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## 1 Mineral characterization of the biogenic Fe(III)(hydr)oxides produced

## 2 during Fe(II)-driven denitrification with Cu, Ni and Zn

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#### 1 Abstract

2 The recovery of iron and other heavy metals by the formation of Fe(III) (hydr)oxides is an important 3 application of microbially-driven processes. The mineral characterization of the precipitates formed 4 during Fe(II)-mediated autotrophic denitrification with and without the addition of Cu, Ni, and Zn by four different microbial cultures was investigated by X-ray fluorescence (XRF), Raman 5 6 spectroscopy, scanning electron microscopy equipped with energy dispersive X-Ray analyzer 7 (SEM-EDX), Fourier transform infrared spectroscopy (FTIR) and X-ray Powder Diffraction (XRD) 8 analyses. Fe(II)-mediated autotrophic denitrification resulted in the formation of a mixture of Fe(III) 9 (hydr)oxides composed of amorphous phase, poorly crystalline (ferrihydrite) and crystalline phases 10 (hematite, akaganeite and maghemite). The use of a *Thiobacillus*-dominated mixed culture enhanced 11 the formation of akaganeite, while activated sludge enrichment and the two pure cultures of T. 12 denitrificans and Pseudogulbenkiania strain 2002 mainly resulted in the formation of maghemite. 13 The addition of Cu, Ni and Zn led to similar Fe(III) (hydr)oxides precipitates, probably due to the 14 low metal concentrations. However, supplementing Ni and Zn slightly stimulated the formation of 15 maghemite. A thermal post-treatment performed at 650°C enhanced the crystallinity of the 16 precipitates and favored the formation of hematite and some other crystalline forms of Fe associated 17 with P, Na and Ca.

18

19 Keywords: ferrous iron, nitrate, autotrophic denitrification, heavy metals, biorecovery

#### 20 1. Introduction

The need for iron and other metals from secondary sources is nowadays gaining increasing 21 22 attention. Most of the recovery options currently used is not economically feasible and, therefore, it 23 is critical to develop low-cost and environmental-friendly methods to recover metals from 24 wastewaters (Pat-Espadas and Cervantes, 2018). Fe(II)-mediated autotrophic denitrification is a 25 promising biotechnology to be employed for iron recovery when nitrogen and metal contamination coexists. Chemoautotrophic denitrifiers are capable of coupling Fe(II) oxidation with denitrification 26 27 resulting in the bioreduction of nitrate to nitrogen gas and the production of Fe(III) precipitates, 28 which can be subsequently removed and recovered (Kiskira et al., 2017a). Different biogenic 29 Fe(III)(hydro)oxide mineral phases such as goethite, lepidocrocite, hematite, magnetite, maghemite, 30 2-line ferrihydrite and mixed-phase Fe(II)-Fe(III) minerals (i.e. green rust) have been reported as 31 potential Fe-based products from Fe(II)-driven denitrification (Senko et al., 2005a, b; Weber et al., 32 2006; Konhauser et al., 2011). 33 A lot of environmental and operating parameters affect the formation of Fe(III) precipitates. Among them, the use of different microbial species and their metabolic rates significantly result in 34

35 diverse biogenic Fe(III) (oxyhydr)oxide minerals from soluble Fe(II) (Posth et al., 2014). Other

36 factors are the medium composition, the concentration of possible co-substrates, the incubation

37 conditions and the presence of nucleation sites (Kappler and Straub, 2005; Senko et al., 2005a;

38 Larese-Casanova et al., 2010; Pantke et al., 2012). A slower Fe(II) oxidation rate produces

39 crystalline minerals, such as goethite, while a rapid Fe(II) oxidation rate leads to the formation of

40 less crystallized phases such as ferrihydrite (Lack et al., 2002; Senko et al., 2005b). pH is another

41 factor that affects the formation and composition of Fe(III) precipitates. For instance, goethite

formation was reported at pH 7.7, while the production of lepidocrocite occurred at pH 6.3 in the
presence of *Acidovorax* sp. BoFeN1 (Larese-Casanova et al., 2010).

44 Due to their larger specific surface area and higher binding energy, biogenic oxides are more 45 efficient for the recovery of other metals and metalloids than chemically produced oxides through co-precipitation or adsorption mechanisms (Hennebel at al., 2009). For instance, arsenic (As) has 46 47 been reported to effectively co-precipitate with biogenic goethite and lepidocrocite (Senn and 48 Hemond, 2002; Hohmann et al., 2009; Ahoranta et al., 2016; Xiu et al., 2016), with As being 49 adsorbed after its oxidation from As(III) to As(V) (Papirio et al., 2014). The microbial oxidation of 50 As(III) and Fe(II) linked to denitrification resulted in an enhanced immobilization of As(V) onto 51 biogenic Fe(III) (hydr)oxides in sand-packed bed columns (Sun et al., 2009). Pantke et al. (2012) 52 reported that green rust is a highly reactive iron mineral and can be used for the recovery of 53 chromium (Cr) and arsenic. Furthermore, Fe(III) oxides can be recycled and used for phosphorus 54 removal (Nielsen and Nielsen, 1998; Zhang et al., 2015).

55 The determination of the mineral identity of the Fe(III) precipitates formed is crucial in order 56 to decide the adoption of the "reuse strategy" and the possible applications (Velenturf and Jopson, 57 2019). Iron oxides are valuable minerals for many applications due to their low cost, abundance and 58 physico-chemical properties (Pouran et al., 2014; dos Santos et al, 2016). Among Fe(III) 59 (hydr)oxides, hematite, magnetite, magnetite and goethite have the greatest interest in industry 60 (Cao et al., 2015). Some of the major applications of recovered biogenic Fe(III) minerals are the 61 production of pigments for paints and the construction industry, magnetic pigments and ferrites, 62 catalysts for industrial synthesis, photo-chemicals and fertilizers, heavy metals immobilization, 63 adsorbents for water and gas purification as well as radioactive waste decontamination (Cornell and 64 Schwertmann, 2003; Chen et al., 2017; Fajardo et al., 2019).

65 A lack of knowledge exists on the mineralogical characterization of the precipitates 66 produced during Fe(II)-mediated autotrophic denitrification by different microbial species. Indeed, 67 most of the studies only dealt with the use of a single microbial culture, or did not investigate a 68 strategy to evaluate the potential of biogenic Fe(III) precipitates to incorporate other metals (Kiskira 69 et al., 2017a). To the best of authors' knowledge, only one study besides this investigated the 70 removal of other metals/metalloids (i.e. As) through the use of biogenic Fe(III) precipitates obtained 71 by Fe(II)-driven denitrification (Sun et al., 2009). In addition, this study is the first proposing the use 72 of a thermal post-treatment of amorphous phases or poorly crystalline biogenic Fe(III) precipitates 73 to increase their crystallinity, depending on their reuse strategy. 74 In particular, the objectives of this work were to: a) evaluate and characterize the different 75 Fe(III) precipitates produced using different microbial inocula by different mineral characterization 76 methods, i.e. X-ray Powder Diffraction (XRD), X-ray fluorescence (XRF), Raman spectroscopy, scanning electron microscopy equipped with an energy dispersive X-Ray analyzer (SEM-EDX) and 77 78 Fourier transformation infrared spectroscopy (FTIR); b) use a thermal post-treatment to improve the

crystallinity of the precipitates; c) investigate the potential adsorption/co-precipitation of Cu, Ni and
Zn on Fe(III) precipitates.

