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New radiolarian assemblages from La Désirade Island basement complex (Guadeloupe, Lesser Antilles Arc) and Caribbean tectonic implications

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Key words : Radiolaria, La Désirade, Guadeloupe, Lesser Antilles, Caribbean plate, Jurassic

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Abstract

New radiolarian assemblages of Late Jurassic age were extracted from seven chert localities of the basement complex of La Désirade Island (Guadeloupe), the oldest geological unit of the Lesser Antilles Arc and the eastern Caribbean region. These fauna range in age from Kimmeridgian or early Tithonian to early or early late Tithonian, showing that La Désirade cherts are not all coeval, contrary to what was established by previous studies. Furthermore, the close spatial association of cherts and volcanic elements composed of basaltic flows and volcanoclastics suggests a mixed setting of hydrothermal siliceous deposits and biogenous sedimentation consistent with an arc-related oceanic setting rather than an ophiolite-type oceanic crust. As the youngest age is obtained to the northwest of the basement complex whereas older data are found to the southeast, it also suggests that La Désirade basement complex may have undergone some structural stacking consistent with an accretion-type geodynamic setting. Our biochronological data provide a preliminary tectonic model of an eastward subduction zone for Lesser Antilles arc basement emplacement during the Mesozoic.

Nouveaux assemblages de radiolaires du complexe de base de La Désirade (Guadeloupe, Petites Antilles) et implications tectoniques caraïbes

Mots-clés : Radiolaria, La Désirade, Guadeloupe, Petites Antilles, plaque Caraïbe, Jurassique

Résumé

De nouveaux assemblages de radiolaires d'âge Jurassique supérieur sont extraits de sept localités de jaspes siliceux associés au complexe de base de l'île de La Désirade (archipel de Guadeloupe), l'unité géologique la plus ancienne de l'arc des petites Antilles et de la région est-caraïbe. Ces faunes ont un âge qui s'étend du Kimméridgien supérieur-Tithonien inférieur au Tithonien inférieur-Tithonien supérieur basal montrant que, contrairement à ce qui avait été établi précédemment, les sédiments océaniques de La Désirade sont diachrones. Par ailleurs, l'association très particulière des jaspes à radiolaires et des éléments volcaniques du complexe de base suggère l'existence d'une sédimentation siliceuse hydrothermale associée à des dépôts biogènes dans un contexte de sédimentation océanique d'arc plutôt que de dépôts sur une croûte océanique vraie. Nos résultats suggèrent que le complexe de base de La Désirade a subi un raccourcissement tectonique du nord-ouest au sud-est, compatible avec une structuration de complexe d'accrétion associé à une subduction vers l'est, fournissant un modèle préliminaire de mise-en-place tectonique de ce segment de l'arc des Antilles au Mésozoïque.

INTRODUCTION

The scope of this study is to reassess the age and significance of La Désirade Island oceanic sediments and associated volcanic rocks belonging to the Lesser Antilles arc basement (fig. 1). The geodynamic history and plate tectonic evolution of the Caribbean region have been a subject of intense debate for over 40 years, due to the complexity of its elements and their protracted history. This region evolved at the interaction between two major continents (North and South Americas) and two major oceans (Pacific and Atlantic) since the Jurassic, and consequently comprises oceanic rocks of various settings (ophiolites, arcs, back-arcs, accretion complexes). The island of La Désirade is an important element of this Caribbean network of Jurassic oceanic rocks, being the easternmost element of the Lesser Antilles island arc (fig. 2). Furthermore, it is the only element of the Lesser Antilles region exposing Mesozoic crystalline basement. It has generated a great deal of attention since Fink [1968] determined a Late Jurassic radiometric age for a trondhjemite locality whereas the oldest age of the Lesser Antilles was thought to be much younger. The actual nature of the basal complex of La Désirade Island has been much debated and no final model has yet been established with certainty [see recent discussion by Mattinson *et al.*, 2008]. To summarize, it has been considered as part of an ophiolitic complex made of oceanic crust [Mattinson *et al.*, 1973, 1980; Fox & Heezen, 1975; Dinkelman & Brown; 1977; Le Guen de Kerneizon *et al.*; 1979], or a primitive island arc subsequently detached from the Greater Antilles [Fink, 1968, 1970a, 1970b, 1972; Bouysse *et al.*, 1983; Donnelly *et al.*, 1990, Gauchat, 2004]. More recently, Mattinson & Hopson [*in* Mattinson *et al.*, 2008] stood by their original conclusion that the igneous rock assemblage and vertical distribution on La Désirade Island and its submerged foundation are compatible with the upper part of an ophiolite.

Radiolarian-bearing cherts from La Désirade have previously been studied and dated by Bouysse *et al.* [1983], Montgomery *et al.* [1992] and Mattinson *et al.* [2008]. Radiolarians extracted from these rocks have not only been used for biochronology but also for biogeography in order to assess the former location and the tectonic origin of the Caribbean plate [Montgomery *et al.*, 1992, 1994a, 1994b]. However, one problem with former studies is that radiolarian ages provided by Bouysse *et al.* [1983] and Montgomery *et al.* [1992] are not equivalent (Early Cretaceous vs. Late Jurassic, respectively). In order to reassess their age in the light of updated radiolarian biozonations, we undertook a new investigation using field and laboratory techniques previously applied to Cordilleran chert terranes [Cordey & Krauss, 1990]. As a result, we present new micropaleontological and biostratigraphic data from La Désirade cherts combined with new field observations. These results provide some new constraints on the nature of La Désirade oldest rocks and therefore on the preliminary tectonic evolution of the Lesser Antilles and the eastern Caribbean region.

GEOLOGICAL SETTING

La Désirade Island

La Désirade is a 10 km long island located to the east of the main island of Guadeloupe (fig. 2), standing out in the forearc of the Lesser Antilles region. The western part of the Guadeloupe archipelago is a segment of the active Plio-Quaternary arc currently represented by the volcano “La Soufrière”. The eastern part of the Guadeloupe archipelago is composed of Grande Terre, Marie-Galante, and La Désirade, which all belong to the Calcareous Lesser Antilles (“Petites Antilles calcaires”). These islands correspond to eroded fragments of an Eocene-Miocene arc capped by Plio-Quaternary calcareous formations, showing that the activity of the Lesser Antilles arc has shifted to the west since the Eocene-Miocene (fig. 2). La Désirade Island is

bounded to the northeast by a trench 4000 m deep, one major escarpment of the eastern Caribbean plate topographic features.

