

# Timing of Cache Creek Ocean closure: insights from new Jurassic radiolarian ages in British Columbia and Yukon and their significance for Canadian Cordillera tectonics

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### Abstract

The closure of the Cache Creek Ocean during the Jurassic led to the structuration of the Canadian Cordillera. To reassess the timing of this major tectonic event, this study proposes a biochronological review of all the youngest radiolarian-bearing localities of the Cache Creek Terrane in British Columbia and Yukon using updated radiolarian biozonations and taxonomy, complemented by new data from the type locality. The results show that all ages are Early Jurassic and fairly homogenous along the entire terrane. All the youngest siliceous rocks are restricted to the Late Sinemurian-Early Toarcian interval (~195-180 Ma). The youngest welldated pelagic cherts are Early Pliensbachian, with slightly younger siliceous argillites of Pliensbachian and possibly Early Toarcian age. These results suggest that Cache Creek deep pelagic sedimentation ceased fairly synchronously along the entire belt, probably during the Pliensbachian (~191–183 Ma), predating the final closure of the basin when Cache Creek units were thrust over Stikinia and Quesnellia in the Middle and Late Jurassic. This apparent synchronicity should result in re-evaluating the tectonic mechanism of the Cache Creek "entrapment model" in its present form. The study also shows that pelagic sedimentation ceased 22 to 31 m.y. earlier in the Cache Creek Terrane than in the Bridge River Terrane, confirming the stratigraphic discrepancy between the two terranes during the Jurassic.

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#### Résumé

La fermeture de l'océan Cache Creek au Jurassique a conduit à la structuration de la Cordillère Canadienne. Afin de réévaluer la chronologie de cet événement tectonique majeur, l'étude propose une synthèse biochronologique de toutes les plus jeunes localités à radiolaires du terrane de Cache Creek en Colombie-Britannique et au Yukon grâce aux nouvelles biozonations et à une taxonomie actualisée, avec en complément de nouvelles données pour la localité-type. Les résultats montrent que tous les âges sont du Jurassique inférieur et sont homogènes pour l'ensemble du terrane. Toutes les roches siliceuses les plus jeunes sont comprises dans l'intervalle Sinémurien supérieur—Toarcien inférieur (~ 195–180 Ma). Les plus jeunes cherts pélagiques sont du Pliensbachien inférieur, avec des argilites siliceuses légèrement plus jeunes du Pliensbachien et peut-être du Toarcien inférieur. Ces résultats suggèrent que la sédimentation pélagique profonde de l'océan Cache Creek a cessé de manière relativement synchrone sur toute sa longueur, probablement pendant le Pliensbachien (~191–183 Ma), avant la fermeture finale du bassin lorsque les unités de Cache Creek ont été charriées sur la Stikinia et la Quesnellia au Jurassique moyen et supérieur. Ce synchronisme apparent devrait conduire à réévaluer les modalités du modèle tectonique en ciseau sous sa forme actuelle. L'étude montre aussi que la sédimentation pélagique a cessé 22 à 31 Ma plus tôt dans le terrane de Cache Creek que dans celui de Bridge River, confirmant la différence stratigraphique entre les deux unités au Jurassique. Mots-clés: terrane de Cache Creek, Radiolaria, Jurassique, Cordillère canadienne, radiolarite, terrane de Bridge River

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### Introduction

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The Canadian Cordillera (Fig. 1) is a complex collage of arc and oceanic terranes which has been the subject of a vast amount of studies and literature in the past 50 years. Various models describing the tectonic history of the orogen are still debated (Monger & Gibson 2019 and references therein) but they all point to the importance of the Cache Creek Terrane, considered as one major evidence for a large Mesozoic ocean basin in the Canadian Cordillera. One important issue is the timing of the ocean closure during the Jurassic period which led to collision with Stikinia and Quesnellia volcanic arcs, contributing to the structuration of the Cordillera (Nelson and Mihalynuk 1993, Colpron et al. 2015, Monger & Gibson 2019). To determine the age of the Cache Creek Terrane, radiolarian biostratigraphy has been applied over the last 40 years to its deep ocean strata with a particular focus on the youngest localities considered to predate the final stages of convergence. The youngest Jurassic unit of the Cache Creek Terrane was originally found in its type locality in southern British Columbia (Cordey et al. 1987). At the time, it contradicted the scenario in which amalgamation of the Cache Creek and Ouesnellia terranes occurred in the Late Triassic (Monger 1981, Monger et al. 1982). Further investigations confirmed the occurrence of Jurassic radiolarian-bearing oceanic strata in all segments of the Cache Creek Terrane in British Columbia and Yukon (Fig. 2; Cordey et al. 1991, Cordey & Struik 1996, Cordey 1998, Orchard et al. 2001, Struik et al. 2001, Mihalynuk et al. 2004). Although most of these localities are Early Jurassic, some confusion remained due to the "Early or Middle Jurassic" age of the youngest Cache Creek rocks in the type locality. This question also led to conflicting interpretations in comparisons between the Cache Creek Terrane and other Paleozoic-Mesozoic oceanic units of the Canadian Cordillera such as the Bridge River and Hozameen terranes (Cordey et al. 1987, Cordey & Schiarizza 1993, Mihalynuk et al. 1994, Cordey 1998, Monger 2014).

To clarify the stratigraphy of the youngest Cache Creek strata, the goal of this study is to review all the youngest radiolarian localities and, when possible, improve and refine previous ages with new Jurassic radiolarian biozonations and taxonomy (Baumgartner et al. 1995, Carter et al. 1988, 1998, 2010, Goričan et al. 2006, 2018, O'Dogherty et al. 2009). The results reveal that the youngest ages are fairly homogenous along the entire Cache Creek belt, which provides some clues regarding the accuracy of current tectonic models such as the Cache Creek Terrane entrapment in its present form (Mihalynuk et al. 1994). Finally, the study provides some new elements for stratigraphic comparison of the Cache Creek Terrane with other main oceanic units of the Canadian Cordillera during the Jurassic (Cordey and Schiarizza 1993, Monger and Gibson 2019).

### **Material and Methods**

The radiolarian assemblages presented in this paper are part of a large dataset of Devonian to Jurassic radiolarian samples collected and studied in several terranes of the Canadian Cordillera since the 1980s. Among them, the youngest Jurassic assemblages of the Cache Creek Terrane were found at eight localities from southern British Columbia to southern Yukon (Fig. 2: J1, Pavilion; J2, Alkali Lake; J3, Blackwater; J4, Pinchi Lake; J5, French Range; J6-J8, Teslin Plateau). The geological setting and lithostratigraphy of these localities are described in detail in former publications (J1: Mortimer 1987, Cordey et al. 1987, Cordey 1998; J2: Read 1993, Cordey, unpub. data; J3: Cordey and Struik 1996, Orchard et al. 2001; J4: Orchard et al. 2001, Struik et al. 2001; J5: Mihalynuk et al. 2004; J6-J8: Cordey et al. 1991).

