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OPEN

Historically-based run-time bias corrections substantially improve model projections of 100 years of future climate change

Gerhard Krinner ^{1✉}, Viatcheslav Kharin ², Romain Roehrig ³, John Scinocca ² & Francis Codron ⁴

Climate models and/or their output are usually bias-corrected for climate impact studies. The underlying assumption of these corrections is that climate biases are essentially stationary between historical and future climate states. Under very strong climate change, the validity of this assumption is uncertain, so the practical benefit of bias corrections remains an open question. Here, this issue is addressed in the context of bias correcting the climate models themselves. Employing the ARPEGE, LMDZ and CanAM4 atmospheric models, we undertook experiments in which one centre's atmospheric model takes another centre's coupled model as observations during the historical period, to define the bias correction, and as the reference under future projections of strong climate change, to evaluate its impact. This allows testing of the stationarity assumption directly from the historical through future periods for three different models. These experiments provide evidence for the validity of the new bias-corrected model approach. In particular, temperature, wind and pressure biases are reduced by 40–60% and, with few exceptions, more than 50% of the improvement obtained over the historical period is on average preserved after 100 years of strong climate change. Below 3 °C global average surface temperature increase, these corrections globally retain 80% of their benefit.

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For most practical purposes of climate change impact assessment, regional-scale climate change information is required¹. This information begins with large-scale climate change information produced from coupled atmosphere-ocean global climate models (CGCMs). Critical features at these scales include the placement and intensity of major atmospheric centers of action, such as jet streams, and patterns of variability particularly on interannual and monthly time scales. Finer regional scale and boundary-layer climate are then derived from CGCM output by statistical and dynamical downscaling², thereby adding value to the climate change information produced by CGCMs. It is well known that CGCMs can contain substantial biases in large-scale features, which are often of the order of the projected centennial climate change signal itself³. This fact is sometimes referred to as the garbage in-garbage out problem of regional climate modeling⁴, which denotes the fact that model output is necessarily worthless incorrect if the input is wrong. This problem can be partially compensated for by appropriate selection of the driving coupled climate model^{4–6}, and post hoc bias adjustment of the large-scale or downscaled regional climate change information^{6–8}. However, large-scale climate information biases remain and in particular circulation biases continue to undermine the potential value added by the downscaling exercise^{4,9}.

As an additional tool to support the reduction of CGCM biases of large-scale climate change information, an empirical bias-correction (EBC) approach for simulations with atmosphere-only general circulation models (AGCMs) has been proposed and applied¹⁰. This approach uses bias-corrected sea-surface temperatures and sea ice coupled with in situ, or run-time, corrections of large-scale climate model biases, following previous work^{11,12}, where this method had been applied in the context of seasonal predictions. The EBC involves the derivation of an annual cycle of cyclostationary forcings over the historical period for the atmosphere-only (AGCM) counterpart of a CGCM (see “Methods” section). The EBC is then applied to the AGCM for atmosphere-only future climate projections employing bias-corrected sea-surface temperatures and sea ice from the CGCM with the goal of reducing general circulation biases in future projections. As with all bias correction methods, the utility of the EBC approach rests on the assumption that climate biases are essentially stationary between historical and future climate states. It is not at all obvious that the stationarity assumption should hold given the large external forcing perturbations associated with many scenarios of future climate change.

In this study, the utility of the EBC approach is investigated by employing model data as a proxy for observations to allow an explicit testing of the stationarity assumption. We employ coupled atmosphere-ocean climate models (CGCMs) and their associated AGCMs from three independent climate modeling centers, and pair each center’s AGCM with another center’s CGCM (see “Methods” section). For each pair, an EBC is derived for the AGCM using the historical simulations of the CGCM as the observational reference. The EBC approach is then used in the AGCM for future projections. The historical and future climate (in particular the atmospheric circulation) simulated by the corrected and the uncorrected AGCMs is then compared to that of the reference CGCM. In such a perfect model (or pseudo-reality) experiment^{13–16}, one can explicitly test the stationarity assumption and so the utility of the EBC procedure across several models and several reference versions of pseudo observations. It is arguable that employing CGCM model output as a proxy for observations provides an even more stringent test of the EBC methodology as all CGCMs were developed independently, and each have unique historical climatologies and future projections. The fundamental question addressed here is whether, and to what extent, the improvement obtained by construction for the

present-day reference climate is preserved in the future (here, over 100 years under a very strong emission scenario). While the EBC method was initially developed and tested for a seasonal prediction context^{11,12}, more recent work¹⁰ heuristically applied it for Antarctic climate projections. This paper provides the first proof of the validity of this method in the context of projections of strong climate change.