Field studies and geological mapping of the island were originally undertaken by Fink [1970b] and then Westercamp [1980], leading to the differentiation of three units by Bouysse *et al.* [1983]:

- 1) a Central Acid Massif composed of felsic quartz-diorite intrusion associated with rhyolitic lava flows;
- 2) a Northeast Volcanic Complex comprising pillowed metabasalt and interbedded radiolarian chert;
- 3) an Upper Meta-andesitic group represented by east-west trending dike swarm complex.

Mattinson *et al.* [2008] have kept Unit 2 of Bouysse *et al.* [1983] but have renamed the first and third units the Trondhjemitic/Rhyolitic Igneous Complex and the Diabasic/Microdioritic Dike Swarm, respectively. Bouysse *et al.* [1983] interpreted the trondhjemitic/rhyolite assemblage to be the oldest on the island as shown by a Late Jurassic radiometric age (~145 Ma, Fink, 1968; Mattinson *et al.*, 1973, 1980) and considered that their unit 2 was younger as shown by a single locality of Early Cretaceous radiolarians (Hauterivian-Barremian). This micropaleontological datum has since been rightfully questioned by Montgomery *et al.* [1992] who reported five radiolarian assemblages of Late Jurassic age (mid late Tithonian). These radiolarian data have recently been documented again by Mattinson *et al.* [2008] who also produced a new radiometric age of 143.74 +/- 0.33 Ma using CA-TIMS zircon method.

La Désirade cherts

They are exposed within the Northeast Volcanic complex of Bouysse *et al.* [1983] (fig. 3) and are found in various settings (fig. 4): 1) bedded sequence within pillow basaltic flows, locally duplicated, 2) inter- and intra-pillow fillings, 3) thin layers in volcanoclastics. Fink [1970b] first

implied that radiolarian cherts occur within basalts but also overlie them. This interpretation was recently revised by Mattinson *et al.* [2008] who consider that “the sedimentary rocks occur within the submarine lavas, not on top of them”. In their opinion, radiolarian cherts from La Désirade Island are not equivalent to typical radiolarian ribbon chert (RRC) usually composed of rhythmically bedded deposits in which thin chert layers are separated by shaly partings, a feature amplified with diagenesis. Instead, they consider these cherts as intralava sediments derived from heating, oxidation and silicification of initially porous ferruginous hydrothermal sediment when they were buried beneath hot new submarine flows. To support their view, Mattinson *et al.* [2008] state that “radiolarians are sparse and commonly absent from the red cherts”.

BIOSTRATIGRAPHIC RESULTS

Chert localities, faunal assemblages and ages

On the island of La Désirade, chert exposures are essentially found: 1) on cliffs or on the shore between Baie Mahault and Pointe du Grand Abaque (fig. 3), more particularly around Pointe Doublé and Pointe Mancenillier; 2) in small outcrops on topographic knobs located between Baie Mahault and Pointe Mancenillier (fig. 3); these knobs are broadly aligned in a SW-NE direction.

Seven radiolarian localities have yielded diagnostic radiolarians (pl. 1, 2). Some morphotypes are identical to some documented by Bouysse *et al.* [1983], Montgomery *et al.* [1992] and Mattinson *et al.* [2008]. However these studies do not provide any fossil location in coordinates, preventing precise geographic correlations. A complete list of radiolarian morphotypes obtained during this study is presented in Table 1. The following descriptions include details on chert exposures, microfossil content and biochronological assignments. The

nature and composition of our radiolarian assemblages led us to use Baumgartner *et al.* [1995] biozonation, but we also took into account other biochronological data from North America [Pessagno *et al.*, 1984, 1993], Asia [Matsuoka, 1995a; Matsuoka *et al.*, 2002], and Tethys [Gorican, 1994].

LDE3 (coordinates N16°20'02.2" W61°00'11.8"; bedding orientation and dip N078.05W) (fig. 3): at the base of cliff, 30 cm thick red/mauve chert sequence between masses of basalts (layers or pillows); it comprises five chert beds with thin red and green shale interlayers; sequence exposed over about 60 m along the shore. Radiolarians from this locality are badly preserved. Only *Vallupus* sp. was identified, indicating a Late Jurassic age (Kimmeridgian or Tithonian).

LDE7 (coordinates N16°19'56.2" W61°00'19.4"; bedding orientation and dip N120.30NE) (fig. 3): located on the shore below the lighthouse, 50 cm thick red chert sequence with one isolated basalt pillow. Occurring radiolarians are *Eucyrtidiellum ptyctum* (Riedel & Sanfilippo), *Obesacapsula verbana* (Parona) and *Praeconocaryomma* sp.. This association was not identified previously on the island [Bouysse *et al.* 1983, Montgomery *et al.* 1992, Mattinson *et al.* 2008]. It is not easily correlated with the zonation proposed by E.A. Pessagno Jr. [see Mattinson *et al.* 2008 and references therein]. However, the two species *Obesacapsula verbana* and *Eucyrtidiellum ptyctum* are present in UAZ11 of Baumgartner *et al.* [1995] zonation, providing a late Kimmeridgian or early Tithonian age.

LDE9 (coordinates N16°19'44.1" W61°00'41.3") (fig. 3); 790 m to the southwest of the lighthouse, small outcrop of red chert with no clear relationships with basalts. Radiolarians comprise *Archaeodictyomitra excellens* (Tan), *Archaeodictyomitra minoensis* (Mizutani), *Eucyrtidiellum* sp., *Loopus primitivus* (Matsuoka & Yao), and *Pseudodictyomitra carpatica*

(Lozynyak). Again, this assemblage was not found previously on the island, and can be assigned to UAZ11-12 of Baumgartner *et al.* [1995]. The corresponding time interval ranges from the late Kimmeridgian to the early late Tithonian.

LDE12 (coordinates N16°20'00.7" W61°00'35.6"; bedding orientation and dip N002.45E) (fig. 3); 440 m to the northwest of lighthouse, 50 to 80 cm thick chert sequence on a topographic knob; red to dark red chert beds of variable thicknesses. This locality yields a well-preserved radiolarian association comprising *Archaeodictyomitra* sp. aff. *excellens* (Tan), *Archaeodictyomitra minoensis* (Mizutani), *Archaeodictyomitra* sp., *Emiluvia* sp. cf. *salensis* Foreman, *Eucyrtidiellum pyramis* (Aita), *Loopus primitivus* (Matsuoka & Yao), *Pseudodictyomitra carpatica* (Lozynyak), *Pseudodictyomitra* sp., *Pseudodictyomitrella* sp. aff. *tuscanica* (Chiari, Cortese & Marcucci), *Sethocapsa* sp., *Stichocapsa* sp., ?*Stylocapsa* sp., *Williriedellum* sp., *Xitus gifuensis* Mizutani and *Zhamoidellum* sp. aff. *ovum* Dumitrica. This assemblage contains some morphotypes from Tibet and Japan presented by Matsuoka *et al.* [2002; 2005] belonging to their *Pseudodictyomitra carpatica* biozone. For instance, our *Pseudodictyomitrella* sp. aff. *tuscanica* (Chiari, Cortese & Marcucci) (pl. 2, fig. 11) closely resembles the specimens from Shirokawa area (southwest Japan) illustrated by Matsuoka *et al.* [2005; fig. 5, n° 7] and attributed by these authors to *Stichocapsa praepulchella* Hori, a probable junior synonym of *Pseudodictyomitrella tuscanica* (Chiari, Cortese & Marcucci) (see Chiari *et al.* [1997] and Hori [1999]). Overall, this assemblage can be assigned to UAZ12 [Baumgartner *et al.*, 1995] of early to early late Tithonian age.

