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## Locating rockfalls using inter-station ratios of seismic energy at Dolomieu crater, Piton de la Fournaise volcano

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### ► To cite this version:

Julian Kuehnert, Anne Mangeney, Yann Capdeville, Jean-Pierre Vilotte, Eleonore Stutzmann, et al.. Locating rockfalls using inter-station ratios of seismic energy at Dolomieu crater, Piton de la Fournaise volcano. *Journal of Geophysical Research: Earth Surface*, 2021, 126 (4), 10.1002/essoar.10503303.3 . hal-03411343

**HAL Id: hal-03411343**

**<https://hal.science/hal-03411343v1>**

Submitted on 2 Nov 2021

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1           **Locating rockfalls using inter-station ratios of seismic**  
2           **energy at Dolomieu crater, Piton de la Fournaise**  
3           **volcano**

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13           **Key Points:**

- 14           • Rockfalls are located using generated seismic signals at high frequencies for highly  
15           resolved spatial and temporal tracking  
16           • Rockfall location is improved using the signature of surface topography on seis-  
17           mic signals simulated with the 3D Spectral Element Method  
18           • By accounting for topography, all signal components can be used, critical in the  
19           case of sparse station networks or noise

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## Abstract

Rockfalls generate seismic signals that can be used to detect and monitor rockfall activity. Event locations can be estimated on the basis of arrival times, amplitudes or polarization of these seismic signals. However, surface topography variations can significantly influence seismic wave propagation and hence compromise results. Here, we specifically use the signature of topography on the seismic signal to better constrain the source location. Seismic impulse responses are predicted using Spectral Element based simulation of 3D wave propagation in realistic geological media. Subsequently, rockfalls are located by minimizing the misfit between simulated and observed inter-station energy ratios. The method is tested on rockfalls at Dolomieu crater, Piton de la Fournaise volcano, Reunion Island. Both single boulder impacts and distributed granular flows are successfully located, tracking the complete rockfall trajectories by analyzing the signals in sliding time windows. Results from the highest frequency band (here 13-17 Hz) yield the best spatial resolution, making it possible to distinguish detachment positions less than 100 m apart. By taking into account surface topography, both vertical and horizontal signal components can be used. Limitations and the noise robustness of the location method are assessed using synthetic signals. Precise representation of the topography controls the location resolution, which is not significantly affected by the assumed impact direction. Tests on the network geometry reveal best resolution when the seismometers triangulate the source. We conclude that this method can improve the monitoring of rockfall activity in real time once a simulated database for the region of interest is created.

## 1 Introduction

Seismology is increasingly used to study and monitor dynamic processes at the interface between the Earth and its fluid envelopes, a field often more specifically referred as environmental seismology (Larose et al., 2015; K. E. Allstadt et al., 2018). Surface processes can include natural phenomena such as storms (e.g. Ebeling & Stein, 2011; Stutzmann et al., 2012), glaciers (e.g. Tsai et al., 2008; Podolskiy & Walter, 2016; Sergeant et al., 2016, 2019), rivers (e.g. Gimbert et al., 2014), debris flow (e.g. Burtin et al., 2009), snow avalanches (e.g. Norris, 1994; Leprettre & Navarre, 1998; Suriñach et al., 2000, 2001) as well as landslides and rockfalls (e.g. Suriñach et al., 2005; Favreau et al., 2010; Hibert et al., 2011; K. Allstadt, 2013; Bottelin et al., 2014; Vouillamoz et al., 2018).

In the context of landslides (used here as the most general term for gravitational mass movements), seismic signals can be used to identify hazards. Growing networks of seismic stations offer the opportunity to continuously monitor large regions of interest. Landslide events can be detected, characterized, and located using the seismic signals they generate (e.g. Suriñach et al., 2005; Hibert et al., 2014; Provost et al., 2017; E. J. Lee et al., 2019). This helps in creating catalogs of landslides that allow statistical analysis of their spatial and temporal activity and estimation of their probability of occurrence. In this way, triggering mechanisms can be studied by correlating landslide catalogs with meteorological data (Burtin et al., 2009; Helmstetter & Garambois, 2010; Durand et al., 2018) or with volcanic seismicity data (Hibert, Mangeney, et al., 2017; Durand et al., 2018). On volcanoes, rockfall locations can provide insight into volcano summit deformation (Durand et al., 2018), and seismic signals are also used to monitor other processes such as lahars (e.g. Zobin et al., 2009; Zobin, 2012; Vázquez et al., 2016; Coviello et al., 2018) as well as magma migration (e.g. Taisne et al., 2011; Lengliné et al., 2016; Duputel et al., 2019).

Several methods for locating landslides from seismic signals have been proposed and can be divided into two main groups. In the first group, the source location is inferred geometrically by pointing to it from several stations and determining the intersection. This can be done by polarization analysis with three-component seismometers (Vilajosana et al., 2008) or by array analysis methods that estimate the apparent slowness vector (Almendros et al., 2002). In the second group, seismic signal properties are back-projected, optimizing correlation between multiple stations. The back-projection

73 relies either on the decay of amplitudes with distance, using methods such as amplitude  
 74 source location (ASL, e.g. Battaglia & Aki, 2003; Battaglia et al., 2005; Walter et al.,  
 75 2017; Morioka et al., 2017; Walsh et al., 2017; Pérez-Guillén et al., 2019; Walsh et al.,  
 76 2020) and seismic intensity ratios (e.g. Taisne et al., 2011), or on travel time differences  
 77 between stations pairs, using cross-correlation of signal envelopes (Burtin et al., 2009;  
 78 Lacroix & Helmstetter, 2011; Yamada et al., 2012; Bottelin et al., 2014; Dietze et al., 2017)  
 79 or inversion of first arrival times (Hibert et al., 2014; Gracchi et al., 2017; Fuchs et al.,  
 80 2018). Li et al. (2020) reviews recent advances of back-projection methods to locate seis-  
 81 mic sources, including wavefield migration, waveform inversion, semblance and template  
 82 matching.

83 As landslides predominantly occur in mountainous regions, generated seismic waves  
 84 are prone to interact with rough surface topography variations. The influence of topog-  
 85 raphy on seismic wave propagation has long been a subject of study (Geli et al., 1988).  
 86 Topography can affect the wave path, wave polarization (e.g. Ripperger et al., 2003; Méta-  
 87 xian et al., 2009) and seismic amplitudes (e.g. S. J. Lee et al., 2009; Maufroy et al., 2015).  
 88 If not taken into account correctly, topographic effects compromise location methods and  
 89 decrease their accuracy.

90 Assuming elongated wave paths along the topography, back-projection methods can  
 91 take topography into account by adjusting source-receiver distances and thus travel times.  
 92 This was done for example by Hibert et al. (2014) to locate rockfalls at Dolomieu crater,  
 93 Reunion Island, and by Levy et al. (2015) to locate granular flows at Soufrière Hills vol-  
 94 cano, Montserrat. However, adjusting the source-receiver distance does not account for  
 95 diffraction or scattering during the propagation of the seismic wave along the topogra-  
 96 phy.

97 In the following we propose a new location method that accounts for the cumula-  
 98 tive effect of the topography on the recorded signal. The method is based on the work  
 99 of Kuehnert et al. (2020), in which topography effects on the seismic wave field were in-  
 100 vestigated using the 3D Spectral Element Method (SEM, e.g. Komatitsch & Vilotte, 1998;  
 101 Chaljub et al., 2007) in combination with a realistic geological domain. By calculating  
 102 seismic energy ratios between stations pairs and hence removing the signature of the seis-  
 103 mic source, they concluded that observed energy ratios from recorded rockfall signals can  
 104 be reproduced when topography is considered in the simulations and site effects are re-  
 105 moved from the observations. This is used here for locating seismic sources by construct-  
 106 ing a database of simulated energy ratios from a grid of potential source positions with  
 107 10 m spacing which are then compared to the observed energy ratios after site effect cor-  
 108 rection using spectral amplification functions estimated from volcano-tectonic (VT) events.

109 The method is tested on seismic signals generated by rockfalls at Piton de la Four-  
 110 naise volcano, Reunion Island. After analyzing one rockfall in detail to tune the method  
 111 for best resolution, a variety of diverse rockfall events are located. As the method as-  
 112 sumes single sources, its performance for largely distributed sources such as granular flows  
 113 is evaluated. Finally, to investigate the limitations of the method, synthetic rockfall sig-  
 114 nals are constructed from single as well as multiple source positions. A resolution proxy  
 115 is defined to test the station coverage and identify network geometries with enhanced  
 116 resolution. Furthermore, the sensitivity of the locating method to the ambient noise level  
 117 as well as to the underlying model assumptions such as the topography resolution and  
 118 the source impact direction is assessed.

## 119 2 Rockfall seismic signals at Dolomieu crater, Reunion Island

120 The study site is located on Piton de la Fournaise volcano, Reunion Island, shown  
 121 in Figure 1a. Rockfalls occur frequently on the unstable flanks of Dolomieu crater, which  
 122 was formed during the caldera collapse in 2007 (Michon et al., 2007). The volcano is mon-  
 123 itored by the *Observatoire Volcanologique du Piton de la Fournaise* (OVPF). The instru-  
 124 mentation, which includes both seismic stations and cameras, allows rockfall analysis by  
 125 correlating recorded seismic signals with video recordings. For the present study, four

126 stations around the Dolomieu crater with a sampling frequency of 100 Hz are used, namely  
 127 the three-component stations BON, BOR, and SNE and the vertical component station  
 128 DSO. BON and SNE are broadband (i.e. corner frequency  $\leq 0.1$  Hz), BOR and DSO are  
 129 short-period (i.e. corner frequency  $> 0.1$  Hz) stations. The three cameras CBOC, DOEC,  
 130 and SFRC are located on the summit of Piton de la Fournaise, look into the Dolomieu  
 131 crater and continuously sample two images per second. The supporting information of  
 132 this article provides the seismograms and the videos of all analyzed rockfalls. To eval-  
 133 uate the results of the present location method, rockfall trajectories are manually esti-  
 134 mated from the videos by determining landmarks visible on both the videos and the avail-  
 135 able Digital Elevation Model (DEM) of 10 m resolution. This way, the trajectory of a  
 136 rockfall on December 13, 2016, is reconstructed and marked in red in Figure 1a. An un-  
 137 certainty of  $\pm 50$  m is assumed.

138 The recorded ground velocity generated by this rockfall on the southwestern crater  
 139 wall is shown in Figure 1b. The most abrupt signals are observed at BOR and DSO, the  
 140 two closest stations, whereas the signals at stations BON and SNE at larger distances  
 141 slowly rise in amplitude. The temporal evolution of the recorded signals can be charac-  
 142 terized by a proxy of the seismic energy  $E_{ij}$  measured at each station  $i$  and component  
 143  $j$  that we define as the square of the recorded ground velocity  $v_{ij}^2(t)$ , integrated over a  
 144 sliding window of width  $d$ :

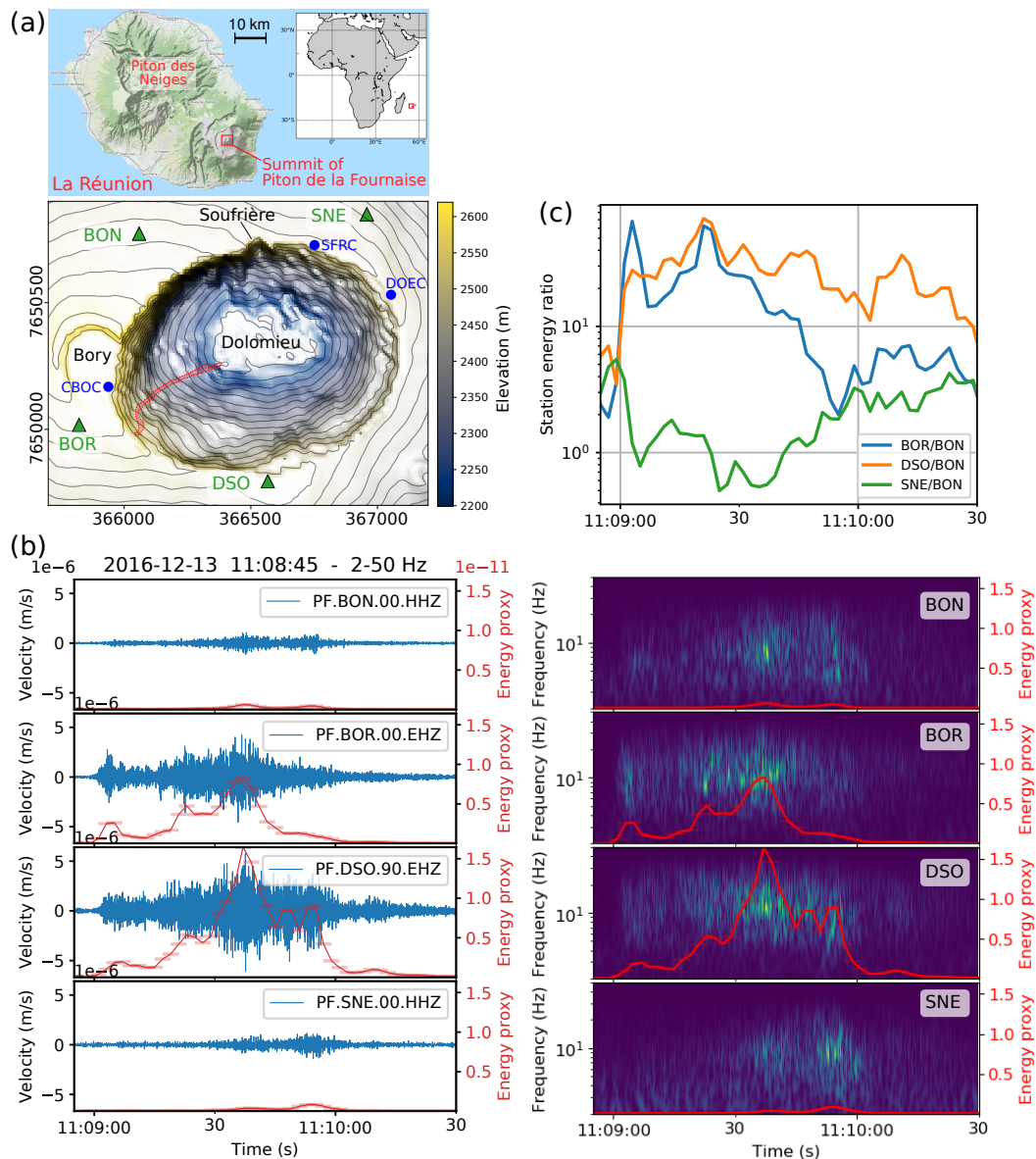
$$E_{ij} = \int_d v_{ij}^2(t) dt. \quad (1)$$

145 Energy proxies  $E_{iz}$ , calculated from vertical component  $j = z$ , are shown in Figure 1b.  
 146 Their inter-station ratios are shown in Figure 1c, where BON is chosen as the reference  
 147 station. The beginning of the seismic signal generated by the rockfall is marked by an  
 148 abrupt increase of the ratios BOR/BON and DSO/BON, whereas the ratio SNE/BON  
 149 decreases. Subsequently, the ratios evolve differently as the rockfall moves towards the  
 150 bottom of the crater.

