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Major coral extinctions during the early Toarcian global warming event

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18

19 ABSTRACT

20

The current loss of taxa from our planet is considered by many scientists to be the sixth mass
extinction; in particular, coral reefs are experiencing significant damage. Scleractinian coral reefs

23 constitute the framework of some of the most diverse ecosystems on Earth today as well as 24 throughout the last 250 million years. The Pliensbachian-Toarcian (Pl-To) transition in the Early 25 Jurassic is an important geologic analogue for present-day global disruptions as it provides insights about the ability of corals to survive significant changes in climatic and oceanic 26 27 conditions. Based on a revised and greatly expanded taxonomy of Tethyan coral faunas for both 28 stages, we reveal a catastrophic extinction (49% at genus and 90.9% at species level) and 29 subsequent radiation for the group during the early Toarcian. The devastation during the late 30 Pliensbachian and early Toarcian biotic crisis suggests that this interval is potentially the most 31 important extinction event for scleractinian corals and fundamentally shaped the future of 32 diversity and morphological disparity within the coral clades that flourished during the Jurassic 33 and Cretaceous. Consequently, this interval should be considered a key period in the evolutionary dynamics of the group and reef ecosystems more broadly. These data on coral 34 35 survival and recovery during the Early Jurassic event are particularly pertinent for questions 36 about present day ecosystem collapse and conservation of coral habitats.

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38 Keywords: early Toarcian, Ocean anoxic event, corals, crisis, Scleractinia,

39

40 1. INTRODUCTION

41

The history of life is punctuated by extinction events that drove evolutionary dynamics, and scientists posit that the Earth is currently entering a similar mass extinction (Ceballos et al., 2015). Five major "first-order" biotic crises are reported for the Phanerozoic, followed by "second-order" crises, such as the Pliensbachian-Toarcian crisis at ca. 183 Ma (e.g., Raup and 46 Sepkoski, 1982; Bambach et al., 2004). The distinction between first and second-order crises 47 remains unclear as the ranking of crises varies according to the considered characters, such as 48 taxonomic scale (i.e., families, genera, species), functional niche loss, consideration of associated 49 palaeoenvironmental upheaval, and the different ways to statistically weigh raw data (Benton, 50 1995).

51 The Pliensbachian-Toarcian (Pl-To) transition was marked by drastic and repeated 52 palaeoenvironmental disturbances associated with the emplacement of the Karoo-Ferrar igneous 53 province in southern Gondwana (Burgess et al., 2015). These environmental changes included 54 major shifts in the temperature, chemistry, and oxygenation of the oceans, culminating in the 55 Toarcian Oceanic Anoxic Event (T-OAE) (e.g., Suan et al., 2010; Gómez and Goy, 2011; Caruthers et al., 2011; Dera et al., 2011, Reolid et al., 2020). Following the middle Pliensbachian 56 57 warm period (Gómez et al., 2016), the latest Pliensbachian is marked by a significant cooling 58 event with important sea-level fluctuations (de Graciansky et al., 1998; Aurell et al., 2003; Suan 59 et al., 2010; Pittet et al., 2014; Krencker et al., 2019). The PI-To boundary coincides with a pronounced warming event (Suan et al., 2010), enhanced continental weathering, and a 60 61 widespread demise of neritic carbonate ecosystem (Bodin et al., 2016; Krencker et al., 2020), 62 which often led to strong condensation of the sedimentary record or even a hiatus in sedimentation (Morard et al., 2003; Léonide et al., 2012; Krencker et al., 2014; Pittet et al., 63 64 2014). Climate cooling is observed during the remainder of the earliest Toarcian (Polymorphum 65 zone; Ruebsam et al., 2019; Krencker et al., 2020), allowing renewed neritic carbonate 66 production in several areas across the Tethys Ocean (Léonide et al., 2012; Trecalli et al., 2012; 67 Krencker et al., 2020). At the onset of the T-OAE, a second demise of neritic carbonate ecosystems occurred, likely triggered by a profound carbon cycle perturbation, soaring global 68

temperature (Suan et al., 2010; Korte et al., 2015), and a second pulse of continental weathering (Bodin et al., 2010; Krencker et al., 2020; Ettinger et al., 2021). Neritic carbonate production resumed along the southern Tethyan margin during the recovery of the carbon isotope excursion associated with the T-OAE (Trecalli et al., 2012; Krencker et al., 2014, 2015) but not until at least the middle Toarcian along the northern Tethyan margin (Blomeier and Reijmer, 1999; Aurell et al., 2003; Léonide et al., 2012; Brigaud et al., 2014).

75 The Pl-To transition is one of the five most significant reefal crises in the geological record 76 (Flügel and Kiessling, 2002; Kiessling and Simpson, 2011; Kiessling, 2011) and this event is 77 correlated with significant extinctions across a wide number of marine taxa (Little and Benton, 78 1995; Ruban, 2004; Mattioli et al., 2009; Dera et al., 2010; Caruthers et al., 2013; Brame et al., 79 2019) as well as continental plants (Slater et al., 2019). The existing literature on corals during 80 the Pl-To transition was highly dependent on the taxonomic work of Beauvais (1986). Indeed, 81 previous studies (Kiessling et al., 2007; Lathuilière and Marchal, 2009) identified the need for 82 more original field data followed by a revision of the taxonomy based on these new specimens 83 and with consideration of possible intraspecific variability. This study provides a stronger 84 analysis of Pliensbachian and Toarcian coral assemblages based in extensive new collections 85 from Morocco (Middle and High Atlas), Italy (Calabria), France (Calvados and Pyrenees) and a deep revision of taxa from the literature (Vasseur, 2019; Vasseur and Lathuilière, 2020 in press). 86 87 It also offers the first opportunity to examine coral assemblages in between the 88 Pliensbachian/Toarcian stage boundary and T-OAE.

