



HAL
open science

Constraining the Solomon Sea as a source of Al and Mn to the Equatorial Undercurrent

Susanna Michael, Joseph Resing, F. Lacan, Nathaniel Buck, Catherine Pradoux,
Catherine Jeandel

► **To cite this version:**

Susanna Michael, Joseph Resing, F. Lacan, Nathaniel Buck, Catherine Pradoux, et al.. Constraining the Solomon Sea as a source of Al and Mn to the Equatorial Undercurrent. *Deep Sea Research Part I: Oceanographic Research Papers*, 2021, 174, pp.103559. <10.1016/j.dsr.2021.103559>. <hal-03358780>

HAL Id: hal-03358780

<https://hal.science/hal-03358780v1>

Submitted on 4 Oct 2021

HAL is a multi-disciplinary open access archive for the deposit and dissemination of scientific research documents, whether they are published or not. The documents may come from teaching and research institutions in France or abroad, or from public or private research centers.

L'archive ouverte pluridisciplinaire **HAL**, est destinée au dépôt et à la diffusion de documents scientifiques de niveau recherche, publiés ou non, émanant des établissements d'enseignement et de recherche français ou étrangers, des laboratoires publics ou privés.



HAL Authorization

1 **Constraining the Solomon Sea as a Source of Al and Mn to the Equatorial**
2 **Undercurrent**

3 Authors: Susanna Michael^{a,b,*}, Joseph Resing^{b,c}, Francois Lacan^d, Nathaniel Buck^{b,c}, Catherine
4 Pradoux^d, Catherine Jeandel^d

5 Affiliations:

6 ^a School of Oceanography, University of Washington, Seattle, WA 98195 USA
7 (smicha@uw.edu)

8 ^b Joint Institute for the Study of the Atmosphere and Ocean, University of Washington, 3737
9 Brooklyn Ave NE, Seattle, WA 98105 USA

10 ^c NOAA Pacific Marine Environmental Laboratory, 7600 Sand Point Way NE, Seattle, WA
11 98115 USA
12 (joseph.resing@noaa.gov; nathan.buck@noaa.gov)

13 ^d LEGOS, Université de Toulouse, CNRS, CNES, IRD, UPS (Toulouse), France
14 (francois.lacan@legos.obs-mip.fr; catherine.pradoux@legos.obs-mip.fr;
15 Catherine.Jeandel@legos.obs-mip.fr)

16 *Corresponding author: smicha@uw.edu
17

18 **Highlights**

19 Al and Mn are not significantly enriched during transit through the Solomon Sea.

20 Fluxes of Al and Mn into and out of the Solomon Sea are almost equal.

21 Al and Mn are elevated near continental shelves and margins in the Solomon Sea.

22 Local enrichments must be balanced by boundary exchange and scavenging processes.

23 Water exiting the Solomon Sea accounts for ca. half the flux of Al and Mn in the EUC.

24 **Abstract**

25 Total dissolvable and dissolved aluminum (TDAI, DAI) and manganese (TDMn, DMn)
26 concentrations were measured at 12 stations in and around the Solomon Sea in 2012 as part of
27 the GEOTRACES GP-12 cruise. These data were used to determine the potential for the
28 Solomon Sea to act as a source of Al and Mn to the Equatorial Undercurrent (EUC). From a net
29 budget perspective, waters entering the Solomon Sea at the time of the cruise were already
30 enriched in Al and Mn, and as that water transited through the Solomon Sea, further net
31 enrichments were small compared to overall concentrations of these metals. Despite this overall
32 balance, on a local scale, we observed enrichment of Al and Mn at stations located near
33 coastlines, most likely caused by sediment scouring by strong currents. Calculated fluxes of DAI,
34 and TDAI out of the Solomon Sea relative to the EUC are large enough to account for about
35 three quarters of their respective budgets within the EUC, while the DMn and TDMn fluxes
36 exiting the Solomon Sea can only account for about half of their respective budgets in the EUC.
37 These fluxes are subject to high temporal variability and to uncertainty of the relative
38 contributions of Northern and Southern Hemisphere water mass to the EUC.

39 **Keywords:**

40 Aluminum, Manganese, Solomon Sea, Equatorial Undercurrent, GEOTRACES

41

42

43 **1. Introduction**

44 The Equatorial Undercurrent (EUC) flows at ~200 m depth along the equator at a rate of 20-30
45 Sv, transporting water, nutrients, and trace elements (e.g., aluminum (Al), and manganese (Mn),
46 and iron (Fe)) from the Western Pacific to the Eastern Pacific in less than a year (Tsuchiya et al.,
47 1989). There, it shoals in the photic zone of the High Nutrient, Low Chlorophyll (HNLC) region
48 of the eastern equatorial Pacific—where ~20% of the world's new primary productivity takes
49 place (Coale et al., 1996). Understanding sources of trace elements to the EUC is thus important
50 to better understanding the factors that contribute to primary productivity in the eastern
51 equatorial Pacific and its contribution to the global carbon cycle.

52 Concentrations of Al, Mn, (and Fe) are elevated in the EUC, relative to open ocean values
53 (Coale et al., 1996; Gordon et al., 1997; Kaupp et al., 2011; Slemmons et al., 2010, 2012) with
54 concentrations increasing westward. In the western equatorial Pacific, the major water sources to
55 the EUC are two low-latitude western boundary currents, the New Guinea Coastal Undercurrent
56 (NGCU) coming from the south and the Mindanao Current from the north, with the NGCU being
57 the more important of the two (e.g., Tsuchiya et al., 1989; Grenier et al., 2011, 2013). The
58 NGCU originates in the Solomon Sea (Fig. 1) where it comes in contact with the coastlines of
59 volcanic islands with abundant natural and anthropogenic runoff. As a result, the Solomon Sea is
60 considered to be an important source of trace metals, especially Al and Fe, to the NGCU and
61 thus the EUC. This conclusion is supported by both modeling studies and geochemical
62 measurements within the current (Lacan and Jeandel, 2001, 2005; Mackey et al., 2002a,b;
63 Slemmons et al., 2009; Kaupp et al., 2011; Qin et al., 2015; Pham et al., 2019). Here we report on
64 dissolved and total dissolvable Al (DAI; TDAI) and Mn (DMn; TDMn) collected from seven
65 stations within the Solomon Sea and five stations just outside of it during the 2012 PANDORA
66 cruise (GEOTRACES GP-12). While several trace metal profiles have been collected within the
67 Solomon Sea and in the neighboring Bismarck and Coral Seas (Mackey et al., 2002a,b; Obata et
68 al., 2008), those studies were not part of a broader interdisciplinary study. The PANDORA

69 cruise discussed here included a major physical oceanographic component that examined the
70 major currents flowing through the Solomon Sea (Ganachaud et al., 2017). The combination of
71 these chemical and physical oceanographic data enables examination of the trace metal budget of
72 the Solomon Sea, and in turn, its importance to the trace metal budget of the EUC. The data
73 presented here suggest that the waters entering the Solomon Sea were enriched in trace metals
74 prior to entering the basin and that their transit through the basin resulted in only a minor net
75 increase of Mn and Al to these waters.

76 *1.1 Geographic Setting: The Solomon Sea*

77 The Solomon Sea is a semi-enclosed basin bounded by the islands of Papua New Guinea (PNG)
78 to the west, New Ireland and New Britain to the north, and the Solomon Islands to the east, and
79 is open to the southeast (Fig. 1). An important oceanographic characteristic of this region is the
80 flow of the New Guinea Coastal Current/ Undercurrent system (NGCC/NGCU) through the
81 basin. This large current system transports water into the Solomon Sea at depths of 0 - 1400 m
82 with the strongest transport (40-80 cm/s) being within the thermocline waters that feed the EUC,
83 at ~ 200 m (Lindstrom et al., 1987; Tsuchiya et al., 1989; Cravatte et al., 2011; Germaineaud et
84 al., 2016; Alberty et al., 2019). The large transport and associated current speeds result in
85 relatively short residence time for waters in the Solomon Sea (e.g., ~ 4 months for thermocline
86 waters). The waters that make up the NGCU originate as the Southern Equatorial Current (SEC),
87 which flows west between ~5°N and 20°S. As the SEC encounters islands in the Coral Sea, it
88 branches into various currents, including the North Vanuatu Jet (NVJ) and the New Caledonia
89 Jet (NCJ, Kessler and Cravatte, 2013; Germaineaud et al., 2016). The NCJ further bifurcates
90 around 18°S into the North Queensland Current (NQC), which flows north through the northern
91 Coral Sea and the Gulf of Papua, and around the southern coast of PNG before joining the NVJ
92 to form the NGCU (Fig. 1; Sokolov and Rintoul, 2000). In the surface ocean (0 – ~150 m), in
93 addition to the NGCU and NVJ, waters from the SEC flow into the Solomon Sea through gaps
94 between the Solomon Islands (Hristova and Kessler, 2012) and through the Solomon Strait

95 (Germineaud et al., 2016; Albery et al., 2019). Most of the surface water exits through the
96 Vitiaz Strait, with some flowing through the St. George's Channel and/or the Solomon Strait
97 (e.g., Albery et al., 2019). The direction of flow through the Solomon Strait varies seasonally,
98 and on a net annual basis flows into the Solomon Sea (Albery et al., 2019). The thermocline
99 waters of the NGCU flow at a depth of ~200 m along the eastern PNG coast and bifurcate south
100 of New Britain where northwestward flow exits through the Vitiaz Strait (Tsuchiya et al., 1989);
101 the remainder flows eastward as the New Britain Coastal Undercurrent (NBCU) (Melet et al.,
102 2010). The NBCU bifurcates around New Ireland, with the western limb flowing out of Saint
103 George's Channel as the Saint George's Undercurrent (SGU), while the eastern limb exits out
104 the Solomon Strait as the New Ireland Coastal Undercurrent (NICU; Butt and Lindstrom, 1994).
105 The NICU then combines with the EUC (Fig. 1; Germineaud et al., 2016). The deeper sub-
106 thermocline (450-1400m, $26.9\sim 27.5\sigma_\theta$) circulation in the Solomon Sea is also dominated by the
107 NGCU, however at much reduced levels of transport. Here, as the NGCU approaches the
108 Woodlark Archipelago it is pushed eastward over this depth range; below 1000 m it passes
109 through, over and around the submarine extension of the Woodlark Archipelago. This deep water
110 exits through the Vitiaz Strait (~65%) and Solomon Strait (~30%; Albery et al., 2019), but with
111 some seasonal variability.

112 The western equatorial Pacific is temporally variable on both seasonal (monsoonal) and
113 interannual (ENSO) scales. This variability leads to fluctuations in transport through the
114 Solomon Sea and the EUC, and is accompanied by changes in river input and surface circulation
115 (Cresswell, 2000; Melet et al., 2013; Delcroix et al., 2014). Seasonal differences in transport and
116 circulation in the Solomon Sea have been studied during the PANDORA (austral winter 2012)
117 and MOORSPICE (austral summer 2014) cruises. Both cruises occurred during a neutral El Niño
118 phase, and demonstrated that the overall flow into and out of the Solomon Sea was enhanced
119 during austral winter (PANDORA) (Germineaud et al., 2016). While overall transport through
120 the Solomon Sea varies seasonally, modeling studies have shown that transport through the
121 Vitiaz Strait exhibits less temporal variability, presumably because transport through the strait is

122 mainly controlled by bathymetry (narrow, 1200 m deep channel) which restricts water flow,
123 while seasonality changes are a result of changes in flow of the NVJ and NQC upstream (Melet
124 et al., 2010). These changes are observed further downstream, in strong temporal variability in
125 the current flow through Saint George's Channel and the Solomon Strait. Historical observations,
126 however, do show variation in the transport through the Vitiaz Strait, with intensification during
127 positive El Nino phase (Lindstrom et al., 1987; Butt and Lindstrom, 1994; Murray et al., 1995).
128 Modeling studies have also observed intensification of the NGCU during El Nino events, which
129 creates eddies, causing increased contact with the PNG shelf (Ryan et al., 2006). While models
130 predict intensification of the NGCU, during the El Nino event of 1991/1992, Murray et al. (1995)
131 observed a weakening of the EUC, where the maximum velocity during neutral El Nino was
132 observed at 90 cm/s, and then dropped to 20 cm/s during peak El Nino. The EUC also varies
133 seasonally, and is stronger in austral winter and weaker in austral summer (Melet et al., 2010).

134

135 The waters that flow through the Solomon Sea are subject to many potential chemical inputs
136 from the margins, rivers, runoff, mine tailings, and volcanic and hydrothermal activities; it is
137 thus thought that the NGCU should undergo significant chemical enrichments as it passes
138 through the Solomon Sea (Lacan and Jeandel, 2001, 2005). Islands in this region are young and
139 easily eroded, leading to large lithogenic inputs via rivers (e.g., Milliman et al., 1999; Sholkovitz
140 et al., 1999). Two very large rivers empty into the ocean along the pathway of the major currents
141 flowing through the region: the Sepik River empties into the Bismarck Sea, thus contributing to
142 the NGCU as it exits the Solomon Sea, and the Fly River empties into the northern Coral Sea
143 where it alters the chemistry of the North Queensland Current (Fig. 1). The Sepik River and Fly
144 River estuaries are two very different systems: The Sepik river system is located over a steep and
145 narrow shelf, such that river sediments are discharged directly into the ocean (Milliman et al.,
146 1999, Sholkovitz et al., 1999; Kineke et al., 2000). Sedimentation occurs on the shelf through the
147 settling of sinking particles from the surface plume, as well as via hyperpycnal flows, which

148 transport sediment to intermediate depths along isopycnals (Kineke et al., 2000). The Fly River
149 is a shallow estuary where sediment deposition and resuspension impacts the chemical makeup
150 of the surface water. The Fly River has a high sediment load (85×10^9 kg/year) relative to its
151 discharge (220 km^3 /year) (Salomons and Eagle, 1990). These sediments are dominated by
152 resuspension due to intense tidal activity, and bioturbated muds in this region have been shown
153 to have elevated Al, Mn, and Fe fluxes (Harris et al., 1993; Alongi et al., 1996).

154 The Fly river, in addition to having a large sediment load due to its strong relief, high rainfall,
155 and easily erodible rock, (e.g., Harris et al., 1993) is impacted at its head waters by increased
156 contamination from the Ok Tedi mine. This ore deposit, and others in the area, are a result of
157 elevated tectonic and volcanic activity in the region. Over the life of the Ok Tedi mine it is
158 estimated that it was responsible for the input of 66 million tons per year of mine tailings,
159 including 24 million tons per year of mill fines (e.g., Hettler et al., 1997). Tailings from mines
160 throughout the region are delivered to the ocean through run off, smaller rivers, and erosion, and
161 a portion of this waste reaches the coastal ocean, where it might be entrained into the NQC.

162 A seafloor spreading-center in the Woodlark Basin in the eastern Solomon Sea hosts
163 hydrothermal vent systems at $>2500\text{m}$ (e.g., Laurila et al., 2012) and undersea-volcanos in the
164 region are hydrothermally (McConachy et al., 2002; Laurila et al., 2012) and volcanically active
165 (McConachy et al., 2002); the eruption of Kavachi volcano in the surface ocean is notable (Baker
166 et al., 2002; Phillips et al., 2016). In the shallow ocean both diffuse and focused venting have
167 been observed at Tutum Bay on the northeastern side of the Solomon Sea, enriching the surface
168 waters in Fe and Mn, though these concentrations only persist near the vent sites (Pichler et al.,
169 1999). In coastal regions, runoff from Rabaul volcano ($4^{\circ}14'25''\text{S}$, $152^{\circ}11'45''\text{E}$, east of New
170 Britain) was measured to be high in Fe (Labatut et al., 2014), and ash from Rabaul's eruptions
171 are deposited onto the ocean surface with precipitation being slightly acidic from the volcanic
172 SO_2 (Ganachaud et al., 2017). The elevated tectonic and magmatic activity in this region
173 suggests that other, undiscovered hydrothermal vent sites likely exist, and this region has been

174 extensively explored for deep-sea mining of hydrothermally-sourced minerals (e.g., Jankowski,
175 2011).

176 ***1.2 Sampling locations***

177 The PANDORA cruise (GEOTRACES GP-12) took place during austral winter 2012 (28 June–6
178 August) aboard the R/V *l'Atalante* and was led by the Laboratoire d'Etudes en Geophysique et
179 Oceanographie Spatiales (LEGOS, Chief Scientist G. Eldin). A more detailed description of
180 hydrographic data and geochemical sampling, as well as preliminary findings about this cruise,
181 can be found in Ganachaud et al. (2017). A total of 170 casts were collected at 83 stations in and
182 around the Solomon Sea. The 12 stations that are examined in this study (Fig. 1a) were sampled
183 using a trace-metal clean rosette according to GEOTRACES protocols. Five of these stations
184 were located outside of the Solomon Sea: three to the south of the entrance (stations 4, 10, and
185 82), one to the northeast near the Solomon Strait (station 43), and one northeast of the Solomon
186 Archipelago (station 13). Within the Solomon Sea, one station was located on the west side of
187 the Solomon Islands, off San Cristobal (station 21). Two stations were located within the flow
188 path of the NGCU near Trobriand Island (stations 39 and 71). In the northern Solomon Sea,
189 stations are located in the Vitiaz Strait (station 77), near the Solomon Strait (station 42), and
190 along the flow path of thermocline waters as they exit via Saint George's Channel (station 60).
191 One station is located at the southern entrance of the Solomon Sea, off the coast of Rossel Island,
192 PNG (station 34) and was only sampled to 350m.

193 **2. Methods**

194 Water sampled for Al and Mn was collected using an epoxy-coated titanium CTD Rosette with
195 twelve 12-L Go-Flo bottles and described in more detail in Ganachaud et al. (2017) . Go-Flo
196 bottles spent minimal time on-deck, and after filling were transported to a clean-air environment
197 for sub-sample removal and storage between casts. Total acid-soluble metal samples, which

198 includes the dissolved fraction plus the acid soluble portion of particles present in unfiltered
199 samples (abbreviated as TD – total dissolvable, based on the terminology of Mackey et al. 2002a
200 and Slemons et al. 2010) were collected directly from Go-Flo bottles into acid-cleaned 100 mL
201 LDPE bottles (with LDPE caps). Dissolved metal (D) samples were collected from the Go-Flo
202 bottles using slight overpressure of filtered N₂ for filtration through 0.2 µm Sartobran-300
203 capsule filters into acid-cleaned 100 mL LDPE bottles (with LDPE caps). All samples were
204 acidified to 0.024 N using Optima HCl one month prior to Al analysis, and six months prior to
205 Mn analysis.

206 Al was analyzed by flow injection analysis (FIA) using direct injection and fluorescent detection
207 of the Al-lumogallion complex, following Resing and Measures (1994). This method had an
208 average detection limit of 0.85 nM, which is 34% of the lowest concentration measured, and 7%
209 of the average concentration of all samples. GEOTRACES GD (consensus ± 1 SD = 17.7 ± 0.2
210 nM) reference standard was run daily, with an average value of 19.49 ± 0.88 nM (1 SD, n = 17).
211 We acknowledge that this value is higher than the reported consensus value, but we are not able
212 to identify any blanks in our system based on the methodology that was used here. We note that
213 the reference material is fairly old (GEOTRACES GD was collected 11 years prior to these
214 analyses), and has been stored with HDPE caps, which are known to cause contamination for Al
215 (Brown and Bruland, 2008). Other analysts have found consistently elevated Al values for the
216 low nM level GEOTRACES standards (e.g. Resing et al., 2015; Singh et al., 2020). Daily
217 precision for standards was on average 3.4% relative standard deviation (RSD) at 1 nM and 2.2%
218 RSD at 20 nM.

219 Mn was analyzed by FIA using in-line preconcentration of Mn onto an 8-hydroxyquinoline
220 column and spectroscopic detection of leucomalachite green, based on the method of Resing
221 and Mottl (1992), with the addition of 4 g of nitrilo tri-acetic acid per liter to the ammonium
222 acetate reaction buffer. This method had a detection limit of 0.027 nM, which is 14% of the
223 lowest sample measured and 3% of the average concentration of all samples measured.

224 GEOTRACES GD (consensus \pm 1 SD = 1.50 ± 0.11 nM) standards were run at least once a day
225 and measured to be 1.80 ± 0.19 (1 SD, n = 17). Daily precision for standards was on average
226 3.0% RSD at 0.1 nM and 2.0% RSD at 1.0 nM. An internal consistency standard was run at least
227 twice daily and found to be $0.32 \text{ nM} \pm 0.02 \text{ nM}$ (1 SD, n = 94).