81

## 82 2. Materials and methods

#### 83 2.1. Sources of microorganisms and mineral growth media

Fe(III) precipitates were obtained from the batch experiments carried out by Kiskira et al.
(2018) with four chemolithotrophic denitrifying cultures in serum bottles. A *Thiobacillus*-dominated
mixed culture, previously used in experiments with thiosulfate (Di Capua et al., 2016; Zou et al.,

87 2016) and Fe(II) (Kiskira et al., 2017b) as electron donors, and an activated sludge enrichment, 88 originally collected from the municipal wastewater treatment plant in Cassino (Italy), were used as 89 mixed bacterial consortia. Pseudogulbenkiania sp. 2002 (DSM 18807) and T. denitrificans (DSM 90 12475) were obtained from the 'Leibniz-Institute DSMZ - German collection of microorganisms 91 and cell cultures' and used as pure cultures. All the cultures were supplemented to the serum bottles 92 with a 10% (v/v) amount. The initial concentration of *Thiobacillus*-mixed, *T. denitrificans* pure, 93 activated sludge and *Pseudogulbenkiania* sp. 2002 cultures in terms of volatile suspended solids 94 (VSS) was 200, 320, 410 and 750 mg VSS/L (Kiskira et al., 2018).

95 The denitrification potential of all cultures with Fe(II) was preliminarily evaluated in the 96 absence (Kiskira et al., 2017b) and in the presence of Cu, Ni and Zn (Kiskira et al., 2018) in 97 experiments performed at room temperature (i.e.  $22 \pm 2^{\circ}$ C), under helium atmosphere and with an 98 initial pH of 7.0. A circumneutral pH was maintained along the 10-days experiments due to the 2 99 g/L NaHCO<sub>3</sub> provided as both buffer and inorganic carbon source. Therefore, in this study the effect 100 of pH on the formation and composition of Fe(III) precipitates was not investigated. In addition to 101 NaHCO<sub>3</sub>, the basal medium also contained (g/L) 0.25 NH<sub>4</sub>Cl, 0.30 KH<sub>2</sub>PO<sub>4</sub>, 0.40 K<sub>2</sub>HPO<sub>4</sub>, and 0.10 102 NaCl. The trace elements solution was added from a sterile stock solution as described in Weber et 103 al. (2001). Cu, Ni and Zn were added in concentration of 40 mg/L in the form of copper chloride 104 (CuCl<sub>2</sub>), nickel chloride (NiCl<sub>2</sub>·6H<sub>2</sub>O) and zinc chloride (ZnCl<sub>2</sub>), all purchased from Sigma Aldrich 105 (Germany). The feed Fe(II) and NO<sub>3</sub><sup>-</sup> concentrations were 600 and 120 mg/L, supplemented in the 106 form of iron(II) chloride (FeCl<sub>2</sub>·4H<sub>2</sub>O) and sodium nitrate (NaNO<sub>3</sub>), respectively.

#### 107 **2.2. Sample preparation and analytical methods**

108 The mineral characterization of the Fe(III) precipitates formed during Fe(II)-mediated
109 autotrophic denitrification was performed on: i) four samples obtained in the absence of metals, with

110 the use of the Thiobacillus-dominated mixed culture (TM), T. denitrificans (TDP), activated sludge 111 enrichment (AS) and Pseudogulbenkiania strain 2002 (PG); ii) twelve samples obtained in the 112 experiments with Cu (TMC, TDPC, ASC, PGC), Ni (TMN, TDPN, ASN, PGN) and Zn (TMZ, 113 TDPZ, ASZ, PGZ). Furthermore, the precipitates from abiotic control (AC) experiments without 114 metals (Cu, Ni and Zn) were also characterized. Abiotic controls were also performed with the 115 addition of Cu, Ni and Zn as described by Kiskira et al. (2018), but the authors observed an 80% 116 lower precipitation of metals in the abiotic controls than that achieved in the biotic tests. Due to that, 117 the precipitates obtained in the abiotic controls with Cu, Ni and Zn were not analyzed in this study 118 because considered similar to those from abiotic tests without metals.

119 All the Fe(III)(hydr)oxides production tests were performed in duplicate. The precipitates 120 obtained from two duplicates were mixed prior to being analyzed with the different methods due to 121 the limited amount of precipitate obtained in each bottle. Afterwards, Fe(III) oxides were 122 pasteurized (80°C for 10 min), collected by centrifugation under anaerobic conditions, and washed 123 twice with a NaHCO<sub>3</sub> buffer (pH 6.8) as previously described (Weber et al., 2001). Then, the 124 precipitates were washed several times with double-distilled water by centrifugation, in order to 125 remove the impurity ions associated with the procedure, as described by Schwertmann and Cornell 126 (2000). Finally, the samples were dried at 40°C for 48 h prior to analysis.

127 2.2.1. X-ray fluorescence, scanning electron microscopy and x-ray diffraction analyses

The samples were firstly analyzed by using X-ray fluorescence (XRF) with a Niton XL3t GOLDD + XRF analyzer (Thermo Fisher Scientific, USA) with  $\pm$  0.50 weight % precision in measurement. The loose powder method was chosen to analyze these samples, as described in detail by Takahashi (2015). The morphology of the samples after sputter coating with Au was also examined by scanning electron microscopy (SEM) using an XL series microscope (Philips, The

Netherlands). Additionally, SEM-EDX analysis was performed with a Vega 3 microscope 133 134 (TESCAN, Czech Republic), equipped with an energy dispersive QUANTAX X-Ray analyzer 135 (EDX) (Bruker, Czech Republic) for chemical analysis. Elemental identification and quantitative 136 compositional analysis were performed by XRF and, in some cases, were supported by SEM-EDX. 137 The analysis of total Fe oxides in the sample was made by XRF. The mineralogical characterization 138 of the powdered samples was performed by X-ray diffraction analysis (XRD) using a Panalytical 139 diffractometer (Almelo, the Netherlands) and Cu Ka radiation. The XRD interpretation software 140 was PDF-2 (Powdered Diffraction File-2) from the International Centre for Diffraction Data.

#### 141 **2.2.2.** Fourier transformation infrared spectroscopy

142 FTIR spectra were recorded with an IRAffinity-1 Fourier transformation infrared 143 spectrophotometer (Shimadzu, Japan), equipped with a deuterated-triglycine sulfate (DTGS) 144 detector using a wavelength between 400 - 7000 cm<sup>-1</sup>. Prior to being analyzed with FTIR, the 145 samples were transformed into pellets (diameter 7 mm) by using potassium bromide (KBr). This 146 method exploits the property that alkali halides (e.g. KBr) become plastic when subjected to 147 pressure and form a sheet that is transparent in the infrared region. KBr does not contain bands in 148 the mid-IR region of the spectrum and, therefore, the preparation as halide disks allows to lose less 149 information. FTIR was used to obtain a rapid and useful identification of Fe(III) precipitates and 150 information about the extent of metal substitution (Gotic and Music, 2007). Furthermore, with FTIR 151 spectroscopy, information about the adsorbed molecules and the nature of surface complexes could 152 be obtained (Cornell and Schwertmann, 2003).