89

90 **2. MATERIAL:**

92 Data have been gathered by the authors from five field missions in Morocco (Dadès Valley, 93 Amellagou, Ziz Valley and Guigou Plateau) and one in Italy (Caloveto, Calabria). In addition to 94 corals, ammonites were also collected to improve the age calibrations for this material using the 95 most recent revisions of ammonite determination and biozones (Dommergues and Meister, 96 2017). New specimens collected by Marc Chesnier from Normandy (France) and Philippe Fauré 97 from the French Pyrenees were also included. In addition, the dataset integrates updated coral 98 descriptions from the literature with revised descriptions of the corresponding samples when 99 available (collections of the Muséum d'Histoire Naturelle in Paris, Muséum d'histoire naturelle 100 de Genève). These taxa are described in Vasseur (2019) and in a monograph for Pliensbachian 101 corals (Vasseur and Lathuilière 2020, in press). The dataset thus compiled includes records from 102 Algeria, France, Italy, Morocco, Slovenia, Belgian, Germany, England, Argentina, Indonesia, Spain, Switzerland, Tajikistan. 103

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105 **3. RESULTS AND ANALYSES:**

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107 **3.1. Pliensbachian coral assemblages**

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In the field, Pliensbachian coral faunas are widely dominated in both diversity and quantity by solitary and phaceloid forms, while highly integrated forms (i.e., plocoid, cerioid, thamnasterioid, and meandroid) are rare and often occur as diminutive colonies (Tab. 1, Fig.1). One exception to this general trend is in the Middle-Atlas and Calabria, where large, highintegration corals occur with higher diversity and skeletal volume in quite distal position.

	Solitary			Phaceloid		Plocoid		Cerioid			Thamnasterioid			Meandroid				
	Pli	To 1	To 2	Pli	To 1	To 2	Pli	To 1	To 2	Pli	To 1	To 2	Pli	To 1	To 2	Pli	To 1	To 2
Gu	7	-	3	7	-	6	1	-	1	7	-	7	1	-	3	1	-	5
Am	25	-	1	2	-	3	1	-	1	3	-	8	1	-	-	-	-	3
Da	4	1	2	2	2	2	1	1	-	2	3	-	-	1	-	-	-	-
Zi	-	-	-	1	-	3	-	-	-	-	-	-	1	-	-	-	-	1
Ca	6	-	-	5	-	-	2	-	-	5	-	-	2	-	-	-	-	-
Br	1	-	-	-	-	-	-	-	-	-	-	-	-	-	-	-	-	-
Es	3	-	-	2	-	-	-	-	-	-	-	-	-	-	-	-	-	-
Li	18	-	7	3	-	3	-	-	-	-	-	3	2	-	1	-	-	2

Table 1. Number of species of each morphological type (solitary, phaceloid, plocoid, cerioid,
thamnasterioid, meandroid) in each studied time zone (Pliensbachian [Pli], Toarcian pre-OAE
[To1], and Toarcian post-OAE [To2]) for every studied locality with additional information from
the published literature (Gu: Guigou Plateau/Aïn Leuh (Morocco); Am: Amellagou (Morocco);
Da: Dades Valley (Morocco); Zi: Ziz valley (Morocco); Ca: Calabria (Italy); Br: Brûlon
(France); Es: Estivère Pass (France); Li: literature).



Figure 1. Coral integration types as a percentage of total species among Tethyan coral faunasduring the Pliensbachian, pre-T-OAE Toarcian, and post-T-OAE Toarcian.

126

130 In proximal environments, corals sometimes lived in close relationship with lithiotid bivalves; 131 such reefs and biostromes can be found in the Moroccan High-Atlas Mountains, Spain, Oman, 132 Slovenia, Italy, Switzerland, the Himalayas, and Timor (Brame et al., 2019) but we found no 133 such association in Calabria and the Middle-Atlas Mountains. In the Middle-Atlas Mountains, 134 coral populations were established in giant offshore bioherms detached from any inner platform 135 on a gently sloping tilted block in Guigou Plateau (El Arabi et al., 1987; Vasseur, 2019). In the 136 Longobucco basin (Calabria), reefs are smaller and occur on steep slopes (Santantonio, 2012; 137 Santantonio et al., 2016).

138

Pliensbachian coral populations are composed predominantly of genera with Triassic or Early
Jurassic roots but very few of these genera persist through the T-OAE (Fig. 2). The

Pliensbachian is characterized by a surprising radiation event: 17 genera appeared, including Solitary and 5 phaceloidgenera. Among these 17 genera, 11 persisted after the T-OAE. All genera that are exclusively Pliensbachian are solitary or phaceloid with one exception.



144

Figure 2. Range chart of Tethyan Pliensbachian and Toarcian coral genera. New data are in red; thick line: observed taxon within the stage, dotted line: interpolated occurrence. The colored dots refer to the integration types presented in Fig. 1: red for solitary; green for phaceloid; yellow for plocoid; blue for cerioid; purple for thamnasterioid; grey for meandroid. See supplemental data (Appendix) for taxonomy.