228 *2.1 Calculation of a trace metal budget in the Solomon Sea*

229 To examine the net impact on water passing through the Solomon Sea, we construct a budget to
230 evaluate the fluxes into and out of the Solomon Sea at three density intervals, based on transport
231 estimates calculated for the PANDORA cruise by Germaineaud et al. (2016): Surface layer (surf –
232 $24\sigma_\theta$), Thermocline layer or NGCU ($24\text{--}26.9\sigma_\theta$), and Deep layer ($>26.9\sigma_\theta$; our maximum
233 sampled σ_θ is 27.54/1300 m). For the thermocline layer ($24\sigma_\theta\text{--}26.9\sigma_\theta$), we can also compare
234 these fluxes to the flux of the trace metals at the Equatorial Undercurrent to evaluate the
235 contribution of the Solomon Sea trace metal pool to that of the EUC (section 4.2).

236 The depth-weighted average concentration over each density interval is calculated using a
237 trapezoidal integration over the depth range corresponding to potential density for each station
238 and is reported in Table 1. The corresponding depths for each density interval can be found in
239 Table S1 and the full dataset can be found in Table S2. Errors reported in the text and in the
240 tables represent one standard deviation calculated using the analytical error on each
241 measurement.

242 This budget considers that there is a background flux of Al and Mn entering the Solomon Sea
243 over each of the three potential density ranges, referred to here as the Solomon Sea Inflow. The
244 outflow budget considers outputs through the Vitiaz Strait (station 77), St. George's Channel
245 (station 60) and the Solomon Strait (station 42), which we refer to collectively as the Solomon
246 Sea Outflow. Germaineaud et al. (2016) reported outflow values for the Vitiaz Strait and the sum

247 of the outflow for the water leaving via the Solomon Strait and St. George's Channel. Because
248 they do not report individual outflows for these two straits, we assume their outflows to be equal.
249 While Alberty et al. (2019) estimate transport through each channel, it is on an annual basis and
250 not over shorter time intervals. This is important because the flows are seasonally variable with
251 water flowing both into and out of the Solomon Strait based on season. During PANDORA,
252 surface water was flowing out of the Solomon Sea via Solomon Strait, rather than into it, as it
253 does on a net annual basis (Alberty et al., 2019). For these reasons, we find that the estimates
254 provided by Germaineaud et al. (2016) are the best to use for this budget.

255 Chemical influx and efflux are calculated by considering transport and the concentrations of each
256 DAl, DMn, TDAI, and TDMn at appropriate stations as follows:

257
$$\text{Flux (moles/s)} = \text{volume transport (m}^3/\text{s)} \times [\text{element}] \text{ (moles/m}^3\text{)}$$

258 For the efflux, transport out of the Solomon Sea is considered for each of the exit straits
259 combined with the depth-weighted average concentration at stations closest to them.

260 The specific assumptions and processes used to determine average concentration for each flux
261 are described below.

262 2.1.1 Inflow

263 We assume the Al and Mn concentrations at stations 10 and 82 to be representative of those
264 flowing into the Solomon Sea. Station 10 is located where water flows across the Vanuatu
265 Archipelago via the North Vanuatu Jet, while station 82 is located in the middle of the Coral Sea,
266 farther away from potential shelf inputs, and likely is dominated by water flowing via the NQC.
267 Station 34, which is the station likely to be most representative of the NGCU inflow, was not

268 sampled through the potential density range of the EUC due to bad weather, and also shows
269 elevated Al relative to other profiles within the Solomon Sea, suggesting shelf input, and we
270 have therefore chosen not to consider it as a background profile. While a full profile of trace
271 metal samples was not collected for station 34, a standard CTD-rosette package collected a full
272 depth CTD profile, which shows that temperature, salinity, oxygen, and potential density profiles
273 at this station appear to be intermediate between stations 10 and 82. Station 82 generally has
274 lower concentrations of Mn and Al (by ~0.2 nM Mn and up to 5 nM Al) than Station 10. To
275 calculate influx, the average Al and Mn concentrations over each density interval for station 82
276 and 10 were calculated and multiplied by transport rate. Error was determined using the
277 combined analytical error (1SD) of the measurements used to calculate the average.

278 2.1.2 Outflow

279 The depth-weighted average concentrations of DAl, TDAI, DMn, and TDMn over each density
280 interval for each strait/station were multiplied by water transport rate (Eq. 1). The chemical
281 distributions at station 77 (Vitiav Strait), station 60 (St. George's Channel), and station 42
282 (Solomon Strait) are assumed to be representative of the water leaving the Solomon Sea. Station
283 60 is not located within St. George's Channel; however we assume that it is representative of
284 water leaving the Solomon Sea via this channel for two reasons. First, the temperature-salinity
285 profile of other stations sampled using the standard CTD rosette in Saint George's Channel are
286 similar to Station 60 (Germineaud et al., 2016). Second, ADCP data at station 60 show that
287 currents flow toward both Saint George's Channel and the Solomon Strait, meaning that at least
288 some portion of the water that flows across station 60 must ultimately be directed out of the
289 Solomon Sea (Germineaud et al., 2016).

290 **3. Results**

291 DAI, TDAI, DMn, and TDMn in and around the Solomon Sea from the 2012 PANDORA cruise
 292 are shown versus potential density in Figures 2-5, and versus depth (supplemental Fig. S2-S5).
 293 Here we define the surface layer as being from the surface ($\sim 21\sigma_\theta$) to $24\sigma_\theta$ (0- ~ 150 m), the
 294 thermocline layer from ~ 150 - ~ 450 m (24 - $26.9\sigma_\theta$), and deep waters at depths $>\sim 450$ m- 1300 m
 295 ($>26.9\sigma_\theta$). Each plot includes a profile of the average of stations 10 and 82, which are
 296 considered to be representative of water entering the Solomon Sea at its southern entrance.
 297 However, we note that these profiles are not representative of all surface water entering the
 298 Solomon Sea, as there is additional inflow through the Solomon Islands or through the Solomon
 299 Strait (e.g., Hristova and Kessler, 2012; Albery et al., 2019).

Table 1a. Depth-weighted average DAI and TDAI concentrations. Values are calculated by integrating concentration data over the depth interval corresponding to the potential density range of each station. Reported error represents one standard deviation calculated from the analytical error of each measurement.

		Surface layer Surf – $24\sigma_\theta$		Thermocline layer /NGCU 24 – $26.9\sigma_\theta$		Deep layer $<26.9\sigma_\theta$	
	Station	DAI (nM)	TDAI (nM)	DAI (nM)	TDAI (nM)	DAI (nM)	TDAI (nM)
Inflow	82	12.7 \pm 0.6	12.7 \pm 0.6	6.2 \pm 0.8	7.8 \pm 0.8	6.8 \pm 0.8	7.3 \pm 0.8
	10	10.4 \pm 0.7	10.1 \pm 0.8	5.8 \pm 0.8	5.1 \pm 0.8	8.0 \pm 0.9	7.5 \pm 0.9
	Average	11.4\pm1.3	11.2\pm1.5	6.0\pm0.8	6.5\pm1.5	7.5\pm0.7	7.4\pm0.7
NGCU Flow path	34	12.5 \pm 0.6	12.5 \pm 0.6	10.5 \pm 0.6	11.5 \pm 0.5		
	39	13.3 \pm 0.5	13.2 \pm 0.5	6.5 \pm 0.8	6.4 \pm 0.8	8.3 \pm 0.7	7.9 \pm 0.7
	71	12.5 \pm 0.8	12.6 \pm 0.8	6.0 \pm 1.3	8.0 \pm 1.3	6.0 \pm 1.4	9.6 \pm 1.4
Outflow/ Straits	77	15.5 \pm 0.7	19.1 \pm 0.8	9.0 \pm 0.7	10.8 \pm 0.5	8.1 \pm 0.8	11.5 \pm 0.8
	42	11.0 \pm 0.7	13.0 \pm 0.7	8.9 \pm 0.7	9.5 \pm 0.7	9.3 \pm 0.7	12.9 \pm 0.8
	60	12.3 \pm 0.7	12.3 \pm 0.7	6.6 \pm 0.7	7.5 \pm 0.7	6.1 \pm 0.7	9.3 \pm 0.7
Outside of Solomon Sea	43	6.3 \pm 0.5	5.7 \pm 0.7	6.0 \pm 1.0	4.7 \pm 1.1	5.9 \pm 1.0	6.0 \pm 0.7
	13	9.6 \pm 1.0	9.2 \pm 0.9	5.6 \pm 1.2	5.6 \pm 1.2	8.6 \pm 1.1	7.6 \pm 1.1
	21	14.4 \pm 0.7	23.5 \pm 0.8	9.4 \pm 0.8	15.0 \pm 1.1	10.2 \pm 0.7	15.3 \pm 1.1
	4	11.9 \pm 0.5	11.5 \pm 0.5	10.2 \pm 0.8	10.1 \pm 0.8	7.7 \pm 0.8	8.8 \pm 0.9

300

301

Table 1b. Depth-weighted average DMn and TDMn concentrations. Values are calculated by integrating concentration data over the depth interval corresponding to the potential density range of each station. Reported error represents one standard deviation calculated from the analytical error of each measurement.

	Station	Surface layer Surf – 24 σ_θ		Thermocline layer/NGCU 24 – 26.9 σ_θ		Deep layer <26.9 σ_θ	
		DMn (nM)	TDMn (nM)	DMn (nM)	TDMn (nM)	DMn (nM)	TDMn (nM)
Inflow	82	1.52±0.11	1.61±0.11	0.27±0.02	0.41±0.03	0.23±0.02	0.41±0.03
	10	1.21±0.11	1.30±0.09	0.27±0.02	0.39±0.02	0.37±0.02	0.56±0.05
	Average	1.32±0.22	1.43±0.19	0.27±0.02	0.40±0.02	0.30±0.07	0.48±0.07
NGCU Flow path	34	0.59±0.12	0.84±0.06	0.30±0.04	0.48±0.04	0.27±0.04	0.53±0.05
	39	1.32±0.09	1.51±0.10	0.25±0.05	0.43±0.05	0.27±0.04	0.53±0.05
	71	1.05±0.06	1.15±0.06	0.24±0.03	0.40±0.04	0.34±0.04	0.63±0.06
Outflow/ Straits	77	1.22±0.10	1.45±0.12	0.26±0.01	0.49±0.02	0.51±0.04	0.88±0.07
	42	0.77±0.06	1.07±0.09	0.27±0.05	0.49±0.05	0.46±0.06	0.81±0.06
	60	1.26±0.10	1.35±0.11	0.28±0.02	0.44±0.03	0.43±0.03	0.74±0.06
Outside of Solomon Sea	43	0.72±0.05	0.79±0.05	0.32±0.14	0.42±0.04	0.33±0.03	0.50±0.05
	13	0.70±0.05	0.80±0.06	0.29±0.10	0.40±0.09	0.39±0.08	0.60±0.05
	21	0.93±0.07	1.38±0.11	0.31±0.09	0.56±0.04	0.38±0.08	0.68±0.05
	4	1.54±0.09	1.58±0.09	0.38±0.02	0.54±0.05	0.36±0.02	0.63±0.06

302

303 In the surface layer, average DA1 and TDA1 inflowing water are 11.4 ± 1.4 nM and 11.2 ± 1.3
304 nM, respectively. Within the Solomon Sea (stations 34, 39, 71) and in St. George’s Channel
305 (station 60), average DA1 and TDA1 in the surface layer are ~1-2 nM greater than in the
306 inflowing water. The Solomon Strait (station 42) shows a decrease in DA1 (~0.5 nM), but an
307 enrichment of ~1nM in TDA1. By contrast the Vitiav Strait (station 77), shows a much greater
308 enrichment in both DA1 and TDA1 (~4 nM and ~8 nM, respectively). Outside of the Solomon
309 Sea to the north and northeast (stations 13 and 43; Figs. 2,3 j,k) DA1 and TDA1 are ~2-5 nM
310 lower than the waters entering the Solomon Sea. Station 21, which is just outside of the eastern
311 opening of the Solomon Sea and south of the Solomon Islands, is fed by waters from north of the
312 Solomon Islands (e.g., from station 13) that flows through gaps in the islands, resulting in
313 average concentrations at station 21 being enriched by 5 nM for DA1 and 13 nM for TDA1
314 relative to their concentrations at station 13.

315 In the thermocline layer which, within the Solomon Sea, is largely made up of the NGCU (~150-
316 450 m; 24-26.9 σ_θ), DAI and TDAI generally decrease from surface concentrations to lower
317 values, often reaching mid-depth minima at different depths for different stations (Fig. 2 and Fig.
318 S2). Over this density range, the inflowing waters have depth-weighted average DAI and TDAI
319 concentrations of 6.0 ± 0.8 nM, and 6.5 ± 1.5 nM, respectively. Along the flow path of the
320 NGCU (stations 39 and 71), concentrations remain roughly constant relative to inflow, with
321 slight enrichments of DAI (~0.5 nM) at station 39 and TDAI (~1.6 nM) at station 71. In the exit
322 straits (stations 42, 60, and 77) DAI and TDAI are enriched relative to the inflow by ~0.65 nM –
323 4 nM. To the north and northeast of the Solomon Sea (station 13 and 43) DAI is similar to the
324 inflowing water, while TDAI is less at station 13 (~1 nM) and station 43 (~2 nM; however, it
325 should be noted that DAI at station 43 exceeds TDAI, suggesting that these samples may be
326 slightly contaminated). Stations 4 and 21, which are close to local bathymetry (e.g., sills, straits)
327 are enriched in DAI at both stations by ~4 nM, and in TDAI by ~4 nM and ~8 nM at station 4
328 and 21, respectively.

329 In the deep layer (>450 m; >26.9 σ_θ) the depth-weighted average inflowing DAI is 7.5 ± 0.7 nM,
330 and TDAI is 7.4 ± 0.7 nM. Within the basin (stations 39 and 71), in the Vitiaz Strait (station 77),
331 as well as outside the Solomon Sea (station 13), depth-weighted average DAI concentrations are
332 within the observed variability of the inflowing water. There is slight enrichment in DAI (~2-3
333 nM) in the Solomon Strait (station 42) and near the Solomon Islands (station 21). There are
334 small (2-3 nM) enrichments of TDAI within the basin at station 71 and in St. George's channel at
335 station 60, and larger enrichments relative to the inflow (4-5 nM) in Vitiaz (station 77) and
336 Solomon (station 42) Straits, as well ~8 nM enrichment at station 21 on the eastern edge of the
337 basin, again reflecting the proximity to local bathymetry. Northeast of the Solomon Sea (station
338 43), both DAI and TDAI are depleted by ~1 nM, relative to the waters flowing into the Solomon
339 Sea.

340 Mn has a scavenged distribution at all stations, with the highest concentrations at the surface and
341 generally decreasing with depth. From the surface to $24\sigma_\theta$, the depth-weighted average
342 concentration of inflowing waters for DMn is 1.32 ± 0.22 nM, while TDMn is 1.43 ± 0.19 nM.
343 Within the Solomon Sea, DMn and TDMn at most stations do not vary from these inflow values
344 (stations 39, 60, 77), or are depleted relative to the inflowing waters by $\sim 0.3 - 0.7$ nM along the
345 flow path of the NGCU (stations 34, 42). Station 21 is also depleted in DMn by ~ 0.4 nM relative
346 to inflowing waters. Outside the Solomon Sea (stations 13, 43), DMn and TDMn are depleted by
347 ~ 0.6 nM relative to inflowing water, while south of the Solomon Sea at station 4, DMn and
348 TDMn concentrations are ~ 0.2 nM higher than the waters flowing into the Solomon Sea.

349 Over $24-26.9 \sigma_\theta$, the depth-weighted average DMn of inflowing water is 0.27 ± 0.02 nM and
350 TDMn is 0.40 ± 0.02 nM. In this density interval, most stations have DMn and TDMn
351 concentrations that are within the variability of the inflow. There are a few stations where there
352 is enrichment (~ 0.1 nM) of TDMn – along the flow path of the NGCU at station 34, in the Vitiaz
353 (station 77) and Solomon (station 42) Straits. TDMn is enriched ~ 0.2 nM near the Solomon
354 Islands (station 21). South of the Solomon Sea (station 4) DMn and TDMn are higher than those
355 waters entering the Solomon Sea by ~ 0.1 nM and ~ 0.2 nM respectively.

356 In the deeper ocean ($>450\text{m}$; $>26.9 \sigma_\theta$), on average, DMn and TDMn generally increase with
357 depth at all stations. Below $26.9 \sigma_\theta$, the inflowing waters have a depth-weighted average DMn
358 concentration of 0.30 ± 0.07 nM and TDMn concentration of 0.48 ± 0.07 nM. In the straits,
359 (stations 77, 60, and 42), enrichments in DMn and TDMn relative to the inflow are on the order
360 of $0.1-0.2$ nM and $0.3-0.4$ nM, respectively. Within the basin (stations 34, 39, 71) the average
361 concentrations of DMn are generally similar to those in the inflow, while smaller enrichments
362 (~ 0.1 nM) of TDMn are found at stations 4, 10, 13, 71, and of 0.2 nM at station 21.

363 **4. Discussion**

364 **4.1 Al and Mn enrichments within the Solomon Sea**

365 4.1.1 Surface Layer (surf - $24\sigma_\theta$)

366 Over this potential density interval, average Al and Mn concentrations are lowest in the surface
367 ocean at the two stations (13 and 43) located in the westward flowing South Equatorial Current
368 (SEC) north and east of the Solomon Islands just outside of the Solomon Sea, relative to
369 concentrations elsewhere in the Solomon Sea basin. This must reflect their westward transport
370 from the open ocean. This is consistent with the Al values observed at the eastern entrance of the
371 Solomon Sea (station 10; Table 1a, Table S2). Within the Solomon Sea and the straits, average
372 concentrations over this density interval tends to be higher.

373 The elevated surface concentrations (Table 1) that are observed in Mn and Al within the
374 Solomon Sea could be derived from riverine, aeolian, or coastal/margin sources. Based on
375 salinity, our data reveal no large inputs of fresh water to the Solomon Sea during this cruise and
376 consequently no correlation between salinity and trace metals is observed. This means either that
377 at the time of the cruise, trace metal inputs to the Solomon Sea via rivers were small or that trace
378 metal inputs were large relative to freshwater input, which is more consistent with observations
379 of high sediment loads, relative to fluvial discharge in this region (e.g., Milliman et al., 1999).
380 Dust input to the Solomon Sea can be predicted using the MADCOW model (Measures and
381 Brown, 1996) based on regional dust fluxes and empirical dust solubilities. Using a dust flux of
382 $1.0 \text{ g m}^{-2} \text{ y}^{-1}$ to the Solomon Sea (Shank and Johansen, 2008), an empirical dust solubility of
383 6.0% (Buck et. al., 2006), a residence time of surface water in the Solomon Sea of 0.6-2 months
384 (estimated based on an average velocity of $20\text{-}60 \text{ cm s}^{-1}$ over the $\sim 1000 \text{ km}$ distance between the
385 inflow and outflow, and consistent with literature estimates (Melet et al., 2011, Hristova and

386 Kessler, 2012; Alberty et al., 2019)) and a mixed layer depth <100 m (based the definition of de
387 Boyer Montégut et al., 2004), it can be estimated that DA_I added from dust deposition accounts
388 for only 0.30 nM to 1.14 nM of the DA_I added to the surface waters as they transit through the
389 Solomon Sea which is only $\sim 5.4\% \pm 2.5\%$ (1SD) of DA_I present there. Thus, dust is a relatively
390 small source of DA_I to the waters in this region. A similar estimation for DMn can be made
391 using a fractional solubility of Mn in dust of 45.1% from Buck et al. (2013), and a crustal
392 abundance of Mn of 954 ppm from Taylor, (1964) to be consistent with the reference used by
393 Measures and Brown (1996). We estimate that DMn added from dust deposition as the waters
394 transit through the Solomon Sea to be only 0.01 nM to 0.05 nM or $\sim 1.9\% \pm 0.9\%$ (1SD) of DMn
395 present in the surface Solomon Sea.

396 Elevated Al and Mn concentrations in the surface ocean within the straits and Solomon Sea,
397 relative to inflowing waters or stations outside the Solomon Sea, suggest that these elements
398 become enriched when waters interact with local bathymetry as they enter the basin. The SEC
399 transports surface waters into the Solomon Sea passing through both the Solomon Strait and
400 other gaps in the Solomon archipelago ultimately exiting through the Vitiaz Strait (Cravatte et al.,
401 2011; Hristova and Kessler, 2012). TDA_I of the surface samples in both the Solomon Strait
402 (station 42 – 18.7 ± 0.4 nM) and Vitiaz Strait (station 77 – 38.7 ± 0.4 nM) are enriched compared to
403 DA_I (13.5 ± 0.3 nM, 28.7 ± 0.4 nM respectively) suggesting the input of sediments or rapid removal
404 of DA_I by scavenging. A similar but muted effect is seen for Mn at these stations. The high Mn
405 (2.5 nM) and Al (>20 nM) in the surface waters that exit through the Vitiaz Strait (station 77)
406 must result from scouring and entrainment of local sediments into the water column due to rapid
407 geostrophic current velocities (20-60 cm/s; e.g. Hristova and Kessler 2012) coupled with the
408 narrow strait. A maximum in Rare Earth Elements (REE), including dissolved Cerium (Ce), was
409 also observed in the surface waters at station 77 during the PANDORA cruise (Pham et al.,
410 2019). At the eastern end of the Solomon archipelago, where waters flow from the SEC (station
411 13) through the islands to station 21, Mn and Al concentrations are also greater than those in the
412 SEC (see results) reflecting input from the islands and the scouring of coastal sediments. This is

413 also largely consistent with enrichments in REEs observed during PANDORA (Pham et al.,
414 2019). The surface water flowing through the Solomon Sea has a residence time of ~ 53-103
415 days (Melet et al., 2011), so input and removal processes for Mn and Al must be both strong and
416 rapid to produce the changes in concentrations observed here.

