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Daval

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Luca Stigliano, Martina Siena, Philippe Ackerer, Alberto Guadagnini, Damien Daval. Statistical Description of Calcite Surface Roughness Resulting from Dissolution at Close-to-Equilibrium Conditions. ACS Earth and Space Chemistry, 2021, 10.1021/acsearthspacechem.1c00226 . hal-03382908

HAL Id: hal-03382908

<https://hal.science/hal-03382908>

Submitted on 18 Oct 2021

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Statistical description of calcite surface roughness resulting from dissolution at close-to-equilibrium conditions

Journal:	<i>ACS Earth and Space Chemistry</i>
Manuscript ID	sp-2021-00226q.R2
Manuscript Type:	Article
Date Submitted by the Author:	n/a
Complete List of Authors:	Stigliano, Luca; Universite de Strasbourg; Universite Grenoble Alpes Siena, Martina; Politecnico di Milano Ackerer, Philippe; Universite de Strasbourg Guadagnini, Alberto; Politecnico di Milano Daval, Damien; Universite de Strasbourg; Universite Grenoble Alpes

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6 **dissolution at close-to-equilibrium conditions**
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Abstract

Linking the evolution of the surface area (as quantified, e.g., through its spatial roughness) of minerals to their dissolution rate is a key aspect of mineral reactivity. Unraveling the nature of its main features requires relying on approaches yielding a quantitative characterization of the temporal evolution of surface topography/roughness. Here, a mechanically-polished {104} calcite surface was dissolved at room temperature and at close-to-equilibrium conditions ($\Omega=0.6$) with an alkaline solution (pH = 8) across a temporal window of 8 days. Surface topography images were acquired daily using vertical scanning interferometry, the ensuing topography data being then embedded within a statistical analysis framework aimed at describing comprehensively the surface roughness evolution. The strongest system variations were observed after one day: the probability density function of surface roughness was observed to transition from being approximately Gaussian to being left-skewed and leptokurtic, exhibiting a dramatic increase in the variance and a significant change in the semi-variogram structure. After a relaxation time of approximately two days, the reacting surface appeared to attain a steady-state configuration, being characterized by values of the statistical moments characterizing surface roughness which become virtually independent of time. Attempting to unravel the underlying dissolution mechanism, an original numerical model able to reproduce satisfactorily the statistical behavior observed experimentally was developed and tested. Our results suggest that, under the investigated conditions, dissolution may be characterized as a spatially correlated random process, with the areas most exposed to the flowing fluid being prone to preferential dissolution. The numerical model was also used to obtain insights on the influence of the initial surface roughness and of the fluid composition on the steady-state statistical characterization of the surface roughness. Our results suggest that the influence of the initial surface roughness is limited. The present study suggests that potential empirical relations linking the surface roughness of reacted crystals to the saturation state at which they

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3 dissolved may be developed, which would allow to back-estimate the reacting conditions only
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5 based on topography data.
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10 **Keywords:** Calcite, Near equilibrium dissolution, Surface roughness, Spatial statistics,
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1. Introduction

Chemical weathering is a central part of the Earth system, as it is one of the key processes that contributes to the redistribution of mass on Earth. The dissolution of rock-forming minerals controls the formation of soil and supplies nutrients to sustain ecosystems ¹, it is responsible for atmospheric CO₂ drawdown, regulating the long-term climate of the Earth ², and is involved in various geo-engineered strategies such as enhanced rock weathering for carbon dioxide removal from the atmosphere ³, radioactive waste storage ⁴, geothermal energy recovery ⁵, or CO₂ sequestration ⁶. The development of rate laws describing the dissolution kinetics of minerals is critical to an effective strategy for upscaling of mineral reactivity, with the ultimate goal of modeling fluxes associated with chemical weathering across a variety of space and time scales. The classical hypothesis underlying the rate laws that are employed to interpret mineral reactivity (through modeling frameworks based on either a bottom-up (e.g., ⁷) or a top-down (e.g., ⁸) approach) rests on a conceptual picture according to which the dissolution flux is proportional to the contact surface area between the fluid and the solid.

While such a simplification rests on the assumption that reactivity is immutable in time and homogeneous at the surface of the minerals, the heterogeneous nature of the surface reactivity appears as a universal feature of solid dissolution, as documented through a variety of materials degradation studies ranging from the corrosion of metals ⁹⁻¹¹, glasses ¹², and ceramics ¹³⁻¹⁶ to the weathering of rocks and minerals (e.g., ¹⁷⁻²⁸). At the heart of the heterogeneous nature of surface reactivity is the concept of *surface energy*, and its distribution in space and time throughout the evolution of the dissolution process (e.g., ^{15, 29}), which is then mirrored by the gradual development of surface roughness. The heterogeneous distribution of structural defects and the seemingly random occurrence of hydrolysis events at the mineral surface drive variability of surface reactivity, the latter possibly varying across two orders of magnitude for a single mineral (e.g., ^{15, 29}). These observations thus support the need and opportunity for the

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2
3 development of an alternative approach for the assessment of mineral reactivity focused on the
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5 heterogeneous space-time dynamics of surface roughness.
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8 As an example, the rate spectra approach ²⁰, is based on the evaluation of the probability
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10 density function of the dissolution rates at each given location of the mineral surface. From a
11
12 practical standpoint, spatially-resolved dissolution rates at the surface of a given material can
13
14 be quantified upon monitoring the time-resolved evolution of surface topography resulting from
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16 dissolution (e.g., ^{15, 20, 22, 27, 29-33}). One of the most promising aspects of the rate spectra approach
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18 is that the deconvolution of a rate spectrum provides information about the modes (i.e., the main
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20 contributors to the overall rates), which can then be compared against outputs of atomic-scale
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22 simulations such as kinetic Monte Carlo modeling, and refine our knowledge of the
23
24 corresponding reaction mechanisms (e.g., ^{34, 35}).
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29 A notable limitation of the rate spectra approach is that determining the distribution of
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31 dissolution rates at the mineral surface requires analyzing across time a laterally fixed and
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33 motionless surface. Such measurements can only be achieved under well controlled laboratory
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35 conditions, using techniques such as *in situ* atomic force microscopy (AFM) conducted in a
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37 fluid reaction cell (e.g., ³⁰) or vertical scanning interferometry (VSI) with a reaction cell fixed
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39 to the VSI stage (e.g., ^{32, 36} and references therein), so that a given dissolution rate can be
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41 assigned to each pixel (or group of pixels) depicting the topography data collected for a given
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43 field of view. When considering crystals subject to weathering in the field (e.g., ³⁷) or samples
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45 with a low reactivity (which makes them incompatible with measurements of the dissolution
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47 progress at room temperature over short time periods), obtaining insights into the temporal
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49 evolution of the key statistics of the overall roughness of a given surface remains of significant
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51 interest. As documented by Fischer et al. ¹⁹, the investigation of surface roughness evolution
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53 could provide information about the reaction history across the dissolution process. As such,
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55 time-resolved statistical analyses of dissolving mineral surfaces are of critical importance, as
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3 they underpin the development of empirical and/or theoretically-supported relationships
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5 between surface topography and the extent of reaction, or between the surface topography and
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7 the reaction conditions.
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10 Pollet-Villard et al.³⁸ showed that a statistical treatment of etch pit nucleation and growth
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12 monitored by AFM could be used to successfully develop a model accounting for the long-term
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14 evolution of the (001) surface of a dissolving K-feldspar at far-from-equilibrium conditions.
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16 By tuning the model parameters based on the comparison between the statistics of modeled and
17
18 measured surface topography at the early stage of dissolution, these authors could forecast the
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20 subsequent retreat rate of the surface, as well as evaluate, through a sensitivity analysis, the
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22 impact of defect density (i.e., screw dislocation) on K-feldspar dissolution rates. To the best of
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24 our knowledge, a detailed analysis combining statistical approaches and process models to
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26 characterize the evolution of the surface roughness of a crystal dissolved at close-to-equilibrium
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28 conditions, where the nucleation of etch pits is no longer spontaneous, is still lacking. The
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30 present study aims at bridging this gap by providing a detailed statistically-based analysis of
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32 the topography evolution of calcite surface dissolved at close-to-equilibrium conditions. In their
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34 preliminary study, Siena et al.³⁹ demonstrated that single snapshots of the spatial distribution
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36 of surface topography of calcite under such conditions can be successfully characterized by
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38 jointly modeling the probability distribution functions of surface heights and their spatial
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40 increments as a Generalized Sub-Gaussian (GSG) random field. Since single snapshots of
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42 surface roughness are not sufficient to predict or reconstruct reaction rates during dissolution
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44 (see, e.g.,²²), the present study extends their work by providing a time-resolved analysis of the
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46 evolution of surface topography. We then assess the ability of three empirical numerical models
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48 of dissolution to reproduce the observed behavior of the key statistical descriptors associated
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50 with the experimental data. The model which is most successful for reproducing the observed
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52 statistical description of the system is then employed to assess the relative impact of initial
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3 surface roughness and fluid composition on the statistical evolution of the surface. The outputs
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5 of the simulations suggest that a bijective relationship might exist between saturation state and
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7 the steady-state statistics of surface topography data.
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10 **2. Materials and methods**

11 12 13 **2.1 Materials and experimental methods**

14 15 **2.1.1 Sample preparation**

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17 The fragment used in the present study was obtained from an optically transparent calcite
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19 sample originating from the Massif des Ecrins (Alps, France). The chemical composition of the
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21 sample was determined, as in Bouissonnié et al. ²⁵, from 3 g of calcite by loss on ignition and
22
23 alkaline fusion, and only very few impurities (0.2 wt. % MgO, 0.008% MnO, and trace amounts
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25 of Sr) were revealed. The calcite sample was cleaved following the natural {104} calcite plane
26
27 to obtain a 3×5 mm² size fragment (Figure 1). The crystal was then embedded in epoxy resin to
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29 be subjected to a multi-step abrasive grinding (the finest grit size being 8.4 μm) to establish an
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31 initially flat surface and to remove the effects of sample cutting, followed by a polishing
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33 sequence (the finest grit size being 0.25 μm), using an aqueous solution saturated with respect
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35 to calcite as lubricant to avoid etching of the surface. More details are provided as Supporting
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37 Information (Table S1). After polishing, the crystallographic orientation of the sample was
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39 verified over the entire surface using electron backscatter diffraction (EBSD) on a TESCAN
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41 VEGA II scanning electron microscope (SEM). The sample was analyzed as it was (i.e., without
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43 carbon coating) and no grain boundaries were observed. The resin was then removed using a
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45 precision saw and the crystal surface was cleaned with ethanol. Finally, a small portion (~ 1-2
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47 mm²) of the surface of the crystal was masked with RTV glue spots (Figure 1) to provide a non-
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49 reacting reference surface to enable one to measure the height difference (i.e., surface retreat)
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51 between the reference and the unmasked portions using VSI (see Section 3.1). The average
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53 roughness (R_a), for a given observation window, is defined as:
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$$R_a = \frac{1}{N} \sum_{i=1}^N |Z_i(\mathbf{x}) - \langle Z \rangle| \quad (1)$$

where, $Z(\mathbf{x})$ corresponds to the height (evaluated with respect to a given reference) of the cell whose centroid is associated with vector position \mathbf{x} , $\langle Z \rangle$ is the average height (evaluated with respect to the same reference) and N is the total number of sampled height data. Based on VSI data ($50\times$ magnification) collected over $100\times 100 \mu\text{m}^2$ areas and following a second-order degree surface de-trending, which was required to minimize biases stemming from the tilt caused by the VSI acquisition and from the waviness of the sample resulting from polishing (see Section 2.1.5), values of R_a were found to range between 1.0 and 1.3 nm.