150

151 **3.2. Toarcian coral assemblages**

152

153 No coral reefs have been observed in the earliest Toarcian strata (i.e., before the onset of the T-

154 OAE) and this paucity of sampling localities complicates the assessment of the dominant forms

155 in terms of volume in the assemblage. Nevertheless, the analyses of reworked corals and corals 156 that occur in lithiotid-coral biostromes reveal faunas composed of taxa that evolved during the 157 Early Jurassic mixed with taxa that are classically observed in Middle/Upper Jurassic strata 158 (Fig.2). These observations suggest a transitional period when coral extinctions occurred in 159 association with radiations. During the pre-T-OAE, highly integrated forms are more common, 160 both in terms of diversity and biomass, than they are in the Pliensbachian. These strata contain 161 the last important associations of lithiotid bivalves before their extinction at the onset of the T-162 OAE (Brame et al., 2019, Posenato et al., 2018).

Large reefs have been observed in post-T-OAE intervals (in Dades Valley and Aïn Leuh, Morocco) and in these reefs the diversity and biomass of highly integrated corals are equal to phaceloid forms. Most of the genera that constitute these assemblages persisted and diversified during the Middle/Late Jurassic (Vasseur, 2019). 14 genera survived through the entire Toarcian, including 4 exclusively phaceloid and 3 solitary ones. The evolution of generic richness through the Pliensbachian, Toarcian pre-OAE, and Toarcian post-OAE is presented in figure 3.

169



Figure 3. Evolution of generic richness through the three time intervals studied: Pliensbachian, Toarcian pre-OAE, and Toarcian post-OAE with and without Lazarus taxa (Lazarus taxa correspond to taxa that have not been found in the considered stage but exist in previous and later stages).

176

177 **3.3. Accounting for biases**

178

Analyses have been performed on normalized data as well as the raw data. The normalized data have been weighted either by the duration of stages to calculate the diversity or by the total diversity to calculate the extinction rates, according to the methods used in Lathuilière and Marchal (2009). Stage durations have been calculated using the 2018 international stratigraphic chart and the age of the T-OAE from Krencker et al. (2014) and considering a duration of 450 kyr for the T-OAE (Boulila et al., 2019): Pliensbachian: 190.8 - 182.7 = 8.1 Ma;

186 Toarcian: 182.7 - 174.1 = 8.6 Ma;

187 Toarcian pre-T-OAE: 182.7 - 181.7 = 1 Ma;

188 Toarcian post-T-OAE: 8.6 - 1.45 = 7.15 Ma.

189

190 Pre-T-OAE Toarcian and post-T-OAE Toarcian have been discriminated in the analyses, despite

191 the important difference of sampling quality between these two periods.

192

Sampling biases have also been minimized by removing the following singletons from the analyses: Araiophyllum, Archaeosmiliopsis, Astraeomorpha, Dimorpharaea, Gillismilia, *Microsolena*. In the present generic analysis, singletons correspond to genera including one single species, itself represented by one single specimen (Foote 1994, 2000; Aberhan and Kiessling, 2012). Lazarus taxa correspond to taxa that have not been found in the considered stage but exist in previous and later stages.

199

Weighted analyses have been produced on the data with (1) and without (2) the singletons and also without both singletons and Lazarus taxa (3):

202

	Weighted analysis of coral diversity as a function of stage duration										
	Pliensbachian stage	Toarcian stage	Pre-OAE Toarcian	Post-OAE Toarcian							
Weighted analysis* with singletons	45 / 8.1 = 5.56	39 / 8.6 = 4.53	27 / 1 = 27	36 / 7.15 = 5.03							
Weighted analysis* without singletons	41 / 8.1 = 5.06	37 / 8.6 = 4.3	26 / 1 = 26	34 / 7.15 = 4.76							
Weighted analysis* without singletons or Lazarus taxa	41 / 8.1 = 5.06	33 / 8.6 = 3.84	22 / 1 = 22	30 / 7.15 = 4.2							

*Number of genera in the stage / stage duration



Figure 4. Weighted diversity for the Pliensbachian (Pl), pre-OAE Toarcian and post-OAE
Toarcian calculated with (1) and without (2) singletons, and without both singletons and Lazarus
taxa (3).

208

	Weighted ana	lysis of extinctions as	a function of stage	e duration
	Pliensbachian stage	Toarcian stage	Pre-OAE Toarcian	Post-OAE Toarcian
Weighted extinction rate* with singletons	20 / 18 = 1.11	3 / 14 = 0.21	3/2= 1.5	4 / 12 = 0.33
Weighted extinction rate* without singletons	17 / 18 = 0.94	2 / 12 = 0.17	2/2= 1	4 / 10 = 0.4

*Number of extinctions / number of occurrences



210

Figure 5. Weighted extinction rates for the Pliensbachian (Pl), pre-OAE Toarcian and post-OAE
Toarcian calculated with (1) and without (2) singletons.

The results of the raw data and weighted data are comparable (Fig. 4, 5); in fact, Pliensbachian and Toarcian stages show similar durations with taxa and singletons equally distributed between the two stages.

217

218 **3.4. Family scale analysis**

219

Analyses at the family scale provide good appreciation of disparity *sensu* Gould (1991). Moreover, fossil coral families are primarily defined by the microstructures of skeletal septa that are considered as relevant evolutionary novelties. According to our results (Fig.6), 6 families of Triassic affinity disappear at the Pl-To boundary and 2 families of Middle/Late Jurassic affinity appear just after the T-OAE. In addition, four coral families that are key builders of Middle and Upper Jurassic coral reefs radiate during the earliest Toarcian (i.e., before the OAE): Thecosmiliidae (= ex Montlivaltiidae), Comoseridae (=ex Microsolenidae), Stylinidae, and Thamnasteriidae. Furthermore, this study is the first to document the Deltocyathiidae family in the Mesozoic fossil record; this family was hypothesized to split at the base of the "robust" coral clade in the scleractinian phylogeny (Kitahara et al., 2016).

