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Mapping the intensity function of a non-stationary point process in unobserved areas

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Abstract

Seismic networks provide data that are used as basis both for public safety decisions and for scientific research. Their configuration affects the data completeness, which in turn, critically affects several seismological scientific targets (e.g., earthquake prediction, seismic hazard...). In this context, a key aspect is how to map earthquakes density in seismogenic areas from censored data or even in areas that are not covered by the network. We propose to predict the spatial distribution of earthquakes from the knowledge of presence locations and geological relationships, taking into account any interactions between records. Namely, in a more general setting, we aim to estimate the intensity function of a point process, conditional to its censored realization, as in geostatistics for continuous processes. We define a predictor as the best linear unbiased combination of the observed point pattern. We show that the weight function associated to the predictor is the solution of a Fredholm equation of second kind. Both the kernel and the source term of the Fredholm equation are related to the first- and second-order characteristics of the point process through the intensity and the pair correlation function. Results are presented and illustrated on simulated non-stationary point processes and real data for mapping Greek Hellenic seismicity in a region with unreliable and incomplete records.

KEYWORDS: Conditional intensity, Earthquakes, Fredholm equation, Nonstationarity, Second-order characteristics, Spatial point processes.

1 Introduction

Mapping is a key issue in environmental science. A common and first example lies in ecology when mapping species distribution. When the location of individuals is known, we estimate the local density (usually by kernel smoothing), the so-called intensity in point process theory. However, point locations are usually unreachable at the survey scale, so that sampling methods are used; distance sampling or quadrat sampling approaches are possibly the most common ones. When no covariate is available, a global density estimation is then performed.

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However, species distribution characteristics vary spatially as they are governed by environmental data. Several approaches have been developed in that way for species data formed by reported presence locations, also called occurrence-only records (pure records of locations where a species occurred). Generally called Species Distribution Models (SDM), they aim to explain species occurrences from environmental variables. If they are used to gain ecological and evolutionary insights, they are also widely used for model-based interpolation across landscapes or to predict distributions to new geographic regions (Elith and Leathwick, 2009). SDMs are often based on multivariate statistical analysis methods, such as Generalized Linear/Additive Models (GLM/GAM). The most popular models are Maxent (Phillips et al., 2006) and Maxlike (Royle et al., 2012). In these models point locations are aggregated to grid cells and whether one or several individuals are observed in a cell, a one is recorded. Then, their aim is to estimate occurrence probability (species' probability of presence in a grid cell) maps. See e.g. Merow and Silander (2014) for a comparison of the two models and recommendations about their use.

Point process models offer a natural framework for species distribution modelling. Key concerns about SDMs lie in the loss of significant information about the spatial distribution when aggregating point locations, and the dependence of the results on the spatial resolution (Renner and Warton, 2013). Although point process models are connected to Maxent (Aarts et al., 2012; Renner and Warton, 2013), they use a continuous landscape rather than a discretized one, and the number of records is observed and comes from a random process rather than fixed (number of cells/quadrats). Renner et al. (2015) showed that using point process models presents many advantages, including some clarification about the response variable and model assumptions which, in addition, can be checked. Furthermore, because they operate at the individual level, point process models can incorporate interaction between individuals and dependence to environmental covariates.

Another concern about all these approaches is that they are Poisson model-based: their intrinsic definition does not account for relationship between individuals. These interactions nevertheless exist. Competition among individuals often leads to empty areas around each point, so-called exclusion by distance, mimicked by inhibition models. The American Redstarts, for example, compete with conspecifics for habitat in their winter grounds: the re-occupation of the empty areas supports the hypothesis that territoriality in this species acts to exclude conspecifics from certain winter habitat (Marra et al., 1993). On the contrary, individuals can be arranged by groups as with gregarious animals, such groups can also describe some local dispersion of the species around parents (as with plants). This arrangement is achieved in cluster models. The Shorea congestiflora is a dominant species in a 25–ha forest dynamics plot in a rain forest at Sinharaja (Sri Lanka), which apparently shows clustering at several scales (Wiegand et al., 2007). These effects can be mixed, with individuals arranged in groups, but at certain distance of each other inside each group. This can be the case for the spatial distribution of Northern Gannet (Chadœuf et al., 2011).

