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Giovanni Bacco, Barbara Delmonte, Elena Di Stefano, Giannantonio Cibin, Ilaria Crotti, et al.. Deep ice as a geochemical reactor: Insights from iron speciation and mineralogy of dust in the Talos Dome ice core (East Antarctica). *The Cryosphere*, 2021, 15 (10), pp.4807-4822. 10.5194/tc-15-4807-2021 . hal-03434954

**HAL Id: hal-03434954**

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Submitted on 19 Nov 2021

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# Deep ice as a geochemical reactor: insights from iron speciation and mineralogy of dust in the Talos Dome ice core (East Antarctica)

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Received: 25 May 2021 – Discussion started: 29 June 2021

Revised: 15 September 2021 – Accepted: 21 September 2021 – Published: 12 October 2021

**Abstract.** Thanks to its insolubility, mineral dust is considered a stable proxy in polar ice cores. With this study we show that the Talos Dome ice core (TALDICE, Ross Sea sector of East Antarctica) displays evident and progressive signs of post-depositional processes affecting the mineral dust record below 1000 m deep. We apply a suite of established and cutting-edge techniques to investigate the properties of dust in TALDICE, ranging from concentration and grain size to elemental composition and Fe mineralogy. Results show that through acidic/oxidative weathering, the conditions of deep ice at Talos Dome promote the dissolution of specific minerals and the englacial formation of others, affecting primitive dust features. The expulsion of acidic atmospheric species from ice grains and their concentration in localized environments is likely the main process responsible for englacial reactions. Deep ice can be seen as a “geochemical reactor” capable of fostering complex reactions which involve both soluble and insoluble impurities. Fe-bearing minerals can efficiently help in exploring such transformations.

## 1 Introduction

Antarctic ice cores are a valuable archive which allows the reconstruction of the climatic history of the Earth during the last 800 000 years (Wolff et al., 2010). Mineral dust is one of the most extensively studied proxies in ice cores. Its importance stems from its role in the Earth's climate system: production, transport and deposition of dust are controlled by climate-related processes, but at the same time dust affects the climate (Maher et al., 2010). Studying the properties of dust trapped in ice cores, it is possible to obtain information on how the climate influences the dust cycle (Delmonte et al., 2004) and about the effects of dust under different climatic regimes (Mahowald et al., 1999; Wolff et al., 2006; Potenza et al., 2016).

It is known that the physical and compositional properties of dust trapped in ice cores are influenced by climatic, environmental and atmospheric processes (Sugden et al., 2009; Delmonte et al., 2017; Markle et al., 2018). The dust concentration in ice is strictly controlled by climate (Delmonte et al., 2002); dust grain size is related to its atmospheric transport (Delmonte et al., 2017; Albani et al., 2012a) and geochemistry to dust sources (Delmonte et al., 2004).

Recently, a growing number of studies have focused on the preservation and decay of climatic proxies in ice cores over time (Barnes et al., 2003; De Angelis et al., 2013; Ohno et al., 2016; Eichler et al., 2019; Bacco et al., 2021). After the incorporation in snow, the proxies are affected by alterations produced by post-depositional processes. Generally, the influence of such changes progressively increases with ice age and depth (De Angelis et al., 2013; Tison et al., 2015). Post-depositional processes in deep ice are related to three main causes: (1) the interaction between ice flow and bedrock (Goossens et al., 2016); (2) the metamorphism of ice, intended as the set of physical transformations to which ice is subject at increasing pressure and age (Faria et al., 2010); and (3) the diffusion of impurities (Barnes et al., 2003). An irregular bedrock can produce stratigraphic disturbances affecting the original ice succession for several hundred meters. On the contrary, ice metamorphism and englacial diffusion act at the ice-crystal scale (mm/cm scale) (Rempel et al., 2002; Faria et al., 2010). The effects of such phenomena on the proxies investigated in ice cores are not yet fully understood, but improving their comprehension is essential considering the “Oldest Ice” challenge (Fischer et al., 2013).

Dust is considered relatively immobile and stable in ice, and its concentration is used to synchronize deep ice cores when other proxies have deteriorated (Ruth et al., 2007; Kawamura et al., 2017). However, mineral particles may be altered by post-depositional changes. Impurities in ice are affected by small-scale relocation resulting from ice metamorphism, i.e., the re-crystallization and orientation of ice grains (Faria et al., 2010; Marath and Wettlaufer, 2020). With most of the impurities incompatible with respect to the ice lattice (Wolff, 1996), during the re-crystallization the impurities are expelled from the ice crystalline structure and accumulated at ice grain junctions or within intra-grain micro-inclusions (Mulvaney et al., 1988; De Angelis et al., 2005; Sakurai et al., 2017; Stoll et al., 2021). The accumulation of soluble and insoluble impurities forms eutectic mixtures whose pressure melting point is below the ice temperature, promoting the localized formation of liquid brines which lead to *in situ* chemical reactions (Fukazawa et al., 1998; De Angelis et al., 2005, 2013; Sakurai et al., 2017). Such small-scale environments are dominated by sulfur-rich acidic species, strongly affected by re-mobilization and concentration because of their high incompatibility with the ice lattice (Mulvaney et al., 1988; Wolff, 1996; Fukazawa et al., 1998). The interaction between acidic brines and concentrated impurities, including dust, leads to acid–base reactions (Traversi et al., 2009; Ohno et al., 2016). Considering dust, the most common reactions happening in ice are the dissolution of carbonates, the precipitation of gypsum (Ohno et al., 2006; Iizuka et al., 2008; Eichler et al., 2019) and of other uncommon sulfates (Ohno et al., 2014), and the englacial formation of secondary iron minerals (De Angelis et al., 2013; Bacco et al., 2021). Below 3000 m deep in the EPICA Dome C ice core, the re-precipitation of carbonates has also been reported, suggest-

ing that once acidity is consumed, additional reactions take place (Tison et al., 2015).

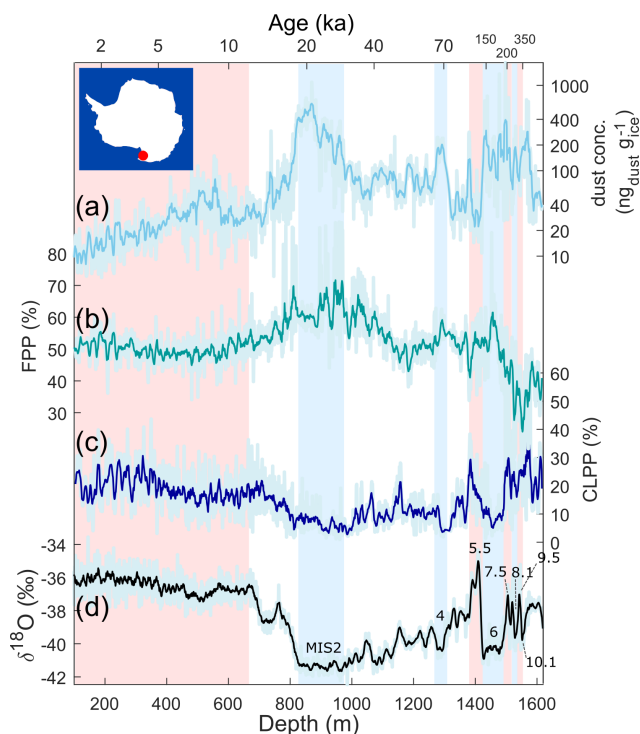
The iron fraction of dust is particularly sensitive to englacial transformations. De Angelis et al. (2013) have identified secondary Fe-bearing minerals in the deep part of the EPICA Dome C ice core, while Eichler et al. (2019) have detected a few particles in the EPICA Dronning Maud Land ice core with a Raman signature compatible with jarosite, an Fe–K sulfate forming from mineral weathering (Papike et al., 2006). This finding is confirmed by a thorough investigation of the Talos Dome ice core (TALDICE, East Antarctica), where jarosite is found below 1000 m deep and interpreted as the result of acidic, water-limited weathering of dust (Bacco et al., 2021). Since Fe biogeochemistry is strongly coupled with the global carbon cycle, this element receives considerable attention from the ice core community (Wolff et al., 2006; Conway et al., 2015; Hooper et al., 2019), and methods have been developed to measure its concentration and speciation (Spolaor et al., 2013; Burgay et al., 2019). Nevertheless, the effects of post-depositional processes in ice cores on its geochemistry are still poorly investigated.

