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# 1 ESR/U-series and ESR dating of several Middle Pleistocene Italian sites: comparison with $^{40}\text{Ar}/^{39}\text{Ar}$ chronology

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## 17 Abstract

18 The stratigraphic sequences of numerous Palaeolithic sites of Central and Southern Italy, very rich in both ar-  
19 chaeological and palaeontological remains, have also recorded Pleistocene volcanic events through volcanic ash de-  
20 posits (tephra). They allow the establishment of an accurate chronological framework by comparing results obtained  
21 by  $^{40}\text{Ar}/^{39}\text{Ar}$  dating method on single volcanic K-feldspar crystals, with those derived from ESR and ESR/U-series anal-  
22 yses on fluvial bleached quartz and tooth enamel respectively.

23 Since 2009, these three methods were hence applied on samples of several Middle Pleistocene sites of Central  
24 and Southern Italy including both volcanic and archaeological levels (from the west to the east): La Polledrara di Ce-  
25 canibbio, Isoletta, (Latium), Guado San Nicola, Isernia La Pineta (Molise), Valle Giumentina (Abruzzo) and Venosa  
26 Notarchirico (Basilicata).

27 The obtained results, covering a time range from 350 to 660 ka, prove that such a multi-method approach when  
28 possible is essential to constrain the chronology of each site and allow the recognition of the specific limitations due  
29 to the lack of quartz for ESR or to complex geochemical histories in teeth rendering difficult the ESR /U-series meth-

30 od. Despite these limitations, the ESR framework is globally in agreement with the  $^{40}\text{Ar}/^{39}\text{Ar}$  chronology, while  
31 ESR/U-series dates can be underestimated for the oldest sites. In such cases, an isochron approach attests however  
32 of the quite good reliability of the palaeodosimetric reconstruction and the observed age underestimation could be  
33 related to other factors affecting the ESR age determination.

34 Key-words – ESR/U-series dating, ESR dating,  $^{40}\text{Ar}/^{39}\text{Ar}$  dating, Middle Pleistocene, Palaeolithic, Italy

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## 36 Introduction

37 The dating of Middle Pleistocene archaeological sites can usually be performed using a relatively restricted num-  
38 ber of geochronological methods. The opportunity to use these technics relies on both the considered time range  
39 and available datable materials found on these sites. These parameters greatly depend on the geological context  
40 and, consequently, the establishment of accurate chronological framework of the Middle Pleistocene human evolu-  
41 tionary scheme is very heterogeneous in terms of precision and accuracy in the various areas of the world. In West-  
42 ern Europe, this framework was mainly built by uranium-series (U-series) dates on speleothems and by palaeodosi-  
43 metric dating methods such as thermoluminescence (TL) on heated flints, optically stimulated luminescence (OSL) on  
44 sediments, electron spin resonance (ESR) on fluvial sediments and coupled ESR and U-series on teeth. For open-area  
45 localities, at least one of these palaeodosimetric methods can generally be applied but, if several methods are used  
46 at the same time and show geochronological discrepancy, the lack of completely independent age control drastically  
47 complicates the chronological and archaeological interpretation.

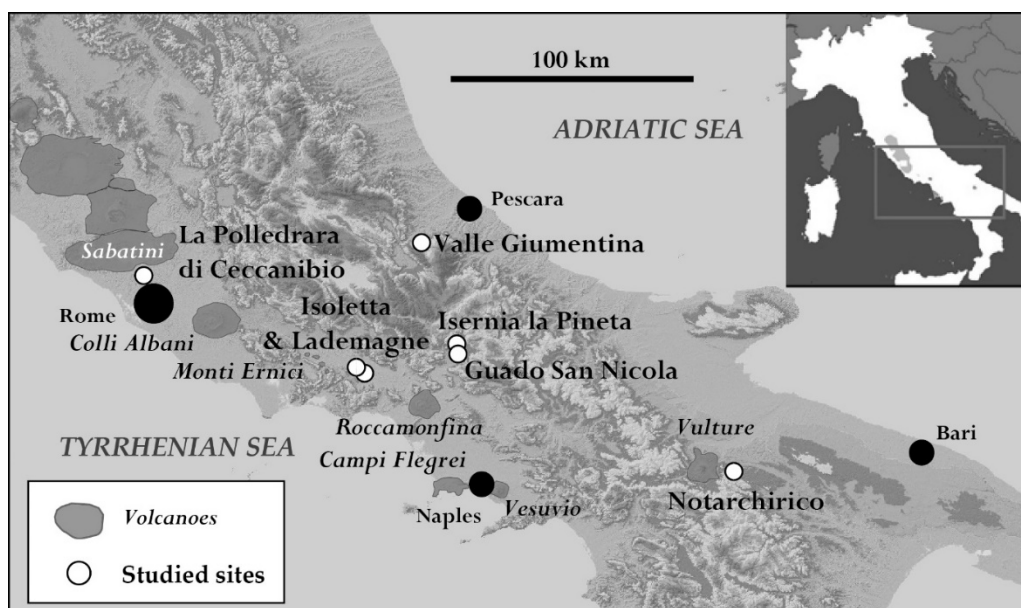
48 However, it is possible in some areas of the European continent to benefit of the combine application of inde-  
49 pendent dating methods. It is the case in Central and Southern Italy where the stratigraphic sequences of numerous  
50 Palaeolithic sites, very rich in both archaeological and palaeontological remains, contain tephra deposits and record  
51 hence the regional volcanic history. The study of such sequences allows hence the comparison of the chronologies  
52 derived from ESR and ESR/U-series analyses on fluvial bleached quartz and tooth enamel respectively with the  
53 chronological framework obtained by  $^{40}\text{Ar}/^{39}\text{Ar}$  dating method on single volcanic K-feldspar crystals. The  $^{40}\text{Ar}/^{39}\text{Ar}$   
54 method became a reference method for the Quaternary chronology as evidenced by the geochronological frame-  
55 work of human evolution in East Africa that was established thanks to this technic (see for example Brown et al.,  
56 2012 or McDougall, 2014).