LDE22 (coordinates N16°20'12.8" W61°00'13.5"; bedding orientation and dip N100.30NE) (fig. 3); 230 m to the northeast of the meteorological station building, 30 to 100 cm thick red chert sequence with 'chevron' folds. This locality has a fairly well-preserved radiolarian

association comprising *Emiluvia chica* Foreman, *Eucyrtidiellum* sp. aff. *nodosum* Wakita, *Homoeoparonaella argolidensis* Baumgartner, *Pantanellium whalenae* Pessagno & MacLeod, *Pantanellium squinaboli* (Tan), *Parvicingula* sp., *Praeconocaryomma* sp., *Pseudictyomitra* sp., *Sethocapsa accincta* Steiger, *Sethocapsa horokanaiensis* Kawabata, *Sethocapsa* sp., *Stichocapsa* sp., *Vallupus hopsoni* Pessagno & Blome, *Zhamoidellum ovum* Dumitrica and *Zhamoidellum* sp.. This assemblage comprises some morphotypes previously documented by Matsuoka *et al.* [2002] and Gorican [1994] such as *Sethocapsa accincta* and *S. horokanaiensis*. It is also correlative with UAZ11 of Baumgartner *et al.* [1995] of late Kimmeridgian or early Tithonian age. However, the occurrence of *Vallupus hopsoni* (absent from Baumgartner *et al.* [1995]) would favor an early Tithonian age [Matsuoka, 1995] for this locality.

LDE27 (coordinates N16°20'25.9" W61°00'21.9", no clear orientation) (fig. 3); 680 m to the north-northwest of the meteorological station; along shore, small outcrop of red chert within basalts. Radiolarians from this locality are badly preserved, with only *Archeodictyomitra* sp. being identified. It provides a Jurassic age without further precision.

LDE30 (coordinates N 16°20'11.5" W61°00'15.9") (fig. 3); 200 m from meteorological station, subexposure of red chert. Radiolarians are moderately well-preserved and comprise *Angulobracchia* sp., *Eucyrtidiellum* sp., *Hsuum* sp. aff. *cuestaense* Pessagno, *Obesacapsula* sp., *Orbiculiformella* sp., *Paronaella* spp., *Parvicingula* sp., *Praeconocaryomma* sp., *Sethocapsa* sp. cf. *accincta* Steiger, *Sethocapsa* sp. cf. *horokanaiensis* Kawabata, *Tethysetta* sp. cf. *dhimenaensis* (Baumgartner) and *Williriedellum* sp.. A tentative correlation for LDE30, which lacks well-preserved diagnostic morphotypes, would be with UAZ11 of Baumgartner *et al.* [1995] of late Kimmeridgian or early Tithonian age.

Age range of La Désirade cherts

In summary, our results show that La Désirade cherts are Late Jurassic, ranging from late Kimmeridgian or early Tithonian to early or early late Tithonian. These localities do not carry identical faunal associations, neither do they all have the same age. Localities LDE7, LDE22 and LDE30 are assigned to biozone UAZ11 of Baumgartner *et al.* [1995] of late Kimmeridgian or early Tithonian age. The occurrence of *Vallupus hopsoni* at LDE22 would constrain this locality to the early Tithonian. LDE12 is assigned to the younger biozone UAZ12 of early of early late Tithonian age. This brings new and critical information to the knowledge of La Désirade Island geology, as previous studies by Bouysse *et al.* [1983], Montgomery *et al.* [1992] and Mattinson *et al.* [2008] considered chert exposures to be coeval (revised age: mid late Tithonian; see below).

DISCUSSION AND INTERPRETATION

Previous radiolarian studies

As mentioned earlier, Bouysse *et al.* [1983] considered their unit 2 to carry Early Cretaceous (Hauterivian-Barremian) radiolarians. Their single assemblage was wrongly assigned at the time to the *Eucyrtis tenuis* Zone of Foreman [1975], as shown by Montgomery *et al.* [1992] who documented radiolarians of Late Jurassic age at five localities (M3, M4, M6, TD and PM7 from their fig. 2). These assemblages were assigned to Subzone 4 beta or the lower part of Subzone 4 alpha of zonation from Pessagno *et al.* [1987] and Yang & Pessagno [1989] of late early to mid late Tithonian age. Recently, Mattinson *et al.* [2008] documented again the original results of Montgomery *et al.* [1992] with one additional and coeval radiolarian locality from Pointe Frégule (“PF” from their fig. 5) and revised the original biochronological correlation of Montgomery *et al.* [1992] to the upper part of Subzone 4 beta of mid late Tithonian age.

Several morphotypes reported by Montgomery *et al.* [1992] are present within Baumgartner *et al.* [1995] UA biozonation, which allows to compare biochronological assignments. Their best and most representative locality “M6” includes *Acanthocircus* sp. cf. *dicranacanthos* (UAZ10-UAZ17), *Eucyrtidiellum ptyctum* (UAZ5-UAZ11), *Loopus primitivus* (UAZ7-UAZ12), *Mirifusus baileyi* (UAZ9-UAZ11) and *Tritrabs ewingi* (UAZ7-UAZ12), an assemblage which can be assigned to UAZ 10-11. Although Montgomery *et al.* [1992] and Mattinson *et al.* [2008] do not provide precise locality coordinates, their map suggests that M6 is located between our localities LDE22 and LDE30 which we assign to biozone UAZ11. Therefore the radiolarian assemblage reported by Montgomery *et al.* [1992] and Mattinson *et al.* [2008] would be equivalent to our oldest localities (LDE7, LDE22, LDE30) interpreted here as late Kimmeridgian or early Tithonian in age. Although we agree on the potential occurrence of late Tithonian radiolarians on the island, we believe that they come from a younger chert sequence not reported by previous authors (LDE12, fig. 3).