This work presents and discusses a set of geochemical evidence showing that below 1000 m deep the dust record of TALDICE is affected by post-depositional transformations which alter its particle-size distribution, elemental composition and mineralogy. A detailed analysis of speciation and mineralogy of the Fe fraction of dust reveals that deep ice at Talos Dome acts as a geochemical reactor, favoring reactions and transformations which involve both soluble and insoluble impurities.

## 2 Materials and methods

### 2.1 The TALDICE ice core

TALDICE has been drilled at Talos Dome (72°49′ S, 159°11′ E; 2315 m a.s.l.; Fig. 1), a peripheral ice dome in the Ross Sea sector of East Antarctica (Frezzotti et al., 2004). It is 1620 m long, and its ice age at 1438 m deep (depth refers to the distance between the ice surface and the considered section of the core) is ~ 150 ka according to the AICC2012 chronology (Veres et al., 2013; Bazin et al., 2013), but there is evidence that the water stable isotope record is preserved up to ~ 1550 m deep, where the estimated ice age is ~ 343 ka, corresponding to Marine Isotope Stage 10.1 (Crotti et al., 2021). Dust concentration in TALDICE during the last climatic cycle reflects high atmospheric loads in glacial periods and lower ones during interglacials, as observed at other East Antarctic sites (Delmonte et al., 2010). A peculiarity of TALDICE is given by the influence of local Antarctic dust sources, corresponding to the ice-free sites of the Victoria Land region, close to Talos Dome (Albani et al., 2012a; Bacco et al., 2018b).



**Figure 1.** The mineral dust record of TALDICE. From the upper to the lower curve: (a) total dust concentration of insoluble particles (Baccolo et al., 2021); (b) FPP (fine particle percentage); (c) CLPP (coarse local particle percentage); (d)  $\delta^{18}\text{O}$  (Stenni et al., 2011). Light blue bands highlight glacial culminations (MIS 2, MIS 4, MIS 6 and MIS 8), the red bands the Holocene and MIS 5.5, MIS 7.5 and MIS 9.5. In the upper left corner is the position of Talos Dome in Antarctica.

## 2.2 Sample preparation

Fifty-four samples are prepared using 191 TALDICE ice sections ( $\sim 25 \times 3 \times 2$  cm). They consist of insoluble mineral particles extracted from meltwater and deposited on filtration membranes. The preparation takes place in a clean room (ISO6) installed at the University of Milano-Bicocca. Ice sections are decontaminated with three baths in ultra-pure Milli-Q water (Merck Millipore). They are stored in clean tubes under an ISO 5 laminar flow bench until melted. Meltwater is split into two aliquots, one ( $\sim 10$  mL) for the Coulter counter method and the remnant for synchrotron radiation analysis.

The aliquots for the Coulter counter (CC), corresponding to single ice sections, are stored in plastic cuvettes (rinsed with Milli-Q water) and added with a NaCl solution (final  $\text{Na}^+$  concentration of samples  $\sim 1\%$  m/m). The solution is prepared using Milli-Q water and high-purity solid NaCl, and before use it is filtered with  $0.22\ \mu\text{m}$  pore-size filters. This procedure makes liquid samples electrically conductive, a requisite for the CC. The aliquots for synchrotron radiation analysis are merged considering multiple ice sec-

tions (191 in total) so as to prepare 54 samples with a total dust mass of at least  $1\ \mu\text{g}$ . Merged samples are filtrated using hydrophilic PTFE membranes ( $\varnothing 13$  mm, pore-size  $0.4\ \mu\text{m}$ ). Before filtration, membranes are rinsed for 2 weeks in a high-purity  $\text{HNO}_3$  solution (concentration  $5\%$  m/m, weekly renewed). Filtration is performed with a micro-pipette to concentrate the particles on the membrane in the smallest possible area (Macis et al., 2018). After filtration the membranes are placed in dedicated clean PTFE holders and sealed in plastic bags.

A total of 7 samples correspond to the Holocene (0–673 m, 0–11.7 ka), 7 to the last deglaciation (674–827 m, 11.7–18 ka), 5 to Marine Isotope Stage (MIS) 2 (828–951 m, 18–30 ka), 15 to MIS 3 (952–1259 m, 30–60 ka), 2 to MIS 4 (1260–1292 m, 60–80 ka), 6 to MIS 5 (1293–1418 m, 80–146 ka), 1 to MIS 6 (1419–1438 m, 146–154 ka) and 11 to the deep part of TALDICE not dated by AICC2012 chronology (1439–1620 m) but partially dated by the new deep chronology (Crotti et al., 2021).

## 2.3 Coulter counter

To determine the concentration and grain size of insoluble dust particles, the CC method is used. The method is well-known for the analysis of insoluble particles in ice cores. It relates the changes in the electrical conductance of meltwater with the size and number of insoluble particles present in the suspension (Delmonte et al., 2002). Samples are measured with a Beckman Multisizer 4, equipped with a  $30\ \mu\text{m}$  orifice to measure the concentration of particles between  $0.6$  and  $18\ \mu\text{m}$  divided into 400 channels.

## 2.4 Synchrotron radiation spectroscopic measurements

The application of synchrotron light to determine the elemental and mineralogical composition of TALDICE dust is performed at beamline B18 of Diamond Light Source (Cibin et al., 2019). A glovebox is connected to the experimental chamber of the beamline to handle the samples in clean conditions. Additional precautions are adopted to limit contamination and increase the signal-to-noise ratio: the application of plastic sheets inside the experimental chamber to limit radiation backscattering, the defocusing of the incidental beam to illuminate the largest part of the samples and the preservation of a high vacuum during the acquisition. Further details are available in Baccolo et al. (2018a).

### 2.4.1 X-ray fluorescence spectroscopy

Major elements in dust are investigated through X-ray fluorescence spectroscopy, using synchrotron radiation as the excitation source (Iida, 2013). Samples are irradiated with a  $10\ \text{keV}$  beam (cross section  $\sim 1 \times 1\ \text{mm}$ ) for 600 s, and the fluorescence signal is acquired with a silicon drift detector, allowing the quantification of the following elements: Na, Mg, Al, Si, K, Ca, Ti, Mn and Fe. Analytical accu-

racy is evaluated analyzing NIST standard reference materials (SRM 2709a); it decreases from light to heavy elements (standard deviation of the replicates for Na is 25 % and for Fe is 10 %). Recovery factors are evaluated comparing certified concentrations of standard reference materials (SRMs) with calculated values: they range from 85 % to 115 % except for Ca and Na (133 % and 129 %). Full details are given in Baccolo et al. (2018a, b). Elemental concentrations are converted into oxide concentrations and are close to 100 % (Rudnick and Gao, 2003).