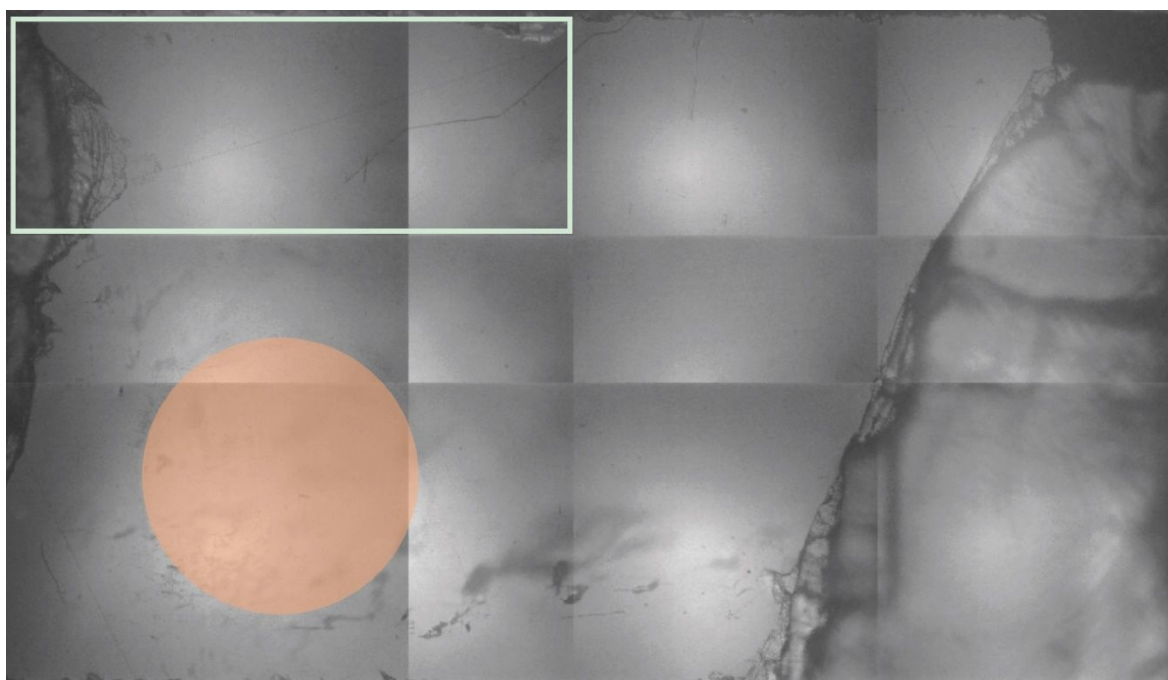


Figure 1 Intensity map of the polished $3\times 5 \text{ mm}^2$ size fragment employed in the experiment as imaged by VSI at $5\times$ magnification (corresponding to a lateral resolution $dl = 2.2 \mu\text{m}$) after 188 hours of dissolution. The orange area indicates where the mask was placed during the experiment before its removal. The highlighted rectangle corresponds to the region where VSI data were acquired also at $50\times$ magnification (corresponding to a lateral resolution $dl = 0.22 \mu\text{m}$; see also Figure 2).

2.1.2 Aqueous solution preparation

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3 Reagent-grade NaCl, NaHCO₃, and CaCl₂ were added to ultra-pure water equilibrated with
4 atmospheric CO₂ (18MΩ·cm) at concentrations similar to those documented in previous the
5 studies (40 and 25) to provide a baseline for the analysis of rate data. The concentration of NaCl
6 was set to 5.85 g/L to yield the ionic strength to approximately 0.1 molal, and that of NaHCO₃
7 was set to 0.095 g/L to fix the alkalinity. The measured pH of the resulting solution was 8.0 ±
8 0.1. The concentration of CaCl₂ was adjusted to maintain a constant saturation index of the
9 solution with respect to calcite ($\Omega_{calcite}$), defined as:

$$\Omega_{calcite} = \frac{a_{Ca^{2+}} a_{CO_3^{2-}}}{K_{sp,calcite}} \quad (2)$$

10 where $a_{Ca^{2+}}$ and $a_{CO_3^{2-}}$ are calcium and carbonate ion activities, respectively, and $K_{sp,calcite}$ is
11 the solubility product of calcite. Since the nucleation of etch pits on the {104} calcite faces is
12 no longer spontaneous for $\Omega > 0.45$ 25, the experiment was conducted at a saturation index of
13 0.60 ($[Ca^{2+}] = 0.00135$ molal).

14 2.1.3 Experimental approach

15 The experimental set-up was analogous to the one used in Bouissonié et al. 25. In brief,
16 the experiment was conducted in a mixed-flow reactor set-up at room temperature ($T = 22 \pm 1$
17 °C). A calcite fragment of size ~ 10 mm³ (which is small in comparison with the volume of the
18 fluid in the reactor, i.e., ~ 100 mL) has been used to maintain a negligible concentration of Ca
19 released throughout calcite dissolution as compared to the background concentration in the inlet
20 solution, thus maintaining a nearly constant saturation index across the reaction progress. The
21 experiment was conducted at a constant flow rate (0.25 mL/min) for a total duration of 188 h
22 (~ 8 days). The peristaltic pump was periodically stopped (every ~ 23 hours) to recover the
23 calcite fragment, clean it with ethanol, and analyze the evolution of the surface roughness as a
24 function of time through VSI. After each analysis, the sample was placed back into the reactor
25 and the pump was restarted for additional 23 hour cycles until the experiment was ultimately

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3 stopped. The aqueous solution was also regularly sampled (once a day) to evaluate pH and Ca
4 concentration, the latter being measured with an inductively coupled plasma atomic emission
5 spectroscopy (ICP-AES, Thermo iCAP 6000 Series). Based on these fluid analyses, the
6
7 saturation index was evaluated to be 0.58 ± 0.01 over the duration of the experiment. ICP-AES
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9 analyses are given in the Supporting Information (Table S2).

14 2.1.4 VSI analyses

16 As emphasized by Fischer et al.¹⁹, a robust statistical analysis of surface roughness requires
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18 relying on an experimental method that yields a sufficiently high vertical resolution (similar to
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20 that provided by AFM) together with a large field of view (e.g., comparable to that provided by
21
22 light optical microscopy). The technical specifications of VSI, for which the vertical resolution
23
24 can range between 0.5 and 2 nm and fields of view can be as large as several tens of mm² in
25
26 stitching mode, are fully compatible with this goal. Measurements of surface topography, $Z(x,$
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28 $y)$, (x, y) being spatial coordinates in the horizontal plane, were collected using a Zygo
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30 NewView 7300 VSI. As roughness is a scale-dependent quantity¹⁹, all data were acquired and
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32 processed at a fixed magnification (50×) over a collection of 9 regions (see Figure 2), each
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34 associated with the same planar area and with lateral resolution $dl = 0.22 \mu\text{m}$. In addition,
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36 topography measurements acquired at 5× across the whole crystal surface were used to evaluate
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38 the overall dissolution rate of calcite over the duration of the experiment according to (see e.g.,
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$$r = \frac{\Delta h}{\Delta t V_m} \quad (3)$$

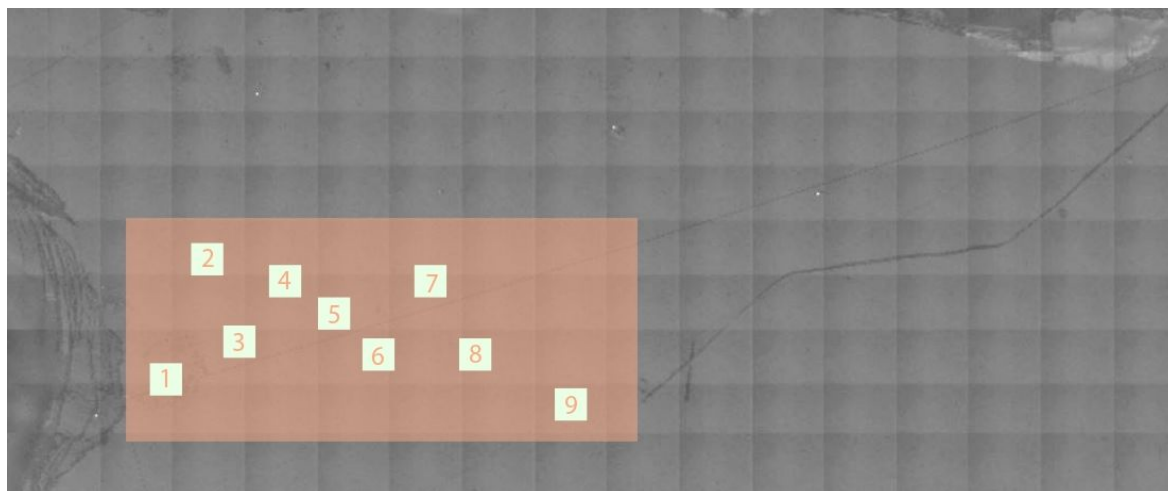
51 where r (mol/m²/s) is the overall dissolution rate, Δh (m) is the spatially-averaged height
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53 difference (i.e., surface retreat) between the unreacted reference surface and the reacted mineral
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55 surface, Δt (s) is the time interval over which the height difference is evaluated, and V_m
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57 (m³/mol) is the molar volume of calcite.

