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## **Quantifying and simulating carbon and nitrogen**

## **mineralization from diverse exogenous organic matters**

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- **Running Title**: C and N mineralization from EOM
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#### **Abstract**

 The potential contributions of exogenous organic matters (EOMs) to soil organic C and mineral N supply depend on their C and N mineralization, which can be assessed in laboratory incubations. Such incubations are essential to calibrate decomposition models, because not all EOMs can be tested in the field. However, EOM incubations are resource- intensive. Therefore, easily measurable EOM characteristics that can be useful to predict EOM behavior are needed.

 We quantified C and N mineralization during the incubation of 663 EOMs from five groups (animal manures, composts, sewage sludges, digestates, and others). This represents one of the largest and diversified set of EOM incubations. The C and N mineralization varied widely between and within EOM subgroups. We simulated C and N mineralization with a simple generic decomposition model. Three calibration methods were compared. Individual EOM calibration of the model yielded good model performances, while the use of a unique parameter set per EOM subgroup decreased the model performance, and the use of two EOM characteristics to estimate model parameters gave an intermediate model performance (average RMSE-C values of 32, 99 and 35 65 mg C  $g^{-1}$  added C and average RMSE-N values of 50, 126 and 110 mg N  $g^{-1}$  added N, respectively).

 Because of the EOM variability, individual EOM calibration based on incubation remains the recommended method for predicting most accurately the C and N mineralization of EOMs. However, the two alternative calibration methods are sufficient for the simulation of EOMs without incubation data to obtain reasonable model performances.

 **Keywords:** organic amendment, fertilizer, model, organic matter, soil, decomposition, N mineralization

#### **Highlights:**

- C and N mineralization in 663 exogenous organic matters was quantified under controlled conditions
- C and N mineralization varied widely between and among the subgroups of exogenous
- organic matters
- A simple generic model can predict the variability in C and N mineralization from EOMs
- A calibration was proposed for 26 EOM subgroups using their biochemical characteristics.

#### **1 Introduction**

 Exogenous organic matter (EOM) is characterized as various residual organic matters applied to soil as organic fertilizers or organic amendments. They encompass animal manures and urban and industrial organic wastes and can be used raw or after treatments (e.g., composting or anaerobic digestion). The contribution of EOM carbon (C) to soil organic carbon (SOC) and EOM nitrogen (N) to the mineral N supply available to crops has been extensively studied, especially in long-term field experiments (Gerzabek et al., 1997; Gómez-Muñoz et al., 2017). A better characterization of the large diversity of EOM is however needed to better quantify their potential for soil carbon storage, especially EOMs that have been submitted to treatments such as composting or anaerobic digestion (Chenu et al., 2019). Due to the limited number of EOMs, cropping systems, and climate and soil conditions that can be tested *in situ*, many authors have used EOM incubations to quantify the C and N mineralization of EOMs in soil under controlled conditions (Lazicki et al., 2020; Mondini et al., 2017; Noirot-Cosson et al., 2017). The kinetics obtained during laboratory incubations can be used to predict C and N mineralization in field or greenhouse conditions with additional accounting for the effect of environmental factors (soil temperature, water content, etc.) (Delin et al., 2012; Gale et al., 2006; Pinto et al., 2020).

 However, a comparison of the C and N mineralization kinetics of a wide range of common types of EOM, integrating their variability, is still lacking. EOM incubations can also be used to calibrate EOM decomposition modules in soil-crop models (Noirot-Cosson et al., 2016), which allows us to predict the impact of EOM applications under field conditions. Calibrations for many EOMs have yet to be proposed for the decomposition modules of soil-crop models, which often do not consider the diversity of and variability in EOMs. Due to the workload involved in laboratory incubations (several months of incubation are needed), the use of easily available EOM characteristics (e.g. biochemical fractions and C:N ratios) to predict C and N mineralization should be encouraged (Delin et al., 2012; Lashermes et al., 2009, 2010; Lazicki et al., 2020; Morvan & Nicolardot, 2009; Pansu & Thuriès, 2003; Parnaudeau et al., 2004). The use of these characteristics for calibrating decomposition modules in soil-crop models should allow prediction not only of the final mineralization but also the mineralization dynamics; therefore, they need to be validated for a wide range of EOMs.

 The aim of this work was to quantify the C and N mineralization of a wide range of EOMs during laboratory incubations, to test the ability of a simple generic decomposition model to simulate their C and N mineralization, and to propose some simple calibration methods of the model in order to avoid the use of long and costly laboratory incubations.