417 Using transport estimates reported in Germaineaud et al. (2016) and the average concentrations of
418 metals at stations in/near the straits, we can estimate the fluxes entering the Solomon Sea from
419 the southern entrance and exiting via the three straits. However, transport estimates
420 (Germaineaud et al., 2016) are unbalanced with a transport entering the Solomon Sea via the
421 southern entrance of 3.3 ± 0.4 Sv, and exiting the Solomon Sea of 6.9 ± 0.6 Sv (4.5 ± 0.4 Sv
422 through the Vitiaz Strait and a combined 2.4 ± 0.3 Sv through St. George's Channel, and the
423 Solomon Strait). The missing flow likely enters through the gaps in the Solomon Islands. As a
424 result, in the surface layer, our budget is unbalanced, and the larger exit fluxes reflect the
425 differences in flow. Table 2 reports the flux in through the southern entrance and out through the
426 straits. While we cannot calculate a balanced flux value for the inflow, the depth-weighted
427 average concentrations at the exit straits all suggest that water leaving the Solomon Sea is within
428 the variability of, or enriched in, Al and Mn relative to the inflowing surface layer (Table 1).

429 Trace metal fluxes based on transport in the surface layer are reported in Table 2. This budget
430 does not consider diapycnal/vertical mixing between layers, because vertical mixing does not
431 appear to be reflected in the chemical profiles that we see in Al and Mn in the upper NGCU.
432 While diapycnal mixing in the Solomon Sea is important (Melet et al., 2011), based on the
433 erosion of the salinity maximum in the thermocline between the entrance and the exit of the
434 Solomon Sea and the salinity of the surface layer, we calculate that no more than 33% of water is
435 mixed between our density intervals. Additionally, several water masses converge upon entrance
436 into the Solomon Sea (e.g., Kessler et al., 2019) and there is evidence of vertical mixing due to
437 internal tides in the region (e.g., Melet et al., 2011), making diapycnal mixing in the surface layer
438 hard to resolve.

Table 2: Transport and trace metal flux estimates for the surface (surf-24) in the Solomon Sea. Transport estimates from Germaineud et al. (2016). Solomon Strait and St. George’s Channel transport are estimated based on total transit through the two outflow straits. Errors represent the combined error from the transport estimate and the analytical error (1 SD) of the average concentration.

Station	Transport (Sv)	DAI (mol/s)	TDAI (mol/s)	DMn (mol/s)	TDMn (mol/s)
Inflow	3.3 ± 0.4	38 ± 6	37 ± 7	4.4 ± 0.9	4.71 ± 0.8
Vitiaz Strait (station 77)	4.5 ± 0.1	70 ± 4	86 ± 4	5.5 ± 0.5	6.5 ± 0.5
Solomon Strait (station 42)	1.2 ± 0.3	13 ± 3	15.6 ± 4.0	0.9 ± 0.2	1.3 ± 0.3
St. George’s Channel (station 60)	1.2 ± 0.3	15 ± 4	14.7 ± 3.8	1.5 ± 0.4	1.6 ± 0.4
Total Outflow	6.9 ± 0.3	98 ± 11	116 ± 12	7.9 ± 1.1	9.4 ± 1.3

440

441

442 4.1.2 Thermocline layer – ~150-450 m and 24-26.9 σ_θ

443 The water between ~150-450 m and 24-26.9 σ_θ is dominated by the core of NGCU, which
 444 originates from the south and east of the Solomon Sea (see section 1.1). At this depth/density
 445 range, dissolved and total dissolvable Al and Mn are lower than in the surface ocean (see Table 1
 446 and results), however there are enrichments in DAI, TDAI, and TDMn (but not DMn) relative to
 447 the water flowing into the Solomon Sea (Sta. 10 and 82). The pattern of enrichments is similar,
 448 in part, to those in the surface layer with enrichments in DAI, TDAI, and TDMn in the Solomon
 449 (station 42) and Vitiaz (station 77) straits. At station 21 where water flows roughly from station
 450 13 to 21 through the gaps in the Solomon Islands (Hristova and Kessler, 2012) and near a
 451 shallow sill, enrichments in DAI, TDAI, DMn, and TDMn are observed (Figures 2-5, k and l).

452 Sampling during PANDORA included measurements of Rare Earth Elements, and our data is
453 consistent with increases in dissolved Lanthanum (La), Neodymium (Nd), Europium (Eu), and
454 Ytterbium (Yb) between stations 13 and 21, reported by Pham et al. (2019). Dissolved Ce and
455 the Ce anomaly are not modified between stations (Pham et al., 2019), which suggests that the
456 water mass is the same between the two stations, but that particle inputs are recent and
457 modifications do not cause changes in dissolved Ce, because it is insoluble. The best explanation
458 for these observations is sediment resuspension which is supported by larger enrichments (~9
459 nM TDAI, 0.2 nM TDMn) in total dissolvable metals versus dissolved metals (~4 nM DAi, ~0
460 nM DMn). Coincident enrichment of Mn and Al also occurs in the Vitiaz Strait (station 77),
461 similarly suggesting sediment resuspension, rather than reductive release from shelf sediments or
462 hydrothermal input at these stations, as these other two processes would impact Mn distributions
463 to a greater extent than Al distributions. This is also consistent with Ce anomaly data, which
464 shows decreasing solubility, rather than remineralization in the Vitiaz Strait (Pham et al., 2019).
465 Physical resuspension/non-reductive dissolution is implicated for the enrichment of Fe in the
466 Vitiaz Strait and the NICU based on Fe isotope measurements (Labatut et al., 2014). These Fe
467 enrichments reflect particulate iron transported from the continent across shelves and slopes,
468 followed by release of DFe from suspended particles (Labatut et al., 2014). This has also been
469 observed for Al and Mn in other regions. Al remobilized from benthic nepheloid layers is an
470 important source for Al in the deep waters of the North Pacific (Moran and Moore, 1991).
471 Lateral transport of Mn remobilized from shelves has been observed off the coast of California
472 (Martin et al., 1987) who show that major sub-surface maxima and minima in the Mn
473 distributions are controlled primarily by sedimentary release combined with water mass
474 movement and physical mixing, rather than through scavenging and remineralization processes.

475 The NGCU is expected to flow from station 34 to 39 and then to 71 after which it bifurcates,
476 going WNW to the Vitiaz strait (station 77) and NE to station 60 and then on to the Solomon
477 strait (station 42) (Fig. 1). Two of the stations along the predicted flow path of the NGCU (34
478 and 71) show enrichments in DAi, TDAi, and TDMn that exceed the variability of the average

479 inflow profile, while station 39 shows no enrichments in either Al or Mn (see Table 1 and
480 results). At station 71, the coherence between DAl and TDAI provides confidence in sample
481 quality and thus the lack of a mid-depth enrichment at station 39 indicates that this station is
482 either not in the NGCU, or that local trace metal input and removal are highly dynamic and
483 variable. Based on transport and current speeds, the NGCU/NBCU is estimated to take 2-4
484 months to transit through the Solomon Sea, which is also consistent with model estimates of
485 residence times (Melet et al., 2011). This is much shorter than reported residence times of Al in
486 the open ocean. Because this region has large coastal influence, it is possible that input and
487 removal of Al are rapid, and riverine inputs and coastal sediments must release and scavenge Al
488 on shorter timescales than the residence time of NGCU waters in the Solomon Sea. If station 34
489 is not in the flow path of the NGCU, then the enrichments seen at station 39 are local. However,
490 if station 39 is not in the flow path of the NGCU, then the enrichments at station 71 might be
491 sourced from the upstream waters coming from station 34. In either case, the small increases in
492 TDAI without enrichments in DAl and Mn at station 60 may be sourced from station 71.

493 In the thermocline layer, observations show that water transport into and out of the Solomon Sea
494 was roughly equal at the time of the PANDORA cruise with transport into the Solomon Sea at its
495 southern entrance $\approx 23.6 \pm 0.8$ Sv, and transport out of the Solomon Sea through the exit straits
496 $\approx 22.8 \pm 1.0$ Sv (Germineaud et al., 2016). This allows us to construct a steady state budget with
497 regard to transport over this density interval. Despite strong seasonal and interannual variability
498 in flow through the Solomon Sea, (e.g., Germaineaud et al., 2016; Albery et al., 2019), estimates
499 of the transport of the EUC (e.g., Grenier et al., 2011; Lindstrom et al., 1987; Tsuchiya et al.,
500 1989) are similar to the Solomon Sea transport over this density interval (22-23 Sv) observed by
501 Germaineaud et al. (2016) during PANDORA. In our steady state model, chemical influx is thus
502 calculated by multiplying the depth-weighted average inflow profile for the Solomon Sea by
503 22.8 ± 1.0 Sv. We also consider that DAl, TDAI, DMn, and TDMn distributions in the Vitiaz
504 Strait (station 77), St. George's Channel (station 60), and Solomon Strait (station 42) are

505 representative of the water leaving the Solomon Sea via the NGCU, SGU, and NICU,
 506 respectively.

507 The inflow, outflow, and EUC trace metal fluxes are reported in Table 3.

Table 3. Transport and trace metal flux estimates for the thermocline layer (24-26.9 σ_θ) in the Solomon Sea. Transport estimates are from Germaineaud et al. (2016). Solomon Strait and St. George's Channel transport are estimated based on total transit through the two outflow straits. Errors represent the combined uncertainty from the transport estimate and the analytical error (1 SD) of the average concentration.

	Transport (Sv)	Dal mols/s	TDAI mols/s	DMn mols/s	TDMn mols/s
Inflow* (mol s ⁻¹)	22.8 ± 1.0	136 ± 19	147 ± 35	6.1 ± 0.5	9.1 ± 0.6
Vitiaz Strait (station 77)	10.3 ± 0.3	93 ± 8	111 ± 5	2.7 ± 0.1	5.0 ± 0.3
Solomon Strait (station 42)	6.25 ± 0.4	56 ± 6	59 ± 6	1.7 ± 0.3	3.1 ± 0.4
St. George's Channel (station 60)	6.25 ± 0.5	42 ± 5	47 ± 4	1.7 ± 0.2	2.7 ± 0.3
Outflow (mol s ⁻¹)	22.8 ± 1.0	190 ± 19	217 ± 16	6.1 ± 0.6	10.8 ± 1.0

508

509 These data show that at the time of the PANDORA cruise, fluxes of DAL, TDAI, and TDMn out
 510 of the Solomon Sea in the thermocline waters were only slightly higher (8-90 mol/s DAL, 7-130
 511 mol/s TDAI, 0.3-2.4 mol/s TDMn) than fluxes into the Solomon Sea, and that DMn fluxes
 512 remained constant through the basin.

513 4.1.3 Deep layer – 450-1400 m; $>26.9 \sigma_\theta$

514 In the deeper ocean (450m – 1300m; $26.9 \sigma_\theta - \sim 27.5 \sigma_\theta$) Al and Mn generally show enrichment
515 within the straits or near local bathymetry, which is similar to what was observed in the
516 thermocline layer, with some exceptions. The enrichments in TDAI and TDMn relative to
517 background concentrations that are observed in the Vitiaz and Solomon Straits (stations 77 and
518 42) reflect the proximity of these stations to local bathymetry, combined with flow through these
519 straits by the NGCU (station 77) and NICU (station 42). Similarly at station 21, increases in
520 DAI, TDAI, and TDMn relative to concentrations at station 13 reflect enrichments due to
521 scouring as the SEC flows through the Solomon Islands.

522 Other factors, such as accumulation of sinking particles may be responsible for the enrichments
523 in Mn and Al at stations 39, 71, and 60, which lie deep within the Solomon Sea and are isolated
524 from local features. The Woodlark Rise, extending 250 km NE from the Woodlark Island
525 Archipelago to within ~ 60 km from the Solomon Archipelago, is a ridge line of bathymetric
526 highs that reach ≤ 1000 m with gaps between highs reaching ~ 1500 m. This feature creates a semi-
527 enclosed basin below ~ 1000 m that encompasses stations 71 and 60 in the NE Solomon Sea.
528 Below 400m, the NGCU (Alberty et al., 2019) transits from the Coral Sea and flows around the
529 Louisiade Archipelago and then around, through, and/or over the Woodlark Rise; below 1000 m
530 it likely passes through the channel between the Woodlark Rise and the Solomon Islands (Fig.
531 1b). At stations 42 and 60, enrichments in TDAI and TDMn, the HCl-labile portion of the
532 particulate fraction, likely represent sinking particles accumulating along the flow path of the
533 NGCU/NBCU and within the semi-enclosed basin due to longer residence time for waters there.
534 The residence time of water in the basin between ~ 450 and ~ 1400 m is estimated to be <1 year
535 based on models (Melet et al., 2011) or transport estimates, however transport (and current
536 speed) significantly decreases with depth below 450m (Alberty et al., 2019; Gasparin et al.,
537 2012) resulting in increased residence times with depth.

538 The relative increases in TDAI and TDMn also arise from the scavenging of dissolved phases
539 sourced along the flow path or from particles resuspended as water flows over the Woodlark rise
540 and other bathymetry. Through the Vitiaz Strait (station 77), there is likely sediment
541 resuspension from the bottom and walls of the channel ($\sim 1100 \text{ m}/27.4 \sigma_\theta$) as water passes
542 through. The larger enrichments in the exit straits may also be due to accumulation of sinking
543 particles, or other local, shorter-lived phenomena (e.g., increased river runoff), though we do not
544 have enough information to say this definitively. It is possible that at these deeper depths,
545 especially through the Solomon Strait (station 42), decreased current speeds lessen the likelihood
546 of scouring and sediment resuspension, supporting a role for sinking particles from shallow
547 sources. Additional measurements in this region would help better constrain this hypothesis.

548 The straits in the Solomon Sea also exhibit increases in DMn with depth, which, in the deep
549 layer, generally indicates recent inputs from sedimentary and/or hydrothermal sources but can
550 also be associated with oxygen deficient zones. Prior to reaching the Woodlark Rise, within the
551 Woodlark basin, hydrothermal activity is present at $>2500 \text{ m}$ (Laurila et al., 2012) with
552 shallower hydrothermal activity inferred at Kana Keoki seamount at 650 m near the Solomon
553 Archipelago (InterRidge Database: <https://vents-data.interridge.org/>). These are unlikely to be
554 the source of the Mn enrichments between 700 and 1350 m , however. Given the highly tectonic
555 and magmatic character of the region, hydrothermal hot springs likely exist throughout the basin.
556 Diffusion from or resuspension of sediments are also possible source mechanisms for DMn
557 whose input coincides with transport through and along the Woodlark Rise. Dissolved oxygen
558 remains relatively elevated throughout the basin at these depths, suggesting that particle
559 remineralization does not play a large role in maintaining DMn levels. This is supported by
560 dissolved Ce concentrations (Pham et al., 2019), which decrease as water flows through the
561 basin, and Ce anomalies in the deeper ocean that do not indicate any shifts towards either
562 remineralization or increased solubility due to changes in redox chemistry. The increase in DMn
563 concentrations between stations 71 and 60 versus those in the straits suggest that sediments may
564 be the predominant sources of DMn in the regional deep layer. The absence of an enrichment in

565 DMn, DA1, and TDA1 at station 39 (and small TDMn enrichment) suggests that there is little to
566 no Mn input as the NGCU transits through and around bathymetry prior to reaching station 39 at
567 these depths. While it is possible that the NGCU is bathymetrically steered away from station
568 39, it seems more likely that Mn and/or Al are added after passing this station. The presence of
569 DMn at this station suggests that the TDMn is sourced at depth from the oxidation and
570 scavenging of DMn. Station 21 also shows enrichments in TDMn, TDA1, and DA1 relative to
571 stations outside of the Solomon Sea suggesting the accumulation of dissolved metals and
572 sediments as water flows through the Solomon Archipelago. While various small enrichments
573 are observed at stations outside of the Solomon Sea, it is not possible to evaluate background
574 values for these stations. However, we do note that the SEC flows past many shallow seamounts
575 en route to the Solomon Sea.

576 Estimated fluxes of trace metals in the deep layer are determined based on transport estimates for
577 the PANDORA cruise calculated by Germaineaud et al. (2016) and are reported in Table 4.
578 Because there is less mixing deeper in the ocean, transport estimates are more consistent between
579 the inflow (9.3 ± 1.4 Sv) and outflow (8.4 ± 2.6 Sv) than they are in the surface layer, but there are
580 still large uncertainties on transport estimates deeper in the water column, as well as a slight
581 imbalance between inflowing and outflowing water transport. Fluxes reported here for total
582 outflow show that trace metal modifications are minimal relative to inflowing water, despite
583 regions of local enrichment observed in profile data.

584

Table 4: Transport and trace metal flux estimates for the deep layer ($>26.9\sigma_\theta$) in the Solomon Sea. Transport estimates from Germaineaud et al. (2016). Solomon Strait and St. George's Channel transport are estimated based on total transit through the two outflow straits. The combined error from the transport estimate and the analytical error (1 SD) of the average concentration.

Station	Transport (Sv)	DAI (mol/s)	TDAI (mol/s)	DMn (mol/s)	TDMn (mol/s)
Inflow	9.3 ± 1.4	69.6 ± 12.1	68.6±12.0	2.76±0.78	4.51±0.96
Vitiaz Strait (station 77)	4.6± 0.2	37.3 ± 4.0	53.0 ± 4.2	2.33 ± 0.20	4.05 ± 0.37
Solomon Strait (station 42)	1.9 ± 1.2	17.7 ±11.6	24.5 ± 15.5	0.87 ±0.56	1.54 ±0.98
St. George's Channel (station 60)	1.9 ± 1.2	11.6 ±7.5	18.9 ±12.0	0.81 ±0.51	1.40 ±0.89
Total Outflow	8.4 ± 2.6	66.61 ± 22.7	96.3 ±31.7	4.01 ±1.27	6.99 ±2.24

585

586

587 ***4.2 Constraining the Solomon Sea Flux to the EUC***

588 The EUC is largely composed of water originating from the NGCU/thermocline waters (24-26.9
589 σ_θ), and so we can compare the flux of trace metals out of the Solomon Sea (Table 3) in this
590 layer to the flux of trace metals in the EUC over the same density interval by using average DAI,
591 TDAI, DMn, and TDMn concentrations from the EUC collected at 0°, 156 °E by Slemmons et al.
592 (2010; Station 22). At this longitude, water exiting the Solomon Sea would have been entrained
593 in the EUC. These measurements were collected six years before the PANDORA cruise, but
594 during the same season, and in a similar ENSO phase. The implicit assumption in the
595 comparison of flux out of the Solomon Sea and the flux into the EUC is that everything leaving
596 the Solomon Sea at that depth range enters the EUC, as shown in Fig. 6. This is clearly not the
597 case, but this assumption establishes an upper bound for trace metals contributed from the
598 Solomon Sea to the EUC. For flux estimates several additional assumptions are made: the
599 Solomon Sea is in steady state and not temporally variable; stations in the straits are
600 representative of water leaving the Solomon Sea; and that the two profiles used for inflow waters

601 are, on average, representative water entering the Solomon Sea. The trace metal fluxes of the
 602 Solomon Sea Inflow, Outflow, and flux at the EUC are shown in Table 5.