2.1.5 Topography data processing

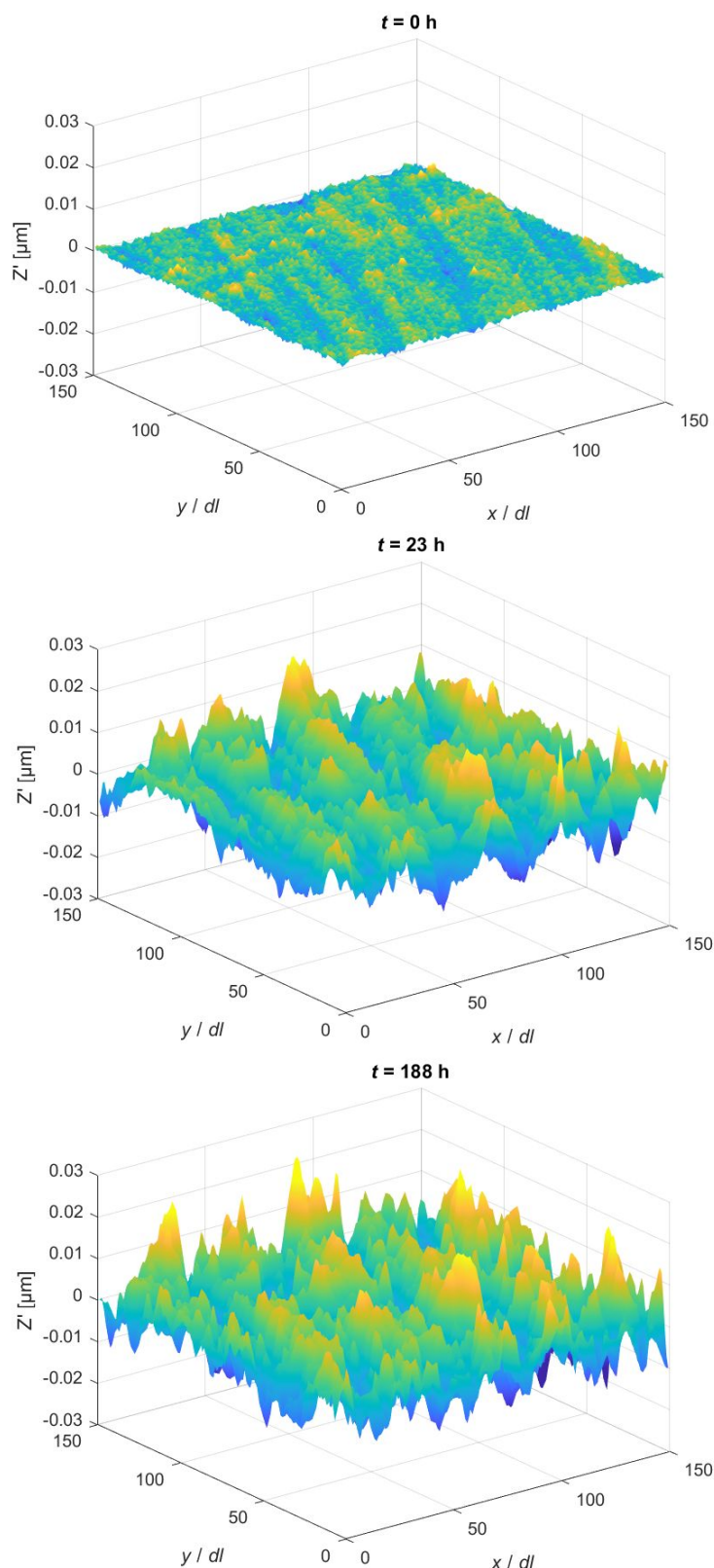
The analysis of topography data is focused on the $710 \times 325 \mu\text{m}^2$ area delineated in Figure 2. This region has been identified as the most suitable for this study, as it has the lowest density of surface defects (such as macro-steps or cracks), which may lead to optical artifacts that would strongly affect the statistical analyses. Thus, the only features characterizing the initial surface in the selected region were polishing scratches. Nine sub-regions of size $33 \times 33 \mu\text{m}^2$ have been selected to provide a representative statistical description of the roughness of this area while attempting to avoid most of the VSI-related artifacts (e.g., artificial spikes arising as a result from a sharp (sub-vertical) height transition between two adjacent locations), which would lead to an overestimation of the average surface roughness parameter R_a ⁴². As such, for all 9 sampled regions, and at all available times, the total amount of ‘dead pixels’ (i.e., pixels for which the VSI was not able to acquire a height value) and of VSI-related artifacts was negligible (< 0.1%) as compared to the total number of data. Each of these ‘dead pixels’ and VSI-related artifacts was replaced by the arithmetic mean of the height values of the 8 neighboring pixels. If one (or more) neighboring pixel was either a ‘dead pixel’ or a VSI-related artifact, then these were excluded from the calculation of the mean value. The following two criteria have been considered for the selection of the sample window size: (i) convergence of the average surface roughness parameter R_a ¹⁹, which was observed to remain approximately constant within windows ranging from 25×25 to $35 \times 35 \mu\text{m}^2$; and (ii) the presence of a sufficient number of data for a statistical analysis.

We note that the raw topography data in the sampled regions could not be analyzed as acquired. Identification of the actual surface roughness required separating the effects of (i) the waviness generated by the polishing process and (ii) the tilt of the sample support applied during the acquisition phase to adjust the alignment between the sample surface and the VSI microscope objective. For this purpose, the topography data of each sample window at each

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3 observation time have been de-trended by removing the best-fitting second order degree surface
4 from all data points (Figure 3). Otherwise, polishing scratches have been considered as part of
5 the initial surface roughness (and, hence, were not removed with data processing) to be able to
6 observe their influence on the surface roughness evolution.
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29 **Figure 2** Intensity map of the portion of the crystal surface imaged at 50× magnification
30 (depicted in Figure 1). The 710×325 μm^2 area delineated in the figure (orange color box)
31 represents the region identified as most suitable for the study and contains the 9 square sub-
32 regions (33×33 μm^2) investigated across time during dissolution. The initial average surface
33 roughness was evaluated (at 50× magnification) to range between 0.6 and 0.8 nm across the 9
34 sampling windows.
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54 **Figure 3** Exemplary three-dimensional (3D) representation of the mean-removed topography
55 data, Z' , (see Section 2.2) associated with a selected sample window, after replacement of 'dead
56 pixels' and VSI-related artifacts (see Section 2.1.5), and second order degree de-trending at
57 observation times $t = 0, 23,$ and 188 hours.
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2.2 Statistical approaches

A statistical characterization confined to the univariate analysis of the frequency distribution of heights shadows the information associated with the degree of spatial correlation of the data. We rely on statistical tools to characterize the degree of spatial correlation of surface topography by (i) interpreting the investigated random field, $Z(\mathbf{x})$, as the sum of an ensemble mean, $\langle Z \rangle$, and a zero-mean fluctuation, $Z'(\mathbf{x})$; and (ii) evaluating key statistical features of both $Z'(\mathbf{x})$ and spatial increments, $\Delta Z(\mathbf{h}) = Z(\mathbf{x}) - Z(\mathbf{x} + \mathbf{h})$, \mathbf{h} being separation distance (or lag).

We characterize the temporal evolution of sample probability density functions of Z' (Figure 4) through the time-resolved analysis of the corresponding sample variance, skewness (Equation 4), and kurtosis (Equation 5). We quantify the temporal variation of the degree of spatial correlation of the roughness field through the analysis of the experimental semi-variogram (Equation 6).

We recall that the skewness and the kurtosis of a probability density function are defined as:

$\text{skew} = \frac{\sum_i^M (z_i - \langle Z \rangle)^3}{(M - 1) s^3}$	(4)
$\text{kurt} = \frac{\sum_i^M (z_i - \langle Z \rangle)^4}{(M - 1) s^4}$	(5)

where z_i is the i -th observed value of Z , $\langle Z \rangle$ is the ensemble average, M is the number of available observations of Z , and s is the standard deviation of Z . We recall that skewness is a measure of the degree of asymmetry of a probability distribution, whereas kurtosis imbues the combined weight of the tails relative to the rest of the distribution. We recall that a Gaussian distribution is characterized by a skewness and kurtosis equal to 0 and 3, respectively. A

negative skewness indicates a left skewed distribution; a kurtosis lower or higher than 3 denotes a distribution with tails which are respectively thinner or heavier than Gaussian.

The experimental semi-variogram is defined as half of the sample variance of the spatial increments, i.e.,:

$$\gamma(h) = \frac{1}{2 n(h)} \sum_{i=1}^{n(h)} [Z(x_i) - Z(x_i + h)]^2 \quad (6)$$

where $n(h)$ is the number of data pairs separated by a given omni-directional lag $h = |\mathbf{h}|$. As stated above, it is a measure of the spatial dependency of Z .

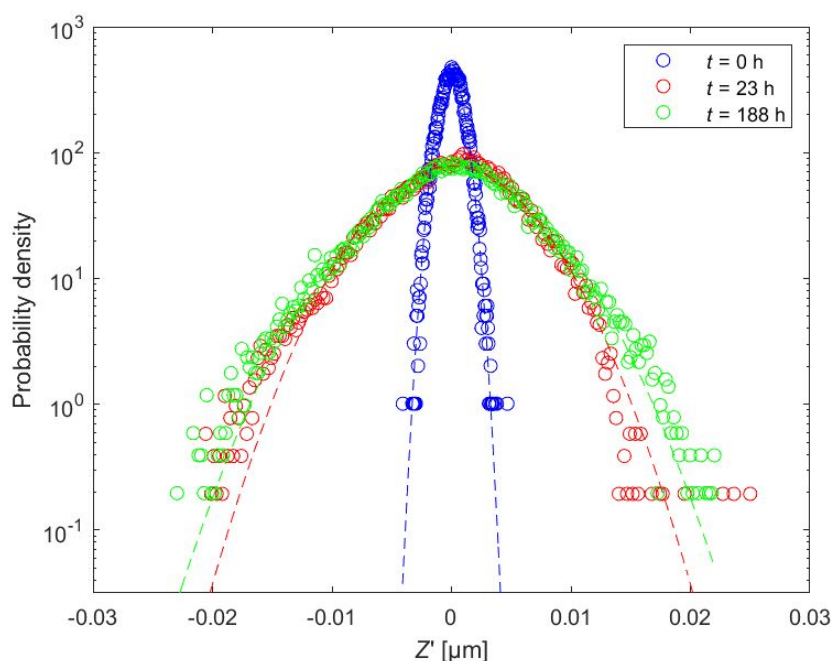


Figure 4 Comparison of the sample distributions probability densities of the variable Z' at observation times $t = 0, 23,$ and 188 hours (open circles) with Gaussian distributions densities (dashed lines) characterized by the same mean and standard deviation of Z' . The distribution of Z' appears to be symmetric at $t = 0$ hours and slightly asymmetric (left-skewed) at longer times.