153 **2.2.3. Raman spectroscopy** 

Raman measurements were carried out with an INVIA spectrometer (Renishaw, United
Kingdom) equipped with a microscope and a charge coupled device (CCD) detector. A 532 nm

solid-state green (Nd:YAG) laser was used with a maximum power of 5 mW. The acquisitions were performed through a x50 magnification objective (Leica, Germany) after achieving a calibration on a silicon standard. The Rayleigh scattering component was removed by an Edge filter, and the Raman-scattered light was dispersed by a holographic grating with 1800 lines/mm. The integration time was set at 60 s. Specific preparation of the sample was not required. The theoretical spot size was below 1 micrometer. Several spots on each sample were measured to avoid the measurement of a small impurity instead of the main mineral.

163 **2.2.4. Thermal post-treatment** 

A thermal post-treatment in a furnace at 650°C for 2 h was performed for some of the samples to favor the crystallization of amorphous phases (calcination). The weight loss was calculated according to (Eq. 1):

167 
$$Weight Loss (\%) = \frac{weight(25 \ ^{\circ}C) - weight(650 \ ^{\circ}C)}{weight (25 \ ^{\circ}C)} \times 100 \quad (Eq. 1)$$

168 The calcined products were characterized by XRD analysis.

169

#### 170 **3. Results and Discussion**

#### 171 **3.1.** Effect of the activity of Fe(II)-oxidizing denitrifying bacteria on the mineralogy of

172 precipitates

Table 1 reports the color, FTIR/Raman bands and the most intense XRD peaks (in interreticular spacings) of the minerals observed in this study. The FTIR, Raman and XRD peaks/bands were similar, indicating that the medium composition and the incubation conditions were likely important factors for the Fe(III) (hydr)oxides formation. Many impurities and different minerals
apart from Fe(III) (hydr)oxides were present in all precipitates.

178 Raman spectroscopy is an easy and fast method for the identification of iron oxides, but also 179 different iron hydroxides and iron oxyhydroxides especially in poorly crystallized materials, which 180 are often formed in bioprocesses (Hanesch, 2009). None of the samples observed with the 181 microscope equipped in the Raman spectroscopy was homogeneous. A mixture of several spots 182 (dark, white, red and orange) in microscale was observed (Fig. 1). The differences in color 183 corresponded to different minerals. The single spot size for the measurement was approximately 1 184 μm, and several spots on each sample were measured to better characterize the entire Fe(III) 185 (hydr)oxides. Raman spectra of some characteristic bands are shown in Fig. 2. Some spectra also showed a band at around 150-200 cm<sup>-1</sup> (Fig. 2B). Fig. 2A shows a typical 2-line ferrihydrite 186 187 (Hanesch, 2009) of a dark spot in the sample PGN.

188 SEM inspection showed varying sizes and diverse geometrical shapes such as irregular, rod, 189 modular and spherical in diameter ranging from few nanometers to 15 µm (Fig. 3). Similarly, 190 Kappler and Newman (2004) found that Fe(III) precipitates, from bottles inoculated with the R. 191 ferrooxidans strain SW2, showed a variety of particles in the range of hundreds of nm to 20 µm. 192 These particles showed various shapes, e.g. needles and plates, as well as irregular shapes, as also 193 reported by Kappler et al. (2005). Fig. 3C, D shows the crystalline phases of the biogenic Fe(III)-194 hydroxides obtained with the two pure cultures of *T. denitrificans* and *Pseudogulbenkiania* sp. 2002. 195 The biogenic minerals were in close association with the bacterial cell, forming a common globular 196 shape of biogenic minerals known as cell-mineral aggregate (Fig. 3C, D, Fig. 4B, C) (Schaedler et 197 al., 2009; Posth et al., 2014; Schmid et al., 2014). In comparison with the mixed cultures, the use of 198 pure cultures resulted in a higher amount of cell-mineral aggregates (Fig. 3).

199 All Raman spectra showed a band at around 1000 cm<sup>-1</sup>, most likely indicating the presence 200 of siderite, i.e. an iron carbonate mineral (FeCO<sub>3</sub>) (Fig. 2B, C) in agreement with the observations of 201 SEM-EDX analysis of some precipitates and the prediction of Visual MINTEQ reported by Kiskira 202 et al. (2018). Siderite spectra consisted of bands at 184, 287, 731 and 1090 cm<sup>-1</sup> (Rull et al., 2004), 203 with the most common band being at 1090 cm<sup>-1</sup>. The characteristic peak of siderite at  $32.2^{\circ}2\theta$  was 204 not observed by XRD analysis, as it likely overlapped with the hematite peak at  $33.2^{\circ}2\theta$  or the 205 siderite amount was not sufficient to be detected by XRD. Amounts of phosphate and carbonate 206 minerals were recorded by SEM-EDX. XRD and SEM analyses revealed the presence of vivianite, 207 i.e. Fe<sub>3</sub>(PO<sub>4</sub>)<sub>2</sub>·8H<sub>2</sub>O, in many experiments with microorganisms (Fig. 3A, 5, Table 1). Similar to this 208 study, the use of the Acidovorax strain BoFeN1 in the presence of Fe(II) and acetate resulted in a 209 mixture of Fe(III) (hydr)oxides, vivianite and amorphous Fe(III)-phosphate as reported by Miot et 210 al. (2009) and Schaedler et al. (2009).

211 Ferrihydrite was most probably contained in all the samples, in agreement with the majority 212 of studies on nitrate-dependent ferrous oxidation (Straub et al., 1996, 2004; Hafenbradl et al., 1996; 213 Lack et al., 2002; Senko et al., 2005 a, b; Hohmann et al., 2009; Chakraborty et al., 2011; Zhao et 214 al., 2013; Miot et al., 2014; Li et al., 2014). FTIR was not a suitable method for the identification of 215 ferrihydrite precipitates that are impure and poorly crystalized (2-line ferrihydrite). The Mössbauer 216 spectroscopy might have supported a better characterization of such precipitates (Cornell and 217 Schwertmann, 2003). However, poorly crystalline phases in aqueous solution have been reported to 218 transform into goethite, hematite, lepidocrocite and other crystalline forms depending on the pH and 219 solution chemistry (Cornell and Schwertmann, 2003; Kappler and Newman, 2004). The 220 transformation of poorly crystalline Fe(III) into more crystalline phases requires atomic 221 reorganization, dissolution and re-precipitation (Kappler and Newman, 2004). The process can be

222 catalyzed by reducing agents such as cysteine or Fe(II) (Cornell and Schwertmann, 2003). The 223 presence of reducing agents or Fe(II) in combination with surface bridging ligands can induce a 224 reductive dissolution of poorly crystalline Fe(III)-oxides and the subsequent re-precipitation of more 225 highly ordered phases (Cornell et al., 1989; Kappler and Newman, 2004). This stands in line with 226 this study, with ferrihydrite being always present with other more crystalline phases (Fig. 3A, Table 227 1). The transformation of the initial precipitates into more crystalline phases is generally faster in the 228 presence of bacteria, through acidification of the local environment followed by an acid-catalyzed 229 dissolution of Fe(III)-minerals or a reduction of Fe(III) (Kappler and Newman, 2004).