### 2.4.2 X-ray absorption spectroscopy

Speciation and mineralogy of the Fe fraction of TALDICE dust are investigated through X-ray absorption near-edge structure spectroscopy (XANES), performed at the Fe K-edge transition, that is the excitation energy of the innermost electrons of Fe. XANES relates the spectral features of X-ray absorption spectra to chemical and molecular characteristics of specific elements. The sample is irradiated with a monochromatic beam of photons whose energy finely changes with time. The response of the sample depends on features such as oxidation, coordination and mineralogy (Calvin, 2013). For each sample three measurements are carried out, acquiring the fluorescence signal of the samples at steps of 0.15 eV and considering the interval between 7000 and 7400 eV. Spectra are calibrated, normalized and averaged using the Athena software (Ravel and Newville, 2005). Three spectral features are gathered: (1) the energy of the Fe K-edge transition, (2) the energy of the pre-edge peak centroid and (3) the intensity of the pre-edge peak (Fig. S1 in the Supplement) (Baccolo et al., 2018b). The energy of both the pre-edge peak and the Fe K-edge transition is directly related to the oxidation state of Fe, while the intensity of the pre-edge peak depends on its coordination (Berry et al., 2003).

### 2.4.3 Relative abundance of Fe-bearing minerals

Comparing dust samples with the ones corresponding to 14 Fe-mineral references (biotite, chlorite, glaucophane, goethite, hematite, hornblende (ferro-hornblende), jarosite, magnetite, muscovite, fayalite, pargasite, pyrite, schorl, siderite), it is possible to estimate the contribution of single Fe-bearing minerals to the samples (Shoenfelt et al., 2018). XANES spectra of minerals are collected following the protocol adopted for TALDICE samples (Fig. S2). XANES spectra of ice core dust are reproduced through ordinary least-squares regression (OLS), using linear combinations constructed with mineral spectra (Fig. S3). For each sample all the combinations defined by four, three, two and one mineral references (1456 combinations per sample) are calculated through OLS, and the best one (in terms of  $R^2$ , it always exceeded 0.9) is selected to represent the sample. In some cases, the second best-fit combination has  $R^2$  value

close to the best fit, but the difference between the two combinations always regards the less abundant of the four selected references, with negligible effects on the interpretation. The combinatoric package of the Athena software is used (Ravel and Newville, 2005). Following the procedure adopted by Shoenfelt et al. (2018), the relative abundance of Fe-bearing minerals is estimated considering the percentage of linear coefficients obtained from OLS.

## 3 Results and discussion

### 3.1 The TALDICE dust record

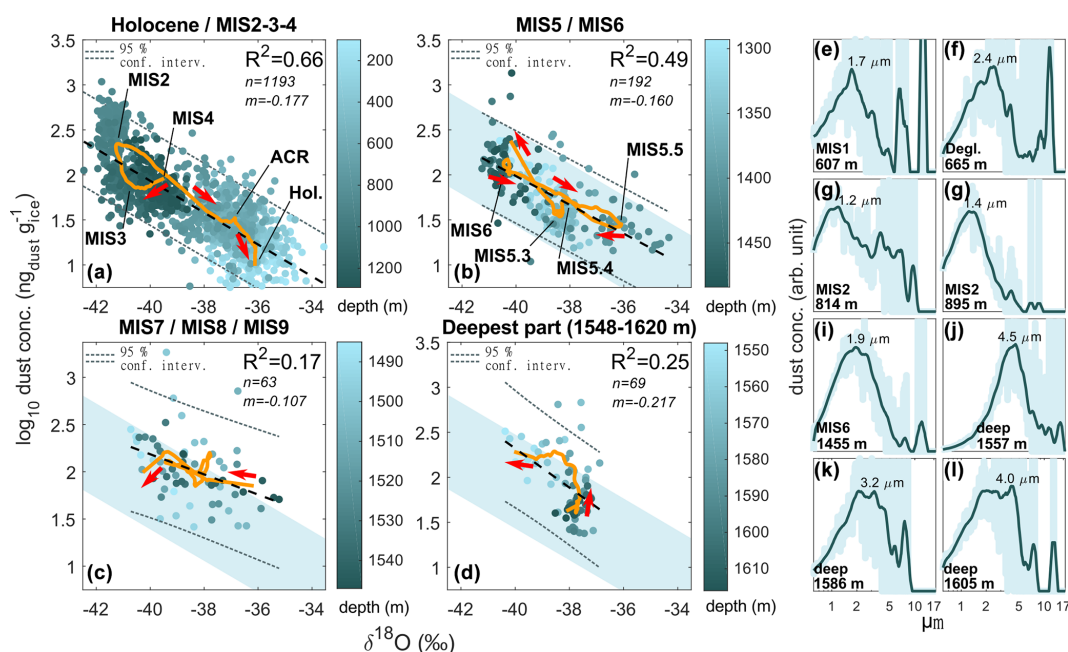
Considering the last climatic cycle (the Holocene and MIS 2, MIS 3 and MIS 4), well-known features characterizing the relationships between the atmospheric dust cycle and Antarctic climate are visible (Fig. 1). The most evident is the negative correlation between dust concentration and  $\delta^{18}\text{O}$ . In accordance with the suppression of dust production and transport from remote sources during interglacials (Albani et al., 2012b), the mean dust concentration in TALDICE Holocene ice is  $\sim 25 \text{ ng g}^{-1}$ , while during MIS 2 it exceeds  $300 \text{ ng g}^{-1}$ , as a consequence of the activation of South American sources (Sugden et al., 2009) and the enhanced atmospheric transport toward Antarctica (Markle et al., 2018).

Such a shift from interglacial to glacial conditions affects not only dust concentration but also its grain size, as revealed by the fine particle percentage (FPP) and coarse local particle percentage (CLPP) indexes. The first one is the relative concentration of particles between 0.6 and  $2 \mu\text{m}$  with respect to the 0.6– $5 \mu\text{m}$  interval; CLPP is the ratio between the concentration of particles between 5 and  $10 \mu\text{m}$  and the total concentration of particles between 0.6 and  $10 \mu\text{m}$ . During the Holocene, FPP has a mean value of 50 %, while during MIS 2 it increases to 63 %, revealing that under glacial conditions dust particles deposited at Talos Dome are smaller than in interglacial periods. Similar evidence has already been observed in other East Antarctic sites (Dome C, Vostok) and interpreted considering that dust transported to East Antarctica during glacials is subject to long-range and high-altitude atmospheric pathways, allowing the efficient removal of coarse particles (Delmonte et al., 2002). CLPP has a mean value of 19 % during the Holocene and of 6.5 % during MIS 2. This index is indicative of the relative abundance of particles larger than  $5 \mu\text{m}$ , which are related to local Antarctic sources (Albani et al., 2012a; Baccolo et al., 2018b). The decrease in the index in glacial periods must be interpreted in relative terms. CLPP is lower in glacials not because of a reduction in coarse particles but due to an increase in the fine ones from South America (Baccolo et al., 2018b).

MIS 4 displays features similar to MIS 2, while MIS 5.5 is similar to the Holocene, confirming increased deposition of fine dust during glacial culminations and reduced fluxes of relatively coarse dust during peak interglacial warmth. In the

**Table 1.** Dust concentration in TALDICE. For each climatic period mean concentrations are reported along with standard deviations.