In the following, the EBC approach is assessed for its ability to improve climatological (20-year) mean values and statistics of interannual and synoptic time scale variability. The results show that using this bias correction method consistently improves the simulated climate on this large range of time scales, and more importantly, that a large part of the improvement obtained for the present reference period is preserved under strong climate change.

Results

Climatological means. Figure 1 displays time series of the global mean square error of 20-year running-mean averages of monthly climatological means of fundamental atmospheric variables at representative levels in the free troposphere (air temperature T , zonal wind speed u , meridional wind speed v , geopotential height z_g) and surface air pressure p_s . For each corrected AGCM, the error is calculated with respect to the monthly climatology of the target CGCM at each level, vertically averaged if applicable, and normalized by dividing by the corresponding error of the uncorrected version of the AGCM. The figure shows clearly that except for surface pressure in ARPEGE after 2080, the corrected AGCMs continue to outperform the uncorrected versions systematically, and that for almost all variables and models, there is no substantial degradation of the error scores after the 1981–2000 calibration period. For most variables and models, the mean square error of the corrected AGCMs are fairly stable in time, with typical values of about 40–60% of the MSE of their uncorrected versions (see also Supplementary Fig. 1).

The degradation of the surface pressure and geopotential height scores for ARPEGE in the latter half of the twenty-first century is due to a positive bias of the target IPSL-coupled model over the Southern Ocean during the reference period and a negative sea-level pressure trend in that model over the twenty-first century (see Supplementary Fig. 2). While the corrected version of the ARPEGE AGCM better represents the present sea-level pressure over the Southern Ocean of the IPSL-coupled model, both the corrected and uncorrected versions of the AGCM do not reproduce the strong twenty-first century trend simulated by the target model that compensates for its present-day bias, leading to a strong error of the corrected AGCM by the end of the twenty-first century. This is consistent with the fact that the IPSL and CNRM coupled models represent opposite extremes of the Coupled Model Intercomparison Project phase 5 (CMIP5) spectrum regarding the trends of Southern Hemisphere jet and Southern Annular Mode¹⁷. The ARPEGE/IPSL-CM combination is thus a test of the most unfavorable possible situation. For the other AGCMs and other variables of the ARPEGE model, 80% of the MSE reduction obtained for the 1981–2000 reference period is preserved at least until about 2070, when the global mean surface air temperature change with respect to the reference period attains about 3 °C. A striking example is the 500 hPa geopotential height in the CanAM4 model, displayed in Fig. 2, which emulates very well both the present and future global geopotential height distribution of the CNRM-CM coupled model. Supplementary Figs. 3–6, which display similar maps for all models and additional variables, confirm that a general and consistent improvement is obtained for the present and, importantly, largely conserved in the projections. We note that while the amplitude of