Coral families	Pl	To pre-OAE	To post-OAE
Astraeomorphidae Frech 1890		pre-OAE	post OTHE
Corvphylliidae Beauvais 1981			
Margarophylliidae Cuif 1977			
Protoheterastraeidae Cuif 1977			
Reimaniphylliidae Melnikova 1975			
Zardinophyllidae Montanaro Gallitelli 1975			
Hispaniastraeidae Boivin et al. 2019			
Archaeosmiliidae Melnikova 1975			
Actinastreidae Alloiteau 1952			
Axosmiliidae Geyer 1955			
Cladophylliidae Morycowa & Roniewicz 1990			
Dermosmiliidae Koby 1889			
Epismiliidae Alloiteau 1952			
Deltocyathidae Kitahara, Cairns, Stolarski and Miller 2012			
Latomeandridae Alloiteau 1952			
Comoseriidae Fromentel 1861			
Procyclolitidae or Conophylliidae			
Stylinidae d'Orbigny 1851			
Stylophyllidae Frech 1890			
Thamnasteriidae Vaughan and Wells 1943			
Thecosmiliidae Duncan 1884			
Aplosmiliidae Fromentel 1861			
Caryophylliidae Dana 1846			

230

231 Figure 6. Worldwide known range chart of coral families that occur within the Pliensbachian and

232 Toarcian stages.

233

4. DISCUSSION:

235

236 The results presented here illustrate the importance of the Pl-To crisis in coral communities; the

- 237 Pliensbachian Toarcian crisis marks the disruption between an early period of high disparity
- 238 (represented by the amount of Pliensbachian families) and the later diversification of surviving
- taxa during the Toarcian and later. Furthermore, the Pliensbachian and post-T-OAE intervals

240 mark the development of coral pioneers in the colonisation of deeper environments, with many 241 specimens recovered from hemipelagic deposits (Deltocyathiidae and Caryophylliidae), a 242 phenomenon that implies the conquest of new ecological niches. Some of the genera that survived the crisis are assumed to be zooxanthellate corals because of their highly-integrated 243 244 morphology, high growth rates, living environments and colony shape variations (e.g., Dimorphastrea, Microphyllia, Thamnasteria - Vasseur, 2019). This is inconsistent with the 245 246 hypothesis that photosymbiosis is lost during crises periods (Rosen, 2000), unless these genera 247 developed an ability to live without symbionts during crisis time. On the other hand, it is 248 consistent with the high correlation between coloniality and photosymbiosis (Simpson, 2013) as 249 all of these genera are colonial.

250

4.1. Extinction Rates

252

253 According to the revised and enhanced scleractinian coral dataset, 23 genera among 45 went 254 extinct during the early Toarcian (Fig. 2), including extinctions at the Pl-To boundary and the T-OAE, which corresponds to an extinction rate of 51.1% at the genus scale, 48.7% without 255 256 singletons (19 genera among 39). Compared to other extinction events, these values are 257 substantial and exceed the revised values of $39 \pm 8\%$ calculated for the Cretaceous/Paleogene 258 (K/Pg) crisis by Kiessling and Baron-Szabo (2004) (evaluated by counting genera that exist just 259 before the extinction but have no record in the Cenozoic). At a finer scale, only 6 of 66 species 260 identified in the Pliensbachian stage survived through the early Toarcian, an extinction rate of 261 90.9%.

262 Another major implication of these results is that Triassic/Jurassic (T/J) coral extinction rates 263 must be recalculated as many Triassic genera thought to go extinct at the T/J were found in the 264 Pliensbachian. Nevertheless, before a T/J extinction rate recalculation can be done, the taxonomy 265 needs revision to minimize taxa overestimation. Kiessling et al. (2007) considered taxa 266 overestimation for the T/J and estimated an extinction rate of 51%-53% at the genus level. All in 267 all, the newly obtained value for PI-To is still significantly above the mean extinction 268 background signal of $12 \pm 4\%$ for scleractinian coral genera calculated by Kiessling and Baron-269 Szabo (2004) from Sepkoski (2002) over the total range of Scleractinia. In summary, our new 270 data highlight the significance of the Pl-To coral crisis and elevate it to the level of other mass 271 extinctions as one of the three most severe scleractinian coral crises ever. Moreover, this event 272 was likely the most severe extinction event for Scleractinia throughout their entire evolutionary 273 history.

274

275 Concerning the issue of "dead clade walking" (i.e., clades surviving the extinction but becoming 276 extinct soon after, Kiessling *et al.*, 2004), 16.67% (4 genera among 24) survived the early 277 Toarcian but went extinct after the T-OAE and are absent in Aalenian strata. This value is higher 278 than the 3% calculated in Sepkoski's compendium (Kiessling et al., 2004) and even higher than 279 the 13% value calculated by Kiessling et al. 2004 for the K/Pg. The subsequent environmental 280 crises in the middle and late Toarcian (Dera et al., 2010; Krencker et al., 2014) might also 281 account for the extinction of these coral genera.