**2 Materials and methods**

## *2.1 EOM database*

 The EOM database compiles the results of 663 EOM incubations in soil over at least 90 days under controlled laboratory conditions. The EOM incubations were performed in various research projects for many years and were not all conducted with exactly the same procedure. However, most incubations were carried out according to a standard method

 (AFNOR, 2009): the equivalent of 25 g of dry soil was mixed with the EOM added at a 92 rate corresponding to 2 g organic C  $kg^{-1}$  dry soil. The EOM was dried and ground to a particle size of 1 mm. Mineral N was added in excess to avoid any mineral N deficiency, which could have limited the EOM decomposition rate, and to highlight the potential N immobilization (Recous et al., 1995). The temperature was maintained at 28°C, and the 96 gravimetric water content was equivalent to the field capacity ( $pF = 2.5$ ). A few incubations were performed under different conditions (e.g., a temperature of 15°C and a longer duration and/or with the addition of fresh EOM) (Table S1, Appendix B, online supporting information). Most soils used in the incubations were loamy and noncalcareous 100 with a low carbon content (mean soil carbon content equal to 11.6 g kg<sup>-1</sup>, mean soil organic C:N equal to 9.8) (Table S2, Appendix B, online supporting information).

 In each incubation, the CO<sup>2</sup> evolved and the soil mineral N were measured to determine the mineralized C and N. The net C and N mineralized from each EOM were computed by subtracting the mineralized C and N of an unamended soil (control). The proportions of the net mineralized C and N from the EOMs were obtained by dividing the net C and N mineralized from each EOM by the total amount of added organic C and organic N by the EOM, respectively. We thus hypothesized the absence of a priming effect (Bol et al., 2003).

 The EOMs were classified into 5 groups: livestock manures (n=103), composts (n=337), anaerobic digestates (n=54), sewage sludges (n=70) and others (n=99). Each EOM group was divided into different EOM subgroups (from 1 to 8). Only digestates were not subdivided into different subgroups because of the limited number of digestates in the database compared to the diversity of digestates according to process (solid or liquid state), post-treatment (phase separation or not), and digested waste (animal manures, urban wastes) and the non-significant differences in C and N mineralization among these factors.

 Each EOM was described by its organic C content, its organic and mineral N contents and its Corg:Norg ratio [\(Table](#page-30-0) 1). For 608 of the 663 incubated EOMs, the biochemical fractions of the EOM (Van Soest & Wine, 1967) were also available. These biochemical fractions combined with the proportion of EOM organic C mineralized during 3 days of incubation in soil allowed the computation of the *IROC* value, which is an indicator of the proportion of the EOM organic matter remaining in soils over the long term after application (Lashermes et al., 2009) (Equation 1).

$$
123 \tI_{ROC} = 44.5 + 0.5 \, SOL - 0.2 \, CEL + 0.7 \, LIC - 2.3 \, C_{3d} \tag{1}
$$

 where *SOL*, *CEL*, and *LIC* are the soluble, cellulose-like and lignin-like fractions (percentage of the total organic matter) and *C3d* is the percentage of organic C mineralized during 3 days of incubation in soil (percentage of the total C)

 Finally, to estimate the contributions of organic C from the EOMs to the soil organic matter and of organic N from the EOMs to the soil mineral N, the mineralized organic carbon and the mineralized organic nitrogen after a period of incubation equivalent to one year in field conditions based on mean annual temperature were computed. This equivalent incubation duration was determined by using the concept of "normalized" time (Mary et al., 1996), which modifies the incubation time with correction factors based on soil temperature and soil water content. The correction factors of the STICS model were used (Brisson et al., 2008) (see section [2.2](#page-7-0) and Appendix A). The mean annual temperature in the field applied was 12°C, corresponding to the mean annual temperature in central France, while a constant soil water content equivalent to field capacity was used. The observed quantity of C and N mineralized were not modified, only the corresponding incubation time. For example, 89 days of incubation at 28°C and field capacity corresponded to 365 days in the field at 12°C and field capacity.

<span id="page-7-0"></span>*2.2 EOM decomposition model*

 We used a simple generic decomposition model based on the residue decomposition module of the STICS model. This model was initially developed for crop residues (Nicolardot et al., 2001). In this study, we used a modified version of the STICS decomposition module, adapted from Levavasseur et al. (2021), which allowed a better simulation of C and N mineralization from EOMs of various origins.

 The soil organic matter is subdivided into five pools [\(Figure](#page-34-0) 1). EOM C is divided into labile (*RES1*) and recalcitrant (*RES2*) pools (with RES standing for "residues"). The allocation of nitrogen between these two pools can be different, with each pool having its own C:N ratio (*CNres1* and *CNres2*). This allocation is defined by a single parameter, *aCN1*, which is the ratio between *CNres1* and *CNres*, where *CNres* is the C:N ratio of the EOM as a whole. The labile pool decomposes according to first-order kinetics with the decomposition constant *Kres*, while the recalcitrant pool is directly incorporated into the active organic matter pool. The decomposed labile pool is either mineralized or assimilated by the zymogenous microbial biomass pool (with an assimilation yield *Y*). Soil mineral N may be immobilized if the C:N ratio of the labile pool (*CNres1*) greatly exceeds that of the microbial biomass (*CNbio*). The microbial biomass decays according to first-order kinetics, with the decomposition constant *Kbio*. The decomposed biomass is either mineralized or incorporated in the active soil organic matter, with the humification yield *H*. The soil mineral N may be immobilized during this decay process depending on the C:N ratio of the newly formed active soil organic matter, which is set as equal to the initial C:N ratio of the soil organic matter. The active soil organic matter decomposes according to first-order 162 kinetics, with the decomposition constant  $K_a$  that depends on soil type (Clivot et al., 2017). In this study, however, the *K<sup>a</sup>* constant was optimized to minimize the simulation error of the mineralized C and N in the control treatment, which did not include EOM in the

