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# COMBINING GEODETIC AND GEOMORPHIC METHODS TO MONITOR RESTORED SIDE CHANNELS: FEEDBACK FROM THE UPPER RHINE

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## Abstract

15 Producing accurate and rapid geomorphic surveys is a key issue for the growing scientific and operational area of physical river restoration. A geodetic survey using 3-D modeling (lasergrammetry, photogrammetry, tacheometry) was combined with a geomorphic survey (RFID bedload tracking, survey of grain size, geomorphic units evolution and geometrical changes) in order to monitor morphodynamic adjustments in a restored anastomosing channel of the Upper Rhine. On this basis, functioning indicators were developed to survey morphological changes at different spatio-temporal scales. Because

20 Structure from Motion (SfM)-photogrammetry is a fast and low-cost method able to produce high-resolution point clouds, the method is particularly well-suited for monitoring complex fluvial environments that have been subject to rapid and intense changes. A hybrid method was developed to complement bank point-clouds with bathymetric data obtained by Total Station leveling. By using the Multiscale Model to Model Cloud Comparison method (M3C2), volumes of erosion-deposition were determined for all surveys. Combining geodetic and geomorphic approaches resulted in a detailed

25 assessment of channel adjustments. This was achieved by creating a set of indicators (e.g., related to vertical and longitudinal evolutions, sediment budget, etc.) that allowed us to both characterize geomorphic adjustments and identify morphodynamic limiting factors. These indicators may be used in the future in a wide range of restoration surveys.

## Key words

Morphological 3-D modelling; Morphodynamic functioning indicators; Sediment budget; Mid-sized river hydro-morphological restoration.

## 1 Introduction

5 The geomorphic monitoring of restored hydrosystems is a key issue in surveying channel adjustments (Hickin, 1983; Gurnell *et al.*, 2006), habitat dynamics (Raven *et al.*, 2002; Palmer *et al.*, 2010; Belletti *et al.*, 2014) and thereby assessing the efficiency of restoration actions (Fausch *et al.*, 2002; Bernhardt *et al.*, 2005; Stammel *et al.*, 2011; Buchanan *et al.*, 2012; Rinaldi *et al.*, 2017). Conventional morphological surveying techniques, which are usually based on cross sections, are time consuming and require a trade-off between temporal and spatial resolution, accuracy, and survey frequency (Lane *et al.*,  
10 1994; Barker *et al.*, 1997; Brasington *et al.*, 2000; Bangen *et al.*, 2014a). The first applications of remote sensing in fluvial environments date back to the 1960s (Painter *et al.*, 1974; Collins and Moon, 1979). However, no significant technological innovations in topographic techniques that would have improved and simplified field surveys appeared before the beginning of the twenty-first century (James & Robson, 2012; Westoby *et al.*, 2012; Nadal-Romero *et al.*, 2015; Smith *et al.*, 2015; Dietrich, 2016; Eltner *et al.*, 2016).

15 Numerous studies have analyzed the accuracy and resolution (Smith *et al.*, 2015; Eltner *et al.*, 2016; Grussenmeyer *et al.*, 2016; Murtiyoso and Grussenmeyer, 2017) of Terrestrial Laser Scanning (TLS; Telling *et al.*, 2017) and Structure from Motion (SfM)-derived topographic datasets for a wide range of geomorphic environments (Table 1). Compared to the first TLS instrument that was developed two decades ago, recent technologies (e.g., FARO Focus 3D, Z+F IMAGER 5016) use smaller, more lightweight devices, and provide high quality colorimetric clouds (Remondino and El-Hakim, 2006). The TLS  
20 method is still expensive because of the high cost of the recording system, and that it requires special care in remote areas (targets, spheres, etc.). These constraints explain why only a few surveys aiming at monitoring geomorphic changes in dynamic channels have been based on TLS (Table 1). Besides, SfM photogrammetry is a low-cost method that generates high spatial resolution and 3-D data. It also allows for accurate surveys of frequent river changes, potentially for the short- (<5 yr) and/or mid-term (>5 yr; Fonstad *et al.*, 2013). Additionally, we consider in Table 4 that long-term periods cover >10  
25 yr). However, despite the well-known high accuracy of 3-D topographic monitoring obtained by TLS and SfM-

photogrammetry, these methods have been rarely used in river restoration, especially in the context of monitoring morphodynamic changes in side channel restoration (Table 1). Recent geomorphic studies also use newly developed remote sensing applications such as Unmanned Aerial Systems (UAS; Lejot *et al.*, 2007; Casado *et al.*, 2015; Prosdocimi *et al.*, 2015; Woodget *et al.*, 2015; Belletti *et al.*, 2017), tethered kites, blimps (Boike and Yoshikawa, 2003; Smith *et al.*, 2009; Vericat *et al.*, 2009) and near-census approaches to identify and map a wide range of morphological changes at any spatial scale (Wyrick *et al.*, 2014, 2016).

Recent remote sensing applications provide fine Digital Elevation Models (DEM) combined with bathymetric modelling for a large range of channel types (Gao, 2009; Flener *et al.*, 2013; Williams *et al.*, 2014; Eltner *et al.*, 2016). They are, however, limited by water depth, riparian canopy effects, and logjams, which are frequent along dynamic side channels located in dense alluvial forests. Within this context, the combination of survey technologies (e.g., mixing GPS and LiDAR data) is useful to generate complete and accurate datasets (Wheaton *et al.*, 2009; Bangen *et al.*, 2014b). This provides a way to visualize and precisely quantify channel adjustments (e.g., the sediment budgets) by applying DEM of Difference (DoD) (Martin and Church, 1995; Brasington *et al.*, 2000a; Lane *et al.*, 2000; Wheaton *et al.*, 2009; Milan, 2012; Wyrick *et al.*, 2016).

With the increased number of river restoration projects during the last three decades (Morandi *et al.*, 2014), many studies tried to develop indicators that would allow assessing the efficiency and sustainability of such management actions (Berger, 1997; Palmer *et al.*, 2005; Brierley and Fryirs, 2008; Dufour and Piégay, 2009; Pander and Geist, 2013; Rinaldi *et al.*, 2017), but only a few studies have combined geodetic surveys and conventional geomorphic methods (McLean and Church, 1999; Chandler *et al.*, 2002; Chapuis *et al.*, 2015; MacVicar *et al.*, 2015). An indicator in fluvial geomorphology is a quantified specific metric allowing the characterization of the evolution of fluvial forms (generally in longitudinal, lateral and/or vertical dimensions) and processes (erosion, sediment deposition, bedload dynamics; Corbonnois *et al.*, 2011). The global objective of the paper is to assess the impacts of a restoration project (morphological changes, bedload dynamics, habitats diversification, etc.) by combining geodetic (TLS, SfM) and geomorphic monitoring methods. More specifically, a threefold aim is targeted in the framework of the functional restoration performed in the Rohrschollen island (Upper Rhine River, France): (i) to adapt geodetic survey to monitor intense and frequent morphological changes along a dynamic side channel,

- (ii) to combine this geodetic survey with classical geomorphic methods to develop a set of morphodynamic indicators and
- (iii) to interpret results in terms of efficiency and sustainability of the restoration project.