151 As the seismic source is identical for all stations, the temporal evolution of energy  
 152 ratios is caused by the wave propagation path. First of all, as the source position moves,  
 153 the source-receiver distances vary, which modifies the amplitude of the signals because  
 154 of geometric spreading and attenuation. Moreover, soil heterogeneities and topography  
 155 can affect the wave propagation between the source and the receiver. By modeling the  
 156 influence of the topography on the energy ratios through direct numerical wave simu-  
 157 lation and by taking into account local heterogeneities using empirical site amplification  
 158 factors, the present study aims to locate rockfalls with high spatial and temporal res-  
 159 olutions.

### 160 3 Methodology

161 The proposed methodology for estimating rockfall locations uses energy ratios be-  
 162 tween stations to predict source positions. The ratios characterize the path effects, while  
 163 the energy of the source itself can be ignored. This was used for example by Taisne et  
 164 al. (2011) to map magma propagation. Here we use this strategy to compare the observed  
 165 energy ratios with simulated ones, which was shown by Kuehnert et al. (2020) to be pos-  
 166 sible considering the topography. Instead of using spectral ratios at single frequencies,  
 167 we average the energy ratios within frequency bands of 4 Hz. This makes the method more  
 168 robust to fluctuations in spectral values. In order to explore all potential rockfall sources,  
 169 reciprocal simulations are carried out, where the synthetic source is placed at the po-  
 170 sition of the seismometer and the wave field is recorded over the entire crater area. A  
 171 grid search is then performed to find the source positions that best fit the observed en-  
 172 ergy ratios.



**Figure 1.** (a) Map of Reunion Island (top) and hillshaded elevation map of Dolomieu crater on Piton de la Fournaise volcano (bottom). The smaller Bory and Soufrière craters are located west and north of Dolomieu crater, respectively. Seismic stations are marked by green triangles and cameras by blue dots. The red zone marks a rockfall trajectory estimated from the video. (b) Seismograms (left) and corresponding spectrograms (right) of vertical velocity generated by a rockfall on December 13, 2016, corresponding to the red trajectory indicated in a). The signals were recorded at the four seismic stations surrounding Dolomieu crater. The red curves are the seismic energy proxy  $E_{iz}$  (according to Eq. 1), calculated from a sliding time window in steps of 2s and of width  $d = 4$ s. Ambient seismic noise can only be detected in the spectrogram at furthest station SNE below 3 Hz. (c) Inter-station energy ratios from vertical ground velocity. The beginning of the seismic signal emitted by the rockfall is marked by an abrupt change of the ratios.

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### 3.1 SEM simulations

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The propagation of the seismic wave field is simulated using the 3D Spectral Element Method (SEM, e.g. Komatitsch & Vilotte, 1998; Chaljub et al., 2007) in a numerical domain of dimensions  $x = 2100$  m (easting),  $y = 1800$  m (northing), and  $z = 600$  m (depth) as shown in Figure 2a, identical to the domain of the simulations presented by Kuehnert et al. (2020).

179

The domain is meshed in the top 150 m with hexahedral elements of 10 m side length to correctly accommodate the surface topography of Dolomieu crater provided by a 2009 Digital Elevation Model (DEM) of 10 m resolution. Further below in depth, the element size is increased to 30 m to reduce computational costs. A *Zone of refinement* connects the two different element sizes, while a smooth *Buffer layer* filters out short wavelength variations of the fine mesh that cannot be represented in the coarse mesh. A polynomial order of 5 is used in all elements. To simulate an unbounded domain, absorbing PMLs (Perfectly Matched Layers, e.g. Festa & Vilotte, 2005) of 160 m thickness are attached on the sides and the bottom of the domain.

188

The subsurface is parametrized using the generic velocity model proposed by Lesage et al. (2018) for the shallow structure of volcanoes. It is characterized by a velocity profile gradually increasing with depth as illustrated in Figure 2b. It is implemented on the 3D domain by laterally following the surface topography (i.e. each point at the surface is defined by depth  $z = 0$  m), deforming the horizontal layers of equal velocity. Kuehnert et al. (2020) validated that this velocity model represents reasonably well the present study site at Piton de la Fournaise volcano by comparing simulated and recorded seismic signals of different rockfalls. The rock density is set to  $\rho = 2000 \text{ kg m}^{-3}$  and quality factors are set to  $Q_P = 80$  and  $Q_S = 50$  for P- and S-wave velocity, respectively. These values are based on previous studies on Piton de la Fournaise and similar volcanoes (Battaglia & Aki, 2003; O'Brien & Bean, 2009; Hibert et al., 2011).

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Seismic sources are represented by a point force and a Ricker wavelet of 7 Hz dominant frequency. This corresponds to a frequency range between 2 Hz and 20 Hz, in agreement with the predominantly observed seismic spectrum associated with rockfalls at the Dolomieu crater (see Fig. 1b and e.g. Hibert et al., 2014). The source-time function as well as its spectral content is displayed in Figure 2c. A typical wave simulation is shown in Figure 2d with a snapshot at time  $t = 3.2$  s illustrating the wave field radiated by a vertical point source located at the southern crater wall (yellow arrow). It can be observed that the majority of seismic energy is located close to the surface as a result of the shallow low seismic velocity. The surface topography causes a highly scattered wave field. Synthetic seismograms recorded at the surface along the cross-section are shown in a time-offset representation. The wave field originates at the source location (0 km offset) and travels outwards in all directions. Wave scattering caused by the topography is detectable here, especially close to the crater rim at around 0.6 km offset with reflections back-propagating towards the bottom of the crater.

213

Concerning computational efforts, it takes a CPU time of around 10,000 CPU hours (10 cores per CPU) for one simulation on the presented domain (i.e. duration: 6 s, number of elements: 915,704, number of GLL points: 6, max. frequency: approx. 20 Hz, min. velocity:  $320 \text{ m s}^{-1}$ ). We run the simulations in parallel on 200 CPUs, leading to 2.3 days per simulation.

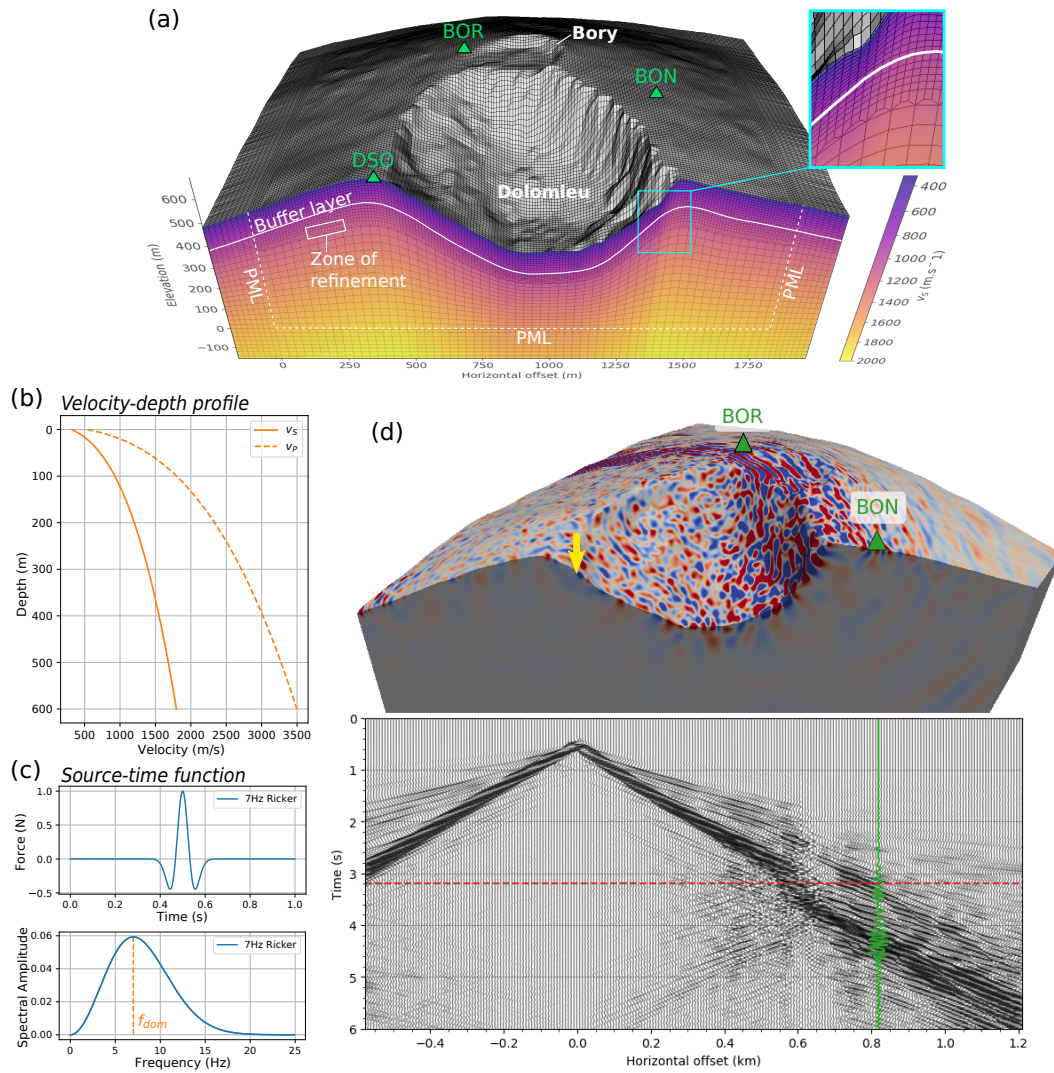
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To efficiently explore different potential positions of the rockfall source without performing a simulation for each of them, the reciprocity principle is used (Aki & Richards, 2002): the synthetic source is located at the station location and the wave field is recorded at the source location. Potential rockfall source positions are confined within a rectangular area at Dolomieu crater, shown in Figure 3a. The area is sampled by a grid of measurement points (in blue) with 10 m spacing.

224

The principle of reciprocity is illustrated in Figure 3b. It is shown that performing reciprocal simulations by interchanging source and receiver (and their corresponding directions) results in identical synthetic seismograms. In order to collect all necessary information, simulations for each component of all seismometers are carried out, i.e.

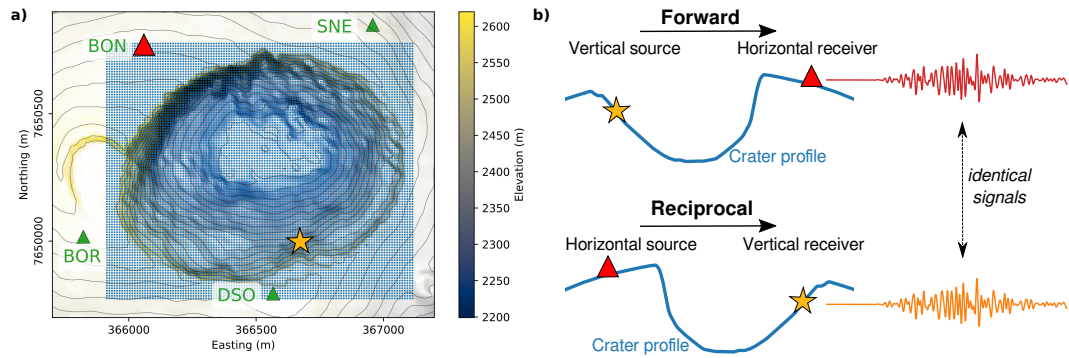
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**Figure 2.** (a) Cross-section through the meshed domain with 10 m resolution surface topography from Dolomieu crater. The color map depicts the seismic velocity model. The elements have a side lengths of 10 m in the top 150 m, increasing to 30 m below the *Zone of refinement*. Absorbing PMLs (Perfectly Matched Layers) of 160 m width are attached on the sides and bottom of the domain. (b) Generic velocity-depth profile for S-wave velocity  $v_S$  (solid line) and P-wave velocity  $v_P$  (dashed line) as proposed by Lesage et al. (2018) for the shallow structure of volcanoes. (c) Ricker wavelet of 7 Hz dominant frequency: source-time function (top) and corresponding spectrum (below). (d) Simulation of the wave field (vertical velocity) from a vertical source on the southern crater wall (yellow arrow). On the top, a snapshot of the wave field is shown at time  $t = 3.2$  s, where positive amplitudes are denoted in red, negative in blue. The graph below shows the seismic traces recorded at the surface along the same cross-section. The green trace corresponds to the signal recorded at station BON.

228 a point source is placed at the position of a given seismometer while the input force di-  
 229 rection is aligned with the component of the seismometer. In total, ten simulations are  
 230 carried out:  $3 \times 3$  simulations for the three-component seismometers BON, BOR, and  
 231 SNE and one simulation for the single-component seismometer DSO. This is done for





**Figure 3.** (a) Grid of receivers (in blue) for reciprocal simulations. The yellow star and red triangle are used to illustrate reciprocity in b). The sampled area measures 1200 m × 1000 m (east × north). Sample spacing is 10 m, resulting in 121 × 101 = 12221 grid points. (b) Illustration of the principle of reciprocity. *Top:* Forward simulation where source (vertical) and receiver (horizontal) are placed at the position of the true source and the true receiver, respectively. *Bottom:* Reciprocal simulation, where a synthetic horizontal source replaces the true horizontal receiver and a synthetic vertical receiver replaces the true vertical source, resulting in identical synthetic seismograms.

232 both the model with Dolomieu crater topography and for a model with a flat surface for  
 233 comparison.