 incubation. The stable soil organic matter pool is assumed to be inert at the century time scale. With this model, the proportion of mineralized C and N is the sum of three exponential terms (Appendix A, online supporting information).

*2.3 Model calibration*

 Three methods of calibration (i.e., the adjustment of the model parameters) were tested. M1 included a calibration specific to each EOM, wherein the unknown decomposition parameters were optimized for each EOM separately. M2 included a calibration per EOM subgroup, assuming that all EOMs of the same subgroup had the same optimized decomposition parameters. M3 included a calibration per EOM subgroup based on measured EOM characteristics. For the M3 method, the results of the M1 calibration method were used to derive linear relationships between optimized decomposition parameters and two EOM characteristics (*IROC* and *C3d*) for all the EOMs together. The parameters of these linear relationships were further optimized in calibration method M3, i.e., the decomposition parameters were predicted by using linear relationships with EOM characteristics. Simple linear functions with only *IROC* and *C3d* were selected; other types of functions and other EOM characteristics available in the database (e.g., lignin content) did not significantly improve the prediction of the EOM parameters. As for the M2 method, the interest of the M3 method is to avoid the need for an additional experimental incubation for each new EOM (i.e., an EOM not in the database). In addition to M2, the objective of the M3 method is to use easily available EOM characteristics to improve the calibration within each EOM subgroup.

 The upper and lower limits of the optimized parameters were mainly determined according 187 to the literature.  $K_{res}$  varied between 0.005 and 0.7 day<sup>-1</sup>. These values prevented both an accumulation of the labile pool in soil (which would not be consistent) and unrealistic 189 decomposition in only a few days.  $K_{bio}$  was set at 0.0076 day<sup>-1</sup> according to Justes et al.

 (2009), assuming that the zymogenous biomass had the same decay rate for EOMs and crop residues. The assimilation yield (*Y*) varied between 0.1 and 0.6 (Lee & Schmidt, 2014; Sauvadet et al., 2018; Spohn et al., 2016). The mean value of the humification yield (*H*) optimized with calibration method M1 was 0.88. Preliminary tests indicated that fixing *H* at this value for all EOMs did not significantly decrease model performances (Table S3, Appendix B, online supporting information) due to compensation between *H*, *Y* and *RES1*. This high and constant *H* value is consistent with the primary role of microbial necromass in soil organic matter formation (Miltner et al., 2012). *CNbio* was fixed at 7, according to the mean value reported in the literature and because the C:N ratio of microbial biomass is not expected to change markedly (Mooshammer et al., 2014). Fixing *CNbio* at this value for all EOMs did not significantly decrease model performances (Table S3, Appendix B, online supporting information). *RES<sup>1</sup>* and *RES<sup>2</sup>* varied between 0 and 1, allowing EOM to be completely labile or recalcitrant. *CNres1* and *CNres2* were defined as positive and to ensure the nitrogen mass balance (relative to the total amount of N in the EOMs).

 For each method of model calibration, the parameters were optimized to minimize the sum of the root mean square errors (*RMSEs)* for the mineralized C and N (Equation 2).

206 
$$
RMSE = \sqrt{\frac{1}{n} \cdot \sum_{i=1}^{n} (O_i - S_i)^2}
$$
 (2)

207 where  $S_i$  and  $O_i$  are the simulated and observed values for the same measurement date (*i*) and *n* is the number of measurement dates. The optimized (*Kres*) or fixed (*Kbio*) decomposition rates were corrected with the temperature and soil water content functions (Brisson et al., 2008; Mary et al., 1996) (Appendix A) to take into account the effect of 211 these factors.

 The "optim" function in R with the "L-BFGS-B" method was used for optimization (Byrd et al., 1995), giving a lower and an upper bound to each parameter, as mentioned above.

### 214 *2.4 Model validation*

 Calibration method M1 could not be validated on an independent dataset, and the calibration was specific to each EOM. We assessed the performances of calibration methods M2 and M3 by using V-fold cross-validation with three replicates. Each EOM subgroup was randomly divided into three different subsamples. Two subsamples were used for calibration, and the third subsample was used for validation. We repeated this operation three times, changing the validation subsample each time, and then repeated these three operations three times by changing the random division of the sample to avoid any sampling effects in the results.