Authors Date	Context	Approach	Geomorphological dimensions & components										DEM methods		Hydromorphological methods						
			Channel types	River restoration	Lateral channel restoration	Geomorphological survey	Methodological studies <sup>a</sup>	Lateral adjustments	Vertical adjustments	Bedload dynamics	Hydrodynamics	Geomorphic units evolution	Lasergrammetry	Photogrammetry	Levelling (cross-section)	RFID tracking	Hydraulic modelling	Sediment budget	Other methods		
Collins and Moon, 1979	WR																				
Dickinson and Scott, 1979	MR																				
Lane et al., 1994	BR																				
Martin and Church, 1995	BR																				
Barker et al., 1997	MR																				
Pyle et al., 1997	PGR																				
Heritage et al., 1998	MR																				
Westaway et al., 2000	BR																				
Brasington et al., 2000	BR																				
Chandler and Shiono, 2001	TF																				
Lane et al., 2001	TF																				
Chandler et al., 2002	PGR																				
Fuller et al., 2003	WR																				Bar
Formann et al., 2007	BR																				Bar
Lejot et al., 2007	BR																				Bar
Wheaton et al., 2009	BR																				
Lague et al., 2013	MR																				
Bangen et al., 2014a	SAC																				
Javernick et al., 2014	BR																				
Casado et al., 2015	WR																				
Chapuis et al., 2015	WR																				
Kuo et al., 2015	BR																				
MacVicar et al., 2015	SAC																				
Brunier et al., 2016	BR																				Bar
Dietrich, 2016	SAC																				
Jozkow et al., 2016	MS																				
Belletti et al., 2017	WR																				
Marteau et al., 2017	MS																				
Cook, 2017	WR																				
Wyrick et al., 2016	WR																				
Jugie et al., 2018	SAC																				

**BR** Braiding River  
**SAC** Small Alluvial Channel  
**TF** Tilting Flume  
**MR** Meandering River  
**MS** Mountain Stream  
**PGR** ProGlacial River  
**WR** Wandering River

<sup>a</sup> Results focused on methodological innovation (accuracy, methodological comparison...)  
<sup>b</sup> Bathymetric models  
<sup>c</sup> Unmanned Aerial Vehicles (UAV) and Artificial Neural Networks (ANN)  
<sup>d</sup> Geomorphic Units Survey (GUS) and Morphological Quality Index (MQI)  
<sup>e</sup> Near-census approach and decision-tree classification  
<sup>f</sup> Photo-Electronic Erosion Pin (PEEP)

**Table 1: Studies developed to monitor geomorphic channel changes by using DEM approaches.**

## 2 Study site

The Rohrschollen artificial Island is located about 8 km South East of the City of Strasbourg (Fig. 1a). It results from the construction of a power plant on the Rhine canal in 1970 (Fig. 1b). It has been highly impacted by engineering works since the beginning of the nineteenth century (Eschbach *et al.*, 2018). A diversion dam located at the southern side of the island (Fig. 1b) divert up to 1550 m<sup>3</sup>.s<sup>-1</sup> towards the power plant by means of an artificial canal (named “Rhine canal”). When discharge is below this threshold, only 13 m<sup>3</sup>.s<sup>-1</sup> feed the Old Rhine (East part of the island; Fig. 1b). When the Rhine discharge exceeds 1550 m<sup>3</sup>.s<sup>-1</sup>, overflows are streamed into the Old Rhine. The agricultural dam located in the northern side of the island allows progressive floods when discharges exceed 2800 m<sup>3</sup>.s<sup>-1</sup> (CIPR-ICPR, 2012). Consequently, during floods, the island is submerged by static flow. Significant flood retention rarely occurs, about once per decade in average. In addition, flowing water during floods does not occur and active morphodynamic processes in the floodplain and the anastomosing channel of the island. In this study, this channel is called “old channel” (Fig. 1b).

In the framework of the European LIFE+ project, launched in 2012, a floodgate was constructed, as well as a stabilized rockfill canal and a new upstream non-stabilized channel (called “artificial side channel”), which were excavated in the southern part of the island to reconnect the old channel to the Rhine’s hydrological regime. Depending on the discharge of the Rhine, water injected by the floodgate ranges from 2 to 80 m<sup>3</sup>.s<sup>-1</sup> during an average of about 51 days annually when the Rhine discharge is higher than 1550 m<sup>3</sup>.s<sup>-1</sup>. This new mode of functioning should allow the reestablishment intense and frequent morphodynamic processes (bank erosion, bedload transport, etc.), especially in the artificial side channel, which was intentionally undersized to induce self-forming adjustments. The artificial side channel shows a mean longitudinal water slope of 0.94‰, a bankfull discharge of about 20.5 m<sup>3</sup>.s<sup>-1</sup>, a mean bankfull width of about 12 m (except on sector C; see Section 3.1) and a mean initial bankfull specific stream power of 15-22 W.m<sup>-2</sup> (February 2014).

Fig. 1 (Study\_site)

Figure 1: (a) Location of the Rohrschollen Island in the Upper Rhine basin, (b) DEM of the whole study site and location of the main engineering works, (c) three survey sectors located on curvatures of the artificial side channel (white linear), (d) pictures of the diachronic evolution close to the three curvature sectors, (e) hydrogram of the reporting period.

## 5 3 Material and methods

### 3.1 Experimental spatial and flow regime designs

Our investigations focused on two spatial scales: (1) the whole artificial side channel and (2) three curvature sectors of this channel corresponding to reaches where lateral erosion was expected to be the highest (noted A, B and C in Fig. 1c). The morphological evolution of the old channel will not be discussed in this paper. Short and low intensity ecological floods were conducted during summer 2014 (less than  $20 \text{ m}^3 \cdot \text{s}^{-1}$  during about 1-2 h; see the top of the Fig. 2: trial flow pulse). Several ecological floods occurred during May and June 2015. Two main experimental floods were designed (Fig. 1e). Flood 1 was based on an incremental progress with three sporadic peaks in order to test the floodgate ( $40\text{-}70 \text{ m}^3 \cdot \text{s}^{-1}$ ) and flood 2 was characterized by a high discharge of  $60 \text{ m}^3 \cdot \text{s}^{-1}$  during 48 h (in accordance with the discharge of the Rhine). Diachronic pictures in Fig. 1d illustrate the effects of floods (enlargement, lateral erosion, bar formation, etc.) in the vicinity of the three studied curvature sectors during the total survey period. The total period covers two morphological changes (called surveying periods in Fig. 2) highlighted by three states that resulted from flood events that occurred from February 2014 to March 2016 (Fig. 2).

### 3.2 Geodetic survey

#### 3.2.1 TLS and SfM comparison

20 Focusing on the three main curvature sectors of the artificial side channel (Fig. 1c), the initial post-restoration state (state 0 = before floods and morphological evolution) was recorded before the first floods by a combination of TLS and SfM methods. TLS was set up with a FARO Focus 3-D with complete vertical and horizontal  $360^\circ$  view. This system can deliver an accuracy of 2 mm at 25 m. The system was setup to provide a spatial sampling of 3 mm at 10 m (Faro resolution setting:  $\frac{1}{2}$ ). Photogrammetric images were recorded using a calibrated Canon EOS 5D Mark II camera with a focal length of 28 mm. The camera was fixed on a tripod and translated along the riverbanks to satisfy a forward overlap of 70-80% between the successive parallel images. Additional oblique images were recorded to improve the geometry of the photogrammetry

datasets. Circular coded control targets permanently fixed on trees and regularly distributed on both riverbanks were used to process the images and georeference each curvature sector. To avoid errors related to the movement of trees (e.g., tree growth), the XYZ data of each target were collected at least once a year, if possible, during the SfM acquisition. Dense point clouds (Medium quality setting, i.e., ~1 pt/cm after post-processing) were generated using the *Photoscan* SfM software (version 1.2.6) and the optimization tool was used to georeference each sector. The internal projection errors for the GCPs were less than 5 cm and below 0.5 pixel (Table 2). The accuracy of the SfM method was assessed by a cloud to cloud registration algorithm considering TLS as a reference. Cloud to cloud accuracy ranged from 2 to 5 cm. The SfM method was therefore considered satisfactory and well-suited to monitor important morphological adjustments and rapid changes in the three curvature sectors.

10

	Sector A		Sector B		Sector C	
	Left bank	Right bank	Left bank	Right bank	Left bank	Right bank
<b>Number of image</b>	128	90	129	101	74	83
<b>Number of GCP</b>	5	6	7	8	10	7
<b>Reprojection error (Standard deviation m)</b>	0.001	0.02	0.02	0.05	0.03	0.01
<b>Reprojection error (Standard deviation pixels)</b>	0.18	0.23	0.49	0.2	0.19	0.12

**Table 2: *Photoscan* parameters for each sector. Example for state 2 (March 2016).**

### 3.2.2 Topographical data

A Total Station recording was used to detect morphological changes beneath the water surface, as proposed by Lane *et al.* (1994), and beyond the water between the curvature sectors. Regularly spaced (20 m) cross sections composed of 5 to 8 points were recorded in XYZ along the whole artificial side channel. Supplementary thalweg points were collected between each cross section for the longitudinal profile. In fact, such a method is time-consuming and can only generate a relatively small set of bathymetric points (Chandler *et al.*, 2002), but provides a reliable terrain-modeling method to quantify changes at the scale of the whole channel (Brasington *et al.*, 2000). Differential GNSS positioning was used to set up reference points upstream and downstream of the channel (Leica TS02 Total Station and Leica GPS 1200 System used in GNSS RTK Mode). They were connected with the polygonal network in order to provide a comparable diachronic database. Planimetric and

20

altimetric accuracy were about 3 cm. Thalweg trajectories were extracted using the free extension to the desktop *ArcGis* software from ESRI “River bathymetry Toolkit” with a 0.5 m vector resolution (ESSA, 2016).