234 **3.2 Optimization method for source location**

235 A source location probability estimate is associated with each point of the grid in  
 236 Figure 3a by considering the inverse of the misfit between the synthetic energy ratio  $E_{ij}^{\text{simu}}/E_{\text{ref},j}^{\text{simu}}$ ,  
 237 computed when the source is actually located at that specific grid point, and the observed  
 238 energy ratio  $E_{ij,tw}^{\text{obs}}/E_{\text{ref},j,tw}^{\text{obs}}$ . Here ‘ref’ refers to the reference station, while  $i$  designates  
 239 the station and  $j$  the component considered. Given that the rockfall source is moving,  
 240 the observed energy ratio is evaluated over a time window ‘tw’. The misfit  $e_{tw}$  for each  
 241 time window is defined as follows:

$$e_{tw} = \frac{1}{N_{\text{Sta}}} \sum_{ij} \left| \log_{10} \left( \frac{E_{ij}^{\text{simu}}}{E_{\text{ref},j}^{\text{simu}}} \div \frac{E_{ij,tw}^{\text{obs}}}{E_{\text{ref},j,tw}^{\text{obs}}} \right) \right|, \quad (2)$$

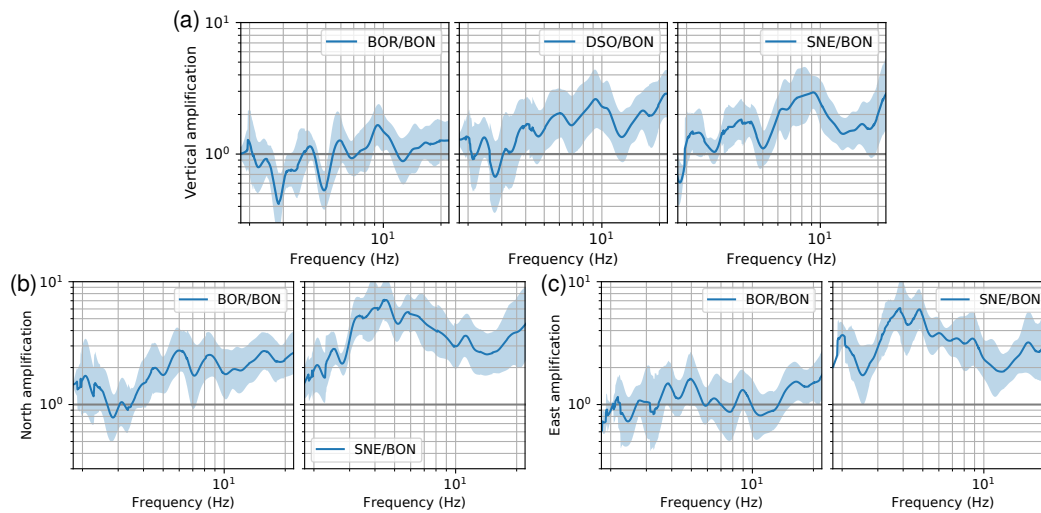
242 where  $N_{\text{Sta}}$  is the number of station-channel pairs to be considered, with each compo-  
 243 nent counted separately. Zero misfit is achieved when simulated and observed energy ra-  
 244 tios are equal. Using the logarithm in equation (2) distributes their relative values equally  
 245 around zero. This, combined with the absolute value, ensures that the misfit estimation  
 246 is not biased by the reference station. The probability of the source location is calculated  
 247 by the inverse of misfit  $e_{tw}$  and scaled to a probability density function (PDF) with re-  
 248 lative values between 0 and 1. Alternative formulas were investigated, for example the  
 249 relative difference between simulated and observed energy ratios or approaches with con-  
 250 ditional statements. However, the estimate in equation (2) was evaluated to be the most  
 251 appropriate for the location problem, notably because it results in spatially smooth vary-  
 252 ing probability values and because it is not biased towards the reference station.

253 In order to consider the frequency dependency of the energy ratios, location is car-  
 254 ried out in different frequency bands, namely at 3-7 Hz, 8-12 Hz, and 13-17 Hz. This se-  
 255 lection is defined to cover a large part of the available frequency content from the sim-  
 256 ulations and the observations. A bandwidth of 4 Hz is assumed to be narrow enough to  
 257 respect the dispersive character of the energy ratios and broad enough to average over

258 fluctuations of the spectral ratios. Noise levels at Dolomieu crater are very low at fre-  
 259 quencies above 3 Hz and can be ignored in the tests. Tests of the location method with  
 260 added synthetic with the noise are performed in section 5.3.

261 Selecting the width of time window ‘tw’ over which the observed seismic energy is  
 262 estimated requires special attention. On the one hand, the width has to be chosen as small  
 263 as possible in order to sample the moving source. On the other hand, as the same time  
 264 window is used for all stations, most of the seismic signal generated by a given rockfall  
 265 source has to arrive at each of the stations within the time window. In order to respect  
 266 both criteria, a window width of 4 s is defined, and confirmed by simulations to be an  
 267 appropriate compromise.

268 To allow comparison between observed and simulated energy ratios, the recorded  
 269 signals must first be corrected for local site amplification, not considered in the simu-  
 270 lations. Therefore, site amplification functions were estimated for each station channel  
 271 using thirty-six volcano-tectonic (VT) events that were centered around 2 km below Dolomieu  
 272 crater. Station BON is used as the reference station given its low spectral amplitudes  
 273 from VT recordings as well as low spectral H/V noise ratios. The resulting site ampli-  
 274 fication functions are shown in Figure 4. The site effect correction is performed prior to  
 275 locating by deconvoluting the recorded signals with the spectral amplification functions.



**Figure 4.** Spectral amplification functions estimated from volcano-tectonic (VT) events relative to reference station BON. Amplification is calculated from smoothed spectral ground velocity recorded by: a) vertical components, b) north components, and c) east components. Smoothing is performed as proposed by Konno and Ohmachi (1998). The blue-shaded area indicates the standard deviation of the amplification as calculated from all VTs. Figure adapted from Kuehnert et al. (2020).

276 To test the influence of the above parameters and site effects on the location method,  
 277 a hands-on Jupyter notebook (Kluyver et al., 2016) is published on [https://github.com/Jubeku/RF\\_localization](https://github.com/Jubeku/RF_localization)  
 278 (Kuehnert et al., 2019).  
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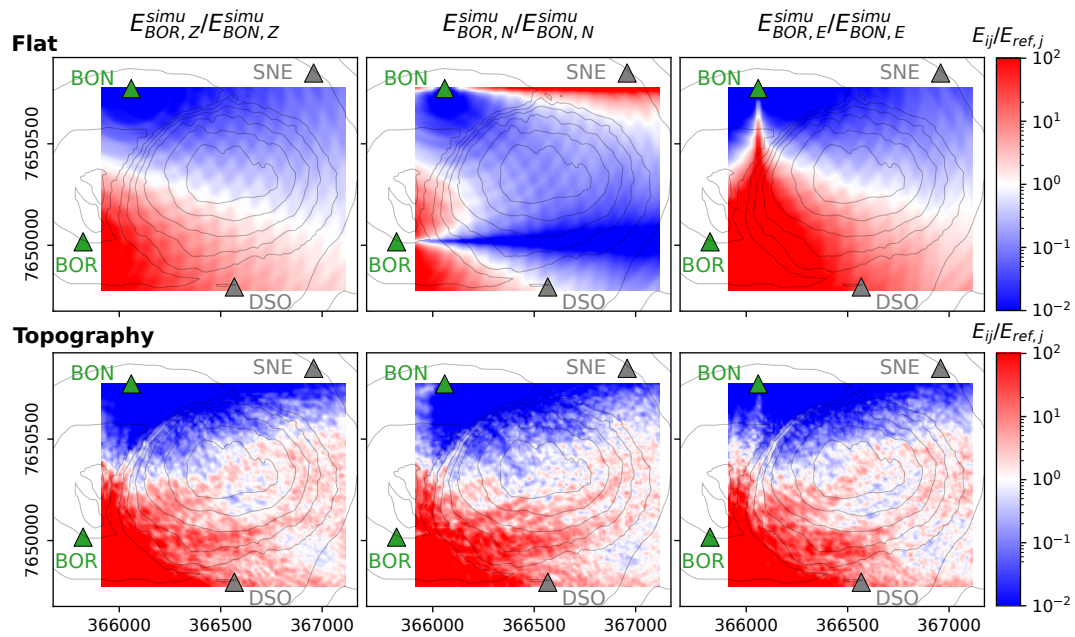
### 280 3.3 The influence of topography on inter-station energy ratios

281 The relative amplitudes recorded at various stations can be influenced by the to-  
 282 pography (e.g. Kuehnert et al., 2020), thus modifying the energy ratios in equation (2).  
 283 Having built databases of the simulated energies  $E_{ij}^{\text{simu}}$  for a domain with a flat surface

284 and a domain with topography, we can gain a first insight into the influence of topog-  
 285 raphy by comparing the resulting synthetic energy ratios.

286 This is done in Figure 5, where the energy ratio between station pair BOR and BON  
 287 at each grid point of potential source locations is shown for a flat surface (top) and for  
 288 the Dolomieu topography (bottom). The energy ratios  $E_{ij}^{simu}/E_{ref,j}^{simu}$  are calculated re-  
 289 spectively from all three components  $j = Z, N, E$  and reference station  $i = ref$  is cho-  
 290 sen to be BON.

291 For the domain with a flat free surface, when vertical signal component  $j = Z$  is  
 292 measured, we can observe a bipolar pattern of the energy ratios with values  $> 1$  towards  
 293 station BOR and values  $< 1$  towards reference station BON, while unity is reached at  
 294 equidistant positions between the station-pair. Values are determined purely by the source-  
 295 receiver distances. The energy ratios from the horizontal signal components  $i = N, E$   
 296 result in a more complicated spatial pattern. This is because the radiation pattern from  
 the vertical source in the horizontal plane is not radially isotropic.



**Figure 5.** Seismic energy ratios between station pair BOR and BON (in green) from simu-  
 lations on a domain with a flat surface (*top*) and a domain with topography (*bottom*). At each  
 grid point (see Fig 3a), representing a potential source position, the ratio is computed from  
 vertical-component seismic energy  $E_{i,Z}$  (*left*), north-component seismic energy  $E_{i,N}$  (*middle*),  
 and east-component seismic energy  $E_{i,E}$  (*right*). Unfiltered synthetic seismograms were used for  
 the calculation.

297 For the domain with surface topography, in the case of vertical component  $i = Z$ ,  
 298 the pattern of energy ratios becomes distorted because of the signature of the topogra-  
 299 phy in the wave field. In general, a bipolar spatial distribution of the energy ratios per-  
 300 sists, indicating that the decay of seismic amplitude remains influenced by the source per-  
 301 ceiver distance. The pattern of energy ratios from the horizontal signal components  
 302  $i = N, E$  is comparable to the vertical pattern, indicating that the wave propagation  
 303 along the topography dominates over source-characteristic radiation patterns. This to-  
 304 pographic path effect (e.g. Kumagai et al., 2011; Kuehnert et al., 2020) is similar to the  
 305 distortion of radiation patterns by the scattering of the wave field by small-scale soil het-  
 306 erogeneities, validating the assumption of an isotropic radiation at high frequencies above  
 307

308 around 5 Hz (e.g. Takemura et al., 2009; Kumagai et al., 2010). As a consequence, the  
 309 presented method can be implemented independently of the source impact direction used  
 310 in the simulations (here we have chosen a vertical surface traction), whereby both ver-  
 311 tical and horizontal component signals can be used for location. We show here that lever-  
 312 aging horizontal component signals can improve the locating results. Typically, only ver-  
 313 tical component signals are used in rockfall location methods, except for polarization ap-  
 314 proaches such as proposed by Vilajosana et al. (2008).

## 315 4 Application to rockfalls at Dolomieu crater

316 The proposed formalism to evaluate the relative probability of potential source lo-  
 317 cations on a predefined grid of positions is now applied to rockfall seismic signals recorded  
 318 at Piton de la Fournaise volcano. After analyzing individual time windows, all proba-  
 319 bilities derived from a sliding time window are combined in the attempt to reconstruct  
 320 the full rockfall trajectory.

### 321 4.1 Rockfall location at given time steps

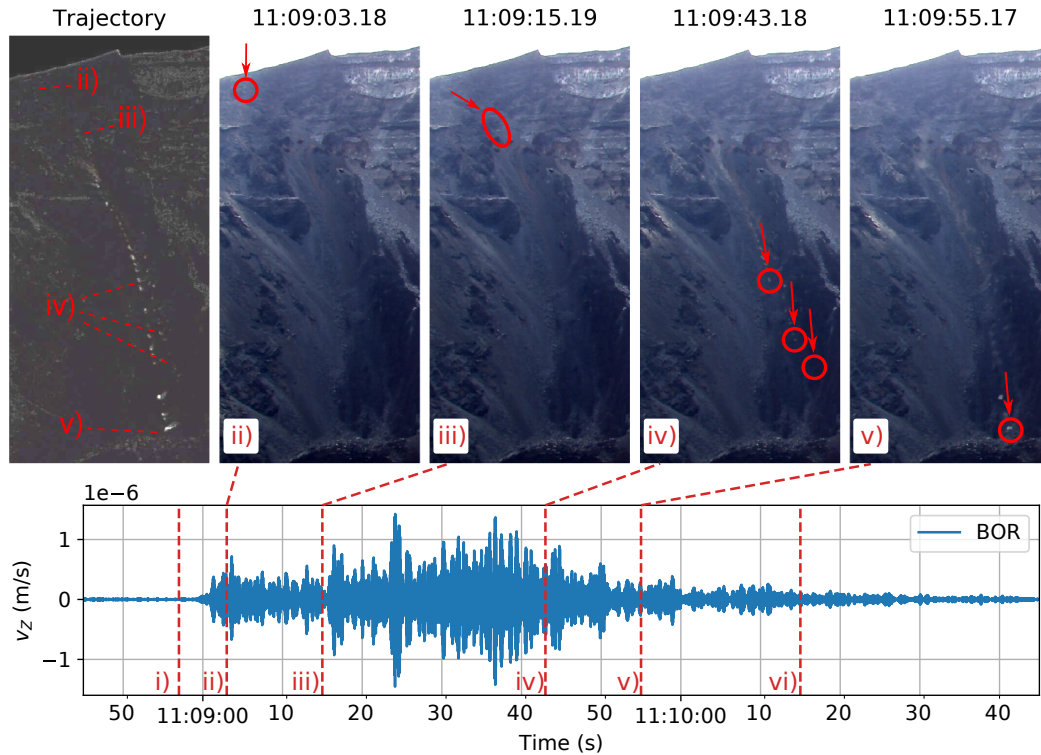
322 The location method is initially tested for a rockfall that occurred on December  
 323 13, 2016, corresponding to the event presented in Figure 1. The analysis is carried out  
 324 at six different times i) to vi) as indicated on the seismogram in Figure 6. Above the seis-  
 325 mogram, the whole trajectory is shown as well as camera snapshots of times ii) to v).  
 326

327 Time i) is just before the start of the rockfall. Time ii) is after the detachment, when  
 328 movements can be detected on the video. At time iii), the rockfall appears from behind  
 329 a small valley at the top of the crater wall. Thereafter, the rockfall accelerates, which  
 330 leads to stronger impacts and thus to the highest signal amplitudes. A total of three boul-  
 331 ders are detected at time iv) on their way down towards the crater bottom. At time v),  
 332 the third boulder arrives at the bottom. No movement is detected later on the video at  
 333 time vi). Nevertheless, it can be assumed that smaller granular material is still active  
 334 on the flank, causing small amplitude seismic signals.

335 Rockfall location is performed here in the highest frequency band (13-17 Hz) with  
 336 simulated energy ratios from the model with Dolomieu topography and using all avail-  
 337 able station-channel pairs, i.e.  $N_{\text{Sta}} = 7$ , adding up three station pairs for the verti-  
 338 cal component and two station pairs each for the north and east component (DSO con-  
 339 tains only the vertical component). Figure 7 shows the resulting source location prob-  
 340 ability maps at the six successive time steps i) to vi).

341 At time i), most probable seismic sources are located at the center and the south-  
 342 eastern side of the crater. As the rockfall has not yet started at that time, the distribu-  
 343 tion must be related to ambient seismic noise. At time ii), the source probability abruptly  
 344 moves southwest, marking the beginning of the rockfall. The position of maximum prob-  
 345 ability is around 100 m from the estimated location of detachment. Then the area of prob-  
 346 able source locations moves north at time iii) with the most probable source location ap-  
 347 proaching the estimated rockfall position. The predicted source location continues to move  
 348 along the rockfall trajectory at time iv). At time v) it arrives at the bottom of the crater.  
 349 At this time, the position of maximum probability is at around 200 m from the estimated  
 350 location. However, the distribution also shows densely populated high probabilities close  
 351 to the estimated rockfall location. At time vi), after the last boulder visible on the video  
 352 has reached the crater floor, a zone of probable source positions remains in the lower part  
 353 of the trajectory. This may be explained by the late movement of granular material that  
 354 is not visible on the video.