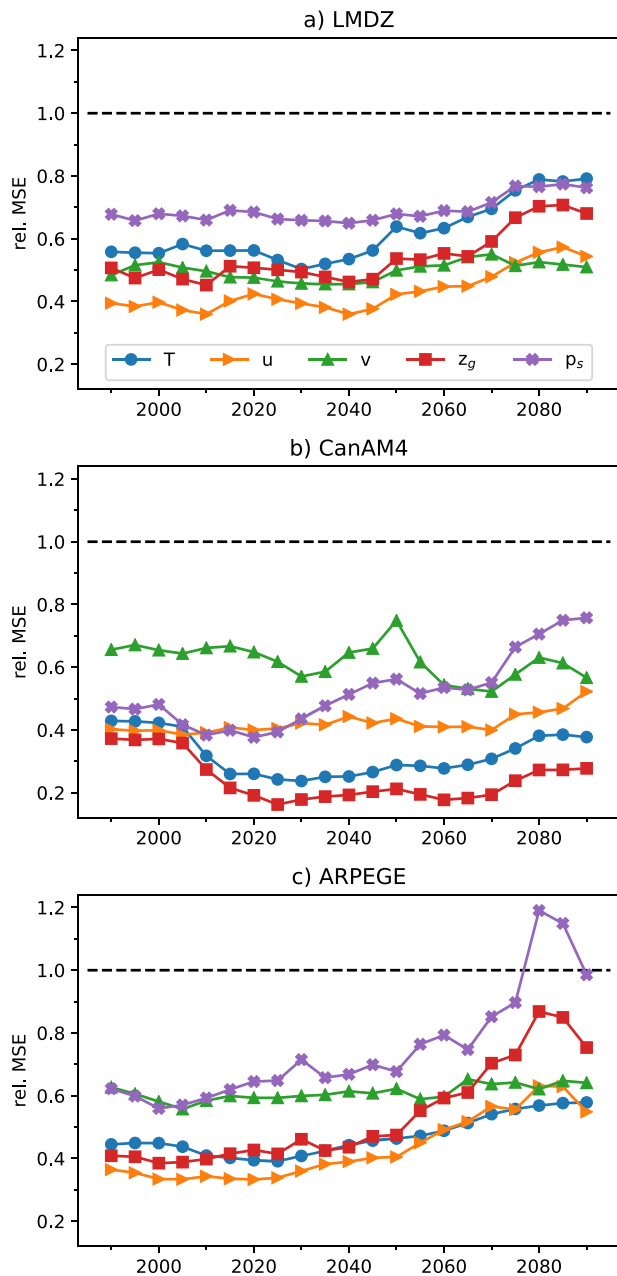


Fig. 1 Time evolution of global errors. Global average mean square error (MSE) of the corrected AGCM runs relative to the uncorrected reference runs for 20-year running means of fundamental atmospheric circulation variables (air temperature T , zonal wind speed u , meridional wind speed v , geopotential height z_g , and surface air pressure p_s (over areas with surface height below 1000 m)). MSE of T , u , v , and z_g are averaged over four representative pressure levels (850, 700, 500, and 300 hPa (200 hPa for LMDZ)). **a** LMDZ; **b** CanAM4; **c** ARPEGE.

the biases is quite systematically reduced in the corrected runs, the spatial patterns of the biases are often similar between the uncorrected and the corrected versions, and between the reference and projection period.

While free atmosphere temperature and wind fields were nudged and corrected in our simulations, atmospheric mass (that is, surface pressure and geopotential height fields) was not. Good and stable performance for these variables (at least up to about 3 °C global mean surface air temperature change), linked to the extratropical wind fields by the near-geostrophic relation, is

therefore an indicator of consistent and robust improvement of the simulated atmospheric circulation. These improvements, both for the reference and the projection periods, are also visible in the location and intensity of atmospheric centers of action such as the Southern Westerlies (Supplementary Fig. 7) and the Aleutian low (Supplementary Fig. 8).

Interannual variability. For the climatological means of wind and temperature, and, indirectly, geopotential heights and sea-level pressure, the improvement during the 1981–2000 reference period is obtained by construction, while the preservation of the benefit of the bias correction beyond that period is an indicator of the validity of the method. Conversely, emergent climate system properties such as patterns of interannual circulation variability are not necessarily improved by construction even during the reference period. However, as can be seen in Fig. 3, the corrected AGCMs do better represent the spatial patterns of the interannual extratropical circulation variability of the target coupled models than the uncorrected AGCMs. Across the three models, the squared spatial correlation coefficient r^2 between the first 500 hPa geopotential empirical orthogonal functions (EOFs) of the AGCMs and the target coupled models is higher for the corrected than the uncorrected models during 83% of the simulation period, and no systematic degradation can be seen in the latter part of the period. This includes a number of situations, where improvement is particularly challenging because the spatial patterns of the first EOF in the AGCM and in its reference coupled model are very similar (that is, the spatial r^2 is close to 1 already in the uncorrected version). For the 500-hPa zonal wind speed, the corresponding successful proportion is 78%. This indicates that the corrected AGCMs quite systematically simulate interannual circulation variability patterns that are in better agreement with those of the target pseudo-reality coupled model, and this improvement is preserved under strong climate change. This ability of the EBC to improve second-order climate statistics is consistent with a previous study¹² that demonstrated improved interannual variability for ten fields in the context of EBC seasonal predictions.