Depositional environment

La Désirade cherts have been previously considered as volcanic jaspers rather than radiolarian ribbon chert (RRC), based on spatial association with basalts as well as poor radiolarian recovery [Mattinson *et al.*, 2008]. We propose here a different interpretation: the depositional environment of La Désirade cherts is probably a combination of factors comprising volcanic-related hydrothermal jasper as well as slow radiolarian deposition generating ribbon chert sediments. This interpretation is based on the following observations: 1) true radiolarian ribbon chert do occur on the island, more particularly inland where we document our youngest radiolarian assemblage (LDE12); 2) as shown here, La Désirade chert deposition is not restricted to the mid late Tithonian as previously thought. Therefore, diachronous volcanic

flows could have provided, after cooling, topographic basements for temporary planktonic deposition; 3) the quality of La Désirade radiolarians is not significantly different from that of typical ribbon cherts found for instance within Paleozoic and Mesozoic accretionary complexes [Cordey, 1998], as shown by the fair preservation of diagnostic morphological elements (see pl. 1, 2). In some cases, this may have been linked to the presence of volcanic-related silica-rich seawater [see discussions in De Wever *et al.*, 1994; Racki & Cordey, 2000].

Paleobiogeographic affinities

Previous paleobiogeographic analysis on radiolarian associations from La Désirade Island are confusing: Montgomery *et al.* [1992] assigned them to Northern Tethyan to Boreal paleolatitudes, whereas Mattinson *et al.* [2008] propose a Southern or Northern Tethyan affinity [*sensu* Pessagno *et al.*, 1993]. According to Mattinson *et al.* [2008], all Upper Jurassic red cherts from the basement complexes of Puerto Rico, the Dominican Republic and Cuba have the same paleolatitudinal character. The presence of genus *Vallupus* has been used as a marker of paleolatitude by Matsuoka [1995b] who considers it characteristic of low and middle paleolatitudes (within 25° of the Jurassic paleoequator). In combining these two models, the depocenters of La Désirade cherts could correspond to a fairly large ocean domain located at ~ N 25°-15° or S 15°-25°.

According to Baumgartner *et al.* [2004], radiolarians from La Désirade resemble coeval faunas extracted from cherts of the Californian Coast Range Ophiolite which they infer to have formed in the upwelling area along the Western American Margin. In their opinion, the presence of these faunas does not allow for a paleolatitudinal interpretation, but rather suggests a paleolongitudinal displacement from the Eastern Pacific and/or the western American Margin into their present position in Central America and the Eastern Caribbean.

Our faunal associations partly differs in composition from those documented previously by Bouysse *et al.* [1983] and Mattinson *et al.* [2008], showing the difficulty of reliably identifying radiolarian faunal provincialism. At La Désirade, it is actually possible to come up with different conclusions depending on the studied localities. For instance, LDE12 bears a significant resemblance with Tethyan fauna and does not contain *Vallupus* which has been observed in all well-preserved samples from Mattinson *et al.* [2008], whereas LDE22 contains *Parvicingula* sp. and *Vallupus* sp. which are important components of Mattinson *et al.* [2008], and key elements on discussions by authors on paleogeography [Matsuoka, 1995b; Pessagno *et al.*, 1993].

Although La Désirade Island and elements of the proto-Caribbean basement may have undergone significant displacements from the Eastern Pacific into the Caribbean region [Montgomery *et al.*, 1994a, 1994b; Baumgartner *et al.*, 2004], this interpretation may be more a result of consensual tectonic scenarios rather than a constrained model strictly based on paleobiogeography. Obviously, significant tectonic convergence occurred at the interaction between the Pacific and the Caribbean regions, but it doesn't necessarily imply that radiolarian depocenters had to be restricted to the Pacific realm. It was, after all, the same oceanic domain linking the Pacific to the Atlantic. Moreover, the absence of long-lived oceanic successions within the basement complexes of the lesser Antilles would tend to favour fairly restricted displacements of oceanic terranes prior to their entrapment in the Caribbean region. Similarly, Bortolotti & Principi [2005] suggested that some Jurassic oceanic elements of the eastern border of the Greater and Lesser Antilles could have originated within the Central Atlantic oceanic realm.

Structure of La Désirade basement

Our biochronological data suggest that the Northeast Complex is composed of basalt/chert successions of slightly different ages. In terms of spatial distribution, our oldest localities LDE7, LDE22 and LDE30 (late Kimmeridgian or early Tithonian) are found to the east of the complex whereas the youngest exposure LDE12 (early or early late Tithonian) is located to the west of the study area, showing an age polarity perpendicular to the fault structures exposed on the slope dipping eastward between Pointe Doublé and Pointe Mancenillier (fig. 3). Although the age coverage is far from being conclusive, these results suggest that the Northeast Complex may contain some structural stacking of low metamorphic grade. This is also supported by the occurrence of “chevron” folds and tectonic duplication within chert sequences on the eastern shore of the island. Therefore the Northeast Complex may have undergone an accretionary-type tectonic emplacement. This would be consistent with former interpretations that La Désirade eastern units corresponds to a basement complex trapped within the forearc of the Lesser Antilles in a subduction zone in the Mesozoic [Ghosh *et al.*, 1984; Meschede & Frisch, 1998]. The age distribution of chert localities shows a southeast to northwest-directed polarity from oldest to youngest exposures, which could be related to accretion within a former eastward-directed subduction setting. Obviously this hypothesis must be considered with caution because we do not know the range of post-emplacement tectonics (rotations and/or overturns). However, this model may be tested in the event of future underwater dredging on the northern and eastern slopes of the island.

Until recently, it was unclear which was the oldest rock unit of La Désirade Island, which would also be the oldest unit of the entire Lesser Antilles region and the eastern part of the Caribbean plate. The new radiometric age of 143.74 +/- 0.33 Ma recently produced by Mattinson *et al.* [2008] from the Trondhjemitic/Rhyolitic Igneous Complex (LDIC) would

correspond to either late Tithonian [Palfy *et al.*, 2000] or early Berriasian [Gradstein *et al.*, 2004]. If late Tithonian, this radiometric datum would partly overlap the youngest radiolarian assemblage, also of possibly late Tithonian age. Based on geological observations, Mattinson *et al.* [2008] favoured the interpretation that the basalt/chert association of the Northeast Complex is older than the Igneous Complex (LDIC). Our results not only support this view but also provide a clearer understanding of the age relationships among these units, as our oldest radiolarian assemblages (biozone UAZ11) are probably no younger than early Tithonian. Therefore the chert/basalt association of the Northeast Complex is, so far as known, the oldest unit of La Désirade Island, the Lesser Antilles region, and the eastern part of the Caribbean plate.