2.3 Numerical modeling

2.3.1 Numerical approach

In addition to providing a statistical description of the evolution of the surface roughness resulting from dissolution, we also aim at comparing such depiction to results associated with three dissolution models, as described in the following. The objective of the development of

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3 these models is two-fold: (i) demonstrating that a statistical characterization of surface
4 roughness evolution may help to rule out some of the existing models describing the
5 mechanisms of mineral dissolution; (ii) providing a qualitative assessment of the impact of
6 input parameters such as initial surface topography and fluid composition on the calculated
7 statistical descriptors.
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12 Each of the three dissolution models we analyze is characterized by a given mathematical
13 rendering of dissolution processes, which are viewed in a probabilistic context. The parameters
14 of each of these models (see Section 2.3.2) were adjusted to assess whether spatial statistics
15 based on modeling results were compatible with the corresponding statistics based on the
16 experimentally-observed temporal evolution of calcite roughness. The models considered differ
17 in the way one assigns the probability that dissolution takes place at a given location on the
18 simulated surface, as described in the following (additional details being provided in Section
19 2.3.2):
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33 (i) Uniform and spatially independent dissolution probability (hereafter termed *Model*
34 *1*): the occurrence of dissolution at each location of the simulated crystal surface
35 and at each time is associated with a given probability, the latter being uniformly
36 distributed across the system, independent of location. Initiation of dissolution at a
37 given location is therefore unpredictable and independent of the neighborhood. Such
38 a conceptual framework is typically common to *pitting corrosion* models ^{9, 10}.
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46 (ii) Equiprobable dissolution with presence of volume defects dissolving preferentially
47 (hereafter termed *Model 2*). At close-to-equilibrium conditions, the strain field of a
48 dislocation does not open up hollow cores to form etch pits (e.g., ⁴³) and dissolution
49 is driven by point defects, pre-existing edges, and corners ³⁶. As such, whereas the
50 dissolution taking place at defect-free zones is considered to be equally likely,
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3 locations termed as ‘defects’ are associated with a higher probability to dissolve,
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5 this being intended to simulate such highly localized point defects.
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- 8 (iii) Dissolution probability of individual locations in the (x, y) plane controlled by fluid
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10 accessibility (hereafter termed *Model 3*). The model is based on a conceptual picture
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12 according to which dissolution is locally proportional to the contact surface between
13
14 the crystal and the fluid ⁴⁴, thus rendering the sites that are more accessible to the
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16 fluid more likely to be dissolved.
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20 **2.3.2 Details of the Numerical models**

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22 For all models, the simulated crystal can be interpreted as a three-dimensional collection
23
24 of voxels, each with a lateral length equal to the VSI lateral resolution ($dl = 0.22 \mu\text{m}$), their
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26 vertical size being assigned a value of $dv = 2 \text{ nm}$, consistent with the VSI vertical resolution
27
28 range (0.5 - 2 nm; note that we selected the largest value to reduce computational time).
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31 To ensure that the outputs of the numerical simulations were representative of the
32
33 experimental scenarios, several settings were considered for all models: (i) the size of the
34
35 simulation window was taken equal to the size of a single experimental sample window
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37 (150×150 in terms of pixels, corresponding to a crystal surface area of 33×33 μm^2); (ii) model
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39 evaluations were stopped when the simulated surface retreat, averaged over the whole surface,
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41 was equal to its experimental counterpart at the end of the experiment (at 5× magnification);
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43 and (iii) the initial topography employed in the simulations coincided with one of the
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45 experimental sample windows at $t = 0$ (as the sample windows were characterized by similar
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47 initial average surface roughness R_a , as described in Section 2.1.5). Finally, the outputs of the
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49 simulations were compared with the experimental data in terms of the statistics of Z' and of the
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51 associated semi-variogram function to identify the numerical model which was most
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53 compatible with the statistical results stemming from the experimental observations.
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3 The simulation of the dissolution process at a given time is based on the generation of a
4 random number sampled from a uniform distribution at each (x, y) location of the simulated
5 solid matrix. Such number was then compared to the probability value assigned to the
6 corresponding voxel: if the generated number was lower than the corresponding probability,
7 the voxel was released from the crystal volume; otherwise, it remained in place.
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14 *Model 1* involved the presence of 5 types of surface voxels assigned at the beginning of the
15 simulation, each one associated with a given probability of dissolution, being randomly
16 distributed over the entire simulated crystal volume (i.e., the frequency of each voxel type was
17 20% and they were distributed without any spatial dependency; see Figure 5a).
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24 *Model 2* was initiated upon assigning to the whole crystal volume the same probability of
25 dissolution (P_{min}). Then, a given number of voxel-defects (characterized by a given probability
26 $P_{max} > P_{min}$) were randomly and uniformly distributed over the whole crystal volume. Each of
27 these defects was associated with a degree of spatial influence, i.e., all neighboring voxels up
28 to a given distance (L , which is intended as an adjustable model parameter) were characterized
29 by a probability of dissolution ($P_{min} < P < P_{max}$, P decreasing linearly with distance from the
30 voxel-defect, until the non-disturbed region was reached; see Figure 5b for an example of a
31 voxel-defect).
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42 *Model 3* was characterized by 2 parameters (here denoted as P_1 and P_1^*), and is based on
43 the concept that the voxels having the largest number of faces exposed to the fluid are the most
44 reactive ones. All voxels having only their top faces exposed to the fluid are characterized by a
45 probability of dissolution P_1^* . All voxels having also one or more lateral faces exposed to the
46 fluid are characterized by a higher probability of being associated with dissolution: the higher
47 the number of faces exposed to the fluid, the higher the probability that dissolution takes place
48 at that voxel. Model parameter P_1 corresponds to the contribution to the probability of
49 dissolution of a single voxel given by a single lateral face exposed to the fluid. Hence, the
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probability of dissolution linked to a voxel having j lateral faces in contact with the fluid is $P_1 \times j$. This model also includes the possibility that various voxels which are superimposed (henceforth termed voxel spines, see Figure 5c) dissolve simultaneously. Consistently, the larger the voxel spines lateral area in contact with the fluid, the higher the probability that one or more voxels are dissolved: for a given voxel within a voxel spine, the probability that the voxel considered is dissolved along with all the upper voxels is $P_1 \times \sum_{j=1}^4 j \times n_j + P_1^*$, where n_j is the number of lower voxels (the one considered being included) having j faces exposed to the fluid and P_1^* is nonzero only when considering the top voxel (see Figure 5c for a schematic example).

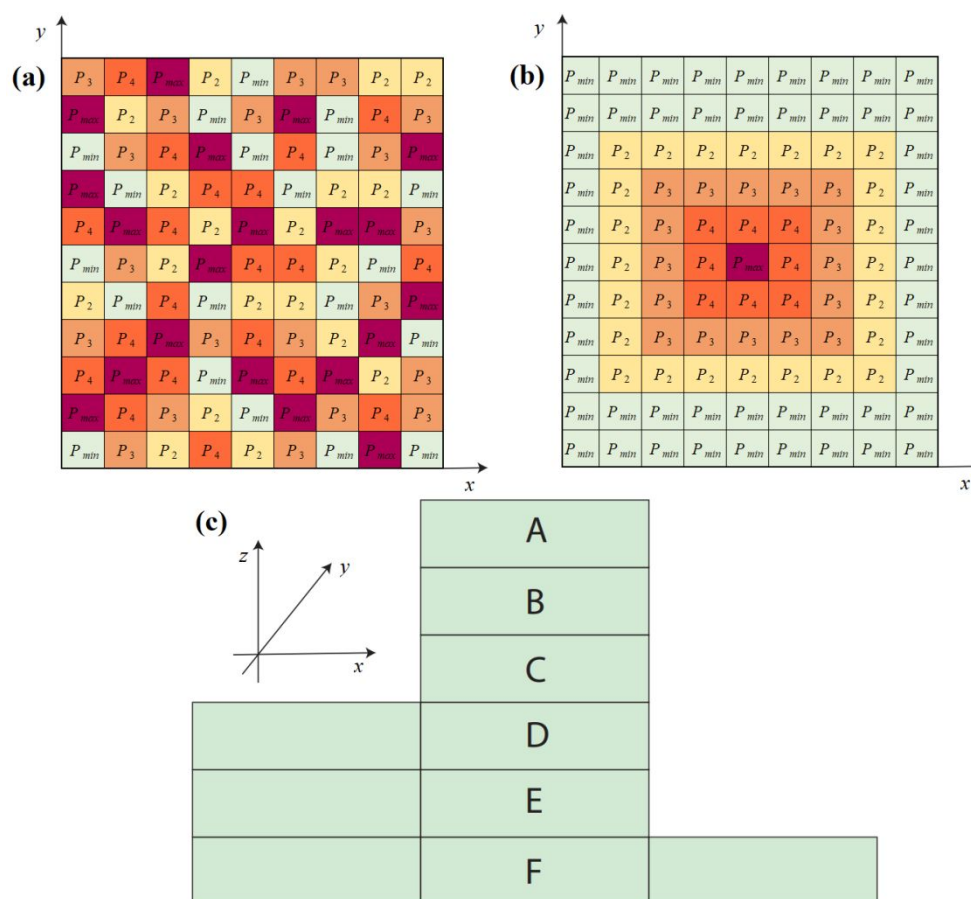


Figure 5 (a-b) Conceptual representation of the main features of *Model 1* (a) and *Model 2* (b) in a horizontal section. In these examples, values of dissolution probabilities are represented as $P_{\min} < P_2 < P_3 < P_4 < P_{\max}$. (c) Schematic representation of a voxel spine in a vertical section. As an example, the probability $P_{A+B+C+D+E+F}$ that voxels *A*, *B*, *C*, *D*, *E*, and *F* are all subject to dissolution (and hence are released from the solid matrix) at the same time is zero, as it is given by the probability P_F of voxel *F* (with no lateral faces exposed to the fluid) to be released, which is null (note that the configuration of the voxel layers along the y -axis is the same as the

one here depicted). Following a similar reasoning: $P_{A+B+C+D+E} = P_{A+B+C+D+E+F} + P_E$
 $= 0 + P_1 \times 1$; $P_{A+B+C+D} = P_{A+B+C+D+E} + P_D = P_1 \times 1 + P_1 \times 1$; $P_{A+B+C} =$
 $P_{A+B+C+D} + P_D = P_1 \times 2 + P_1 \times 2$; $P_{A+B} = P_{A+B+C} + P_C = P_1 \times 4 + P_1 \times 2$; $P_A = P_1$
 $\times 6 + P_1 \times 2 + P_1^*$.