230 In this study, the mineralogy of the precipitates was only analyzed at the end of the 231 experiments, and no additional information can be given on the production rate of the minerals in 232 the presence of microorganisms. In terms of nitrate removal and Fe(II) oxidation, Kiskira et al. 233 (2018) observed the highest specific rates with the *Thiobacillus*-dominated mixed culture, while the 234 lowest specific rates were obtained in the bioassays with the *Pseudogulbenkiania* sp. 2002. A higher 235 Fe(II) oxidation rate has been related to the formation of less crystallized phases, while a lower 236 Fe(II) oxidation rate produces crystalline minerals (Lack et al., 2002; Senko et al., 2005b). However, 237 the effect of the reaction rate needs further investigation due to a mixture of Fe(III) (hydr)oxides of 238 amorphous phase, poorly crystalline and crystalline phases that are generally produced 239 simultaneously.

Table 2 reports the amount (%) of Fe oxides in each sample obtained in this study. The use of the *Thiobacillus*-mixed culture resulted in 61% of Fe(III) oxides in the corresponding precipitate. A slightly lower percentage of 57 and 55% was observed when the pure *T. denitrificans* culture and the activated sludge enrichment were used, respectively. The use of *Pseudogulbenkiania* strain 2002 resulted in 50% of Fe(III) oxides in the precipitate. An eye-catching color method was also used for

an easy and fast identification of the iron oxides (Cornell and Schwertmann, 2003). However,
Fe(III) precipitates consisted of a mixture of (hydr)oxides with many impurities and the color
identification was, thus, not reliable. When transformed into a powder, the precipitates showed a
different color, which is reported in Table 1.

249 When the activated sludge enrichment was used, the liquid phase turned into a reddish and 250 brownish color throughout the experiments (Fig. S1). In the bottles with the *Thiobacillus*-mixed 251 culture, the color was more red than brown. The pure T. denitrificans and Pseudogulbenkiania strain 252 2002 resulted in a darker reddish brown or even black color (Fig. S1). Other studies observed a 253 white fluffy structure that turned into a greenish/gray or reddish/orange substance within a week 254 after the start of the incubation (Chaudhuri et al., 2001; Miot et al., 2009; Pantke at al., 2012). Green 255 rusts as major products cause a grayish/green appearance (Straub et al., 1996) and Fe(III) 256 oxyhydroxide or ferrihydrite generally give an orange/brown color to the liquid phase (Lack et al., 257 2002), as also observed in this study.

258

#### 3.1.1. Thiobacillus-mixed (TM) culture

The precipitates of the experiments carried out with the *Thiobacillus*-dominated mixed culture were a mixture of hematite, akaganeite and/or ferrihydrite. The infrared spectrum of the precipitates with the *Thiobacillus*-mixed culture showed strong bands at 540 and 3380 cm<sup>-1</sup> (Fig. S2A, Table 1), which are typical of hematite (Cornell and Schwertmann, 2003). Hematite was also detected with Raman spectroscopy and XRD. In the dark spots having a diameter of approximately 2.5  $\mu$ m, the combination of bands at 225, 290-300 and 412 cm<sup>-1</sup> was observed (Table 1, Fig. S3A), as well as a typical peak at around 24°20 in the XRD spectrum (Fig. 5A).

266 Strong bands at 1630 and 1050 cm<sup>-1</sup> in the infrared spectrum (S2A, Table 1) most likely 267 corresponded to akaganeite, as confirmed by the two broad bands at 700 and 400 cm<sup>-1</sup> obtained for 268 the red spots in Raman spectra (diameter 3 µm) (Fig. S3B) (Boucherit et al., 1989). However, the 269 band at around 700 cm<sup>-1</sup> might also have been due to ferrihydrite. In six-line ferrihydrite, three bands 270 at 370, 510 and 710 cm<sup>-1</sup>, while in two-line ferrihydrite only the strong 710 cm<sup>-1</sup> band can be seen 271 (Mazzetti and Thistlethwaite, 2002). However, none of the FTIR bands of ferrihydrite at 3615, 3430, 272 650, 450 cm<sup>-1</sup> was observed in the infrared spectrum, indicating that the red spots probably 273 corresponded to akaganeite. No XRD peak of akaganeite at 11.9°20 was observed, presumably due to overlapping with the peak of vivianite at  $11.2^{\circ}2\theta$  (Cornell and Schwertmann, 2003), in agreement 274 275 with SEM observations that showed the presence of vivianite.

The formation of akaganeite was likely stimulated by the presence of chloride ions. The structural network of akaganeite consists of tunnels with large cavities surrounded by double rows of edge-shared octahedral with Fe<sup>3+</sup>, Cl<sup>-</sup> or H<sub>2</sub>O molecules. The structure has two double chains that are corner-shared to neighboring double chains, forming a 3D network that differs from all other Fe(III) hydr(oxides) (body-centered cubic array) and stabilizes the tunnel structure of β-FeOOH crystals (Cornell and Schwertmann, 2003; dos Santos et al., 2016).

282 **3.1.2.** *T. denitrificans* (TDP) pure culture

283 When the pure culture of T. denitrificans was used, maghemite was observed on top of 284 hematite and/or ferrihydrite (Table 1). Raman spectra of the red spots (diameter ~5 µm) identified bands at 350, 700, 1300 and 1600 cm<sup>-1</sup> that corresponded to maghemite (Fig. S3D, Table 1) 285 (Hanesch, 2009). No other maghemite bands at 512 and 664 cm<sup>-1</sup> were identified, in contrast to what 286 287 was reported by Jacintho et al. (2007). The FTIR, Raman and XRD analyses confirmed also the 288 presence of hematite. Raman spectra of the dark spots showed bands at 225, 290-300, 412 and, 289 additionally, 611 cm<sup>-1</sup>, all corresponding to hematite (Fig. S3C). The band at 700 cm<sup>-1</sup> could also be ascribed to the presence of two-line ferrihydrite. The additional FTIR bands at 1400 and 2360 cm<sup>-1</sup> 290

were impurities (Fig. S2D), probably due to the presence of nitrate (Cornell and Schwertmann,
2003).

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### 3.1.3. Activated sludge (AS) enrichment

The activated sludge enrichment allowed the production of hematite and maghemite and/or
magnetite (and/or ferrihydrite). The FTIR spectrum showed bands as TDP and TM cultures (Fig.
S2B), and Raman spectra as well as XRD (24°20) confirmed the presence of hematite (Table 1, Fig.
5B). Similarly, Sun et al. (2009) reported the formation of a mixture of Fe(III) oxides dominated by
crystalline hematite with the use of a sludge inoculum in a lab-scale As(III)-oxidizing denitrifying
bioreactor.

The Raman bands at 1300 and 1600 cm<sup>-1</sup> in dark and red spots likely corresponded to 300 301 maghemite and/or magnetite (Fig. 2C). The Raman spectrum of white spots showed bands at 350, 302 500, 700 and 1350 cm<sup>-1</sup>, very similar to the maghemite bands at 350, 512, 664, 726 and 1330 cm<sup>-1</sup> 303 reported by Hanesch (2009). However, the FTIR bands at 3740 and 3725 cm<sup>-1</sup> corresponding to the 304 OH groups of maghemite were not recorded, indicating that the spectrum was likely a mixture of 305 magnetite and maghemite (Busca et al., 1993). De Faria et al. (1997) reported three Raman broad 306 peaks at 350, 500 and 700 cm<sup>-1</sup> in the maghemite spectrum. The Raman broad band from 653 to 750 307 cm<sup>-1</sup> that can be observed in all samples can be caused by magnetite and maghemite (Hanesch et al., 308 2009). Raman bands of magnetite were observed at 309 and 540 cm<sup>-1</sup>, whereas maghemite peaks at 344 and 490 cm<sup>-1</sup> (Table 1). The Raman band above 1300 cm<sup>-1</sup> of maghemite was relatively weak 309 310 compared to the other bands, probably due to the presence of magnetite. The red spots might also 311 have contained six-line or two-line ferrihydrite, corresponding to the three bands observed at 350, 312 500 and 700 cm<sup>-1</sup>.