282

4.2. What factors explain evolutionary patterns in the PI-To interval?

285 During the Pliensbachian, corals in proximal environments lived in association with lithiotids 286 (corals have been found within or near lithiotid-bearing strata in Morocco, Spain, Oman, 287 Slovenia, Italy, Switzerland, the Himalayas, and Timor; see Brame et al. 2019), which probably 288 lived in organic-rich, turbid waters due, at least in part, to their filter-feeder activity and own 289 feces production (Brame et al., 2019). Solitary and phaceloid forms were probably able to 290 accommodate a shared ecological niche and cohabitation with lithiotids because of their 291 biological ability to cope with high sediment loads (Klaus et al., 2011; Brame et al., 2019). This 292 niche likely favoured the origination of many solitary forms in proximal, lagoonal environments 293 during the Pliensbachian, whereas more integrated forms would have been living in offshore 294 environments. Important extinctions occurred in coral and lithiotid faunas at the Pl-To boundary 295 and during the T-OAE. The irreversible demise of lithiotid communities at the T-OAE left 296 vacant ecological niches, which encouraged the radiation of many massive coral genera. In 297 Morocco, both the Pl/To boundary and T-OAE coral extinctions are associated with neritic 298 carbonate productivity shutdown (Krencker et al. 2020), likely associated with ocean 299 acidification events (Müller et al., 2020). They also correspond to periods of increased delivery 300 of coarse siliciclastic material, highlighting the intimate link between carbonate factory crisis and 301 increased continental weathering correlated with climate changes (Krencker et al. 2020).

302

303 **5. CONCLUSION**

304

This study highlights the significance of both the Pl/To boundary and the T-OAE crises in coral evolution. These events resulted in extinction rates as high as 51.1% at the generic level and 90.9% at the species level, and combined (i.e., the Pl-To interval) rank as one of the three most 308 important extinction and recovery periods in the geological history of Scleractinia, on par with 309 (or even more devastating than) the T/J and K/Pg extinctions. Although it is not considered a 310 first order biotic crisis, the second-order Pl-To extinction is as significant for corals as a first-311 order biotic crisis. The extinction resulted in a loss of coral diversity and disparity, as well as a 312 significant reef collapse. This crisis shaped the future of diversity and morphological disparity 313 within the coral clades that flourished during the Jurassic and Cretaceous. Consequently, it must 314 be considered a key period in the evolutionary dynamics of Scleractinia. Given the significance 315 of the Pl-To crisis to coral and reef communities, their survival and recovery during this event 316 are particularly pertinent for questions about present day ecosystem collapse and conservation of 317 coral habitats. Unique characteristics of this extinction, such as the survival of zooxanthellate 318 coral, may prove important for understanding current coral bleaching and conservation.

319

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340 **REFERENCES CITED**

341

Aberhan, M.,Kiessling, W., 2012. Phanerozoic marine biodiversity: a fresh look at data,
methods, patterns and processes.Earth and life (pp. 3-22). Springer, Dordrecht.
https://doi.org/10.1007/978-90-481-3428-1_1

345

Aurell, M., Robles, S., Bádenas, B., Rosales, I., Quesada, S., Meléndez, G., Garcia-Ramos, J. C.,
2003. Transgressive – regressive cycles and Jurassic palaeogeography of northeast Iberia.
Sediment. Geol., 162(3-4), 239-271.

349

Bambach, R. K., Knoll, A. H., Wang, S. C., 2004. Origination, extinction, and mass depletions of
marine diversity. Paleobiology, 30(4), 522-542.

353	Beauvais,	L.,	1986.	Monographie	des	madréporaires	du	Jurassique	inférieur	du	Maroc.
354	Palaeontog	gr. A	bt., A 1	94(1/3):1-68.							

- 355
- Benton, M. J., 1995. Diversification and extinction in the history of life. Science, 268(5207), 5258.
- 358
- 359 Blomeier, D.G., Reijmer, J.G., 1999. Drowning of a Lower Jurassic carbonate platform: Jbel
- Bou Dahar, High Atlas, Morocco. Facies, 41, 81–110.
- 361
- Bodin, S., Mattioli, E., Fröhlich, S., Marshall, J. D., Boutib, L., Lahsini, S., Redfern, J., 2010.
 Toarcian carbon isotope shifts and nutrient changes from the Northern margin of Gondwana
 (High Atlas, Morocco, Jurassic): palaeoenvironmental implications. Palaeogeogr.
 Palaeoclimatol. Palaeoecol., 297(2), 377-390.
- 366
- Bodin, S., Krencker, F.-N., Kothe, T., Hoffmann, R., Mattioli, E., Heimhofer, U., Kabiri, L.,
 2016. Perturbation of the carbon cycle during the late Pliensbachian Early Toarcian: New
 insight from high-resolution carbon isotope records in Morocco. J. African Earth Sci., vol. 116:
 89-104. https://doi.org/10.1016/j.jafrearsci.2015.12.018
- 371
- Boulila, S., Galbrun, B., Sadki, D., Gardin, S., Bartolini, A., 2019. Constraints on the duration of
 the early Toarcian T-OAE and evidence for carbon-reservoir change from the High Atlas
 (Morocco). Glob. Planet. Chang., 175, 113-128.
- 375