3. Results

3.1 Calcite dissolution rate

The average surface retreat (Δh) measured after 188 hours of dissolution at $\Omega = 0.60$ was determined to be $(0.20 \pm 0.01) \mu\text{m}$ (mean value \pm standard deviation), which resulted in an estimate of the {104} overall dissolution rate equal to $(8.0 \pm 0.4) \times 10^{-9} \text{ mol m}^{-2} \text{ s}^{-1}$. We note that this value is in good agreement with close-to-equilibrium values documented in previous studies (⁴⁰ and ²⁵) using similar experimental and analytical protocols (these values ranging from $1.8 \times 10^{-8} \text{ mol m}^{-2} \text{ s}^{-1}$ at $\Omega = 0.45$ to $5.6 \times 10^{-9} \text{ mol m}^{-2} \text{ s}^{-1}$ at $\Omega = 0.80$). This imbues us with confidence that the desired close-to-equilibrium thermodynamic conditions were attained in our experimental setting.

3.2 Statistical analyses

3.2.1 Time-resolved statistical analysis of surface roughness

Key results of the time-resolved statistical analysis of the mean-removed surface topography (Z') are collected in Figure 6, which depicts the temporal dependence of variance (Fig. 6a), skewness (Fig. 6b), and kurtosis (Fig. 6c) of Z' , as obtained upon averaging the corresponding quantities evaluated for each of the 9 samples analyzed. Intervals of width equal to \pm one standard deviation are also depicted to provide an appraisal of the variability of such statistical moments among the samples. These plots reveal that temporal variations of the main statistics of the surface topography take place mostly at early times (reaction times shorter than 46 hours). All of the investigated statistical moments are seen to exhibit a plateau after such a relaxation time. Fig. 6a clearly shows a steep increase of the variance for time $t < 46$ h, until an

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3 asymptotic value of $(2.85 \pm 0.6) \times 10^{-5} \mu\text{m}^2$ is attained. The sharp variation observed at the
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5 very beginning of the experiment is possibly related to the initial condition rendered by the
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7 polishing process which might have locally weakened the surface. Polishing-resulting scratches
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9 are likely to induce the formation of surface defects which are in turn prone to preferential
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11 dissolution (see e.g., ^{29, 32}). Attainment of constant values of these statistical moments suggests
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13 that the rate at which the surface topography dissolves tends to become spatially constant as
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15 time progresses. Analysis of the third and fourth moments (i.e., skewness and kurtosis) of Z'
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17 reveal that the initial surface configuration is characterized by an approximately Gaussian
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19 probability density function, whereas the surface configuration at late times emerges as left-
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21 skewed ($skew < 0$) and leptokurtic ($kurt > 3$). Fig. 6b shows a transition from a Gaussian pattern
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23 skewed ($skew < 0$) and leptokurtic ($kurt > 3$). Fig. 6b shows a transition from a Gaussian pattern
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25 ($skew \approx 0$) to an asymptotic negative value ($skew \approx -0.2$), displaying a minimum value at early
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27 times ($t \approx 23\text{h}$). Similarly, Fig. 6c illustrates a deviation from values typical of a Gaussian
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29 distribution at $t = 0\text{h}$ to a leptokurtic behavior at later times. This implies that dissolution leads
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31 to the formation of more pronounced peaks and valleys as the reactions starts, the valleys being
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33 predominant. This behavior may also provide a quantitative criterion to distinguish, at steady-
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35 state, between reacted and unreacted (mechanically-polished) calcite surfaces based only on
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37 surface topography measurements. Consistent with the observation of Siena et al. ³⁹,
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39 (mechanically-polished) dissolved calcite surfaces are characterized, at late times, by left-
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41 skewed topography distributions. As such, while the probability distribution of the surface
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43 roughness across the mechanically-polished unreacted surface is symmetric, reaction is seen to
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45 be associated with a larger frequency of low rather than high heights values, implying that
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47 dissolution was very active in the majority of the area. As long as initial symmetric unreacted
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49 surfaces are concerned, the asymmetry of Z' distributions can then be considered as a distinctive
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51 feature of dissolution, being also consistent with the right-skewed rate-spectra observed in
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53 several studies related to a variety of minerals altered through various dissolution regimes (e.g.,
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3 ²⁰ or ³⁰). Furthermore, the decrease in skewness and the increase in kurtosis as the reaction
4 starts, may provide some insights regarding surface reactivity. Respectively: (i) a skewness
5 value that decreases from 0 at $t = 0\text{h}$ to a negative value at $t = 23\text{h}$ might be the result of a non-
6 uniform distribution of dissolution rates; (ii) a kurtosis value that increase from 3 at $t = 0\text{h}$ to a
7 larger value at $t = 23\text{h}$ might be related to dissolution rates that are, respectively, extremely
8 lower and extremely higher with respect to the average dissolution rate.
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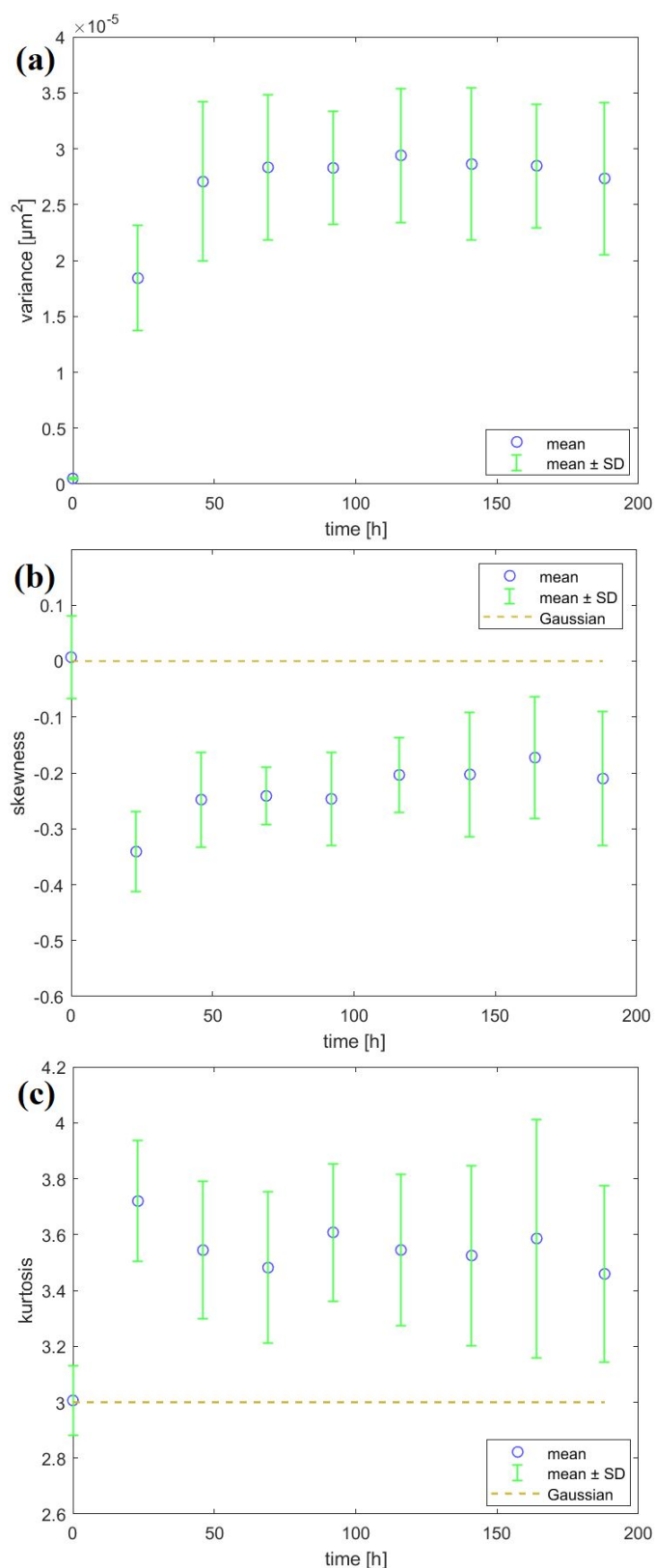


Figure 6 Temporal evolution of the key higher-order statistics of Z' , i.e., variance (a), skewness (b) and kurtosis (c). The figure depicts mean values (circles) evaluated across the 9 sample windows identified in Figure 2 and intervals of width corresponding to \pm one standard deviation (SD).

3.2.2 Time-resolved semi-variogram analyses

Figure 7 depicts sample omni-directional semi-variograms of surface roughness evaluated as a function of time across a selected observation window, results associated with the remaining windows being qualitatively similar.

Consistent with the results of Section 3.2.1, the most pronounced variations take place at early times. Three regimes can be identified. The first regime (corresponding to $0 \text{ h} \leq t \leq 23 \text{ h}$) is characterized by a strong increase of the experimental semi-variogram sill with time, consistent with the sample variance rise (Figure 6a and Figure 7a). The semi-variogram at $t = 0 \text{ h}$ exhibits higher values at short lags and a more persistent spatial correlation than its counterpart at $t = 23 \text{ h}$ (Figure 7b). This feature, combined with the (previously discussed) observation that the distributions of Z' are left-skewed, may be employed to quantitatively distinguish between reacted and unreacted surfaces. Semi-variograms evaluated for $23 \text{ h} < t \leq 69 \text{ h}$ are characterized by a mild temporal increase of their sill (see also Fig. 6a) and a virtually constant level of spatial correlation across time (Figure 7c). Our results show that differences between semi-variograms evaluated at subsequent times tend to diminish as time increases (see also Fig. 7d), consistent with the observation that the system tends to attain a certain degree of steady-state, in terms of the statistical characterization of surface roughness.