313 **3.1.4.** *Pseudogulbenkiania* (PG) strain 2002

314 The use of the pure culture of Pseudogulbenkiania strain 2002 resulted in hematite and 315 maghemite formation (and/or ferrihydrite). FTIR spectra showed the bands of hematite (Fig. S2C), such as the Raman spectra with the additional band at 1600 cm<sup>-1</sup> corresponding to maghemite (Fig. 316 317 2B, Table 1). Zhao et al. (2013) reported that *Pseudogulbenkiania* strain 2002 allowed the formation 318 of vivianite, ferrihydrite and magnetite. Xiu et al. (2016) used *Pseudogulbenkiania* strain 2002 and 319 reported that lepidocrocite was the only crystalline biogenic solid formed. In this study, only 320 amorphous phases were reported by XRD analyses and similar to those obtained with the activated 321 sludge enrichment (Fig. 5B, Table 1). Crystalline phases were observed by SEM, although it was not 322 possible to distinguish between the different hydr(oxides) (Fig. 2C).

#### 323 **3.2.** Precipitates in abiotic controls

Chemical Fe(II) oxidation mainly resulted in the formation of hematite and another mineral between goethite or akaganeite. Also, a well crystallized vivianite ( $Fe_3(PO_4)_2 \cdot 8H_2O$ ) in the typical monoclinic crystal system and siderite ( $FeCO_3$ ) were easily identified by EDX (Fig. 3A and 4A). White, grey and reddish precipitates were observed in abiotic controls in the first 24 hours. In agreement, white precipitates from chemical Fe(II) oxidation in both biotic and abiotic experiments were reported elsewhere (Miot et al., 2009).

The infrared spectrum of the precipitates in the abiotic controls (AC) showed strong bands at 470 and 3380 cm<sup>-1</sup>, which are typical of hematite (Fig. 6). The band at 3380 cm<sup>-1</sup> most likely corresponded to the hydroxyl groups in the hematite crystals (Rochester and Topham, 1979). The presence of hematite was also confirmed by the reddish color of the AC sample (Table 1).

The obtained particles were aggregates of larger globular shaped particles than those
 obtained in experiments with microorganisms (Fig. 3). Dark and red spots (diameter ~10 µm) were
 observed with the microscope equipped in the Raman spectrometer. The Raman spectra of the dark

spots showed bands at 225, 290-300, 412 and 611 cm<sup>-1</sup>, which permitted to identify hematite (Table
1) (Oh et al., 1998) as also confirmed by FTIR analysis (Fig. 6). Furthermore, a band at 1321 cm<sup>-1</sup>,
corresponding to hematite, was also observed due to the presence of red spots.

340 The FTIR band at 1630 cm<sup>-1</sup> was probably related to the position of H<sub>2</sub>O bending vibrations of goethite or the presence of the typical O-H bending of akaganeite, whereas the band at 1050 cm<sup>-1</sup> 341 342 could be ascribed to a significant amount of adsorbed impurities in goethite particles (Gotic and 343 Music, 2007) (Table 1). Chemical Fe(II) oxidation results in the formation of goethite (Park et al., 344 2014), also in the presence of high concentrations of HCO<sub>3</sub><sup>-</sup> (Larese-Casanova et al., 2010). 345 However, the band at 1630 cm<sup>-1</sup> can also be related to the formation of akaganeite, and the additional band observed at 1050 cm<sup>-1</sup> has been reported in akaganeite samples in KBr discs, which 346 347 were considered to be artefacts (Murad and Bishop, 2000). However, no goethite or akaganeite 348 bands were observed in Raman spectra.

### 349 **3.3.** Mineralogy of the precipitates in the presence of metals

350 Copper, nickel and zinc have been reported to be essential for the structure of several 351 precipitates, such as akaganeite and maghemite (Deliyanni et al. 2007; Drofenik et al., 2008; Nador 352 et al, 2013; dos Santos et al., 2016; Hu et al., 2016; Suzuki et al., 2017). However, no previous 353 studies have investigated the effects of these metals on the formation of biogenic minerals by 354 nitrate-dependent Fe(II)-oxidizing microorganisms. In this study, all the precipitates contained Cu, 355 Ni, Zn, indicating metal precipitation, co-precipitation and sorption onto iron (hydr)oxides (Table 356 2). This is in agreement with the decrease of the soluble Cu, Ni and Zn concentration and the prediction of Visual MINTEQ reported by Kiskira et al. (2018). 357

The addition of Cu resulted in a red to brown liquid phase and a browner liquid phase was observed with the addition of Ni and Zn. The elemental identification and quantitative

360 compositional analysis performed by XRF and SEM-EDX of the precipitates obtained with the three 361 heavy metals are shown in Table 2. When Cu was supplemented, a higher amount of Fe(III) oxides 362 by 20 and 15% was obtained with the pure cultures of T. denitrificans and Pseudogulbenkiania 363 strain 2002, respectively, compared to the precipitates obtained from the bottles without metals. In 364 contrast, a decrease by 14 and 13% was observed in the samples with the *Thiobacillus*-mixed culture 365 and activated sludge enrichment (Table 2). The addition of Ni resulted in an increase of the Fe(III) 366 oxides formation in the range of 5-20% with all the microbial cultures. The precipitates obtained in 367 the presence of Zn had a lower amount of Fe oxides by 6 and 2% with the Thiobacillus-mixed 368 culture and the *Pseudogulbenkiania* strain 2002, respectively, whereas an increase by 22 and 9% 369 was observed in bottles with activated sludge and the *T. denitrificans* pure culture (Table 2).

370 The addition of metal salts such as Ni, Cu and Zn likely prevented the crystallization of 371 vivianite, thus determining the prevalent formation of amorphous phases (Fig. 5C, D). Hematite and 372 K<sub>2</sub>CaP<sub>2</sub>O<sub>7</sub> have been detected in ASC and TMC samples by XRD, respectively (Table 1). With all 373 the techniques used, the presence of Cu, Ni and Zn resulted in similar Fe(III) (hydr)oxide 374 precipitates as those obtained in the absence of metals, probably due to their low concentrations (Table 1). No significant differences by SEM were observed with the addition of metals (Cu, Ni, 375 376 Zn), except the formation of covellite (CuS) in the experiments with the activated sludge enrichment 377 and the addition of Cu (Fig. 4E). Traces of sulfate were present in the basal medium, and covellite 378 was formed likely due to the sulfidogenic activity of sulfate reducing bacteria present in the activated sludge enrichment (Gramp et al., 2006; Villa-Gomez et al., 2012; Papirio et al., 2013). No 379 380 significant differences were observed between the PG and PGC samples with *Pseudogulbenkiania* 381 sp. 2002.