The radiolarian faunas from the Cache Creek Terrane are usually poorly or moderately preserved due to silica recrystallization as a result of diagenesis associated with P/T conditions during

tectonic emplacement and local magmatic intrusions. To maximize the quality of the microfossils, the samples were selected with specific field detection techniques applied to pelagic strata from orogenic belts, using a strong hand lens and investigating specific microfacies such as argillaceous layers within chert beds (Cordey & Krauss 1990, Cordey 1998). Radiolarians were then extracted by repetitive leaching of samples with low-concentration hydrofluoric acid (HF) and then hand-picked and mounted on aluminum stubs for SEM observation and taxonomical identifications (Tabletop SEM Phenom ProX, Laboratoire de Géologie de Lyon). Although most radiolarian assemblages presented here are described in previous publications, new SEM images have been obtained for a selection of samples to improve the quality of previous illustrations and complement former studies with new taxa. When microfossil assemblages are not well-preserved, radiolarians are usually identified at the genus level only (J1, J2, J5, J7, J8; Table 1). In this case, the age is established from overlapping ranges of genera (O'Dogherty et al. 2009). Other assemblages show a better state of preservation with radiolarians identified at the species level (J3, J4, J6), allowing correlations to more precise biozones (Carter et al. 1988, Cordey 1998, Whalen and Carter 2002, Goričan et al. 2006, Carter et al. 2010). A selection of biostratigraphically-significant radiolarians is presented in Figures 5 and 6. Radiolarian assemblages and ages are synthesized in Table 1.

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# 134 Results

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### New radiolarian assemblage from the type locality

In southern British Columbia, the Cache Creek Terrane is divided into three belts (Figs. 3, 4)

(Duffell and McTaggart 1952, Trettin 1980, Shannon 1981, Monger and McMillan 1984, Monger 1985, Mortimer 1987). The central belt is composed of massive limestone and marble of the

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Marble Canyon Formation where schwagerinid fusulinids of Late Permian age (Colania assemblage) belong to a Tethyan faunal province and are cited as evidence that carbonate units of the Cache Creek Terrane are far-travelled (Monger and Ross 1971; Monger et al. 1982; Danner 1985; Orchard et al. 2001). To the west of Marble Canyon, the western belt contains the youngest strata of the type locality as shown by Early or Middle Jurassic radiolarians found in siliceous argillites above Hambrook Creek ("R2", Cordey et al. 1987, Cordey 1998). With the aim of improving this previous radiolarian dating, this locality was resampled (NEO-R2, Figs. 3, 4). A new radiolarian assemblage J1 (Table 1) is composed of the following taxa: Archaeodictyomitra sp., Homoeoparonaella sp. cf. reciproqua Carter, Hsuum sp., Pseudocrucella? sp., Thurstonia sp., Trillus sp., Xiphostylus? sp., Triactoma sp., and Zartus sp.. The previous collection was dated as Pliensbachian–Bajocian based on the age range of Zartus sp. (Cordey et al. 1987). In the new sample NEO-R2, the co-occurrence of *Thurstonia* sp., *Trillus* sp. and Zartus sp. provides an age restricted to the Early Pliensbachian–Early Toarcian interval (O'Dogherty et al. 2009). Although the preservation of the assemblage is poor, the specimens of Thurstonia sp. display well-developed nodes at pore frame intersections (Fig. 5, no. 1), showing affinities with *Thurstonia* specimens from the northern part of the Cache Creek Terrane (J6, Fig. 5, no. 3) as well as *Thurstonia* sp. B documented in Early Jurassic radiolarian assemblages from the Bridge River and Hozameen terranes (Cordey 1998). Refined Jurassic radiolarian ages Southern British Columbia An Early Jurassic radiolarian assemblage was found in gray-green radiolarian chert of the Riske Creek Formation in the Alkali Lake map area (locality J2, Fig. 2) as a result of mapping and sampling by Read (1993). Such as J1, it is located within the western belt of the terrane. This assemblage is composed of Canutus sp., Lantus sp., Orbiculiformella sp., Parahsuum sp. and

Praeconocaryomma sp. (Table 1). Originally dated as Sinemurian or Pliensbachian, its age is refined to the Early Pliensbachian-Early Toarcian interval (Table 1) based on co-occurrence of the genera Canutus sp. and Lantus sp.. (O'Dogherty et al. 2009). It is noteworthy that the age ranges of J1 and J2 assemblages are identical, although based on different radiolarian taxa. Central British Columbia Geological mapping of the Cache Creek Terrane in central British Columbia (Struik et al. 2007) and references therein) resulted in an important number of Paleozoic and Mesozoic radiolarian collections (Cordey and Struik 1996, Struik 1998, Orchard et al. 2001). The youngest radiolarian chert succession of this region is exposed along the Blackwater River in Punchaw Lake map area (J3, Fig. 2), 25 km to the south-southwest of the town of Prince George (Cordey and Struik 1996, Orchard et al. 2001, Struik et al. 2001). This succession is broadly continuous and contains Middle Triassic, Late Triassic, and Early Jurassic radiolarians. The youngest radiolarian assemblage found near the top of the section was initially dated as Pliensbachian (Cordey and Struik 1996, Orchard et al. 2001). Complementary chemical processing and new SEM images reveal an abundant and well-preserved assemblage (Table 1) composed of *Atalantria* sp., Beatricea? argescens Cordey, Bipedis japonicus Hori, Broctus sp., Canoptum rugosum Pessagno & Poisson, Canoptum sp. cf. dixoni Pessagno & Whalen, Charlottea sp., Katroma sp. cf. elongata Carter, Lantus sp. cf. praeobesus Carter, Laxtorum sp., Orbiculiformella? trispina s.l. (Yeh), Orbiculiformella sp., Palaeosaturnalis sp., Parahsuum simplum Yao, Parahsuum vizcainoense Whalen & Carter, Praeconocaryomma sp. cf. sarahae Carter, Thurstonia sp. and Spumellaria gen. sp. indet. (Figs. 5, 6). This assemblage corresponds to the interval from the middle part of the Canutus tipperi – Katroma clara Zone to the base of the Gigi fustis – Lantus sixi Zone of Early Pliensbachian age (Carter et al. 2010).

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In the Fort-St-James area, the Tezzeron succession (originally called the Pinchi succession by Struik et al. (1996) and Cordey and Struik (1996)) is exposed to the north of Stuart and Pinchi lakes and bears some similarities with clastic assemblages from the upper part of the Cache Creek Terrane in southern British Columbia (Shannon 1981, Monger 1985) and in Yukon (Cordey et al. 1991, Gordey and Stevens 1994). It is interpreted as belonging to the Cache Creek Terrane (Bellefontaine et al. 1995, Struik et al. 1996, Orchard et al. 2001, Struik et al. 2001). Within the Tezzeron succession, a limy siltstone unit located near the south end of Pinchi Lake contains an Early Jurassic radiolarian assemblage (J4, Fig. 2, Table 1) previously reported by Orchard et al. (2001) and Struik et al. (2001). It is composed of *Emiluvia*? sp. A (Carter et al. 1988), Hsuum lucidum Yeh (formerly interpreted as H. sp. cf. optimum Carter), Paronaella grahamensis Carter, Praeconocaryomma decora gr. Yeh (formerly P. sp. A), and Praeparvicingula tlellensis Carter (formerly Parvicingula sp. E, Carter et al. 1988). This assemblage was originally assigned to the biozones 2 to 5 of Carter et al. (1988). It correlates with the *Elodium pessagnoi – Hexasaturnalis hexagonus* Zone of Carter et al. (2010) of Middle-Late Toarcian age. The corresponding sample, which is a limy siltstone, cannot be considered part of "deep ocean" Cache Creek strata. It shows, however, that radiolarian pelagic input lasted until the Toarcian, at least locally. Northern British Columbia In the French Range to the west of the Dease Lake area, a unit belonging to the Cache Creek Terrane is composed of cherty argillites with local tuffaceous layers that contain Middle Triassic and Early Jurassic radiolarians (Mihalynuk et al. 2004). This succession is correlative to the rocks on the Teslin Plateau, ~200 km to the north (Cordey et al. 1991). One Early Jurassic radiolarian assemblage (J5, Fig. 2, Table 1) was extracted from a black cherty argillite exposure located 2.5

km to the north of Slate Creek. It is from blueschists, and therefore places a maximum age limit