Period (ka)	Holocene (0–11.7)	Degl. (11.7–19)	MIS 2 (19–31)	MIS 3 (31–58)	MIS 4 (58–68)	MIS 5 (68–132)	MIS 6 (132–190)	MIS 7 (190–246)	MIS 8–MIS 9 (246–337)	Deep part (337–unk.)
Conc. ( $\text{ng g}^{-1}$ )	$26 \pm 18$	$104 \pm 25$	$317 \pm 174$	$80 \pm 31$	$116 \pm 59$	$61 \pm 55$	$158 \pm 190$	$182 \pm 302$	$121 \pm 73$	$110 \pm 123$
FPP (%)	$50 \pm 4$	$57 \pm 8$	$63 \pm 8$	$55 \pm 7$	$53 \pm 5$	$52 \pm 5$	$54 \pm 7$	$45 \pm 9$	$35 \pm 10$	$36 \pm 9$
CLPP (%)	$19 \pm 7$	$14 \pm 7$	$6.5 \pm 3.4$	$10 \pm 5$	$9.6 \pm 3.9$	$13 \pm 8$	$9.1 \pm 5.2$	$21 \pm 14$	$21 \pm 6$	$24 \pm 11$



**Figure 2.** (a–d) Linear correlation between  $\delta^{18}\text{O}$  and dust concentration (logarithmic scale) in TALDICE. Dashed lines highlight the 95 % confidence level interval of linear regression in each period; the blue band in panels (b–d) refers to the 95 % confidence level interval of the linear regression calculated considering the last climatic cycle (a). The orange curve is the trajectory showing the evolution with time of the  $\delta^{18}\text{O}$ –dust pair (arrows from deep to shallow). It was obtained applying a first-order (40 % sample window) Savitzky–Golay filter to both variables. (e–l) Dust grain size distributions from TALDICE.

deeper layers of the core, the relationship between dust and climate is less evident. Figure 1 and Table 1 show that below 1430 m deep, dust concentration exceeds  $100 \text{ ng g}^{-1}$  with dampened oscillations. CLPP exceeds 20 % and FPP drops below 45 %, highlighting that in deep TALDICE dust particles are coarser. This can be interpreted as a consequence of dust aggregation in deep ice, a process already observed in Antarctic ice below 2500 m deep (Lambert et al., 2008; De Angelis et al., 2013).

Size distributions of insoluble particles also point to dust aggregation. Dust from the upper part of the core (Fig. 2e–h) presents a mode between 1 and  $2.5 \mu\text{m}$  and a tail of particles larger than  $5 \mu\text{m}$ , suggesting a mix from remote and local sources (Albani et al., 2012a). Dust from deep TALDICE (Fig. 2j–l) is characterized by a higher abundance of coarse particles, a lack of fine ones and modal values exceeding  $4 \mu\text{m}$ . Such features are not encountered in shallow sections of Antarctic ice cores, where the effects of post-depositional alterations are limited (Royer et al., 1983; Delmonte et al.,

2002; Wegner et al., 2015); they result from in situ aggregation of particles in deep ice (Lambert et al., 2008; De Angelis et al., 2013; Baccolo et al., 2021).

Post-depositional processes in TALDICE ice also influence the climatic significance of the dust record. This is shown in Fig. 2a–d, where the correlation between dust and ice stable isotopes is analyzed. The correlation is high during the last climatic cycle but decreases in older periods ( $R^2$  decreases from 0.66 during the last cycle, to less than 0.3 in the deep core). The same conclusion is drawn looking at the trajectory describing the evolution of the  $\delta^{18}\text{O}$ –dust pair. During the first cycle it well reproduces the transition from glacial to interglacial conditions, including the Antarctic Cold Reversal and the partition of the last glacial period into MIS 2, MIS 3 and MIS 4. This is partially true for the previous climatic cycle, but the correlation decreases ( $R^2$  from 0.66 to 0.49). The degradation continues in the deepest part of TALDICE, where the coefficient does not exceed 0.25 and the trajectories follow irregular paths, high-

lighting a substantial decoupling of dust concentration and isotopic signals. This is confirmed by Crotti et al. (2021), who showed that at Talos Dome below 1548 m deep, the climatic signals are not preserved.

### 3.2 Major element composition of mineral dust

TALDICE dust composition during the last climatic cycle resembles the signature of post-Archean Australian shale (PAAS; Taylor and McLennan, 1985; Fig. 3). With respect to the upper continental crust (UCC; Rudnick and Gao, 2003), PAAS is depleted in mobile oxides (such as CaO and Na<sub>2</sub>O) and enriched in Fe, Ti and Al oxides as a result of chemical weathering. PAAS is in fact representative of surficial sedimentary rocks subject to chemical weathering (Taylor and McLennan, 1985), while UCC of the whole upper continental crust. The similarity between TALDICE dust and PAAS is not unexpected, since atmospheric dust is produced at the Earth surface, where sedimentary and weathered rocks dominate.

Another feature emerging from Fig. 3 is represented by the change in dust composition with depth (see also Fig. S4). The oxides showing the most evident trends are SiO<sub>2</sub> (increasing) and CaO (decreasing). Their average concentration varies between 64 % and 1.7 % in the Holocene to 74 % and 0.4 %, respectively, in the deep part of the core. Other oxides showing minor variations are Na<sub>2</sub>O (decreasing), MgO (decreasing), Al<sub>2</sub>O<sub>3</sub> (decreasing) and K<sub>2</sub>O (increasing). These variations are related to depth and not to climatic cycles. This is an indication that they are likely related to post-depositional processes and not to primary changes in dust sources. A disturbance from the bedrock must be also discarded since the ice stratigraphy at Talos Dome is uninterrupted until 1548 m deep (Crotti et al., 2021). Ca, Mg and Na, the elements showing the strongest decrease, are mobile and typically affected by chemical weathering (Nesbitt and Young, 1982). Their reduction suggests a progressive alteration of dust with depth. In particular, the deepest samples lack MgO and CaO, indicating carbonate dissolution. The reaction between acidic species and carbonates and the consequent precipitation of gypsum are well-known post-depositional processes in deep ice (Ohno et al., 2006; Iizuka et al., 2008; Traversi et al., 2009; Eichler et al., 2019). Our evidence suggests that this reaction also occurs at Talos Dome. The increase in SiO<sub>2</sub> is likely relative and reflects the progressive loss of labile species.

### 3.3 Iron oxidation and coordination symmetry

XANES reveals a progressive oxidation of Fe with depth (Fig. 4). Figure 4a shows that in the first 1000 m of the core, Fe in dust consists of a mixture of Fe<sup>2+</sup> and Fe<sup>3+</sup>, reflecting the typical composition of mineral aerosols (Schroth et al., 2009). Deep samples show a pure Fe<sup>3+</sup> signature. Regarding Fe coordination, samples display an octahedral symmetry

(coordination number 6), with secondary inputs from other geometries. This is also in accordance with observations concerning aerosols (Wilke et al., 2001; Formenti et al., 2014). Figure 4b shows the correlation between the energy of the pre-edge peak and of the K-edge transition, both related to Fe oxidation (Berry et al., 2003). Figure 4c shows the variation in the K-edge energy transition along the core. It increases by ~ 2 eV, pointing to the oxidation of Fe in mineral dust found in deep ice. The process is rather continuous, regardless of the climatic oscillations, suggesting that it relates to post-depositional changes. This is confirmed by the similarity of the oxidation trend and the growth of ice grains with depth (Fig. 4c). There are two exceptions: (1) in the deepest part of the core the Fe K-edge reaches a stable energy, pointing to a complete oxidation, and (2) in correspondence with MIS 2 and MIS 4 the trend shows two slowdowns, as if oxidation is inhibited.