### 3.2.3 Point cloud comparison

For each sector, accurate TIN models based on the dense point clouds were merged with DEM bathymetric data to generate  
5 “combined models”. The *CloudCompare M3C2* (version 2.8.1) method was used to record differences between point clouds from two consecutive surveys (Lague *et al.*, 2013; Fig. 2). The normal-scale parameter has been set to 1-2 m. We consider this parameter satisfactory to average the topographical discontinuities of the riverbanks. To provide a reliable comparison of high morphological volume changes (especially with regards to lateral erosion), a gridding interpolation was defined before merging two point clouds. Finally, a script implementation based on *MatLab* functions developed by TRIO group was used  
10 to calculate volume differences automatically between two clouds at two spatial scales: the whole channel and the three sectors.

## 3.3 Geomorphic survey

### 3.3.1 Grain size survey

To characterize the grain size at the channel bottom, the Wolman (1954) method was used on bars, riffles and an inherited  
15 gravel bar exhibiting particularly coarse grain sizes. Homogeneous grain size patches were delimited and b-axis measurements of 100 particles on each patch were taken, as recommended by Kondolf *et al.* (2016). This was used to generate grain size cumulative curves and extract statistical parameters such as the median diameter and standard deviation.

### 3.3.2 RFID tracking

The method involves attaching passive Radio Frequency IDentification (RFID) tags into gravel particles in order to track  
20 bedload movement (Lamarre *et al.*, 2005). We employed low-frequency tags consisting of glass cylinders 3.8 mm in diameter and 23 mm in length operating at 134.2 kHz and distributed by Texas Instruments (TI-RFID, 2006). A total of 413 particles with a b-axis of 22.6 to 64 mm were instrumented. RFID tags were attached to each particle by creating a lateral notch for the smaller particles and by drilling into the larger particles. Implementation was realized before state 1. Tracers were distributed homogeneously between the three transects at three injection-points located on both sides and in the middle

of the channel bottom. The experiment took place during two periods: the first one occurred between the two ecological floods of surveying period 2 and the second one was performed during state 2 (Fig. 2).

During field surveys, bars and shallow-water areas were monitored on foot using a 0.46 m diameter loop antenna and tracer positions were recorded using a GPS Trimble 6000 GNSS RTK. An accuracy of 0.05 to 0.10 m for such GPS points was reached after post-processing. In addition, horizontal errors include the detection limit of the antenna of around 0.34 m. Maximum cumulative horizontal errors were estimated to be  $\pm 1$  m depending on the relative position of the antenna during the survey, RFID orientation and GPS survey (Schneider *et al.*, 2010; Chapuis *et al.*, 2014; Arnaud *et al.*, 2015). According to operating controls carried out by Arnaud *et al.* (2015), the vertical detection error for the 0.46 m detection system is about 0.69 m. Mean tracer recovery rate attained 41% after the first flood and decreased to 31% for state 2. Transport distances were calculated automatically by applying the vectorization function in *ArcGis 10.3* to relay starting and ending points corresponding to the GPS points of each particle between two floods. Displacement trajectory was improved manually by using the thalweg pattern.

### 3.3.3 Geomorphic unit survey

Following Malavoi and Souchon (2002) and Schmitt *et al.* (2011), the channel was divided into different in-channel Geomorphic Units (GU): riffles, pools, and runs. Bars were added as a GU because they represent an important ecological habitat for pioneer alluvial vegetation. Surveys were conducted during low flow. Points were recorded by GPS RTK to localize the main limits between each GU. Aerial images, acquired during state 1 for another study (Eschbach *et al.*, 2017), were used to improve the delineation. For state 2, we used a ground control photographic interpretation as well as longitudinal and transverse profiles to improve the delineation between units. For each surveyed state, GU surfaces were measured using the “compute geometry” function in *ArcGis 10.3*. In addition, logjams were located for each state. No survey was implemented for state 0 because of the artificial state of the channel (no water, no natural fluvial forms like pool-riffle sequences, bars, etc.).

### 3.3.4 Survey of lateral and vertical adjustments

Lateral channel mobility for each sector was determined using a sinuosity index that was calculated by using state 0 as the reference situation. An indicator of bankfull geometry evolution was calculated using the width/depth ratio (Schmitt *et al.*, 2007) and the lateral erosion rate ( $I_{ler}$ ; Fig. 2) mentioned by Malavoi and Bravard (2010). The latter corresponds to the eroded width between two floods ( $w_e$ ) divided by the width at the initial stage of the surveyed event ( $w_{in}$ ), as shown in the following formula:

$$I_{ler} = \left[ \frac{w_e}{w_{in}} \right] * 100 \quad (1)$$

To analyze vertical adjustment between the two states, we developed an indicator of thalweg vertical variation ( $I_{vv}$ ; Fig. 2) by calculating positive residuals ( $R^+$ ) and negative residuals ( $R^-$ ) between the second longitudinal profile (state 2) and the straight trend line of the first longitudinal profile (state 1). Thalweg aggradation of the longitudinal profile is marked by  $\Sigma R^+ > \Sigma R^-$ . Conversely,  $\Sigma R^+ < \Sigma R^-$  indicates thalweg incision. The thalweg longitudinal variability was calculated using the standard deviation of the residuals of both surveys ( $I_{vv}$ ).

**Fig. 2 (Methods)**

**Figure 2: Top: details of the hydrological events during the reporting period and features of fieldwork acquisition for the two states. Bottom: process workflow (grey arrows) and results produced to achieve a global hydromorphological understanding of the channel functioning. Because the artificial side channel did not show any natural fluvial forms (i.e., pool-riffle sequences, bars, etc.), only the topographic and grain size surveys were used for State 0.**

## 4 Results

### 4.1 Multiscale sediment budgets ( $I_q$ )

#### 4.1.1 Sediment budget close to the three curvature sectors

Comparisons between the 3-D combined models are locally affected by data gaps (white holes in Fig. 3a), which results from removing of riverine vegetation and logjams during data processing. However, the colorimetric visualization of diachronic evolutions based on cloud to cloud comparisons allowed us to locate and quantify accurately erosion versus deposition processes at the channel boundaries. Suspended load was determined for 10 L of water sampled in the middle of the channel during an ecological flood (the upstream part of the water column was sampled).

After the first short and low-intensity hydrological trial flow pulse (surveying period 1, top of the Fig. 3a), the bottom of the channel in the three curvature sectors showed little morphological changes (i.e., vertical dynamics). Lateral dynamics were more pronounced, especially on concave banks (locally bank erosion reached up to 2 m; Fig. 3a, sectors A and B) and on convex banks where deposition (formation of convex bars) attained up to 1.5 m (sectors A and B). Volumes of deposition (D) reached  $131.3 \pm 16$  and  $268.7 \pm 15$  m<sup>3</sup> on sectors A and B, respectively, while volumes of erosion (E) attained on these sectors were  $246.1 \pm 16$  and  $225.6 \pm 26$  m<sup>3</sup>, respectively (Table 3). Sector C recorded the most important adjustments, with maximum erosion-deposition attaining 2.5 m. Similarly, volumes of erosion (E) in this sector were more than two times higher than for sectors A and B (Fig. 3a).  $I_q$  abstracted from sediment budgets indicates that the deposition volume (D) at sector C is 171 m<sup>3</sup> higher than at sector 1 and 33 m<sup>3</sup> higher than at sector B. Additionally, sediment output (O) reached  $235.8 \pm 22$  m<sup>3</sup> in sector C while being twice as low in sector A ( $114.8 \pm 22$  m<sup>3</sup>) and zero in sector B (because deposition was  $43.1 \pm 14$  m<sup>3</sup> higher than erosion).