57 In 2009, in collaboration with Italian and French archaeologists and geologists, we initiated an inter-comparison  
58 geochronological project aiming to date by different methods Middle Pleistocene sites of Central and Southern Italy,

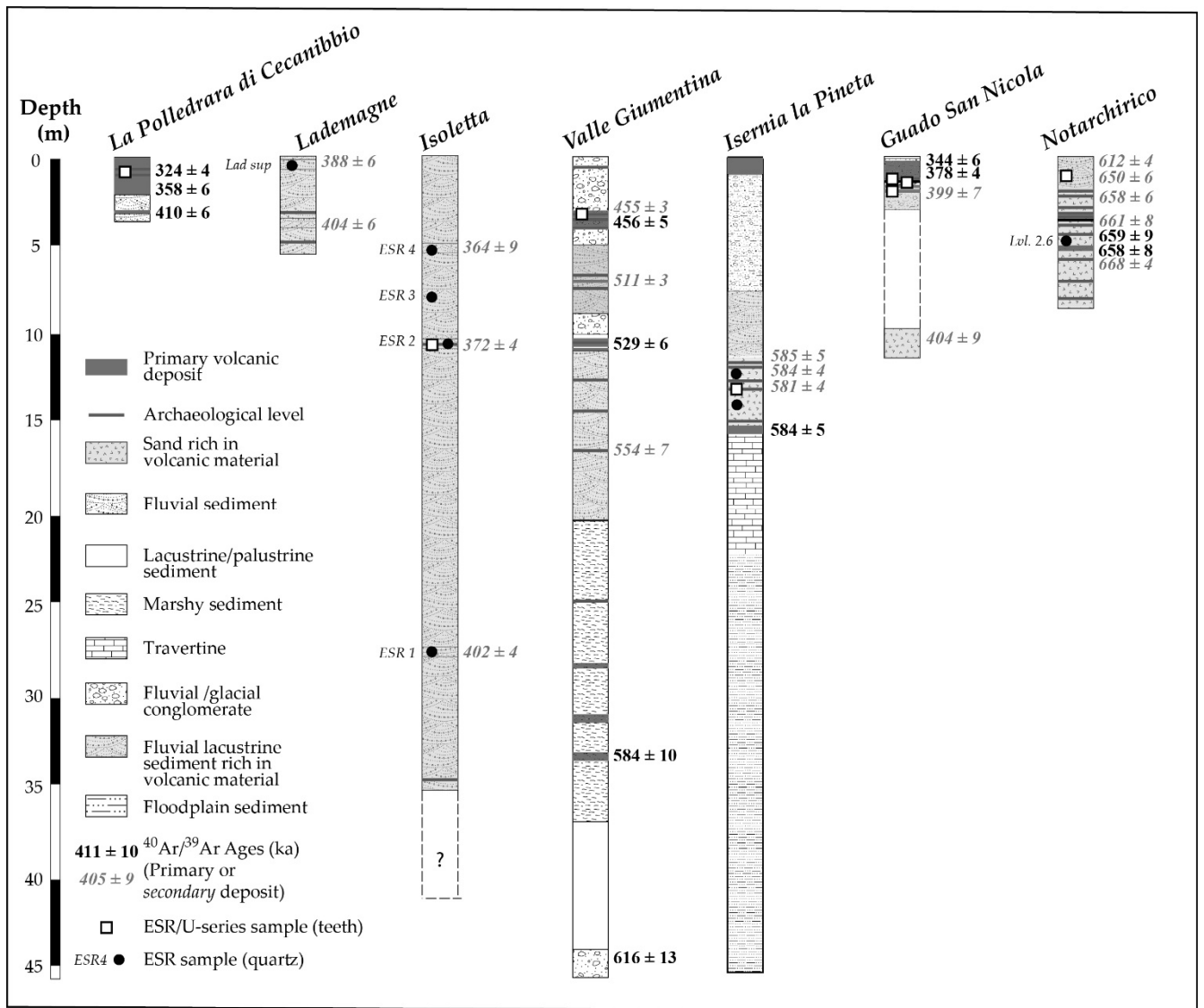
59 which displayed both volcanic materials and archaeological levels. On each site, teeth and sediments containing vol-  
60 canic minerals, both in primary fallout deposits and in reworked sedimentary layers, as well as sediments containing  
61 quartz when present, were systematically sampled. The main results obtained in this inter-comparison project are  
62 displayed in this paper and compared to the data derived from paleo-environmental and geological studies in order  
63 to improve the chronology of the sites dated by only one method or/and only from the available paleo-  
64 environmental data. The present paper synthesizes hence for the first time the data obtained during the project by  
65 ESR/U-series and ESR on these Italian sites and compares them with the  $^{40}\text{Ar}/^{39}\text{Ar}$  chronology established in parallel  
66 for each site, allowing a general discussion on the potential and limits of ESR/U-series and ESR technics applied to  
67 date Middle Pleistocene sites. Even if the main part of the displayed results are already published in different papers,  
68 the main issue of the present methodological article is to compare systematically the dating results obtained by  
69 ESR/U-series and ESR with those derived from  $^{40}\text{Ar}/^{39}\text{Ar}$  analyses, such systematic discussion with comparison having  
70 not really realized before.