Similar issues arise in environmental science. The case we consider in this paper concerns seismic mapping in the Aegean Sea region. This is an area with high seismic density due to the presence of numerous volcanoes and plate movements. Earthquakes have a heterogeneous spatial distribution and we might be interested in understanding why certain regions are more favorable than others. Modeling earthquakes will therefore require to take into account geological information and interactions between events, clustering being often linked to aftershocks. Evaluating the seismic hazard requires a reliable monitoring network and sufficient coverage, what is rarely the case, and a lack of recording may not be due to the absence of an earthquake, but to an insufficient or unreliable network. The question then arises of mapping seismic activity in areas where the observation is unreliable or even when the area is not covered by the network, in order to access a relevant map where the seismicity is particularly high.

From a statistical point of view, the method developed in Gabriel et al. (2017) aimed to predict the local intensity of a spatial point process accounting for the individual relationships modeled by the pair correlation function (which is related to the probability to find a second point of the process at a given distance from a known point of the process). The main interest of this approach is that it estimates, using only first- and second-order characteristics of the point process, the local intensity outside the observation window, hereafter called prediction. The prediction of the local intensity is obtained conditionally to the records in the observation window. However, this method did not allow to consider environmental covariates and thus did not take into account potential spatial variations driven by these covariates at large scale, which may lead to unrealistic predictions in ecology and environment.

To fill this potential unwished situation, we propose to predict the spatial distribution of earthquakes from the knowledge of presence locations and geological relationships, taking into account any interactions between records. Namely, in a more general setting, we aim to estimate the intensity function of a point process, conditional to its censored realization, as in geostatistics for continuous processes. We define a predictor as the best linear unbiased combination of the observed point pattern, where the weight function associated to the predictor is the solution of a Fredholm equation of second kind related to the first- and second-order moments of the point process. We describe our approach in Section 2. We evaluate the goodness of our predictions through a simulation study (Section 3) for several cluster models. In Section 4, our methodology is applied to predict and map Hellenic seismicity in a region with unreliable and incomplete records.

2 Predicting the intensity conditionally to the observation

We consider in the following a spatial point process Φ in \mathbb{R}^2 , i.e. a random pattern of points for which both the number of points and their locations are random. Let us denote $\Phi(B)$ the number of points of Φ in any borelian set B, and Φ_B their locations in B. The intensity function, denoted by $\lambda(x)$, is defined as the function such that $\int_B \lambda(x) \, dx = \mathbb{E}[\Phi(B)]$, for Bany borelian set; this corresponds to the local probability to observe a point of Φ at a fixed location (if dx is an infinitesimal volume around location x, then $\mathbb{P}[\Phi(dx) = 1] = \lambda(x) \, dx$). The intensity function provides a trend in the spatial variation of points density, and we suppose that the intensity is driven by spatial covariates Z, such that $\lambda(x) = h(Z_x)$, where both h and Z are known.

The interaction between points is described through the pair correlation function g(x, y), which gives the extend at which the probability to find a point at a location y changes by the presence of a point of the process at location x (if dy is an infinitesimal volume around y, then $\mathbb{P}\left[\Phi(dy) = 1 \mid x \in \Phi\right] = g(x, y)\lambda(y) \, dy$). We assume that the process is second-order intensity reweighted stationary (Baddeley et al., 2000). This assumption means that its intensity varies in space, but the pair correlation function between two locations depends only on their difference vector. More generally, a point process can be described through another kind of intensity, the so-called (Papangelou) conditional intensity. This conditional intensity corresponds to the conditional probability of finding a point of the process inside an infinitesimal neighborhood of the location x, given the complete point pattern at all other locations. See Cœurjolly and Lavancier (2019) for a thorough presentation of the intensity and the conditional intensity, including theoretical definition, estimation methods and formal expression (when available) for the most popular point processes.