Table 5. Average metal fluxes into and out of the Solomon Sea at $24-26.9\sigma_{\theta}$ compared to metal fluxes in the EUC 156°E . This assumes that transport in and out of the Solomon Sea is $22.8 \text{ Sv} \pm 1.0 \text{ Sv}$. Errors are calculated using uncertainty reported on transport estimates (1SD) by Germaineaud et al. (2016) and analytical error of the concentration data (1SD)

	DAI	TDAI	DMn	TDMn
Inflow (mol s^{-1})	136 ± 19	147 ± 35	6.1 ± 0.5	9.1 ± 0.6
Outflow (mol s^{-1})	190 ± 19	217 ± 16	6.1 ± 0.6	10.8 ± 1.0
In EUC (mol s^{-1})	160 ± 9	177 ± 10	7.8 ± 0.4	17 ± 0.9

603

604 4.2.1 Sources of water to the EUC

605 It is important to note that the EUC does not receive its water solely from the Solomon Sea, and
 606 the breakdown of how much water comes from the Solomon Sea versus from other Southern and
 607 Northern Hemisphere sources is still unclear. Hydrographic parameters measured during the
 608 Western Equatorial Pacific Ocean Study (WEPOCS, 1985/1986) (e.g., Lindstrom et al., 1987;
 609 Tsuchiya et al., 1989) show the EUC was fed predominantly by water from the Southern
 610 Hemisphere with $\sim 2/3$ from southern low latitude western boundary currents, including the
 611 NGCU (Tsuchiya et al., 1989) and $\sim 1/4$ of the water originating in the Northern Hemisphere
 612 (Lindstrom et al., 1987; Tsuchiya et al., 1989). However, this study also observed a seasonally
 613 variable NGCU, including one season in which it flowed at the same rate as the EUC. Since then,
 614 other studies have investigated the relative importance of water from northern and southern
 615 sources and have come to differing conclusions about the proportion of water from the north and
 616 south. Grenier et al. (2011) estimate about 58% of the water in the EUC at 156°W passed
 617 through the Solomon Sea, and about 33% comes from the Mindanao Current in the north, with

618 the remainder coming from other sources south of the EUC. A modeling study by Izumo et al.
619 (2002) indicates roughly equal contributions from the Northern and Southern Hemispheres,
620 while other studies suggest the dominance of a southern source (Liu and Huang, 1998). Most
621 recently, biogeochemical tracer data indicate that the nutrients in the EUC are primarily derived
622 from Southern Hemisphere water: based on isotopic tracers (^{15}N and ^{18}O measured in nitrate) and
623 oxygen measurements combined with the nutrient data (silicic acid and nitrate), the fraction of
624 water in the EUC derived from the Mindanao Current (in the Northern Hemisphere) is estimated
625 to be much smaller than that from the Southern Hemisphere (Lehmann et al., 2018).

626 The relative contributions of trace metals from the Solomon Sea to the EUC trace metal pool can
627 then be estimated for the different estimates of northern versus southern water contribution to the
628 EUC (Table 6). Averaging these estimates, our data suggest that about 70% of the DA1, 80% of
629 the TDA1, 50% of DMn, and 40% of the TDMn in the EUC must come from waters exiting the
630 Solomon Sea. It is also important to note that there is seasonal variability in these estimates, and
631 that as the EUC flows east, it entrains additional water, and the relative importance of the waters
632 originally feeding the EUC diminishes during its eastward transit (Qin et al., 2015). Our results
633 indicate that the Solomon Sea can supply a majority of the required Al to the EUC. However,
634 relative to Al, there is proportionally less Mn in the EUC coming from the Solomon Sea,
635 suggesting other sources to the EUC must be more enriched in Mn and have higher Mn/Al ratio.

636 Because there are other sources of water to the EUC, there are other potential sources of trace
637 metals to the EUC that are not accounted for with these Solomon Sea data. Waters exiting the
638 Solomon Sea through the Vitiaz Strait pass along the PNG coast/shelf and through the Bismark
639 Sea. However Mn data collected at stations just beyond the Vitiaz Strait along the PNG coast by
640 Slemons et al. (2010) are remarkably consistent with Mn concentrations reported here. By
641 comparison Al concentrations along the coast vary greatly, especially in the surface, but are
642 generally lower by $\sim 2\text{-}5$ nM (and in the surface up to 35 nM) compared to the data for the
643 Solomon Sea reported here. These findings indicate that the PNG coastline does not provide the

644 missing Mn and may be a sink for Al. This is consistent with the findings of Mackey et al.
 645 (2002a) who show that while concentrations of Fe and Mn are high off the coast of PNG
 646 especially near the outflow of the Sepik River, this riverine source of Mn and Fe to the Bismarck
 647 Sea is insufficient to produce the concentrations of these elements observed in the EUC. Just
 648 beyond the Solomon Strait, enrichments observed for Rare Earth Elements indicate hydrothermal
 649 and/or lithogenic inputs to the waters along the flow path of the NICU (Behrens et al., 2020)
 650 between the Solomon Sea and EUC, however there are no Al and Mn data in this region and
 651 thus, while we might anticipate inputs of Al and Mn to the NICU, conclusions about their inputs
 652 to the NICU are not possible. As noted above, the EUC is derived from waters originating both
 653 from the south and north of the equator and the balance of Al and Mn required to balance the
 654 trace metal budget of the EUC may have their source there.

Table 6. Contribution of the Solomon Sea contribution to EUC trace metal pool based on variations in source waters. The flux of each trace metal out of the Solomon Sea is multiplied by the percentage of water contributed by the Solomon Sea to the EUC at 156°E based on different studies.

Reference	% of EUC derived from Southern Hemisphere	Method of Estimate	% trace metals contributed to EUC based on magnitude of source			
			DAI	TDAI	DMn	TDMn
	100%	Upper Bound	119	123	79	64
Grenier et al., 2011 ¹	59%	Model	70	72	47	38
Tsuchiya et al., 1989	67%	WEPOCS	80	82	53	43
Izumo et al., 2002	52%	Model	62	64	41	33
Lehman et al., 2018 ²	<70%	O ₂ , Si, N isotopes	83	86	55	45
Qin et al., 2015 ³	63%	Model	75	77	50	40
Average			74	76	49	40

¹ Fluxes are contribution from Solomon Sea, specifically, ² Looks at upper and lower EUC, ³40% Solomon Strait, 23% Vitiaz Strait at 165°E

655

656 **5. Conclusions**

657 Our data show that water exiting the Solomon Sea is important in supplying aluminum, and to a
658 lesser degree, manganese, to the Equatorial Undercurrent. However, from a net budget
659 perspective, the amount of Al and Mn enrichment that occurs within the basin is small, relative
660 to the inflow concentrations, indicating that most of the Al and Mn was acquired prior to
661 reaching the Solomon Sea or that inputs are approximately balanced by scavenging within the
662 basin. This is also true for deeper water in the Solomon Sea. The trace metal pool in the surface
663 layer does appear to be enriched relative to inflowing waters, but because of large seasonal
664 variability and disparities in water transport between entrance and exit, more studies are needed
665 to conclude this definitively.

666 Our work is consistent with previous studies and supports the idea that boundary exchange
667 through sediment resuspension, non-reductive release of dissolved species from suspended
668 particles, and scavenging onto these suspended particles are the dominant processes providing
669 additional metals to the Solomon Sea (e.g., Lacan and Jeandel, 2005; Grenier et al., 2013;
670 Labatut et al., 2014; Jeandel, 2016). These highly localized input processes indicate that a higher
671 resolution study is required to better understand these boundary processes, and in particular, their
672 potential to impact larger-scale cycling.

673 The importance of the Solomon Sea as a source of trace nutrients to the EUC is heavily
674 influenced by the proportion of water in the EUC that is derived from the Southern Hemisphere.
675 Improving our estimates of the EUC water mass breakdown from important regions of input in
676 the Northern Hemisphere (e.g., North Equatorial Current, Mindanao Current) will be valuable in
677 better constraining these fluxes. Higher sampling resolution (spatial and temporal) within the
678 region, including within the EUC, at potential source regions south of the Solomon Sea
679 (including near Vanuatu, and in the Coral Sea), and from the SEC will be important in
680 constraining the impact of the waters passing through the Solomon Sea on the chemistry of the
32

681 EUC. This increased resolution would allow for a more thorough understanding of trace metal
682 and nutrient dynamics in the western Pacific and their impact on the HNLC eastern equatorial
683 Pacific.

684 **Acknowledgements**

685 M. Grand, M. Labatut and F. Qu  rou   are thanked for their hard work with trace element
686 sampling. Thanks to the crew of the R/V l'Atalante and Chief Scientist G. Eldin, for technical
687 assistance, and expertise during the PANDORA Cruise. W. Kessler provided helpful discussion
688 about the circulation of the Solomon Sea. Most geochemical graphics were created using ODV
689 (Schlitzer, R., Ocean Data View, <http://odv.awi.de>, 2015).

690 **Funding**

691 This publication is partially funded by the Joint Institute for the Study of the Atmosphere and
692 Ocean (JISAO) under NOAA Cooperative Agreement NA15OAR4320063, Contribution No.
693 2018-0188, and PMEL Contribution No. 4904. JAR, NB, SM were funded by US NSF award
694 OCE-1237011 and OCE-0649505. The PANDORA cruise was co-funded by NSF grant
695 OCE1029487, and by ANR project ANR- 09-BLAN-0233-01 and INSU/LEFE project Solwara
696 (IDAO and CYBER).

697

698 **References**

- 699 Alberty, M., Sprintall, J., MacKinnon, J., Germaineaud, C., Cravatte, S., Ganachaud, A., 2019.
700 Moored Observations of Transport in the Solomon Sea. *J. Geophys. Res. Ocean.* 124, 8166–
701 8192. <https://doi.org/10.1029/2019JC015143>
- 702 Alongi, D.M., Boyle, S.G., Tirendi, F., Payn, C., 1996. Composition and behaviour of trace
703 metals in post-oxic sediments of the Gulf of Papua, Papua New Guinea. *Estuar. Coast. Shelf Sci.*
704 42, 197–211. <https://doi.org/10.1006/ecss.1996.0015>
- 705 Baker, E.T., Massoth, G.J., de Ronde, C.E.J., Lupton, J.E., McInnes, B.I.A., 2002. Observations
706 and sampling of an ongoing subsurface eruption of Kavachi volcano, Solomon Islands, May
707 2000. *Geology* 30, 975–978. [https://doi.org/10.1130/0091-
708 7613\(2002\)030<0975:OASOAO>2.0.CO;2](https://doi.org/10.1130/0091-7613(2002)030<0975:OASOAO>2.0.CO;2)
- 709 Behrens, M.K., Pahnke, K., Delpech, A., Cravatte, S., Marin, F., Jeandel, C., 2020. Trace
710 element sources and fluxes in the zonal current system of the western tropical Pacific Ocean:
711 Evidence from combined rare earth element, Nd isotope distributions and physical observations.
712 Ocean Sciences Meeting, San Diego, CA.
- 713 Brown, M.T., Bruland, K.W., 2008. An improved flow-injection analysis method for the
714 determination of dissolved aluminum in seawater. *Limnol. Oceanogr. Methods* 6, 87–95.
715 <https://doi.org/10.4319/lom.2008.6.87>
- 716 Buck, C.S., Landing, W.M., Resing, J.A., 2013. Pacific Ocean aerosols: Deposition and
717 solubility of iron, aluminum, and other trace elements. *Mar. Chem.* 157, 117–130.
718 <https://doi.org/10.1016/j.marchem.2013.09.005>

- 719 Buck, C.S., Landing, W.M., Resing, J.A., Lebon, G.T., 2006. Aerosol iron and aluminum
720 solubility in the northwest Pacific Ocean: Results from the 2002 IOC cruise. *Geochem. Geophys.*
721 *Geosyst.* 7, 1–21. <https://doi.org/10.1029/2005GC000977>
- 722 Butt, J., Lindstrom, E.J., 1994. Currents off the east coast of New Ireland, Papua New Guinea,
723 and their relevance to regional undercurrents in the western equatorial Pacific Ocean. *J.*
724 *Geophys. Res.* 99, 12503-12514. <https://doi.org/10.1029/94JC00399>
- 725 Coale, K.H., Fitzwater, S.E., Gordon, R.M., Johnson, K.S., Barber, R.T., 1996. Control of
726 community growth and export production by upwelled iron in the equatorial Pacific Ocean.
727 *Nature* 379, 621–624. <https://doi.org/10.1038/379621a0>
- 728 Cravatte, S., Ganachaud, A., Duong, Q.P., Kessler, W.S., Eldin, G., Dutrieux, P., 2011. Observed
729 circulation in the Solomon Sea from SADC data. *Prog. Oceanogr.* 88, 116–130.
730 <https://doi.org/10.1016/j.pocean.2010.12.015>
- 731 Cresswell, G.R., 2000. Coastal currents of northern Papua New Guinea, and the Sepik River
732 outflow. *Mar. Freshw. Res.* 51, 553–564. <https://doi.org/10.1071/MF99135>
- 733 de Boyer Montégut, C., Madec, G., Fischer, A.S., Lazar, A., Iudicone, D., 2004. Mixed layer
734 depth over the global ocean: An examination of profile data and a profile-based climatology. *J.*
735 *Geophys. Res.* 109, C12003. <https://doi.org/10.1029/2004JC002378>
- 736 Delcroix, T., Radenac, M.-H., Cravatte, S., Alory, G., Gourdeau, L., Leger, F., Singh, A.,
737 Varillon, D., 2014. Sea surface temperature and salinity seasonal changes in the western
738 Solomon and Bismarck Seas. *J. Geophys. Res. Oceans* 119, 2642–2657.
739 <https://doi.org/10.1002/2013JC009733>

- 740 Ganachaud, A., Cravatte, S., Sprintall, J., Germineaud, C., Albery, M.S., Jeandel, C., Eldin, G.,
741 Metzl, N., Bonnet, S., Benavides, M., Heimburger, L.-E., Lefèvre, J., Michael, S., Resing, J.A.,
742 Quéroùé, F., Sarthou, G., Rodier, M., Berthelot, H., Baurand, F., Grelet, J., Hasegawa, T.,
743 Kessler, W., Kilepak, M., Lacan, F., Privat, E., Send, U., Van Beek, P., Souhaut, M., Sonke, J.E.,
744 2017. The Solomon Sea: its circulation, chemistry, geochemistry and biology explored during
745 two oceanographic cruises. *Elem. Sci. Anth.* 5, 33. <https://doi.org/10.1525/elementa.221>
- 746 Gasparin, F., Ganachaud, A., Maes, C., Marin, F., Eldin, G., 2012. Oceanic transports through
747 the Solomon Sea: The bend of the New Guinea Coastal Undercurrent. *Geophys. Res. Lett.* 39, 1–
748 6. <https://doi.org/10.1029/2012GL052575>
- 749 Germineaud, C., Ganachaud, A., Sprintall, J., Cravatte, S., Eldin, G., Albery, M.S., Privat, E.,
750 2016. Pathways and water mass properties of the thermocline and intermediate waters in the
751 Solomon Sea. *J. Phys. Oceanogr.* 46, 3031–3049. <https://doi.org/10.1175/JPO-D-16-0107.1>
- 752 Gordon, R.M., Coale, K.H., Johnson, K.K.S., 1997. Iron distributions in the equatorial Pacific:
753 Implications for new production. *Limnol. Oceanogr.* 42, 419–431.
754 <https://doi.org/10.4319/lo.1997.42.3.0419>
- 755 Grenier, M., Cravatte, S., Blanke, B., Menkes, C., Koch-Larrouy, A., Durand, F., Melet, A.,
756 Jeandel, C., 2011. From the western boundary currents to the Pacific Equatorial Undercurrent:
757 Modeled pathways and water mass evolutions. *J. Geophys. Res. Oceans* 116, C12044.
758 <https://doi.org/10.1029/2011JC007477>
- 759 Grenier, M., Jeandel, C., Lacan, F., 2013. From the subtropics to the central equatorial Pacific
760 Ocean: Neodymium isotopic composition and rare earth element concentration variations. *J.*
761 *Geophys. Res. Oceans* 118, 592–618. <https://doi.org/10.1029/2012JC8239>

- 762 Harris, P.T., Baker, E.K., Cole, A.R., Short, S.A., 1993. A preliminary study of sedimentation in
763 the tidally dominated Fly River Delta, Gulf of Papua. *Cont. Shelf Res.* 13, 441–472.
764 [https://doi.org/10.1016/0278-4343\(93\)90060-B](https://doi.org/10.1016/0278-4343(93)90060-B)
- 765 Hettler, J., Irion, G., Lehmann, B., 1997. Environmental impact of mining waste disposal on a
766 tropical lowland river system: A case study on the Ok Tedi Mine, Papua New Guinea. *Miner.*
767 *Depos.* 32, 280–291. <https://doi.org/10.1007/s001260050093>
- 768 Hristova, H.G., Kessler, W.S., 2012. Surface circulation in the Solomon Sea derived from
769 Lagrangian drifter observations. *J. Phys. Oceanogr.* 42, 448–458. [https://doi.org/10.1175/JPO-D-](https://doi.org/10.1175/JPO-D-11-099.1)
770 [11-099.1](https://doi.org/10.1175/JPO-D-11-099.1)
- 771 Izumo, T., Picaut, J., Blanke, B., 2002. Tropical pathways, equatorial undercurrent variability
772 and the 1998 La Niña. *Geophys. Res. Lett.* 29, 2080. <https://doi.org/10.1029/2002GL015073>
- 773 Jankowski, P., 2011. Independent Technical Assessment of Sea Floor Massive Sulphide
774 Exploration Tenements in Papua New Guinea, Fiji and Tonga. SRk Consult. Jeandel, C., 2016.
775 Overview of the mechanisms that could explain the ‘Boundary Exchange’ at the land–ocean
776 contact. *Philos. Trans. R. Soc. A* 374, 20150287. <https://doi.org/10.1098/rsta.2015.0287>
- 777 Jeandel, C., 2016. Overview of the mechanisms that could explain the ‘Boundary Exchange’ at
778 the land–ocean contact. *Philos. Trans. R. Soc. A Math. Phys. Eng. Sci.* 374, 20150287.
779 <https://doi.org/10.1098/rsta.2015.0287>
- 780 Kaupp, L.J., Measures, C.I., Selph, K.E., Mackenzie, F.T., 2011. The distribution of dissolved Fe
781 and Al in the upper waters of the eastern equatorial Pacific. *Deep-Sea Res. Part II Top. Stud.*
782 *Oceanogr.* 58, 296–310. <https://doi.org/10.1016/j.dsr2.2010.08.009>

- 783 Kessler, W.S., Cravatte, S., 2013. Mean circulation of the Coral Sea. *J. Geophys. Res. Ocean.*
784 118, 6385–6410. <https://doi.org/10.1002/2013JC009117>
- 785 Kessler, W.S., H.G. Hristova and R.E. Davis, 2019. Equatorward western boundary transport
786 from the South Pacific: Glider observations, dynamics, and consequences. *Prog. Oceanogr.*, 175
787 (2019) 208–225. <https://doi.org/10.1016/j.pocean.2019.04.005>
- 788 Kineke, G.C., Woolfe, K.J., Kuehl, S.A., Milliman, J.D., Dellapenna, T.M., Purdon, R.G., 2000.
789 Sediment export from the Sepik River, Papua New Guinea: Evidence for a divergent sediment
790 plume. *Cont. Shelf Res.* 20, 2239–2266. [https://doi.org/10.1016/S0278-4343\(00\)00069-8](https://doi.org/10.1016/S0278-4343(00)00069-8)
- 791 Labatut, M., Lacan, F., Pradoux, C., Chmeleff, J., Radic, A., Murray, J.W., Poitrasson, F.,
792 Johansen, A.M., Thil, F., 2014. Iron sources and dissolved-particulate interactions in the
793 seawater of the western equatorial Pacific, iron isotope perspectives. *Global Biogeochem. Cycles*
794 28, 1044–1065. <https://doi.org/10.1002/2014GB004928>
- 795 Lacan, F., Jeandel, C., 2001. Tracing Papua New Guinea imprint on the central Equatorial
796 Pacific Ocean using neodymium isotopic compositions and Rare Earth Element patterns. *Earth*
797 *Planet. Sci. Lett.* 186, 497–512. [https://doi.org/10.1016/S0012-821X\(01\)00263-1](https://doi.org/10.1016/S0012-821X(01)00263-1)
- 798 Lacan, F., Jeandel, C., 2005. Neodymium isotopes as a new tool for quantifying exchange fluxes
799 at the continent-ocean interface. *Earth Planet. Sci. Lett.* 232, 245–257.
800 <https://doi.org/10.1016/j.epsl.2005.01.004>
- 801 Laurila, T.E., Petersen, S., Devey, C.W., Baker, E.T., Augustin, N., Hannington, M.D., 2012.
802 Tectonic and magmatic controls on hydrothermal activity in the Woodlark Basin. *Geochemistry,*
803 *Geophys. Geosystems* 13, 1–13. <https://doi.org/10.1029/2012GC004247>