355 The spatially scattered distribution of the predicted sources and the discrepancy  
 356 between position of maximum probability and actual rockfall location imply a lack of  
 357 resolution that can have several reasons. Firstly, the source positions are somewhat am-  
 358 biguous, i.e. different locations can explain the observed seismic energy ratios equally



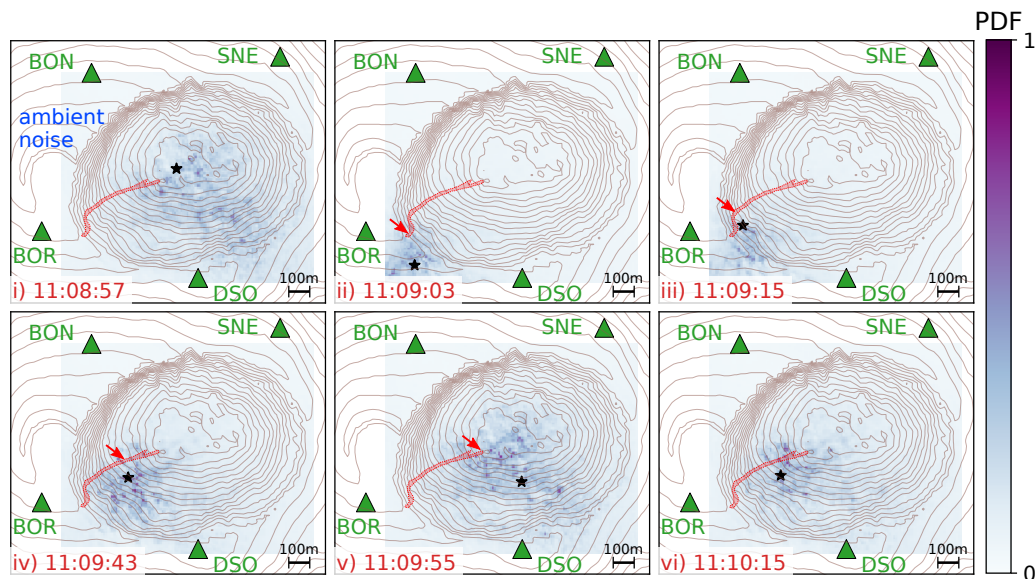
**Figure 6.** Camera images and seismic signal of a rockfall on December 13, 2016. *Top:* Images taken by camera DOEC. The full rockfall trajectory on the left is reconstructed from differences between successive images. Towards the right, snapshots at times ii) to v) are displayed. Rockfall positions are indicated by red circles and the direction of movement by red arrows. *Bottom:* Vertical ground velocity  $v_z$  recorded at closest station BOR. Time steps i) to vi) are marked by red vertical dashed lines. Rockfall location is performed in time windows of  $\pm 2$ s around these time steps. The signal is bandpass filtered at 13-17 Hz.

359 well. Secondly, as observed on the video, the rockfall does not consist of a single boulder.  
 360 The resulting seismic signal is hence a superposition of multiple sources shifted in  
 361 time and space. Given that it is based on the assumption of a single source, the loca-  
 362 tion method is flawed, a problem that will be studied in section 5.1 using synthetic sig-  
 363 nals. The general southward shift of the predicted source locations compared to the true  
 364 trajectory may also be caused by soil heterogeneities that affect seismic wave propaga-  
 365 tion and are not considered by either the simulations or the site impact correction. An-  
 366 other cause could be an inaccurate representation of the topography, which is possible  
 367 since the DEM used here is from 2009 and the rockfall analyzed occurred in 2016.

## 368 4.2 Spatio-temporal rockfall evolution

369 In order to reconstruct the full rockfall trajectory, the location method of equation  
 370 (2) is used with a sliding time window. Results from all time windows are combined at  
 371 each potential source position by selecting the maximum probability over time. For each  
 372 grid point, the minimum misfit  $e$  between observed and simulated energy ratios is de-  
 373 fined by

$$e = \min_{tw} e_{tw}, \quad (3)$$



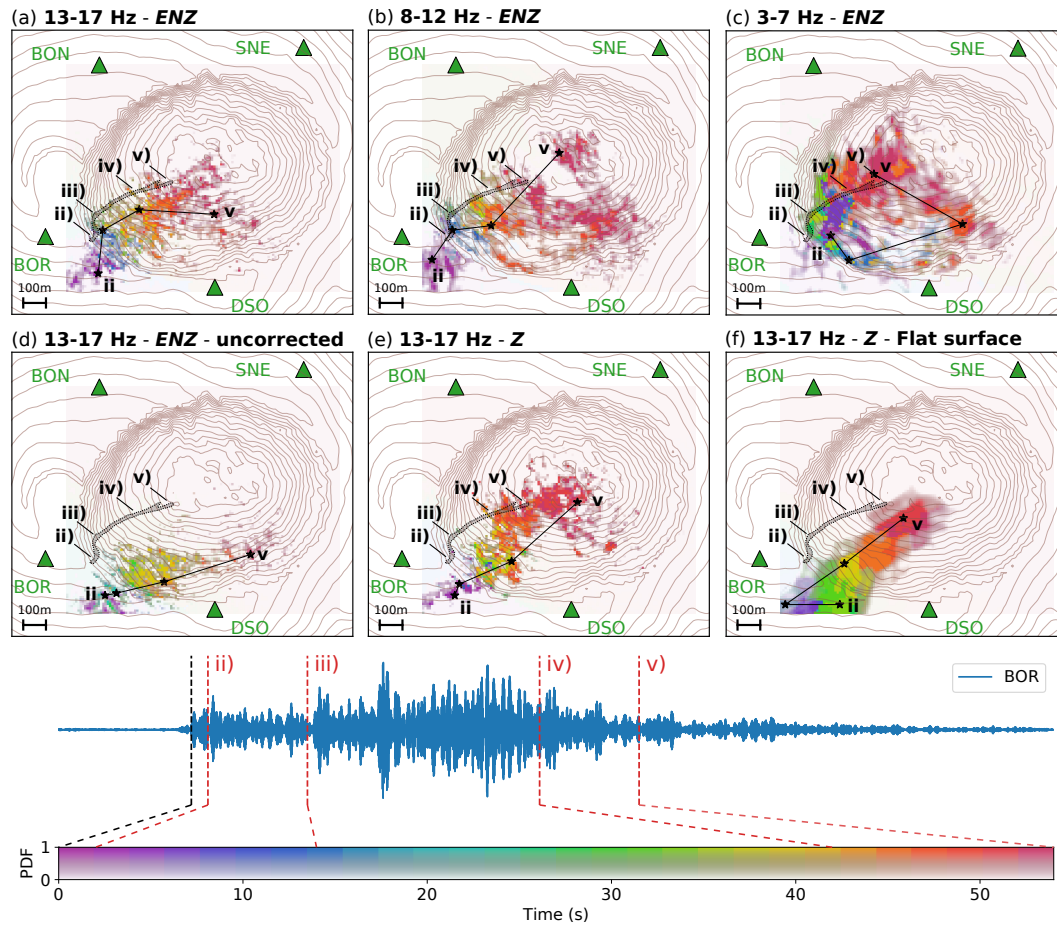
**Figure 7.** Location of seismic source at time steps i) to vi) as defined in Figure 6. The color-scale represents the source location probability. Black stars denote the position of maximum probability. Red shaded zone marks actual rockfall trajectory as estimated from the video and the red arrows approximate the current rockfall location. Locating is carried out in frequency band of 13-17 Hz using simulations from the model with Dolomieu topography. All stations and components are used, i.e.  $N_{Sta} = 7$ .

374 where  $e_{tw}$  is the misfit in each time window ‘tw’ defined in equation (2). The maximum  
 375 probability is the inverse of misfit  $e$  and can be plotted for each spatial point. In this way,  
 376 the temporal evolution of the rockfall trajectory can be displayed on a single graph.

377 Figure 8a shows the resulting location map of the previously analyzed rockfall, using  
 378 the same method configuration (i.e. at high frequency 13-17 Hz and with all available  
 379 station-channel pairs). Thanks to the color sequence, we can track how the rockfall  
 380 moves from top to bottom of the crater over time in agreement with the observed  
 381 rockfall trajectory from the video. Black stars denote the positions of maximum prob-  
 382 ability at time steps ii) to v), identical to those shown in Figure 7. Again, a general south-  
 383 ward shift of around 100 m with respect to the video-estimated trajectory is observed.

384 Rockfall location is performed in different frequency bands. In the intermediate fre-  
 385 quency band at 8-12 Hz, Figure 8b, the predicted source locations follow the movement  
 386 of the actual trajectory and the positions of maximum probability are at distances com-  
 387 parable to those in Figure 8a. However, the resolution decreases as the spatial distribu-  
 388 tion of probable sources becomes much wider, covering large parts of the crater. The res-  
 389 olution is even worse in the lowest frequency band at 3-7 Hz, Figure 8c, where the gen-  
 390 eral downward movement of the rockfall is hardly noticeable, with large discrepancies  
 391 of the maximum probability positions in the time steps ii) to v). The observed decrease  
 392 of resolution towards lower frequencies can be explained by the increase of the seismic  
 393 wavelength. Assuming mainly fundamental-mode Rayleigh waves (Kuehnert et al., 2020),  
 394 the wavelength for the velocity model used here increases from 26 m at 15 Hz by a fac-  
 395 tor of 1.7 to 44 m at 10 Hz and by a factor of 4.5 to 116 m at 5 Hz; resolution can be ex-  
 396 pected to decrease accordingly.

397 A reduction in location error when higher frequencies are used in the location pro-  
 398 cess is also reported by Lacroix and Helmstetter (2011). When analyzing single impact  
 399 signals with frequency contents up to 30 Hz, they achieve locating accuracies of 50 m us-



**Figure 8.** Spatio-temporal rockfall evolution. Color represents time of seismic record and intensity representing probability of source location. Black shaded zone indicates rockfall trajectory from video. Black stars denote positions of maximum probability at time steps ii) to v). The seismogram underneath was recorded at closest station BOR at 13-17 Hz, associating time with color using a 2D colorbar (MMesch, 2016). Signals recorded at all stations for all components are shown with scales in the supplementary material. Using the results from the same simulation including topography, the pre-processing of synthetic and observed signals for location changes as follows: (a) Location at high frequency band 13-17 Hz using all available station channels (*ENZ*, i.e. east, north, and vertical components), thus  $N_{Sta} = 7$ . (b) Location at intermediate frequency band 8-12 Hz, with all components *ENZ*, i.e.  $N_{Sta} = 7$ . (c) Location at low frequency band 3-7 Hz, with all components *ENZ*, i.e.  $N_{Sta} = 7$ . (d) Location at 13-17 Hz, with all components *ENZ*, i.e.  $N_{Sta} = 7$ , without site effect correction. (e) Location at 13-17 Hz, using only vertical component *Z*, i.e.  $N_{Sta} = 3$ . (f) Location at 13-17 Hz, using vertical component *Z*, i.e.  $N_{Sta} = 3$ , and using simulations from a model with a flat surface.

400 ing beamforming and a priori measured seismic velocities. Similarly, when analyzing single  
 401 impacts and frequency contents of 5-25 Hz, Dietze et al. (2017) achieves average lo-  
 402 cation accuracies of 81 m, comparing the results of back-projecting seismic envelopes with  
 403 those of TLS-based measurements. For continuous and distributed sources, the reported  
 404 location accuracies decrease. Pérez-Guillén et al. (2019) use the ASL technique in a slid-  
 405 ing window to track snow avalanches and slush flows. When comparing the locations from  
 406 seismic signals to numerical flow simulations, they report mean locating accuracies be-

407 tween 85 m and 271 m, which is of a similar order of magnitude to the results presented  
 408 here.

409 Figure 8d shows the location results without prior correction of the recorded sig-  
 410 nals from site amplification. The results fail to predict a clear rockfall trajectory and a  
 411 large spatial discrepancy is observed between probable source positions and actual rock-  
 412 fall location.

413 In Figure 8e, the rockfall is located using seismic signals of only vertical compo-  
 414 nent  $Z$ , leading to  $N_{\text{Sta}} = 3$  station-channel pairs. A narrow corridor of high proba-  
 415 bilities can be seen, indicating a well resolved rockfall trajectory. However, compared to  
 416 the results in Figure 8a, a larger discrepancy with the actual rockfall location is observed.  
 417 This suggest that adding additional measurements may reduce the resolution as it be-  
 418 comes harder to keep the misfit, as defined in equation (2), small, however, the predic-  
 419 tions potentially improve as noisy or malfunctioning measurements can be compensated,  
 420 which is in agreement with the network resolution study in section 5.2.

421 In Figure 8f, locating is carried out using simulated energy ratios from a model with  
 422 a flat surface. In this case, only vertical components can be used as energy ratios from  
 423 horizontal components lead to values that are strongly modulated by radiation patterns,  
 424 as shown in Figure 5. The resulting source probability distribution consists of patch-like  
 425 areas that do not show smooth transitions over time (i.e. color), leading to a coarse rock-  
 426 fall location that is not well resolved in time. This is because the energy ratios from the  
 427 flat model are dominated by the source-receiver distances and only these localized patches  
 428 can explain the observed energy ratios. Typically, rockfall location methods attempt to  
 429 account for the effect of the topography during the location process by considering a map  
 430 of elongated travel times (e.g. Hibert et al., 2014; Levy et al., 2015; Dietze et al., 2017),  
 431 assuming straight wave paths along the surface. The method proposed here allows the  
 432 high-resolution topography and its influence on the wave field to be fully accounted for  
 433 by 3D numerical modeling of the seismic wave field. The influence on the location by im-  
 434 plementing a slightly coarser resolution DEM is demonstrated in section 5.3.