Here we consider $W \subset \mathbb{R}^2$, a window of interest, and we assume that Φ has only been observed in some observation window $W_{obs} \subset W$. The conditional distribution of $\Phi_{W \setminus W_{obs}}$ given $\Phi_{W_{obs}}$ can be expressed in terms of the conditional density (cf Cœurjolly et al. (2017)). Unfortunately, the density of Φ restricted to W_{obs} is barely tractable (and hard to handle) except for just some few processes, such as Poisson, Gibbs and determinantal processes; however, this is not the case for Cox and cluster point processes that are often used to model environmental or ecological point patterns.

Our aim is thus to estimate local intensity variations of the point process at any location $x_o \in W \setminus W_{obs}$ conditionally to h, Z and $\Phi_{W_{obs}}$, that is the Papangelou conditional intensity of Φ knowing $(\Phi_{W_{obs},h,Z})$. Following Gabriel et al. (2016, 2017), we refer to it as the spatial *local intensity* of Φ for sake of convenience. Given that we want to predict the local intensity at points $x_0 \in W \setminus W_{obs}$ using a linear predictor, it can be written as

$$\widehat{\lambda}(x_o|\Phi_{W_{obs}}) = \int_{W_{obs}} w(x;x_o) \sum_{y \in \Phi_{W_{obs}}} \delta(x-y) \,\mathrm{d}x = \sum_{x \in \Phi_{W_{obs}}} w(x;x_o),\tag{1}$$

where δ denotes the Dirac delta function.

Mimicking kriging, we want on one hand an unbiased predictor such that $\mathbb{E}[\hat{\lambda}(x_o|\Phi_{W_{obs}}) - \lambda(x_o|\Phi_{W_{obs}})] = 0$, which leads to

$$\int_{W_{obs}} \lambda(x) w(x; x_o) \, \mathrm{d}x - \mathbb{E} \left[\lambda(x_o | \Phi_{W_{obs}}) \right] = 0,$$

and therefore

$$\int_{W_{obs}} \lambda(x) w(x; x_o) \, \mathrm{d}x = \lambda(x_o). \tag{2}$$

On the other hand, the predictor must minimize the prediction error variance, which is equivalent to minimizing the following expression

$$\begin{aligned} \mathbb{V}\mathrm{ar}[\hat{\lambda}(x_o|\Phi_{W_{obs}}) - \lambda(x_o\Phi_{W_{obs}})] &= \int_{W_{obs}} \lambda(x)w^2(x;x_o)\,\mathrm{d}x \\ &+ \int_{W_{obs}} \int_{W_{obs}} \lambda(x)\lambda(y)w(x;x_o)w(y;x_o)(g(x-y)-1)\,\mathrm{d}x\,\mathrm{d}y \\ &- 2\lambda(x_o)\int_{W_{obs}} \lambda(x)w(x;x_o)(g(x_o-x)-1)\,\mathrm{d}x \end{aligned}$$

Using variational calculation and the Riesz representation Theorem similarly to (Gabriel et al., 2017), one gets the following Fredholm equation of the second kind

$$\lambda(x)w(x;x_o) + \lambda(x)\int_{W_{obs}}\lambda(y)w(y;x_o)(g(x-y)-1)\,\mathrm{d}y - \frac{\lambda(x)}{\int_{W_{obs}}\lambda(z)\,\mathrm{d}z} \left[\lambda(x_o) + \int_{W_{obs}}\int_{W_{obs}}\lambda(z)\lambda(y)w(y;x_o)(g(z-y)-1)\,\mathrm{d}z\,\mathrm{d}y\right] = \lambda(x_o)\lambda(x)(g(x_o-x)-1) - \frac{\lambda(x_o)\lambda(x)}{\int_{W_{obs}}\lambda(z)\,\mathrm{d}z}\int_{W_{obs}}\lambda(z)(g(x_o-z)-1)\,\mathrm{d}z.$$
(3)

Its solution is implicitbut it can be numerically approximated using the Galerkin approximation method (Kress, 2013). This projection method convert the Fredholm equation into a discrete problem where a Finite Element Space is used for the approximation space, see Gabriel et al. (2017).