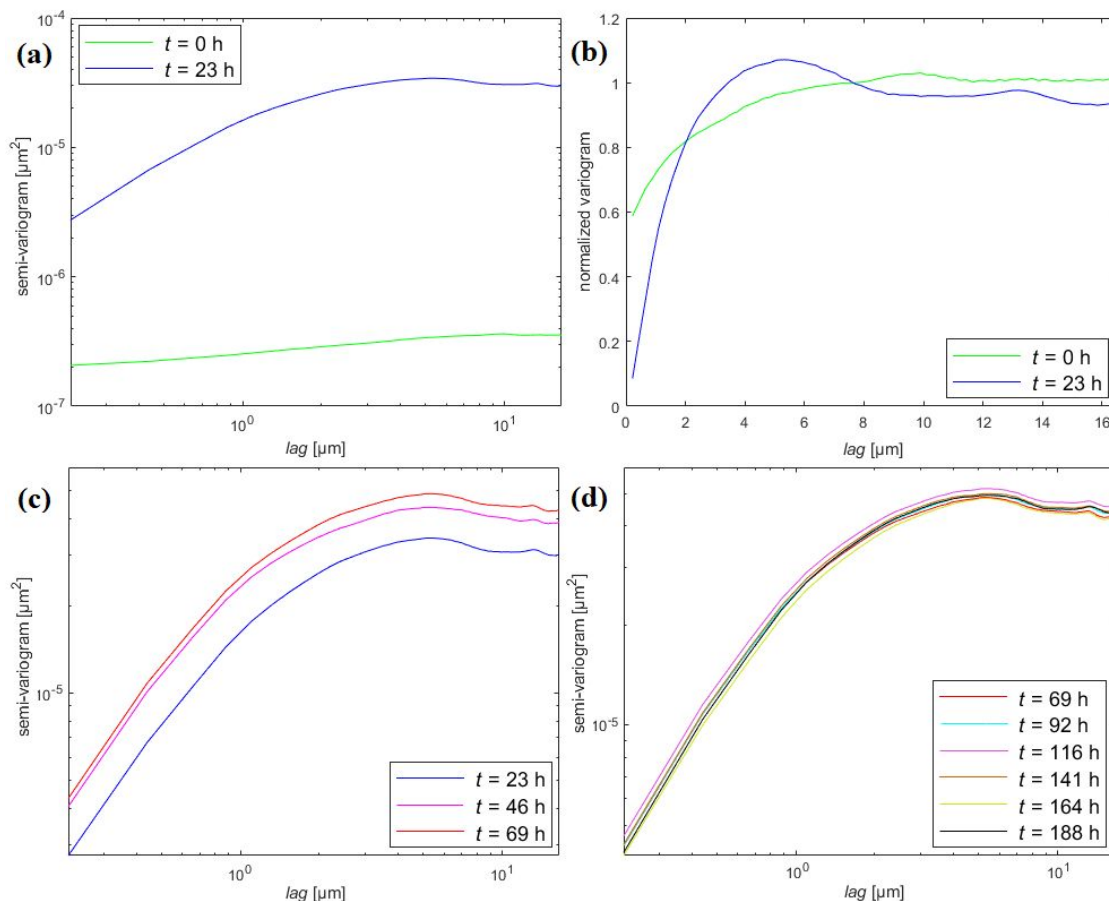


Figure 7 Evolution of experimental semi-variogram in time for a selected sample window (data associated with the remaining windows exhibit a qualitatively similar behavior). (a) Experimental semi-variogram (log-log scale) at observation time $t = 0$ h and 23 h; (b) Normal scored transformed variogram (i.e., semi-variogram normalized to the sample variance) at 0 and 23 hours. The main advantage of this visualization is that it makes it possible to compare the structure of semi-variograms characterized by strongly different sample variances; (c) Experimental semi-variograms for $23 \text{ h} \leq t \leq 69 \text{ h}$; (d) Experimental semi-variograms for $69 \text{ h} \leq t \leq 188 \text{ h}$.

3.3.1 Numerical model selection

As discussed in Section 2.3.1, model parameters are tuned until the best possible agreement with the statistics of the experimental data is attained. The least stringent condition to be met to this end was the constraint on the observed average surface retreat.

Optimal parameter values correspond to minimization of the mean square error (MSE) between (i) the average variance values depicted in Fig. 6a and their model-based counterparts and (ii) the late-time model-based semi-variogram and the corresponding averaged

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3 experimental semi-variogram. We recall that for a stationary random field, as $Z'(\mathbf{x})$ is, the semi-
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5 variogram sill corresponds to the variance.
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8 Model-based results correspond to averages across a collection of 100 simulation runs for
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10 each set of parameters tested. Figure 8a depicts the behavior of the variance of surface
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12 roughness associated with (i) the experimental observations (and obtained upon averaging
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14 results associated with all of the spatial sub-regions investigated) and (ii) their model-based
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16 counterparts corresponding to the optimal parameter value sets versus the extent of reaction
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18 (i.e., the fraction of total iterations, equivalent to time). Figure 8b shows a corresponding
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20 depiction of sample semi-variograms associated with the longest observation time ($t = 188$ h,
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22 corresponding to 100% extent of the reaction).
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26 Results from *Model 1* were deemed as not consistent with the documented experimental
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28 evidences because the model-based semi-variogram in Figure 8b displayed a lack of spatial
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30 correlation, in line with the nature of the underlying physical model. A large number of
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32 parameter value combinations have been investigated for *Model 2*. It was found that, even
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34 though the model does imprint a certain degree of correlation to the spatial field of roughness,
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36 none of the analyzed parameter sets allowed attaining a variance plateau over time. The latter
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38 result is related to the observation that the considered dissolution mechanism implied the
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40 variance to continuously increase with time. Otherwise, results from *Model 3* are consistent
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42 with the experimental observations with reference to both the attainment of the variance plateau
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44 at long times (see Figure 8a) as well as the ability to reproduce the semi-variogram structure
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46 after 188 hours of dissolution (see Figure 8b). As an additional advantage, we note that *Model*
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48 3 requires only two model parameters. All of these features contribute to identify *Model 3* as
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50 the most successful, in terms of consistency with the experimental results, across the collection
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52 of models tested.
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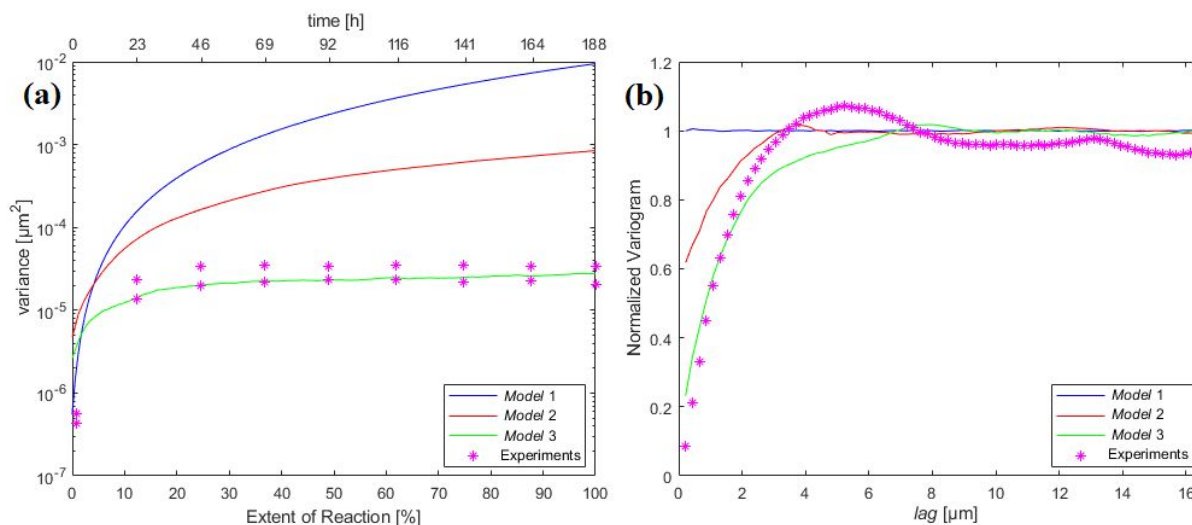


Figure 8 (a) Variance of surface roughness (Z') associated with (i) the experimental observations (symbols; \pm standard deviation evaluated across the 9 spatial sub-regions investigated) and (ii) their model-based counterparts corresponding to the optimal parameter value sets (solid lines) versus the extent of reaction (i.e., the fraction of total iterations, equivalent to time). (b) Experimental and numerically-based sample semi-variograms associated with the longest observation time ($t = 188$ h, corresponding to 100% extent of the reaction).

3.3.2 Parametrization and validation of the selected model

The set of parameters for *Model 3* that results in the best match (based on the minimization of MSE, as described in Section 3.3.1) between numerical and experimental data is: $P_1=0.05$ and $P_1^*=0.1$.

This model is able to capture the observed behavior of the variance, as it causes the variance to steeply increase at early times through the generation of areas dissolving preferentially, until a plateau is reached as these areas merge at long times, while the average retreat increases linearly. The implementation of the dissolution mechanism associated with *Model 3* imprints a certain degree of spatial correlation due to the assumption that the voxels associated with a larger number of faces exposed to the fluid are more reactive (i.e., more likely to be dissolved). As a consequence, the simulated dissolution process advances through time by propagating a number of randomly generated valleys (i.e., areas resulting from dissolution being locally more active than in the remaining portions of the surface). The temporal evolution of the simulated surface is probabilistically related to the locations where these valleys are randomly generated