Bearing on the evolution of the Lesser Antilles and the Caribbean plate

As stated in the Introduction, the actual nature of the basement complex of La Désirade Island has been much debated and no final model has yet been established. The “ophiolitic complex” [Mattinson *et al.*, 1973, 1980; Fox & Heezen, 1975; Dinkelman & Brown, 1977, Le Guen de Kerneizon *et al.*, 1979] has also been interpreted as a primitive island arc detached from the Greater Antilles [Fink, 1968, 1970a, 1970b, 1972; Bouysse *et al.*, 1983; Maury *et al.*, 1990; Donnelly *et al.*, 1990]. A more recent study by Gauchat [2004] suggests that La Désirade pillow basalts have a back-arc geochemical signature. Baumgartner *et al.* [2004] stressed that Caribbean radiolarites are associated with mafic and ultramafic igneous rocks of diverse petrogenetic origins, but rarely do they form the stratigraphic cover of MORB-type ocean floor. Although Mattinson & Hopson [*in* Mattinson *et al.*, 2008] stand by their original conclusion that the igneous rock assemblage and vertical distribution on La Désirade are compatible with the upper part of an ophiolite, we do not support this interpretation. In our opinion, La Désirade oceanic rocks display a peculiar spatial combination of volcanics and oceanic deposits

illustrating an interaction of arc-related material and siliceous biogenous sedimentation. This is consistent with the depth of deposition assessed by Bouysse *et al.* [1983] who analyzed the Grand Abaque eruptive center and its submarine scoriae, proposing a maximum lower limit of about 2000 m based on the equilibrium between hydrostatic and critical water pressures.

CONCLUSION

Our investigation of La Désirade Island basement complex brings the following results and interpretations:

- 1) Radiolarian fauna obtained from seven chert localities range in age from late Kimmeridgian-early Tithonian to early or early late Tithonian. It confirms the occurrence of Late Jurassic oceanic strata associated with volcanics of the basement complex. The basalt/chert unit is the oldest on the island, therefore of the entire Lesser Antilles region.
- 2) Whereas former studies reported a single radiolarian age datum on the island, we show that some chert localities are diachronous. La Désirade Jurassic volcanism probably lasted over a longer period than previously established.
- 3) The spatial association of radiolarian-bearing cherts and volcanics does not favour the model of an ophiolite-type oceanic crust, but rather the occurrence of arc-related submarine volcanic activity combined with planktonic sedimentation.
- 4) Tectonic stacking suggested by biostratigraphic data could be linked to an accretion-type tectonic emplacement. The preliminary age distribution provided by chert localities shows a southeast to northwest-directed polarity, potentially indicative of a former eastward subduction zone in the eastern part of the proto-Caribbean plate. This tentative model needs to be tested by future tectonic studies in the region.

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FIGURE CAPTIONS

Figure 1. Structural map of the Caribbean region and the Lesser Antilles arc with location of the Guadeloupe archipelago.

Figure 1. *Schéma structural de la région Caraïbes et de l'arc des Petites Antilles, et localisation de l'archipel de Guadeloupe.*

Figure 2. a: The Lesser Antilles arc and the location of La Désirade Island; b: geological sketch map of La Désirade [after Westercamp, 1980].

Figure 2. *a: Localisation de l'île de La Désirade dans l'arc des Petites Antilles; b: carte géologique simplifiée de La Désirade [d'après Westercamp, 1980].*

Figure 3. Geological map of the Northeast Complex unit of La Désirade Island, location of radiolarian-bearing localities LDE1 to LDE30 and biochronological assignments [UAZ: Unitary Association biozones, Baumgartner *et al.* 1995].

Figure 3. *Carte géologique de l'unité du Complexe de Base du Nord-Est de La Désirade, localités à radiolaires LDE1 to LDE30 et corrélations biochronologiques [UAZ: biozones d'Associations Unitaires, Baumgartner et al. 1995].*

Figure 4. Exposures of sedimentary successions and associated magmatic units from La Désirade “Northeast Volcanic complex”. a: basaltic flows overlying red and brown chert and tuffs at Pointe Mancenillier; b: pillow basalts overlying a thin red chert sequence at Pointe Doublé; c: close-up of chert sequence; d: cross-section of large pillow basalt or flow with interpillow red chert (to right hand side of hammer), north of Pointe Mancenillier; e, f: pillow basalts into red chert sequence near Pointe Doublé; g: contact between basaltic pillows and

overlying red chert and hyaloclastites, north of Pointe Mancenillier; h: interbedded red chert (dark colour) and graded bedded tuffs (light colour), west of Baie Mahault.

Figure 4. *Roches sédimentaires et magmatiques du complexe de base de La Désirade. a: coulées basaltiques recouvrant des jaspes rouges et bruns ainsi que des tuffs, Pointe Mancenillier; b: basaltes en coussins recouvrant une série peu épaisses de jaspes rouges, Pointe Doublé; c: vue rapprochée d'un niveau de jaspes; d: section de coulée basaltique comprenant de larges coussins et des méas de jaspe rouge à radiolaires (à droite du marteau), nord de la Pointe Mancenillier; e, f: coussin basaltique dans une série de jaspes à radiolaires près de la Pointe Doublé; g: contact entre des coussins basaltiques et des jaspes rouges et hyaloclastites, nord de la Pointe Mancenillier; h: interstratification de jaspes rouges et de tuffs granoclassés, ouest de Baie Mahault.*

TABLE CAPTION

Table 1. List of radiolarian taxa and occurrences from La Désirade cherts. LDE3 to LDE30 are productive localities (see text for location); bottom: correlation with UA biozonation from Baumgartner *et al.* [1995] (n.a.: not available).

Table 1. *Liste des taxons de radiolaires de La Désirade et leur occurrence. LDE3 à LDE30 correspondent aux localités diagnostiques (voir texte) ; base : corrélation aux biozones d'Associations Unitaires de Baumgartner et al. [1995] (n.a.: néant).*

PLATE CAPTIONS

Plate 1. Late Jurassic radiolarians from La Désirade cherts (Scanning Electron Microscope).

Figures: taxon, sample number, database picture number, maximum width.

Planche 1. Radiolaires du Jurassique supérieur de La Désirade (Microscope Electronique à Balayage). Figures : taxon, numéro d'échantillon, numéro de cliché de la base de données (« px »), largeur maximale.