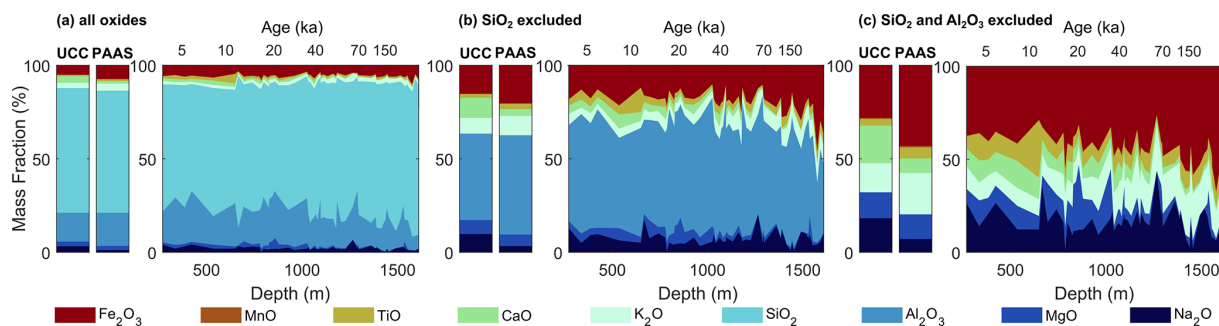
The partial increase in Fe<sup>2+</sup> during glacial culminations is related to the transport of fresh glacial dust from South America (Spolaor et al., 2013), only partially oxidized due to limited atmospheric exposure (Shoenfelt et al., 2017). Another reason for the slowdown of Fe oxidation during glacials is that in these periods the enhanced dust deposition at Talos Dome acts as a buffer and partially neutralizes the acidity of ice, consuming reactive species, as proposed for the Dome Fuji and EDML ice cores (Ohno et al., 2005, 2006; Eichler et al., 2019), and inhibits dust oxidation. Figure 4c shows that the growth of ice grains is also temporarily inhibited during MIS 2 probably because of grain boundary pinning by insoluble particles as suggested by Durand et al. (2006). A similar pattern is visible in MIS 4, while ice corresponding to MIS 6 does not present either an Fe<sup>2+</sup> recovery or a decrease in ice grain size (Fig. 4c), probably because in situ oxidation of Fe minerals and ice re-crystallization are too advanced. On the contrary, in interglacial ice atmospheric acidity is more available for post-depositional reactions favoring oxidation and weathering, also thanks to a more efficient ice re-crystallization (Iizuka et al., 2008; Eichler et al., 2019).

### 3.4 Iron mineralogy

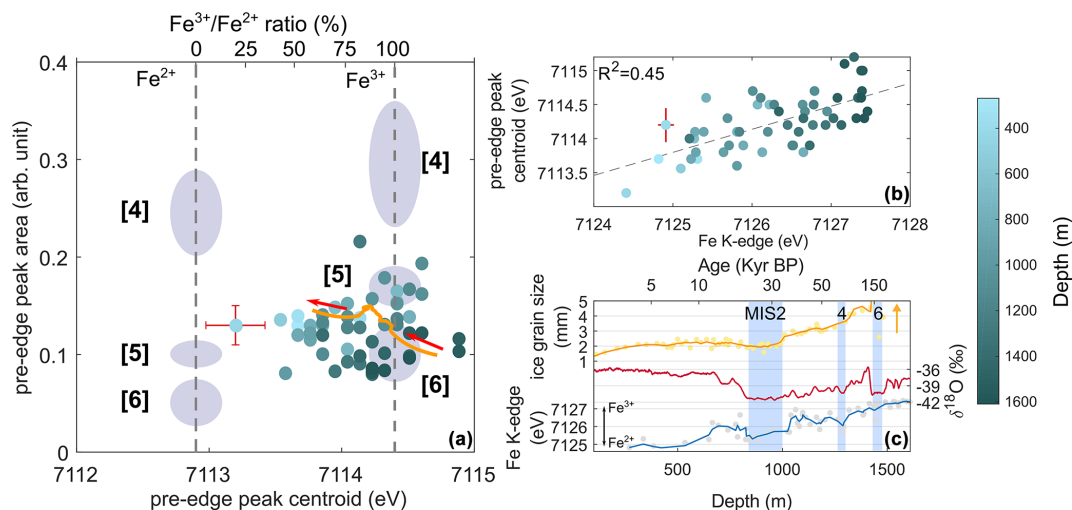
Fe-mineralogy results are shown in Figs. 5 and 6 and Table 3. Only minerals whose average relative abundance exceeds 2 % have been considered in the discussion.

#### 3.4.1 Hornblende and jarosite

Minerals showing the most evident trends are hornblende and jarosite. Hornblende dominates samples in the first 1000 m of TALDICE with a decreasing trend (Fig. 6a); on the contrary jarosite is present only below 1000 m deep, and its concentration increases with depth (Fig. 6h). Trends related to these minerals involve large parts of the core regardless of climatic conditions; we interpret them as a consequence of post-depositional processes. Hornblende,



**Figure 3.** Major element in dust from TALDICE. Data are shown considering all major element oxides (a), excluding  $\text{SiO}_2$  (b), and excluding  $\text{SiO}_2$  and  $\text{Al}_2\text{O}_3$  (c). UCC and PAAS references are shown for comparison (Rudnick and Gao, 2003; Taylor and McLennan, 1985).



**Figure 4.** XANES results from the analysis of dust in TALDICE. (a) Analysis of the Fe K-pre-edge spectral region. The intensity and energy position of the pre-edge peak are shown following the scheme proposed by Wilke et al. (2001). Ellipses and numbers in square brackets refer to the coordination of Fe and vertical lines to its two oxidation states; the orange curve represents the trajectory of the samples along the core, it was calculated as in Fig. 2 (arrows from deep to shallow). (b) A comparison between the energy position of the pre-edge peak and the main K-edge transition (see Fig. S1 for details). (c) The energy position of the Fe K-edge transition (blue line) vs. the isotope composition of ice (red line) (Stenni et al., 2011) and ice grain size (yellow line) (Montagnat et al., 2012); the arrow indicates the observed but not quantified ice crystals larger than 40 cm found below 1481 m deep; blue bands highlight MIS 2, MIS 4 and MIS 6. For each panel one point presents mean error bars (not visible in panel c because of scale).

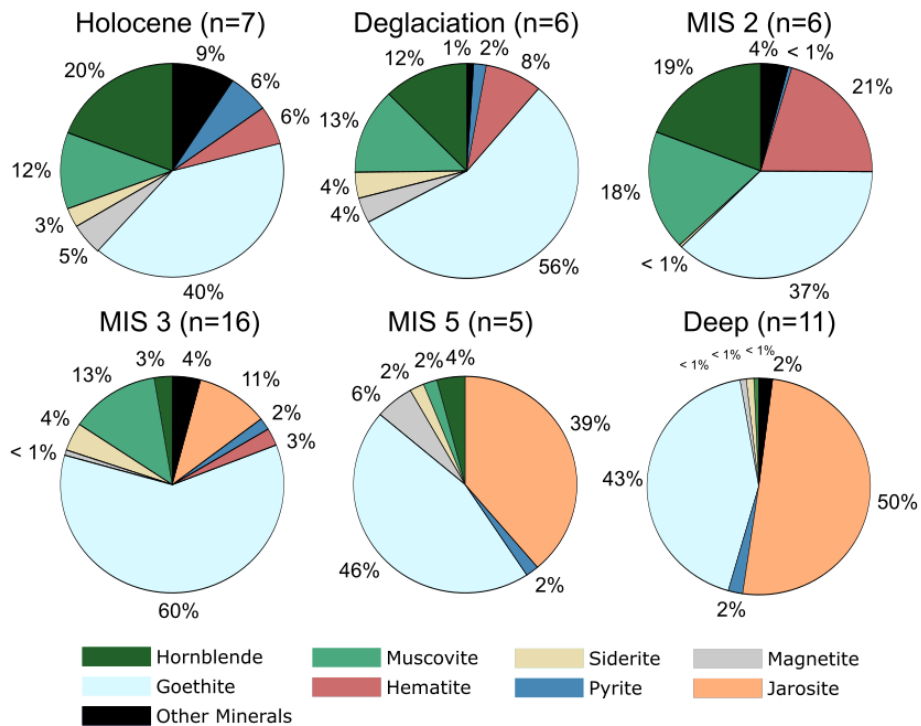
$\text{Ca}_2(\text{Fe}_4^{2+}\text{Al})(\text{Si}_7\text{Al})\text{O}_{22}(\text{OH})_2$ , is the second-most-abundant Fe mineral in Holocene ice (relative abundance 19%), but it rapidly decreases with depth and in the deepest part of TALDICE it is almost absent (relative abundance 0.7%). Jarosite shows an opposite behavior. It is not present in the shallow part of TALDICE; it appears below 1000 m deep, becoming the dominant Fe-bearing mineral below 1400 m (relative abundance 50%). The two trends are linked. Hornblende is a common ferrous mineral present in South American dust (Shoenfelt et al., 2018) and is one of the dominant  $\text{Fe}^{2+}$ -bearing minerals in global aerosols (Schroth et al., 2009). Jarosite,  $\text{KFe}_3(\text{SO}_4)_2(\text{OH})_6$ , is not common in atmospheric dust. It is well-known for being a weathering product, and its widespread identification in the deep layers of TALDICE is regarded as evidence of weathering affecting