Synoptic variability. A similarly fundamental emerging circulation characteristic is the sea-level pressure variability on the synoptic time scale between 2 and 6 days¹⁸, dominated by extratropical storm tracks. As shown in Fig. 4, the corrected AGCMs quite consistently (during 87% of the simulation period across the three models) exhibit a reduced global average mean square error of the temporal standard deviation of band-pass (2–6 days) filtered daily sea-level pressure, and again, a beneficial effect of correcting the mean circulation characteristics remains visible until at least about 2070. In two out of the three corrected AGCMs, only a negligible long-term increase of the relative MSE is observed over the twenty-first century. This improvement in the global-mean synoptic variability is consistent with an improved intensity and location of the variability maxima, as can be seen for example for the Aleutian Low in winter in Supplementary Fig. 8.

Discussion

Taken together, our results show that the run-time empirical bias correction (EBC) method tested here improves simulated atmospheric circulation characteristics on a broad range of temporal scales. In addition to the climatological mean values, which are improved by construction over the historical period, emergent circulation properties such as patterns and intensities of interannual and synoptic-scale circulation variability are also improved. Most importantly, our pseudoreality test clearly shows

zg500 biases, CanAM vs. CNRM-CM

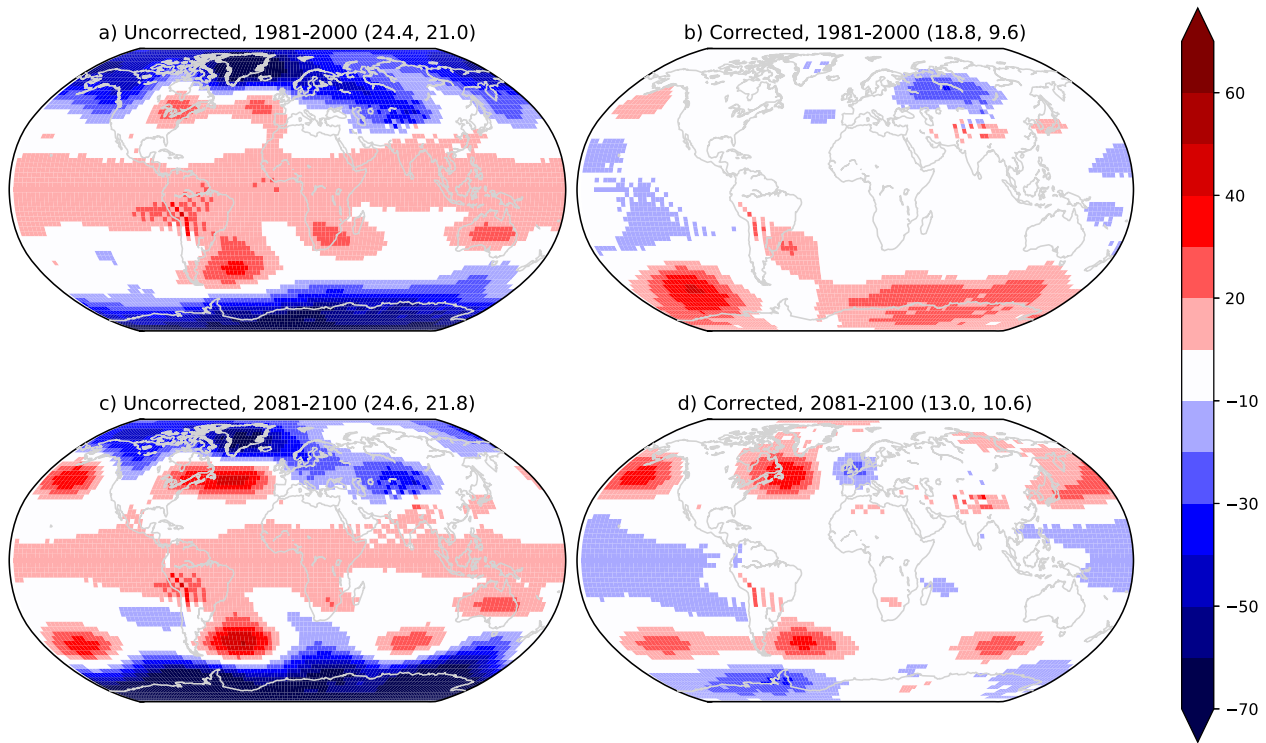


Fig. 2 Mid-tropospheric geopotential height errors. Annual mean 500 hPa geopotential height difference (in m) between CanAM and CNRM-CM. **a** Uncorrected, 1981–2000; **b** corrected, 1981–2000; **c** uncorrected, 2081–2100; **d** corrected, 2081–2100. The global mean error is subtracted to highlight the spatial structures over global mean errors. RMSE of the CanAM4 simulation with respect to the CNRM reference (in m), before and after subtraction of the global mean error, is given in parentheses in the panel titles.

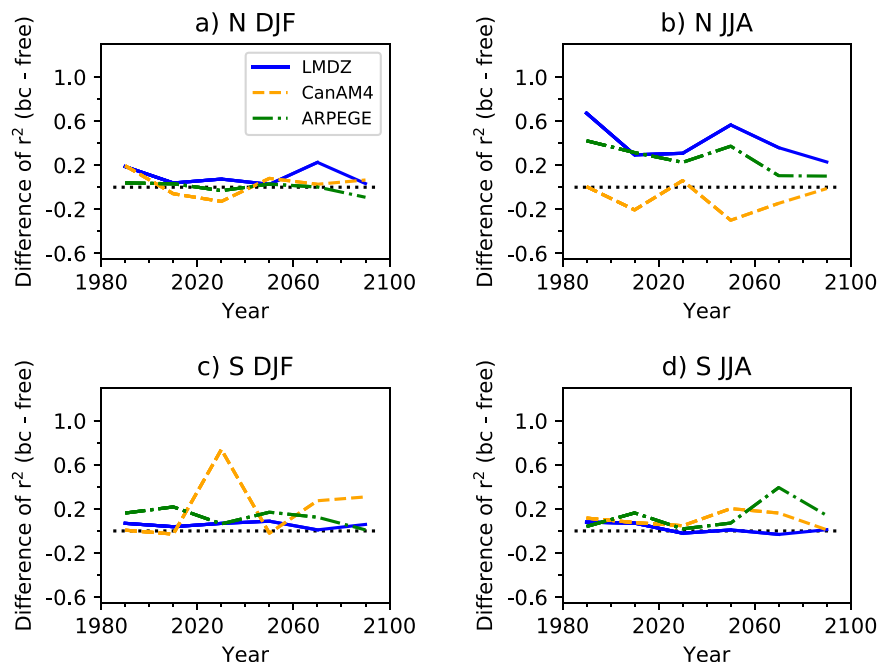


Fig. 3 Interannual variability of extratropical circulation. Squared spatial correlation coefficient (r^2) of the first EOF of monthly mean extratropical (polewards of 30°N and S) 500 hPa geopotential (20-year periods), bias-corrected minus uncorrected simulation. **a** Northern hemisphere winter (DJF); **b** Northern hemisphere summer (JJA); **c** Southern hemisphere winter (DJF); **d** Southern hemisphere summer (JJA).