 The performances of the model for calibration and validation were assessed with three different statistical criteria: the *RMSE*, the mean error (*ME*), and the coefficient of determination (*R²*) (Equations 2, 3 and 4). These criteria were computed for each incubation and then averaged.

227 
$$
ME = \frac{1}{n} \cdot \sum_{i=1}^{n} (O_i - S_i)
$$
 (3)

228 
$$
R^{2} = \left(\frac{\sum_{i=1}^{n} ((o_{i} - \bar{o}) \cdot (S_{i} - \bar{s}))}{\sqrt{\sum_{i=1}^{n} (o_{i} - \bar{o})^{2}} \cdot \sqrt{\sum_{i=1}^{n} (S_{i} - \bar{s})^{2}}}\right)^{2}
$$
(4)

229 where  $S_i$  and  $O_i$  are the simulated and observed values for the same measurement date (*i*) 230 and *n* is the number of measurement dates.  $\overline{O}$  and  $\overline{S}$  are the means of the observations and 231 simulations, respectively.

#### 232 **3 Results**

#### 233 *3.1 Observed C and N mineralization*

234 For a given EOM group, the C and N mineralization varied widely between the different

235 EOMs [\(Figure](#page-34-1) 2). For example, for animal manures, the mineralized C varied between 0

236 and 500 mg C  $g^{-1}$  added C and the mineralized N varied between -500 and 500 mg N  $g^{-}$  <sup>1</sup> added N after twenty days of incubation. Among animal manures, chicken droppings (AM\_CD) produced the greatest net mineralization, while horse manure decomposition resulted in a net N immobilization.

 The calculation of mineralized C after an incubation duration equivalent to one year in the field allows an easy assessment of the potential EOM contribution to C storage in soil (the lower the mineralized C is, the higher the contribution) [\(Figure](#page-35-0) 3). Compared to all other EOMs, the lowest values were obtained for composts, with a median C mineralization 244 equal to 181, 270, 372, 403 and 456 mg C  $g^{-1}$  added C for composts, digestates, animal manures, sewage sludges and other EOMs, respectively. However, a large variability was 246 found within the EOM groups. For example, composts of municipal solid waste (C\_MSW) 247 exhibited a higher C mineralization (median value of  $324 \text{ mg } C \text{ g}^{-1}$  added C) than the average, while composts of green waste (C\_GW) showed a very low C mineralization, 249 equal to  $108 \text{ mg } C \text{ g}^{-1}$  added C. Differences in the mineralized C after an incubation duration equivalent to one year in the field also appeared among animal manures: chicken 251 droppings, pig slurries and horse manures had median values of 545, 448 and 432 mg C g<sup>-</sup> 252 <sup>1</sup> added C, respectively, and were more mineralizable than bovine manure (258 mg C g  $\frac{1}{253}$  added C). If we consider all the EOM subgroups, the mineralized C after an incubation duration equivalent to one year in the field appeared to be well correlated with the *IROC* (R²  $255 = 0.67$ , Table S4, Appendix B, online supporting information). For specific EOM subgroups, the *IROC* was also significantly correlated with mineralized C, except for some EOM subgroups with low mineralization (e.g., green waste compost (C\_GW), Table S4, Appendix B, online supporting information).

 The calculation of net mineralized N after an incubation duration equivalent to one year in the field allowed us to compare the fertilizing values of the various EOMs [\(Figure](#page-35-0) 3).

 Composts and digestates were characterized by a small net mineralization, which varied 262 between -104 and 76 mg N  $g^{-1}$  added N. The short-term N fertilizing value of these EOMs thus relied on their mineral nitrogen content only, which was very low for most composts (except, e.g., C\_GWS and C\_PIS) but high for digestates [\(Table](#page-30-0) 1). Wide variability was found among the animal manures: horse manure decomposition resulted in strong N 266 immobilization (-419 mg N  $g^{-1}$  added N), while chicken droppings produced important net 267 N mineralization (304 mg N  $g^{-1}$  added N). Moderate amounts of N were mineralized from 268 bovine and pig manures (52 and 100 mg N  $g^{-1}$  added N, respectively), which had small mineral N contents. The three sewage sludge subgroups exhibited a high mineralized N 270 (278 to 432 mg N  $g^{-1}$  added N). The last group of EOMs ("others") exhibited variable N mineralization, which is consistent with the highly diverse EOMs within this group, ranging from vegetal residues (e.g., bark) that mainly immobilized N to algae and animal residues (e.g., feather meal) that mainly mineralized N. The variability in the net mineralized N after an incubation duration equivalent to one year in the field for all the EOM subgroups was partly related to that of *C:Norg* (R² = 0.32, Table S4, Appendix B, online supporting information). However, the relationship between the mineralized N and *C:Norg* varied widely among the EOM subgroups: a good correlation was observed for 278 some EOM subgroups (e.g.,  $R^2 = 0.89$  for C\_PIS and  $R^2 = 0.92$  for OTH\_AR) to the 279 absence of a relationship for other EOMs (e.g.,  $R^2 = 0$  for AM HM and SS AGR).