382 Using all the microbial cultures, the mineralogy of the precipitates was not considerably 383 affected by the presence of Cu (Table 1). The precipitates obtained with the Thiobacillus-mixed 384 culture in the presence of Cu (TMC) were most likely a mixture of akaganeite and hematite, as also 385 observed in the TM precipitates. The addition of Cu resulted in a broad band at 3448 cm<sup>-1</sup>, likely 386 corresponding to the 3480+3390 (doublet) (OH and  $H_2O$  stretch) of akaganeite, as also confirmed by 387 the additional band at around 500 cm<sup>-1</sup> that was due to the Fe-O stretch of akaganeite (Table 1). The 388 strong band at 3448 cm<sup>-1</sup> might also be ascribed to the surface hydroxyl groups of goethite (Cornell 389 and Schwertmann, 2003). However, the intense band of hydroxyl stretch of goethite at 3140 cm<sup>-1</sup> and the typical band at 3666 cm<sup>-1</sup> were not observed, most likely due to the presence of phosphate in 390 391 the medium as reported by Russell et al. (1979).

392 When the T. denitrificans pure culture was used, another band of hematite at 470 instead of 393 540 cm<sup>-1</sup> was recorded in the TDPC sample (Table 1). The Raman bands at 225, 412 and 1321 cm<sup>-1</sup> 394 corresponded to hematite and an additional band at 700 cm<sup>-1</sup> probably indicated the presence of twoline ferrihydrite. The bands at 700 and 1330 cm<sup>-1</sup> were likely due to maghemite, in agreement with 395 396 what was observed in the TDP sample. Similarly, the ASC sample collected from the bottle seeded 397 with the activated sludge enrichment in the presence of Cu showed the same FTIR bands as for 398 TMC. Hematite, maghemite and/or magnetite and/or two-line ferrihydrite were identified from the 399 Raman spectra (Table 1). However, SEM micrographs of samples from experiments using the 400 activated sludge enrichment with addition of Cu mostly showed amorphous phases (Fig. 4D).

401 The addition of Ni led to more significant differences in the mineralogy of the precipitates 402 obtained with the four microbial cultures. The infrared spectrum of the precipitate obtained with the 403 *Thiobacillus*-mixed culture in the presence of Ni (TMN) showed some additional FTIR peaks at 404 3430, 940, 650 and 450 cm<sup>-1</sup>, corresponding to bulk OH deformations and stretch of ferrihydrite

405	(Table 1). The band revealed at 940 cm <sup>-1</sup> has already been reported when ferrihydrite contains a few
406	percent of Si (Schwetmann and Thalmann, 1976). In this study, the presence of Si was confirmed by
407	the XRF analysis (Table 2). A strong band at 1630 cm <sup>-1</sup> likely indicated the presence of akaganeite.
408	Hematite was revealed by the Raman spectrum of the dark spots (Fig. 3). The additional broad band
409	of maghemite at about 1300 cm <sup>-1</sup> was observed, indicating that the addition of Ni might stimulate
410	the formation of maghemite (Table 1), also in the T. denitrificans (TDPN), activated sludge
411	enrichment (ASN) and Pseudogulbenkiania strain 2002 (PGN) samples. In contrast to other studies
412	on denitrification in the presence of Ni (Zou et al., 2015), no specific Ni precipitate, such as
413	Ni <sub>3</sub> (PO <sub>4</sub> ) <sub>2</sub> , was observed with XRD likely due to the lower Ni concentrations used.
414	The FTIR spectrum of the precipitates taken from the bottles inoculated with Thiobacillus-
415	mixed culture in the presence of Zn (TMZ) showed a mixture of ferrihydrite and akaganeite (Table
416	1). The FTIR band at 3430 cm <sup>-1</sup> and the broad band at 940 cm <sup>-1</sup> of ferrihydrite were observed, as
417	well as the bands at 1630, 410 and 1050 cm <sup>-1</sup> corresponding to akaganeite (Table 1). An additional
418	band at 1384 cm <sup>-1</sup> might be associated with nitrate in ferrihydrite (Cornell and Schwertmann, 2003).
419	Differently to other samples, orange spots were observed by Raman spectroscopy and likely
420	corresponded to hematite, as hematite peaks at 225, 245, 411 and 1321 cm <sup>-1</sup> were visible. The bands
421	of TDPZ, ASZ and PGZ are reported in Table 1. The FTIR and Raman spectra were similar to those
422	obtained with the addition of Ni for all microbial cultures, indicating that also the presence of Zn
423	likely stimulated the formation of maghemite. XRD analysis of ASZ indicated the possible presence
424	of maghemite with the peak at $35^{\circ}2\theta$ .

**3.4. Thermal post-treatment** 

426 The goal of the thermal post-treatment of mineral phases is of double interest. Firstly, it427 allows to evaluate the weight loss due to desorption and decomposition phenomena of the phases

428 present in the products. Strong endothermic effects are mainly related to the water loss, due to the 429 adsorption and crystallization as well as the reaction of decomposition of hydroxides, oxy-430 hydroxides and nitrate among others. Secondly, it is possible to assess the potential for 431 crystallization of the amorphous phases.

To this end, the precipitates were thermally treated at  $650^{\circ}$ C for 2 h and the products were characterized by XRD analysis (Table 1) and weight loss (Eq. 1) due to calcination. The thermal treatment improved the crystallinity of hematite and favored the formation of some other crystalline phases with P in abiotic controls (Table 1). Similarly, the thermal treatment favored the formation of hematite in all samples (Fig. 7) and crystalline phases of Fe with P and Na, due to their high P content, i.e. FePO<sub>4</sub>. Fe<sub>2</sub>(III)Fe(II)(P<sub>2</sub>O<sub>7</sub>)<sub>2</sub> and NaFePO<sub>4</sub>, in TM and AS samples (Table 1, Fig. 7).

438 According to the chemical analysis, the NaFePO<sub>4</sub> phase appears to crystallize when the 439 amount of Na is higher than Cl. In this case and at low P content, the crystalline NaFeO<sub>2</sub> was 440 detected in the ASC sample (Table 1). When the amount of Na was lower than Ca and the P content 441 was relatively high, the crystalline  $Fe_2(III)Fe(II)(P_2O_7)_2$  was formed (AS and AC samples).

442 The weight loss for some samples was also determined (Table 3) and ranged between 24.4 443 and 61.9%. The thermal treatment allows the generation of samples with a different volumetric bulk. 444 Samples with a higher weight loss have a higher volumetric bulk compared with those having a 445 lower volume. In a comparison between the samples obtained with the different Fe(II)-oxidizing 446 denitrifying microorganisms, the use of Pseudogulbenkiania strain 2002 (PG, PGC, PGN and PGZ 447 samples) resulted in a higher weight loss due to the higher amount of biomass and higher amount of 448 water stored, likely indicating that the post-treatment requires a higher energy consumption 449 necessary for iron recovery.

#### 451 **4. Conclusions**

452 Fe(II)-mediated autotrophic denitrification resulted in the formation of a mixture of 453 amorphous phase or poorly crystalline Fe(III) (hydr)oxides (e.g. ferrihydrite), but also in some 454 crystalline phases such as hematite, akaganeite and maghemite. The use of the Thiobacillus-455 dominated mixed culture mainly led to the formation of akaganeite. Maghemite was more 456 intensively observed when the activated sludge enrichment and the two pure cultures of T. 457 denitrificans and Pseudogulbenkiania strain 2002 were used. The thermal post-treatment showed that the experiments generated iron oxide precipitates with different volumetric bulk. The use of 458 459 *Pseudogulbenkiania* strain 2002 resulted in an almost double weight loss of the precipitates, 460 inducing a higher energy consumption necessary for iron recovery. Cu, Ni and Zn were identified in 461 the precipitates by chemical analyses, indicating that metal precipitation, co-precipitation or sorption onto iron (hydr)oxides occurred. However, no difference in the mineralogy of the iron precipitates 462 463 was observed with the addition of Cu, whereas the addition of Ni and Zn stimulated the formation of 464 maghemite.