## 72 Sites, sampling and $^{40}\text{Ar}/^{39}\text{Ar}$ chronology

73 From the west to the east, the studied sites are La Polledrara di Cecanibbio, Isoletta, Lademagne (Latium), Guado  
74 San Nicola, Isernia La Pineta (Molise), Valle Giumentina (Abruzzo) and Notarchirico (Basilicata) (Figure 1). The strati-  
75 graphic sequences of these sites, display both archaeological levels with, for some of them, abundant palaeontologi-  
76 cal fossils, and either primary volcanic deposits such as tephra, or fluvial or fluvio-lacustrine deposits rich in fresh  
77 volcanic minerals reworked from such primary volcanic deposits (Figure 2).



78  
79 **Figure 1** - Location of the studied Middle Pleistocene sites in Central and Southern Italy.



**Figure 2** – Simplified stratigraphic logs of the studied Middle Pleistocene sites with indication of the levels dated by  $^{40}\text{Ar}/^{39}\text{Ar}$  and positions of the levels sampled for ESR and ESR/U-series studies.  $^{40}\text{Ar}/^{39}\text{Ar}$  ages were obtained directly on tephra levels (primary deposits) or on potassic feldspar grains extracted from fluvial sediments (secondary deposits), the youngest population corresponding then to a maximum age for the dated level.

La Polledrara di Cecanibbio is located only at 20 km from Rome. Discovered in 1984, the site is since excavated yearly and it is currently considered as one of the richest palaeontological and archaeological deposits of the Latium province, Latium (Anzidei et al., 2004, 2012; Santucci et al., 2016). More than 20 000 palaeontological remains, constituting an assemblage attributed to the Aurelian Large Mammal's stage of the Italian bio-chronological sequence (Palombo and Milli, 2005, Marra et al., 2018; Petronio et al., 2019), and at least 600 artefacts were recovered in fluvial and volcanic lahar sediments rich in pyroclastic minerals originated from the Monti Sabatini volcanic complex. Three different stratigraphic layers and seven *Bos primigenius* teeth were sampled for  $^{40}\text{Ar}/^{39}\text{Ar}$  on single grain and ESR/U-series dating methods respectively (Pereira et al., 2017).

94 Isoletta and Lademagne sites are both located 110 km southeast of Rome in the Frosinone-Ceprano tectonic ba-  
95 sin of the Latina Valley, close to the Campogrande locality where was discovered in 1994 the Ceprano human calvar-  
96 ium (Ascenzi et al., 1996, Manzi et al., 2016). Both sequences include mainly fluvial deposits rich in volcanic materi-  
97 als, some tephra layers being deposited also at Isoletta, and several archaeological layers displaying for the upper  
98 one Acheulean artefacts (Biddittu, 2004; Biddittu et al., 2012). Several fluvial levels were sampled at Lademagne  
99 (n=2) and Isoletta (n=4) for both ESR and  $^{40}\text{Ar}/^{39}\text{Ar}$  analyses as well as a cervid tooth from the Acheulean level of  
100 Isoletta (Pereira et al., 2018; Voinchet et al., 2020).

101 Isernia La Pineta, 180 km southeast of Rome, is one of the most famous and ancient palaeolithic sites of Italy. It  
102 was extensively excavated since its discovery in 1978 and has delivered abundant lithic and faunal remains charac-  
103 teristic of the Middle Galerian Large Mammal's stage. The site was found in a complex stratigraphic sequence includ-  
104 ing lacustrine, volcanic, fluvial and slope deposits (Coltorti et al., 1982; Peretto et al., 1983, 2015a). A volcanic unit at  
105 the base of the archaeological sequence and four teeth (bison and rhinoceros) from the main archaeological unit t3c  
106 were sampled as well as sediments (n=2) from the archaeological sequence (Shao et al., 2011; Peretto et al., 2015).

107 The Guado San Nicola site, only 10 km south of Isernia, was discovered in 2005. It site was excavated between  
108 2008 and 2012 allowing the recovering of more than 4000 lithic artefacts and 1500 faunal remains. The palaeonto-  
109 logical assemblage is characteristic of the final Italian Galerian stage while the lithic industry includes both Mode 2  
110 and Mode 3 artefacts (Peretto et al., 2016). The archaeological layers are included into a fluvial sequence covered by  
111 several volcanic units attributed to the Roccamonfina volcanic complex (Peretto et al., 2014, 2016). Three different  
112 layers of the stratigraphic sequence (tephra and fluvial deposits) and six teeth (horses and rhinoceros) were sampled  
113 for  $^{40}\text{Ar}/^{39}\text{Ar}$  and ESR/U-series (6 teeth in total) dating methods respectively (Peretto et al., 2016; Pereira et al.,  
114 2016).

115 Valle Giumentina is located about 120 km east of Rome and displays a 70m thick sedimentary sequence, mainly  
116 constituted by lake shore deposits, in which were found several Palaeolithic levels attributed to the Clactonian and  
117 Acheulean in the lower part of the sequence and Levalloisian cultures In the uppermost part respectively (Demange-  
118 ot and Radmilli, 1953; Nicoud et al., 2015). Numerous tephra layers were also recognized in the stratigraphic (Nicoud  
119 et al., 2015; Villa et al., 2016; Degeai et al., 2018). Seven of these volcanic primary deposits sampled all along the  
120 sequence and one cervid tooth extracted from a mandible carried out from an Acheulean level (ALB-42) were dated  
121 (Nicoud et al., 2015; Villa et al., 2016; Degeai et al., 2018).