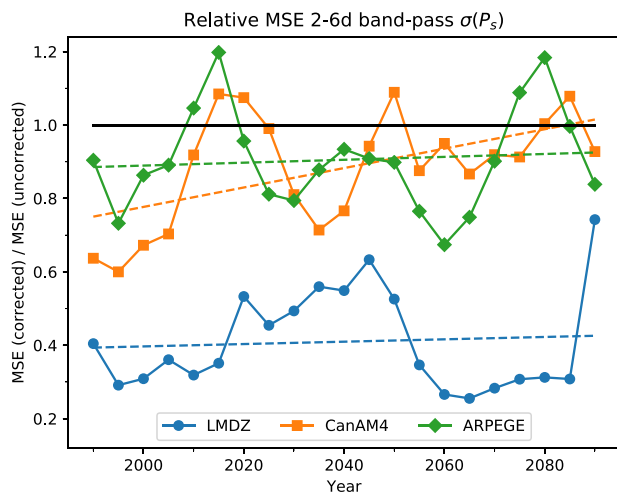


Fig. 4 Synoptic variability. Global average mean square error (MSE) of the corrected AGCM runs relative to the uncorrected reference runs for 20-year running averages of the temporal standard deviation of band-pass (2–6 days) filtered daily sea-level pressure (full lines). Values are calculated every 5 years. Dashed lines indicate the linear regression over the full period. The black constant line at $y = 1$, indicating the limit above which the benefit of the bias correction vanishes, is provided for visual guidance.

that the improvements on climatological, interannual, and synoptic time scales are to a very large part preserved under strong climate change of the order of an about 3 °C global surface air temperature change. In some cases, the relative biases of the corrected models even decrease over time, suggesting that the bias correction, as implemented here, does not over-constrain the models. This method for run-time bias corrections of large-scale atmospheric circulation models thus remains valid under strongly nonstationary conditions. Validity of this bias-correction method is further supported recent work¹⁹, which based on the analysis of CMIP5 results, has shown striking stationarity of large-scale climate model bias patterns under strong climate change.

The EBC approach is by construction expected to have little impact on the magnitude of the climate change response of the uncorrected CGCM as its sea-surface temperatures (SSTs) and sea ice are utilized as boundary conditions for the AGCM future projections (see “Methods” section). Benefit comes rather from circulation improvements about the CGCM’s changing climate signal, which should then improve its subsequent downscaling. The ability of the EBC to alter the control circulation upon which climate-change external forcings are applied, while leaving all other properties of the model unchanged, means that it additionally provides an interesting tool to disentangle the influences of model formulation vs. basic state on the circulation response itself. As such the EBC methodology will also provide a direct means to probe such questions and will compliment existing approaches (e.g., emergent constraints).

The bias correction does not seem to fix the location of features of the atmospheric circulation in space, as can be seen in Supplementary Fig. 7, which shows that the southward shift of the Southern Westerlies over the twenty-first century is similar in the free and in the bias-corrected versions for all AGCMs, confirming previous results¹⁰. This means that the simulated regional-scale climate change is not damped by the bias correction.

Our results immediately open broad perspectives for improved climate change projections on a large range of spatial scales beyond the $> \sim 100$ km scale already covered by typical AGCMs. Applying the EBC method, AGCMs can be used to re-analyze climate change information from large-scale coupled climate

model projections. As demonstrated by this study, the atmospheric run-time bias correction technique can be combined with the more usual bias correction of the prescribed oceanic boundary conditions for AGCM climate change experiments, which consists of imposing the ocean surface condition change (SST and sea-ice change) from a coupled climate model on observed ocean surface conditions^{20–22}. This results in present-day AGCM control runs with substantially reduced biases, and in consistently corrected climate change runs¹⁰. Such sets of consistent control runs and projections come at modest numerical cost, because only AGCM simulations without long ocean spinups are required. Output of EBC simulations with bias-corrected large-scale tropospheric circulation characteristics can then be used to drive higher-resolution limited area atmosphere models, or to directly drive regional or global ocean models, and land surface or ice sheet models leading to an overall reduction in the uncertainty of their climate products.