3 Simulation study

The objective of this section is twofold: first, we want to visualize how the prediction is affected by both the distance from the prediction point x_o to the observed window W_{obs} and the point process structure, and second, we want to measure the variability between our estimator $\hat{\lambda}(x_o | \Phi_{W_{obs}})$ and the conditional intensity $\lambda(x_o | \Phi_{W_{obs}})$.

In order to see how the prediction estimator behaves with respect to the pattern structure we use the inhomogeneous Matérn cluster model which allows the computation of its conditional intensity. This model is obtained as follows. We first define a stationary Matérn cluster process Φ^{Mat} , i.e. a process where each point of a Poisson parent process is replaced by a poissonnian cluster of offspring uniformly distributed in a disc of radius R around the parent point. We then focus on the independent p(x)-thinned process, where p(x) is a deterministic function on \mathbb{R}^2 with $0 \le p(x) \le 1$. Every point x belonging to Φ^{Mat} is deleted with probability 1 - p(x), and again its deletion is independent of locations and possible deletions of any other points (Chiu et al., 2013).

Let Φ be the process of the thinned offspring and Φ^p the parent process (Poisson process with intensity κ). The process Φ is second-order intensity reweighted stationary with intensity $\lambda(x) = \kappa \mu p(x)$, with μ the expected number of offspring per parent. The pair correlation function is the one of Φ^{Mat} :

$$g(r) = 1 + \frac{2}{\kappa (\pi R)^2} \left(\arccos\left(\frac{r}{2R}\right) - \frac{r}{2R} \sqrt{1 - \frac{r^2}{4R^2}} \right), \text{ if } 0 < r < 2R, \text{ and } g(r) = 1 \text{ otherwise.}$$

According to Gabriel and Chadœuf (2021), the local intensity of the thinned Matérn cluster process Φ is

$$\lambda(x_o|\Phi_{W_{obs}}) = \frac{\mu p(x_o)}{\pi R^2} \int_{b(x_o,R) \cap (W_{obs} \cup \partial W)} \lambda^p(y|\Phi_{W_{obs}}) \,\mathrm{d}y + \kappa \mu p(x_o) \nu \left(b(x_o,R) \setminus (W_{obs} \cup \partial W)\right),\tag{4}$$

where ∂W stands for the outside border of thickness R of the observation window, $b(x_o, R)$ the disc of center x_o and radius R, and $\lambda^p(y|\Phi_{W_{obs}})$ is the conditional intensity of parents in $W_{obs} \cup \partial W$ given the realization of offspring in W_{obs} . This intensity is approximated by

$$\widehat{\lambda^p}(y|\Phi_{W_{obs}}) = \frac{1}{\mu p(y)\pi R^2} \sum_{x \in \Phi_{W_{obs}}} \mathbb{I}_{b(x,R)}(y) + \kappa \exp\left(-\frac{\mu}{\pi R^2} \int_{b(y,R) \cap W_{obs}} p(z) \, \mathrm{d}z\right),$$

where the first term is the empirical intensity of parents given the observed offspring, and the second term is the intensity of parents with unobserved offspring. See Gabriel and Chadœuf (2021) for the approximation of the conditional distribution of parent points given the offspring points and its validation. The inhomogeneous Matérn Cluster process (IMCP) Φ depends on four parameters: the thinning probability p(x), the intensity of parents κ , the mean number of points per parent μ , and the radius of dispersion of the offspring around the parent points R. In our simulation study we fix $\kappa = 50$ and $\mu = 40$, and we consider:

- two thinning probabilities: $p_1(x) = p_1(x_1, x_2) = \alpha_1 \mathbb{I}_{\{x_1 \le v\}} + \alpha_2 \mathbb{I}_{\{x_1 > v\}}$, setting $\alpha_1 = 0.8$, $\alpha_2 = 0.2$ and v = 0.5, and $p_2(x) = p_2(x_1, x_2) = 1 - x_1$.