1. *Loopus* sp., LDE12, p10, 345 µm; 2-3. *Pseudodictyomitra* spp., LDE12, 3, 520 µm; p4, 420 µm; 4. *Pseudodictyomitra carpatica* (Lozyniak) LDE12, p5, 380 µm; 5. *Loopus primitivus* (Matsuoka & Yao), LDE12, p6, 330 µm; 6. *Archaeodictyomitra excellens* (Tan), LDE9, p5, 385 µm; 7. *Tethysetta* sp. cf. *dhimenaensis*, LDE30, p7, 430 µm; 8. *Archaeodictyomitra* sp., LDE12, p13, 365 µm; 9. *Thanarla* sp., LDE12, p11, 360 µm; 10. *Loopus* sp., LDE12, p18, 365 µm; 11. *Xitus gifuensis* Mizutani, LDE12, p8, 460 µm; 12. *Sethocapsa accincta* Steiger, LDE22, p19, 410 µm; 13-14. *Sethocapsa* sp., LDE22, p17, 375 µm; p24, 345 µm; 15. *Willriedellum* sp., LDE12, p19, 505 µm; 16-17. *Zhamoidellum* sp., LDE12, p23, 465 µm; LDE22, p15, 400 µm; 18. *Zhamoidellum* sp. aff. *ovum* Dumitrica, LDE12, p20, 395 µm.

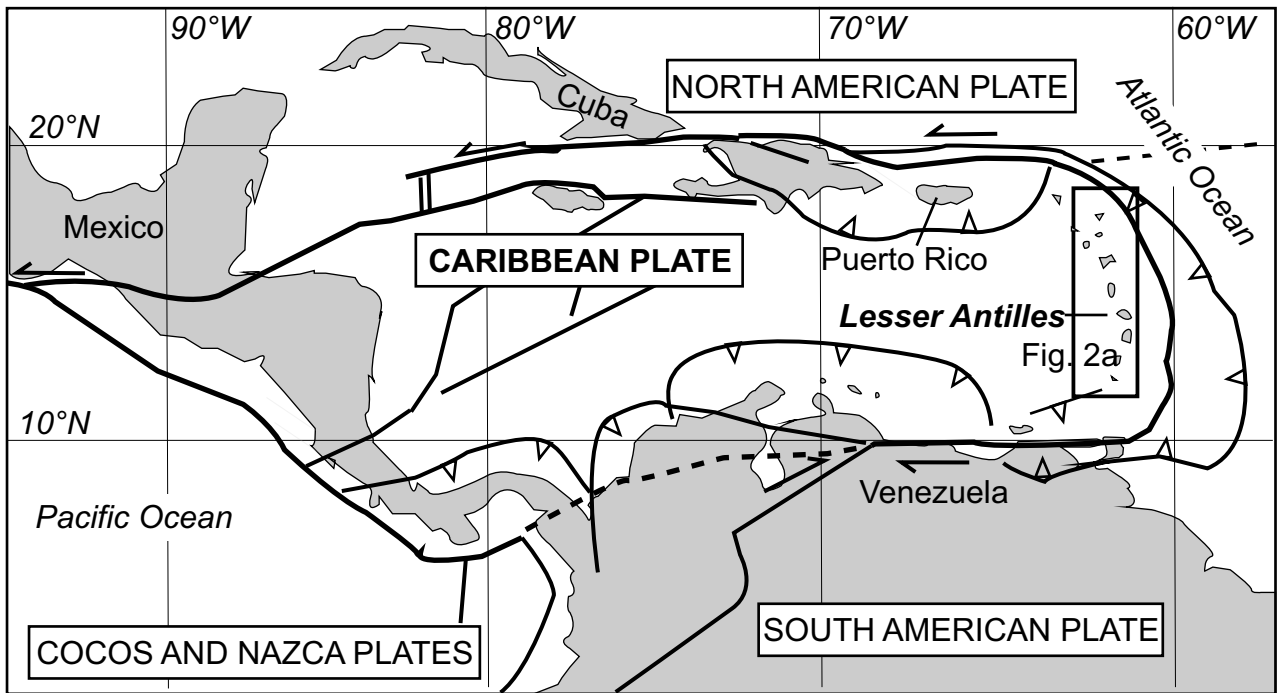
Plate 2. Late Jurassic radiolarians from La Désirade cherts (Scanning Electron Microscope).

Figures: taxon, sample number, database picture number, maximum width.

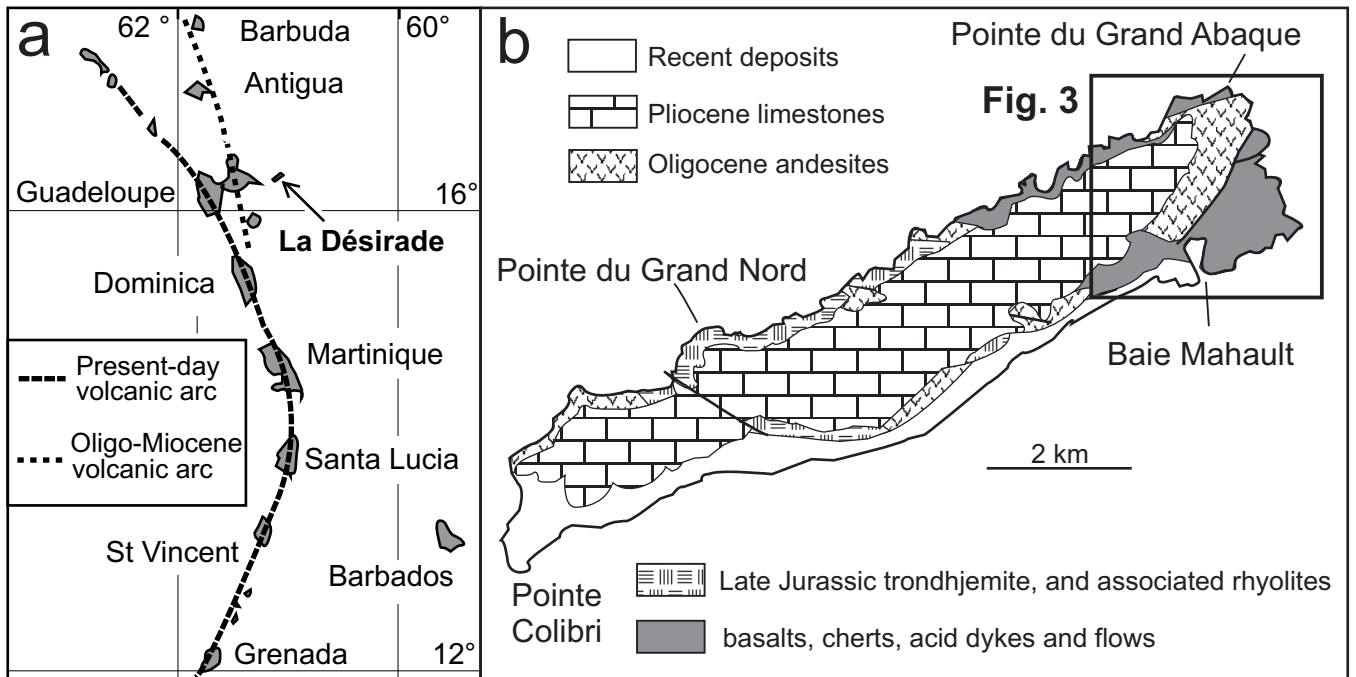
Planche 2. Radiolaires du Jurassique supérieur de La Désirade (Microscope Electronique à Balayage). Figures : taxon, numéro d'échantillon, numéro de cliché de la base de données (« px »), largeur maximale.