Comparison between state 2 and state 1 (surveying period 2) shows an intensification of erosion in sectors A and B (from  $131.3 \pm 16$  to  $166.5 \pm 26$  m<sup>3</sup> and from  $225.6 \pm 16$  to  $310.0 \pm 26$  m<sup>3</sup>, respectively; Table 3). Compared to state 1, the lengths of lateral erosion were shorter, while channel scouring at the feet of concave banks was longitudinally more extended, as were deposition areas in convex banks. Although eroded volumes have increased in the two upstream sectors (A =  $+26 \pm 15$  m<sup>3</sup> and B =  $+84.4 \pm 26$  m<sup>3</sup>), deposited volumes were higher in sector A ( $35.2$  m<sup>3</sup>) and stayed stable in sector B ( $-0.7$  m<sup>3</sup>; Table 3). In sector C, erosion prevails over deposition, which is locally present in the downstream part of the sector. The sediment budget in this sector shows the same tendency, with a large differential between erosion ( $245.6 \pm 28$  m<sup>3</sup>) and deposition ( $66.5 \pm 12$  m<sup>3</sup>), but this trend significantly decreased in comparison with surveying period 1.

<b>Sector A</b>	SP1	SP2	<i>Diff.</i>
Deposition (D)	131,3	166,5	35,2
Output (O)	114,8	105,6	-9,2
Erosion (E)	246,1	272,1	26

<b>Sector B</b>	SP1	SP2	<i>Diff.</i>
Deposition (D)	268,7	268	-0,7
Output (O)	0	42,7	42,7
Erosion (E)	225,6	310	84,4

<b>Sector C</b>	SP1	SP2	<i>Diff.</i>
Deposition (D)	302,1	66,5	-235,6
Output (O)	235,8	176,2	-59,6
Erosion (E)	537,9	245,6	-292,3

**Table 3: Volumes (m<sup>3</sup>) of sediment deposited, eroded and output from the three sectors.**

#### 4.1.2 Sediment budget along the whole channel

At larger temporal (surveying periods 1 and 2) and spatial scales (the whole channel; Fig. 3b), cumulative erosion and deposition curves show no differentiation until 200 m on the x-axis (distance from the beginning of the channel). This distance is the transition point between the rockfill canal and the artificial side channel (Fig. 5e). From this point on, the differential between erosion and deposition rapidly reaches 500 m<sup>3</sup>, before reaching 750 m<sup>3</sup> in sector A. The erosion curve is characterized by three main growth phases corresponding to the three curvature sectors (concave bank erosion). The main one is included in the upstream part of sector C (x = 850 m; Fig. 3b). Cumulative volume of eroded sediment reached up to 3500 m<sup>3</sup> downstream of the whole artificial side channel (x = 1050 m; Fig. 3b). The deposition curve is logically below the curve for erosion, but follows the same general longitudinal tendency. The three major deposition growth peaks are spatially shifted almost 50 m downstream of the erosion growth peaks, excepted for sector B where the deposition curve is almost parallel to the erosion curve. The final difference in values between the two cumulative curves is up to 1000 m<sup>3</sup>, which represents the volume of fine sediments (part of sand, silt, clay) exported towards the old channel, the floodplain and the Old Rhine.

Based on surveying period 2, at the scale of the whole channel, we determined specific sediment budgets (m<sup>3</sup>) for each type of component of the channel boundaries (m<sup>2</sup>): morphological bank types (straight, concave, convex) and channel bottom (Fig. 5b.). Consequently, sediment budgets for this analysis are expressed in m<sup>3</sup>/m<sup>2</sup> (Fig. 5b). These results are based on comparisons of elementary reaches and allowed the linking of vertical and lateral volumetric evolutions. It appears that erosion is higher for banks than for the channel bottom, with a maximum volume of lateral erosion that could attain 0.26 m<sup>3</sup>/m<sup>2</sup> (x = 400 m), while vertical evolution does not exceed 0.05 m<sup>3</sup>/m<sup>2</sup> (x = 390 m). This tendency not only affects the concave banks, but also the straight banks, locally impacting the entire cross section, including also convex banks (especially

in sector C). Conversely, deposition logically occurs mostly at the channel bottom, and locally on banks (convex but also locally concave banks). In addition, it is worth noting that lateral and vertical erosion-deposition volumes are significantly low between 500 and 650 m.

### Fig. 3 (Results\_BS)

5 **Figure 3: (a) Comparison of 3-D combined models using M3C2 during both surveying periods on the three curvature sectors, including sediment budgets. (b) Global cumulative sediment budget for the whole channel during both surveying periods.**

#### 4.2 Lateral channel evolution ( $I_{ler}$ )

In the three curvature sectors, the position of the thalweg shifted locally up to 7 m towards concave banks, while mean relative sinuosity increased from 1.0 to 1.05. This indicates curvature accentuation band concave bank erosion and the lateral  
10 migration of the thalweg during floods (Schumm, 1968; Fig. 4a). Bankfull surfaces tended to converge over the two periods to between 18.60 and 20.89 m<sup>2</sup> (Fig. 4c). Widths increased in all cross sections, especially in CS22 (+1.67 m; +16%) and CS34 (+4.39 m; + 52%). The W/D ratio increased from 5.68 to 7.61 in CS7, from 4.55 to 7.68 in CS22, and from 3.77 to 8.96 in CS34. Lateral erosion ( $I_{ler}$ ) was very low for CS7 with a value of 0.3% during the surveying period 1 and 2.5% during the surveying period 2. For CS22, lateral erosion increased from 6.1% to 9.2% between the two surveying periods.  
15 CS34 recorded the highest lateral erosion related to the first surveying period (44.2%) followed by a lowest value (5.0%) for the second surveying period.

### Fig. 4 (Results\_LER)

**Figure 4: Examples of lateral channel evolution in the three curvature sectors, with (a) thalweg evolution, (b) examples of diachronic cross sections, (c) resulting indicators of lateral channel evolution ( $I_{ler}$ ) and bankfull specific stream power.**

#### 20 4.3 Vertical variation ( $I_{vv}$ )

The comparison of longitudinal profiles during surveying period 2 (Fig. 5e) allows the identification of thalweg vertical variations, characterized by aggradation sectors that correspond notably to riffles (e.g., x = 240; 290; 500; 980 m) and scour sectors that correspond to pools (e.g., x = 225; 260; 425; 725; 875 m). A stable sector, covering mostly sector B, was identified, extending from 530 m to 700 m. No logjams are located between 425 and 830 m. In fact, these are especially  
25 located in the curvature sectors A and C (Fig. 5e). These sectors are also characterized by high vertical evolution, mostly erosion.

Results of the global longitudinal evolution ( $I_{vv}$ ) reveal a longitudinal reversal characterized by three main tendencies (Fig. 5a). The first sector is characterized by a thalweg incision tendency that extends from 225 m to 510 m ( $\Sigma R^+ < \Sigma R^-$ ). Low cumulative residual variations between 510 m and 645 m indicate vertical stability. Conversely, the thalweg is characterized by an aggradation tendency between 645 m and 1150 m. The thalweg longitudinal variability shows an important  
5 augmentation in the difference, from original values around 0.1 for the whole channel at state 0. Comparison between state 1 and 2 shows that the difference increased to 0.15 - 0.41 and 0.15 - 0.28, for the upstream and downstream sectors, respectively. By contrast, these values remained around 0.1 in the identified stable sector.

#### 4.4 Evolution of geomorphic units ( $I_{gu}$ )

The evolution of geomorphic units was surveyed to assess gain or loss of habitat diversity. The total area covered by all  
10 geomorphic units increased from 10 000 to 12 520 m<sup>2</sup> between the two surveys (+25.2%; Fig. 5g). This tendency reveals a general channel enlargement (from 2 to 4 m on average). Among the four geomorphic units: riffle, pool, run and bar, major growth of +1100 m<sup>2</sup> occurred for runs (18.83%). Pools increased by +550 m<sup>2</sup> (60.5%), closely followed by riffles (+450 m<sup>2</sup>; 30.5%) and bars (+420 m<sup>2</sup>; 23.74%).