215 on blueschist mineral formation ((Mihalynuk et al. 2004). The assemblage is composed of Parahsuum izeense (Pessagno & Whalen) (formerly Canutus izeensis Pessagno & Whalen), 216 Orbiculiformella? trispina s.l. (Yeh) (formerly Orbiculiforma silicatilis Cordey), 217 Praeconocaryomma sp. cf. sarahae Carter (formerly P. aff. media Pessagno & Poisson), 218 Praeconocaryomma sp., and Thurstonia sp. (Table 1). Initially, the radiolarian assemblage was 219 dated as Pliensbachian-Toarcian (Mihalynuk et al. 2004). The age is revised to the 220 Pliensbachian-Early Toarcian based on the occurrence of *Parahsuum izeense* Pessagno & 221 Whalen and the refined range of *Thurstonia* (Table 1). The presence of *Praeconocaryomma* sp. 222 223 cf. sarahae Carter suggests that this assemblage could be restricted to the Pliensbachian (Carter et 224 al. 2010), but this cannot be confirmed due to the incomplete nature of the specimen. 225 226 Yukon 227 An assemblage composed of radiolarian cherts and fine-grained clastic rocks is exposed on the 228 Teslin Plateau, ~20 km to the north of the British Columbia-Yukon border (Gordey 1991). This 229 succession has previously been called the Lewes River assemblage (Wheeler et al. 1991, Gordey 1991) or Teenah Lake assemblage (Jackson 1992), and is equivalent in part to the Kedahda 230 231 Formation of Monger (1975). Several localities with Middle Triassic, Late Triassic, and Early 232 Jurassic radiolarian assemblages were discovered along a five-kilometre-long ridge north of Mount Bryde and Sterlin Lake (Gordey 1991, Cordey et al. 1991). The ridge consists of 233 successive outcrops and isolated sections of stacked radiolarian chert, argillite and graywacke 234 235 previously described and discussed in detail (Cordey et al. 1991). The first Early Jurassic radiolarian assemblage J6 (Fig. 2, Table 1) is from siliceous argillite 236 interbedded with graywacke that contain *Bipedis* sp., *Broctus* sp., Canoptum anulatum Pessagno 237 & Poisson, Canoptum sp., Hagiastrum majusculum Whalen & Carter (formerly called H. sp. A, 238 Cordey 1998), Homoeoparonaella sp., Lantus sixi Yeh, Napora cerromesaensis Pessagno, 239

240 Whalen & Yeh, Pantanellium cumshewaense Pessagno & Blome, Parahsuum izeense (Pessagno & Whalen) (formerly described as Canutus giganteus Pessagno & Whalen), Praeconocaryomma? 241 sp. cf. vakounensis Carter (formerly interpreted as *Praeconocaryomma*. sp. aff. P. sp. B Yeh), 242 Thurstonia sp., and Wrangellium sp.. This assemblage was originally dated as 243 Pliensbachian–Early Toarcian (Table 1). It is now correlated with the base of the Gigi fustis – 244 Lantus sixi Zone to the middle part of the Eucyrtidiellum nagaiae – Praeparvicingula tlellensis 245 Zone (Unitary associations UA 12-22) (Carter et al. 2010), establishing a late Early Pliensbachian 246 to Late Pliensbachian age. 247 248 The second youngest Early Jurassic sample on the Teslin Plateau (J7, Fig. 2, Table 1) is from a 249 radiolarian chert succession located 4 km to the south of J6 (Cordey et al. 1991). It was previously dated as Sinemurian-Toarcian (Cordey et al. 1991). It is composed of *Beatricea*? 250 251 argescens Cordey, Canoptum sp. cf. anulatum Pessagno & Poisson (previously interpreted as 252 Wrangellium sp. cf. thurstonense Pessagno & Whalen), Canoptum sp., Praeconocaryomma sp. 253 cf. immodica Pessagno & Poisson, and Wrangellium sp. (Table 1). The age is here refined to the 254 Late Sinemurian-Late Pliensbachian interval based on the co-occurrence of *Beatricea* sp. and 255 Wrangellium sp.. 256 The third sample (J8, Fig. 2, Table 1) is part of the same radiolarian chert succession and location as J7. The radiolarian assemblage is composed of *Beatricea? argescens* Cordey, *Bipedis* sp., 257 Canoptum sp. cf. dixoni Pessagno & Whalen, Gorgansium sp., Pantanellium sp., Parahsuum sp., 258 Praeconocaryomma sp., and Wrangellium sp.. It was previously dated as Sinemurian-Toarcian 259 260 (Cordey et al. 1991). It is here refined to the Late Sinemurian–Late Pliensbachian interval, based on the co-occurrence and refined ranges of *Beatricea* sp., *Bipedis* sp. and *Wrangellium* sp. (Table 261 1). 262 263

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Homogeneity of youngest Cache Creek ages

Revised Jurassic radiolarian ages reveal that the youngest Cache Creek ages are fairly homogenous along the entire Canadian Cordillera (Fig. 7). These ages are comprised within the Late Sinemurian—Late Toarcian interval (~195–174 Ma, Cohen et al. 2013). The siliceous rocks (radiolarian chert and siliceous argillite) are restricted to the Late Sinemurian—Early Toarcian interval (~195–180 Ma), including error margins. The most precise age determinations from central British Columbia and Yukon (Fig. 7) indicate that the youngest true radiolarian cherts (= devoid of clastic intercalations) are Pliensbachian in age (J3), and the youngest siliceous argillites may also be Pliensbachian in age, possibly reaching the Early Toarcian (J1, J5, J6). The youngest Jurassic age obtained on a limy siltstone in the Tezzeron succession (J4) suggests that radiolarian pelagic input possibly lasted until the Middle or Late Toarcian in central British Columbia, although chert sedimentation might have ceased. To date, no pelagic sedimentary strata of Middle Jurassic age have been found in the Cache Creek Terrane.

# Cache Creek Terrane and overlap assemblages

Some of the youngest Cache Creek strata have locally been interpreted as overlap assemblages, for instance the chert-argillite-graywacke successions exposed on the Teslin Plateau (Lewes River assemblage, Wheeler et al. 1991, Gordey 1991; Teenah Lake assemblage, Jackson 1992). More recently, all post-Middle Triassic sequences were proposed to represent overlap assemblages on top of already assembled terranes including the Paleozoic units of the Cache Creek Terrane (Zagorevski et al. 2017).