122 Lastly, the early Middle Pleistocene site of Notarchirico is one of the most famous archaeological localities of  
123 Southern Italy (Piperno, 1997; Lefèvre et al., 2010). Located 180 km east of Naples close to Monte Vulture stratovol-  
124 cano, the site is constituted by a 7m-thick fluvial sequence including eleven archaeological layers. Some of these  
125 archaeo-surfaces have delivered handaxes considered as the oldest evidence of Acheulean settlement presently  
126 known in Italy (Pereira et al., 2015) and one human femur attributed to *Homo heidelbergensis* (Mallegni et al., 1991).  
127 A tephra layer in the lower part of the sequence and several fluvial sediments reworking volcanic minerals were  
128 sampled for  $^{40}\text{Ar}/^{39}\text{Ar}$ , as well as three quartz-rich levels (unit 2.6) and four teeth from the upper palaeo-  
129 anthropological layer (supra- $\alpha$ ) for ESR and ESR-U-series methods respectively (Voinchet et al., 2020).

130 Details on the analytical  $^{40}\text{Ar}/^{39}\text{Ar}$  protocol applied on single volcanic K-feldspar crystals and the main part of the  
131 results have been already published in detail (Pereira et al., 2015, 2016, 2017, 2018; Peretto et al., 2015, Villa et al.,  
132 2016). These data have permitted to greatly refine the chronology of the studied stratigraphic sequences and associ-  
133 ated archaeological levels of the sites.

134 In the present paper,  $^{40}\text{Ar}/^{39}\text{Ar}$  ages are presented as weighted mean ages with related full external error at  $2\sigma$   
135 (analytical + decay constant uncertainties) and as plotted from an inverse isochron diagram ( $^{36}\text{Ar}/^{40}\text{Ar}$  vs  $^{39}\text{Ar}/^{40}\text{Ar}$ ).  
136 Such diagram permits to assess the isotopic composition of  $^{40}\text{Ar}$  trapped at the time of crystallization of the dated  
137 minerals and hence evaluate the possibility of argon excess by verifying the assumption that trapped argon isotopes  
138 at the eruption time presented the composition of modern atmosphere ( $^{40}\text{Ar}/^{36}\text{Ar}$  ratio of 298.56 (Lee et al., 2006)).  
139 Ages were re-calculated in the present contribution according the monitor flux standard ACs-2 at 1.1891 Ma  
140 (Niespolo et al., 2017) and the K total decay constant of Renne et al. (2011). If the dated samples do not show evi-  
141 dence of such contamination, the calculated inverse isochron age should be equivalent to the weighted mean aver-  
142 age one. Inverse isochron  $^{40}\text{Ar}/^{36}\text{Ar}$  initial ratio as well as the related calculated ages at  $2\sigma$  of uncertainties will be  
143 therefore displayed.

## 145 **ESR and ESR/U-series geochronological studies**

### 146 ESR dating on optically bleached quartz

147 In the case of ESR dating of fluvial quartz grains, the dated event corresponds to the sediment deposition by the  
148 river after the exposure of the quartz grains to the sunlight during the river transport phase prior to this deposition.  
149 This exposure leads to an optical bleaching of the quartz ESR signals.

150 In the present work, both aluminium (Al) and titanium (Ti) impurity centers are used for dating purpose. It is im-  
151 portant to keep in mind that the different Ti signals associated to lithium (Ti-Li) and hydrogen (Ti-H) atoms occur-  
152 rence in the vicinity of titanium ones can be quickly and completely reset by light exposure which is not the case of  
153 Al-signals, probably because of the presence in quartz aluminium centers of "Deep Aluminium Traps" (DAT) that  
154 cannot be emptied by the energy provided by sunlight (Tissoux et al 2013). It is therefore necessary to determine for  
155 each sample the level of "residual" ESR signal intensity, corresponding to the maximum bleaching of the Al-centers  
156 into the quartz grains, by exposure of the quartz grains to artificial solar light in the simulator for ca 1600 hours. The  
157 "unbleachable" ESR intensity hence obtained subtracted from the ESR signal intensities of the other aliquots of the  
158 same sample (including natural one) before the construction of the growth curve prior to any age calculation.

159 An multi-center ESR approach, based on the systematic measurements of both aluminium (Al), titanium-lithium  
160 (Ti-Li) and titanium-hydrogen (Ti-H) signals (Toyoda et al., 2000), was therefore used to date the quartz samples. This  
161 method relies on the differences of sensibility of these ESR centers to light. While Ti centers are totally zeroed by  
162 sunlight exposure within few hours (Ti-H) or days (Ti-Li), Al-center cannot be completely reset even after several  
163 months of exposure. In these conditions, the results obtained for a same sample from these three ESR signals permit  
164 to discuss of the quality of the initial bleaching of the quartz grains before deposition on the site or to indicate con-  
165 taminations by unbleached quartz grains from older levels or bedrock (Duval and Guilarte, 2015 ; Voinchet et al.,  
166 2020).

167 Preparation and measurement protocols used during the present work have been recently displayed by Voinchet  
168 et al. (2020). Several sampled sediments did not contain enough quartz to allow an analysis and only 6 quartz sam-  
169 ples have been finally analysed: 4 from Isoletta, 1 from Lademagne and 1 from Notarchirico. The equivalent doses  
170 and derived ages were calculated from Al- and Ti-centers for all the samples except Notarchirico sample, which not  
171 displayed any Ti-signals, and Lademagne one, for which it was not possible to measure Ti-H signal, undistinguishable  
172 to the ESR background noise. For the other samples, the multi-center approach was therefore used. When  $D_E$  values  
173 obtained from the different centers of a sample were close, a weighted average age was calculated using IsoPlot 3.0  
174 (Ludwig, 2003) with 95 % of confidence.