#### 4.5 Grain size variation and bedload dynamics ( $I_{bd}$ )

15 With regards to the combined grain size results from the two surveying periods (Fig. 5c), we observed an average  $D_{50}$  of 15.3  $\pm$  5 mm, except for samples located at a longitudinal distance of 576 and 602 m where  $D_{50}$  is higher than 50  $\pm$  15 mm (standard deviation). Indeed, this sector exhibits a significantly coarser grain size resulting from a large median in-channel gravel bar inherited from past Rhine dynamics, as revealed by the historical study (Fig. 5f; Eschbach *et al.*, 2018). The gravel bar extends from 530 to 815 m under the artificial side channel, and outcrops in the riverbed between 530 and 675 m.  
20 Results of RFID tracking are given on Fig. 6 and used as an indicator for bedload dynamics ( $I_{bd}$ ). According to Lamarre and Roy (2008) and Milan and Large (2014), tracers travel distances recorded in the first survey (S1) are not representative because of an overexposure phenomenon caused by particle setup. This phenomenon is well known and explains why the particles move more easily during the first flood. The results of the two survey periods do not show any significant link between transport distances and particle grain size, indicating a non-selective mobility. The distribution of tracer clouds is

relatively homogeneous through time at transect 8 as well as for survey 1 in transect 9 (Fig. 6b). A significant decrease in travel distances is observed for survey 2 (S2) at transect 9 (Figs. 5d and 6). At this location, mean transport distances drop down from 24.2 to 12.4 m while they increase in transect 8 (from 24.7 to 31.5 m; Fig. 6a). Fig. 6c shows a significant decrease of cumulative distances for transect 9 while they increase in transect 8. This figure shows also that some tracers are slowed down by riffles. By contrast with the recovery rate recorded for transects 8 and 9 (47% and 46%, respectively), the rate for transect 10 (6.5%) is very low, which dampens the robustness of results. Nevertheless, it can be noted that mean transport distances for this transect tended to decrease, from 26.8 to 17.6 m. Results were summarized on Fig. 5d and, to improve readability, consider cumulative travel distance according to the grain size of the tracers.

**Fig. 5 (Results\_PL)**

10 **Figure 5: Longitudinal evolution between (a) vertical variation of cumulative residuals (every 10 m), (b) sediment budget by bank types and channel bottom (every 20 m), (c) grain size, (d) bedload mobility, (e) longitudinal profile evolution, (f) inherited sediment deposits, and (g) geomorphic units evolution.**

**Fig. 6 (Results\_PIT)**

15 **Figure 6: (a) Box-plots of tracer travel distances for the two surveying periods. (b) Travel distances of tracers detected during both surveying periods and according to the grain size of RFID particles. (c) Frequency distribution of the particles location at state 1 and state 2 linked with the evolution of the thalweg longitudinal profiles.**

## 5 Discussion

### 5.1 Combining geodetic and geomorphic methods to survey restored river dynamic side channels

20 Although TLS and SfM photogrammetry methods are able to generate high spatial resolution 3-D data to survey bank erosion, the point clouds obtained must be completed by bathymetric surveys to produce complete 3-D models, which can be chronologically compared. Bathymetric changes can be monitored by a set of cross sections generated by Total Station and GPS measurements, but this low resolution method is time-consuming (Lane *et al.*, 1994; Barker *et al.*, 1997; Brasington *et al.*, 2000a; Bangen *et al.*, 2014b). For instance, covering 1 km length of an artificial side channel with cross sections spaced every 20 m and composed each of 5 to 7 leveling points, requires six people during one to two days, depending on field constraints (riparian vegetation, logjams, weather, etc.). During surveys, the positions of topographic points are accurately determined to reveal lateral slope changes, as recommended by Ferguson *et al.* (1992) and Lane *et al.* (1994). Topo-bathymetric surveys are usually used to infer geomorphic features such as longitudinal profiles and cross sections, and to

delineate geomorphic units (riffles, pools and runs). In addition, diachronic topographic points allow the determination of geomorphic changes with an acceptable resolution along reaches (~0.5 – 1 m). Consequently, hybrid methods based on SfM-photogrammetry and Total Station surveys is a relevant compromise between a spatial resolution that is both satisfactory for managers and allows geomorphic interpretation, and field constraints (Westaway *et al.* 2003; Rayburg *et al.* 2009; Legleiter 5 2012; Bangen *et al.* 2014a), especially to survey dynamic side channels.

The major limitation of 3-D modeling applied to a fluvial context is the presence of dense riparian vegetation and logjams, which may partially cover banks and generate gaps in point clouds. To avoid this constraint, we recommend carrying out surveys during winter, when vegetation is the least developed. Furthermore, to survey intra-annual changes and improve stereoscopic pair acquisition (especially when vegetation is high and where numerous logjams are present in the channel), 10 we recommend combining terrestrial and local aerial photogrammetry obtained using drones (Lejot *et al.*, 2007; Raclot *et al.*, 2005), UAV-Photography (Flener *et al.*, 2013) or L-shaped bracket (Grussenmeyer *et al.*, 1996). However, although this method is effective for geomorphologic monitoring of rugged terrain, it is difficult to implement at large spatial scales (more than one kilometre). Its deployment is optimal for sections of a few hundred meters. Beyond this, the acquisition and processing time would become excessively high.

15 Concerning the process workflow, our study shows that the M3C2 method is a powerful tool to visualize morphological changes and to produce detailed sediment budgets at different spatial and temporal scales. However, in the curvature sectors, we recommend anticipating, as much as possible, future lateral evolution so that the area of lateral channel mobility is included into the point cloud scene.

Despite the implementation constraints in alluvial artificial side channels (dense riparian vegetation, accessibility), it seems 20 promising to explore the potential of topographic/bathymetric Light Detection And Ranging (LIDAR) to survey morphological changes at larger spatial scales (Brasington *et al.*, 2000; Legleiter, 2012; Bizzi *et al.*, 2016). In reaches where the water depth is less than 2 m and the turbidity is low, this method has great potential, especially to improve the resolution of bathymetric point clouds (Gao, 2009; Flener *et al.*, 2013; Lague *et al.*, 2013; Woodget *et al.*, 2015).

## 5.2 Spatio-temporal adjustments identified by using functioning indicators

Main changes were identified in the downstream part of the channel (sector C) where the width was lower (Fig. 1d, sector C in 2014). This induced a high value of stream power ( $22.2 \text{ W.m}^{-2}$ ) explaining why morphological changes, especially enlargement, were the highest in this sector (Fig. 3a, sector C, surveying period 1). In comparison, stream power was lower  
5 farther upstream (sectors A and B), between  $15.4$  and  $18.1 \text{ W.m}^{-2}$  (Fig. 4c). For the survey period 2 (2015-2016), eroded and deposited volumes are slightly more important in sectors A and B but decreased significantly in sector C by comparison to the first survey period (2014-2015). This may be explained by the fact that, for sector C, stream power dropped to about  
10  $15 \text{ W.m}^{-2}$ , which is close to the stream power values of sectors A and B, while for the latter sectors the stream power drop was very low (Fig. 4c). Despite a negative sediment budget for the downstream part of sector C (Fig. 3a, sector C, survey  
15 period 2 and Fig. 5b),  $I_{vv}$  reveals a general tendency of thalweg aggradation (Fig. 5a). This is especially because of the high lateral enlargement that supplied gravel and induced sedimentation on the channel bottom. Locally, both bank erosion and vertical evolution are also strongly linked to the presence of numerous logjams (Fig. 5e).