The notion of an "overlap assemblage" implies that marine sedimentation ceased prior to the deposition of clastic sediments on top of exhumed oceanic units. Regarding the Cache Creek

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Terrane, this view is not consistent with the occurrence of radiolarian chert sedimentation as young as Pliensbachian or of Middle Jurassic blueschists with in part Early Jurassic sedimentary protoliths (Mihalynuk et al. 2004), which demonstrate that deep ocean sedimentation persisted between Stikinia and Quesnellia until the end of the Early Jurassic. In central British Columbia (Blackwater; Figs. 7, 8), a Triassic–Early Jurassic chert succession also suggests that Cache Creek marine sedimentation was continuous prior to the closure of the basin. However, pelagic and clastic regimes are interrelated: in the northern part of the Cache Creek Terrane, the argillites and graywackes are locally intercalated with Late Triassic radiolarian cherts (Cordey et al. 1991). More recently, a Hettangian age (199.5  $\pm$  2.0 Ma) was obtained on detrital zircons from a Cache Creek graywacke east of Carcross (Colpron et al. 2015) and interpreted as a probable distal equivalent of the Richtofen Formation (Laberge Group), showing a probable link between the Cache Creek Terrane and the Whitehorse trough as early as the end of the Triassic or the beginning of the Jurassic. The input of clastic sediments within Cache Creek pelagic strata may be a result of the geometry of the terrane at this time, flanking the Yukon-Tanana Terrane to the north during the counterclockwise rotation of Stikinia ("isolation stage" of Mihalynuk et al. 1994; Colpron et al. 2015). These transitional regimes evolved towards a narrowing Cache Creek basin incorporating thicker and coarser siliciclastic sediments associated with tectonic convergence in the form of synorogenic piggyback basins (Colpron et al. 2015, Bickerton et al. 2020). Numerous examples of pelagic to clastic successions are found in other segments of the circum-Pacific ranges such as in Japan and Siberia where the Permian-Jurassic oceanic terranes display Jurassic intercalations of radiolarian chert, siliceous argillite and sandstone interpreted as ocean environments within the depositional influence of one continental margin or flanking sources (Isozaki et al. 1990, Matsuoka et al. 1996). In general, the intercalations and lateral facies changes between pelagic and clastic rocks make it difficult to delineate a clear boundary where oceanic terrane sedimentation ends and where clastic assemblages begin. This is true for the

Cache Creek Terrane in the Yukon, but also in southern British Columbia (Mortimer 1987, Cordey et al. 1987). The composite nature of Cordilleran oceanic terranes, i.e. the likelihood for these terranes to be the end-result of juxtaposed basinal entities, has previously been suggested for the Bridge River Terrane (Cordey and Schiarizza 1993, Cordey 1998) and the Cache Creek Terrane (Golding 2018) but these hypotheses do not challenge the interpretation that these terranes were dominantly oceanic until the final stages of convergence.

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# Cache Creek closure: fast entrapment?

Several models have been proposed for the tectonic mechanism of Cache Creek Ocean closure (see Golding 2018 for a recent review). The consensus model of Cache Creek "entrapment" from Nelson and Mihalynuk (1993) and Mihalynuk et al. (1994) proposes that the terrane became enclosed between the Stikinia and Quesnellia terranes leading to Cache Creek oceanic crust being subducted during an oroclinal bending event, with a progressive "scissor-like" closure from north to south, and the Yukon-Tanana Terrane acting as the hinge. In this scenario, the Whitehorse trough is considered a northern precursor to the Bowser Basin (Ricketts et al. 1992, Evenchick et al. 2010) accompanying the final stages of convergence and closure (Colpron et al. 2015). The entrapment model implies a diachronous closure of the Cache Creek Ocean, with the youngest oceanic sedimentary rocks progressively younging to the south (basin isolation followed by closure, Mihalynuk et al. 1994). The updated radiolarian ages suggest instead a relative synchronicity in the cessation of pelagic sedimentation along the entire Cache Creek Terrane. While Yukon hemipelagic rocks (J6) are Pliensbachian in age, a similar age (Pliensbachian-Early Toarcian) is obtained on siliceous argillites 1200 km to the south in the type locality (J1). Additionally, radiolarian cherts from central British Columbia (J3) and Yukon (J8) could have a common Pliensbachian age, even though they are ~1000 km apart. One could argue that the speed and geometry of the counterclockwise oroclinal bending of

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Stikinia towards Quesnellia are not precisely known. One way to reconcile the entrapment model with the relative synchronicity of the youngest Cache Creek oceanic strata is through very rapid tectonic rotation. In northern British Columbia, Mihalynuk et al. (1994, 2004) assessed that the final closure of the Cache Creek Terrane in the north occurred before 171 Ma (late Aalenian or early Bajocian, Cohen et al. 2013), the age of the post-kinematic Fourth of July batholith that intrudes deformed Cache Creek units in the Atlin Lake area (Mihalynuk et al. 1992). The minimum age of Cache Creek extrusion in the French Range was constrained to ~174.5–172 Ma (latest Toarcian or Aalenian, Cohen et al. 2013), using the interval between the youngest blueschists and the oldest intrusions (Mihalynuk et al. 2004). In the south, the Cache Creek Terrane was thrust eastward over Quesnellia in the Late Jurassic ( $\leq 160$  Ma) (Travers 1978). The most precise ages of the youngest Cache Creek Terrane strata are Pliensbachian and found in central British Columbia (J3) and Yukon (J6). The Early Pliensbachian chert J3 (Figs. 7, 8) found at the top of a Triassic-Early Jurassic succession along the Blackwater River in central British Columbia is a good time proxy for the age of the youngest and still truly oceanic Cache Creek sedimentation at the center of the Cache Creek belt, suggesting that the youngest sustained hemipelagic sedimentation in southern British Columbia could be slightly younger, possibly Late Pliensbachian or Early Toarcian, predating the final closure of the Cache Creek Ocean. In this case, the final phase of the counterclockwise rotation of Stikinia must have occurred within 8 m.y. (the interval between late Early Pliensbachian and Early Toarcian, ~188–180 Ma, Cohen et al. 2013).

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### Comparison between youngest Cache Creek and Bridge River ages

Since the 1980s, geological and paleontological studies have established that the oceanic strata of the Cache Creek and the Bridge River terranes (Fig. 1) share fairly similar Carboniferous-Jurassic stratigraphic ages (Cordey et al. 1987, Mihalynuk et al. 1994, Monger 2014). The Hozameen