<span id="page-12-0"></span>*3.2 Prediction of C and N mineralization*

 The simulation of C and N mineralization with a calibration specific to each EOM (M1) gave very good results. The observed variability was very well accounted for [\(Figure](#page-35-1) 4): the determination coefficient was 0.92 for C and 0.68 for N, the bias was very low 284 (5 mg C  $g^{-1}$  added C and 2 mg N  $g^{-1}$  added N) and the *RMSE* values were 32 mg C  $g^{-1}$ 



- the model performance based on the validation dataset was better than the performance
- 295 observed for method M2, since the *RMSE* of mineralized C dropped from 99 to 65 mg C g<sup>-</sup>
- 296  $^{-1}$  added C and the *RMSE* of mineralized N decreased from 126 to 110 mg N g<sup>-1</sup> added N.

 Regardless of the calibration method, the model performance strongly varied with EOM subgroup (Tables S3 and S4, Appendix B, online supporting information). For example, the highest *RMSE* was often associated with the EOM that had the highest mineralization (e.g. algae (OTH\_ALG) and agri-industrial wastewater (OTH\_AGRWW)).

## *3.3 Parameters of the calibration per EOM subgroup*

 For calibration method M2, a set of parameters was proposed for each EOM subgroup [\(Table](#page-32-0) 3). In the absence of a specific incubation for a given EOM or of its characteristics (e.g., *IROC*), these parameters can be used to simulate C and N mineralization with acceptable accuracy (section [3.2\)](#page-12-0). The parameters greatly varied according to EOM subgroup. The labile fraction of the residues (*RES1*) varied between 0.13 for the green waste compost (C\_GW) and 1.00 for vinasse (OTH\_VIN) and algae (OTH\_ALG). The microbial assimilation yield (*Y*) ranged from 0.10 (for example, in chicken droppings

309 (AM\_CD)) to 0.60 in green waste and sludge compost (C\_GWS). The parameter  $aCN_I$  was greater than 1 for all EOMs except SS\_AGR, indicating that the C:N ratio of the labile pool (*CNres1*) was greater than the C:N ratio of the recalcitrant pool (*CNres2*). The corresponding simulated C and N mineralization dynamics for the different EOM subgroups are presented in Figure S1 (Appendix C, online supporting information).

*3.4 Parameters of the calibration per EOM subgroup and with EOM characteristics*

 For calibration method M3, the relationships between the model parameters and EOM characteristics were established by considering the significant relationships observed between the optimized parameters of EOM in the first calibration method (M1) and EOM characteristics (Figure S2). The labile fractions of the EOMs (*RES1*) decreased with 319 increasing values of  $I_{ROC}$  ( $R^2=0.61$  with EOM subgroup as a covariable,  $R^2 = 0.49$  for all the EOMs together). The decomposition rate (*Kres*) increased with *C3d* according to a 321 logarithmic relationship ( $R^2 = 0.34$  with EOM subgroup as a covariable,  $R^2 = 0.21$  for all the EOMs together). The other significant relationships were not considered because they did not improve the residual variance. The aforementioned linear relationships were introduced in calibration method M3 to predict the EOM parameters of the model [\(Table](#page-33-0) 4 and Table S7, Appendix B, online supporting information). EOM subgroup was added as a 326 covariable due to the systematic improvement in  $R^2$  when using it.

#### **4 Discussion**

### *4.1 Variability in EOM C and N mineralization*

 The comparison of a wide range of EOMs highlighted some important differences in C and N mineralization among the EOMs [\(Figure](#page-34-1) 2, [Figure](#page-35-0) 3). While C mineralization from the EOMs was generally low for composts, C mineralization can be high for certain EOMs, such as chicken droppings, animal residues or vinasses. The rate of N mineralization from

 the EOMs exhibited even more variability between EOM subgroups (e.g., important net N mineralization for chicken droppings and strong net N immobilization for horse manure). Considering the organic N content of these two EOMs [\(Table](#page-30-0) 1), their typical water content (60% and 30%, respectively, personal data) and their typical application rate (20 337 and  $3 \text{ tha}^{-1}$ , respectively, personal data), the input of organic N would represent 338 approximately 100 kg ha<sup>-1</sup> in both cases, thus leading to a net immobilization of  $\,$  42 kg N ha<sup>-1</sup> for horse manure and a net mineralization of 30 kg N ha<sup>-1</sup> for chicken droppings after one year in the field. Despite the high immobilization caused by horse manure in the first phase, this EOM could have a fertilizing value during the second phase, in which N is slowly released [\(Figure](#page-34-1) 2). Moreover, it is important to highlight that these high values of N immobilization represents only the potential immobilization, which can be reached only when the soil mineral N content is large enough and is a nonlimiting factor in decomposition. The hierarchy of amendment and fertilizing values between the EOM subgroups is in line with existing literature (Lazicki et al., 2020; Mondini et al., 2017; Noirot-Cosson et al., 2017), even though the comparison of so many EOM subgroups is rare. The hierarchy of mineralized C and N after an incubation duration equivalent to one year in the field between the EOM subgroups also reflected their contribution to the SOC (inverse relationship) observed in the field by various authors (Gerzabek et al., 1997; Levavasseur et al., 2020) and to N short-term supply (Gutser et al., 2005).

