reversal of the 277 preceding bore, thus it travels in a greater depth, the bore is partially-advected 278 by the uprush flow and it collapses further landward. As a result, the time 279



Figure 9: Examples of bore-bore interactions. a) represents an extended run-up due to overtaking, b) indicates a suppressing bore-interaction and c) shows a free swash movement. The black line shows the bore-track of the first incident bore and the red represents the second. The circles represent the detected bore collapse point. d)-f) show corresponding relative time between two consecutive bores as the bores propagate inshore. The black and red dashed lines correspond to the cross shore location of the bore collapse for the first (black) and second (red) bore. The green dashed line represents the time between two consecutive bore collapse events and the blue line shows the time between incident bores propagating inshore.

between consecutive bores reduces as they move landward, and the maximum 280 run-up is much greater for the second bore despite the fact that the bores had 281 the same height at x = -17.5 m. In the partially suppressing case (Figure 9 b and 282 e), the second bore propagates on a seaward-directed backwash flow which holds 283 the bore back prior to collapse, making the bore collapse further seaward and 284 reducing the maximum run-up. The time between consecutive bores reduces in 285 the landward direction due to both a lower terminal velocity and smaller bore 286 collapse height of the second bore. Finally, in the free bore case (Figure 9 c 287 and f) the bore collapse position of the second wave is almost unaffected by the 288 preceding bore, and the run-up for both bores is very similar 289

290 4. Discussion

The results above show a majority of bores with a C-value smaller than 291 The results above show a majority of bores with a C-value smaller than 292 2. However, C > 2 is found for a significant number of bores which can be 293 important for extrema-analyses such as run-up predictions. C-values are often 294 greater than predicted by a conversion of potential to kinetic energy during 295 bore collapse and appear to be influenced by the local Froude number. If we 296 assume that the terminal bore celerity $c_{b,c}$ contributes directly to the initial 297 swash velocity U_0 , we can rewrite (2) as:

$$U_0 = c_{b,c} + \alpha \sqrt{gH_{b,c}} \tag{7}$$

by substituting U_0 using (7) instead of (2) in (1), the vertical run-up can be calculated in terms of the local bore related Froude number (8):

$$R = \frac{\left[c_{b,c} + \alpha \sqrt{gH_{b,c}}\right]^2}{2g} \iff R = \frac{\left[\left(Fr_{b,c} + \alpha\right) \sqrt{gH_{b,c}}\right]^2}{2g} \tag{8}$$

The term in parentheses in the right part of (8) effectively represents C, which consists of the bore related Froude number and a new conversion coefficient α , as presented in (9). Considering (9) and by rearranging (3), α is then related to the run-up and bore height as presented in (10).

$$C = Fr_{b,c} + \alpha \tag{9}$$

$$\alpha = \sqrt{\frac{2R}{H_{b,c}}} - Fr_{b,c} \tag{10}$$

Notably, the definition of C compared to Baldock and Holmes [8] has not changed other than that C is now defined by a celerity component and a conversion efficiency component α which in the case of a perfect conversion of potential to kinetic energy will take the value 2 as in the earlier formulation [40]. Also, the left hand side of (8) allows for a component expansion which suggests that R is a function of two physical components: terminal bore celerity and the conversion efficiency in the bore collapse process. In addition to these two distinct ³¹¹ physical processes, (11) and (12) also include a third term which incorporates
 ³¹² non-linear interactions between them:

$$R = \frac{c_{b,c}^2}{2g} + \frac{c_{b,c}\alpha\sqrt{gH_{b,c}}}{g} + \frac{\left(\alpha\sqrt{gH_{b,c}}\right)^2}{2g} \tag{11}$$

$$R = R_c + R_{\alpha,c} + R_\alpha \tag{12}$$

In (12), R_c is the run-up component related solely to the terminal bore 313 celerity, R_{α} is the component related solely to bore collapse and $R_{\alpha,c}$ represents 314 non-linear interaction between the terminal celerity and collapse. C and α are 315 similar conversion coefficients respectively with and without a terminal bore 316 celerity component. The upper plot in Figure 10 shows total vertical run-up 317 as a function of the conversion coefficient C^2 . In the lower plot in Figure 10 318 the component of the vertical run-up due to the terminal bore celerity R_c is 319 subtracted from the total run-up, leaving the components of run-up that are 320 related to the terminal bore height at collapse and this is shown as a function 321 of α^2 . 322