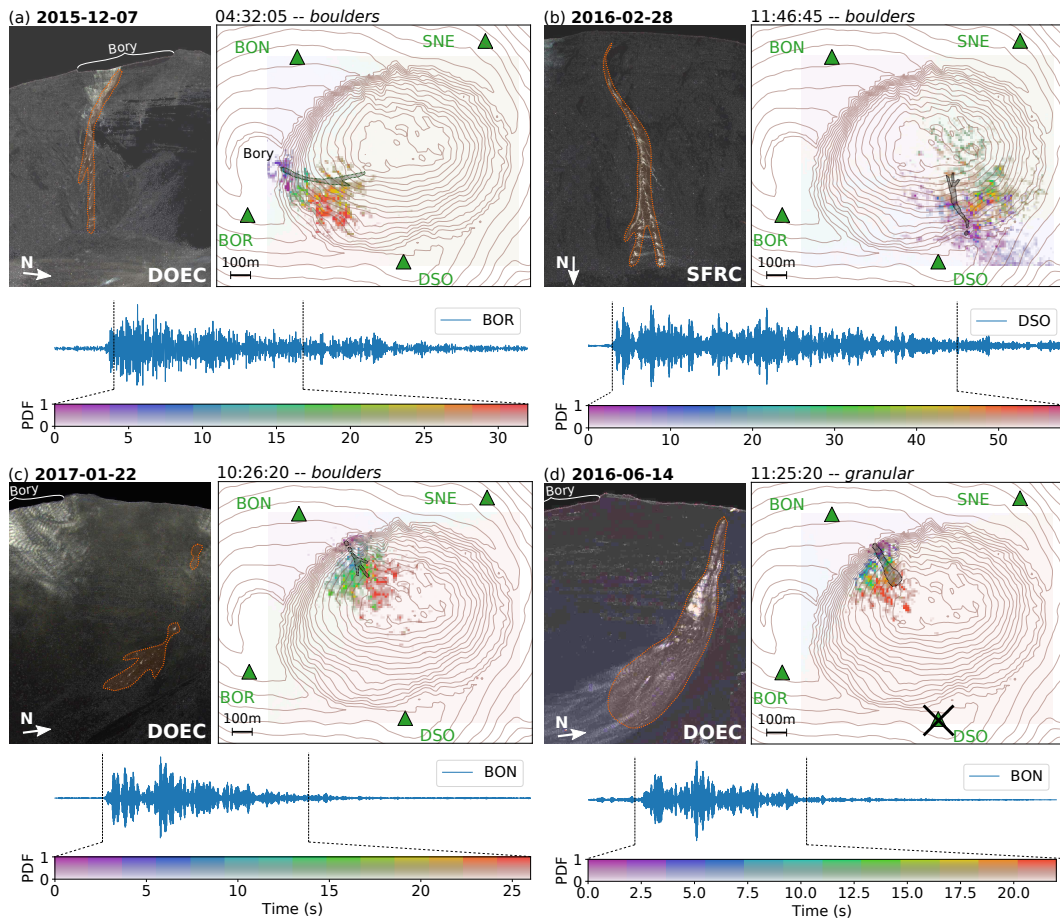
### 435 4.3 Locating other rockfalls

436 The comparison in the previous section suggests that the best locating results for  
 437 the present study site can be achieved in the high frequency range (13-17 Hz), using both  
 438 vertical and horizontal components ( $ENZ$ ), removing site effects from the observed sig-  
 439 nals, and simulating energy ratios on the model with topography. With this configura-  
 440 tion, the location method will now be evaluated using four rockfalls of different types  
 441 and from different locations within the Dolomieu crater. The observed trajectories as  
 442 well as the locations are shown in Figure 9 and described below.

443 Rockfall (a) is again located in the southwestern region, with an initial detachment  
 444 further north compared to the previously analyzed event. Rockfall (b) is in the south-  
 445 east and rockfall (c) is in the north of the Dolomieu crater. Rockfall (d) occurred in the  
 446 same region as rockfall (c) but consisted of fine granular material in contrast to the other  
 447 events which consisted of individual boulders.

448 For rockfall (a), Figure 9a, the most probable sources are all inferred at locations  
 449 close to the observed trajectory in the southwestern region of Dolomieu crater. In par-  
 450 ticular, the location of the detachment phase (in purple) and the observed trajectory to-  
 451 wards the east are well represented. However for the last stage of the rockfall, inferred  
 452 sources are located too far south, at the wall of the crater, and not at the bottom of the  
 453 crater as observed. This might be interpreted as the signature of possible superposition  
 454 of the seismic signals generated by subsequent boulders or an incorrect topography rep-  
 455 resentation. Note that the resolution of the location method makes it possible to iden-  
 456 tify the trajectory and the detachment of this rockfall event distinctly with respect to  
 457 the trajectory and detachment of the event analyzed in the previous section, Figure 8,  
 458 for which the detachment phase is located 100 m further south.





**Figure 9.** Four rockfalls used for locating evaluation. Left: trajectory reconstructed from successive camera images (outline marked in orange and north-direction and camera indicated at the bottom). Right: map of predicted spatio-temporal source evolution (black-shaded video estimated trajectory). Underneath: seismograms recorded at closest station (vertical ground velocity, bandpass filtered at 13-17 Hz). Signals recorded at all stations for all components are shown with scales in the supplementary material. a) Rockfall consisting of individual boulders occurring on December 7, 2015 in the southwest with detachment position just beneath Bory crater. b) Rockfall consisting of individual boulders occurring on February 28, 2016 in the south. c) Rockfall consisting of individual boulders occurring on January 22, 2017 in the north. d) Rockfall consisting of fine granular material occurring on June 14, 2016; traces from dust clouds extend beyond the outlines of the sketched event location and station DSO was malfunctioning and could not be used for locating.

459 For rockfall (b), Figure 9b, the inferred sources are correctly located in the southern  
 460 region of the Dolomieu crater, but with strongly deteriorated resolution in space  
 461 and time. The inferred source locations 30 s after the start of the event (in green),  
 462 are located at the bottom of the crater. The video shows that the first boulder  
 463 arrives at the crater bottom at this time, but other boulders are still moving at  
 464 the top of the crater wall. Multiple sources hamper the ability of the location  
 465 method to determine the trajectory of a single source. As a result, the inferred  
 466 sources at later times (yellow and red colors) are located half-way down at the  
 crater wall. Another explanation for the poor resolu-

467 tion is the station network configuration which will be studied in section 5.2 using syn-  
 468 thetic signals.

469 For rockfall (c), Figure 9c, the inferred sources are well-located at the beginning  
 470 of the event, while locations become more and more scattered in space at later times.  
 471 This time-deterioration of the resolution can be analyzed with the help of the video that  
 472 shows that at beginning the event initially involves a single boulder impacting the crater  
 473 wall, with subsequent distribution of boulders originating from the fragmentation of the  
 474 original boulder or from the mobilisation of basal rock deposits.

475 Finally for rockfall (d), Figure 9d, which occurred in the same region as event (c)  
 476 but consisted of fine granular material flowing down the steep crater wall, the method  
 477 is able to locate the event with high-resolution, in particular the initial activation loca-  
 478 tion. This is quite remarkable given that station DSO was not functioning properly and  
 479 was disregarded for the analysis. Moreover, given that the source is parametrized as a  
 480 single force, this high-resolution of the source locations in the case of granular flow is not  
 481 intuitive as it generates a complex extended source. This might suggest that recorded  
 482 signals are dominated by a localized high-energy radiating source area, which we will fur-  
 483 ther discuss in section 5.4. Using a similar approach based on analysing the seismic sig-  
 484 nals in a sliding time window, Pérez-Guillén et al. (2019) are also able to track the dis-  
 485 tributed and moving seismic sources generated by snow avalanches and slush flows.

## 486 5 Evaluation of the presented location method using synthetic signals

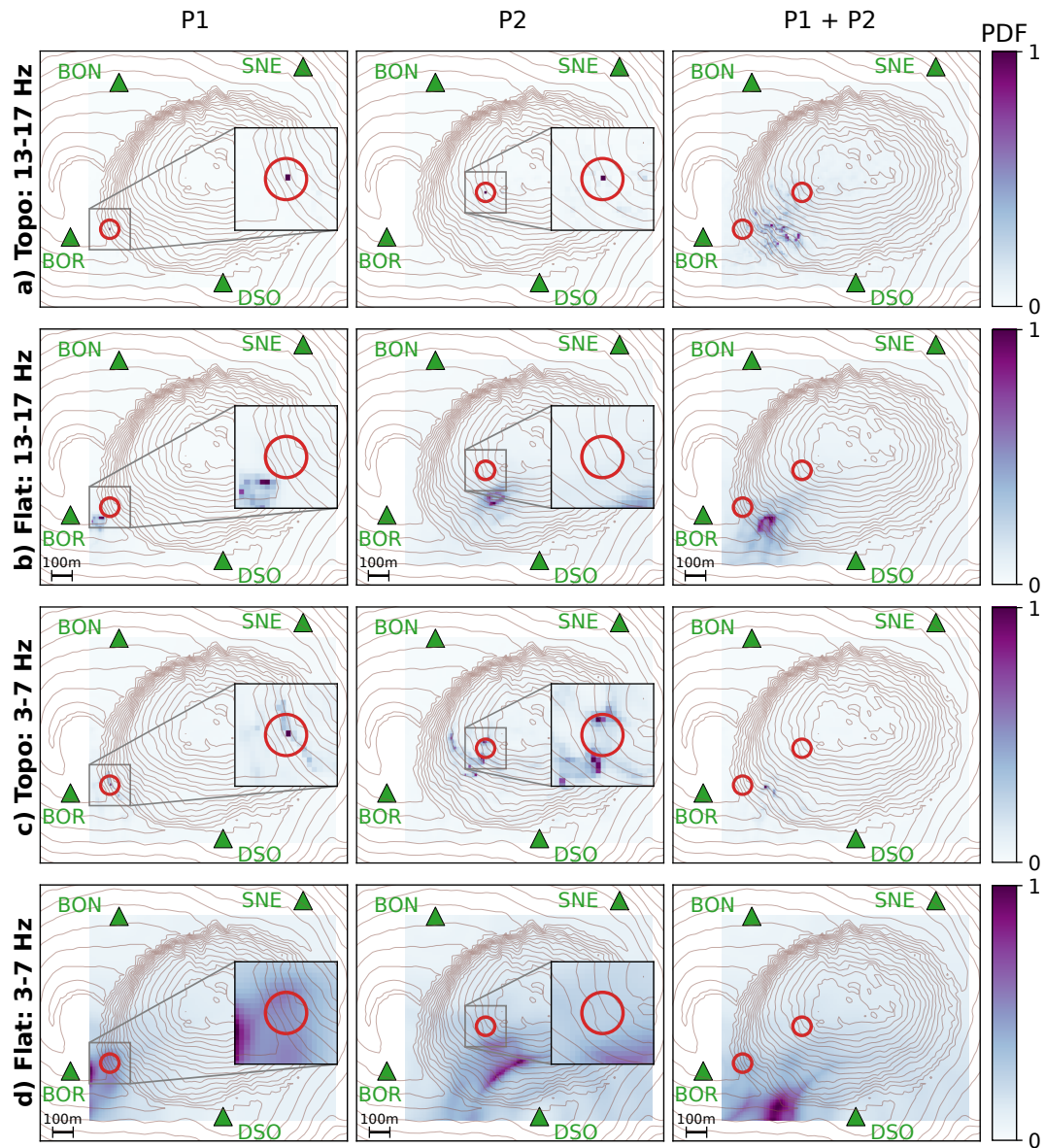
487 Rockfall events generate complex and extended seismic sources, and the resolution  
 488 and the limitations of the proposed location methods need to be assessed through tests  
 489 with synthetic seismograms for which the seismic sources can be controlled, e.g., with  
 490 known source time functions and locations. In this way, we study the problem of the su-  
 491 perposition of spatially distributed sources as well as the performance of the location method  
 492 in different frequency bands and the error introduced when topography is not consid-  
 493 ered. The study with synthetic seismograms additionally offers the possibility to eval-  
 494 uate the influence of the network geometry on the location resolution.

### 495 5.1 Single-point sources mimicking boulder impacts

496 In order to assess the frequency-dependent location error introduced when topog-  
 497 raphy effects are ignored, synthetic seismic signals are generated for two controlled point  
 498 sources (e.g. two distinct boulders impacting at the same time with the same force, lo-  
 499 cated in the southwestern part of Dolomieu crater) through 3D wave simulation in the  
 500 model including topography. Locations of the sources are then inferred from different  
 501 frequency bands of the synthetic signals using the location method with and without top-  
 502 ography effects. For better comparability between these two cases, only vertical com-  
 503 ponent signals are used for location. Figure 10 summarizes the main results. The true  
 504 locations of the two point sources are located in the center of the red circles.

505 The left and middle columns show the inferred locations when considering sources  
 506 at position P1 and P2 separately and the right column when the two sources are act-  
 507 ing simultaneously. Rows a-d show the inferred locations when considering two frequency  
 508 bands of the synthetic signals (i.e. 13-17 Hz and 3-7Hz) and propagating models with  
 509 topography or a flat surface.

510 For the high-frequency band, Figure 10a shows the results when using propagat-  
 511 ing models including topography effects. The true positions of the sources P1 and P2,  
 512 when considered acting separately, are well reconstructed as would be expected, corre-  
 513 sponding to a point with high-probability (i.e. dark purple) in the center of the red cir-  
 514 cles. In contrast, when using a flat propagating model, Figure 10b, sources are recon-  
 515 structed with a 100 m-200 m location error and with a spatial ambiguity (multiple source  
 516 positions with similar probability).



**Figure 10.** Location of point source at positions P1 and P2 and after simultaneously activating P1 and P2. The exact source position is located in the center of the corresponding red circle. Location is performed using vertical component  $Z$ , i.e.  $N_{Sta} = 3$ . In each map, the color is normalized by the maximum probability. (a) Location of signals in frequency band 13-17 Hz using simulations from the domain with topography. (b) Location of signals in frequency band 13-17 Hz using simulations from the domain with a flat surface. (c) Location of signals in frequency band 3-7 Hz using simulations from the domain with topography. (d) Location of signals in frequency band 3-7 Hz using simulations from the domain with a flat surface.

517 For the low-frequency band, when including topographic effects, Figure 10c, the  
 518 location of the source P1 is well reconstructed, while the location is more ambiguous for  
 519 the source P2 and spatially scattered within an area of size up to 300 m. The imperfect  
 520 location could be caused by the 4s time window cutting a part of the low frequency signal.  
 521 As expected, when topographic effects are not included (flat model), Figure 10d,

522 the inferred locations become more blurred. With longer probing wavelengths (i.e. for  
 523 the low frequency band), we would expect reduced location resolution for both models.  
 524 The good reconstruction of source P1 in this low-frequency band for the model with to-  
 525 pography is therefore puzzling. This cannot be explained by the proximity of station BOR,  
 526 since better resolution of the source P1 would in that case also be observed when using  
 527 the flat model. This might be the signature of topography effects, since the source P1  
 528 is located just below the crater rim, one of the steepest regions of the crater geometry,  
 529 leading to generated signal characteristics quite distinct from those of neighboring po-  
 530 tential locations.

531 When both sources are acting simultaneously (Figure 10, right column), positions  
 532 of the individual sources can no longer be determined for all the test cases. Taking into  
 533 account topography effects, Figures 10a and c, the probability distribution of source lo-  
 534 cation inferred from the high-frequency band (Figure 10a) is spatially scattered with rel-  
 535 atively high-probability patches of around 300 m size in the neighborhood of the indi-  
 536 vidual sources, while the distribution inferred from the low-frequency band is focused  
 537 in a single region that seems to best explain the superposed signal from the two impacts.  
 538 In other words, a single source in this region would result in similar relative energy mea-  
 539 surements at the stations as the superimposed signal from the two sources. Interestingly,  
 540 a small shift to the south is observed, similar to what occurs when locating real rock-  
 541 falls in this area (see Fig. 8a and 9a), suggesting that the observed shift could partly be  
 542 caused by the superposition of impacts from several boulders at different locations.

543 When ignoring topographic effects with the flat wave propagating model, Figures  
 544 10b and d, the inferred probability distribution of source location becomes smoother and  
 545 less spatially resolved, since recorded seismic signals then contain only information on  
 546 the source-receiver distances. For the low-frequency band, the relatively high-probability  
 547 areas are loosely defined and shift further away from the actual source positions.