 In addition to the differences between EOM subgroups, a high variability in C and N mineralization existed inside each EOM subgroup [\(Figure](#page-35-0) 3). This result highlighted the need for a detailed characterization of EOMs, either through laboratory incubations, or with indicators that are easier to retrieve, such as the *IROC* indicator based on biochemical fractionation (Lashermes et al., 2009), to determine the contribution of EOMs to SOC. We found a strong correlation between the mineralized C after an incubation duration

358 equivalent to one year in the field and  $I_{ROC}$  ( $R^2 = 0.67$ ), which was expected because  $I_{ROC}$  was defined as a predictor of residual C in soils under laboratory incubation. Our study, however, validates the interest of *IROC* for a wider range of EOMs, including for some EOM subgroups not used in the *IROC* calibration such as digestates (*R²* between mineralized C and *IROC* equal to 0.62). The mineralized N after an incubation duration equivalent to one year in the field was less but significantly correlated with EOM characteristics (*R²* with *CNres* equal to 0.32), and the correlation was higher for some EOM subgroups (e.g., animal residues (OTH\_AR)). This result contradicts a study showing that *CNres* was a good predictor of N mineralization (*R²* = 0.94) (Lazicki et al., 2020) but confirms another study pointing out that *CNres* is not sufficient for predicting the N mineralization of EOMs of various qualities (Bonanomi et al., 2019). The limited capacity of this parameter to predict N mineralization could be explained by the great diversity of EOMs used in our study. For example, some anaerobic digestates exhibited low *CNres* and induced N immobilization, in opposition to other EOM with low *CNres* like sewage sludge. Moreover, the use of incubations realized in different experimental conditions in our study (temperature, water content) and their "standardization" using the concept of normalized time (Mary et al., 1996) could partly weaken the relationship between EOM characteristics and EOM mineralization for all the incubation taken together, even if most of the EOM incubation that we used were realized in similar conditions (Tables S1 and S2).

 Despite the interest in EOM characteristics, they were not sufficient to predict the dynamics of C and N mineralization in various soil and climate conditions, which required the use of a well-calibrated decomposition model.

*4.2 Model performance*

 The calibration of individual EOMs in the model yielded good model performances for all 382 the EOM subgroups, with *RMSE* values equal to  $32 \text{ mg C g}^{-1}$  added C and  $50 \text{ mg N g}^{-1}$ 

 $\frac{1}{2}$  added N and an  $R^2$  equal to 0.92 and 0.68 for C and N mineralization, respectively. These performances are similar to those reported for other models for a less diverse groups of EOMs (Gale et al., 2006; Mohanty et al., 2011; Mondini et al., 2017). It is also comparable to the performance obtained with the STICS model for crop residues (Justes et al., 2009). When the same set of parameters was used for all the EOMs of the same EOM subgroup, the model performances decreased for C and N, due to the diversity of EOMs within each subgroup. The simulation of C mineralization was improved when the EOM characteristics (*IROC* and *C3ds*) were accounted for in the model parameterization. However, the EOM characteristics were not sufficient for the determination of decomposition parameters capable of simulating the N mineralization kinetics very accurately. Such a result was found by Noirot-Cosson et al. (2017) and Monhanty et al. (2011).

 Regarding the model performances within the three calibration methods, we recommend that a specific EOM laboratory incubation is used to calibrate the model (M1 method) and to most accurately predict its potential behavior in the field, when the user considers a specific EOM (Figure 5). However, model users often do not have any information on the particular EOM, except the quantity applied. In that case, we showed that a calibration depending only on its subgroup (M2 method) bring some insights that are sufficient to test some global EOM application scenarios. We do not recommend to use the M2 "default" method to simulate accurately the effects of a specific EOM in case of repeated application so as not to accumulate over time the error in the estimates. As an alternative to the latter case, we highly recommend the addition of EOM easily measurable characteristics to calibrate the model (M3 method), as tested in Levavasseur et al. (2021) for the simulation of a long term experiment. The quality of the calibration methods M2 and M3 varied according to the considered EOM subgroup (Table S5 and S6). We recommend the user to consider the EOM subgroup to determine whether a specific calibration of the modeled

 EOM (M1) is required or not and to put it in perspective with the objective and characteristics of the modeling study (e.g., importance of EOM application in the scenarios).