From Figure 10 it is evident that by removing the component of run-up directly caused by the terminal bore celerity, the relationship between the remaining components of run-up and the bore collapse height is strengthened, as indicated by the reduced scatter. Thus, the results in Figure 10b indicate the energy converted by the bore collapse process through α , and the scatter represents the non-linear interaction with the terminal bore celerity. To further highlight this, we present $R_{\alpha,c}$ and R_c as a function of α in Figure 11.

Figure 11 shows that R_{α} is smaller than the non-linear interaction term $R_{\alpha,c}$ for most range of α , until α exceeds a value of 1.5. The colouring suggests that even when the value of α is small, substantial vertical run-up can occur when there is a large value of the terminal bore celerity. It is also evident that the higher values of terminal bore celerity are typically related to low values of α and for high values of α , terminal bore celerity tends to be relatively small. For the lower (greater) values of α this suggests a larger (lower) relative contribution of



Figure 10: Scatter plots of run-up versus conversion coefficients C and α . a) presents the total run-up as a function of C^2 . b) shows the relation between α^2 and run-up without the terminal celerity component. The colour of the symbols represents the terminal bore celerity. The red lines in both plots represent the linear fit with details presented in the lower right corner of each panel.



Figure 11: Scatter plot of α compared to run-up related to c_{bc} and α represented by the coloured dots $R_{\alpha,c}$, whereas the black dots show the run-up related to the energy conversion R_{α} . The red lines represent the quadratic fit with between α and R_{α} (solid) and the linear fit between α and $R_{\alpha,c}$ (dashed).

the terminal bore celerity to the total run-up in comparison to the terminal bore 337 height for the $R_{\alpha,c}$ term. The solid red line in Figure 11 represents a quadratic 338 fit to the R_{α} component which shows a good and significant correlation to α (r 339 = 0.894 with an associated p-value of 0.004). Since the non-linear interaction 340 term is also dependent on the terminal bore celerity and non-linear in nature, 341 the correlation with α is weaker (R = 0.268 and a p-value of 0.002). On average, 342 for all the bores within this dataset, the contribution to the total run-up from 343 R_{α} 26%, the R_c term accounts for 27% while the contribution of $R_{\alpha,c}$ is 47%. 344 This analysis indicates that while the collapse and terminal celerity mechanisms 345 contribute equally to the total run-up, the non-linear interaction term clearly 346 dominates. This highlights the significance of including celerity component to 347 approximate run-up, suggesting that its direct and indirect impact on the run-348 up and bore collapse process should not be neglected. With the inclusion of 349 bore front celerity, the run-up can be estimated more accurately from measured 350 bore collapse parameters. Using (8) and the average observed α ($\overline{\alpha_{obs}} = 0.889$) 351 the RMS error for the run-up is reduced by approximately 10% to 0.295 m. 352

In the previous model by Baldock and Holmes [8], C can be seen as a reposi-353 tory of all unknown processes and interactions that occur between the inner surf 354 and swash zone [36]. A direct link between α and C with terminal bore celerity 355 or other measured components was sought to enable improved prediction of in-356 dividual wave run-up based on measured bore properties. Attempts were made 357 to relate C and α to the incident bore front slope, the slope of the free-surface 358 behind the bore front (see green dashed line in Figure 4) and the relative angle 359 between the two, but no significant relationship could be found other than a 360 weak trend found between back angle and C. Landward-sloping free surfaces 361 behind the bore front were found to be associated with greater C-values, while 362 seaward-slope free surfaces behind the bore front tended to have lower C-values. 363

364 5. Conclusion

A 2D LiDAR scanner has been used to obtain high spatial and temporal resolution water surface profiles illustrating the complex bore collapse process. From the 2D LiDAR data, it is possible to accurately obtain the bore collapse point in space and time and extract a range of parameters including bore celerity, bore slope and bore height at collapse. It is observed that the terminal bore celerity at collapse is consistently non-zero and the bore collapse front slope is in the range 12-35 degrees to the horizontal.