- Brame, H. M. R., Martindale, R. C., Ettinger, N. P., Debeljak, I., Vasseur, R., Lathuilière, B.,
 Kabiri, L., Bodin, S., 2019. Stratigraphic distribution and paleoecological significance of Early
 Jurassic (Pliensbachian-Toarcian) lithiotid-coral reefal deposits from the Central High Atlas of
 Morocco. Palaeogeogr. Palaeoclimatol. Palaeoecol., 514, 813-837.
- 380 https://doi.org/10.1016/j.palaeo.2018.09.001
- 381
- Brigaud, B., Vincent, B., Carpentier, C., Robin, C., Guillocheau, F., Yven, B., Huret, E., 2014.
 Growth and demise of the Jurassic carbonate platform in the intracratonic Paris Basin (France):
 interplay of climate change, eustasy and tectonics. Mar. Pet. Geol., 53, 3-29.
- 385
- Burgess, S.D., Bowring, S.A., Fleming, T.H., Elliot, D.H., 2015. High-precision geochronology
 links the Ferrar large igneous province with Early-Jurassic ocean anoxia and biotic crisis. Earth
 Planet. Sci. Lett., v. 415, p. 90–99. https://doi.org/10.1016/j.epsl.2015.01.037.
- 389
- 390 Caruthers, A.H., Gröcke, D.R., Smith, P.L., 2011. The significance of an Early Jurassic 391 (Toarcian) carbon-isotope excursion in Haida Gwaii (Queen Charlotte Islands), British 392 Columbia, Canada. Earth Planet. Sci. Lett., 307. 19–26. v. p. 393 https://doi.org/10.1016/j.epsl.2011.04.013.
- 394
- Caruthers, A. H., Smith, P. L., Gröcke, D. R., 2013. The Pliensbachian–Toarcian (Early Jurassic)
 extinction, a global multi-phased event. Palaeogeogr. Palaeoclimatol. Palaeoecol., 386:104-118.
 https://doi.org/10.1016/j.palaeo.2013.05.010
- 398

- Ceballos, G., Ehrlich, P.R., Barnosky, A.D., García, A., Pringle, R.M., Palmer, T.M., 2015.
 Accelerated modern human-induced species losses: Entering the sixth mass extinction. Sci.
 Adv., 1(5), e1400253.
- 402
- 403 Dera, G., Neige, P., Dommergues, J.-L., Fara, E., Laffont, R.,Pellenard, P., 2010. High404 resolution dynamics of Early Jurassic marine extinctions: the case of Pliensbachian–Toarcian
 405 ammonites (Cephalopoda). J.Geol. Soc. London, vol. 167:21–33. https://doi.org/10.1144/0016406 76492009-068
- 407
- 408 Dera, G., Brigaud, B., Monna, F., Laffont, R., Pucéat, E., Deconinck, J.-F., Pellenard, P.,
 409 Joachimski, M.M., Durlet, C., 2011. Climatic ups and downs in a disturbed Jurassic world.
 410 Geology, 39, 215 218.
- 411
- 412 Dommergues, J.-L., Meister, C., 2017. Ammonites du Jurassique inférieur (Hettangien,
- 413 Sinémurien, Pliensbachien) d'Afrique du Nord (Algérie, Maroc et Tunisie). Atlas
- 414 d'identification des espèces. Rev. de Paléobiologie, 36(2):189-367.
- 415
- El Arabi, H., Canerot, J., Charrière, A., 1987, Dynamique récifale et tectonique distensive en
 blocs basculés: l'exemple des récifs liasiques du bloc du Guigou (Causse moyen atlasiqueMaroc). Mem. Soc. Phys. Hist. nat. Geneve, 259-271.
- 419

- Ettinger, N. P., Larson, T. E., Kerans, C., Thibodeau, A. M., Hattori, K. E., Kacur, S. M.,
 Martindale, R. C., 2021. Ocean acidification and photic-zone anoxia at the Toarcian Oceanic
 Anoxic Event: Insights from the Adriatic Carbonate Platform. Sedimentology, 68(1), 63-107.
- Flügel, E., Kiessling, W., 2002, Patterns of Phanerozoic reef crises. SEPM Special Publication,
 72, 691-734.

Foote, M., 1994. Temporal variation in extinction risk and temporal scaling of extinction metrics.
Paleobiology, 20, 424-444.

429

Foote M., 2000, Origination and extinction components of taxonomic diversity: general
problems. Paleobiology 26:74–102.

432

- Gómez, J.J., Goy, A., 2011. Warming-driven mass extinction in the Early Toarcian (Early
 Jurassic) of northern and central Spain. Correlation with other time-equivalent European
 sections. Palaeogeogr. Palaeoclimatol. Palaeoecol., v. 306, p. 176–195.
 https://doi.org/10.1016/j.palaeo.2011.04.018.
- 437
- Gómez, J.J., Comas-Rengifo, M.J., Goy, A., 2016. Palaeoclimatic oscillations in the
 Pliensbachian (Early Jurassic) of the Asturian Basin (Northern Spain). Clim. Past., 12, 1199–
 1214. https://doi.org/10.5194/cp-12-1199-2016.

443	analysis: why we must strive to quantify morphospace. Paleobiology, 17(4), 411-423.
444	
445	Graciansky, P. C., Jacquin, T., Hesselbo, S. P., 1998. The Ligurian cycle: an overview of Lower
446	Jurassic 2nd-order transgressive/regressive facies cycles in Western Europe. Society for
447	Sediment. Geol., 60.
448	
449	Kiessling, W., Baron-Szabo, R. C., 2004. Extinction and recovery patterns of scleractinian corals
450	at the Cretaceous-Tertiary boundary. Palaeogeogr. Palaeoclimatol. Palaeoecol., 214(3), 195-223.
451	
452	Kiessling, W., Aberhan, M., Brenneis, B., Wagner, P.J., 2007. Extinction trajectories of benthic
453	organisms across the Triassic-Jurassic boundary. Palaeogeogr. Palaeoclimatol. Palaeoecol.,, 244,

Gould, S. J., 1991. The disparity of the Burgess Shale arthropod fauna and the limits of cladistic

454 455 201.