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3 at early iterations. As a result, diverse simulation runs may lead to moderately different
4 statistical evolutions of the simulated surface. We investigate this aspect by performing 100
5 runs of *Model 3* with the optimal parameter set. Figure 9 depicts a comparison of Z' statistics
6 between \pm standard deviation resulting from 100 simulation runs of *Model 3* and \pm standard
7 deviation evaluated across the 9 sub-regions investigated.
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10 We observe that the steady-state model-based values of Z' variance are in good agreement
11 with their experimental counterparts (see Figure 9a), whereas mild discrepancies can be
12 observed as the order of the statistical moments considered increases (see Figures 9b and 9c).
13 The main differences in the behavior of the variance of Z' are detected at early times, the model
14 plateau being attained later than the experimental one. Whereas the early time system variation
15 observed experimentally is very sharp (and possibly enhanced by the weakening effect of
16 polishing scratches), numerical results show that merging of areas dissolving preferentially is
17 more gradual. Nevertheless, the main temporal trend of the modeling results is satisfactorily
18 consistent with the experimental observations, thus supporting the selection of *Model 3* as
19 capable of embedding the main statistical traits associated with the observed system dynamics.
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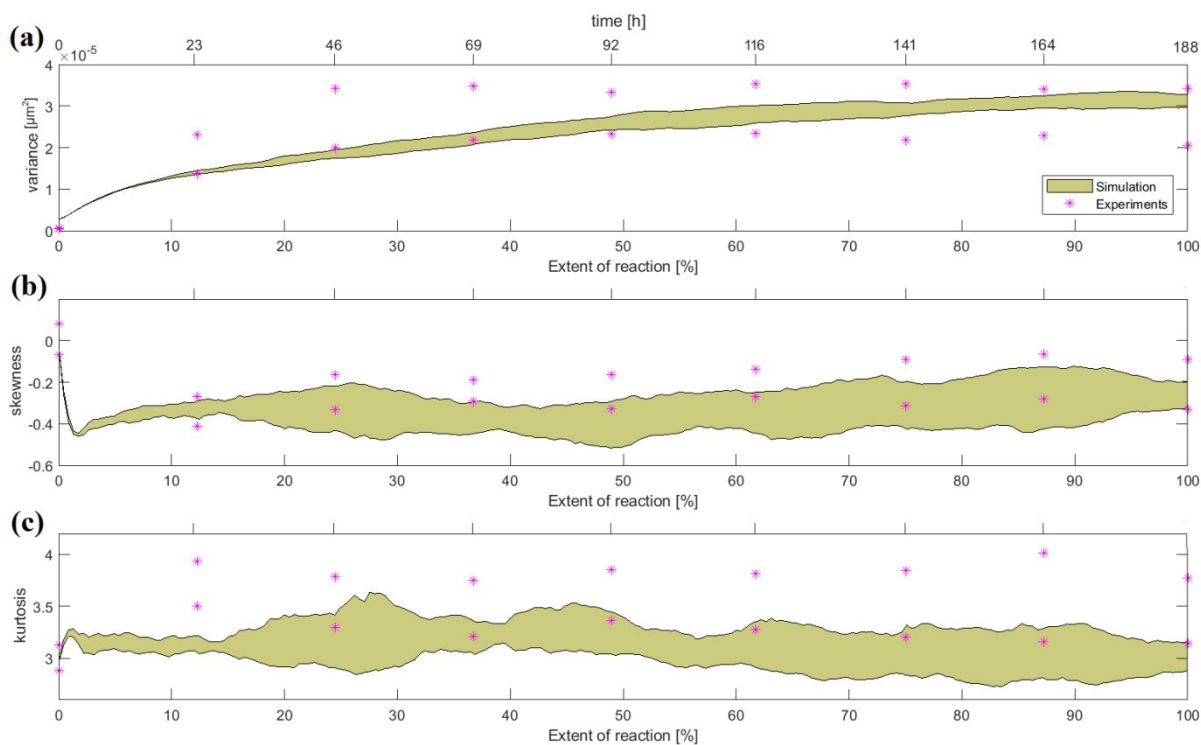


Figure 9 Temporal evolution of the key statistics of Z' evaluated over 100 realizations of *Model 3*. Envelopes of width equal to \pm one standard deviation around the corresponding mean are depicted (shaded areas). Symbols denote the corresponding intervals associated with the experimental data monitored across the 9 sample windows investigated.

4. Discussion

4.1. Interest of a statistical approach to surface topography of dissolved mineral surfaces

Attempting to take into account the spatial heterogeneity of mineral dissolution, Fischer et al. ²⁰ suggested the rate spectra approach, which was then largely applied in several subsequent works (e.g., ^{15, 20, 22, 27, 29-33}). This strategy has opened the possibility for new developments in the field of mineral reactivity, as the deconvolution of rate spectra allows for the determination of the main contributors to a given rate distribution and can shed some light on our ability to identify dissolution mechanisms ^{34, 35, 45} or the specific contribution of microbes to chemical weathering ³³. However, this approach suffers from two main limitations: (i) it requires the presence of a laterally fixed (motionless through time) surface and/or reference points at the crystal surface, which may not be compatible with a number of field applications; and (ii) it does not contain information about the spatial correlation of the surface roughness

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3 distribution, as the evolution of dissolved surfaces characterized by diverse extents of spatial
4 correlation might be associated with the same rate spectra.
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8 Our statistically-based approach, targeting close-to-equilibrium conditions (which are
9 characteristic of most natural settings) can be seen as complementary to the rate spectra
10 approach.
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14 Our observations that the initial surface configuration of a mechanically-polished calcite
15 sample is characterized by a Gaussian probability density function, whereas the surface
16 equilibrium configuration is left-skewed and leptokurtic, may provide a quantitative criterion
17 to distinguish between reacted and unreacted (mechanically-polished) calcite surfaces based
18 only on surface topography measurements. More specifically, the left-skewness of dissolved
19 calcite surfaces (already observed by Siena et al. ³⁹), which intrinsically implies that the surface
20 is characterized by a larger number of low heights values rather than high heights values, could
21 be viewed as a specific feature of dissolution, as it results in a right-skewed rate spectrum
22 consistent with the observations of numerous studies related to various minerals altered through
23 various dissolution regimes (e.g., ²⁰ or ³⁰).
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37 **4.2 Insights on surface dynamics of calcite dissolved at close-to-equilibrium conditions**

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39 Our results show that the surface roughness variance increases with time until a plateau
40 is reached, indicating the achievement of a steady-state surface configuration, as supported by
41 the attainment of a virtually invariant semi-variogram. We note that the numerical simulations
42 and scanning tunneling microscopy observations of silver electro dissolution by Hernández et
43 al. ⁴⁶ yielded similar trends. A qualitatively similar behavior was also observed in previous
44 works (²³ and ⁴⁷) at far-from-equilibrium conditions for fluorite, these authors noting that the
45 value of the variance plateau may be specific to the considered crystallographic orientation.
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56 With the aim of obtaining further insights on the underlying dissolution mechanism,
57 three diverse numerical models have been examined (see Section 2.3). *Model 3* (i.e., Dynamic
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3 dissolution probability based on fluid accessibility) has been identified as the one that best
4 captures the results of our statistical analyses on the evolution of the surface topography. *Model*
5 *1* (i.e., Uniform and spatially independent dissolution probability similar to pitting corrosion
6 models) failed to account for spatial correlation, whereas *Model 2* led to a monotonous temporal
7 increase of the variance of the surface topography and the formation of geometric etch pits,
8 these features being incompatible with dissolution at close-to-equilibrium conditions.
9 Otherwise, the results of *Model 3* (i.e., Equiprobable dissolution with presence of volume
10 defects dissolving preferentially), by involving a dynamic set of dissolution probabilities, were
11 fully consistent with the main statistics and semi-variogram structure of surface roughness.
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24 Pollet-Villard et al. ³⁸ developed a model accounting for the long-term evolution of the
25 (001) surface of a dissolving K-feldspar at far-from-equilibrium conditions, based on the
26 correspondence between modeling and experimentally based statistical behaviors. Similarly,
27 the results of our study enable us to conclude that the dissolution of a {104} calcite surface at
28 close-to-equilibrium conditions progresses in a way which is consistent with a mechanism
29 based on a random generation of a number of valleys, which tend to increase with time, and
30 grow until they finally merge to attain a steady-state surface area. We note that, while the
31 temporal evolution of the roughness variance reaches a plateau, the average retreat increases
32 according to a linear trend, which is also consistent with other works (e.g., ³⁸).
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45 One should emphasize that our results are not intended to imply that *Model 3* can be
46 considered as a universal descriptor of mineral dissolution that would apply to any mineral
47 reacted at close-to-equilibrium conditions. As opposed to mechanistic models such as kinetic
48 Monte Carlo simulations (e.g., ^{26, 28, 48}), *Model 3* is essentially phenomenological and its
49 parameters need to be estimated against experimental observations. Nonetheless, and more
50 importantly, the development and analysis of these various models underlined the need to
51 consider some key elements that have to be implemented into any future model to ensure its
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3 ability to capture major features of the statistics of surface roughness, as described in the
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5 following:
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8 (i) The dissolution probability of a given voxel is dynamic, as it depends on its immediate
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10 neighboring environment. Interestingly, this feature is also observed at a smaller (i.e.,
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12 atomic) scale (e.g., [26, 48-51](#)), although in this latter case, the relationship that links the
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14 dissolution probability of an atom to its environment corresponds to the product of the
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16 probabilities to break the bonds linking the atom to the surface (thus corresponding to the
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18 hydrolysis of all bonds simultaneously). Applying the exact same rule also at the nm-scale
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20 would most likely be unreasonable without the development of upscaling laws, as the
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22 detachment of a nm-scale voxel already involves the dissolution of hundreds of thousands
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24 of atoms. Here, we show that the detachment of a voxel is related, in a statistical sense, to
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26 the number of faces of the voxel exposed to the solution, consistent with a macroscopic
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28 description of the dissolution process where the dissolution rate is scaled to the surface area
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30 in contact with the solution.
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35 (ii) Specific features that are shown to be consistent with the observation that the variance of
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37 the surface roughness reaches a plateau include (a) considering that columns of voxels (i.e.,
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39 voxel spines) can dissolve at the same time, as well as (b) assigning a higher dissolution
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41 probability to voxels located at the top of a given column (see Section 2.3.2). Interestingly,
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43 such a condition is similar to that implemented in one of the models developed to describe
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45 the dissolution of silver [46](#), where a given probability is assigned to each voxel depending
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47 on its distance to the basal plane.
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51 (iii) Finally, we showed that the strength of the spatial correlation associated with the semi-
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53 variogram for late times is not necessarily related to a specific distance between pre-
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55 existing defects in the simulated crystal (as in *Model 2* or in Pollet-Villard et al. [38](#)), and
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3 can be otherwise associated with the specific choice of a dynamic set of dissolution
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5 probabilities.
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7 8 **4.3 Statistical characterizations of surface topography as a metric for identification of** 9 10 **reacting conditions?**