dust in deep ice (Baccolo et al., 2021). The concurrent decrease in hornblende and increase in jarosite confirm Fe oxidation. Hornblende seems to be the principal mineral whose dissolution leads to the consumption of  $\text{Fe}^{2+}$ , while jarosite buildup drives the accumulation of  $\text{Fe}^{3+}$ . The other minerals presenting a ferrous component, with the exception of magnetite, also show a decreasing trend (muscovite, siderite, pyrite) and are almost absent in the deepest part of the core.

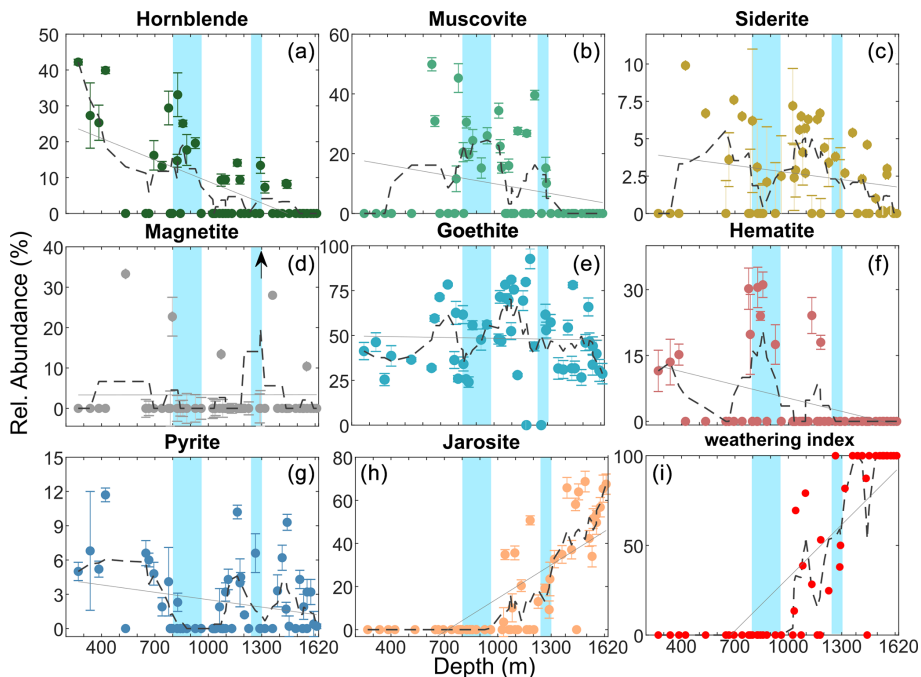
### 3.4.2 Siderite and pyrite

Some minerals show a pattern in correspondence to MIS 2; in some cases it is a relative maximum (muscovite, hematite) and in others a minimum (siderite, pyrite). Considering the correspondence with MIS 2, such features are interpreted as





**Figure 5.** Pie charts showing the Fe mineralogy of TALDICE dust. Each chart refers to a climatic period. Marine isotope stages 4 and 6 have been excluded because of the low number of samples corresponding to these periods.



**Figure 6.** The variation in the major Fe-bearing minerals in TALDICE dust with depth. Dashed lines correspond to a 5-point moving average and solid grey lines to linear trends. In panel (d), one sample point is off the scale (black arrow, magnetite relative abundance 70 %). Panel (i) shows the weathering index, defined as the percent ratio between jarosite concentration and the sum of jarosite, hornblende, muscovite and hematite.

climate-related signals. In the Holocene siderite and pyrite,  $\text{FeCO}_3$  and  $\text{FeS}_2$ , respectively, constitute about 3 % each of Fe minerals in TALDICE ice, but during MIS 2 their abundance drops to 0.5 %. A shift in mineralogy between the Holocene and MIS 2 is expected; as in MIS 2 dust is supplied to Talos Dome mostly by Patagonian sources (Delmonte et al., 2004, 2010), while in the Holocene its origin is local from northern Victoria Land (Delmonte et al., 2010; Baccolo et al., 2018b). The presence of siderite and pyrite in Holocene dust at Talos Dome agrees with the geology of Victoria Land, where they are common accessory minerals, owing to the basaltic/doleritic nature of local rocks (Sturm and Carryer, 1970; Dow and Neall, 1974). In addition to the geologic context, atmospheric transport can also partially explain their absence during MIS 2. Both minerals are easily oxidized when exposed to the atmosphere. Their lack of presence in MIS 2 can be related to their oxidation during the long-range transport from South America (Shi et al., 2012). On the contrary, in the Holocene siderite and pyrite did not undergo chemical reactions because of short-range atmospheric transport.

### 3.4.3 Muscovite

Muscovite reaches its highest relative abundance during MIS 2, presenting a mean value of 18 %. Thanks to the aerodynamic shape of its crystals, muscovite, an Al-phyllsilicate, is common in atmospheric dust subject to long-range transport and is one of the most abundant minerals deposited on East Antarctica in MIS 2 (Delmonte et al., 2017; Paleari et al., 2019). Its abundance in glacial ice is likely the main reason for the slowdowns observed along the oxidation trend of Fe (Fig. 4c), as Fe in muscovite is present as both ferric and ferrous iron. During MIS 2 muscovite is the only mineral, with hornblende, presenting a ferrous component, since siderite, magnetite and pyrite are almost completely absent (Fig. 6). A similar pattern is observed in MIS 4, when the second slowdown in Fe oxidation is observed, again corresponding to an increase in muscovite. In this second case both the slowing of oxidation and the increase in muscovite are less evident, probably because of the more advanced weathering at greater depth. Muscovite is completely absent below 1300 m deep, indicating that it is affected by weathering in deep ice, similarly to hornblende. Muscovite dissolution probably supplies a fraction of the K required for jarosite precipitation.

### 3.4.4 Hematite

Hematite,  $\text{Fe}_2\text{O}_3$ , is a weathering product in soils under dry and warm conditions; it is typical in tropical regions, while it is rarely encountered in cold and wet climates (Schwertmann, 1988). In TALDICE it is mostly found in MIS 2, when the dust signature is fully South American (Delmonte et al., 2010). During glacial culminations an additional source other than Patagonia, the Puna–Altiplano dry region in the

tropical Andes, supplies dust to Antarctica (Delmonte et al., 2010). In the Puna–Altiplano hematite is widely present (Aubry et al., 1996). Accordingly, our results show an excess of hematite in MIS 2 (mean relative abundance 21 %). A previous study focusing on inner East Antarctica observes a higher abundance of hematite during the Holocene than in MIS 2 (Paleari et al., 2019). The difference can be related to the geographic position of the two sites and to the influence of secondary sub-tropical sources during the Holocene. Below 1300 m deep, hematite is not observed, suggesting that this mineral is not stable in deep ice at Talos Dome. It is known that under acidic conditions ( $\text{pH} \sim 4$ ) hematite is not stable and dissolves (Schwertmann and Murad, 1983; Zolotov and Mironenko, 2007), leading to the precipitation of jarosite and goethite (Papike et al., 2006).