442

Kiessling, W., 2011. Patterns and Processes of Ancient Reef Crises. The Paleontological Society
Papers, 17, 1-14. https://doi.org/10.1017/S1089332600002412

458

459 Kiessling, W., Simpson, C., 2011. On the potential for ocean acidification to be a general cause
460 of ancient reef crises: Ancient Reef Crises. Glob. Change Biol., 17, 56–67,
461 https://doi.org/10.1111/j.1365-2486.2010.02204.x.

- Kitahara, M. V., Fukami, H., Benzoni, F., Huang, D., 2016. The new systematics of Scleractinia:
 integrating molecular and morphological evidence. The Cnidaria, past, present and future, pp.
 465 41-59. Springer, Cham.
- 466
- Klaus, J. S., Lutz, B. P., McNeill, D. F., Budd, A. F., Johnson, K. G., Ishman, S. E., 2011. Rise
 and fall of Pliocene free-living corals in the Caribbean. Geology, 39(4), 375-378.
- 469
- 470 Korte, C., Hesselbo, S. P., Ullmann, C. V., Dietl, G., Ruhl, M., Schweigert, G., Thibault, N.,
- 471 2015. Jurassic climate mode governed by ocean gateway. Nat. Commun., 6, 10015.
- 472
- Krencker, F.-N., Bodin, S., Hoffmann, R., Suan, G., Mattioli, E., Kabiri, L., Föllmi, K.B.,
 Immenhauser, A., 2014. The middle Toarcian cold snap: trigger of mass extinction and carbonate
 factory demise. Glob. Planet. Chang., 117, 64-78.
- 476
- Krencker, F. N., Bodin, S., Suan, G., Heimhofer, U., Kabiri, L., Immenhauser, A., 2015.
 Toarcian extreme warmth led to tropical cyclone intensification. Earth Planet. Sci. Lett., 425,
 120-130. https://doi.org/10.1016/j.epsl.2015.06.003
- 480
- 481 Krencker, F. N., Lindström, S., Bodin, S., 2019.A major sea-level drop briefly precedes the
 482 Toarcian oceanic anoxic event: implication for Early Jurassic climate and carbon cycle. Sci.
 483 Rep., 9(1), 1-12.
- 484

- Krencker, F. N., Fantasia, A., Danisch, J., Martindale, R., Kabiri, L., El Ouali, M., Bodin, S.,
 2020. Two-phased collapse of the shallow-water carbonate factory during the late Pliensbachian–
 Toarcian driven by changing climate and enhanced continental weathering in the Northwestern
 Gondwana Margin. Earth-Sci. Rev., 103254.
- 489
- 490 Lathuilière, B., Marchal, D., 2009. Extinction, survival and recovery of corals from the Triassic
 491 to Middle Jurassic Time. Terra Nova, 2, 57-66. https://doi.org/10.1111/j.1365492 3121.2008.00856.x
- 493
- 494 Léonide, P., Floquet, M., Durlet, C., Baudin, F., Pittet, B., Lécuyer, C., 2012. Drowning of a
 495 carbonate platform as a precursor stage of the Early Toarcian global anoxic event (Southern
 496 Provence sub-Basin, South-east France). Sedimentology, 59, 156-184.
- 497
- 498 Little, C. T., Benton, M. J., 1995. Early Jurassic mass extinction: a global long-term event.
 499 Geology, 23(6), 495-498.
- 500
- Mattioli, E., Pittet, B., Petitpierre, L., Mailliot, S., 2009. Dramatic decrease of pelagic carbonate
 production by nannoplankton across the Early Toarcian anoxic event (T-OAE). Glob. Planet.
 Chang., 65(3-4), 134-145.
- 504
- 505 Morard, A., Guex, J., Bartolini, A., Morettini, E., De Wever, P., 2003. A new scenario for the 506 Domerian _ Toarcian transition. Bull. Soc. Géol. Fr., 174. 4, 351-356. https://doi.org/10.2113/174.4.351 507

509	Müller, T., Jurikova, H., Gutjahr, M., Tomašových, A., Schlögl, J., Liebetrau, V., Duarte, L.V.,							
510	Milovský, R., Suan, G., Mattioli, E., Pittet, B., Eisenhauer, A., 2020. Ocean acidification during							
511	the early Toarcian extinction event: Evidence from boron isotopes in brachiopods. Geology, 48,							
512	1184–1188.							
513								
514	Pittet, B., Suan, G., Lenoir, F., Duarte, L. V., Mattioli, E., 2014. Carbon isotope evidence for							
515	sedimentary discontinuities in the lower Toarcian of the Lusitanian Basin (Portugal): Sea level							
516	change at the onset of the Oceanic Anoxic Event. Sediment. Geol., 303, 1-14.							
517								
518	Posenato, R., Bassi, D., Trecalli, A., Parente, M., 2018. Taphonomy and evolution of Lower							
519	Jurassic lithiotid bivalve accumulations in the Apennine Carbonate Platform (southern Italy).							
520	Palaeogeogr.Palaeoclimatol.Palaeoecol.,489,261-271.							
521	https://doi.org/10.1016/j.palaeo.2017.10.017							
522								
523	Raup, D. M., Sepkoski, J. J., 1982. Mass extinctions in the marine fossil record. Science,							
524	215(4539), 1501-1503.							
525								
526	Reolid, M., Mattioli, E., Duarte, L. V., Marok, A., 2020. The Toarcian Oceanic Anoxic Event							
527	and the Jenkyns Event (IGCP-655 final report). Episodes 43(2), 833-844.							
528								
529	Rosen, B. R., 2000. Algal symbiosis, and the collapse and recovery of reef communities: Lazarus							
530	corals across the KT boundary. Biotic response to global change, 164-180.							