- the unit square as study region W. The observation window is $W_{obs} = W \setminus W_{pred}$, where $W_{pred} = [0.35, 0.65]^2$ when using $p_1(x)$ and $W_{pred} = [0.05, 0.95] \times [0.36, 0.64]$ when using $p_2(x)$. - $R \in \{0.05, 0.09, 0.13\}$.

We perform n = 1000 simulations of Φ for all combinations of pairs (p(x), R). Each scenario is denoted IMCP(p(x), R). Realizations of IMCP $(p_1(x), R)$ (resp. IMCP $(p_2(x), R)$) are given in the first column of Figure 1 (resp. Figure 2) for the different values of R, illustrating inhomogeneous patterns with increasing range of clustering from top to bottom.



Figure 1: First column: inhomogeneous Matérn cluster point patterns, $\Phi_{W_{obs}}$, obtained from $p_1(x)$. Second column: approximation of the local intensity in W_{pred} (central square) and kernel smoothing of the pattern in W_{obs} . Third column: predicted local intensity in W_{pred} (central square) and kernel smoothing of the pattern in W_{obs} .

For each simulation, we compute the local intensity from both the predictor (1) and the approximation (4) in W_{pred} . We consider the same mesh for the Galerkin approximation method for all configurations, for which W_{obs} is subdivided in 15,194 triangles. The local intensity and the prediction are respectively plotted in the second and third columns of Figures 1 and 2. For a visualization purpose, we plotted the local intensity in W_{pred} (central square/rectangle), as well as a Gaussian kernel smoothing of the intensity in W_{obs} .



Figure 2: First column: inhomogeneous Matérn cluster point patterns, $\Phi_{W_{obs}}$, obtained from $p_2(x)$. Second column: approximation of the local intensity in W_{pred} (central rectangle) and kernel smoothing of the pattern in W_{obs} . Third column: predicted local intensity in W_{pred} (central rectangle) and kernel smoothing of the pattern in W_{obs} .

We can see that the method reproduces the structures of the point process. In particular, it reproduces clusters, as soon as there are points close enough to the boundary of W_{pred} , and the prediction tends to the intensity of the process, $\lambda(x) = \kappa \mu p(x)$, when it is made at distances larger than R from the boundary of W_{obs} . This is particularly clear for small values of R (first rows in Figures 1 and 2). Results from these realizations show, however, some weak differences between the local intensity and our prediction that we now quantify. We measure the precision and the variability of our prediction through the relative bias (RB) and the relative root mean square error (RRMSE), for $x_o \in W_{pred}$, as follows

$$\operatorname{RB}(x_o) = \frac{\sum_{i=1}^{n} \left(\widehat{\lambda}_i(x_o | \Phi_{W_{obs}}) - \lambda_i(x_o | \Phi_{W_{obs}}) \right)}{\sum_{i=1}^{n} \lambda_i(x_o | \Phi_{W_{obs}})},$$

and

$$\operatorname{RRMSE}(x_o) = \frac{\sqrt{\frac{1}{n} \sum_{i=1}^{n} \left(\widehat{\lambda}_i(x_o | \Phi_{W_{obs}}) - \lambda_i(x_o | \Phi_{W_{obs}})\right)^2}}{\frac{1}{n} \sum_{i=1}^{n} \lambda_i(x_o | \Phi_{W_{obs}})},$$

where $\widehat{\lambda}_i$ and λ_i stand for the prediction and the local intensity of the *i*-th simulation. Figure 3a and 3c illustrate the distribution of $\text{RB}(x_o)$ and $\text{RRMSE}(x_o)$ for all patterns IMCP(p(x), R).



Figure 3: Relative bias (top) and relative RMSE (bottom) computed from 1000 simulations with $p_1(x)$ and $p_2(x)$ within W_{pred} . Distances are measured outwards the boundary of W_{obs} , inwards W_{pred} .