548 To summarize, the individual sources can be well located only from high-frequencies  
 549 and when taking into account topography effects. When the two sources are acting si-  
 550 multaneously, i.e. impacting at the same time and with same force, the individual sources  
 551 can no longer be distinguished and the location probabilities are concentrated somewhere  
 552 in the vicinity between the sources. Similar results were reported by Kumagai et al. (2009)  
 553 who numerically tested the amplitude source location (ASL) method with two simulta-  
 554 neous, spatially separated sources, resulting in the best location being between the two  
 555 sources. Nonetheless it is important to bear in mind that for real rockfalls, radiating sources  
 556 are non-uniform in space and time. This means that recorded signals are dominated by  
 557 the signature of the most strongly radiating sources at a given time. This makes it pos-  
 558 sible to locate the strongest sources in space at each time and to reconstruct rockfall tra-  
 559 jectories reasonably well.

## 560 **5.2 The influence of the network geometry on the location resolution**

561 The previous analysis is extended by quantifying the decay of the location prob-  
 562 abilities as a function of distance from the actual source. In this way, a proxy for the spa-  
 563 tial resolution is defined for each grid point, which can serve as an array response func-  
 564 tion for single-impact sources.

565 Considering each grid point as the actual source position, the location probabili-  
 566 ty of all other grid points is calculated (resulting in probability maps as in Figure 10).  
 567 Then, assuming a circular symmetry of the probability as a first order approximation,  
 568 the source probability  $p$  as a function of distance  $d$  from the actual source is approxi-  
 569 mated by an exponential decay of the following form:

$$p(d) = a \exp(-dk) + b, \quad (4)$$

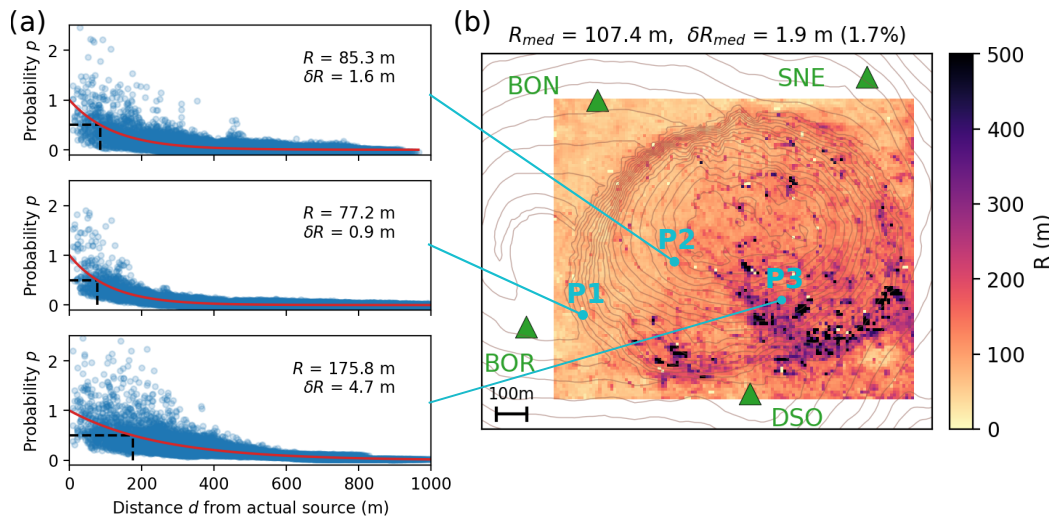
570 where  $a$ ,  $k$ , and  $b$  are fitting parameters. The fit is performed without considering the  
 571 probability value at the actual source position to avoid influence from singularities (such  
 572 as the high probability at the source position in Figure 10a). Finally, a proxy  $R$  for the

573 spatial resolution is defined by the *half-life* of the exponential decay using rate constant  
 574  $k$ :

$$R = \frac{\ln(2)}{k}. \quad (5)$$

575 The shift  $b$  can be ignored, since only relative probability variations are important.

576 Figure 11 presents the analysis performed for a network using the vertical compo-  
 577 nents  $Z$  of all four stations. The decay of probabilities with distance is shown for three  
 578 points as example in Figure 11a, where points P1 and P2 correspond to the points an-  
 579 alyzed in Figure 10. The fitted exponential function is shown in red and half-decay is  
 580 marked with black dashed lines. The uncertainty  $\delta R$  on resolution proxy  $R$  has been prop-  
 agated from the fitting error  $\delta k$  in the rate constant  $k$ .

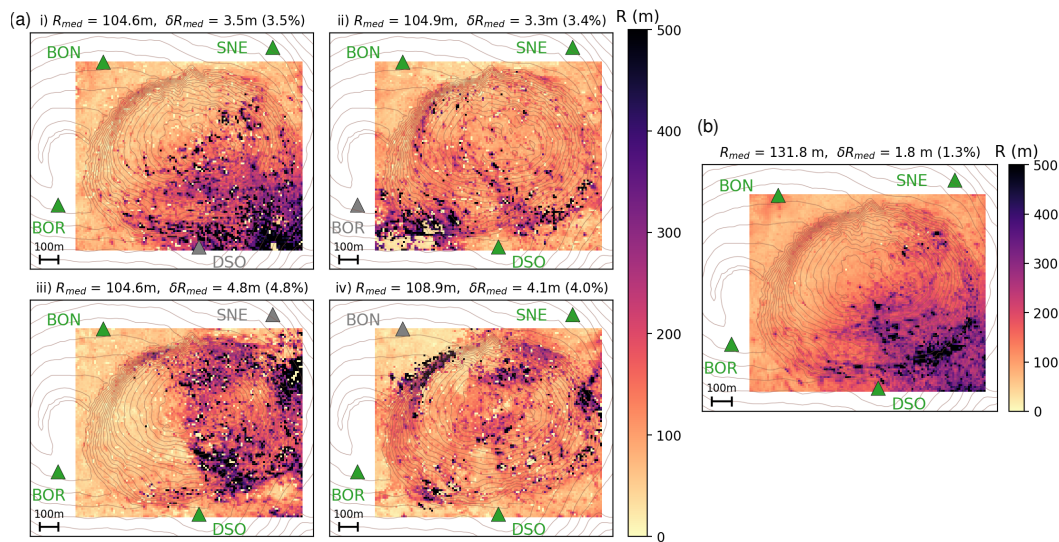


**Figure 11.** Resolution proxy  $R$  from a network of four stations with vertical components. (a) Source location probability as a function of distance from the actual source at P1, P2, and P3, respectively. Blue circles in the distribution correspond to all the other grid points of potential source locations. The fitted exponential function is shown in red. Black dashed lines mark the position of resolution proxy  $R$ , defined by the half-decay. The plots are normalized so that  $p(0) = 1$  and  $p(d \rightarrow \infty) = 0$ . (b) Map of the resolution proxy  $R$ , constructed by calculating the half-decay for each grid point in case it is the actual source, as shown for points P1, P2, and P3.

581 The map of resolution proxy  $R$  in Figure 11b shows a median resolution of  $R_{med} =$   
 582 107.4 m and generally indicates proxy values below 100 m in the northwest (as for points P1  
 583 and P2) and values above 100 m in the southeast (as for point P3). The southeastern  
 584 region is not enclosed by the network geometry which may explain the poorer resolution.  
 585 The poorer resolution in the vicinity of station DSO may be caused by the proximity of  
 586 this station to the crater rim, suggesting that this is not an optimal position for locat-  
 587 ing seismic sources.  
 588

589 To evaluate the influence of the network geometry on the resolution, the analysis  
 590 was performed with one of the stations removed alternately. Figure 12a shows the re-  
 591 sulting maps of resolution proxy  $R$  using a reduced network of three stations with ver-  
 592 tical components only. Generally, the resolution becomes poorer in the direction of the  
 593 removed station while the three remaining stations form a triangle that spans an area  
 594 of enhanced resolution, best seen in i) and iii). This triangle is not so clearly visible in  
 595 ii) and iv) because of the generally poorer resolution in the southeast (see Fig. 11b).

596 The median resolution  $R_{med}$  decreases slightly compared to the previous analysis  
 597 with four stations (except for case iv). If fewer stations are involved, the accumulated



**Figure 12.** Influence of the network geometry on the resolution proxy  $R$ . (a) Map of resolution when using three stations of vertical components  $Z$ . Removed station is shown in grey. Station combinations consist of i) BON-BOR-SNE, ii) BON-DSO-SNE, iii) BON-BOR-DSO, and iv) BOR-DSO-SNE. (b) Map of resolution when using four stations and all available channels, i.e. three component  $ENZ$  of stations BON, BOR, and SNE; vertical component  $Z$  of station DSO.

598 misfit at positions in the vicinity of the actual source is lower, resulting in higher probabilities and steeper decay of the exponential curve. However, the relative median uncertainty  $R_{\text{med}}$  increases by a factor between 2 and 3 (from 1.7% to 3.4%-4.8%), indicating a more scattered probability distribution.

602 On the contrary, adding measurements increases the median resolution as can be seen in Figure 12b with  $R_{\text{med}} = 131.8$  m, where all available station components have been combined (i.e. three components of BON, BOR, and SNE, and vertical-component of DSO). Nevertheless, the relative median uncertainty decreases to 1.3% and the spatial variation of the resolution is smoother compared to the response from vertical components only (Figure 11b). This suggests that the location method is more stable with an increased number of measurements which help to better determine the solution space. This is analogous to findings of Kraft et al. (2013) whose optimal network design algorithm, which takes into account laterally variable noise levels, often extends an established network with stations near existing station locations to further enhance the seismic source resolution. Their algorithm is based on the linearized earthquake location problem ( $D$ -criterion), first implemented by Kijko (1977). Toledo et al. (2020) use the same theory to develop a network design tool for seismic sources in geothermal and volcanic contexts. Their study shows how the first four stations can significantly improve the cost-benefit given optimal locations, while the added value decreases with each additional station in a power-law like manner.

618 For the given network at Dolomieu crater, the above tests indicate enhanced resolution in the area which is enclosed by the network geometry. This is an effect which can be observed in previous rockfall location studies (e.g. Lacroix & Helmstetter, 2011; Gracchi et al., 2017) and agrees with findings from optimal network design studies (e.g. Rabinowitz & Steinberg, 1990). In the present case, the southeast part of the crater shows lower resolution, explaining the poor location of the rockfall on February 28, 2016, in Figure 9b. Adding additional measurements, such as horizontal channels, can increase the stability of the solution with only a slight loss of resolution, which is especially impor-

626 tant in the field when measurements can be contaminated by noise or when the site am-  
627 plification functions are poorly known, which is in agreement to the rockfall location re-  
628 sults in Figure 8a and e using three components ( $ENZ$ ) and the vertical component ( $Z$ ),  
629 respectively.

### 630 **5.3 Multiple-point sources mimicking a down-slope moving rockfall**

631 Synthetic rockfall seismic signals are generated here from a downward moving seis-  
632 mic point source, i.e. parametrized as a single vertical traction, kinematically constrained  
633 by the boulder trajectory observed during one rockfall event at the Dolomieu crater (De-  
634 cember 13, 2016) already discussed in section 4.1. The space-and-time positions of the  
635 seismic point-source is mapped in Figure 13. Another representation of the space-time  
636 trajectory of the seismic source is shown in the graph at the top of Figure 14a.

637 To construct the source space-and-time trajectory of the point-source, the position  
638 and time of seven markers (Figure 14a) during the rockfall were determined manually  
639 from the analysis of the video images of the rockfall event on December 13, 2016. The  
640 time and space positions of the source between the selected markers are interpolated in-  
641 cluding small fluctuations using the  $10\times 10$  m spatial grid covering the observed rockfall  
642 trajectory, leading to a total of 200 impacts. The source can be activated at the same  
643 spatial position a number of times as the trajectory is spatially discretized by only 60  
644 cells.

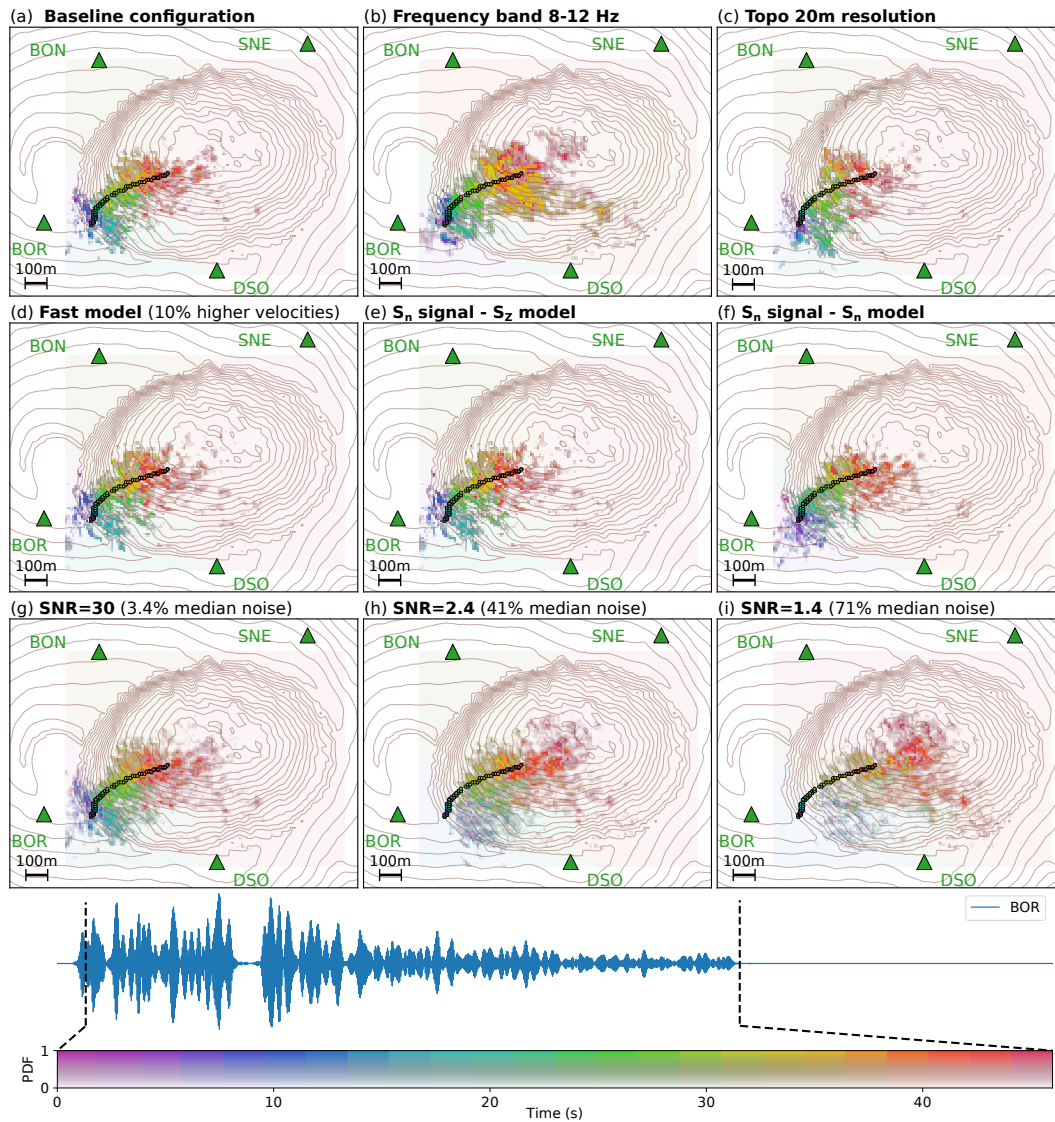
645 Synthetic seismic signals, hereafter designated the reference signals, are generated  
646 at the different stations from all the source positions and activation times, using the wave  
647 propagating model including topography. The source-time function (7 Hz Ricker wavelet)  
648 and the amplitude of the vertical traction are the same for all the impact sources. The  
649 corresponding generated signals can be seen in Figure A1b.