1. *Praeconocaryomma* sp., LDE22, p1, 745 µm; 2-3. *Vallupus hopsoni* Pessagno & Blome, LDE22, p11, 370 µm; p12, 310 µm; 4. *Pantanellium whalenae* Pessagno & MacLeod, LDE22, p21, 365 µm; 5. *Pantanellium squinaboli* (Tan), LDE22, p22, 290 µm; 6-7. *Paronaella* spp., LDE30, p5, 1200 µm; p4, 1200 µm; 8. *Angulobracchia* sp., LDE30, p6, 1250 µm; 9.

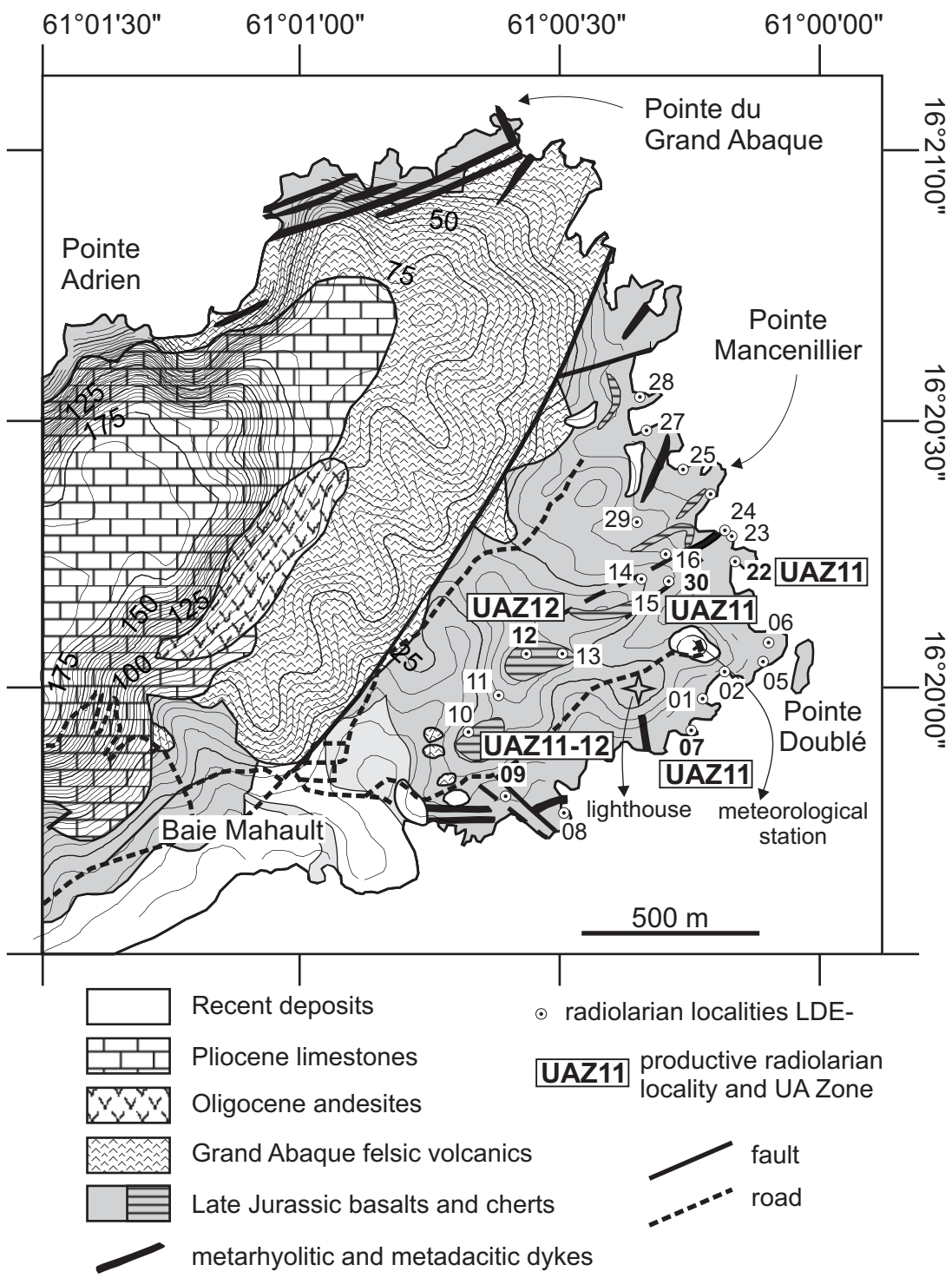
Eucyrtidiellum ptyctum (Riedel & Sanfilippo), LDE7, p5, 120 µm 10. *Eucyrtidiellum pyramis* Aita, LDE12, p14, 200 µm; 11. *Pseudodictyomitrella* sp. aff. *tuscanica* (Chiari, Cortese & Marcucci), LDE12, p2, 550 µm; 12. *Obesacapsula* sp. cf. *verbana* (Parona), LDE30, p3, 1100 µm; 13. *Parvicingula* sp., LDE22, p25, 385 µm; 14. *Hsuum* sp. aff. *cuetaense* Pessagno, LDE30, p11, 470 µm; 15. ?*Eucyrtidiellum* sp., LDE12, p22, 400 µm; 16. ?*Sethocapsa* sp., LDE22, p3, 580 µm; 17. *Zhamoidellum* sp., LDE12, p17, 375 µm.