- 804 Lehmann, N., Granger, J., Kienast, M., Brown, K.S., Rafter, P.A., Martínez-Méndez, G.,
805 Mohtadi, M., 2018. Isotopic evidence for the evolution of subsurface nitrate in the western
806 equatorial Pacific. *J. Geophys. Res. Oceans* 123, 1684–1707.
807 <https://doi.org/10.1002/2017JC013527>
- 808 Lindstrom, E.J., Lukas, R., Fine, R.A., Firing, E., Godfrey, S., Meyers, G., Tsuchiya, M., 1987.
809 The western equatorial Pacific Ocean circulation study. *Nature* 330, 533–537.
810 <https://doi.org/10.1038/330533a0>
- 811 Liu, Z., Huang, B., 1998. Why is there a tritium maximum in the central equatorial Pacific
812 thermocline? *J. Phys. Oceanogr.* 28, 1527–1533. [https://doi.org/10.1175/1520-](https://doi.org/10.1175/1520-0485(1998)028<1527:WITATM>2.0.CO;2)
813 [0485\(1998\)028<1527:WITATM>2.0.CO;2](https://doi.org/10.1175/1520-0485(1998)028<1527:WITATM>2.0.CO;2)
- 814 Mackey, D.J., O’Sullivan, J., Watson, R.J., 2002a. Iron in the western Pacific: a riverine or
815 hydrothermal source for iron in the Equatorial Undercurrent? *Deep-Sea Res. Part I-Oceanogr.*
816 *Res. Pap.* 49, 877–893. [https://doi.org/10.1016/S0967-0637\(01\)00075-9](https://doi.org/10.1016/S0967-0637(01)00075-9)
- 817 Mackey, D.J., O’Sullivan, J.E., Watson, R.J., Dal Pont, G., 2002b. Trace metals in the Western
818 Pacific: temporal and spatial variability in the concentrations of Cd, Cu, Mn and Ni. *Deep-Sea*
819 *Res. Part I-Oceanogr. Res. Pap.* 49, 2241–2259. [https://doi.org/10.1016/S0967-0637\(02\)00124-3](https://doi.org/10.1016/S0967-0637(02)00124-3)
- 820 Martin, J.H., Knauer, G.A., Karl, D.M., Broenkow, W.W., 1987. VERTEX: carbon cycling in
821 the northeast Pacific. *Deep-Sea Res. Part A, Oceanogr. Res. Pap.* 34, 267–285.
822 [https://doi.org/10.1016/0198-0149\(87\)90086-0](https://doi.org/10.1016/0198-0149(87)90086-0)
- 823 McConachy, T., Binns, R., Arculus, R.J., 2002. Submarine hydrothermal activity and volcanic
824 petrogenesis associated with the birth of island arcs in the Solomon Islands (SOLA VENTS -

- 825 2002). Cruise Report of RV Franklin, Commonwealth Scientific and Industrial Research
826 Organisation, Australia.
- 827 Measures, C.I., Brown, E.T., 1996. Estimating dust input to the Atlantic Ocean using surface
828 water aluminium concentrations, in: Guerzoni, S., Chester, R. (Eds.), *The Impact of Desert Dust*
829 *across the Mediterranean*. Springer, Dordrecht, pp. 301–311. [https://doi.org/10.1007/978-94-017-](https://doi.org/10.1007/978-94-017-3354-0_30)
830 [3354-0_30](https://doi.org/10.1007/978-94-017-3354-0_30)
- 831 Melet, A., Gourdeau, L., Kessler, W.S., Verron, J., Molines, J.-M., 2010. Thermocline
832 circulation in the Solomon Sea: A modeling study. *J. Phys. Oceanogr.* 40, 1302–1319.
833 <https://doi.org/10.1175/2009JPO4264.1>
- 834 Melet, A., Gourdeau, L., Verron, J., Djath, B., 2013. Solomon Sea circulation and water mass
835 modifications: Response at ENSO timescales. *Ocean Dyn.* 63, 1–19.
836 <https://doi.org/10.1007/s10236-012-0582-0>
- 837 Melet, A., Verron, J., Gourdeau, L., Koch-Larrouy, A., 2011. Equatorward Pathways of Solomon
838 Sea Water Masses and Their Modifications. *J. Phys. Oceanogr.* 810–826.
839 <https://doi.org/10.1175/2010JPO4559.1>
- 840 Milliman, J.D., Farnsworth, K.L., Albertin, C.S., 1999. Flux and fate of fluvial sediments leaving
841 large islands in the East Indies. *J. Sea Res.* 41, 97–107. [https://doi.org/10.1016/S1385-](https://doi.org/10.1016/S1385-1101(98)00040-9)
842 [1101\(98\)00040-9](https://doi.org/10.1016/S1385-1101(98)00040-9)
- 843 Moran, S.B., Moore, R.M., 1991. The potential source of dissolved aluminum from resuspended
844 sediments to the North Atlantic Deep Water. *Geochim. Cosmochim. Acta* 55, 2745–2751.
845 [https://doi.org/10.1016/0016-7037\(91\)90441-7](https://doi.org/10.1016/0016-7037(91)90441-7)

- 846 Murray, J.W., Johnson, E., Garside, C., 1995. A U.S. JGOFS process study in the equatorial
847 Pacific (EqPac): Introduction. *Deep. Res. Part II* 42, 275–293. [https://doi.org/10.1016/0967-](https://doi.org/10.1016/0967-0645(95)00044-Q)
848 0645(95)00044-Q
- 849 Obata, H., Shitashima, K., Isshiki, K., Nakayama, E., 2008. Iron, manganese and aluminum in
850 upper waters of the western South Pacific Ocean and its adjacent seas. *J. Oceanogr.* 64, 233–245.
851 <https://doi.org/10.1007/s10872-008-0018-0>
- 852 Pham, V.Q., Grenier, M., Cravatte, S., Michael, S., Jacquet, S., Belhadj, M., Nachez, Y.,
853 Germineaud, C., Jeandel, C., 2019. Dissolved rare earth elements distribution in the Solomon
854 Sea. *Chem. Geol.* 524, 11–36. <https://doi.org/10.1016/j.chemgeo.2019.05.012>
- 855 Phillips, B., Dunbabin, M., Henning, B., Howell, C., Deciccio, A., Flinders, A., Kelley, K.A.,
856 Scott, J.J., Albert, S., Carey, S., Tsadok, R., 2016. Exploring the “Sharkcano”: Biogeochemical
857 observations of the Kavachi submarine volcano (Solomon Islands). *Oceanography* 24, 160–169.
858 <https://doi.org/10.5670/oceanog.2016.85>
- 859 Pichler, T., Veizer, J., Hall, G.E.M., 1999. The chemical composition of shallow-water
860 hydrothermal fluids in Tutum Bay, Ambitle Island, Papua New Guinea and their effect on
861 ambient seawater. *Mar. Chem.* 64, 229–252. [https://doi.org/10.1016/S0304-4203\(98\)00076-0](https://doi.org/10.1016/S0304-4203(98)00076-0)
- 862 Qin, X., Sen Gupta, A., Van Sebille, E., 2015. Variability in the origins and pathways of Pacific
863 Equatorial Undercurrent water. *J. Geophys. Res. Oceans* 120, 3113–3128.
864 <https://doi.org/10.1002/2014JC010549>

- 865 Resing, J.A., Measures, C.I., 1994. Fluorometric determination of Al in seawater by flow
866 injection analysis with in-line preconcentration. *Anal. Chem.* 66, 4105–4111.
867 <https://doi.org/10.1021/ac00094a039>
- 868 Resing, J.A., Mottl, M., 1992. Determination of manganese in seawater using flow injection
869 analysis with on-line preconcentration and spectrophotometric detection. *Anal. Chem.* 64, 2682–
870 2687. <https://doi.org/10.1021/ac00046a006>
- 871 Resing, J.A., Sedwick, P.N., German, C.R., Jenkins, W.J., Moffett, J.W., Sohst, B.M., Tagliabue,
872 A., 2015. Basin-scale transport of hydrothermal dissolved metals across the South Pacific Ocean.
873 *Nature* 523, 200–203. <https://doi.org/10.1038/nature14577>
- 874 Ryan, J.P., Ueki, I., Chao, Y., Zhang, H., Polito, P.S., Chavez, F.P., 2006. Western Pacific
875 modulation of large phytoplankton blooms in the central and eastern equatorial Pacific. *J.*
876 *Geophys. Res. Biogeosciences* 111, 1–14. <https://doi.org/10.1029/2005JG000084>
- 877 Salomons, W., Eagle, A.M., 1990. Hydrology, sedimentology and the fate and distribution of
878 copper in mine-related discharges in the fly river system, Papua New Guinea. *Sci. Total Environ.*
879 97–98, 315–334. [https://doi.org/10.1016/0048-9697\(90\)90248-S](https://doi.org/10.1016/0048-9697(90)90248-S)
- 880 Shank, L.M., Johansen, A.M., 2008. Atmospheric trace metal and labile iron deposition fluxes to
881 the equatorial Pacific during EUCFe2006, Ocean Sciences Meeting, Orlando, Fla, USA.
- 882 Sholkovitz, E.R., Elderfield, H., Szymczak, R., Casey, K., 1999. Island weathering: River
883 sources of rare earth elements to the Western Pacific Ocean. *Mar. Chem.* 68, 39–57.
884 [https://doi.org/10.1016/S0304-4203\(99\)00064-X](https://doi.org/10.1016/S0304-4203(99)00064-X)

- 885 Singh, N.D., Chinni, V., Singh, S.K., 2020. Dissolved aluminium cycling in the northern,
886 equatorial and subtropical gyre region of the Indian Ocean. *Geochim. Cosmochim. Acta* 268,
887 160–185. <https://doi.org/10.1016/j.gca.2019.09.028>
- 888 Slemons, L.O., Gorgues, T., Aumont, O., Menkes, C., Murray, J.W., 2009. Biogeochemical
889 impact of a model western iron source in the Pacific Equatorial Undercurrent. *Deep-Sea Res.*
890 *Part I Oceanogr. Res. Pap.* 56, 2115–2128. <https://doi.org/10.1016/j.dsr.2009.08.005>
- 891 Slemons, L.O., Murray, J.W., Resing, J.A., Paul, B., Dutrieux, P., 2010. Western Pacific coastal
892 sources of iron, manganese, and aluminum to the Equatorial Undercurrent. *Global Biogeochem.*
893 *Cycles* 24, GB3024. <https://doi.org/10.1029/2009GB003693>
- 894 Slemons, L.O., Paul, B., Resing, J.A., Murray, J.W., 2012. Particulate iron, aluminum, and
895 manganese in the Pacific equatorial undercurrent and low latitude western boundary current
896 sources. *Mar. Chem.* 142–144, 54–67. <https://doi.org/10.1016/j.marchem.2012.08.003>
- 897 Sokolov, S., Rintoul, S., 2000. Circulation and water masses of the southwest Pacific: WOCE
898 Section P11, Papua New Guinea to Tasmania. *J. Mar. Res.* 58, 223–268.
899 <https://doi.org/10.1357/002224000321511151>
- 900 Taylor, S.R., 1964. Abundance of chemical elements in the continental crust: a new table.
901 *Geochim. Cosmochim. Acta* 28, 1273–1285. [https://doi.org/10.1016/0016-7037\(64\)90129-2](https://doi.org/10.1016/0016-7037(64)90129-2)
- 902 Tsuchiya, M., Lukas, R., Fine, R.A., Firing, E., Lindstrom, E.J., 1989. Source waters of the
903 Pacific Equatorial Undercurrent. *Prog. Oceanogr.* 23, 101–147. [https://doi.org/10.1016/0079-](https://doi.org/10.1016/0079-6611(89)90012-8)
904 [6611\(89\)90012-8](https://doi.org/10.1016/0079-6611(89)90012-8)

905 **Figure Captions**

906 **Fig. 1. a)** Sub-surface currents [dark blue; after Germaineaud et al. (2016)], rivers (blue text),
907 straits (purple text) and potential point sources (yellow) to the Solomon Sea. These currents are:
908 Equatorial Undercurrent (EUC); South Equatorial Current (SEC); New Guinea Coastal
909 Undercurrent (NGCU); New Ireland Coastal Undercurrent (NICU); Saint George's Undercurrent
910 (SGU); North Queensland Current (NQC); New Caledonia Jet (NCJ); North Vanuatu Jet (NVJ);
911 East Australian Current (EAC)
912 Red dots and numbers represent stations sampled for Al and Mn. Orange dot shows the
913 location of station 22 from Slemons et al. (2010). **b)** Bathymetry of the Woodlark
914 Basin and Trobriand Islands.

915 **Fig. 2.** Dissolved Al (DAI) profiles versus potential density from the PANDORA cruise. Light
916 blue circles show measured DAI concentrations. Orange line represents average inflow DAI
917 profile, while grey shading represents the bounds of the average profile. Green box represents the
918 density interval over which budget of the thermocline waters is calculated. **(a–c)** waters that are
919 located south of the Solomon Sea; **(d–f)** profiles found along the NGCU; **(g–i)** profiles that are at
920 the exit straits of the Solomon Sea; **(j–l)** located outside the Solomon Sea.

921 **Fig. 3.** Total Dissolvable Al (TDAI) profiles versus potential density from the PANDORA
922 cruise. Dark blue circles show measured TDAI concentrations. Orange line represents average
923 inflow TDAI profile, while grey shading represents the range of inflow concentrations. Green
924 box represents the density interval over which budget of the thermocline waters is calculated. **(a–**
925 **c)** waters that are located south of the Solomon Sea; **(d–f)** profiles found along the NGCU; **(g–i)**
926 profiles that are at the exit straits of the Solomon Sea; **(j–l)** located outside the Solomon Sea.

927 **Fig. 4.** Dissolved Mn (DMn) profiles versus potential density from the PANDORA cruise. Pink
928 diamonds show measured DMn concentrations. Orange line represents average inflow DMn
929 profile, while grey shading represents the range of inflow concentrations. Green box represents
930 the density interval over which budget of the thermocline waters is calculated. (a–c) waters that
931 are located south of the Solomon Sea; (d–f) profiles found along the NGCU; (g–i) profiles that
932 are at the exit straits of the Solomon Sea; (j–l) located outside the Solomon Sea.

933 **Fig. 5** Total Dissolvable Mn (TDMn) profiles versus potential density from the PANDORA
934 cruise. Red diamonds show measured TDMn concentrations. Orange line represents average
935 inflow TDMn profile, while grey shading represents the range of inflow concentrations. Green
936 box represents the density interval over which the mass balance is calculated. (a–c) waters that
937 are located south of the Solomon Sea; (d–f) profiles found along the NGCU; (g–i) profiles that
938 are at the exit straits of the Solomon Sea; (j–l) located outside the Solomon Sea.

939 **Fig. 6.** Budget for thermocline waters ($24\text{--}26.9 \sigma_\theta$), showing the flux of DA1, TDA1, DMn,
940 TDMn at the inflow, the outflow via currents (NGCU, NICU, and SGU), and the flux out of the
941 EUC at 156°E , as well as the average concentration of the Solomon Sea, calculated from stations
942 most representative of the Solomon Sea. Red dots represent stations used to calculate
943 concentrations at each exit strait over the potential density range of the EUC

944

945 **Supplemental Fig. 1.** Implied labile particle concentrations of aluminum (PA1) and manganese
946 (PMn). Where values are not reported, the dissolved (D) concentration exceeded the total
947 dissolvable (TD) concentration, but within the standard deviation of the measurement. $P = TD -$
948 D.

949 **Supplemental Fig. 2.** Dissolved Al (DAI) profiles versus potential density from the PANDORA
950 cruise. Light blue circles show measured DAI concentrations. Orange line represents average
951 inflow DAI profile, while grey shading represents the range of inflow concentrations. Green box
952 represents the density interval over which budget of the thermocline waters is calculated. (a–c)
953 waters that are located south of the Solomon Sea; (d–f) profiles found along the NGCU; (g–i)
954 profiles that are at the exit straits of the Solomon Sea; (j–l) located outside the Solomon Sea.

955 **Supplemental Fig. 3.** Total Dissolvable Al (TDAI) profiles versus potential density from the
956 PANDORA cruise. Dark blue circles show measured TDAI concentrations. Orange line
957 represents average inflow TDAI profile, while grey shading represents the range of inflow
958 concentrations. Green box represents the density interval over which budget of the thermocline
959 waters is calculated. (a–c) waters that are located south of the Solomon Sea; (d–f) profiles found
960 along the NGCU; (g–i) profiles that are at the exit straits of the Solomon Sea; (j–l) located
961 outside the Solomon Sea.

962 **Supplemental Fig. 4.** Dissolved Mn (DMn) profiles versus potential density from the
963 PANDORA cruise. Pink diamonds show measured DMn concentrations. Orange line represents
964 average inflow DMn profile, while grey shading represents the range of inflow concentrations.
965 Green box represents the density interval over which budget of the thermocline waters is
966 calculated. (a–c) waters that are located south of the Solomon Sea; (d–f) profiles found along the
967 NGCU; (g–i) profiles that are at the exit straits of the Solomon Sea; (j–l) located outside the
968 Solomon Sea.

969 **Supplemental Fig. 5.** Total Dissolvable Mn (TDMn) profiles versus depth from the PANDORA
970 cruise. Red diamonds show measured TDMn concentrations. Orange line represents average
971 inflowing TDMn profile, while grey shading represents the range of inflow concentrations.
972 Green box represents the density interval over which budget of the thermocline waters is

973 calculated. (a–c) waters that are located south of the Solomon Sea; (d–f) profiles found along the
974 NGCU; (g–i) profiles that are at the exit straits of the Solomon Sea; (j–l) located outside the
975 Solomon Sea.

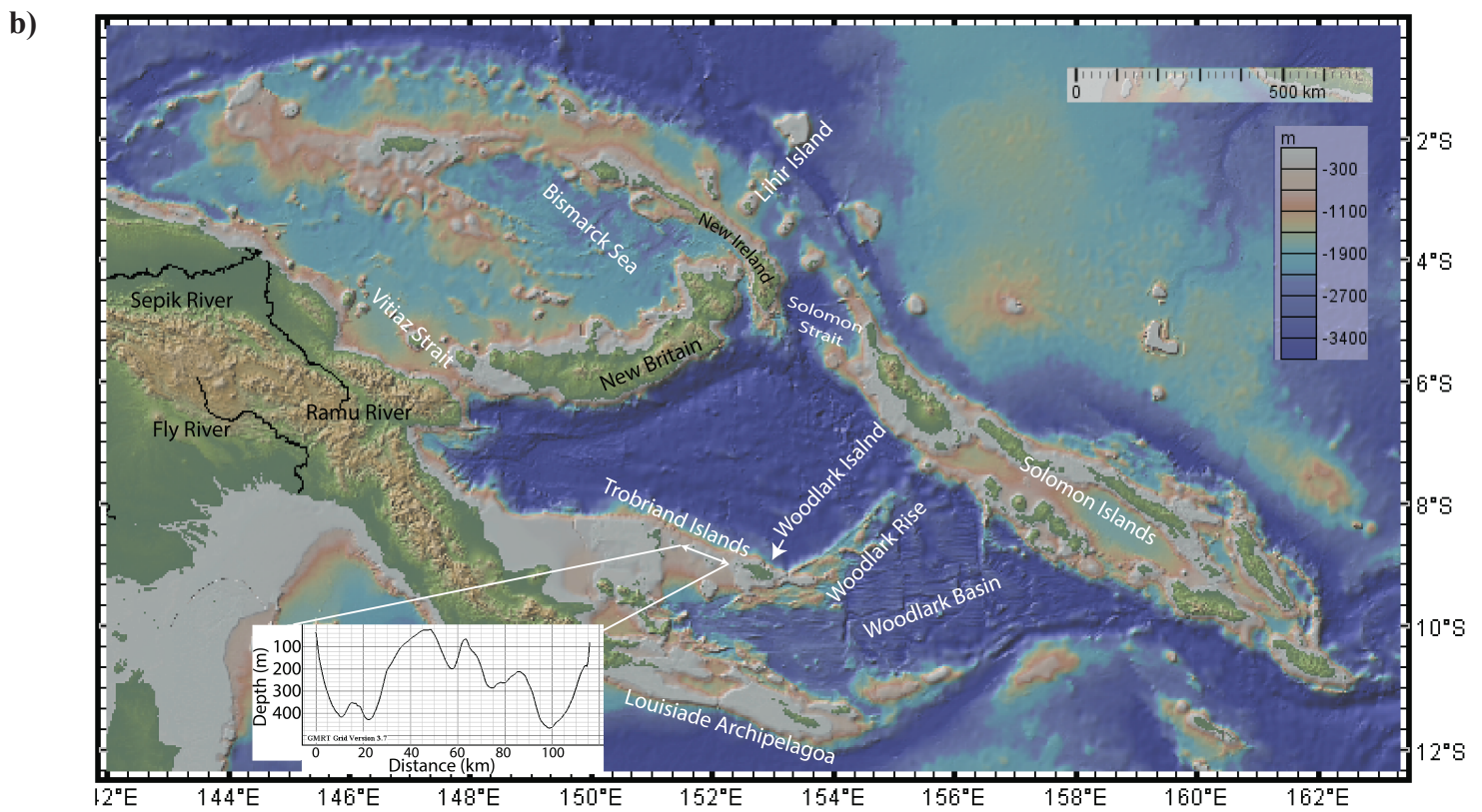
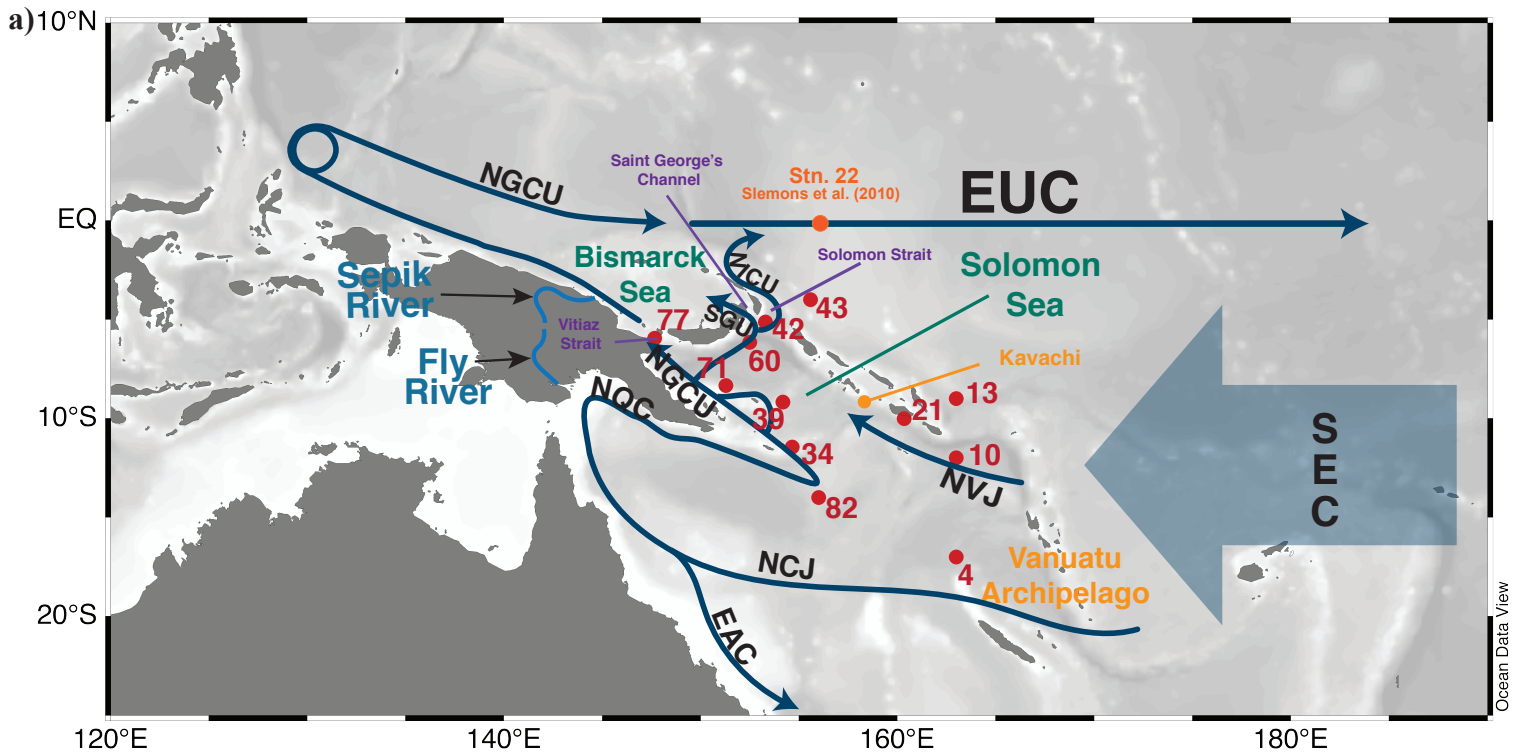