650 In the test, the location method is applied using: 1) the same topographic model  
651 used for the generated reference signals, in a high-frequency (13-17 Hz, defined as base-  
652 line model) and a low-frequency (8-12 Hz) band of the reference signals (Figures 13 a and  
653 b, respectively); 2) a wave propagating model including a low-pass filtered topography  
654 with a 30 m corner wavelength reducing the topography resolution from 10 m to 20 m (Fig-  
655 ure 13c); and 3) a wave propagating model including the original topography but a ve-  
656 locity model increased by 10% (Figure 13d). Further, the influence of the assumption  
657 of vertical rockfall impacts is tested by synthesizing reference signals generated by source  
658 impacts normal to the slope and locating them using the same wave propagation model  
659 as above from vertical sources (Figure 13e) and a wave propagation model from slope-  
660 normal sources (Figure 13f). Finally the method is tested after adding different levels  
661 of white noise to the reference signal (Figure 13g to i).

662 Results in Figure 13a are expected to be best because the synthetic signals are an-  
663 alyzed using the same model used for their generation and because the best spatial res-  
664 olution is expected at high frequencies as already seen before. Nevertheless, in contrast  
665 to a single-point source which could be located exactly (Fig. 10a), we can observe a cor-  
666 ridor of high probability that extends to up to 200 m. This is related to the superposi-  
667 tion of signals from temporally overlapping sources, which compromises the predictions.  
668 Still, the predictions correctly follow the progressively downhill moving active source re-  
669 gion.

670 The second low-frequency test, Figure 13b, demonstrates that valuable informa-  
671 tion is also contained in the low-frequency band, even though larger-wavelengths result  
672 in lower spatial resolution than when using high frequencies, extending the high prob-  
673 ability corridor especially at later times to up to 300 m. The contained information still  
674 suggests that developing methods that can exploit information across different frequency  
675 bands would be a major improvement, but is beyond the scope of this study.

676 In the third test, Figure 13c, in which the forward modeling part of the location  
677 method includes a smoother representation of the topography, inferred source locations  
678 are shifted compared to those inferred from first step, e.g. the source positions between



**Figure 13.** Location of a synthetic signal for a single rock moving down-slope. Color-filled circles mark the space-and-time positions of the vertical point impacts. The bottom graph shows the generated seismogram (vertical velocity) recorded at BOR and filtered at 13-17 Hz, resulting from a total of 200 impacts (the low amplitude gap in the signal is random and corresponds to the gap at around 12s in the offset-delay distribution in Figure 14a). Signals generated at all stations for all components are shown with scales in Figure A1. (a) Baseline configuration for location at 13-17 Hz using a wave propagation model with a topography resolution of 10 m and all station-channel pairs ( $N_{Sta} = 7$ ). (b) Location in a lower frequency band at 8-12 Hz. (c) Location using a wave propagation model with a topography resolution of 20 m. (d) Location using a wave propagation model with a 10% faster medium velocity. (e) Location of a synthetic rockfall signal from sources that are directed normal to the topography, referred to as  $S_n$ , using a wave propagation model with vertical sources  $S_z$ . (f) Location of a synthetic rockfall signal from sources which are directed normal to the topography, referred to as  $S_n$ , using a wave propagation model with normal sources  $S_n$ . (g) Location of a reference signal contaminated by a median white noise level of 3.4%, corresponding to a median signal-to-noise ratio  $SNR \approx 30$ , comparable to the noise level observed in this frequency band for the previously analyzed and relatively small rockfall on December 13, 2016 (see Appendix A). (h) Location of reference signal with median  $SNR \approx 2.4$  (Figure A2b). (i) Location of reference signal with median  $SNR \approx 1.4$  (Figure A2c).



679 20 s and 30 s are shifted by around 50 m towards the south. This stresses the importance  
 680 of properly resolving the topographic effects at the scale of the frequency bands that are  
 681 analyzed. Besides the already discussed superposition of multiple sources at different lo-  
 682 cations, the observed southwards shift of the predicted trajectory in section 4.1 can partly  
 683 be interpreted as possibly resulting from an inaccurate outdated DEM given that the  
 684 surface topography of the Dolomieu crater is continuously reshaped by high rockfall ac-  
 685 tivity (e.g. Hibert, Mangeney, et al., 2017; Durand et al., 2018; Derrien et al., 2019). Which  
 686 of the two effects is stronger is inherently dependent on the location of the rockfall and  
 687 the relative positions and magnitudes of the inferring sources, and cannot be predicted  
 688 in a general way.

689 To better understand the influence of the velocity model for the source predictions,  
 690 the location method is applied using a wave propagation model including the original  
 691 topography but with a modified velocity model in which velocities are globally increased  
 692 by 10 %, Figure 13d. This also influences intrinsic attenuation by decreasing the velocity-  
 693 dependent absorption coefficient (e.g. Aki & Richards, 2002). The inferred source loca-  
 694 tions don't differ significantly from the best reference test with the original velocity model,  
 695 Figure 13a, in the same frequency band. This might appear to be surprising as in this  
 696 test the forward modeling part of the location method is computed using the same to-  
 697 pography resolution but with a different velocity model. However in this modified ve-  
 698 locity model, seismic velocities are uniformly increased by 10%, which does not signif-  
 699 icantly alter the energy ratios between the different stations. More systematic scenar-  
 700 ios, including possible spatially localized velocity perturbations and local site effects at  
 701 the stations, need to be investigated in the future to properly assess the influence of the  
 702 a priori uncertainties in the seismic velocity model on the performance of the location  
 703 method. In the case that information about the subsurface properties is available, it can  
 704 be considered in the spectral-element based 3D propagation model and can therefore taken  
 705 into account in the proposed location method. This is in contrast to other locating meth-  
 706 ods where the seismic velocity is used as a free parameter and optimized during the lo-  
 707 cating process to maximize the correlation between stations (e.g. Burtin et al., 2013; Hi-  
 708 bert et al., 2014; Dietze et al., 2017; Pérez-Guillén et al., 2019), therefore three-dimensional  
 709 velocity models cannot be considered.

710 The energy ratios from the wave propagation model are generally computed un-  
 711 der a vertical source assumption, even though a rockfall can generate forces normal and  
 712 tangential to the slope. Kuehnert et al. (2020) showed that wave propagation along the  
 713 topography dominates over the source direction, suggesting that the source direction and  
 714 the resulting radiation patterns are a second order effect for location. To support this  
 715 assumption and verify that non-vertical forces generated from the rockfall on the ground  
 716 can indeed be ignored, a new reference signal is generated from sources directed normal  
 717 to the slope, which we refer to as the  $S_n$ -signal. In the test, the location method is then  
 718 applied using a wave propagation model with vertical sources, referred to as the  $S_z$ -model,  
 719 Figure 13e; and using a wave propagation model with normal sources, referred to as the  
 720  $S_n$ -model, Figure 13f.

721 Results from the first test with the vertical source assumption, Figure 13e, don't  
 722 differ significantly from the previous test with a reference signal from vertical sources,  
 723 Figure 13a, suggesting that the direction of the rockfall source impact does not influence  
 724 the performance of the location method.

725 This conclusion is further supported by the second test using a wave propagation  
 726 model with normal sources, Figure 13f, where the outline of the predicted source distri-  
 727 bution is displaced by a maximum of 50 m compared to the previous distribution in Fig-  
 728 ure 13e. Given the unpredictability of the source field in the case of real rockfalls, we con-  
 729 clude that the vertical source assumption is very reasonable and the most straightfor-  
 730 ward solution when predicting rockfall trajectories with the proposed location method.

731 Finally, noise contaminated reference signals were located. In Figure 13g, the noise  
 732 causes the probability distribution to be slightly blurred with location probabilities re-  
 733 ducing by around 10 % and the width of the spatial distribution increasing by around

734 100 m compared to the noise-free test in Figure 13a. Here, the added white noise is sim-  
 735 ilar in amplitude as observed at 13-17 Hz in the signals from Dolomieu crater, i.e. a me-  
 736 dian noise level of 3.4 % or median signal-to-noise ratio of  $SNR = 30.0$  for the rockfall  
 737 on December 13, 2016, which does not comprise large volumes, as can be seen in Fig-  
 738 ure 6. Higher noise levels, Figure 13h and i, increasingly blur the predicted source prob-  
 739 ability distribution with location probabilities reducing by around 30 % and 70 %, and  
 740 the width of the spatial distribution increasing by around 200 m and 400 m, respectively,  
 741 compared to the noise-free test in Figure 13a. The tests suggest that the location method  
 742 is robust to noise levels many times higher than those observed at Dolomieu crater, and  
 743 that at an average SNR of 2.4, the rockfall trajectory can still be tracked reasonably well  
 744 with an error of about 200 m.

745 **5.4 Distributed point sources mimicking complex rockfalls and granu-**  
 746 **lar flows**

747 To increase complexity, the downward moving seismic source is activated twice with  
 748 a respective time shift of 20 s, as shown in Figure 14a where the top graph shows the space-  
 749 time trajectory of two successive boulders. As a consequence of the respective time shift,  
 750 the first boulder arrives at the bottom of the crater while the second boulder is still lo-  
 751 cated in the top half of the crater wall, visible by the red-filled circles for times  $> 40$  s  
 752 in the map of Figure 14a.

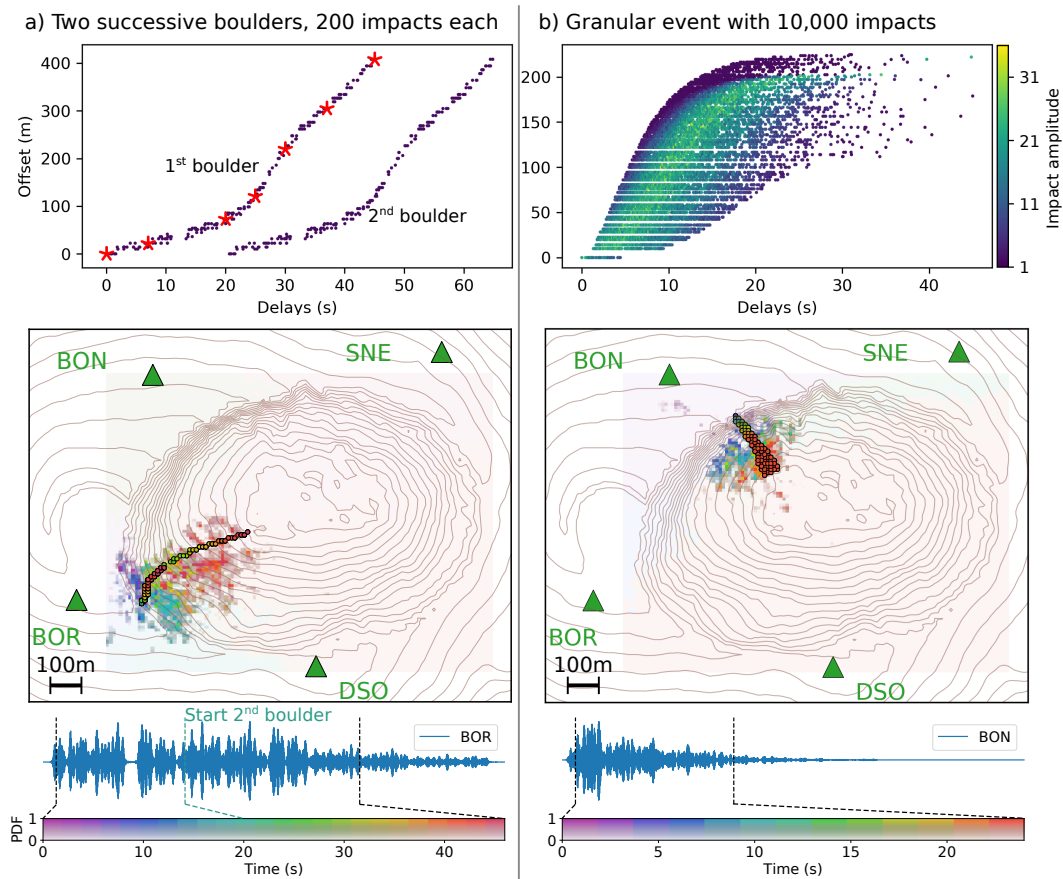
753 Location results show a high-probability corridor of around 200 m width compa-  
 754 rable with the probability distribution from the single-boulder test in Figure 13a. How-  
 755 ever, the superposition of the two simultaneously acting sources compromises the time  
 756 resolution of the method, i.e. the color sequence is mixed so that, for example, the red  
 757 color for times  $> 40$  s is strongly scattered and spread along almost the whole crater wall.  
 758 This loss of spatio-temporal resolution due to superposition of multiple sources was al-  
 759 ready observed in the previous test with single sources P1 and P2, Figure 10, and partly  
 760 explains the poor spatio-temporal location of the rockfall on February 28, 2016, located  
 761 on the southeastern crater wall, Figure 9b.

762 In a second test, a large distributed source with variable impact amplitudes was  
 763 constructed, aiming to synthesize the characteristics of a granular flow. The space-time  
 764 distribution of a total of 10,000 source impacts, presented in Figure 14b, is constructed  
 765 by defining a minimum and a maximum velocity curve as well as a third curve in between  
 766 where impact amplitudes are maximum. The total number of impacts reaches its max-  
 767 imum at around 10 s and decays subsequently towards zero after 45 s. The source area,  
 768 marked by color-filled circles on the map in Figure 14b, is spatially discretized by 87 cells  
 769 (selected from the  $10 \times 10$  m spatial grid), each of which can be activated multiple times  
 770 to simulate a total of 10,000 impacts. The corresponding generated signals are shown  
 771 in Figure A1d.

772 Despite the superposition of the numerous distributed seismic sources, high-probability  
 773 predictions are correctly located on the northwestern crater wall. The width of the prob-  
 774 ability distribution of up to 300 m is very similar to the one from the real granular event  
 775 on June 14, 2016, Figure 9d, and the global downward movement of the sources is well  
 776 captured and can be followed by means of the correctly ordered color sequence.

777 **6 Conclusion**

778 We propose a new rockfall location method based on seismic energy ratios between  
 779 stations. In an optimization routine, observed energy ratios are compared to a database  
 780 of simulated energy ratios in a region of interest. The benefit of the method is that once  
 781 the database has been created, locations can be estimated quickly without the need for  
 782 complicated analyses of the seismic signal such as precise picking of arrival times. The  
 783 rockfall seismic signals are analyzed in sliding time windows, making it possible to fol-  
 784 low the rockfall trajectory over time. The method can therefore potentially be used for



**Figure 14.** (a) The graph on top shows space-time distribution of two source trajectories mimicking two successive boulder tracks of 200 impacts each. Red asterisks mark source locations of the rockfall on December 13, 2016 as estimated from video images that serve as interpolation points. Under the map, the generated reference signal is shown, recorded at BOR and filtered at 13-17 Hz. The map shows picked source positions as well as location results at 13-17 Hz using a wave propagation model with a topography of 10 m resolution. (b) Space-time distribution of 10,000 sources mimicking a granular flow. The sources are distributed within two velocity curves. An additional curve inbetween defines sources of maximum impact amplitude. The amplitude is represented in arbitrary units with a minimum amplitude of 1. Under the map, the generated reference signal is shown, recorded at BON and filtered at 13-17 Hz. Signals generated at all stations for all components are shown with scales in Figure A1. Note that the signal amplitude is controlled by both the amplitude of each individual impact and the number of sources that act simultaneously. The map shows picked source positions as well as location results at 13-17 Hz using a wave propagation model with a topography of 10 m resolution.