### 3.4.5 Goethite

Considering the entire core, goethite,  $\text{FeO}(\text{OH})$ , is the dominant Fe mineral in dust, with a mean relative abundance of 46 % (the second is jarosite, 18 % relative abundance). The distribution of goethite with depth is rather uniform (Fig. 6e). This is indicative of the fact that (1) goethite is a common Fe-bearing phase in mineral aerosols (Formenti et al., 2014) and that (2) goethite is stable in the englacial environment, regardless of depth. The second point is corroborated by previous studies showing that at low temperature and in acidic wet environments, goethite is the most stable Fe oxide-hydroxide (Schwertmann and Murad, 1983; Zolotov and Mironenko, 2007).

### 3.4.6 Magnetite

Magnetite,  $\text{Fe}_3\text{O}_4$ , has a mean relative abundance of about 3 % and does not show evident trends; only a single sample presents a high concentration (70.4 %). Without considering this anomalous value, possibly related to contamination or to the presence of a micrometeorite (Rochette et al., 2008), the mean value drops to 2 %. Despite its low relative abundance, magnetite is relatively stable along TALDICE and it is not depleted in the deepest part, suggesting that it is not affected by oxidation and/or dissolution like the other ferrous minerals considered in this study. This confirms its high chemical resistance to acidic oxidation (Moncur et al., 2009).

### 3.4.7 The weathering index

An index was developed to summarize the information from the different Fe minerals. It is defined as the weathering index and corresponds to the percent ratio between the relative abundance of jarosite and the sum between the relative abundance of jarosite, hornblende, muscovite and hematite. The trend of the index varies between 0 % in the upper part of TALDICE and 100 % below 1300 m deep (Fig. 6i), reflecting the progressive weathering of dust, which consists in the

consumption of some ferrous minerals and the precipitation of jarosite.

### 3.5 Englacial weathering of dust

When considering the pair Holocene–MIS 2, which corresponds to the first 950 m of TALDICE, the chemical and physical properties of TALDICE dust are interpretable in the light of the well-known effects of major climatic swings on atmospheric dust. This is true for concentration, grain size and geochemistry. Below 1000 m deep, trends not related to climatic oscillations appear. The most evident is the increase in jarosite. The mineral is first observed at 1000 m deep, and its concentration increases toward the core bottom (Figs. 5 and 6h).

Jarosite is an Fe–K hydrated sulfate and results from chemical weathering. It has never been reported in global mineral aerosols. These features support the hypothesis that jarosite is a product of englacial diagenesis related to dust chemical weathering (Baccolo et al., 2021). Thanks to the conditions under which jarosite precipitates, its formation supplies information about the deep englacial environment. The precipitation of jarosite is evidence for the presence of liquid water in deep ice (pre-melting), since this mineral forms when a limited quantity of acidic (pH < 4) aqueous solutions rich in solutes (brines) interacts with Fe-bearing minerals (Zolotov and Shock, 2005; Papike et al., 2006). At Talos Dome, ice temperature at 1000 m deep, where jarosite appears, is  $-25.5^{\circ}\text{C}$  (Julius Rix and Carlos Martin, personal communication, 2020). Combining this information with the phase diagram of the sulfuric acid–water solution (Beyer et al., 2003), it is possible to estimate that acidic brines forming in deep ice at Talos Dome have a sulfuric acid concentration between 15 % m/m and 25 % m/m. Liquid water in deep ice has been predicted by theories (Rempel et al., 2002; Marath and Wettlaufer, 2020; Ng, 2021) and confirmed by observations (Fukazawa et al., 1998). The englacial formation of jarosite is further confirmation of deep-ice pre-melting. Moreover, it is also indicative that liquid water found in deep ice is strongly acidic and oxidative.

The relationships between ice re-crystallization and the concentration of impurities in deep ice are not yet completely understood (Eichler et al., 2019; Stoll et al., 2021). Our study suggests that such correlation actually exists and that it has effects on dust geochemistry since Fe oxidation and ice grain growth share a similar trend in TALDICE (Fig. 4c). The formation of the acidic brines required for jarosite precipitation is in fact probably related to ice re-crystallization, since the latter has been invoked as responsible for the concentration of impurities in localized environments and for the local lowering of the pressure melting point (De Angelis et al., 2013; Tison et al., 2015). The acidity of deep brines is explained in the light of the strong incompatibility of acidic atmospheric species with respect to the ice molecular lattice (Wolff, 1996)

and of their concentration at grain boundaries (Mulvaney et al., 1988).

Not only does the formation of jarosite alter the original mineralogical assemblage of dust in TALDICE, but it can also be related to the anomalies in dust grain size observed in deep ice (Fig. 2e–l). Jarosite is known for creating a cementing matrix during weathering (Long et al., 1992). Its precipitation promotes the aggregation of mineral particles, altering the original grain size distribution. This has also been confirmed by microscopic observations, as shown in Baccolo et al. (2021). It is likely that chemical weathering also affects dust concentration, explaining the decrease in the correlation with ice isotopic composition in the deep part of the ice core (Fig. 2).

Jarosite formation is not the only geochemical process occurring in deep TALDICE. Another anomaly is the oxidation of Fe present in dust (Fig. 4). The oxidation of Fe minerals under acidic conditions is a well-known weathering pathway (Jones et al., 2014) which, among others, leads to jarosite precipitation (Papike et al., 2006). Fe oxidation is confirmed by the decline of many ferrous minerals in the deep part of the core (hornblende, muscovite, siderite and pyrite; see Figs. 5 and 6 and Table 3). The disappearance of hematite in deep TALDICE, which is not stable under acidic and wet conditions (Schwertmann and Murad, 1983), further supports oxidative/acidic weathering in deep ice at Talos Dome. The weathering index (Fig. 6i) summarizes these mineralogical changes and highlights the progressive weathering of dust with depth.

The trends of major element oxides in TALDICE dust (Table 2, Fig. 3) agree with the scenario described above. The decrease in Ca, Na and Mg with depth is interpreted as an effect of acidic weathering. In the presence of acidic aqueous solutions, mobile and soluble elements are easily mobilized (Nesbitt and Young, 1982). It is worth mentioning that Al oxide is also depleted in deep TALDICE dust. This suggests that the most stable fractions of dust, such as aluminosilicates, also undergo weathering. This is confirmed by the identification in the deep part of the Dome Fuji ice core of secondary aluminum sulfate (Ohno et al., 2014).

TALDICE presents a number of peculiarities if compared to other East Antarctic ice cores: deposition of local dust presenting a basaltic signature (Baccolo et al., 2018a), relatively warm temperatures and a strong oceanic influence. Such features can explain why post-depositional processes affecting dust are so notable at this site. Basaltic–doleritic rocks are easily weatherable, even at low temperature (Li et al., 2016; Niles et al., 2017). In addition, Talos Dome is located near the Southern Ocean and receives a considerable quantity of marine aerosols rich in reactive acidic species (Iizuka et al., 2013; Mezgec et al., 2017). A further factor to consider is ice temperature. Talos Dome is located near the coast, and its climate is tempered by the ocean. When considering the same depth, ice temperature at Talos Dome is about  $15\text{--}20^{\circ}\text{C}$  warmer than at inner sites (Talalay et al., 2020), and tempera-

**Table 2.** Average major element composition of TALDICE dust. Data are expressed as percent mass fractions of oxides. For each climatic period mean values are reported with standard deviations.