#### 175 ESR/U-series on teeth

177 In the present work, the analysed teeth corresponds to herbivorous cheek teeth (molars or premolars), mainly  
178 from bovids but also from equids, rhinoceros and cervids (Table 1).



Site	Sample	Sample laboratory number	Type of analysed tooth
<b>La Polledrara di Cecanibbio</b>	Po2012-01	IPH-2012-51	Cheek tooth Bovidae
	Po2012-02	IPH-2012-52	Cheek tooth Bovidae
	Po2012-03	IPH-2012-53	Cheek tooth Bovidae
	Po2012-04	IPH-2012-54	Cheek tooth Bovidae
	Po2012-05	IPH-2012-55	Cheek tooth Bovidae
	Po2012-06	IPH-2012-56	Cheek tooth Bovidae
	Po2012-07	IPH-2012-57	Cheek tooth Bovidae
<b>Guado San Nicola</b>	SN1001	IPH-2010-14	Cheek tooth Bovidae
	SN1002	IPH-2010-15	Cheek tooth Bovidae
	SN0902	IPH-2009-27	Cheek tooth Equidae
	SN0906	IPH-2009-31	Cheek tooth Bovidae
	SN1003	IPH-2010-16	Cheek tooth Rhinocerotidae
	SN1004	IPH-2010-17	Cheek tooth Rhinocerotidae
<b>Isoletta</b>	ISOL1401	IPH-2014-28	Cheek tooth Cervidae
<b>Valle Giumentina</b>	VG1501	IPH-2015-04	Cheek tooth Cervidae
<b>Isernia la Pineta</b>	IS0901	IPH-2009-01	Cheek tooth Bovidae
	IS0902	IPH-2009-02	Cheek tooth Bovidae
	IS0903	IPH-2009-03	Cheek tooth Bovidae
	IS0904	IPH-2009-04	Cheek tooth Bovidae
<b>Venosa Notarchirico</b>	VN1201	IPH-2012-38	Cheek tooth Rhinocerotidae
	VN1202	IPH-2012-39	Cheek tooth Rhinocerotidae
	VN1203	IPH-2012-40	Cheek tooth Bovidae
	VN1204	IPH-2012-41	Cheek tooth Bovidae

**Table 1** – List and nature of the analysed teeth from several Middle Pleistocene Italian sites

Post-mortem uranium-uptake into the dental tissues leads to the variation of the dose rate over time. This phenomenon, which depends on the nature of the site and the age of the sample, makes the determination of the ESR ages, strongly linked to uranium content, particularly delicate and it is necessary to use mathematical models to describe the evolution of this parameter over time. In order to calculate the age of a given sample, the various data acquired through ESR and U-series analyses are then used to model post-mortem U-uptake in the different tissues followed in some cases by a loss of uranium (leaching) in order to provide an unique age. One of the main interests of this approach is to allow the determination for each dental tissue of a U-uptake parameter calculated from the whole set of analytical data and hence the description of the uranium-uptake kinetics into the considered tissue.

The mathematical algorithm for calculating combined ESR/U-series ages has been recently described by Shao et al. (2014). From the U-series isotopic data (U contents, isotopic ratios  $^{234}\text{U}/^{238}\text{U}$ ,  $^{230}\text{Th}/^{234}\text{U}$ ,  $^{222}\text{Rn}/^{230}\text{Th}$ ), a relationship between the incorporation parameter and time is first determined for each dental tissue. It then becomes possible to simulate the evolution over time of the contributions of the different dental tissues to the dose rate. By adding the simulation of the environmental external dose, it becomes possible to simulate the evolution of the total dose rate over time and then to calculate an age by comparing the  $D_E$  value obtained experimentally by ESR with this simulation. From this age, a single value corresponding to the incorporation parameter can be calculated for each

196 tissue. In the case of a simple U-uptake, the age is determined with the *Uranium-series (US) model* (Grün et al.,  
197 1988), while in the case of a subsequent slight leaching, the U-uptake can be described with the *accelerating uptake*  
198 (*AU) model* (Shao et al., 2012).

199 It should also be underlined that the main part of the results discussed in the present paper have been soon pub-  
200 lished in previous papers (La Polledrara di Cekanibbio, Pereira et al., 2017; Guado San Nicola, Bahain et al., 2014,  
201 Pereira et al., 2016; Isernia la Pineta, Shao et al., 2011 ; Isoletta, Pereira et al., 2018) and only results of Venosa No-  
202 tarchirico (4 samples) and Valle Giumentina (one sample) are completely unpublished data.

#### 203 204 *Isochron ESR/U-series dating of teeth*

205 The use of isochrons in teeth ESR dating was proposed in 1993 by Bonnie A. Blackwell and Henry P. Schwarcz to  
206 take into account a possible variation of the external dose over time in the age calculation using fixed of U-uptake  
207 models (Blackwell and Schwarcz, 1993). If it is assumed that the external dose has not varied over time, a diagram  
208 can be drawn showing the equivalent doses determined for several sub-samples of the same tooth as a function of  
209 the annual internal dose determined for the same sub-samples using the considered U-uptake model (*isochron*). If  
210 the points of this isochron are linearly correlated, then the slope of the regression gives the age of the tooth and its  
211 intercept with the y-axis gives the accumulated external dose common to all sub-samples. The procedure can then  
212 be repeated for other external dose values until the best possible correlation between the data is obtained for a  
213 given incorporation pattern. It then becomes possible to discuss the appropriateness of using one intake pattern  
214 relative to another for the samples under consideration (Blackwell and Schwarcz, 1993; Blackwell et al., 2001). Very  
215 cumbersome to implement and very time-consuming, this approach has unfortunately been little used thereafter.