In all cases, the relative bias is concentrated around 0 and is much less than 5% (in absolute value) what reveals an excellent precision of the predictor. The distributions of RRMSE(x_o) show a decreasing variability as the interaction radius R increases. Indeed, as shown in Figures 1 and 2, the inter-point distances of the point patterns are greater for larger values of R and the local intensity is thus more diffuse for large values of R, and more picky for very small values. The weight function is closely related to the pair correlation function. It shows a radial wavy behavior around x_o with positive decreasing values at distances less that R, and it has negative values between R and 2R, ..., and then is null (see Figure 4). As the predictor is the sum of the weight at $x \in \Phi_{W_{obs}}$, it reflects this behavior and clusters



Figure 4: Weight function $w(x; x_o)$ (here transformed as $sign(w(x; x_o)) log(|w(x; x_o)|)$) associated to the point x_o (black dot in W_{pred}) for the process IMCP $(p_1(x), 0.05)$. For each realization of the point process, $\Phi_{W_{obs}}$ (e.g. the dark blue dots), the local intensity at x_o is the sum over $x \in \Phi_{W_{obs}}$ of the weights $w(x; x_o)$.

are thus well marked. The approximation of the local intensity (4) is smoother as it is only expressed in terms of first-order moments. Hence, we get more variability in the border ∂W , and even higher for picky intensities as for R = 0.05. That is also shown in Figures 3b and 3d which illustrate the relative bias and relative RMSE computed at different distances from ∂W . They are both higher at very short distances.

4 Predicting the density of earthquakes

In this section we focus on the Greek-Hellenic area, a region of high seismic hazard due to both tectonic and volcanic seismogenic sources. This seismicity is a result of some motion-induced deformations: northward motion of the African lithosphere, westward motion of the Anatolia plate and collision between African and Eurasian plates (Papazachos and Comninakis, 1971; Le Pichôn and Angelier, 1979; McKenzie, 1972; Anderson and Jackson, 1987). A total of

1173 earthquakes of magnitude greater or equal than 4 occurred in the study area W (black square in Figure 5) between 2004 and 2015. Their locations are plotted in Figure 5.



Figure 5: Earthquakes of magnitude greater or equal than 4, faults, plate boundaries and main volcanoes in the Hellenic region. Prediction window is shown in the light yellow area.

Data have been recorded by the Hellenic Unified Seismological Network (HUSN). Seismic networks provide data that are used as the basis both for public safety decisions and for scientific research. They are essential tools for observing earthquakes and assessing seismic hazards that can be described and characterized to assess their degree of coverage (Siino et al., 2020). Indeed, their configuration affects the data completeness, which in turn, critically affects several seismological scientific targets (e.g., earthquake prediction, seismic hazard, etc. (Vamvakaris et al., 2013)). From different indicators (magnitude of completeness, Radius of Equivalent Sphere that estimates an average error of location), D'Alessando et al. (2011) identified some seismogenic areas that probably are not adequately covered by the HUSN. Based on their results we delineate a zone (in light yellow in Figure 5) with unreliable or missing records. We removed the 48 points located in this yellow zone to predict the local intensity of earthquakes.

Modelling earthquakes needs to account for geological information and interactions between earthquakes, usually in terms of clustering. The most popular models of seismological events are the so-called self-exciting, such as Hawkes and ETAS models (Adamopoulos, 1976; Ogata, 1988). Our approach, however, is model-free, it accounts for environmental heterogeneity and interaction between events. Following Siino et al. (2017), we assume that the intensity of earthquakes log-linearly depends on several spatial geological covariates: D_{pb} , the distance to the plate boundary (orange dashed curves in Figure 5); D_f , the distance to the nearest fault (blue lines in Figure 5); and D_v , the distance to the nearest main volcano (red triangles in Figure 5). Thus, the intensity takes the form