In agreement with other studies, a clear relationship between wave run-up 372 and bore height at collapse was observed. However, the measurements obtained 373 by tracking incident bores using the LiDAR enabled further analysis of the 374 underlying mechanisms causing wave run-up. This indicates that terminal bore 375 celerity at the point of bore collapse contributes significantly to individual wave 376 run-up and is strongly influenced by bore-bore interactions in the inner surf 371 zone. Term-expansion of an existing ballistic-type model to describe the run-up 378 of individual waves in combination with the novel measurements showed that 379 the total run-up R could be separated into three different components: bore 380 collapse conversion efficiency, bore celerity and their non-linear interaction. In 381 the dataset presented here, the bore collapse and terminal bore celerity have an 382 equal contribution, while the non-linear interaction between the two dominates 383 the total run-up. This analysis of the driving mechanisms which cause wave 384 run-up, shows that the former conversion coefficient C can be separated into 385 three components: the bore collapse, terminal bore-celerity and their non-linear 386 interaction. Hence, including terminal celerity with collapsing bores cannot be 387 neglected when investigating or predicting wave run-up at sandy beaches. 388

389 6. Acknowledgements

The authors would like to acknowledge the financial assistance and funding for E. Bergsma, C. Blenkinsopp and K. Martins through the WAves into Shallow water (WASH) project provided by the Engineering and Physical Sciences ³⁹³ Research Council (EP/N019237/1). The Nha Trang experiment, R. Almar and

³⁹⁴ L-P Melo de Almeida was funded by the French ANR grant (COASTVAR:

³⁹⁵ ANR-14-ASTR-0019)

396 References

- [1] Almar, R., Marchesiello, P., Almeida, L., Hai, T., Tanaka, H., Nguyen, V.,
 2017. Shoreline response to a sequence of typhoon and monsoon events.
 Water vol. 364, pp. 9(6).
- [2] Almeida, L., Almar, R., Marchesiello, P., Benshila, R., Martins, K., Blenkinsopp, C., Floc'h, F., Ammann, J., Grandjean, P., Viet, N., Thuan, D.,
 Binh, L., Sénéchal, N., Detandt, G., Biausque, M., Garlan, T., Bergsma,
 E., Caulet, C., Tran, H.-Y., 2016. Swash zone dynamics of a sandy beach
 with low tide terrace during variable wave and tide conditions. In: In proceedings of the XIVmes Journées Nationales Génie Cotier Génie Civil.
 p. 2.
- [3] Almeida, L., Masselink, G., Russell, P., Davidson, M., 2015. Observations
 of gravel beach dynamics during high energy wave conditions using a laser
 scanner. Coastal Engineering vol. 228, pp. 15–27.
- [4] Alsina, J., Cáceres, I., 2011. Sediment suspension events in the inter surf
 and swash zone. measurements in large scale and high-energy wave conditions. Coastal Engineering vol. 58, pp. 657–670.
- [5] Alsina, J., Falchetti, S., Baldock, T., 2009. Measurements and modelling of
 the advection of suspended sediment in the swash zone by solitary waves.
 Coastal Engineering vol. 56, pp. 621–631.
- [6] Alsina, J., van der Zanden, J., Cáceres, I., Ribberink, J. S., 2018. The
 influence of wave groups and wave-swash interactions on sediment transport
- and bed evolution in the swash zone. Coastal Engineering, in press.