216 More recently, a slightly modified isochronous approach was used to judge the relevance of the dosimetric re-  
217 construction carried out on samples from the Middle Paleolithic site of Biache-Saint-Vaast, France (Bahain et al.,  
218 2015). In this study, several teeth of the same level excavated in the 1980s were dated by ESR/U-series, using for the  
219 external dose an *in situ* value measured during the excavation as the site is no longer accessible today. Two age  
220 groups were then individualized for the analysed teeth, although the archaeological level is very homogeneous from  
221 a paleontological point of view and a mixture of two stocks of different ages seems unlikely.

222 A diagram showing the equivalent dose values as a function of the total internal dose modelled by ESR/U-Th for  
223 samples assumed to be of the same age (*isochron plot*) was therefore produced. As the data were highly linearly  
224 correlated, it was possible to confirm on the one hand the relevance of using the in situ external dose measured

during the excavation and on the other hand to estimate the age of the level under consideration using the equation of the regression line obtained. The intercept of this line with the Y-axis then corresponds to the proportion of the equivalent dose linked to the external dose. By dividing this value by the annual external dose determined today for the level under consideration, it is possible to estimate the isochron age of the latter and to compare it with the other available geochronological data.

Isochron estimates was also applied successfully on La Polledrara samples (Pereira et al., 2017), leading to a more systematic use on other Italian Middle Pleistocene sites when the number of analysed samples allow such use. In the present work isochron age estimates were therefore determined from Guado San Nicola, Isernia and Notarchirico samples.

## Results

Results are displayed by methods in Tables 2 ( $^{40}\text{Ar}/^{39}\text{Ar}$ ), 3 (ESR), 4 and 5 (ESR/U-series) but will be discussed site by site from the youngest one (La Polledrara) to the oldest one (Notarchirico). The  $^{40}\text{Ar}/^{39}\text{Ar}$  results permit to constrain the ages of the dated archaeological sequences and therefore the ages of associated ESR/U-series and ESR samples.

Concerning the ESR dating of quartz grains (Table 3), except for Isoletta ESR 1 sample for which the age is drastically overestimated for both Al and Ti centers, the age estimates derived from both signals for the other samples are globally in agreement with each other and with the  $^{40}\text{Ar}/^{39}\text{Ar}$  ages obtained for the corresponding levels (Voinchet et al., 2020). While the number of samples is too small to allow a definitive conclusion on this point, these results suggest that the ESR multi-center protocol used here permit to obtain reliable chronology for Middle Pleistocene fluvial sediments. The dose rate was calculated from the radionuclides activities derived both from *in situ* gamma-ray spectrometry measurements using NaI detector and laboratory high resolution and low background gamma spectrometer using a hyper-pure Ge crystal.

Site	Sample	Stratigraphical location of analysed samples	$^{40}\text{Ar}/^{39}\text{Ar}$ ages (ka)	Mean Square Weighted Deviation	Probability fit	Isochron ages (ka)	$^{40}\text{Ar}/^{36}\text{Ar}$ initial ratio
La Polledrara di Cecanibbio (Pereira et al., 2017)	POL 1203 & 1201	Pumices inside the archaeological level	<b>324 ± 4</b>	0.7	0.9	<b>323 ± 7</b>	296 ± 3
	POL1202	Pumices from fluvial deposit just below archaeological level	<b>358 ± 6</b>	0.8	0.5	<b>368 ± 18</b>	249 ± 80
	POL1301	Tephra 1m below archaeological level	<b>410 ± 6</b>	0.3	1.0	<b>410 ± 6</b>	295 ± 9