Fig. 1. a) Currents [dark blue; after Germaineaud et al. (2016)], rivers (blue text), straits (purple text), and potential point sources (yellow text) to the Solomon Sea. These currents are: Equatorial Undercurrent (EUC); South Equatorial Current (SEC); New Guinea Coastal Undercurrent (NGCU); New Ireland Coastal Undercurrent (NICU); Saint George's Undercurrent (SGU); North Queensland Current (NQC); New Caledonia Jet (NCJ); North Vanuatu Jet (NVJ); East Australian Current (EAC). Red dots and numbers represent stations sampled for Al and Mn. b) Bathymetry of the Woodlark Basin and Trobriand Islands

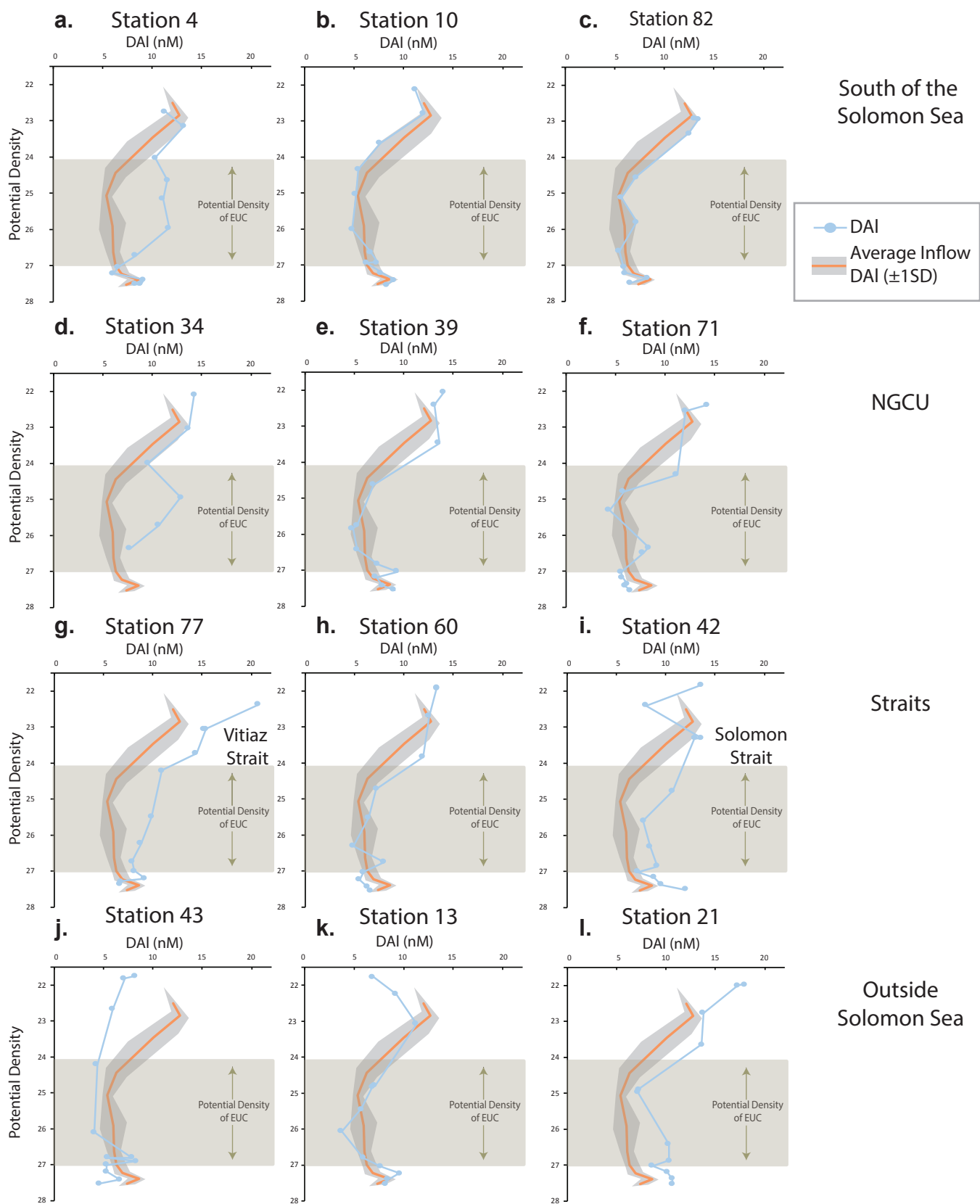


Fig. 2. Dissolved Al (DAI) profiles versus potential density from the Pandora cruise. Light blue circles show measured DAI concentrations. Orange line represents average inflow DAI profile, while grey shading represents the range of inflow concentrations. Green box represents the density interval over which budget of the thermocline waters is calculated. (a–c) waters that are located south of the Solomon Sea; (d–f) profiles found along the NGCU; (g–i) profiles that are at the exit straits of the Solomon Sea; (j–l) located outside the Solomon Sea.

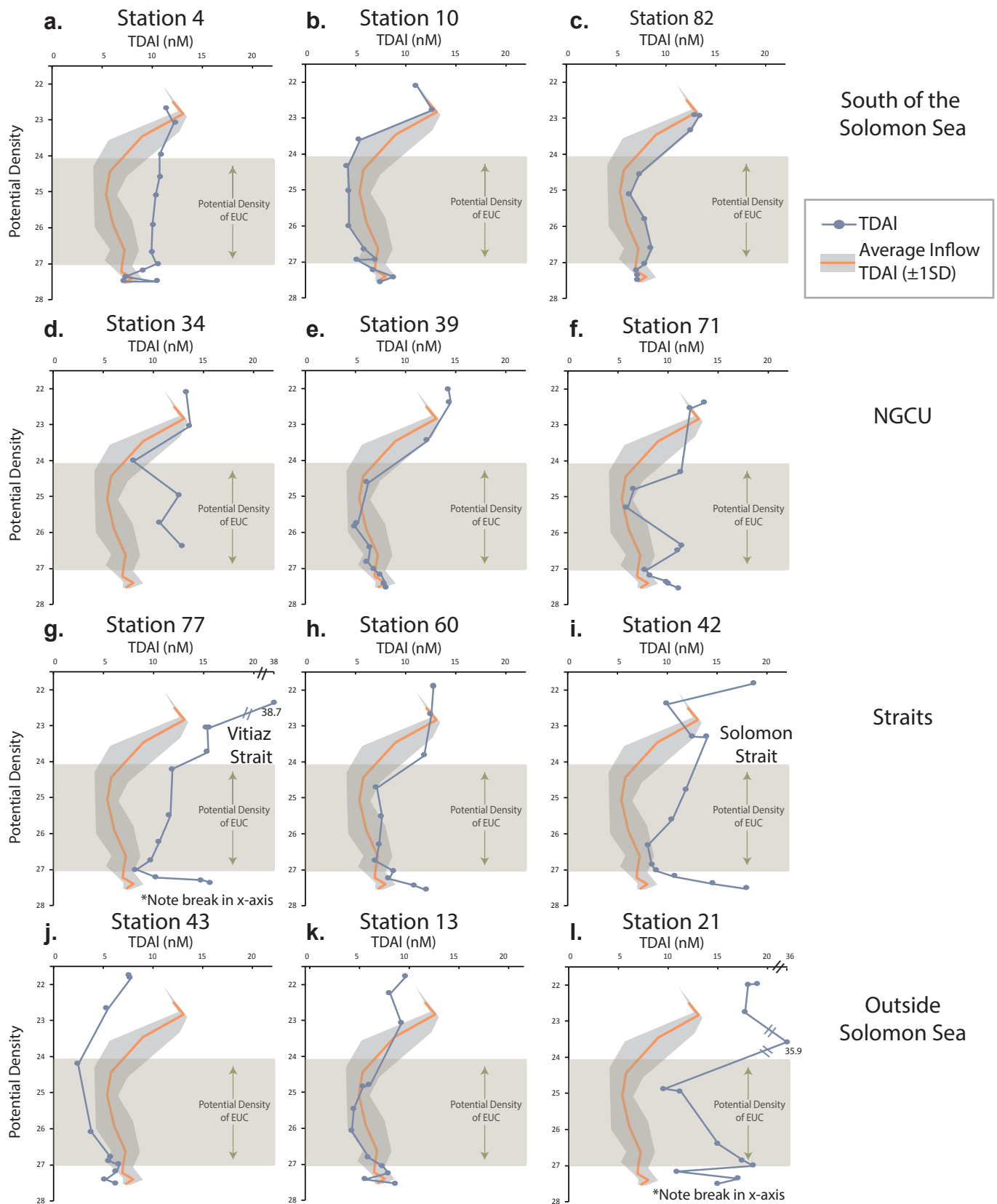


Fig. 3. Total Dissolvable Al (TDAI) profiles versus potential density from the PANDORA cruise. Dark blue circles show measured TDAI concentrations. Orange line represents average inflow TDAI profile, while grey shading represents the range of inflow concentrations. Green box represents the density interval over which budget of the thermocline waters is calculated. (a–c) waters that are located south of the Solomon Sea; (d–f) profiles found along the NGCU; (g–i) profiles that are at the exit straits of the Solomon Sea; (j–l) located outside the Solomon Sea.

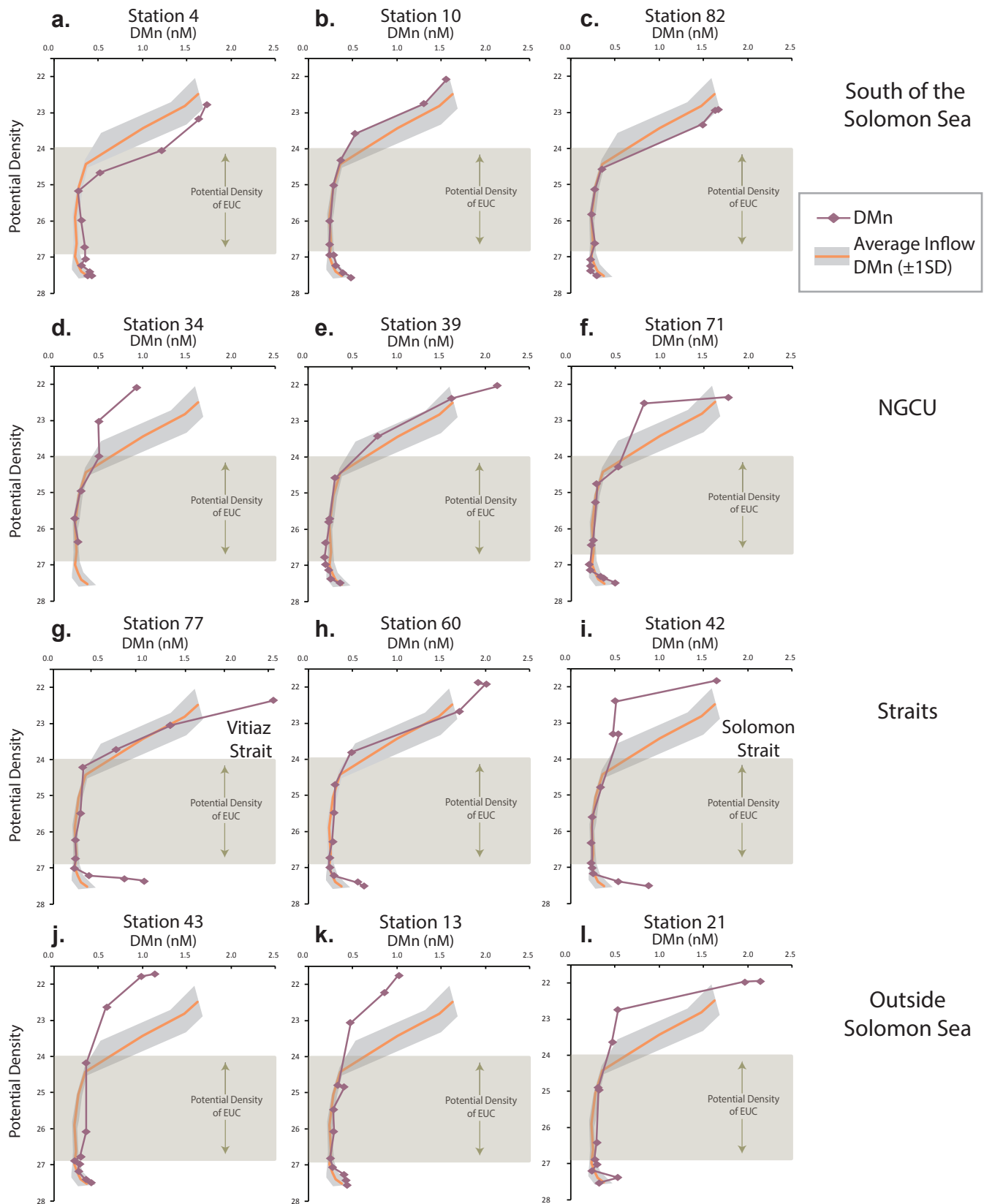


Fig. 4. Dissolved Mn (DMn) profiles versus potential density from the PANDORA cruise. Pink diamonds show measured DMn concentrations. Orange line represents average inflow DMn profile, while grey shading represents the range of inflow concentrations. Green box represents the density interval over which budget of the thermocline waters is calculated. (a–c) waters that are located south of the Solomon Sea; (d–f) profiles found along the NGCU; (g–i) profiles that are at the exit straits of the Solomon Sea; (j–l) located outside the Solomon Sea.

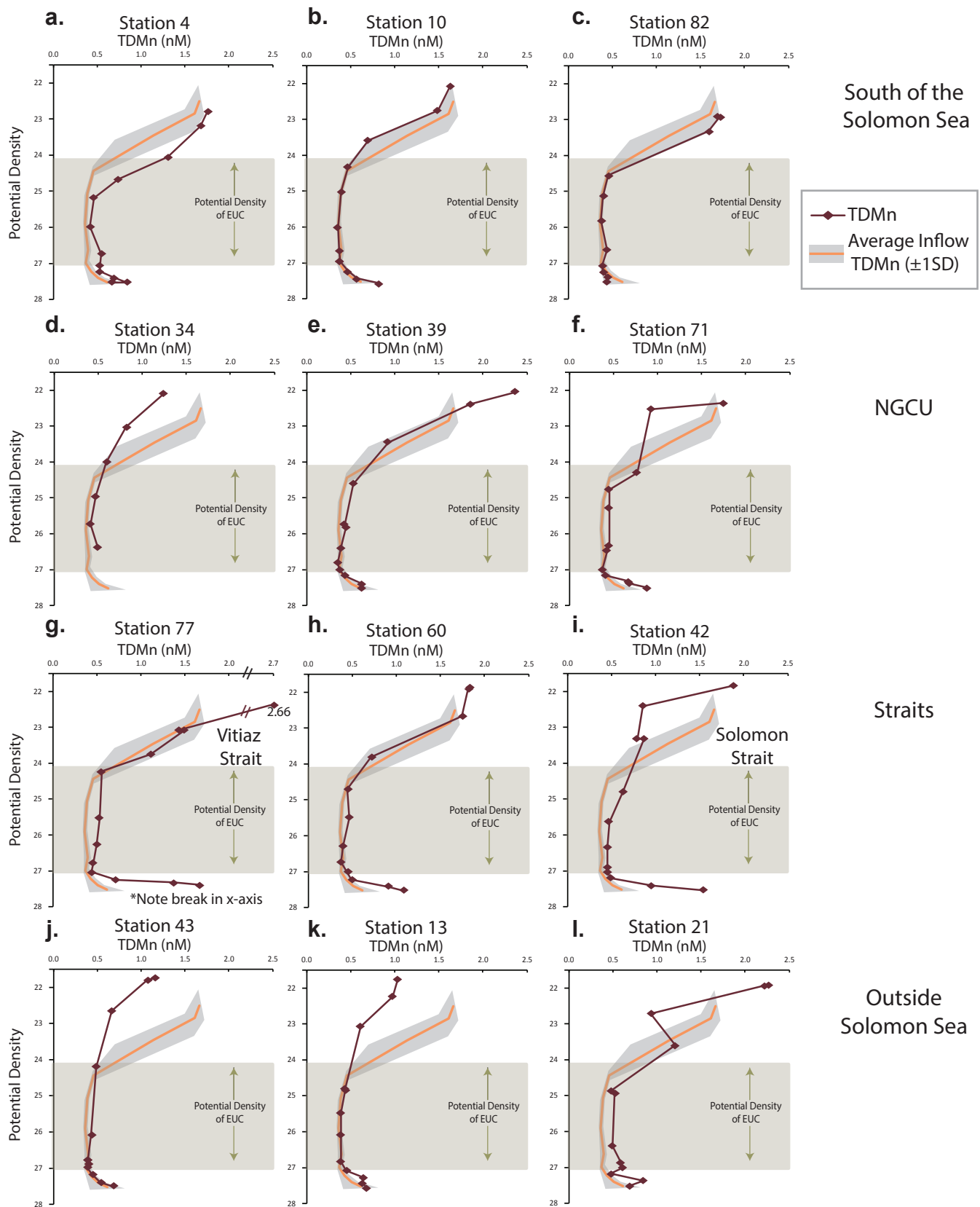


Fig. 5. Total Dissolvable Mn (TDMn) profiles versus potential density from the PANDORA cruise. Red diamonds show measured TDMn concentrations. Orange line represents average inflow TDMn profile, while grey shading represents the range of inflow concentrations. Green box represents the density interval over which budget of the thermocline waters is calculated. (a–c) waters that are located south of the Solomon Sea; (d–f) profiles found along the NGCU; (g–i) profiles that are at the exit straits of the Solomon Sea; (j–l) located outside the Solomon Sea.