Period (ka)	Holocene (0–11.7)	Degl. (11.7–19)	MIS 2 (19–31)	MIS 3 (31–58)	MIS 4 (58–68)	MIS 5 (68–132)	MIS 6 (132–154)	Deep part (154–unk.)
Na <sub>2</sub> O (% m/m)	3.1 ± 1.1	2.1 ± 1.0	1.9 ± 0.5	2.0 ± 0.5	4.3 ± 3.3	2.3 ± 1.0	0.2	1.6 ± 0.9
MgO (% m/m)	1.6 ± 0.4	2.0 ± 1.0	3.0 ± 1.6	0.8 ± 0.3	0.8 ± 1.1	0.4 ± 0.3	0.5	0.6 ± 0.8
Al <sub>2</sub> O <sub>3</sub> (% m/m)	20.7 ± 5.3	19.7 ± 4.8	23 ± 2.5	16.3 ± 3.9	23.5 ± 8.3	14.7 ± 3.9	17.2	13.5 ± 5.3
SiO <sub>2</sub> (% m/m)	63.9 ± 5.0	65.9 ± 6.3	62.9 ± 3.4	72 ± 3.6	63.9 ± 4.5	73.4 ± 3.9	72.9	74.2 ± 5.6
K <sub>2</sub> O (% m/m)	1.6 ± 0.3	1.8 ± 0.5	1.7 ± 0.3	2.0 ± 0.4	2.6 ± 1.7	2.3 ± 0.5	2.5	2.4 ± 0.5
CaO (% m/m)	1.7 ± 0.4	1.3 ± 0.5	1.0 ± 0.4	0.9 ± 0.4	0.4 ± 0.2	0.4 ± 0.2	0.3	0.4 ± 0.3
TiO <sub>2</sub> (% m/m)	1.5 ± 0.4	1.6 ± 1.3	0.9 ± 0.2	0.9 ± 0.2	0.4 ± 0.2	0.9 ± 0.3	0.8	1.0 ± 0.2
MnO (% m/m)	0.06 ± 0.02	0.05 ± 0.01	0.06 ± 0.02	0.04 ± 0.02	0.05 ± 0.04	0.04 ± 0.01	0.04	0.04 ± 0.02
Fe <sub>2</sub> O <sub>3</sub> (% m/m)	5.8 ± 0.6	5.6 ± 1.3	5.4 ± 1.1	5.1 ± 0.8	4.0 ± 1.0	5.5 ± 0.6	5.6	6.3 ± 1.5

**Table 3.** Average relative abundance of Fe-bearing minerals in TALDICE dust. Data are expressed as percent abundances.

Period (ka)	Holocene (0–11.7)	Degl. (11.7–19)	MIS 2 (19–31)	MIS 3 (31–58)	MIS 4 (58–68)	MIS 5 (68–132)	MIS 6 (132–154)	Deep part (154–unk.)
Hornblende (%)	19.2	12.3	19.1	2.6	0	4.1	8.4	0.7
Muscovite (%)	11.5	12.9	18.0	13.0	7.6	2.0	0	0
Siderite (%)	2.9	3.9	0.4	4.2	1.9	2.1	0	1.1
Magnetite (%)	4.8	3.8	0	0.8	35.2	5.6	0	0.9
Goethite (%)	30.8	56.0	37.4	60.3	30.8	45.5	31.7	42.9
Hematite (%)	15.7	8.3	20.6	2.6	0	0	0	0
Pyrite (%)	5.9	1.8	0.5	1.8	3.3	1.9	1.7	2.1
Jarosite (%)	0	0	0	10.5	14.3	38.8	58.1	50.3
Others (%)	9.1	1.0	4.0	4.1	7.0	0	0	2.0

ture is an essential parameter for the relocation of impurities in deep ice (Marath and Wettlaufer, 2020).

All these local features are likely related to the ice metamorphism observed in deep TALDICE (Montagnat et al., 2012), where ice crystals of up to 40–50 cm have been observed below 1480 m deep. High ice temperature and metamorphism could partially explain why dust alteration is so relevant in TALDICE, while at inner sites the original properties of dust seem preserved at greater depth and further back in time (Delmonte et al., 2004; Kawamura et al., 2017). Replicating this study at inner Antarctic sites will be essential to distinguishing the processes depending on the local characteristics of single sites from the ones more deeply related to ice depth and age.

#### 4 Conclusions and perspectives

This study provides a first description of dust chemical weathering in deep polar ice. Grain size, concentration and mineralogy and composition of dust are all affected by post-depositional processes in TALDICE. Fe speciation and mineralogy investigated through synchrotron radiation are efficient probes to explore such transformations in Antarctic ice. The englacial precipitation of jarosite, the oxidation of Fe in dust, the decline of ferrous minerals (hornblende,

pyrite, siderite, muscovite) and of hematite, and the depletion of some major elements (Ca, Mg, Na) suggest that below 1000 m deep, dust in TALDICE is affected by acidic oxidative weathering. The latter results from the formation of acidic brines which interact with dust in deep ice, leading to geochemical reactions. The production of such brines is likely related to ice re-crystallization and to the accumulation of impurities in highly localized environments where the interaction between soluble and insoluble species is favored. From this perspective deep Antarctic ice can be seen as a “geochemical reactor” capable of promoting the precipitation of secondary minerals and the dissolution of others. This study shows that dust-related signals in ice cores, traditionally considered stable and resistant to post-depositional processes, are significantly altered in deep ice, at least in TALDICE.

It would be desirable to replicate the present study considering deep ice cores from inner East Antarctica where it would be possible to unveil additional processes related to the conditions found 2000–3000 m deep in the ice. Another possible implementation is the concurrent analysis of soluble and insoluble impurities, including elements other than Fe. This will help to investigate the dissolution and precipitation of primary and secondary minerals and identify the geochemical reactions responsible for these transformations,

paving the way for the development of englacial geochemistry.

**Data availability.** Data used to meet the aims of the present study are available in the Supplement. Full XANES spectra are found in the PANGAEA open repository, <https://doi.org/10.1594/PANGAEA.924114> (Baccolo et al., 2020).

**Supplement.** The supplement related to this article is available online at: <https://doi.org/10.5194/tc-15-4807-2021-supplement>.

**Author contributions.** GB conceived the idea of this work. GB, BD and EDS prepared the samples and performed Coulter counter analyses. GB, GC, DH and AM carried out X-ray absorption and fluorescence analyses. GB interpreted the data and wrote the manuscript with contributions from all the co-authors.

**Competing interests.** The contact author has declared that neither they nor their co-authors have any competing interests.

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**Special issue statement.** This article is part of the special issue "Oldest Ice: finding and interpreting climate proxies in ice older than 700 000 years (TC/CP/ESSD inter-journal SI)". It is not associated with a conference.

**Acknowledgements.** The Talos Dome Ice Core (TALDICE) project, a joint European program, is funded by national contributions from Italy, France, Germany, Switzerland and the United Kingdom. Primary logistical support was provided by PNRA at Talos Dome. This is TALDICE publication no. 62. This publication was generated in the frame of Beyond EPICA. The project has received funding from the European Union's Horizon 2020 Research And Innovation Programme under grant agreement no. 815384 (Oldest Ice Core). It is supported by national partners and funding agencies in Belgium, Denmark, France, Germany, Italy, Norway, Sweden, Switzerland, the Netherlands and the United Kingdom. Logistic support is mainly provided by PNRA and IPEV through the Concordia Station system. The opinions expressed and arguments employed herein do not necessarily reflect the official views of the European Union funding agency or other national funding bodies. This is Beyond EPICA publication number 22. We thank Paolo Gentile for providing mineral standards and also Paul Niles and Tanya Peretyazhko for the fruitful discussions.

**Financial support.** This research has been supported by the Ministero dell'Istruzione, dell'Università e della Ricerca (grant no. PNRA18-00098); Horizon 2020 (grant no. Beyond EPICA (815384)); and Diamond Light Source (grant nos. sp7314, sp8372 and sp9050).

**Review statement.** This paper was edited by Florent Dominé and reviewed by Nicolas Stoll and one anonymous referee.

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