 Other EOM characteristics not considered here might contribute to improving the prediction of EOM model parameters and, subsequently, the simulation of C and N mineralization. For example, the C:N ratio of biochemical fractions has been shown to be useful in predicting N mineralization (Morvan & Nicolardot, 2009; Parnaudeau et al., 2004) as well as 13C-CPMAS NMR spectral regions (Bonanomi et al., 2019; Pansu et al., 2017). However, these characteristics were not available in the EOM database we used and are not usually available outside research laboratories. The identification of readily available EOM characteristics that can be used to calibrate decomposition models remains a challenge.

#### *4.3 Model hypotheses*

 The modification of the STICS model proposed by Levavasseur et al. (2021) to model EOM decomposition was used to simulate a wide range of EOM decomposition levels in our study. This modification includes the addition of a recalcitrant pool of EOM that is directly incorporated in the soil active OM and is not assimilated by the microbial biomass, assuming that this recalcitrant pool results from the previous digestion of raw matter by microbial biomass (in animal guts or during treatments). This pool is considered to have the same dynamics as the active pool in the STICS model. Several authors have hypothesized that EOM is composed of at least two dynamic pools (Gijsman et al., 2002; Mondini et al., 2017) to simulate EOM decomposition; however, this distinction was not necessary for simulating crop residue decomposition in the STICS model, which considered a single dynamic pool (Justes et al., 2009). Whereas the decay rate of the recalcitrant pool of EOM was optimized and inferior to the decay rate of the soil active

 OM in the study of Levavasseur et al. (2021), this decay rate was set equal to that of the soil active OM in the present study. This simplification caused some overestimation of C mineralization for very stable EOMs, such as composts [\(Figure](#page-35-1) 4). The plateau of C mineralization usually observed in laboratory incubations could not be properly simulated with this model. However, this modification was made to limit the number of parameters in the STICS model. Moreover, several authors have suggested that SOM stability after EOM application is not modified (Liu et al., 2018; Luan et al., 2019), even though another study suggested the opposite (Peltre et al., 2017). Levavasseur et al. (2020) successfully simulated carbon storage in long-term field experiments with EOM application using the AMG model without assuming a greater stability of EOM in comparison to that of SOM. Because the STICS and AMG formalisms for SOC are very similar (i.e., SOC is divided into an active and a stable pool, and the mineralization function is the same) (Clivot et al., 2019), the STICS model should also allow the simulation of C and N dynamics of EOMs without assuming a greater stability of those EOMs, i.e., by allocating EOMs only to soil active OM (either directly, or after soil microbial biomass assimilation, [Figure](#page-34-0) 1). Directly allocating the recalcitrant fractions of EOMs to soil active OM could change the soil C:N ratio after repeated applications of the EOMs. This effect is consistent with the fact that soil C:N usually increases with SOC content (Mooshammer et al., 2014). Other simplifications made in our study include the use of a constant humification yield (*H*) of microbial necromass and a constant C:N ratio of microbial biomass (*CNbio*), in line with recent knowledge about SOM formation (Miltner et al., 2012; Mooshammer et al., 2014). However, the *H* value retained in our study is higher than the value reported in the study of Miltner et al. (2012). Concerning the decay rate of the labile pool *Kres*, the calibration 456 method M1 gave some optimized values equal to the minimal  $(0.005 \text{ day}^{-1})$  or maximal  $(0.7 \text{ day}^{-1})$  value for some EOM (Figure S2). The use of lower minimal and higher

 maximal values could improve the simulation. However, we decided to keep these limits to prevent both an accumulation of the labile pool in soil (which would not be consistent) and unrealistic decomposition in only a few days.

 The use of EOM incubations to calibrate the decomposition module of the STICS model and predict EOM decomposition under field conditions requires some correction factors regarding water content and temperature. The use of EOM incubations run under different conditions (temperature, soil water content, soil type) could have introduced errors into the EOM calibration, because of the uncertainty associated with these correction factors. However, most of the EOM incubations used in this study were realized in similar conditions (loamy soil, 28°C, water content close to or equal to field capacity, Tables S1 and S2). The effects of soil type, water content, temperature and N mineral availability are already taken into account in the STICS soil-crop model to calculate EOM mineralization *in situ* (Brisson et al., 2008). The correction factors should allow the extrapolation of laboratory incubations to field conditions (Gale et al., 2006). However, laboratory incubations are usually realized with dried and crushed EOM, which can influence the mineralization of C and, especially, N (Le Roux et al., 2016). Correction factors that consider EOM preparation before incubation should also be proposed.

 Finally, beyond the STICS users, the decomposition module (Appendix A) and the calibrations proposed in this paper could be used outside the STICS model to predict the C and N mineralization of a large diversity of EOMs.