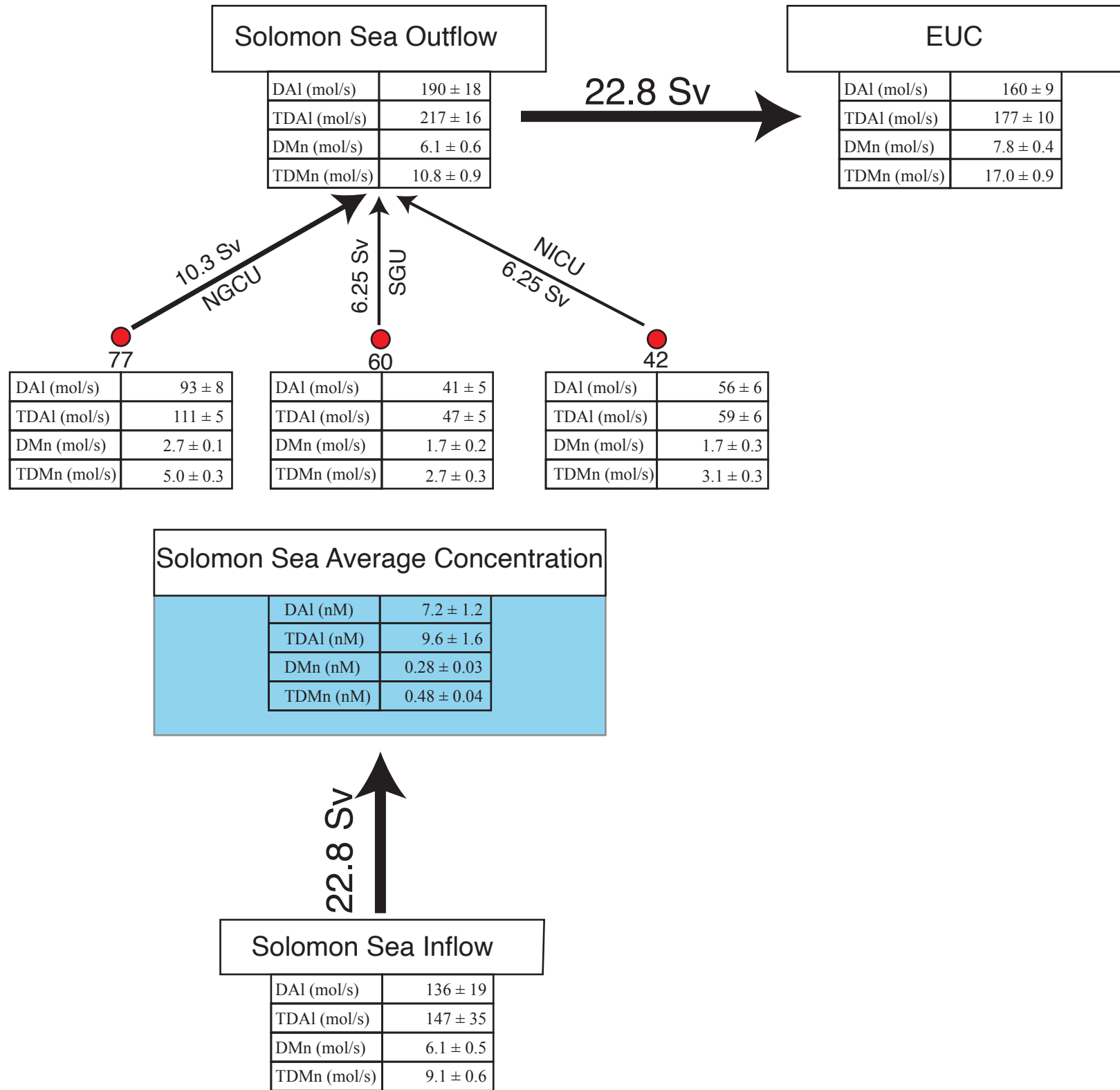
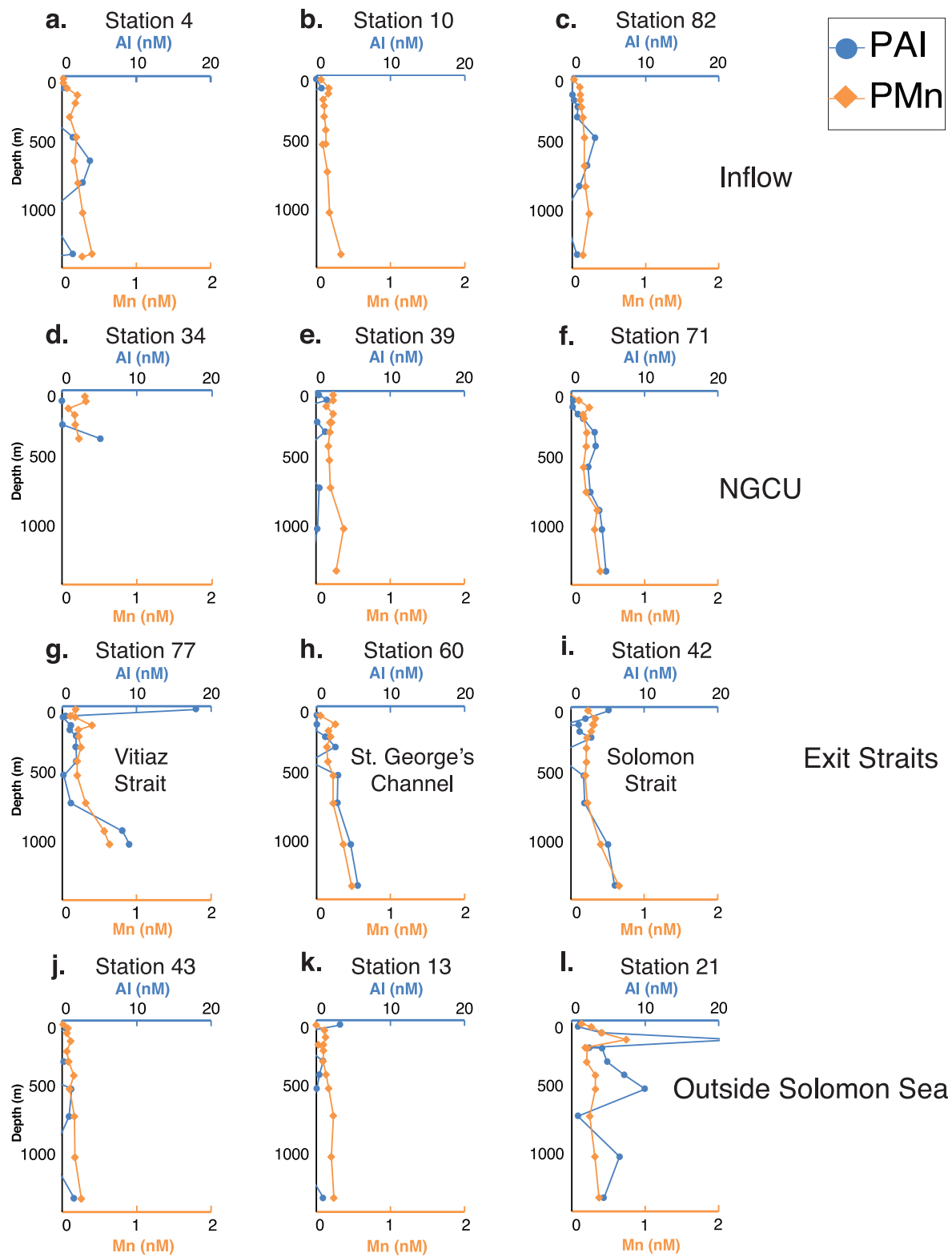
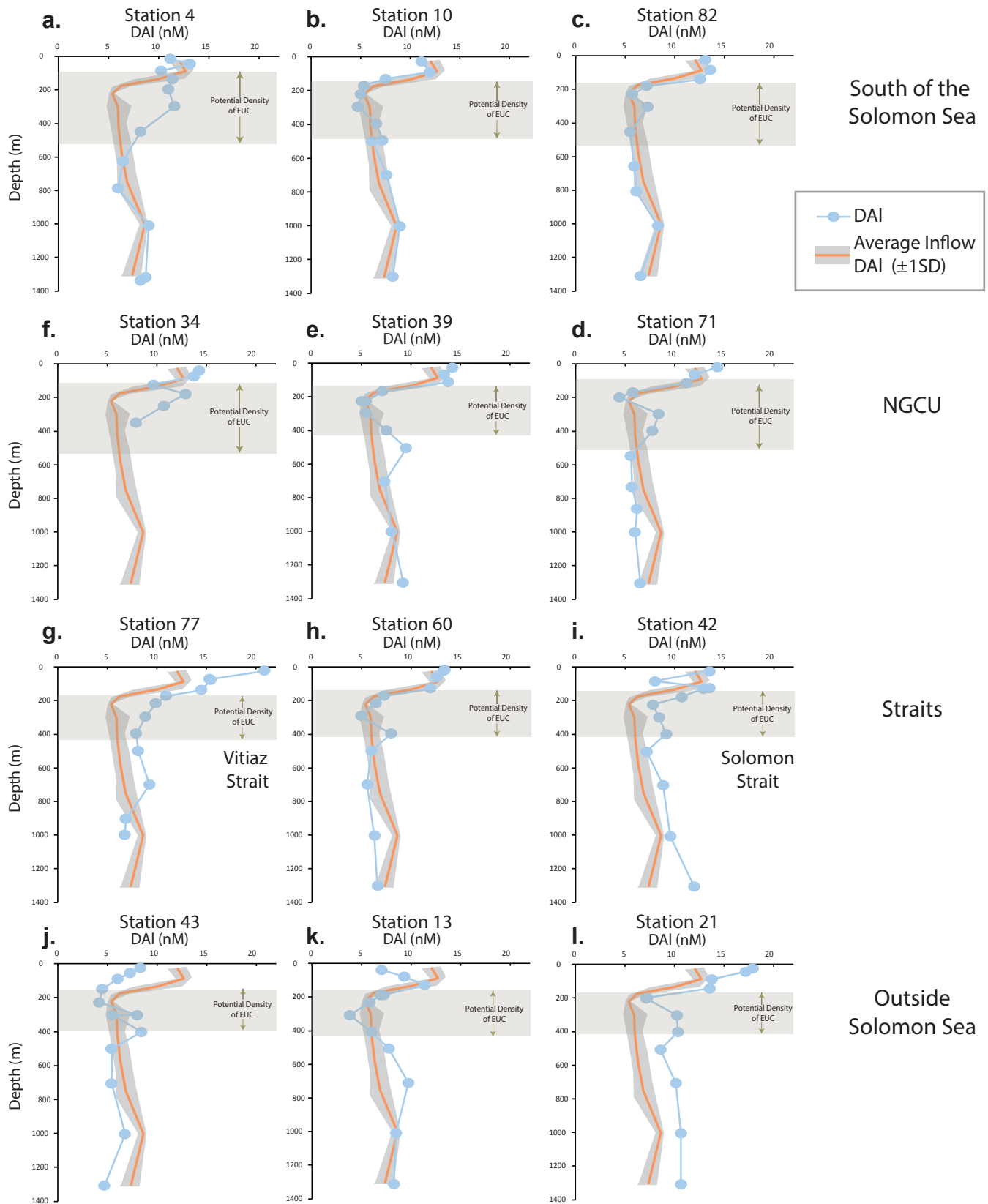


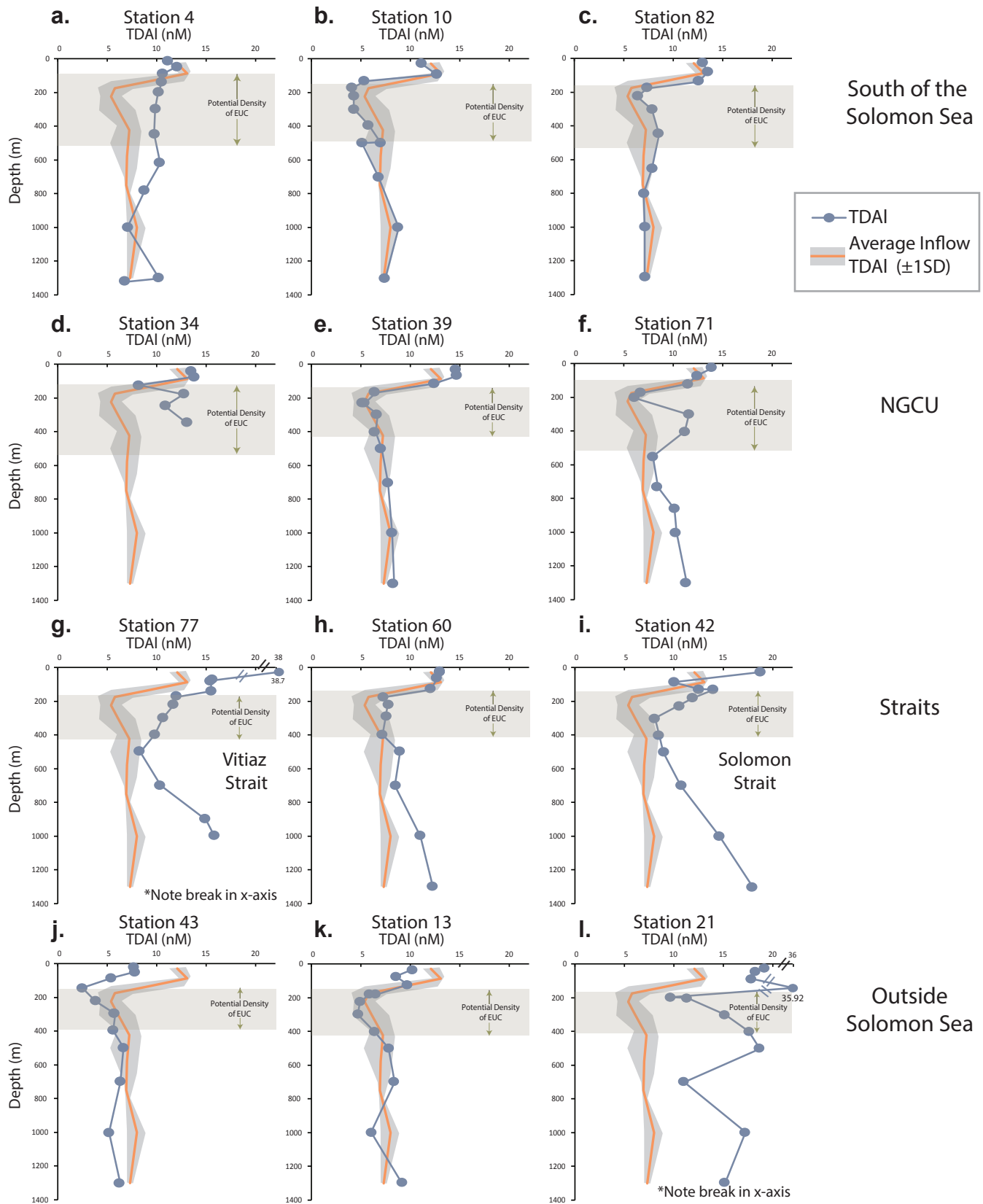
Fig. 6. Budget for 24-26.9 σ_θ , showing the flux of DAI, TDAI, DMn, TDMn at the inflow, the outflow via currents (NGCU, NICU, and SGU), and the flux out of the EUC at 156°E, as well as the average concentration of the Solomon Sea, calculated from stations most representative of the Solomon Sea. Red dots represent stations used to calculate concentrations at each exit strait over the potential density range of the EUC.



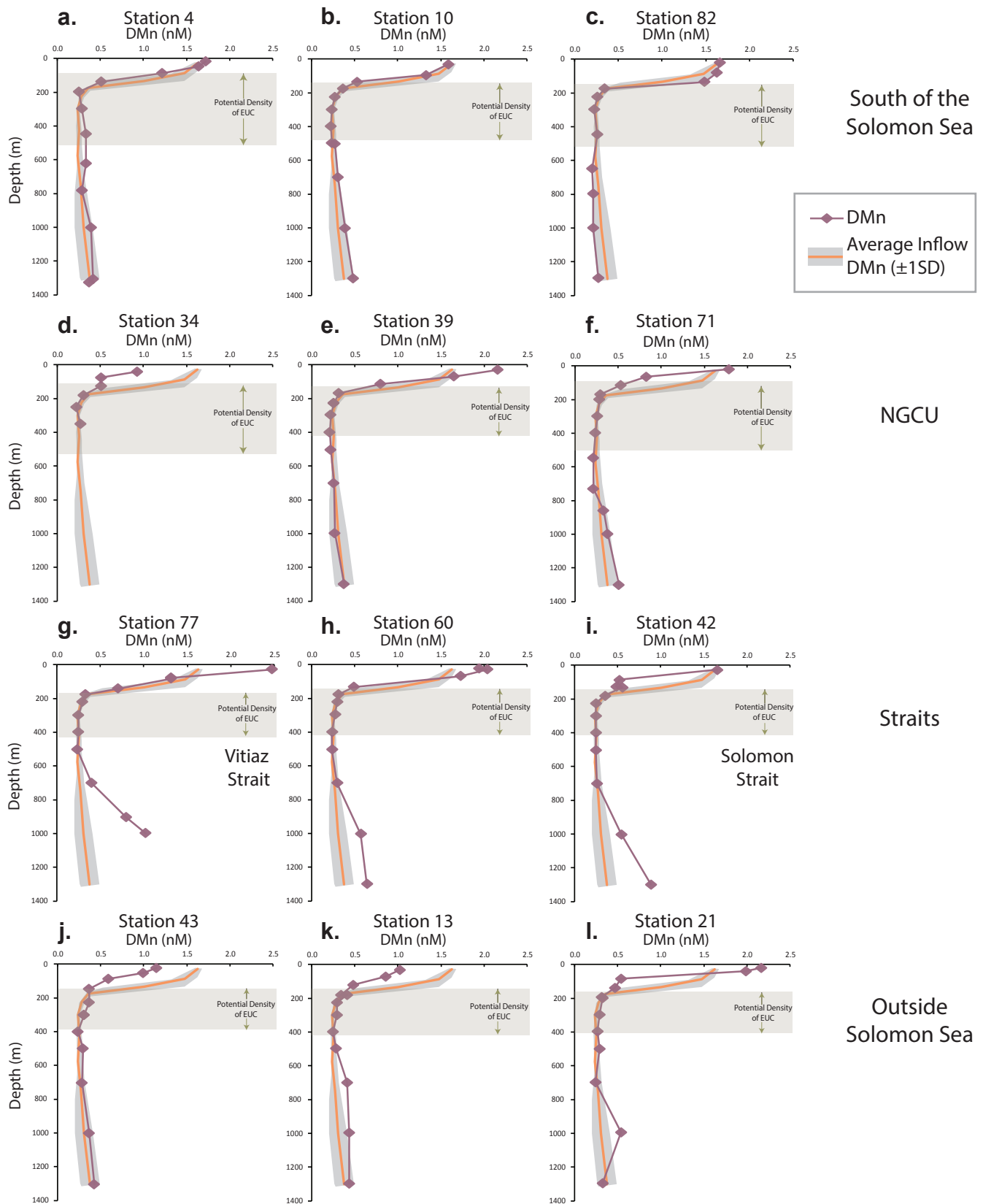
Supplemental Fig. 1. Implied labile particle concentrations of aluminum (PAI) and manganese (PMn). Where values are not reported, the dissolved (D) concentration exceeded the total dissolvable (TD) concentration, but within the standard deviation of the measurement. $P = TD - D$.