$$\lambda(x) = \lambda(x_1, x_2) = \exp\left(\beta_0 + \beta_1 x_1 + \beta_2 x_2 + \beta_3 x_1^2 + \beta_4 x_1 x_2 + \beta_5 x_2^2 + \beta_6 \mathbb{I}_{\left\{D_f(x) \le \phi_1\right\}} D_f(x) + \beta_7 \mathbb{I}_{\left\{\phi_1 < D_f(x) \le \phi_2\right\}} D_f(x) + \beta_8 \mathbb{I}_{\left\{\phi_2 < D_f(x) \le \phi_3\right\}} D_f(x) + \beta_9 \mathbb{I}_{\left\{\phi_3 < D_f(x) \le \phi_4\right\}} D_f(x) + \beta_{10} \mathbb{I}_{\left\{D_f(x) > \phi_4\right\}} D_f(x) + \beta_{11} D_v(x) + \beta_{12} D_{pb}(x)\right), \quad (5)$$

with $\phi_1 = 6.73$, $\phi_2 = 43.48$, $\phi_3 = 54.783$ and $\phi_4 = 112$. The estimated values of model parameters are given in Table 1, and the estimated intensity is depicted in Figure 6 (left panel, log-scaled). This figure shows that the region to be predicted (bordering in black) is across zones of strong and weak intensities. We then calculated the empirical pair correlation

Parameter	β_0	β_1	β_2	β_3	β_4	β_5	β_6
Estimate	-178.483	0.611	9.673	0.034	-0.061	-0.114	-0.034
Parameter	β_7	β_8	β_9	β_{10}	β_{11}	β_{12}	
Estimate	-0.0268	-0.0175	-0.0179	-0.0214	0.0040	-0.0093	

Table 1: Estimated parameters of the intensity model $\lambda(x)$.

function (see Figure 6 (right panel)) and fitted an exponential model of the form $g(r) = \exp(-\alpha_2\sqrt{r})/\alpha_1$, with $\hat{\alpha}_1 = 8.9502$ and $\hat{\alpha}_2 = 0.0266$ (as it is shown in Figure 6 (right panel), for the theoretical fitted model). The high values of the pair correlation function at short distances are indicative of a cluster process with a small range, less than 50 km. Then we have g(r) = 1 (horizontal dark grey line), indicating there is no particular spatial structure or interaction. Note that g(0) > 7.5 corresponds to a high probability of observing an earthquake within a small disc centered at an observed earthquake.



Figure 6: Log-intensity $\hat{\lambda}(x)$ (left panel) and pair correlation function (right panel): empirical (green dashed line), fitted (black solid line).

The prediction of the local intensity is obtained by using a mesh of 22,686 triangles for the Galerkin approximation method. The left panel of Figure 7 shows the predicted local intensity $\widehat{\lambda}(x_o|\Phi_{W_{obs}})$ in $W_{pred} = W \setminus W_{obs}$ and a Gaussian-kernel smoothing of observed earthquake locations in W_{obs} , with bandwidth 22 km. Both intensities are at a log-scale. The blue area in the center of W_{obs} indicates a kernel smoothing value near zero because of the absence of points in this region (see Figure 5). Compared with Figure 6, this plot also emphasizes the differences between the intensity fitted from covariates in Equation (5) and the empirical intensity obtained by kernel smoothing. This leads to conclude that we can expect some influence from point locations on the local intensity. This effect is observed in the middle panel of Figure 7. It focuses on the prediction window in which both the predicted local intensity and the fitted intensity of the point pattern $\widehat{\lambda}(x_{\alpha})$ are plotted at a log-scale. The right panel plots their ratio to highlight their differences. The different range values between $\lambda(x_o|\Phi_{W_{obs}})$ and $\lambda(x_o)$ indicate that if we did not take into account the effect of earthquakes locations in W_{obs} , we would get regular variations of the local intensity in W_{pred} . The knowledge of these points add local hot spots close at the border between W_{obs} and W_{pred} . Hence, the prediction window shows a high and spatially varying local intensity, whose variations reflect the scales of the structures of the underlying process, through both the intensity and the pair correlation function.