#### **5 Conclusions**

 We quantified the C and N mineralization of a wide range of EOMs based on a database of more than 600 EOM incubations, from five groups (animal manures, composts, sewage sludges, digestates, and others) and 26 subgroups (e.g., bovine manure, green waste

 compost). This represents one of the largest and most diversified syntheses of EOM incubation data. The results indicated a wide diversity in EOM contributions to C storage in soil and N supply for crops, both between EOM subgroups and within EOMs of the same subgroup. The EOM incubations were used to calibrate a simple generic decomposition model included in the STICS soil-crop model to simulate EOM decomposition. Individual EOM calibration yielded the best model performances for the simulation of C and N mineralization and is the recommended calibration method of the STICS model for accurate simulations of scenarios of application of a specific EOM in field conditions. In the absence of EOM incubation data, we proposed two calibration methods for the 26 different subgroups of EOMs in the STICS model: using either a unique calibration per EOM subgroup or EOM characteristics as predictors of model parameters. Although these two calibration methods decreased model performances, they allow the prediction with a reasonable performance of the EOM contribution to soil C storage or to mineral N supply to determine the best practices for their use (amount, period of application, fertilized crops, etc.). Future research is needed to better predict the EOM parameters from easily available EOM characteristics and verify the extrapolation of laboratory incubation to field conditions.

#### **6 Acknowledgements**

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## 665 **8 Tables**

## 667 **Table 1** Main characteristics of the EOM database



<span id="page-30-0"></span>\* 668 Indicator of residual organic carbon in soils (Lashermes et al., 2009)

- <span id="page-31-0"></span>669 **Table 2** Model performance per incubation for all the EOMs together according to the
- 670 calibration method (mean calibration and validation performance of the cross-validation
- 671 iterations for methods M2 and M3) (ME: mean error, RMSE: root mean square error, R²:
- 672 coefficient of determination).



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- <span id="page-32-0"></span>675 **Table 3** Optimized parameters for the STICS decomposition module per EOM subgroup
- 676 (obtained with the whole EOM dataset used for calibration method M2). *Kbio*, *H* and *CNbio*
- 677 were fixed at  $0.0076 \text{ day}^{-1}$ , 0.88 and 7.0, respectively. The meaning of the EOM subgroup can
- 678 be found in Table 1.



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- <span id="page-33-0"></span>681 **Table 4** Calibration for the STICS model per EOM subgroup and according to the EOM
- 682 characteristics (M3) (the numerical values of *αres1*, *βres1*, *αkres*, *βkres*, *Y*, *aCN1* per EOM subgroup
- 683 are given in Table S7)

684



#### **9 Figure legends**

- <span id="page-34-0"></span>**Figure 1** Decomposition model of EOM in the modified version of the STICS model. The C
- fluxes are represented as solid black lines, and the N fluxes are represented as dashed gray lines.
- *RES<sup>1</sup>* and *RES<sup>2</sup>* are the proportions of C from the EOMs in the labile pool and recalcitrant
- 691 pools, respectively, with  $RES_2 = 1 RES_1$
- *CNres*, *CNres1* and *CNres2* are the C:N ratios of the EOM and the labile and recalcitrant pools,
- respectively, and *aCN1* is the ratio *CNres1* : *CNres*
- *Kres1* is the decomposition rate of the labile pool
- *Y* is the C assimilation yield of the labile pool
- *CNbio, CN<sup>a</sup>* and *CN<sup>s</sup>* are the C:N ratios of the microbial biomass, and active and stable pools, respectively
- *Kbio* and *K<sup>a</sup>* are the decay rates of the zymogenous biomass and the active soil organic matter pools, respectively
- *H* is the C humification yield.

<span id="page-34-1"></span> **Figure 2** Observed mineralized organic carbon and nitrogen for the different groups of EOMs in laboratory incubations. Each gray line represents an EOM incubation, while the colored lines represent fitted local polynomial regressions per EOM subgroup (performed with the loess function in R). Very few data on N mineralization are lower than -100% and the Y-axis was cut to -100% to improve the visibility. The meaning of the EOM subgroup abbreviations can be found in Table 1.

- <span id="page-35-0"></span>**Figure 3** Observed mineralized carbon and net mineralized nitrogen for each EOM group and
- subgroup after incubation for a duration equivalent to one year under field conditions. The
- meaning of the EOM subgroup abbreviations can be found in Table 1.
- <span id="page-35-1"></span>**Figure 4** Observed and simulated values of carbon and nitrogen mineralization for each EOM
- and sampling date, according to the calibration method (results obtained by using the whole
- EOM dataset for calibration).
- **Figure 5**. Conceptual diagram for the calibration of an EOM in the decomposition model.







**Net mineralized organic C**

### **Carbon mineralization**



#### Experimental data from laboratory incubation for the considered EOM ?