Although the main objective of the EBC is to provide a method for improving the simulated global large-scale circulation, it is worth noting that the corrected simulations also show improved statistics for precipitation, which is a physical quantity that one might suspect could suffer from possible inconsistencies and perturbations introduced by the ad hoc correction terms. For two out of the three AGCMs, the global MSE of precipitation in the corrected AGCM versions is consistently below 85% of the MSE of the uncorrected reference version for the entire 1980–2100 timespan, and for the third AGCM, while there is no strong improvement obtained for the present-day reference period, the precipitation MSE of the corrected version sharply drops after 2020, attaining values below 0.85 after 2050 (see Supplementary Fig. 9). The EBC thus improves even emergent physical properties such as precipitation, reducing errors that are connected indirectly to the reduction of atmospheric circulation biases (type 1 errors²³). However, errors in simulated precipitation rates due to insufficiencies of the physical parameterizations (type 3 errors²³) will not be corrected by the EBC method used here.

There are several areas where the results of the EBC might be improved over those presented here. By retuning free physical parameters of the AGCMs during the nudging stage, the magnitudes of their nudging tendencies and bias correction terms might be reduced. A possible extension of this bias correction approach is its application in a coupled context, with nudging applied in the atmosphere and/or ocean. Furthermore, conditional bias adjustments, frequently suggested and applied in posthoc correction methods^{9,24,25} could also be implemented in our method, possibly based on run-time classification of synoptic situations. This could reduce the remaining mean biases. Finally, mean biases will likely depend on the nudging timescale used in Eq. 1 to derive the EBC and more optimal choices might lead to improvements over the results presented here.

In summary, our work provides strong evidence for the validity of the EBC approach as a consistent, robust, versatile, simple to implement, and numerically affordable methodology to address the garbage in–garbage out problem of regional climate change projections⁴. We conclude that empirical bias corrections have the potential to substantially reduce uncertainty in the output of coupled climate change experiments, particularly in coordinated exercises like CMIP²⁶, and to improve the quality of atmospheric driving data for downstream downscaling activities such as the Coordinated Regional Downscaling Experiment (CORDEX)¹ and climate change impact analyses.

Methods

The run-time bias-correction method for atmospheric models is based on a two-step approach¹¹. The following description of the method is similar to descriptions given in previous work by the authors^{10,12}.

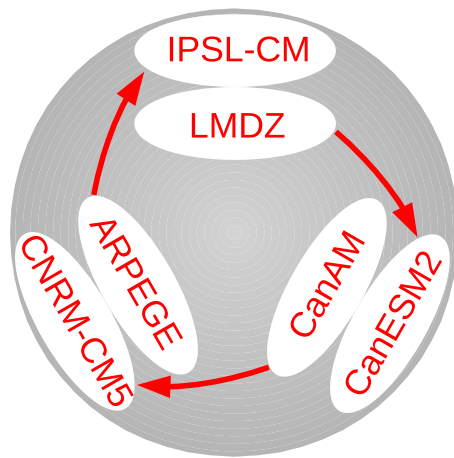


Fig. 5 Setup of the perfect model experiment. Arrows indicate which AGCM emulates which CMIP5 coupled model (CM); for example, LMDZ emulates CanESM2. The AGCMs are corrected towards their respective target coupled model for the 1981–2000 period, using the SST and sea-ice concentration from the target CM. They are then run into the future under the RCP8.5 scenario, using SST and sea-ice anomalies from their own coupled model (for example, IPSL-CM in the case of LMDZ). The uncorrected control runs use the same oceanic boundary conditions.

The first step consists of nudging²⁷ the atmospheric model to a time-varying reference state. At each model time step, the local value of a selected prognostic variable X is adjusted by applying a Newtonian relaxation towards a reference state X_R :