- [7] Baldock, T., Cox, D., Maddux, T., Killian, J., Fayler, L., 2009. Kinematics
 of breaking tsunami wavefronts: A data set from large scale laboratory
 experiments. Coastal Engineering vol. 56, pp.506–516.
- [8] Baldock, T., Holmes, P., 1999. Simulation and prediction of swash oscillations on a steep beach. Coastal Engineering vol. 36, pp. 219–242.
- ⁴²⁴ [9] Battjes, J., 1974. Surf similarity. Coastal Engineering Proceedings 1 (14).
- [10] Blenkinsopp, C., Masselink, G., Turner, I., Russel, P., 2011. Can standard
 energetics models be used to predict net cross-shore sediment flux at the
 beach face? Australian Journal of Civil Engineering vol. 9(1), 16 p.
- [11] Blenkinsopp, C., Matias, A., Howe, D., Castelle, B., Marieu, V., Turner, I.,
 2016. Wave runup and overwash on a prototype-scale sand barrier. Coastal
 Engineering vol. 113, pp. 88103.
- [12] Briganti, R., Torres-Freyermuth, A., Baldock, T. E., Brocchini, M.,
 N.Dodd, Hsu, T., Jiang, Z., Kim, Y., Pintado-Patino, J., Postacchini, M.,
 2013. Advances in numerical modelling of swash zone dynamics. Coastal
 Engineering vol. 115, pp. 26–41.
- [13] Cáceres, I., Alsina, J., 2012. A detailed event-by-event analysis of suspended sediment concentration in the swash zone. Continental Shelf Research vol. 21 (6), pp. 1219–1227.
- [14] Chen, B.-T., Kikkert, G., Pokrajac, D., Dai, H.-J., 2016. Experimental
 study of bore-driven swashswash interactions on an impermeable rough
 slope. Coastal Engineering vol. 108, pp. 10–24.
- ⁴⁴¹ [15] Grilli, S., Svendsen, I., Subramanya, R., 1997. Breaking criterion and char⁴⁴² acteristics for solitary waves on slopes. Journal of Waterway, Port, Coastal,
 ⁴⁴³ and Ocean Engineering vol. 123, pp. 102–122.
- [16] Guard, P., Baldock, T., 2009. The influence of seaward boundary conditions
 on swash zone hydrodynamics. Coastal Engineering vol. 56, pp.506–516.

- [17] Hedges, T. S., Mase, H., 2004. Application of a non-linear shallow water theory to swash following bore collapse on a sandy beach. Journal of
 Waterway, port, coastal and ocean engineering vol. 130, pp. 109–113.
- [18] Hegge, B., Eliot, I., 1991. Swash interactions on sandy beaches. In: Proceedings of the 10th Australian Conference on Coastal and Ocean Engineering.
- [19] Howe, D., 2016. Bed shear stress under wave runup on steep slopes. Ph.D.
 thesis, University of New South Wales, Australia.
- ⁴⁵⁴ [20] Hughes, M. G., 1992. Application of a non-linear shallow water theory to
 ⁴⁵⁵ swash following bore collapse on a sandy beach. Journal of Coastal Research
 ⁴⁵⁶ vol. 8(3), pp. 562–578.
- ⁴⁵⁷ [21] Hughes, M. G., Moseley, A. S., 2007. Hydrokinematic regions within the
 ⁴⁵⁸ swash zone. Continental Shelf Research vol. 27(15), pp. 2000–2013.
- ⁴⁵⁹ [22] Kikkert, G., Pokrajac, D., O'Donoghue, T., Steenhauer, K., 2013. Exper⁴⁶⁰ imental study of bore-driven swash hydrodynamics on permeable rough
 ⁴⁶¹ slopes. Coastal Engineering vol. 79, pp. 42–56.
- ⁴⁶² [23] Martins, K., Blenkinsopp, C., Deigaard, R., Power, H., 2018. Energy dissi⁴⁶³ pation in the inner surf zone: New insights from lidar-based roller geometry
 ⁴⁶⁴ measurements. Jour. of Geophysical Research: Oceans vol. 123, pp. 3386–
 ⁴⁶⁵ 3407.
- ⁴⁶⁶ [24] Martins, K., Blenkinsopp, C., Zang, J., 2016. Monitoring individual wave
 ⁴⁶⁷ characteristics in the inner surf with a 2-dimensional laser scanner (lidar).
 ⁴⁶⁸ Journal of Sensors vol. 2016, 11 pages.
- ⁴⁶⁹ [25] Masselink, G., Puleo, J., 2006. Swash zone morphodynamics. Continental
 ⁴⁷⁰ Shelf Research vol. 26, pp. 661–680.
- ⁴⁷¹ [26] Masselink, G., Russell, P., Turner, I., Blenkinsopp, C., 2009. Net sediment
 ⁴⁷² transport and morphological change in the swash zone of a high-energy