465

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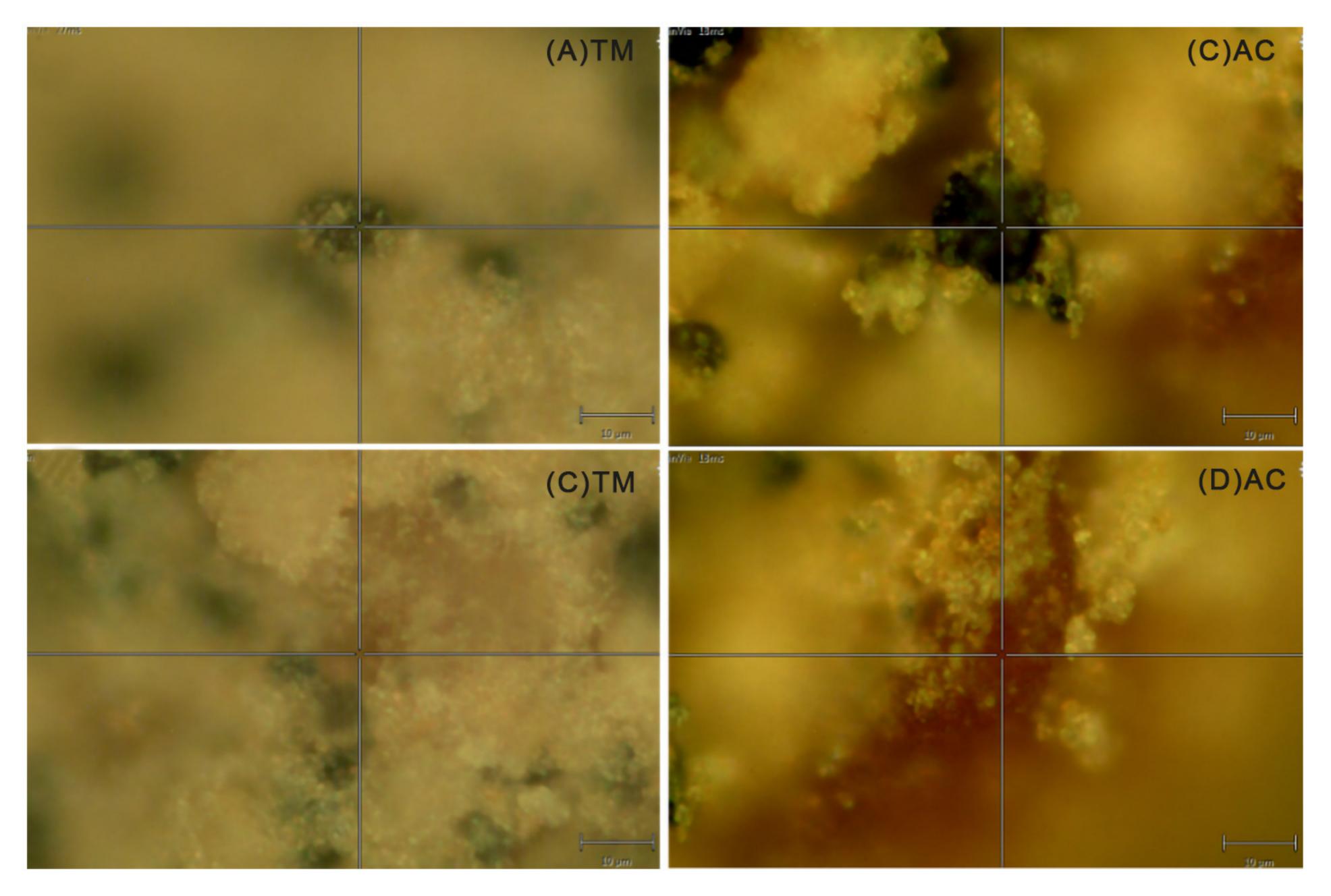
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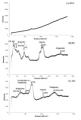
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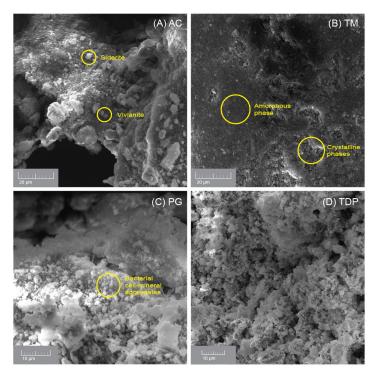
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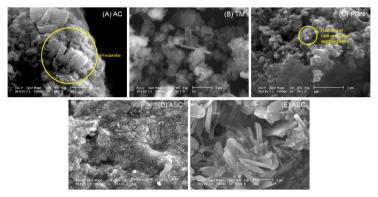
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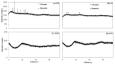
- 700 Fig. 1: Raman microscopy of the dark and red spots in the precipitates obtained from bottles
- inoculated with the (A, B) *Thiobacillus*-mixed culture (TM) and (C, D) from abiotic controls.
- Fig. 2: Raman spectra of the samples obtained during the experiments with (A) *Pseudogulbenkiania*
- strain 2002 in the presence of Ni (PGN), (B) Pseudogulbenkiania strain 2002 (PG) and (C) activated
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- Fig. 3: SEM micrographs without coating of samples were obtained from (A) abiotic controls, and
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- 707 strain 2002 and (D) T. denitrificans pure culture.
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- 711 Fig. 5: Mostly amorphous phases with a little amount of crystalline phases in XRD powder patterns
- of TM (A), AS (B) and TMN (C) and ASN (D) samples (JPCDS card no.30-662)
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- 715 Fig. 7: XRD powder patterns of (A) PGC, (B) PGZ, (C) TM and (D) AS samples calcined at 650°C
- 716 for 2 h (JPCDS card no.30-662).

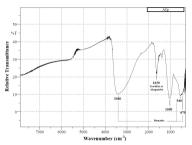


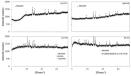












**Table 1:** Color, FTIR/Raman spectroscopy bands and XRD phases of the precipitates obtained during abiotic controls (AC), with different microbial cultures, i.e. *Thiobacillus*-mixed culture (TM), *T. denitrificans* pure culture (TDP), activated sludge (AS) enrichment, and *Pseudogulbenkiania* sp. 2002 (PG), both with and without the addition of Cu, Ni and Zn.