Terrane exposed in southern British Columbia and northern Washington (Fig. 1) is considered as a southern continuation of the Bridge River Terrane (Monger 1985). Early Jurassic radiolarian assemblages found in the Cache Creek Terrane are also present in the Bridge River and Hozameen terranes which contain Pliensbachian and Toarcian radiolarian cherts (Cordey and Schiarizza 1993, Cordey 1998). However, the cherts of the Bridge River Terrane document a pelagic sedimentation that lasted significantly longer: the youngest radiolarian assemblages are from the Noaxe Creek area (Cordey and Schiarizza 1993, Cordey 1998) where six localities are Middle Jurassic (Aalenian–Bajocian, Early–Middle Bathonian, Late Bathonian-Middle Callovian, Late Bathonian-Late Callovian), and one locality is Middle or Late Jurassic (Late Callovian-?Middle Oxfordian) based on biozonations for the Middle-Late Jurassic (Baumgartner et al. 1995, Matsuoka 1995, Goričan et al. 2018). This Bridge River radiolarian chert succession probably grades into Late Jurassic and earliest Cretaceous clastic rocks of the Tyaughton-Methow basin (Mahoney and Journeay 1993, Cordey 1996, Monger 2014, Monger and Gibson 2019). It is hypothesized here that the youngest radiolarian chert of the Cache Creek Terrane is Early Pliensbachian in age (~191–187 Ma, Cohen et al. 2013) based on the well-constrained assemblage at the Blackwater section in central British Columbia (J3, Figs. 7-8). In the Bridge River Terrane, the youngest age for a radiolarian chert is within the "Late Callovian – Middle Oxfordian" interval (~165–160 Ma, Cohen et al. 2013). These ages imply that the Cache Creek Terrane oceanic sedimentation had ceased since 22 to 31 m.y. (minimum and maximum possible durations due to the error margins of biostratigraphic data, corresponding to "Late Pliensbachian-Late Callovian" and "Early Pliensbachian-Middle Oxfordian" intervals, respectively) while the Bridge River Ocean was still recording pelagic sedimentation in the Middle Jurassic and possibly the early Late Jurassic.

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### Cache Creek and Bridge River terranes: one or two sutures?

The question whether the Cache Creek and Bridge River terranes mark two distinct sutures or are the remnants of a single oceanic domain has long been debated (Monger et al. 1972, Coney et al. 1980, Monger et al. 1982, Monger 1985, Cordey 1986, Cordey et al. 1987, Monger et al. 1994, Nokleberg et al. 2000, Dickinson 2004, Johnston and Borel 2006, Gehrels et al. 2009). This issue has been revived by Sigloch and Mihalynuk (2013, 2017) who combined high-resolution tomography, plate reconstructions and the concept of vertical slab walls. Their analysis of the North American Cordillera implies that a cryptic suture zone existed between the Insular and Intermontane superterranes. In this model, the Bridge River Terrane corresponds to the land remnant of an intraoceanic domain identified in the subducted Mezcalera slab wall (named from the Mezcalera oceanic plate hypothesized in Mexico by Dickinson and Lawton 2001). However, some authors question the accuracy of this "two-sutures" model which also implies a west-dipping subduction zone inboard of Wrangellia (see the recent discussion by Pavlis et al. 2019). Monger (2014) and Monger and Brown (2016) have proposed instead that the Bridge River and Hozameen terranes are the southern continuation of the Cache Creek Terrane, prior to dextral displacements of ~115 km on the latest Cretaceous-Paleocene Yalakom Fault (Umhoefer and Schiarizza 1996) and ~140 km of Eocene movement on the Fraser-Straight Fault (Monger and Brown 2016). The age discrepancy between the two terranes is interpreted by implying that the Bridge River Terrane faced open ocean to the west until it was trapped in the Early Cretaceous (≤130 Ma) behind the arc rocks in the southwestern Coast Mountains, whereas in northern British Columbia, the Cache Creek Terrane was thrust southwestward over Stikinia in the earliest Middle Jurassic (~174 Ma) (Mihalynuk et al. 2004), and in southern British Columbia was thrust eastward over Quesnellia probably in the Late Jurassic (≤ 160 Ma) (Travers 1978). The likelihood of differentiating the two terranes based on provincial radiolarian signatures has been discussed by Cordey (1998) and Orchard et al. (2001) and present some limitations. Some

elements of correlation with Tethyan assemblages may be found in older Paleozoic and Triassic assemblages of the Cache Creek Terrane (Cordey 1998, Orchard et al. 2001), but also in the Bridge River Terrane (Cordey and Schiarizza 1993, Cordey 1996, Cordey 1998). During the Jurassic, one can only demonstrate that the Bridge River Terrane has a longer record of chert sedimentation than the Cache Creek Terrane.

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### **Conclusions**

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So far as known, all the youngest radiolarian assemblages of the Cache Creek Terrane are Early Jurassic in age, including in the type locality in southern British Columbia. All refined ages of the youngest radiolarian cherts and siliceous argillites are within the Late Sinemurian-Early Toarcian interval (~195–180 Ma) with a possible Pliensbachian common age (~191–183 Ma). Well-dated Early Pliensbachian radiolarian cherts are exposed in the central part of the terrane and may correspond to the youngest pelagic strata. These results suggest that Cache Creek deep pelagic "true chert" sedimentation ceased fairly synchronously along the entire belt, probably during the Pliensbachian, followed by hemipelagic sedimentation in the Pliensbachian and/or the Toarcian, predating the final closure of the basin that occurred by tectonic thrusting of Cache Creek units over Stikinia and Quesnellia in the Middle and Late Jurassic. These results should lead to re-evaluating the tectonic mechanism of the Cache Creek entrapment model in its present form, either in speed or geometry. In addition, the study shows that pelagic sedimentation ceased ~22 to 31 m.y. earlier in the Cache Creek Terrane than in the Bridge River Terrane, confirming the stratigraphic discrepancy between the two terranes during the Jurassic. These refined chronological data now need to be more precisely taken into account by the current tectonic models which integrate the two terranes and the suture(s) they represent at the core of the 440 discussion. 441 442 Acknowledgements 443 444 This study results from a long collaboration with regional geologists and paleontologists who 445 investigated the Cache Creek Terrane: J. Monger, M. Orchard, P. Read, N. Mortimer, H. 446 Gabrielse, S. Gordey, B. Struik, M. Mihalynuk, J. English, C. Floriet, H. Sano, J. Jackson, S. 447 448 Johnston, P. Schiarizza, and A. Zagorevski. Former fieldwork was supported by the Geological 449 Survey of Canada (Vancouver) and the British Columbia Geological Survey Branch (Victoria). 450 Laboratory facilities were provided by the Laboratoire de Géologie de Lyon (Centre National de 451 la Recherche Scientifique CNRS - Unité Mixte de Recherche UMR5276). Reviewers S. Goričan 452 and J. Nelson as well as Editor A. Polat and Associate Editor B. Pratt provided insightful 453 comments which significantly improved the first version of the manuscript. 454 455 456 References 457 Baumgartner, P.O., Bartolini, A., Carter, E.S., Conti, M., Cortese, G., Danelian, T., De Wever, P., 458 Dumitrica, P., Dumitrica-Jud, R., Goričan, Š., Guex, J., Hull, D.M., Kito, N., Marcucci, M., 459 Matsuoka, A., Murchey, B., O'Dogherty, L., Savary, J., Vishnevskaya, V., Widz, D., and Yao, A. 460 1995. Middle Jurassic to Early Cretaceous radiolarian biochronology of Tethys based on Unitary 461 Associations. In Middle Jurassic to Lower Cretaceous Radiolaria of Tethys: Occurrences, 462 Systematics, Biochronology. *Edited by P.O. Baumgartner*, L. O'Dogherty, Š. Goričan, E. 463 Urquhart, A. Pillevuit and P. De Wever. Mémoires de Géologie (Lausanne), Vol. 23, pp. 1013– 464

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# Table 1. Jurassic radiolarian taxa and ages from the Cache Creek Terrane