Although only short floods occurred during the trial period (survey period 1 - 2014-2015) with a moderate initial bankfull stream power of about  $15$  to  $22 \text{ W.m}^{-2}$ , the channel experienced important morphological adjustments. Part of this result is  
20 explained by the fact that the channel was newly excavated and featured no sediment sorting as, e.g., surface pebble/gravel pavement. In absence of upstream bedload inputs,  $I_q$  revealed that bank erosion provided large bedload volumes into the channel, especially in concave and straight banks (Fig. 5b). While fine sediments are exported out of the channel towards the old channel, the adjacent floodplain and/or the Rhine (Fig. 3b), mean bedload transport distances and dispersion are relatively low at around  $20$  m and  $85$  m, respectively. Consequently, coarse sediment inputs by lateral erosion are deposited  
25 into the channel and contributed significantly to the diversification of the geomorphic units ( $I_{gu}$ ; Fig. 5g), as also shown by Eschbach *et al.* (2017) and studies on other rivers (Hassan *et al.*, 2017; MacWilliams *et al.*, 2006; Sear 1996). On the other hand, diversification of fluvial forms (lateral or median bars, riffles) controlled gravel mobility, as was also showed by Chardon *et al.* (in revision) in the Old Rhine. At transect 8, gravel was trapped in a large pool and in the upstream part of a riffle (Fig. 6c, x-axis =  $450$  to  $470$  m). By contrast, travel distances decreased in transect 9 (from  $24.2$  to  $12.4$  m) while the  
longitudinal profile, as revealed by  $I_{vv}$  (Fig. 5a), showed no important changes (Fig. 5e). This results from the presence of the

old bar formed by coarser sediments (mean b-axis size is 50 mm; Fig. 5c), which crops out on the channel bottom and the banks, close to transect 9 (Fig. 5f, between 530 and 675 m). A significant part of PIT tags remained trapped in the pavement for state 2, regardless of the particle size (Fig. 6c). This inherited bar also significantly influenced all morphodynamic indicators, including bedload transport. On the one hand grain size is 2.5 higher than the one along the channel (bottom and bank footslopes) and controls both vertical ( $I_{vv}$ ; see stability sector in Fig. 5a) and lateral evolution ( $I_{ler}$ ; see bank erosion in Fig. 5b). On the other hand, bedload transport of all grain size classes is considerably lower ( $I_{bd}$ ; Fig. 5d). The second specific sector is located in the upstream part of the artificial side channel, directly downstream of the “*Rockfill canal*”. It is characterized by the presence of a sandy palaeochannel, explaining the predominance of sand in both banks and the channel bottom, thereby favoring lateral and vertical erosion (see Fig. 5f, between 190 and 245 m). This is highlighted by the comparison of longitudinal profiles (Fig. 5e) and by the  $I_{vv}$  indicator (Fig. 5a) which showed the formation of an important scour pool (Fig. 5e). This is also enhanced by a coarse sediment deficit in this upstream extremity of the channel, linked to the absence of bedload input at the floodgate.

The set of functioning indicators listed in Table 4 summarizes the evolutionary trends that occurred in the artificial side channel of the Rohrschollen Island. Initial restoration state refers to the survey campaign that took place immediately after the digging of the artificial side channel and before the first ecological floods (see “Initial State” column in Table 4). The state “targeted” is based on the restoration objectives defined by the manager and scientists working on the restoration environmental monitoring. This highlights the fact that it is essential to define (quantifiable) objectives for restorations using the same indicators than for post-restoration monitoring (Pander and Geist, 2013; Stammel *et al.*, 2011), as well as for the pre-restoration survey. Results of each indicator developed in the present study are summarized by arrows determining the observed evolutionary trends (see “Post-Resto” column in Table 4). Identified long-term trends are based on the interpretation of results presented in this paper and on the historical study developed in Eschbach *et al.* (2018). Table 4 highlights the systematic increase of all indicators, reflecting the significant morphological diversification that occurred during the first two years following the restoration. Obviously, our results currently reflect post-restoration transitory adjustments, which were exacerbated in the newly excavated, undersized, side channel. This tendency should decrease in the coming years when a dynamic equilibrium will be reached. This research will be continued with additional surveys to

confirm the observed trend. In this study, we mainly focused on geomorphological indicators, but ecological indicators should also be considered to assess the “dynamic equilibrium” targeted by the project, as shown for instance by Woolsey *et al.* (2007), Nielsen and Jorgensen (2013) or Pander and Geist (2013).

INDICATORS	NAME	METRICS	METHODS	State		Evolutionary trend	
				Initial state	Target	Post-Resto <sup>o</sup>	Long-term Target
Geomorphic units evolution	$I_{gu}$	Surfaces of geomorphic units (habitats)	Mapping of riffle, pool, run, logjams (GPS; ArcGis)	-	+	↗	→
Bedload dynamics	$I_{bd}$	Transport distances and dispersion	RFID tracking	o	+	↗	→
Vertical variation	$I_w$	Vertical evolution trend	<ul style="list-style-type: none"> <li>Thalweg longitudinal profile</li> <li>Cumul of the residuals</li> </ul>	-	+	↗	→
Lateral erosion	W/D $I_{ler}$	Width/depth ratio $I_{ler} = \left[ \frac{w_e}{w_{in}} \right] * 100$	Cross section (Total Station)	-	+	↗	→
Sediment budget	$I_q$	Volumes (differences) of eroded and deposited sediments	TLS/SfM-photogrammetry/Total Station and DoD	o	+	↗	→

o Any survey      ↘ Decrease  
 + High            ↗ Increase  
 - Low             → Stable

5

**Table 4: Functioning indicators used to survey morphodynamic adjustments in Rohrschollen artificial side channel. Metrics, methods and evolutionary trend are summarized as a management toolkit to monitor the restoration-related changes. Long-term means >10 yr.**

### 5.3 Guidelines to improve restoration efficiency and sustainability

10 Restoration efficiency can be assessed by using functioning indicators (Rinaldi *et al.*, 2017). Such indicators allow investigators to (1) determine the optimal functioning considering the targeted restoration goals (indicators should be used to define quantifiable restoration objectives, and to implement the pre-restoration survey), (2) assess morphological adjustments and identify the controlling factors, and (3) help make the restoration action more sustainable. Managers are thus encouraged to use a set of functioning indicators to monitor long-term post-restoration changes, in order to check if the

15 restoration effects are sustainable and, if not, to engage in a subsequent restoration step (if possible). Thus, indicators are useful to monitor processes at different spatial and temporal scales to follow changes but also to adapt management principles if needed. The indicators developed in our study contribute to the construction of an easily reproducible,

transposable (for current or future restoration) and transferable (to managers) toolbox, as suggested by Fryirs *et al.* (2018) or Díaz-Redondo *et al.* (2018). This is essential to assess both the efficiency and sustainability of any restoration. Using both geodetic and geomorphological surveys makes the monitoring strongly relevant, especially in case of a process-based restoration for which morphological dynamics are intense and frequent. In this context, hybrid methods can provide an accurate understanding of morphodynamics and identify main controlling (and limiting) factors. Because the proposed monitoring program can be relatively time-consuming and costly, managers may also simplify it for annual surveys, and use the full set for accurate surveys that could be realized every 3 to 5 yr and/or after intense or long floods.

This study underlines the importance of self-forming channel processes through bank erosion and vertical adjustments (bedload dynamics, riffle-pool sequences, bars development, etc.), especially in the case of artificial side channels mimicking natural forms and functioning. Bank erosion favors (i) sediment supply to the channel (including the coarse fraction), (ii) frequent deposition and (iii) longitudinal and lateral diversification. This underscores that it is preferable to excavate artificial side channels composed by a thin layer of fine sediments, or allow dynamic floods to excavate new channels into the floodplain, rather than reactivating clogged palaeochannels mostly filled by large quantities of fine sediments that can even be polluted (Eschbach *et al.*, 2018). Consequently, it enhances ecological potential because bed grain size diversification, bedload dynamics and geomorphic unit evolution represent key processes for aquatic habitats (Beisel *et al.*, 1998; Riquier *et al.*, 2017), especially in restored channels (Morandi *et al.*, 2014, Fryirs *et al.*, 2018). All these processes are favored by logjams, which in turn may intensify, and/or be intensified by, bank erosion (Beechie *et al.*, 2010; Stammel *et al.*, 2011; Choné and Biron, 2016). In any case, taking into account long-term temporal trajectories (>10 yr) can help in identifying areas for which potential post-restoration adjustments may be significantly different (areas composed by relatively small or large grain size). This should improve defining restoration targets and help interpreting post-restoration evolution thanks to morphodynamic indicators.

Indicators could help identifying sources of weaknesses in long-term functioning related to process-based restoration, for example, a deficit in the sediment budget or an area where erosion rates are particularly high. This raises the question of potential gravel augmentation, which is one of the main limitations of the kind of projects studied in this paper, especially in terms of management cost and sustainability (Harvey *et al.*, 2005).