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12 Dissolution at far-from-equilibrium conditions results in a final topography which is
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14 markedly different from the steady-state surface observed in our study at close-to-equilibrium
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16 conditions [38](#), [52](#), [53](#). As suggested by Godinho et al. [23](#), the temporal evolution of surface
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18 topography may be associated with decreasing dissolution rates. This opens the possibility that
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20 the observed steady-state surface topography is closely related to the fluid composition. The
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22 experimental investigation of such conceptual element remains the focus of ongoing research.
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26 The calibrated numerical model developed in this work which has been shown to
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28 satisfactorily reproduce the experimentally observed statistical behavior at given conditions of
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30 initial roughness and fluid composition (i.e., *Model 3*), is then employed to explore additional
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32 scenarios with the aim of discriminating the contribution to the surface equilibrium
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34 configuration given by the initial roughness from that resulting from the fluid composition.
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38 Results obtained considering various initial topographies in the absence of macro-
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40 features suggest that the initial surface roughness has only a negligible impact on the simulated
41
42 steady-state surface. Our simulation results obtained upon considering the effect of a random
43
44 noise associated with the initial surface document that there are no significant variations in the
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46 fluid-mineral contact surface (and, hence, in preferential dissolution), whereas considering the
47
48 presence of surface macro-features (e.g., cracks or fractures) would possibly lead to different
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50 results. As an example, Figure 10 a, b depict the temporal evolutions of variance and the steady-
51
52 state semi-variograms resulting from 4 different initial surface topographies: (i) a perfectly flat
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54 surface, (ii) the experimental topography measured at $t = 0$ hours for one of the sample
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56 windows, (iii + iv) the topography associated with (ii) which is then corrupted through a white
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3 noise with standard deviation equal to: (iii) 1 standard deviation and (iv) 2 times the standard
4 deviation of the experimental topography considered. If supported by experimental data, these
5 results would enable one to rule out the starting topography as a major parameter controlling
6 the steady-state variance of the surface topography resulting from mineral dissolution.
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12 Conversely, differences in “fluid compositions” were shown to have a major impact on
13 the steady-state topography. The impact of fluid composition was simulated based on the
14 observation that P_1 and P_1^* are proportional to the dissolution rates, hence inversely
15 proportional to the saturation state. Here, we limit our analysis to conditions where no etch pits
16 would actually be observed, so that only dissolution rates corresponding to saturation states
17 ranging between $\Omega=0.45$ and $\Omega=0.8$ are simulated²⁵. As an example, Figure 10c, d illustrate
18 the variance temporal evolutions and the steady-state semi-variograms related to 3 different
19 simulated saturation states (and hence different surface retreats Δh after $t = 188$ hours): starting
20 from the values of the model parameters which were found to best interpret the experimental
21 conditions ($P_1=0.05$ and $P_1^*=0.1$), the saturation state of the solution was simulated by
22 increasing ($P_1=0.08$ and $P_1^*=0.13$) and decreasing ($P_1=0.02$ and $P_1^*=0.07$) the values of the
23 model parameters, which resulted respectively in higher ($\Delta h = 0.26 \mu\text{m}$) and lower ($\Delta h = 0.14$
24 μm) surface retreats. Model results suggest that higher saturation states (i.e., closer to
25 equilibrium conditions) result in higher values of the variance plateau due to the delayed
26 merging of the areas subject to preferential dissolution. This observation may seem counter-
27 intuitive: Fischer et al.²² showed that kink site density at the atomic scale is positively correlated
28 with an increase in surface roughness, the latter being in turn strongly correlated to the
29 dissolution rate (see their Fig. 6). On the other hand, nm-scale experimental and modeling
30 results of previous relevant studies^(23 and 47) showed that an increase in roughness can be
31 associated with decreasing dissolution rates, which would be in close agreement with our
32 overall analysis. We note that the results from our simulations, albeit constrained to the
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3 available experimental settings, suggest that the steady-state variance of the topography and the
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5 fluid composition may be related. From a practical aspect, generalizing this result would rest
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7 on the assumption that steady-state surface configuration would be achieved also at saturation
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9 states differing from the one here investigated: if supported through additional experimental
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11 observations, direct measurements of surface roughness would allow to unambiguously
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13 distinguish surfaces that have reacted from surfaces that have not, while possibly getting
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15 insights into the saturation state of the fluid at which dissolution took place. This justifies the
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17 need for further experimental investigation on the relationship between mineral dissolution rate
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19 and surface roughness, to enhance the ability of our model to characterize these processes.
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22 While further experimental studies are required to explore the behavior seen in our numerical
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24 simulations, our results suggest that the reacting conditions could be back-estimated through a
25
26 simple statistical characterization of topography data of a reacted surface, paving the way to the
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28 development of empirical transfer functions linking saturation states and reacted surface
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30 topographies.
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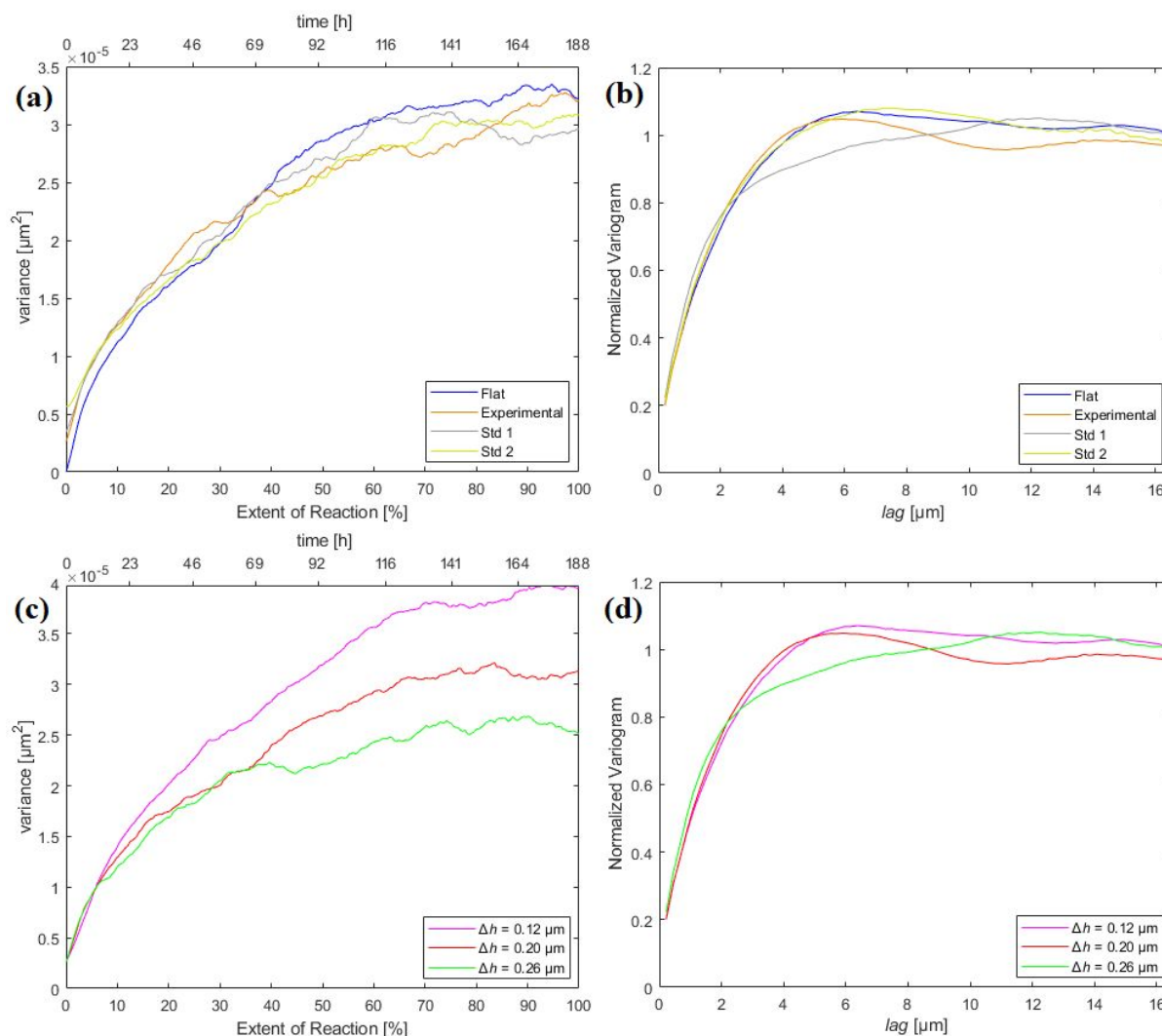


Figure 10 Temporal evolution of the variance of surface roughness as a function of the initial surface roughness (a) and of the fluid composition (i.e., saturation state) (c) resulting from numerical simulations associated with *Model 3*. Corresponding results for the normal scored transformed experimental semi-variogram evaluated at the last observation time (b, d) are also included. No significant variation in the semi-variogram was observed in both cases, suggesting that the semi-variogram structure may significantly change only at far from equilibrium conditions regardless of the initial surface roughness.

5. Conclusions

Time-resolved statistical analyses of a mechanically polished calcite surface reacting at close-to-equilibrium conditions were demonstrated to be a powerful tool to assess attainment of steady-state configuration for the reacting surface. This is reflected through the temporal dynamics of the key statistics and spatial semi-variograms of surface roughness which tend to stabilize after a given relaxation time (approximately two days in our experiments). Our

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3 approach can be used as a quantitative reference to compare (i) differently polished (or
4 unpolished, e.g., cleavage) calcite surfaces reacting at the same physico-chemical conditions,
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6 or (ii) mechanically polished calcite surfaces reacting at different physico-chemical conditions.
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8 In both cases, spatial analysis of semi-variogram functions can provide a quantitative appraisal
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10 of the strength of the space-time correlation of surface roughness as it evolves subject to the
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12 progress of the reaction. In this context, additional investigations may contribute to possibly
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14 distinguish between the individual contribution given by the initial surface roughness and the
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16 reacting conditions to the plateau values of the statistical moments characterizing the dissolving
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18 surface area at steady-state. As a consequence, a thorough characterization and understanding
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20 of the statistical evolution of surface roughness could possibly be key to back-estimate the
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22 reacting conditions (e.g., saturation state of the solution) by simply measuring the topography
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24 of a reacted surface after attainment of steady-state.
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30 We developed and implemented an original conceptual model which enabled us to
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32 reproduce the experimentally observed space-time evolution of the key statistical descriptors of
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34 surface topography. The documented consistency between modeling- and experimentally-based
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36 statistical analyses may imply that, under the investigated conditions, the dissolution
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38 mechanism is characterized by an overall space-time organization, as reflected by the semi-
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40 variogram structure, with the areas most exposed to the fluid being prone to preferential
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42 dissolution. The results of our statistical analyses document that such a dissolution mode is fully
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44 compatible with the achievement of a steady-state surface area observed through the
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46 experiments in this work.
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51 Ongoing work is geared towards analyzing whether this model, while being conceptually
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53 and mathematically simple, can be a viable tool to reliably predict the statistical evolution of
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55 calcite surface roughness as a function of the fluid composition.
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Acknowledgements

This work was funded by the project ANR-20-ERC8-0006-01_t erc Mobidic awarded to D.D.

We thank Gilles Morvan and René Boutin for their help with EBSD and ICP-AES measurements. Finally, L.S. thanks Arnaud Bouissonnié for the assistance in the implementation of experimental protocols.

Supporting Information

The Supporting Information is available at: ...

- (1) Grit sizes of the multi-step grinding and polishing sequence.
- (2) ICP-AES analyses of the aqueous solution at different times.

Data Availability

Data are available from the corresponding author upon request.

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TOC Graphic