Locality, Region	Subterrane or unit, map area & UTM coordinates	Lithology, Sample, GSC locality number when available	Radiolarian assemblage	Original age, Reference(s)	New or revised age	Generic assemblage or correlation to biozones (when applicable)
J1 southern British Columbia	western belt Pavilion NTS 092I13 UTM zone 10 581045E 5646753N	siliceous argillite NEO-R2 GSC C-300438 (previously R2, Cordey et al. 1987, Cordey 1998)	Archaeodictyomitra sp., Homoeoparonaella sp. cf. reciproqua, Hsuum sp., Pseudocrucella? sp., Thurstonia sp., Trillus sp., Xiphostylus? sp., Zartus sp.	Early Jurassic (Pliensbachian- Bajocian) (Cordey et al. 1987, Cordey 1998)	Early Jurassic (Early Pliensbachian - Early Toarcian)	Co-occurrence of Thurstonia, Trillus and Zartus (O'Dogherty et al. 2009)
J2 southern British Columbia	western belt Alkali Lake NTS 092016 UTM zone 10 548240E 5749410N	gray-green radiolarian chert 92-PBR-C92-78F GSC C-210033	Canutus sp., Lantus sp., Orbiculiformella sp., Parahsuum sp., Praeconocaryomma sp.	Early Jurassic (Sinemurian- Pliensbachian) (Cordey, unp. data; Read 1993)	Early Jurassic (Early Pliensbachian - Early Toarcian)	Co-occurrence of Canutus and Lantus (O'Dogherty et al. 2009)
J3 central British Columbia	Blackwater River Punchaw Lake NTS 093G6 UTM zone 10 490750E 5904150N	gray radiolarian chert 95FC-21-4 GSC C-303032	Atalantria sp., Beatricea? argescens, Bipedis japonicus, Broctus sp., Broctus? sp., Canoptum rugosum, Canoptum sp. cf. dixoni, Charlottea sp., Katroma sp. cf. elongata, Lantus sp. cf. praeobesus, Laxtorum sp., Orbiculiformella? trispina s.l. (Yeh), Orbiculiformella sp., Parahsuum simplum, Parahsuum vizcainoense, Praeconocaryomma sp. cf. sarahae, Thurstonia sp., Spumellaria gen. sp. indet.	Early Jurassic (Pliensbachian) (Cordey and Struik 1996, Orchard et al. 2001)	Early Jurassic (Early Pliensbachian)	Interval from the middle part of the Canutus tipperi – Katroma clara Zone to the base of the Gigi fustis – Lantus sixi Zone (Unitary associations UA 04-12) (Carter et al. 2010).
J4 central British Columbia	Tezzeron succession Pinchi Lake NTS 093K9 UTM zone 10, 418200E 6049406N	limy siltstone 97-SCB-3602 GSC C-209935	Emiluvia? sp. A, Hsuum lucidum, Paronaella grahamensis, Praeconocaryomma decora gr., Praeparvicingula tlellensis	Early Jurassic (Toarcian) (Struik et al. 2001, Orchard et al. 2001)	Early Jurassic (Middle - Late Toarcian)	Interval represented by the Elodium pessagnoi – Hexasaturnalis hexagonus Zone (Carter et al. 2010)
J5 northern British Columbia	French Range Little Dease Lake NTS 104J09 UTM zone 10, 419000E 6508000N	siliceous argillite MMI99-34-12	Parahsuum izeense, Orbiculiformella? trispina s.l., Praeconocaryomma sp. cf. sarahae, Praeconocaryomma sp., Thurstonia sp.	Early Jurassic (Pliensbachian- Toarcian) (Mihalynuk et al. 2004)	Early Jurassic (Pliensbachian - Early Toarcian)	Co-occurrence of Parahsuum izeense and Thurstonia (Carter et al. 2010, O'Dogherty et al. 2009).
J6 Yukon	Teslin Plateau  Mount Bryde NTS 105C03 UTM zone 10, 599028E 6676847N	siliceous argillite interbedded with sandstone 90FC-56-6 GSC C-177569	Bipedis sp., Broctus sp., Canoptum anulatum, Canoptum sp., Hagiastrum majusculum, Homoeoparonaella sp., Lantus sixi, Napora cerromesaensis, Pantanellium cumshewaense, Parahsuum izeense, Praeconocaryomma? sp. cf. yakounensis, Thurstonia sp., Wrangellium sp.	Early Jurassic (Pliensbachian- Early Toarcian) (Cordey et al. 1991)	Early Jurassic (late Early Pliensbachian - Late Pliensbachian)	Interval from the base of the Gigi fustis – Lantus sixi Zone to the middle part of the Eucyrtidiellum nagaiae – Praeparvicingula tlellensis Zone (Unitary associations UA 12-22; Carter et al. 2010)
J7 Yukon	Teslin Plateau  Mount Bryde NTS 105C03 UTM zone 10, 600094E 6673025N	black radiolarian chert 90FC-63-2 GSC C-177580	Beatricea? argescens, Canoptum sp. cf. anulatum, Canoptum sp., Praeconocaryomma sp. cf. immodica, Wrangellium sp.	Early Jurassic (Sinemurian- Toarcian) (Cordey et al. 1991)	Early Jurassic (Late Sinemurian - Late Pliensbachian)	Co-occurrence of Beatricea and Wrangellium (O'Dogherty et al. 2009)
J8 Yukon	Teslin Plateau  Mount Bryde NTS 105C03 UTM zone 10, 600094E 6673025N	dark gray radiolarian chert 90FC-63-9 GSC C-177587	Beatricea? argescens, Bipedis sp., Canoptum sp. cf. dixoni, Gorgansium sp., Pantanellium sp., Parahsuum sp., Praeconocaryomma sp., Wrangellium sp.	Early Jurassic (Sinemurian- Toarcian) (Cordey et al. 1991)	Early Jurassic (Late Sinemurian - Late Pliensbachian)	Co-occurrence of Beatricea, Bipedis, and Wrangellium (O'Dogherty et al. 2009)

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759 Figure captions 760 Figure 1 (colour online and printed). Paleozoic to early Mesozoic terranes of the Canadian Cordillera and location of the Cache Creek Terrane. Modified from Piercey and Colpron (2009). 761 762 Figure 2. Geographic location of Jurassic radiolarian assemblages from the Cache Creek Terrane 763 (J1 to J8, Table 1). Black frame: Fig. 3a (type locality of the terrane). 764 765 Figure 3. a: Distribution of Paleozoic and Mesozoic rock units in the type locality of the Cache 766 Creek Terrane and position of the study area for locality J1 (Fig. 2). Geology from Monger 767 768 (1985), b: Geological map of the Pavilion area and location of the Early Jurassic radiolarian sample NEO-R2. Geology from Monger and McMillan (1984), Mortimer (1987), and Cordey et 769 al. (1987). 770 771 Figure 4. Synthetic lithostratigraphic successions of the Cache Creek Terrane in its type locality 772 (southern British Columbia). The locality numbers (C-) refer to the Geological Survey of Canada 773 database and the radiolarian localities from Cordey (1998). Black lines (eastern belt): age of 774 blocks (Cache Creek mélange). Conodont data from Orchard (1981, 1984), Beyers and Orchard 775 (1989, 1991); fusulinid data from Monger and Ross (1971). Geological time scale from Cohen et 776 al. (2013). G: Guadalupian; L: Lopingian. LO: Lower. 777 778 779 Figure 5. Early Jurassic radiolarians from the Cache Creek Terrane. For each picture: taxon, sample, locality (Table 1), length of scale bar. 1. Thurstonia sp., NEO-R2, J1, 150 µm. White 780 arrows point to fifth and sixth spines. 2. Thurstonia? sp., NEO-R2, J1, 170 µm. 3-7: Thurstonia 781 782 sp.; 3-5: 90FC-56-6, J6, 160 μm; 6: MMI99-34-12, J5, 110 μm; 7: 90FC-56-6, J6, 180 μm. White