$$\frac{\partial X}{\partial t} = F(X) - \frac{1}{\tau}(X - X_R). \quad (1)$$

Here, the first part of the equation ($\partial X/\partial t = F(X)$) represent the prognostic evolution of the unconstrained AGCM and τ is the nudging time constant, chosen to be 3 days in this study. While the time-varying reference state X_R is usually taken from 6-hourly output of atmospheric (re-)analyses, here it is prescribed to be 6-hourly three-dimensional time-varying output of a coupled model. The nudged variables are air temperature and zonal and meridional wind above the atmospheric boundary layer. The detailed implementation of the smooth transition from no nudging at the surface to full nudging above the boundary layer varies among the participating models and is not critical, since the focus of this work is on free-atmosphere circulation characteristics, but typically the models are fully nudged above about 1500 m for grid points with surface altitude close to sea level. Specifically, the LMDZ5 AGCM²⁸ was nudged towards the first ensemble member of the CanESM2²⁹ historical run of the CMIP5 coordinated experiment³⁰ over the 1981–2000 reference period; the CanAM4 AGCM³¹ was nudged towards the CNRM-CM5.1³² CMIP5 historical run (first ensemble member); and the ARPEGE-Climate v5.2 AGCM³² was nudged towards the first ensemble member of the IPSL-CM5³³ CMIP5 historical run (see Fig. 5). These AGCM nudging reference runs are performed over the 20 year period 1981–2000 and employ the SST and sea-ice concentration of the respective target/reference coupled models as lower boundary conditions. This nudged simulation, by construction, has almost vanishing circulation biases.

For the second step, the nudging tendencies in Eq. 1 from the reference runs are time averaged to produce a climatological seasonal cycle of the applied correction term resulting in:

$$G = -\frac{1}{\tau}(\overline{X_N} - \overline{X_R})^{AC}. \quad (2)$$

The operator $\overline{(\cdot)}^{AC}$ designates the annual cycle of \overline{Y} ¹². The climatological, but seasonally and spatially varying correction terms G correspond to the mean nudging tendencies required to maintain atmospheric conditions close to those of the reference coupled model. The bias correction then consists of adding these cyclo-stationary temporally and spatially varying terms to the prognostic equations of the AGCM:

$$\frac{\partial X}{\partial t} = F(X) + G. \quad (3)$$

This yields an empirically bias-corrected solution for X .

The empirically bias-corrected run for the reference period used the same oceanic boundary conditions as the nudged run. The RCP8.5 run until 2100 then used the same atmospheric correction terms, and, during the CMIP5 projection period (2006–2100), month-by-month sea-surface condition (SST and sea ice) anomalies from the AGCM's own coupled model CMIP5 projection run (e.g., IPSL-CM5 for LMDZ5) superimposed on the 1981–2000 average sea-surface

conditions from the target coupled model (e.g., CanESM2 for LMDZ5) using an anomaly method²⁰. Standard CMIP5 atmospheric boundary conditions (greenhouse gas concentrations etc.) are used in the AGCM runs. The corrected AGCM reference and projections runs are then evaluated against the climate simulated by the target coupled model.

It is important to recognize that the EBC methodology imposes the same perturbative sea-surface forcing change between the future and historical periods in the AGCM future projections as that of its own coupled model CMIP5 projection run (i.e., sea-surface bias corrections cancel out in the difference). Consequently, the global surface atmospheric temperature change in these EBC experiments is strongly constrained to be that of its own coupled model. The climate sensitivities of the three target coupled CMIP5 models are broadly comparable³⁴: 4.1 °C for IPSL-CM5; 3.3 °C for CNRM-CM5.1; 3.7 °C for CanESM2 and so, we have not attempted to compensate for such differences in our analyses.

Data availability

The CMIP5 output used to nudge the LMDZ, ARPEGE, and CanAM models (global 6-hourly atmospheric temperature and winds, and global monthly SST and sea-ice fields for 1981–2000 from the first run of the historical ensembles of the IPSL, CanESM, and CNRM-CM CMIP5 models) is available from the ESGF (see <https://www.earthsystemcog.org/projects/cog/>). AGCM output that supports the findings of this study are available from the corresponding author upon reasonable request. The scripts and prepared data used to produce Figs. 1–4 of this work are available on <https://doi.org/10.5281/zenodo.4018860>.

Code availability

CanAM is freely available on <https://gitlab.com/ccma/canam>. LMDZ is freely available on <https://lmdz.lmd.jussieu.fr/>. Documentation of the Météo France ARPEGE model is available on https://www.umr-cnrm.fr/IMG/pdf/arp62ca_july2017.pdf.

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Author contributions

G.K., V.K., and J.S. designed the study. G.K., V.K., R.R., J.S., and F.C. contributed to the analysis and discussion of the model output and to the drafting and revision of the text.

Competing interests

The authors declare no competing interests.

Additional information

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