532	Ruban, D. A., 2004. Diversity dynamics of Early-Middle Jurassic brachiopods of Caucasus, and
533	the Pliensbachian-Toarcian mass extinction. Acta Palaeontol. Pol., 49(2).
534	
535	Ruebsam, W., Mayer, B., Schwark, L., 2019. Cryosphere carbon dynamics control early Toarcian
536	global warming and sea level evolution. Glob. Planet. Chang., 172, 440-453.
537	
538	Santantonio, M., 2012. Mesozoic evolution of the Caloveto Group, NE Calabria. Rendiconti
539	Online Soc. Geol. Ital., 21, 295-297.
540	
541	Santantonio, M., Fabbi, S., Aldega, L., 2016. Mesozoic architecture of a tract of the European-
542	Iberian continental margin: Insights from preserved submarine palaeotopography in the
543	Longobucco Basin (Calabria, Southern Italy). Sediment. Geol., 331, 94-113.
544	https://doi.org/10.1016/j.sedgeo.2015.10.010
545	
546	Sepkoski, J. J., 2002. A compendium of fossil marine animal genera. Bull. Am. Paleontol., 363,
547	1-563.
548	
549	Simpson, C., 2013. Species selection and the macroevolution of coral coloniality and
550	photosymbiosis. Evolution, 67(6), 1607-1621.
551	

- Slater, S.M., Twitchett, R.J., Danise, S., Vajda, V., 2019. Substantial vegetation response to
 Early Jurassic global warming with impacts on oceanic anoxia. Nat. Geosci., v. 12, p. 462–467.
 https://doi.org/10.1038/s41561-019-0349-z.
- 555
- Suan, G., Mattioli, E., Pittet, B., Lécuyer, C., Suchéras-Marx, B., Duarte, L. V., Philippe, M.,
 Reggiani, L., Martineau, F., 2010. Secular environmental precursors to Early Toarcian (Jurassic)
 extreme climate changes. Earth Planet. Sci. Lett., vol. 290, n° 3-4: 448-458.
 https://doi.org/10.1016/j.epsl.2009.12.047
- 560
- Trecalli, A., Spangenberg, J., Adatte, T., Föllmi, K. B., Parente, M., 2012. Carbonate platform
 evidence of ocean acidification at the onset of the early Toarcian oceanic anoxic event. Earth
 Planet. Sci. Lett., vol. 357-358: 214-225. https://doi.org/10.1016/j.epsl.2012.09.043
- 564
- 565 Vasseur, R., 2019. Extinctions et recouvrements de coraux au cours de la crise Pliensbachien –
 566 Toarcien. Nancy, Université de Lorraine, 542 pp.

- Vasseur, R., Lathuilière, B., 2020. Pliensbachian corals from the Western Tethys. Geodiversitas,
 in press.
- 570

571 FIGURE CAPTIONS

572

Table 1. Number of species of each morphological type (solitary, phaceloid, plocoid, cerioid,
thamnasterioid, meandroid) in each studied time zone (Pliensbachian [Pli], Toarcian pre-OAE

[To 1], and Toarcian post-OAE [To 2]) for every studied locality with additional information
from the published literature (Gu: Guigou Plateau/Aïn Leuh (Morocco); Am: Amellagou
(Morocco); Da: Dades Valley (Morocco); Zi: Ziz valley (Morocco); Ca: Calabria (Italy); Br:
Brûlon (France); Es: Estivère Pass (France); Li: literature).

579

Figure 1. Coral integration types as a percentage of total species among Tethyan coral faunasduring the Pliensbachian, pre-T-OAE Toarcian, and post-T-OAE Toarcian.

582

Figure 2. Range chart of Tethyan Pliensbachian and Toarcian coral genera. New data are in red; thick line: observed taxon within the stage, dotted line: interpolated occurrence. The colored dots refer to the integration types presented in Fig. 1: red for solitary; green for phaceloid; yellow for plocoid; blue for cerioid; purple for thamnasterioid; grey for meandroid. See supplemental data (Appendix) for taxonomy.

588

589 Figure 3. Evolution of generic richness through the three time intervals studied: Pliensbachian, 590 Toarcian pre-OAE, and Toarcian post-OAE with and without Lazarus taxa (Lazarus taxa 591 correspond to taxa that have not been found in the considered stage but exist in previous and 592 later stages).

593

Figure 4. Weighted diversity for the Pliensbachian (Pl), pre-OAE Toarcian and post-OAE
Toarcian calculated with (1) and without (2) singletons, and without both singletons and Lazarus
taxa (3).

- 598 Figure 5. Weighted extinction rates for the Pliensbachian (Pl), pre-OAE Toarcian and post-OAE
- 599 Toarcian calculated with (1) and without (2) singletons.
- 600
- 601 Figure 6. Worldwide known range chart of coral families that occur within the Pliensbachian and
- 602 Toarcian stages.