- arrows point to sixth spine. 8-9. Spumellaria gen. sp. indet., 95FC-21-4, J3, 165 and 140 µm. 10.
- Pantanellium cumshewaense Pessagno & Blome, 90FC-56-6, J6, 110 μm. 11-12. Trillus spp.,
- NEO-R2, J1, 120 μm. 13. Zartus sp., NEO-R2, J1, 115 μm. 14. Charlottea sp., 95FC-21-4, J3,
- 786 115 μm. 15. Palaeosaturnalis sp., 95FC-21-4, J3, 160 μm. 16. Hagiastrum majusculum Whalen
- & Carter, 90FC-56-6, J6, 170 μm. 17. Homoeoparonaella sp. cf. reciproqua Carter, NEO-R2, J1,
- 788 220 μm. 18. *Homoeoparonaella* sp., 90FC-56-6, J6, 140 μm. 19. *Orbiculiformella* sp., 95FC-21-
- 4, J3, 170 μm. 20. Orbiculiformella? trispina s.l. (Yeh), 95FC-21-4, J3, 150 μm. 21. Beatricea?
- argescens Cordey, 90FC-63-2, J7, 160 μm. 22. Praeconocaryomma? sp. cf. yakounensis Carter,
- 791 90FC-56-6, J6, 150 μm. 23-24. *Praeconocaryomma* sp. cf. sarahae Carter; 23: 95FC-21-4, J3,
- 792 130 μm; 24: MMI99-34-12, J5, 160 μm.

- Figure 6. Early Jurassic radiolarians from the Cache Creek Terrane. For each picture: taxon,
- sample, locality (Table 1), length of scale bar. 1. *Lantus sixi* Yeh, 90FC-56-6, J6, 110 μm. 2.
- 796 Lantus sp. cf. praeobesus Carter, 95FC-21-4, J3, 110 μm. 3. Broctus sp., 95FC-21-4, J3, 100 μm.
- 797 4. *Broctus* sp., 90FC-56-6, J6, 130 μm. 5. *Xiphostylus*? sp., NEO-R2, J1, 140 μm. 6. *Katroma* sp.
- cf. elongata Carter, 95FC-21-4, J3, 150 μm. 7. Parahsuum vizcainoense Whalen & Carter, 95FC-
- 799 21-4, J3, 90  $\mu$ m. 8. *Parahsuum izeense* (Pessagno & Whalen), 90FC-56-6, J6, 140  $\mu$ m. 9.
- 800 Canoptum sp., 95FC-21-4, J3, 100 μm. 10. Canoptum rugosum Pessagno & Poisson, 95FC-21-4,
- 301 J3, 90  $\mu$ m. 11-12. *Canoptum* spp., 90FC-56-6, J6, 120  $\mu$ m. 13. *Broctus*? sp., 95FC-21-4, J3, 110
- 802 μm. 14. Canoptum sp. cf. anulatum Pessagno & Poisson, 90FC-63-2, J7, 120 μm. 15. Canoptum
- sp. cf. dixoni Pessagno & Whalen, 95FC-21-4, J3, 120 μm. 16. Atalantria sp., 95FC-21-4, J3, 120
- 804 μm. 17. *Laxtorum* sp., 95FC-21-4, J3, 130 μm. 18. *Wrangellium* sp., 90FC-56-6, J6, 120 μm. 19.
- 805 *Hsuum* sp., NEO-R2, J1, 110 μm. 20. *Archaeodictyomitra* sp., NEO-R2, J1, 130 μm. 21. *Bipedis*
- sp., 90FC-56-6, J6, 110 μm. 22. Bipedis japonicus Hori, 95FC-21-4, J3, 130 μm. 23. Napora
- 807 *cerromesaensis* Pessagno, Whalen & Yeh, 90FC-56-6, J6, 105 μm.

808 Figure 7. Synthesis of the youngest Jurassic radiolarian ages of the Cache Creek Terrane. 809 Previous age determinations: J1: Cordey et al. (1987), Cordey (1998); J2: Read (1993), Cordey 810 811 (unpub. data); J3: Cordey and Struik (1996), Orchard et al. (2001); J4: Struik et al. (2001), 812 Orchard et al. (2001); J5: Mihalynuk et al. (2004); J6-J8: Cordey et al. (1991). 813 Figure 8. Radiolarian age ranges and correlations of youngest Cache Creek strata in British 814 815 Columbia and Yukon. The hatched band materializes the youngest common age range of 816 siliceous rocks (radiolarian cherts and siliceous argillites); limy siltstones at J4 are not included in 817 the correlation. The gray band shows the hypothetical age range of the youngest siliceous argillites overlying the youngest well-dated radiolarian cherts (J3, Early Pliensbachian). The 818 819 oblique line represents the maximum age difference of the youngest well-dated hemipelagic 820 rocks (siliceous argillites) between southern British Columbia (J1) and Yukon (J6); this interval 821 (late Early Pliensbachian to Early Toarcian) is ~8 m.y. (~188–180 Ma, Cohen et al. 2013). 822