- sandy beach from swash event to tidal cycle time scales. Marine Geology
 vol. 267, pp. 18–35.
- ⁴⁷⁵ [27] Mau, L., 2014. Overview of Natural Geographical Conditions of Nha Trang
 ⁴⁷⁶ Bay. Nha Trang: Institute of Oceanography.
- ⁴⁷⁷ [28] Mory, M., Abadie, S., Mauriet, S., Lubin, P., 2011. Run-up flow of a col⁴⁷⁸ lapsing bore over a beach. European Journal of Mechanics B/Fluids vol.
 ⁴⁷⁹ 30, pp. 565–576.
- ⁴⁸⁰ [29] Pujara, N., Lui, P., Yeh, H., 2015. An experimental study of the interaction
 ⁴⁸¹ of two successive solitary waves in the swash: A strongly interacting case
 ⁴⁸² and a weakly interacting case. Coastal Engineering vol. 105, pp. 66–74.
- [30] Puleo, J., Beach, R., Holman, R., Allen, J., 2000. Swash zone sediment
 suspension and transport and the importance of bore-generated turbulence.
 Journal of Geophysical Research vol. 105, pp. 17021–17044.
- [31] Puleo, J., Butt, T., 2006. The first international workshop on swash-zone
 processes. Continental Shelf Research vol. 26, pp. 556–560.
- [32] Puleo, J., Holland, K., 2001. Estimating swash zone friction coefficients on
 a sandy beach. Coastal Engineering vol. 43(1), pp. 25–40.
- [33] Puleo, J., Torres-Freyermuth, A., 2016. The second international workshop
 on swash-zone processes. Coastal Engineering vol. 115, pp. 17.
- [34] Shen, M. C., Meyer, R. E., 1963. Climb of a bore on a beach, part 3 run-up.
 Journal of Fluid Mechanics vol. 14, pp. 305–318.
- ⁴⁹⁴ [35] Stockdon, H., Holman, R., Howd, P., Sallenger, A., 2006. Empirical pa⁴⁹⁵ rameterization of setup, swash, and runup. Coastal Engineering vol. 53 (7),
 ⁴⁹⁶ pp. 573–588.
- ⁴⁹⁷ [36] Svendsen, I., 2006. Introduction to Nearshore Hydrodynamics, Advanced
 ⁴⁹⁸ series on ocean engineering. World Scientific.

- ⁴⁹⁹ [37] Thuan, D., Binh, L., Viet, N., Hanh, K., Almar, R., Marchesiello, P.,
 ⁵⁰⁰ 2016. Typhoon impact and recovery from continuous video monitoring: a
 ⁵⁰¹ case study from nha trang beach, vietnam. In: Vila-Concejo, A.; Bruce,
 ⁵⁰² E.; Kennedy, D.M., and McCarroll, R.J. (eds.), Proceedings of the 14th
 ⁵⁰³ International Coastal Symposium (Sydney, Australia).
- ⁵⁰⁴ [38] Whitham, G., 1958. On the propagation of shock waves through regions of ⁵⁰⁵ non-uniform area or flow. Journal of Fluid Mechanics vol. 4, pp. 520–539.
- [39] Yeh, H., Ghazali, A., 1988. On bore collapse. Journal of Geophyical Re search vol. 93 (C), pp. 6930–6936.
- [40] Yeh, H., Ghazali, A., Marton, I., 1989. Experimental study of bore run-up.
 Journal of Fluid Mechanics vol. 206, pp. 563–578.
- [41] Zhang, Q., Liu, P. L.-F., 2008. A numerical study of swash flows generated
 by bores. Coastal Engineering vol. 55, pp. 1113–1134.