## 6 Conclusion and perspectives

Two major outcomes are obtained in our study. First, the results showed that TLS and SfM-photogrammetry are relevant techniques to monitor morphological changes in fluvial environments, especially for large river side channels. SfM-photogrammetry is recommended for its cost and the flexibility of the acquisition, especially in dynamic side channels showing important and rapid changes. Producing and comparing complete 3-D models by combining terrestrial point clouds with bathymetric data provide a reliable way to identify and assess accurately morphological changes following floods at different spatial scales. Second, we demonstrate the importance of combining geodetic tools with geomorphical methods to develop straightforward morphodynamic indicators able to capture morphodynamic adjustments and processes. This is relevant for assessing both efficiency and sustainability of restoration projects. Indeed, monitoring fluvial changes in restoration activities requires using relevant functioning indicators, which should be determined in the early steps of restoration projects to define specific quantifiable objectives. Such indicators allow investigators to (i) monitor short- (<5 yr) and mid-/long- (at least 5 or >10 yr) term morphological changes, (ii) determine on this basis if the objectives are achieved and over which temporal scale, and (iii) allow inter-project comparison. As 1-D or 2-D hydraulic and morpho-sedimentary modelling is an increasingly used tool to predict changes and support manager decisions, it seems relevant that these modelling approaches integrate the proposed functioning indicators. In the case of an alluvial artificial side channel with water depths less than 2 m, topographic/bathymetric LIDAR should strengthen bathymetric datasets and simultaneously reduce the duration of field work.

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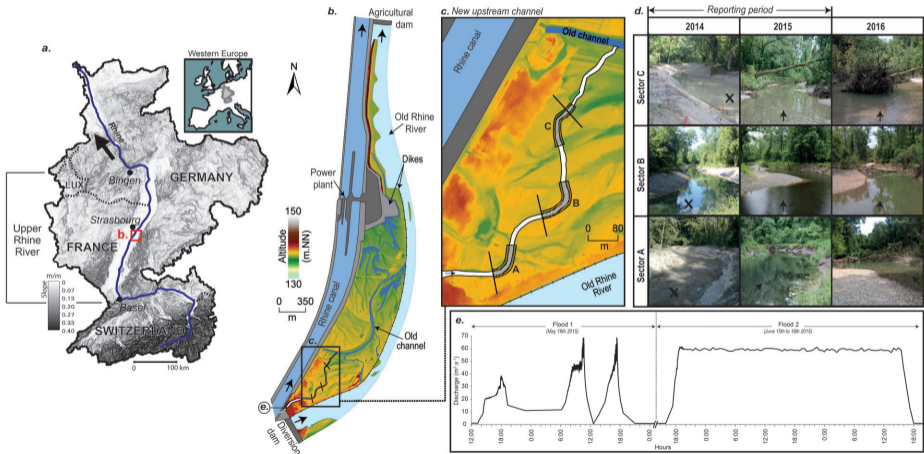
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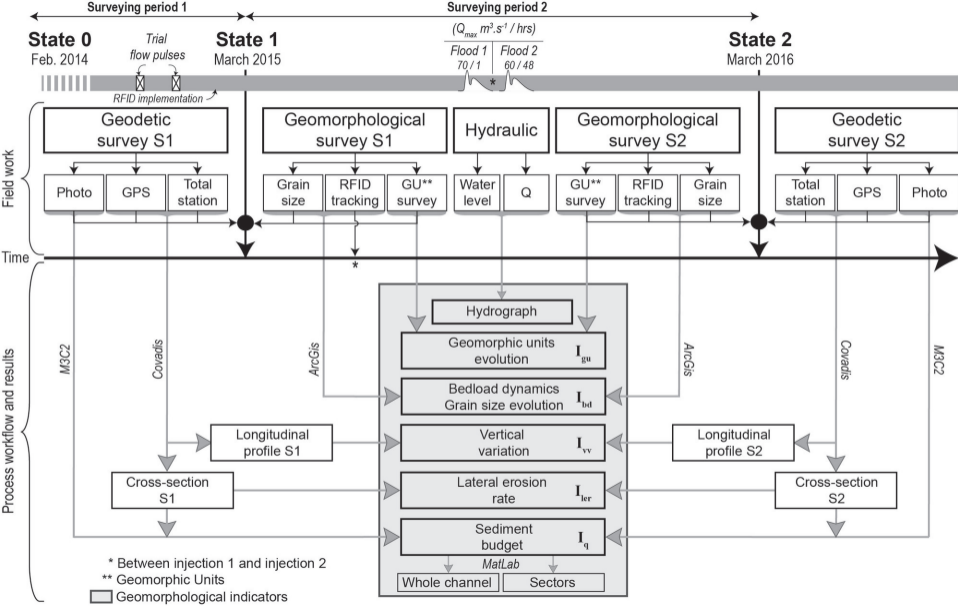
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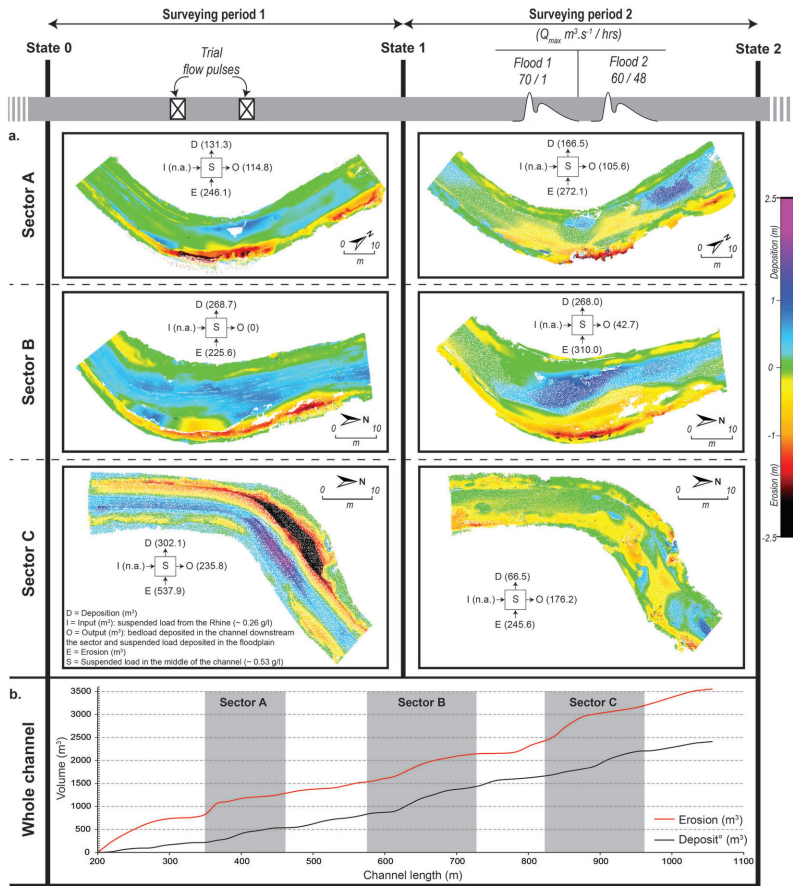
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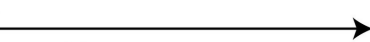
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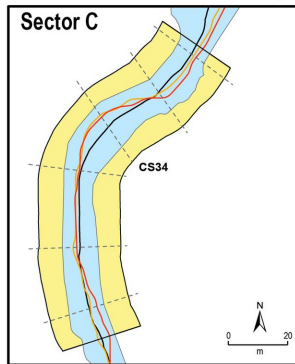
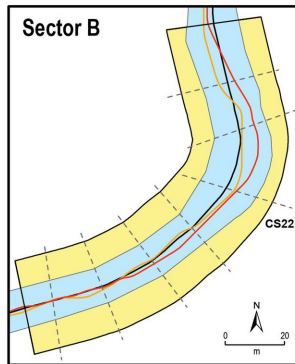
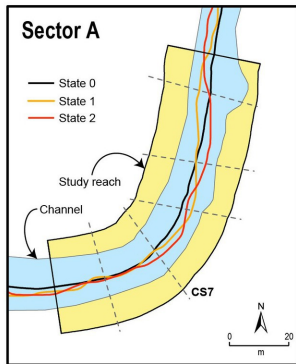




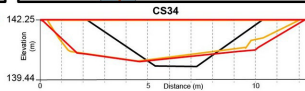
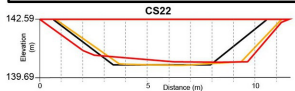
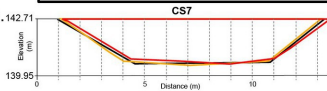


Flow direction 

a.