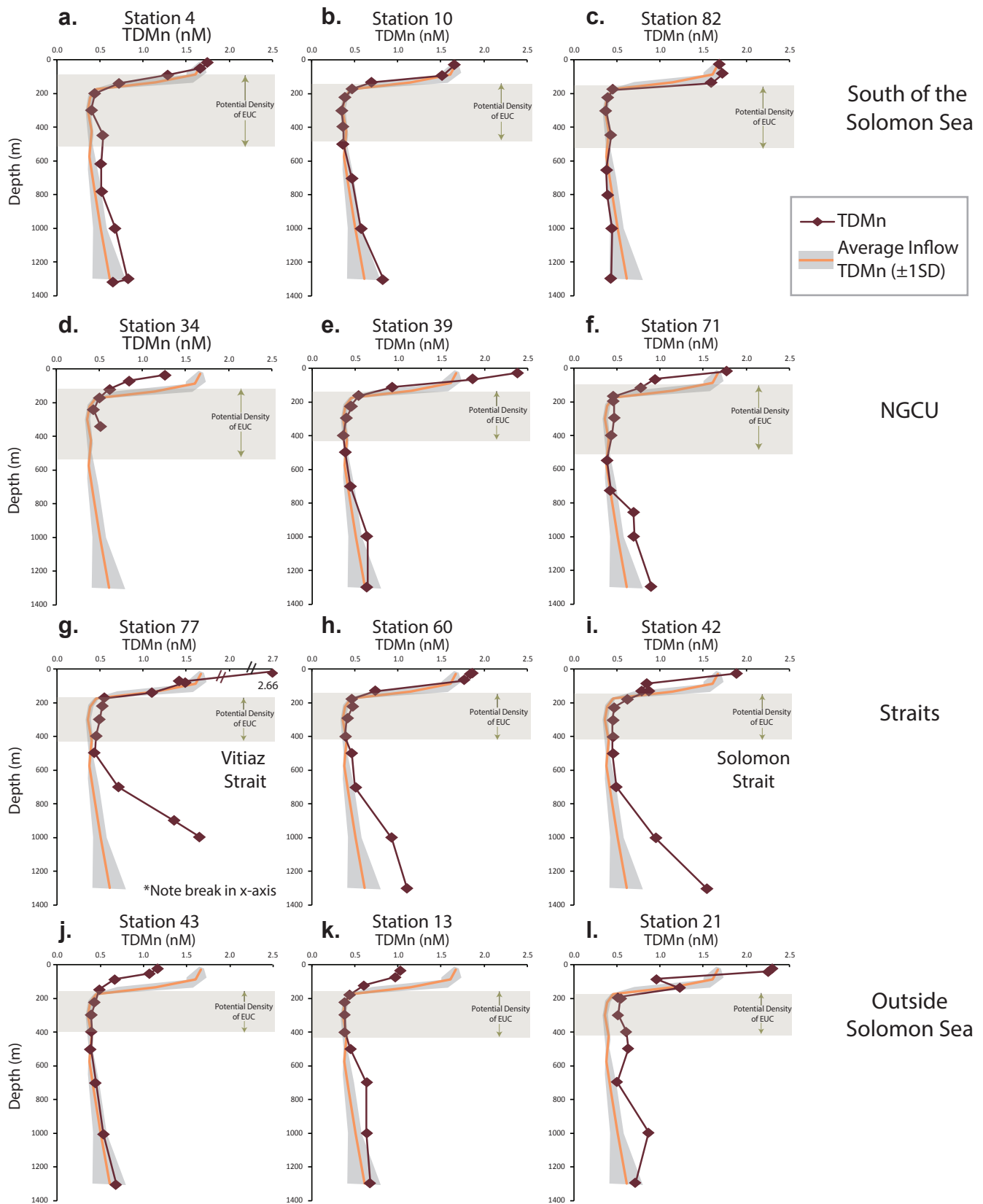
Supplemental Fig. 2. Dissolved Al (DAI) profiles versus depth from the PANDORA cruise. Light blue circles show measured DAI concentrations. Orange line represents average inflow DAI profile, while grey shading represents the range of inflow concentrations. Green box represents the density interval over which budget of the thermocline waters is calculated. (a–c) waters that are located south of the Solomon Sea; (d–f) profiles found along the NGCU; (g–i) profiles that are at the exit straits of the Solomon Sea; (j–l) located outside the Solomon Sea.



Supplemental Fig. 3. Total Dissolvable Al (TDAI) profiles versus depth from the PANDORA cruise. Blue circles show measured TDAI concentrations. Orange line represents average inflow TDAI profile, while grey shading represents the range of inflow concentrations. Green box represents the density interval over which budget of the thermocline waters is calculated. (a–c) waters that are located south of the Solomon Sea; (d–f) profiles found along the NGCU; (g–i) profiles that are at the exit straits of the Solomon Sea; (j–l) located outside the Solomon Sea.



Supplemental Fig. 4. Dissolved Mn (DMn) profiles versus potential density from the PANDORA cruise. Pink diamonds show measured DMn concentrations. Orange line represents average inflow DMn profile, while grey shading represents the range of inflow concentrations. Green box represents the density interval over which budget of the thermocline waters is calculated. (a–c) waters that are located south of the Solomon Sea; (d–f) profiles found along the NGCU; (g–i) profiles that are at the exit straits of the Solomon Sea; (j–l) located outside the Solomon Sea.



Supplemental Fig. 5. Total Dissolvable Mn (TDMn) profiles versus depth from the PANDORA cruise. Red diamonds show measured TDMn concentrations. Orange line represents average inflow TDMn profile, while grey shading represents the range of inflow concentrations. Green box represents the density interval over which budget of the thermocline waters is calculated. (a–c) waters that are located south of the Solomon Sea; (d–f) profiles found along the NGCU; (g–i) profiles that are at the exit straits of the Solomon Sea; (j–l) located outside the Solomon Sea.

Table S1: Potential density thresholds and corresponding depths used to calculate average concentrations over each depth range.

Station	Surface (m)	24 σ_θ (m)	26.9 σ_θ (m)	Deep ($\sim 27.5\sigma_\theta$) (m)
4	18	87	525	1321
10	31	150	490	1303
13	34	152	430	1299
21	26	170	415	1302
34	47	130	*	
39	26	137	430	1299
42	25	150	420	1301
43	25	152	405	1302
60	23	145	420	1300
71	23	105	515	1300
77	25	173	435	999
82	24	158	530	1298

*Profile collected only to 353 m (26.4 σ_θ)

Table S2: PANDORA Al and Mn dataset. Error reported is 1 standard deviation of the measurement

Latitude	Longitude	STATION	Depth (m)	DAI (nM)	DAI_1SD_(nM)	TDAI_nM	TDAI_1SD_nM	DMn (nM)	DMn_1SD_(nM)	TDMn (nM)	TDMn_1SD_(nM)
-17.00	163.00	4	18	11.2	0.4	11.2	0.4	1.72	0.07	1.74	0.07
-17.00	163.00	4	50	13.1	0.4	12.1	0.4	1.63	0.07	1.66	0.07
-17.00	163.00	4	88	10.3	0.4	10.7	0.4	1.21	0.05	1.28	0.05
-17.00	163.00	4	139	11.5	0.4	10.6	0.4	0.51	0.02	0.72	0.03
-17.00	163.00	4	198	11.0	0.4	10.2	0.4	0.26	0.00	0.44	0.02
-17.00	163.00	4	298	11.6	0.4	9.9	0.4	0.29	0.00	0.40	0.02
-17.00	163.00	4	448	8.2	0.4	9.8	0.4	0.33	0.00	0.53	0.02
-17.00	163.00	4	619	6.5	0.4	10.3	0.4	0.34	0.00	0.51	0.02
-17.00	163.00	4	780	6.0	0.4	8.8	0.4	0.29	0.00	0.51	0.02
-17.00	163.00	4	1000	9.0	0.4	7.1	0.4	0.39	0.02	0.67	0.03
-17.00	163.00	4	1300	8.7	0.4	10.3	0.4	0.42	0.02	0.82	0.03
-17.00	163.00	4	1321	8.2	0.4	6.9	0.4	0.37	0.00	0.65	0.03
-12.00	163.00	10	31	11.1	0.4	11.1	0.4	1.59	0.06	1.65	0.07

-12.00	163.00	10	95	11.9	0.4	12.7	0.4	1.33	0.05	1.50	0.06
-12.00	163.00	10	135	7.6	0.4	5.4	0.4	0.53	0.02	0.70	0.03
-12.00	163.00	10	175	5.4	0.4	4.2	0.4	0.37	0.02	0.47	0.02
-12.00	163.00	10	224	5.1	0.4	4.3	0.4	0.28	0.00	0.39	0.00
-12.00	163.00	10	301	4.7	0.4	4.3	0.4	0.24	0.00	0.35	0.00
-12.00	163.00	10	399	6.6	0.4	5.9	0.4	0.24	0.00	0.37	0.00
-12.00	163.00	10	500	7.2	0.4	7.0	0.4	0.24	0.00	0.37	0.01
-12.00	163.00	10	501	6.2	0.4	5.1	0.4	0.28	0.00	0.37	0.01
-12.00	163.00	10	700	7.6	0.4	6.8	0.4	0.32	0.00	0.47	0.02
-12.00	163.00	10	1002	8.9	0.4	8.8	0.4	0.39	0.00	0.57	0.02
-12.00	163.00	10	1303	8.3	0.4	7.5	0.4	0.49	0.02	0.83	0.03
-9.00	163.00	13	34	6.9	0.5	10.2	0.5	1.02	0.04	1.02	0.04
-9.00	163.00	13	74	9.3	0.5	8.6	0.5	0.85	0.03	0.96	0.04
-9.00	163.00	13	125	11.3	0.5	9.7	0.5	0.47	0.02	0.60	0.02
-9.00	163.00	13	179	7.1	0.5	6.5	0.5	0.33	0.05	0.43	0.02
-9.00	163.00	13	181	6.9	0.5	5.9	0.5	0.40	0.05	0.44	0.02
-9.00	163.00	13	225	5.8	0.5	4.9	0.5	0.28	0.05	0.38	0.05
-9.00	163.00	13	298	3.8	0.5	4.7	0.5	0.29	0.05	0.38	0.05
-9.00	163.00	13	400	6.0	0.5	6.4	0.5	0.25	0.05	0.38	0.05
-9.00	163.00	13	501	7.7	0.5	7.8	0.5	0.28	0.05	0.44	0.02
-9.00	163.00	13	700	9.6	0.5	8.4	0.5	0.40	0.05	0.63	0.03
-9.00	163.00	13	1000	8.5	0.5	6.1	0.5	0.42	0.02	0.63	0.03
-9.00	163.00	13	1299	8.2	0.5	9.2	0.5	0.43	0.02	0.67	0.03
-10.01	160.36	21	26	17.9	0.4	19.1	0.4	2.15	0.09	2.28	0.09
-10.01	160.36	21	45	17.2	0.4	18.1	0.4	1.97	0.08	2.24	0.09
-10.01	160.36	21	89	13.8	0.4	17.8	0.4	0.54	0.02	0.95	0.04
-10.01	160.36	21	140	13.6	0.4	35.9	0.4	0.48	0.02	1.22	0.05
-10.01	160.36	21	198	7.3	0.4	9.6	0.5	0.32	0.05	0.50	0.02
-10.01	160.36	21	202	7.2	0.4	11.3	0.4	0.33	0.05	0.54	0.02
-10.01	160.36	21	303	10.3	0.4	15.1	0.5	0.30	0.02	0.51	0.02
-10.01	160.36	21	401	10.3	0.4	17.5	0.4	0.28	0.05	0.60	0.02
-10.01	160.36	21	504	8.6	0.4	18.6	0.6	0.30	0.02	0.62	0.02
-10.01	160.36	21	700	10.1	0.4	10.9	0.5	0.25	0.05	0.49	0.02
-10.01	160.36	21	1002	10.7	0.4	17.1	0.5	0.55	0.02	0.86	0.03
-10.01	160.36	21	1302	10.7	0.4	15.1	0.5	0.34	0.02	0.71	0.03
-11.45	154.67	34	47	14.3	0.3	13.4	0.3	0.93	0.18	1.24	0.05
-11.45	154.67	34	79	13.7	0.3	13.8	0.3	0.51	0.02	0.83	0.03
-11.45	154.67	34	129	9.6	0.3	8.2	0.3	0.51	0.02	0.60	0.02
-11.45	154.67	34	181	12.9	0.3	12.7	0.3	0.31	0.02	0.48	0.02
-11.45	154.67	34	249	10.7	0.3	10.8	0.3	0.24	0.02	0.42	0.02

-11.45	154.67	34	353	7.8	0.3	13.0	0.3	0.28	0.02	0.50	0.02
-9.17	154.19	39	26	14.2	0.3	14.6	0.3	2.16	0.09	2.38	0.10
-9.17	154.19	39	63	13.3	0.3	14.7	0.3	1.64	0.07	1.87	0.07
-9.17	154.19	39	111	13.7	0.3	12.4	0.3	0.80	0.03	0.93	0.04
-9.19	154.17	39	163	7.1	0.3	6.4	0.3	0.31	0.02	0.54	0.02
-9.19	154.17	39	221	5.4	0.3	5.4	0.3	0.25	0.02	0.44	0.02
-9.19	154.17	39	227	5.0	0.3	5.1	0.3	0.25	0.02	0.46	0.02
-9.19	154.17	39	295	5.5	0.3	6.7	0.3	0.21	0.02	0.40	0.02
-9.19	154.17	39	397	7.6	0.3	6.4	0.3	0.20	0.02	0.36	0.02
-9.19	154.17	39	500	9.5	0.3	7.1	0.3	0.21	0.02	0.39	0.02
-9.19	154.17	39	700	7.3	0.3	7.7	0.3	0.25	0.02	0.44	0.02
-9.19	154.17	39	998	8.0	0.3	8.2	0.3	0.26	0.02	0.64	0.03
-9.19	154.17	39	1299	9.2	0.3	8.3	0.3	0.37	0.02	0.64	0.03
-5.14	153.30	42	25	13.5	0.4	18.7	0.4	1.64	0.07	1.88	0.08
-5.14	153.30	42	83	8.0	0.4	9.9	0.4	0.51	0.02	0.84	0.03
-5.14	153.30	42	129	13.5	0.3	12.4	0.4	0.48	0.02	0.78	0.03
-5.14	153.30	42	130	12.9	0.4	13.9	0.4	0.55	0.02	0.86	0.03
-5.15	153.29	42	181	10.6	0.4	11.9	0.4	0.35	0.02	0.62	0.02
-5.15	153.29	42	227	7.7	0.4	10.5	0.4	0.24	0.02	0.46	0.02
-5.15	153.29	42	297	8.4	0.4	8.0	0.4	0.24	0.02	0.45	0.02
-5.12	153.33	42	401	9.1	0.3	8.4	0.4	0.24	0.02	0.45	0.02
-5.12	153.33	42	501	7.1	0.3	8.9	0.4	0.25	0.02	0.45	0.02
-5.12	153.33	42	700	8.8	0.3	10.7	0.4	0.26	0.02	0.49	0.02
-5.12	153.33	42	1000	9.5	0.3	14.6	0.4	0.54	0.02	0.94	0.04
-5.12	153.33	42	1301	11.9	0.3	17.9	0.4	0.88	0.04	1.53	0.06
-4.00	155.59	43	25	8.3	0.4	7.7	0.5	1.15	0.05	1.16	0.05
-4.00	155.59	43	56	7.2	0.4	7.9	0.5	0.99	0.04	1.07	0.04
-4.00	155.59	43	91	6.0	0.4	5.5	0.5	0.60	0.02	0.66	0.03
-4.00	155.59	43	160	4.5	0.5	2.5	0.3	0.37	0.01	0.49	0.02
-4.00	155.59	43	224	4.2	0.4	3.9	0.5	0.37	0.08	0.44	0.02
-4.00	155.59	43	300	8.0	0.4	5.8	0.5	0.31	0.08	0.39	0.02
-4.00	155.59	43	300	5.6	0.4	5.8	0.5	0.30	0.08	0.39	0.02
-4.00	155.59	43	401	8.4	0.5	5.6	0.3	0.24	0.01	0.40	0.02
-4.00	155.59	43	500	5.4	0.5	6.7	0.3	0.29	0.01	0.39	0.02
-4.00	155.59	43	701	5.5	0.5	6.4	0.3	0.28	0.01	0.45	0.02
-4.00	155.59	43	1002	6.7	0.5	5.3	0.3	0.37	0.01	0.54	0.02
-4.00	155.59	43	1302	4.7	0.5	6.4	0.3	0.42	0.02	0.68	0.03
-6.17	152.50	60	23	13.3	0.3	12.9	0.3	1.92	0.08	1.85	0.07
-6.17	152.50	60	29	13.3	0.3	12.9	0.3	2.01	0.08	1.83	0.07
-6.17	152.50	60	65	12.5	0.3	12.6	0.3	1.71	0.07	1.76	0.07

-6.17	152.50	60	127	11.9	0.3	12.0	0.3	0.48	0.02	0.73	0.03
-6.17	152.50	60	176	7.2	0.3	7.2	0.3	0.30	0.01	0.46	0.02
-6.17	152.50	60	220	6.4	0.3	7.7	0.3	0.29	0.01	0.48	0.02
-6.17	152.50	60	293	4.9	0.3	7.4	0.3	0.27	0.01	0.41	0.02
-6.17	152.50	60	401	8.0	0.3	7.1	0.3	0.23	0.01	0.38	0.02
-6.17	152.50	60	499	5.9	0.3	8.8	0.3	0.23	0.01	0.46	0.02
-6.17	152.50	60	699	5.5	0.3	8.4	0.3	0.28	0.01	0.50	0.02
-6.17	152.50	60	1000	6.3	0.3	10.9	0.3	0.56	0.02	0.92	0.04
-6.17	152.50	60	1300	6.6	0.3	12.2	0.3	0.63	0.03	1.10	0.04
-8.34	151.29	71	23	14.3	0.6	13.9	0.6	1.79	0.07	1.77	0.07
-8.34	151.29	71	69	12.0	0.6	12.3	0.6	0.84	0.03	0.94	0.04
-8.34	151.29	71	118	11.2	0.6	11.4	0.6	0.54	0.02	0.78	0.03
-8.34	151.29	71	168	5.8	0.6	6.7	0.6	0.30	0.01	0.46	0.02
-8.34	151.29	71	201	4.3	0.6	6.0	0.6	0.28	0.01	0.46	0.02
-8.33	151.29	71	299	8.3	0.6	11.5	0.6	0.26	0.01	0.46	0.02
-8.33	151.29	71	400	7.7	0.6	11.1	0.6	0.24	0.01	0.44	0.02
-8.33	151.29	71	551	5.5	0.6	7.9	0.6	0.21	0.01	0.38	0.02
-8.33	151.29	71	730	5.7	0.6	8.3	0.6	0.22	0.01	0.42	0.02
-8.33	151.29	71	860	6.2	0.6	10.0	0.6	0.34	0.01	0.69	0.03
-8.33	151.29	71	1000	6.0	0.6	10.2	0.6	0.38	0.01	0.70	0.03
-8.33	151.29	71	1300	6.5	0.6	11.2	0.6	0.51	0.02	0.90	0.04
-5.95	147.67	77	25	20.7	0.4	38.7	0.4	2.48	0.10	2.66	0.11
-5.95	147.67	77	75	15.3	0.4	15.7	0.4	1.31	0.05	1.42	0.06
-5.95	147.67	77	81	15.4	0.4	15.5	0.4	1.32	0.05	1.49	0.06
-5.95	147.66	77	142	14.4	0.4	15.6	0.4	0.70	0.03	1.10	0.04
-5.95	147.66	77	174	11.0	0.4	12.0	0.4	0.33	0.01	0.54	0.02
-5.95	147.66	77	220	9.9	0.4	11.8	0.4	0.30	0.01	0.52	0.02
-5.95	147.66	77	299	8.9	0.4	10.7	0.4	0.24	0.01	0.49	0.02
-5.95	147.66	77	398	8.0	0.4	9.9	0.4	0.25	0.01	0.45	0.02
-5.95	147.66	77	501	8.2	0.4	8.4	0.4	0.23	0.01	0.43	0.02
-5.95	147.66	77	702	9.3	0.4	10.4	0.4	0.40	0.01	0.71	0.03
-5.95	147.66	77	902	6.9	0.4	15.0	0.4	0.80	0.03	1.36	0.05
-5.95	147.66	77	999	6.8	0.4	15.8	0.4	1.02	0.04	1.65	0.07
-14.00	156.01	82	24	13.0	0.4	13.0	0.4	1.66	0.07	1.68	0.07
-14.00	156.01	82	80	13.5	0.4	13.4	0.4	1.63	0.07	1.72	0.07
-14.00	156.01	82	135	12.5	0.4	12.5	0.4	1.48	0.06	1.59	0.06
-14.00	156.01	82	174	7.1	0.4	7.3	0.4	0.34	0.01	0.45	0.02
-14.00	156.01	82	224	5.6	0.4	6.4	0.4	0.27	0.01	0.39	0.02
-14.00	156.01	82	300	7.2	0.4	7.8	0.4	0.23	0.01	0.37	0.01
-14.00	156.01	82	447	5.4	0.4	8.5	0.4	0.27	0.01	0.43	0.02

-14.00	156.01	82	651	5.9	0.4	7.9	0.4	0.21	0.01	0.38	0.02
-14.00	156.01	82	799	6.1	0.4	7.0	0.4	0.22	0.01	0.39	0.02
-14.00	156.01	82	1001	8.3	0.4	7.1	0.4	0.22	0.01	0.44	0.02
-14.00	156.01	82	1298	6.5	0.4	7.1	0.4	0.28	0.01	0.43	0.02