785 continuous monitoring in real time, in parallel with existing methods that detect and clas-  
 786 sify rockfall seismic signals (e.g. Dammeier et al., 2016; Dietze et al., 2017; Provost et  
 787 al., 2017; Maggi et al., 2017; Hibert, Provost, et al., 2017; E. J. Lee et al., 2019).

788 By direct numerical modeling of the wave field on a domain representing the study  
 789 site, no assumptions about the wave type of the recorded signal are required, high-resolution  
 790 surface topography and its influence on the wave field can be accounted for, and a pri-  
 791 ori information about subsurface properties and a corresponding three-dimensional seis-

mic velocity model can be considered and are not required to be estimated during the location process.

Here, location was performed for rockfalls at Dolomieu crater, Reunion Island. All analyzed rockfall events could be located in the correct area of the crater. Generally, the best spatial resolution (below 100 m) is observed in the beginning of the rockfall when the seismic source is very confined in space. Thereafter, the predicted source locations become more scattered. This is linked to the spatial distribution of the seismic source, comprising multiple simultaneous impacts at different positions. The superposition of multiple sources is not considered in the method and hence compromises the location results. Nonetheless, the method performs remarkably well in this regard and is even able to locate a downward moving granular flow, likely because the signals are dominated by the signature of the most radiating sources at a given time.

It is shown that the influence of the assumed source impact direction on the location is of the second order, since propagation along the Dolomieu crater topography dominates over source-characteristic radiation patterns for the investigated frequencies above about 3 Hz. Thus, a vertical surface traction can be assumed, even though the actual source field of real rockfalls remains unknown. Furthermore, the insignificance of the source-characteristic radiation patterns makes it possible to use all vertical and horizontal component signals for location, which makes the method more robust against ambient noise or poorly known site amplifications.

Experiments with synthetic rockfall sources confirmed that the best spatial resolution is achieved at high frequencies (here in the frequency band of 13-17 Hz). For future development of the method, a combination of multiple and possibly also narrower frequency bands should be considered. The synthetic tests also revealed that a precise representation of the surface topography is crucial to the quality of the location results.

Investigations on the influence of the network geometry on the resolution suggests that best resolution (below 100 m) is achieved when the source area is triangulated by the seismic stations. The method currently assumes that the signal of a seismic source arrives fully within the defined time window at all stations. This is possible because of the positions of the seismometers with respect to the rockfalls at Dolomieu crater but might be a limitation for other source-receiver geometries. In order to overcome this limitation, a time shift can potentially be introduced at each station with respect to the region of interest after estimating the approximate arrival times.

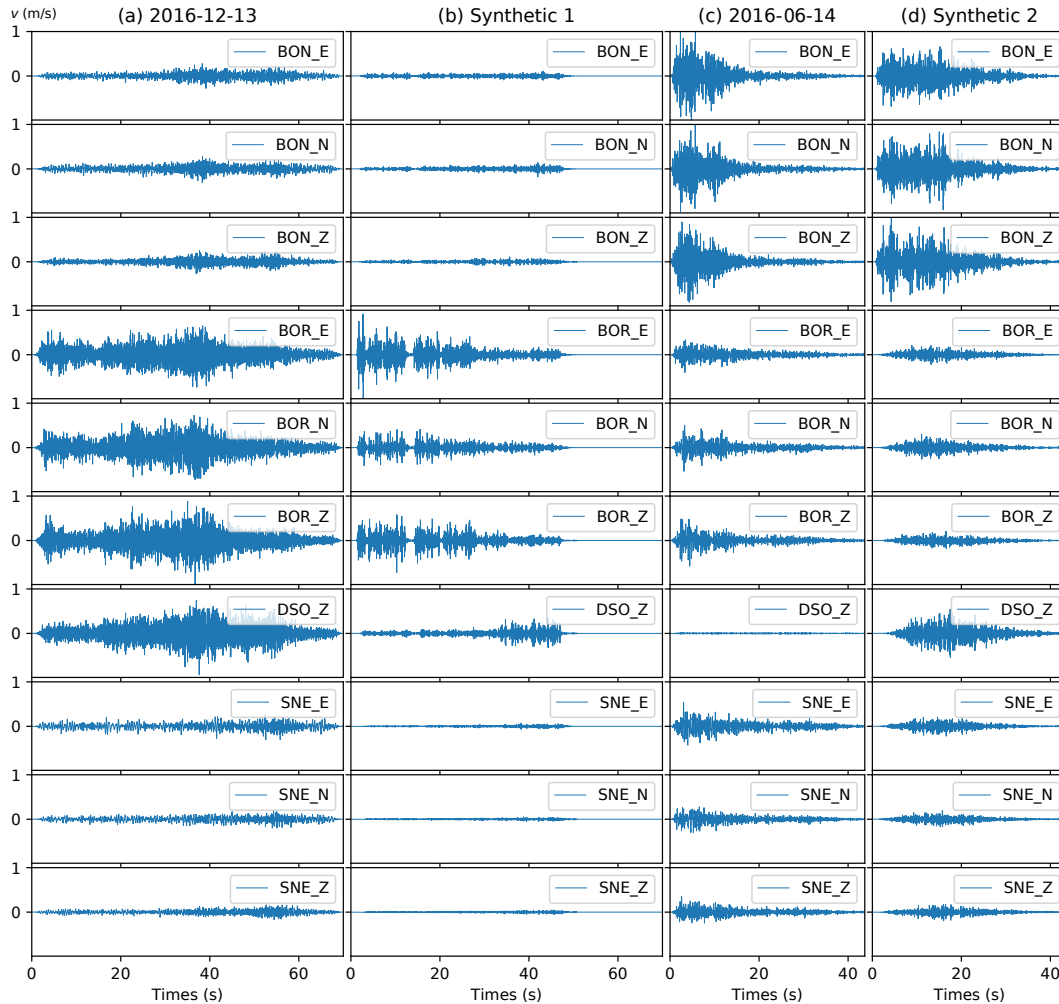
Comparisons with other location methods that are able to track moving seismic surface sources, as for example the approach of amplitude source location (ASL, e.g. Pérez-Guillén et al., 2019), need to be carried out at the same study site and using the same station network to assess the benefits of each method and compare their resolution.

Noise levels at Dolomieu crater are very low at the here studied frequencies above 3 Hz and could be ignored when locating the observed rockfalls. Tests with added white noise showed that the location method is robust against noise levels that are considerably higher than those observed at Dolomieu crater, and that using signals in the 13-17 Hz frequency band with an average SNR of 2.4, rockfall trajectories can still be tracked with an error of about 200 m.

No significant effects on the location results were found when modifying the subsurface velocity model. However, seismic velocities in the test were uniformly increased by 10%, which does not significantly alter the energy ratios between the different stations. More systematic scenarios, including possible spatially localized velocity perturbations and local site effects at the stations, will need to be investigated in future studies to properly assess the influence of the a-priori uncertainties in the seismic velocity model on the performance of the location method.

842 **Appendix A Comparison of real, synthetic and noise-contaminated syn-**  
 843 **thetic rockfall signals**

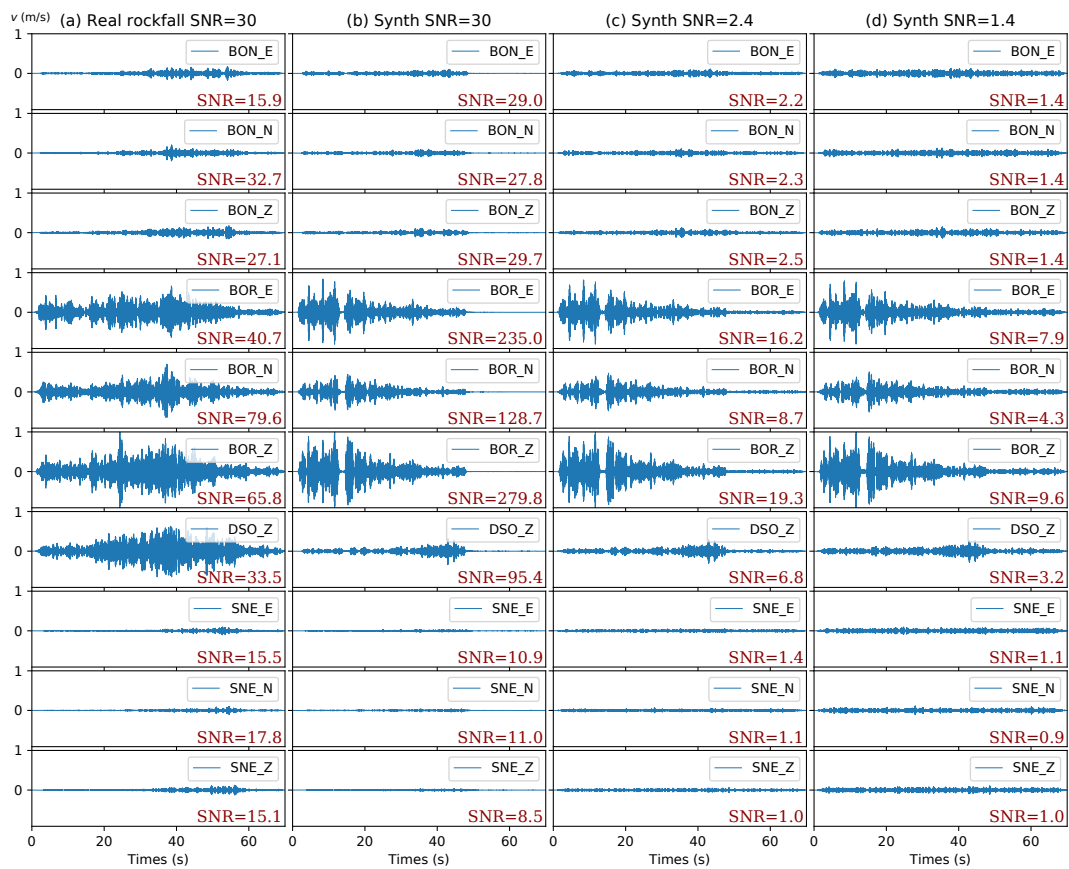
844 Seismograms generated by the above analyzed rockfall events on December 13, 2016  
 845 and on June 14, 2016 are shown in Figures A1a and c, respectively. They are shown to-  
 846 gether with synthetic rockfall signals, mimicking the real events, in Figure A1b for the  
 847 boulder-type event synthesized and analyzed in section 5.3 and in Figure A1d for the granular-  
 type event synthesized and analyzed in section 5.4.



**Figure A1.** Seismograms of all station channels for real rockfalls at Dolomieu crater and synthetic signals mimicking these events. Signals of each event are normalized and real seismograms are bandpass filtered at 1-35 Hz. (a) Signals of boulder-type event on December 13, 2016, analyzed in Figures 6 to 8. (b) Synthetic signals generated by 200 point impacts mimicking a down-slope moving rock, analyzed in Figure 13. (c) Signals of granular-type event on June 14, 2016, analyzed in Figure 9d. (d) Synthetic signals generated by 10,000 point impacts mimicking a granular flow, analyzed in Figure 14.

848 To test the location method with noise-contaminated synthetic signals, the noise  
 849 levels on the observed signals at Dolomieu crater are analyzed using the rockfall event  
 850 on December 13, 2016. Observed signals filtered at 13-17 Hz and corresponding signal-  
 851 to-noise ratios (SNR) are shown in Figure A2a. The minimum SNR is 15.1 while the me-  
 852

853 dian SNR is 30.0, which corresponds to a median noise level of 3.4%. For the relatively  
 854 small rockfall, the SNR is high, which indicates a low noise level at the Dolomieu crater  
 at these high frequencies.



**Figure A2.** Comparison of signal-to-noise ratios (SNR) for real and synthetic signals, filtered at 13-17 Hz and with normalized amplitudes. (a) Observed signal at Dolomieu crater generated by rockfall on December 13, 2016, with minimum SNR of 15.1 and median SNR of 30.0, which corresponds to a median noise level of 3.4%. (b) Synthetic signal from the test in section 5.3, contaminated with a median noise level of 3.4% (i.e. median SNR  $\approx 30$ ), similar to the level on the observed rockfall signal. (c) Synthetic signal contaminated with a median noise level of 41.5% (i.e. median SNR  $\approx 2.4$ ), 12 times higher than the noise level on the observed rockfall signal. SNR values are close to 1 at station SNE, hiding the signals almost entirely. (d) Synthetic signal contaminated with a median noise level of 71.0% (i.e. median SNR  $\approx 1.4$ ), 21 times higher than the level on the observed rockfall signal. SNR values are close to 1 at both station BON and SNE, hiding the signals almost entirely.

855 White noise is now added to the synthetic rockfalls in Figure A1b. Figure A2a shows  
 856 the synthetic signals contaminated with a noise level so that the median SNR  $\approx 30$ ,  
 857 comparable to the SNR observed at Dolomieu crater. Noise levels are then increased by  
 858 a factor of 12 and a factor of 21, resulting in the synthetic signals shown in Figures A2b  
 859 and c, respectively. A factor of 12 increases the median noise level to 41.5% (i.e. median  
 860 SNR  $\approx 2.4$ ), hiding almost entirely the signal at station SNE. A factor 21 increases  
 861 the median noise level to 71.0% (i.e. median SNR  $\approx 1.4$ ), hiding not only signals at  
 862 SNE, but also at BON.  
 863

864 **Acknowledgments**

865 We are very grateful to Emma Suriñach and two anonymous reviewer for their critical  
 866 and constructive remarks which contributed to improve the quality of the present pa-  
 867 per. We want to thank the whole team at the OVPF observatory that provided the ex-  
 868 cellent field data used in this study. We thank the whole team at the OVPF that pro-  
 869 vided the excellent field data for this study. The seismic data were acquired by the Vol-  
 870 canological and Seismological Observatory of Piton de la Fournaise (OVPF)/Institut de  
 871 Physique du Globe de Paris (IPGP) via the VOLOBSIS Portal: <http://volobsis.ipgp.fr/query.php>.  
 872 Camera data are stored in <http://doi.org/10.5281/zenodo.4031816>. Data from the sim-  
 873 ulations are available from <http://doi.org/10.5281/zenodo.3949826>. Numerical compu-  
 874 tations were partly performed at S-CAPAD (Service de calcul parallèle et de traitement  
 875 de données en sciences de la Terre), IPGP, France, as well as at CCIPL (Centre de Cal-  
 876 cul Intensif des Pays de la Loire), Université de Nantes, France. This work was funded  
 877 by ERC Contract No. ERC-CG-2013-PE10-617472 SLIDEQUAKES.

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