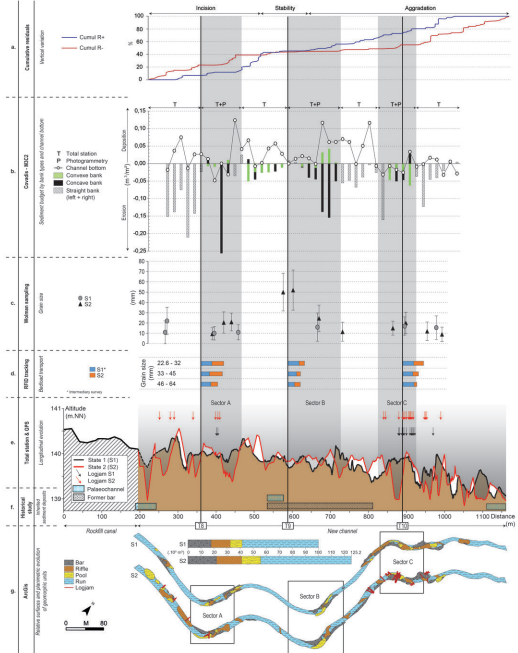


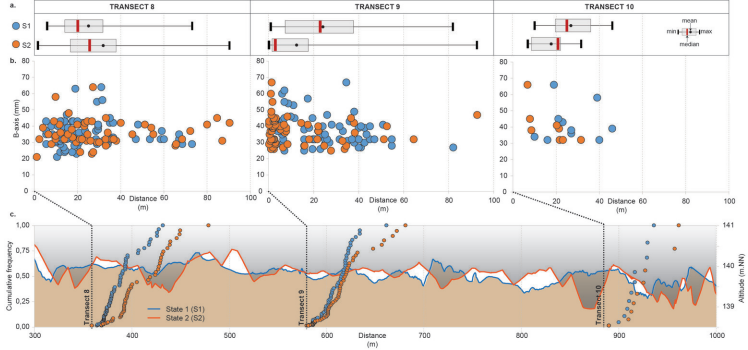
b.



c.

	State 0	State 1	State 2	State 0	State 1	State 2	State 0	State 1	State 2
Surface (m <sup>2</sup> )	20.39	19.90	20.89	16.56	17.31	19.17	11.16	16.19	18.60
Width (m)	12.26	12.30	12.61	10.46	11.10	12.13	8.52	12.29	12.91
Average depth (m)	2.16	1.64	1.66	2.30	1.56	1.58	2.21	1.32	1.44
w/d	5.68	7.39	7.61	4.55	7.12	7.68	3.77	9.33	8.96
Specific stream power (W/m <sup>2</sup> )	15.40	15.40	15.00	18.10	17.00	15.60	22.20	15.40	14.60
Lateral erosion (%)	0.3		2.5	6.1		9.3	44.2		5.0





Authors	Date	Context	Approach	Geomorphological dimensions & components						DEM methods		Hydromorphological methods						
				Channel types	River restoration	Lateral channel restoration	Geomorphological survey	Methodological studies <sup>a</sup>	Lateral adjustments	Vertical adjustments	Bedload dynamics	Hydrodynamics	Geomorphic units evolution	Lasergrammetry	Photogrammetry	Levelling (cross-section)	RFID tracking	Hydraulic modelling
Collins and Moon,	1979	WR			•						•							
Dickinson and Scott,	1979	MR		•							•							
Lane et al.,	1994	BR			•						•	•						
Martin and Church,	1995	BR		•									•					•
Barker et al.,	1997	MR			•						•							
Pyle et al.,	1997	PGR			•						•							
Heritage et al.,	1998	MR			•						•							
Westaway et al.,	2000	BR			•						•							
Brasington et al.,	2000	BR			•		•				•	•						
Chandler and Shiono,	2001	TF			•						•							
Lane et al.,	2001	TF			•						•							
Chandler et al.,	2002	PGR			•						•	•						
Fuller et al.,	2003	WR		•							•							•
Formann et al.,	2007	BR	•								•	•					•	
Lejot et al.,	2007	BR	•		•	•					•	•						• <sup>b</sup>
Wheaton et al.,	2009	BR			•						•	•	•					•
Lague et al.,	2013	MR			•						•							
Bangen et al.,	2014a	SAC			•								•					•
Javernick et al.,	2014	BR			•						•							
Casado et al.,	2015	WR			•						•							• <sup>c</sup>
Chapuis et al.,	2015	WR		•								•		•	•			
Kuo et al.,	2015	BR		•							•							•
MacVicar et al.,	2015	SAC	•		•								•		•			
Brunier et al.,	2016	BR		•							•							•
Dietrich,	2016	SAC	•		•	•					•	•	•					
Jozkow et al.,	2016	MS		•							•							
Belletti et al.,	2017	WR			•													• <sup>d</sup>
Marteau et al.,	2017	MS	•		•							•						•
Cook,	2017	WR			•						•	•						
Wyrick et al.,	2016	WR			•													• <sup>e</sup>
Jugie et al.,	2018	SAC		•							•	•						• <sup>f</sup>

**BR** Braiding River

**SAC** Small Alluvial Channel

**TF** Tilting Flume

**MR** Meandering River

**MS** Mountain Stream

**PGR** ProGlacial River

**WR** Wandering River

<sup>a</sup> Results focused on methodological innovation (accuracy, methodological comparison...)

<sup>b</sup> Bathymetric models

<sup>c</sup> Unmanned Aerial Vehicles (UAV) and Artificial Neural Networks (ANN)

<sup>d</sup> Geomorphic Units Survey (GUS) and Morphological Quality Index (MQI)

<sup>e</sup> Near-census approach and decision-tree classification

<sup>f</sup> Photo-Electronic Erosion Pin (PEEP)

	Sector A		Sector B		Sector C	
	Left bank	Right bank	Left bank	Right bank	Left bank	Right bank
<b>Number of image</b>	128	90	129	101	74	83
<b>Number of GCP</b>	5	6	7	8	10	7
<b>Reprojection error</b> <i>(Standard deviation m)</i>	0.001	0.02	0.02	0.05	0.03	0.01
<b>Reprojection error</b> <i>(Standard deviation pixels)</i>	0.18	0.23	0.49	0.2	0.19	0.12

<b>Sector A</b>	SP1	SP2	<i>Diff.</i>
Deposition (D)	131,3	166,5	35,2
Output (O)	114,8	105,6	-9,2
Erosion (E)	246,1	272,1	26

<b>Sector B</b>	SP1	SP2	<i>Diff.</i>
Deposition (D)	268,7	268	-0,7
Output (O)	0	42,7	42,7
Erosion (E)	225,6	310	84,4

<b>Sector C</b>	SP1	SP2	<i>Diff.</i>
Deposition (D)	302,1	66,5	-235,6
Output (O)	235,8	176,2	-59,6
Erosion (E)	537,9	245,6	-292,3

INDICATORS	NAME	METRICS	METHODS	State		Evolutionary trend	
				Initial state	Target	Post-Resto°	Long-term Target
Geomorphic units evolution	$I_{gu}$	Surfaces of geomorphic units (habitats)	Mapping of riffle, pool, run, logjams (GPS; ArcGis)	-	+	↗	→
Bedload dynamics	$I_{bd}$	Transport distances and dispersion	RFID tracking	○	+	↗	→
Vertical variation	$I_w$	Vertical evolution trend	{ Thalweg longitudinal profile Cumul of the residuals	-	+	↗	→
Lateral erosion	W/D $I_{ler}$	Width/depth ratio $I_{ler} = \left[ \frac{w_e}{w_{in}} \right] * 100$	Cross section (Total Station)	-	+	↗	→
Sediment budget	$I_q$	Volumes (differences) of eroded and deposited sediments	TLS/SfM-photogrammetry/Total Station and DoD	○	+	↗	→

○ Any survey  
+ High  
- Low

↘ Decrease  
↗ Increase  
→ Stable