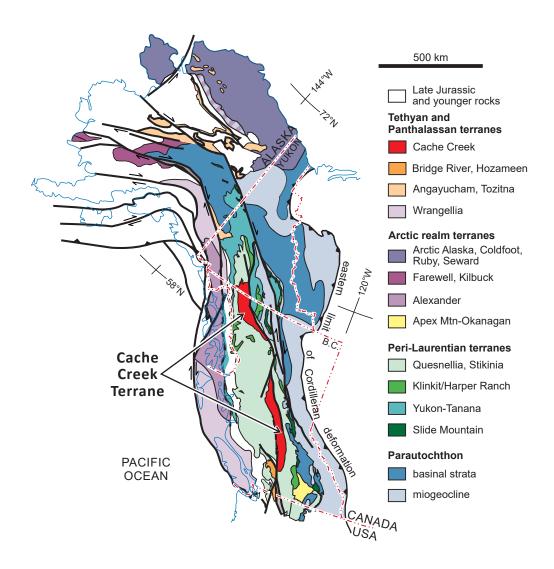


Figure 1

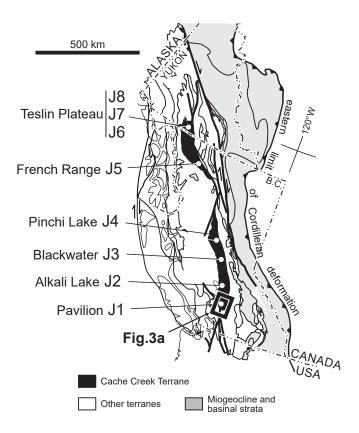


Figure 2

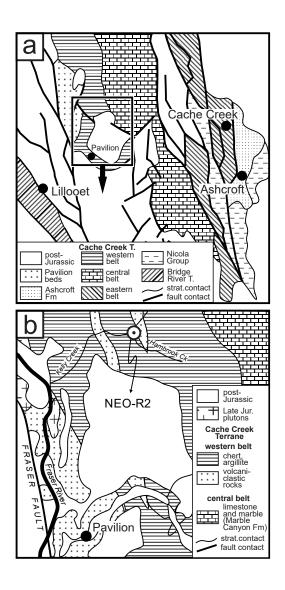


Figure 3

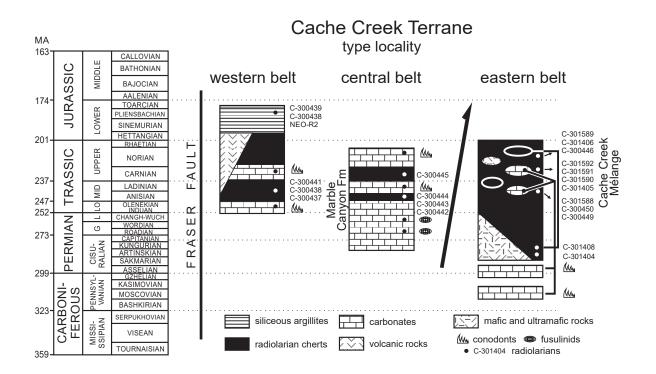


Figure 4

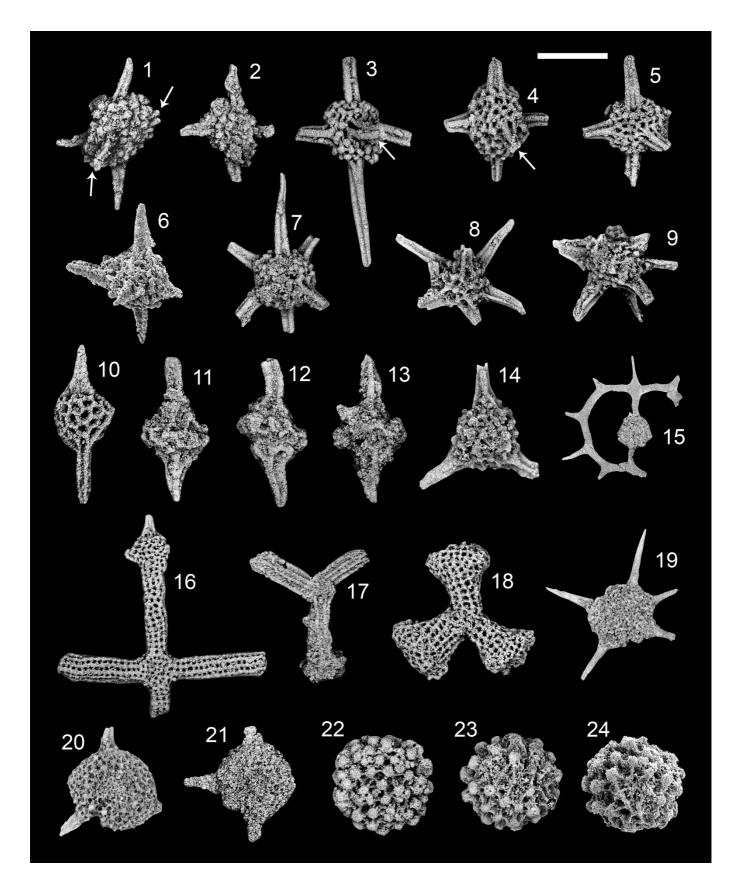


Figure 5

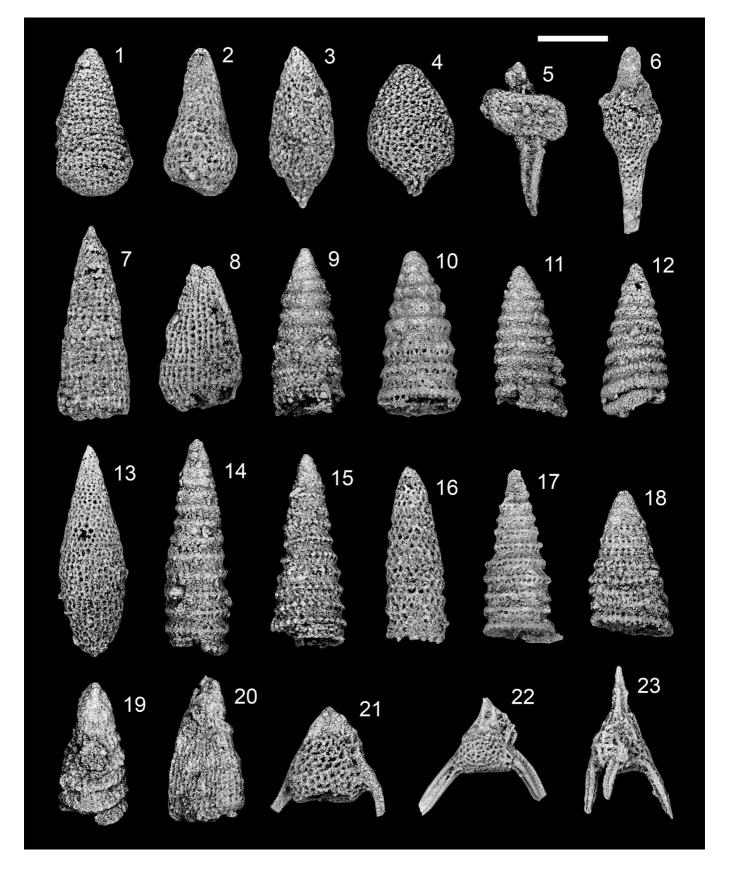


Figure 6

#### southern BC central BC north. BC Yukon Pinchi French Alkali Black-Pavilion Teslin Plateau Lake water Lake Range Ма J2 J3 J4 J5 J7 J1 J6 J8 163.5 ±1.0 Callovian Middle 166.1 ±1.2 Bathonian Bajocian Aalenian Toarcian 182.7 ±0.7 Lower Pliensbachian.

190.8 ±1.0 ····

Sinemurian L

Hettangian

Cache Creek Terrane

new or refined age

Figure 7

previous age determination

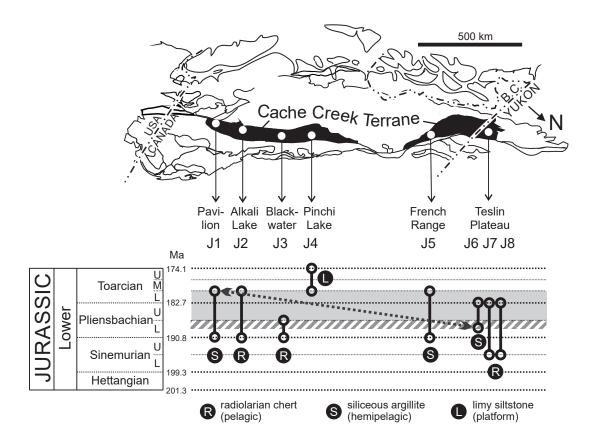


Figure 8