Figure 7: Prediction of the local intensity $\hat{\lambda}(x_o | \Phi_{W_{obs}})$ in W_{pred} , and a Gaussian kernel smoothing in W_{obs} (left panel). Prediction of the local intensity and fitted intensity $\hat{\lambda}(x_o)$ given in Equation (5) in W_{pred} (middle panel) and their ratio (right panel). All plots are at a log-scale.

Note that from a practical point of view, the approximation of the weight functions has been implemented in FreeFem++ (Hecht, 2012) a high level integrated development environment for numerically solving partial differential equations. All other computations are implemented in R (R Core Team, 2021). We used the R package spatstat (Baddeley and Turner, 2005) for fitting the intensity $\lambda(x)$ and for computing the empirical pair correlation function. Because interfacing R and FreeFem++ is not straightforward and not feasible on any operating system, codes are only available upon request to the first author.

5 Discussion

We have proposed a predictor of the local intensity of point processes conditionally to some censored observations, and accounting for the individual relationships and considering environmental covariates. In this general context we can distinguish two situations: in a first case, censoring is a continuous function defined over the study region W, so that observed and unobserved points share the same space; in a second case, observations are taken in some windows and we want to predict outside the observation window. In this paper we presented the latter case and considered that the spatial variation driving the point density is related to an intensity governed by environmental covariates.

When observed and unobserved points share the same space, the study region is the observation window and the observation is not exhaustive. In the example of earthquakes, we know the network reliability. In some sense it can be viewed as a probability of detection, say $\pi(x)$, which is independent of the underlying process of earthquakes. This process is the union of the observed process, Φ^{π} , and the unobserved process, $\Phi^{1-\pi}$. Then the local intensity is the sum of the intensity of the observed process and the intensity of the unobserved process given the observation, what can be predicted as previously. But now, weights also depend on the observation probability.

The interaction with other processes provides another form of spatial variation. We can easily imagine that the presence or the absence of a species depend on the presence/absence of other species. The extension of our approach to multi-type processes is an on-going work. In that case, we consider several processes, say $\Phi^{(k)}$, $k = 1, \ldots, K$, observed respectively in some windows W_1, \ldots, W_K . To predict the intensity of the first process given the realization of the others, we can define the predictor as follows

$$\widehat{\lambda_1}(x_o | \Phi_{W_1}^{(1)}, \dots, \Phi_{W_K}^{(K)}) = \sum_{k=1}^K \sum_{x \in \Phi_{W_k}^{(k)}} w_k(x; x_o).$$

The unbiasedness constraint on the predictor modifies the constraint on the weight function,

$$\sum_{k=1}^{K} \int_{W_k} \lambda_k(x) w_k(x; x_o) \, \mathrm{d}x = \lambda_1(x_o),$$

and minimizing the error prediction variance under this constraint leads to a new Fredholm equation which now depends on the pair correlation function of each process and on the cross-pair correlation functions.

A question remains: how to control the approximation accuracy when solving the Fredholm equation? There are two ways to increase the accuracy of the computed solution of the Fredholm equation: either refine the mesh, i.e. lower the size of the triangles, or use another approximation basis than the usual one (P1={piece-wise Linear functions}). Both options significantly increase the dimension of the resulting dense linear problem and from the point of view of computing efficiency we can wonder which is the most suitable approximation. Note that for a given mesh, we get a more precise representation of the regularity of the solution if we use a finite element basis involving more regular functions e.g. P2={piece-wise Quadratic functions} or P3={piece-wise Cubic functions}...However, changing the basis of approximation won't reduce the approximation errors made with respect to the true solution of the Fredholm equation, errors that are intrinsically linked to the typical size of a triangle of the mesh. Up to our knowledge, no rule has been published that relates the size of the mesh to the characteristics of the equation. A too crude size can even lead to negative estimated local intensities. We propose to verify empirically the approximation quality by simulating point realizations using an easy to manipulate point process with the same first- and second-order characteristics, as the log-Gaussian Cox point process, and check discrepancy between reestimated local intensities and the estimated one.

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