<b>Guado San Nicola</b> (Pereira et al., 2016)	S.U. tufi	Volcanic level at the top of the sequence	<b>344 ± 6</b>	0.2	0.6	<b>339 ± 15</b>	299 ± 8
	S.U. A	Volcanic level just capping the archaeological sequence	<b>378 ± 4</b>	1.0	0.4	<b>376 ± 5</b>	307 ± 11
	S.U. C	Fluvial deposits inside the Lower archaeological level	<b>399 ± 7</b>	0.5	0.9	<b>400 ± 6</b>	292 ± 15
	Volcanic sands	Fluvial deposits 10m below the archaeological sequence	<b>404 ± 9</b>	1.3	0.2	<b>405 ± 10</b>	282 ± 11
<b>Isoletta</b> (Pereira et al., 2018)	ESR 4	Fluvial deposits 5m above the Acheulean archaeological level	<b>364 ± 9</b>	0.3	0.8	/	/
	GA6Z	Sediments of the Acheulean archaeological level	<b>374 ± 4</b>	1.4	0.2	<b>374 ± 4</b>	298 ± 2
	ESR 1	Fluvial sands at the bottom of the studied sequence	<b>402 ± 4</b>	0.9	0.6	<b>400 ± 10</b>	300 ± 4
<b>Lademagne</b> (Pereira et al., 2018)	LAD sup	Sediments at the top of the archaeological sequence	<b>388 ± 6</b>	1.6	0.1	<b>390 ± 7</b>	289 ± 11
	LAD inf	Sediments from a layer located between two archaeological levels	<b>404 ± 6</b>	1.1	0.4	<b>405 ± 7</b>	293 ± 6
<b>Valle Giumentina</b> (Nicoud et al., 2015; Villa et al., 2016; Degeai et al., 2018)	T103b	Sediments of the archaeological level LABM	<b>455 ± 3</b>	0.7	0.7	<b>454 ± 3</b>	297 ± 3
	T109b	Sediments of the archaeological level LAN2	<b>529 ± 6</b>	1.3	0.3	<b>528 ± 6</b>	300 ± 3
	T115	Sediments of the archaeological level LN	<b>554 ± 7</b>	0.6	0.8	<b>554 ± 6</b>	294 ± 1.4
	T32	Tephra below the archaeological sequence	<b>584 ± 10</b>	0.4	0.9	<b>581 ± 17</b>	305 ± 30
	T45	Tephra at the bottom of the sedimentary sequence	<b>616 ± 13</b>	0.2	1.0	<b>617 ± 14</b>	298 ± 4
<b>Isernia la Pineta</b> (Peretto et al., 2015)	3s6-9	Sediments sampled above the main archaeological level t3a	<b>585 ± 5</b>	1.0	0.5	<b>585 ± 5</b>	291 ± 7
	3s10	Sediments sampled above the main archaeological level t3a	<b>581 ± 7</b>	1.4	0.2	<b>581 ± 8</b>	296 ± 4
	3coll	Sediments capping the main archaeological level t3a	<b>581 ± 4</b>	1.2	0.3	<b>583 ± 5</b>	290 ± 10
	U4T	Tephra just below the archaeological sequence	<b>584 ± 5</b>	0.5	1.0	<b>583 ± 7</b>	305 ± 30
<b>Venosa Notarchirico</b> (Pereira et al., 2015; Voinchet et al., 2020)	NOT1.6	Top of the archaeological sequence	<b>612 ± 4</b>	1.5	0.2	<b>612 ± 4</b>	296 ± 2
	NOT 1.5	Sediments capping supra α level	<b>650 ± 6</b>	0.6	1.1	<b>651 ± 6</b>	295 ± 3
	NOT 1.3	Just below supra α level	<b>658 ± 6</b>	0.1	1.2	<b>657 ± 11</b>	298 ± 36
	NOT 2.6	Just below D level	<b>661 ± 8</b>	0.7	0.6	<b>658 ± 9</b>	294 ± 9
	NOT 2.2	Tephra between level E and F	<b>659 ± 9</b>	0.3	1.0	<b>663 ± 10</b>	291 ± 6
	NOT 2.1	Tephra between level E and F	<b>658 ± 8</b>	0.6	0.6	<b>657 ± 8</b>	304 ± 14
	U.3	Sediments just below F level	<b>668 ± 4</b>	1.8	0.1	<b>668 ± 4</b>	297 ± 4

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**Table 2** –  $^{40}\text{Ar}/^{39}\text{Ar}$  ages obtained from volcanic minerals recovered from several Middle Pleistocene Italian sites.

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The indicated  $^{40}\text{Ar}/^{39}\text{Ar}$  age corresponds to the age of the youngest homogeneous potassic feldspars population. These ages

252

were re-calculated in the present contribution according to the monitor flux standard ACs-2 at 1.1891 Ma (Niespolo et al., 2017)

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and the K total decay constant of Renne et al. (2011) from the original data published in the various referred papers. A homogeneous

254

population is considered relevant when the weighted mean of these crystals has the following statistical characteristics:

255

MSWD < 1.5, Probability fit  $\geq 0.1$ . The weighted average ages were calculated using IsoPlot 3.0 (Ludwig, 2003) and given at 95%

256

(2 $\sigma$ ) of confidence.

257

Site	Sample	ESR signal	$D_{\alpha}$ ( $\mu\text{Gy/a}$ )	$D_{\beta}$ ( $\mu\text{Gy/a}$ )	$D_{\gamma}$ ( $\mu\text{Gy/a}$ )	$D_{\text{cosmic}}$ ( $\mu\text{Gy/a}$ )	W (%)	$\delta_{\text{bl}}$ (%)	$D_{\text{a}}$ ( $\mu\text{Gy/a}$ )	$D_{\text{E}}$ (Gy)	Age (ka)	Mean Age (ka)	$^{40}\text{Ar}/^{39}\text{Ar}$ Age (ka)
<b>Isoletta</b> (Pereira et al., 2018; Voinchet et al., 2020)	Isoletta ESR 1	Ti-H	19±1	190±11	168±10	24±1	15	100	401±15	325 ± 29	810 ± 47	-	402 ± 4
		Ti-Li	19±1	190±11	168±10	24±1	15	100	401±15	715±75	1786 ± 327		
		Al	19±1	190±11	168±10	24±1	15	48	401±15	462±24	1154 ± 109		
	Isoletta	Ti-H	57±2	1786±36	1008±28	33±2	14	100	2884±15	295 ± 20	102 ± 16	442±58	374 ± 4

	ESR 2	Ti-Li	57±2	1786±36	1008±28	33±2	14	100	2884±15	1155±110	401 ± 59	349±26	364 ± 9 < X < 374 ± 4
		Al	57±2	1786±36	1008±28	33±2	14	42	2884±15	1315±62	456 ± 34		
	Isoletta ESR 3	Ti-H	30±2	497±20	354±16	53±3	12	100	934±26	349±24	374±17		
		Ti-Li	30±2	497±20	354±16	53±3	12	100	934±26	314±73	336 ± 66		
	Isoletta ESR 4	Al	30±2	497±20	354±16	53±3	12	46	934±26	338±60	362 ± 64		
		Ti-H	20±1	303±12	220±10	101±5	15	100	644±16	243±15	374±84		
<b>Lademagne</b> (Pereira et al., 2018 ; Voinchet et al., 2020)	Lademagne Sup	Ti-H	Unmeasurable									402±71	388 ± 6
		Ti-Li	77±3	980±35	919±52	111±6	5	100	2088±70	831±70	398 ± 51		
		Al	77±3	980±35	919±52	111±6	5	44	2088±70	847±100	406 ± 51		
<b>Venosa Notarchirico</b> (Pereira et al., 2015; Voinchet et al., 2020)	Level 2-6	Ti-H	Unmeasurable									-	612 ± 4 < X < 661 ± 8
		Ti-Li	Unmeasurable										
		Al	84±1	1999±24	1803±21	166±8	10	57	4052±32	2660±122	657 ± 31		