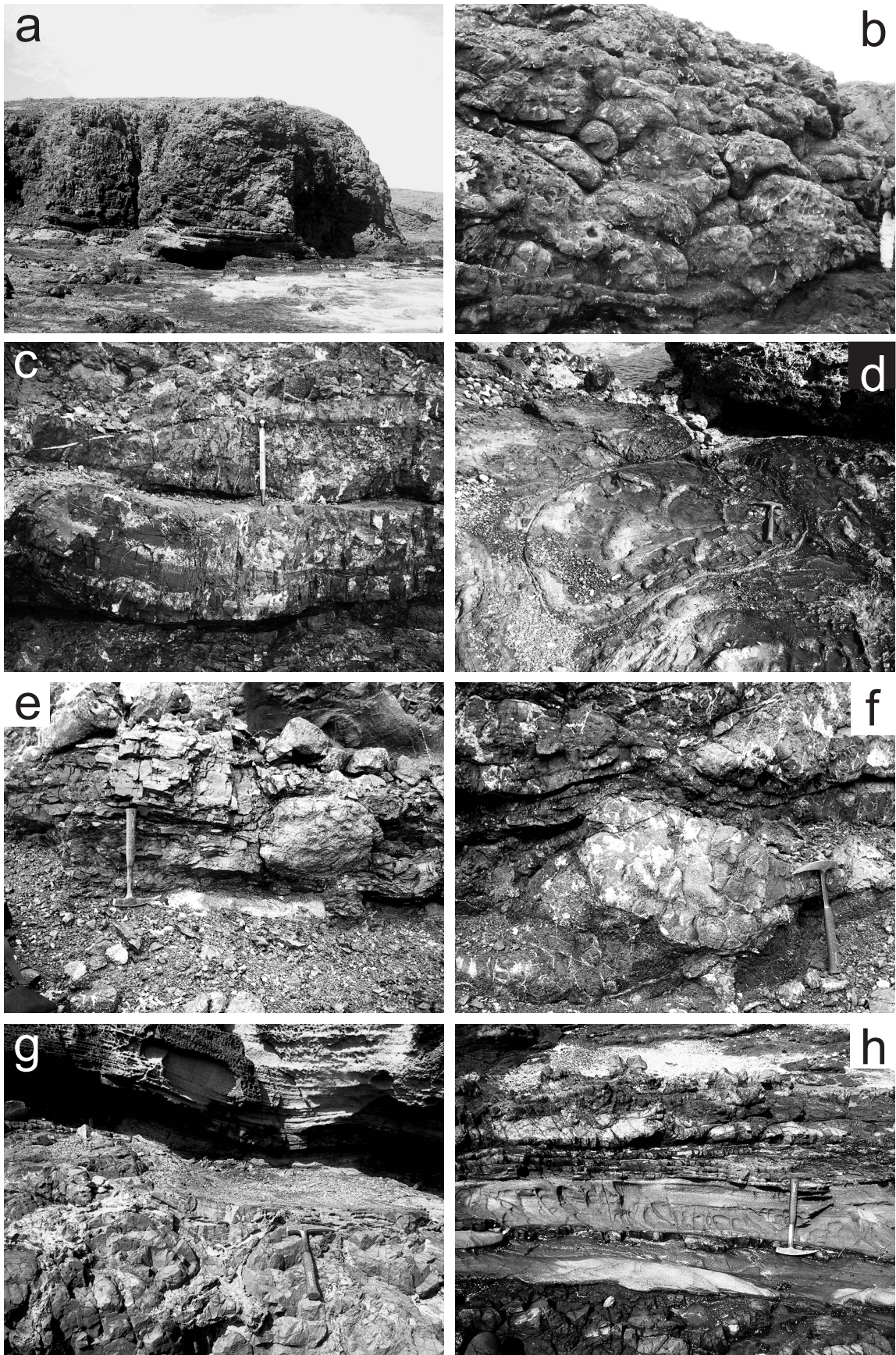
Cordey & Cornée, Fig. 1



Cordey & Cornée, Fig. 2



Cordey & Cornée, Fig. 3



Cordey & Cornée, Fig. 4

Radiolarian taxa	LDE3	LDE7	LDE9	LDE12	LDE22	LDE27	LDE30
<i>Angulobracchia</i> sp.							X
<i>Archaeodictyomitra excellens</i> (Tan) 11-22			X				
<i>Archaeodictyomitra</i> sp. aff. <i>excellens</i> (Tan)				X			
<i>Archaeodictyomitra minoensis</i> (Mizutani) 9-12			X	X			
<i>Archaeodictyomitra</i> sp.				X		X	
<i>Emiluvia chica</i> Foreman 3-18					X		
<i>Emiluvia</i> sp. cf. <i>salensis</i> Foreman 4-13				X			
<i>Eucyrtidiellum</i> sp. aff. <i>nodosum</i> Wakita					X		
<i>Eucyrtidiellum ptyctum</i> (Riedel & Sanfilippo) 5-11		X					
<i>Eucyrtidiellum pyramis</i> (Aita) 12-13				X			
<i>Eucyrtidiellum</i> sp.			X				X
<i>Homoeoparonaella argolidensis</i> Baumgartner 4-11					X		
<i>Hsuum</i> sp. aff. <i>cuetaense</i> Pessagno 10-13							X
<i>Loopus primitivus</i> (Matsuoka & Yao) 7-12			X	X			
<i>Obesacapsula verbana</i> (Parona) 11-20		X					
<i>Obesacapsula</i> sp.							X
<i>Orbiculiformella</i> sp.							X
<i>Pantanellium whalenae</i> Pessagno & MacLeod					X		
<i>Pantanellium squinaboli</i> (Tan) 11-22					X		
<i>Paronaella</i> spp.							X
<i>Parvicingula</i> sp.					X		X
<i>Praeconocaryomma</i> spp.		X			X		X
<i>Pseudodictyomitra carpatica</i> (Lozynyak) 11-21			X	X			
<i>Pseudodictyomitra</i> sp.				X	X		
<i>Pseudodict.</i> sp. aff. <i>tuscanica</i> (Chiari, Cortese & Marcucci)				X			
<i>Sethocapsa accincta</i> Steiger					X		
<i>Sethocapsa</i> sp. cf. <i>accincta</i> Steiger							X
<i>Sethocapsa horokanaiensis</i> Kawabata					X		
<i>Sethocapsa</i> sp. cf. <i>horokanaiensis</i> Kawabata							X
<i>Sethocapsa</i> sp.				X	X		
<i>Stichocapsa</i> sp.				X	X		
? <i>Stylocapsa</i> sp.				X			
<i>Tethysetta</i> sp. cf. <i>dhimenaensis</i> (Baumgartner) 3-11							X
<i>Thanarla</i> sp.				X			
<i>Vallupus hopsoni</i> Pessagno & Blome					X		
<i>Vallupus</i> sp.	X						
<i>Williriedellum</i> sp.				X			X
<i>Xitus gifuensis</i> Mizutani 11-18				X			
<i>Zhamoidellum ovum</i> Dumitrica 9-11					X		
<i>Zhamoidellum</i> sp. aff. <i>ovum</i> Dumitrica				X			
<i>Zhamoidellum</i> sp.				X	X		
Unitary Association Zones	n.c.	11	11-12	12	11	n.c.	11

Cordey & Cornée Table 1