Sample	Color	FTIR, bands (cm <sup>-1</sup> )	Raman spectroscopy bands (cm <sup>-1</sup> )	Minerals (by FTIR/Raman Spectroscopy)	XRD phases	XRD phases at 650 °C
AC	Reddish	470, 540, 1050, 1630, 3380	Dark/Red: 225, 290-300, 412, 470, 611, 1000, 1321	Hematite (Goethite or Akaganeite)	Amorphous phases, Vivianite	Amorphous phases, Hematite, Fe <sub>2</sub> (III)Fe(II)(P <sub>2</sub> O <sub>7</sub> ) <sub>2</sub> (1-89- 876)
TM	Red to brown	540, 1050, 1400, 1630, 3380	Dark: 225, 290-300, 412,1000 Red: 400, 700	Hematite and Akaganeite (and/or Ferrihydrite)	Amorphous phases, Vivianite	Amorphous phases, Hematite, FePO4, NaFePO4
TDP	Brownish- reddish, yellow	540, 1050, 1400, 1630, 3380	Dark: 225, 290-300, 412, 611,1000 Red: 350, 700, 1300, 1600	Hematite and Maghemite and/or 2-line Ferrihydrite	Amorphous phases, Vivianite	Amorphous phases, Hematite, FePO4, NaFePO4
AS	Dark reddish brown	540, 1050, 1630, 3380	Dark1:225, 290-300, 412, 611 Dark2: 225, 290-300, 412, 700, 1000, 1300, 1600 Red: 350, 700, 1300, 1600 White: 350, 500, 700, 1350	Hematite, Maghemite and/or magnetite (and/or Ferrihydrite)	Amorphous phases, Vivianite, hematite	Amorphous phases, Hematite, Fe2(III)Fe(II)(P2O7)2 (1-80- 2315))
PG	Brownish- reddish,	540, 1050, 1630, 3380	Dark: 225, 290-300, 470, 1000, 1600	Hematite and maghemite	Amorphous phases,	-

	yellow				Vivianite	
TMC	Red to brown (yellow)	470, 500, 540, 1050, 1400, 1630, 3448	Dark: 225, 290-300, 412, 1000 Red: 400, 700	Hematite and Akaganeite	Amorphous phases, K2CaP2O7	-
TDPC	Brownish- reddish	470, 1050, 1400, 1630, 3380	Dark/red: 225, 400, 700, 1000, 1300	Hematite and Maghemite and/or 2-line Ferrihydrite	Amorphous phases	-
ASC	Brownish- reddish	470, 500, 1050, 1400, 1630, 3448	Dark1:225, 290-300, 412, 611 Dark2: 225, 290-300, 412, 700, 1000, 1300, 1600 Red: 400, 700, 1300, 1600 White: 400, 500, 700, 1350	Hematite, Maghemite and/or magnetite (and/or Ferrihydrite)	Amorphous phases, Hematite, K2CaP2O7	Amorphous phases, Hematite, NaFeO <sub>2</sub> (1-76- 600), Ca <sub>3</sub> Cl <sub>3</sub> P (1-84-1125)
PGC	Black	470, 1050, 1400, 1630, 3380	Dark: 400, 700, 1000, 1300	Hematite and maghemite	Amorphous phases	Amorphous phases, Hematite
TMN	Red to brown	450, 650, 940, 1630, 3430	Dark, red: 400, 700, 1000, 1321	Hematite Akaganeite/Ferr ihydrite, and (possible Maghemite)	Amorphous phases	-
TDPN	Red to brown	450, 1050, 1400, 1630, 3380	Dark, red: 400, 700, 1000, 1321 Dark: Additional 1600	Hematite and Maghemite and/or 2-line Ferrihydrite	Amorphous phases	-

ASN	Red to brown	450, 1050, 1400, 1630, 3380	White: 400, 700, 1000, 1321, 1600	Hematite, Maghemite, (and/or Ferrihydrite)	Amorphous phases	-
PGN	Red to brown	450, 1050, 1400, 1630, 3430	Dark, red: 400, 700, 1000, 1321	Hematite and maghemite	Amorphous phases	Amorphous phases, Hematite
TMZ	Red to brown	410, 940, 1630, 3440	Dark, red: 225, 245, 411, 700, 1000, 1300	Hematite Akaganeite/Ferr ihydrite, and (possible Maghemite)	Amorphous phases	-
TDPZ	Red to brown	450, 1050, 1400, 1630, 3380	Dark, red: 400, 700, 1000, 1321 Red/White: Additional 1600 (Maghemite)	Akaganeite and Hematite (possible Maghemite)	Amorphous phases	-
ASZ	Red to brown	450, 940, 1050, 1400, 1630, 3380	Dark, red: 400, 700, 1000, 1321 Red/White: Additional 1600 (Maghemite)	Hematite, Maghemite and/or magnetite (and/or Ferrihydrite)	Amorphous phases	-
PGZ	Black	450, 940, 1050, 1400, 1630, 3380	Dark, red: 225, 290-300, 412, 611, 700, 1000, 1321 Red/White: Additional 1600	Hematite and maghemite	Amorphous phases	Amorphous phases, Hematite

	Total Fe						<b>F</b> 1			(61)					
	oxides in the sample					]	Elemen	tal comj	position	.(%)					
Sample	Fe <sub>2</sub> O <sub>3</sub> (%)	Fe	Р	Cl	K	Na	Mg	Si	Al	Ca	Mn	S	Cu	Ni	Zn
AC	61	54.40	8.14	1.27	0.56	1.40	0.19	0.14	0.13	0.13	0.12				
ТМ	61	56.69	10.97	0.69	0.51	1.60	0.43	0.15	0.05	0.25	0.05	0.15			
TDP	57	55.82	10.19	0.42	0.81	1.22	0.20	0.00		0.10	0.05				
AS	55	53.21	9.53	2.04	0.78	1.60	0.69	0.81		1.48	0.12				
PG	50	43.28	7.61	4.39	0.76	2.20		0.15		0.20	0.05				
ТМС	47	41.16	8.09	4.95	1.09	5.66	0.25	1.17		0.19	0.14		0.40		
TDPC	77	70.79	1.31	1.61	0.10	0.00	0.42	1.97		0.10	0.08		0.78		
ASC	44	40.69	1.82	7.52	1.32	9.14	0.40			0.16	0.04		0.45		
PGC	65	58.14	0.80	1.43	0.10	1.10	0.30	1.00		0.08	0.11		0.54		
TMN	81	76.84	1.16	2.03	0.22	1.30	0.20	1.00		0.00				0.31	
TDPN	77	69.68	2.92	0.78	0.29	4.57	1.05			0.04	0.11			0.28	
ASN	73	66.21	0.86	0.71	0.09	0.00		1.30	0.10	0.08	0.10			0.25	
PGN	55	47.48	0.62	5.27	0.68	7.26		1.43		0.15				0.35	
TMZ	54	50.10	2.39	7.61	0.61	5.00	0.44	0.08		0.00		0.15			0.20
TDPZ	66	55.75	3.10	1.32	0.20	3.20	0.30	1.22		0.12	0.11				0.23
ASZ	77	66.68	1.50	1.54	0.15	5.02	0.30	1.23		0.07	0.13				0.19
PGZ	48	37.42	0.49	1.88	0.00			1.01							0.18

**Table 2:** Elemental identification and quantitative compositional analysis by XRF and SEM-EDX of the precipitates obtained in abiotic controls and in the batch bioassays operated with and without heavy metals (Ni, Cu and Zn).

Sample	АСр	PGp	ТМр	ASp	TDPp	PGCp	PGNp	PGZp	ТМСр	ASCp
wt %	25.1	49.5	25.1	31.5	24.4	57.5	57.4	61.9	30.1	26.5

**Table 3:** Weight loss of calcined samples after thermal treatment at 650°C for 2 hours.