**Table 3** – ESR data and ages for sediments recovered from several Middle Pleistocene Italian sites.

Equivalent doses ( $D_e$ ) were derived from the obtained intensity growth curves using an exponential + linear function for Al and Ti-Li centers (Duval et al., 2009; Voinchet et al., 2013) and by a single saturating exponential function for Ti-H from the six first points of the growth curves (Voinchet et al., 2020) with Microcal OriginPro 8 software, both with  $1/I^2$  weighting (according with Yokoyama et al., 1985). Age calculations were performed using the following parameters: dose-rate conversions factors from Guérin et al. (2011); a k-value of  $0.15 \pm 0.1$  (Laurent et al., 1998); alpha and beta attenuations from Brennan (2003) and Brennan et al. (1991); water attenuation formulae from Grün (1994); cosmic dose rate estimated from the Prescott and Hutton's equations (1994). The internal dose rate was considered as negligible because of the low contents of radionuclides usually found in quartz grains (Murray and Roberts 1997; Vandenbergue et al. 2008). ESR age estimates are given with one sigma error range.

Concerning the ESR/U-series results (Tables 4 and 5), when several teeth from a same archaeological level were analysed, the obtained age estimates are displayed both by age density probability plots and isochron plots (Figure 3) and are compared with the available  $^{40}\text{Ar}/^{39}\text{Ar}$  and ESR dates (Figure 4). The age density plots build from the ESR/U-series age estimated with IsoPlot 3.0 software (Ludwig, 2003) permit to observe the homogeneity of the obtained age results and to determine mean ages for each eventual age populations. In complement, the isochron approach (Blackwell and Schwarcz, 1993) allows the evaluation of the quality of the dose rate reconstruction by comparison of the respective weight of accumulated external and internal dose contributions, the last one evolving with time according with the U-uptake modelling in each dental tissue. It permits to estimate an age estimate for a given level by plotting the accumulated internal doses modelled for the teeth vs the  $D_e$  and by dividing the intercept value ( $x=0$ , that represents the corresponding accumulated external doses) by the "real" external dose rates used for the age calculation. These two graphic representations permit to better discuss of the results.

Site	Samples	Tissue	U content (ppm)	$^{234}\text{U}/^{238}\text{U}$	$^{230}\text{Th}/^{232}\text{Th}$	$^{230}\text{Th}/^{234}\text{U}$	Apparent U-series age (ka)	$^{222}\text{Rn}/^{230}\text{T}$ h	Initial thickness ( $\mu\text{m}$ )	Removed thickness Internal side ( $\mu\text{m}$ )	Removed thickness External side ( $\mu\text{m}$ )
La Polledrara di Cecanibio (Pereira et al., 2017)	Po2012-01	dentine	342.83 $\pm$ 9.12	1.215 $\pm$ 0.008	20	0.658 $\pm$ 0.030	112 $\pm$ 9	0.351	1270 $\pm$ 159	172 $\pm$ 21	32 $\pm$ 4
		enamel	9.89 $\pm$ 0.26	1.131 $\pm$ 0.021	> 100	0.760 $\pm$ 0.034	149+16/-13	0.612			
	Po2012-02	dentine	294.18 $\pm$ 9.96	1.172 $\pm$ 0.033	48	0.617 $\pm$ 0.027	101 +8/-7	0.395	1323 $\pm$ 165	116 $\pm$ 14	6 $\pm$ 1
		enamel	10.11 $\pm$ 0.33	1.148 $\pm$ 0.039	13	0.798 $\pm$ 0.039	164+23/-18	1.000			
	Po2012-03	dentine	283.76 $\pm$ 8.81	1.215 $\pm$ 0.031	> 50	0.575 $\pm$ 0.025	90+7/-6	0.390	1124 $\pm$ 140	191 $\pm$ 24	86 $\pm$ 11
		enamel	23.52 $\pm$ 0.83	1.153 $\pm$ 0.044	46	0.629 $\pm$ 0.032	105+10/-9	0.321			
	Po2012-04	dentine	326.73 $\pm$ 8.79	1.209 $\pm$ 0.026	> 100	0.646 $\pm$ 0.026	108+8/-7	0.426	1394 $\pm$ 174	43 $\pm$ 5	114 $\pm$ 14
enamel		13.80 $\pm$ 0.49	1.387 $\pm$ 0.053	> 50	0.704 $\pm$ 0.034	122+12/-11	0.497				
Po2012-05	dentine	273.00 $\pm$ 9.22	1.239 $\pm$ 0.036	> 50	0.633 $\pm$ 0.033	105+10/-9	0.533	1238 $\pm$ 155	109 $\pm$ 14	78 $\pm$ 10	
	enamel	9.98 $\pm$ 0.28	1.185 $\pm$ 0.023	> 50	0.699 $\pm$ 0.033	125+12/-11	0.557				
Po2012-06	dentine	340.53 $\pm$ 11.61	1.252 $\pm$ 0.034	> 100	0.503 $\pm$ 0.027	74 $\pm$ 6	0.385	1339 $\pm$ 167	225 $\pm$ 28	155 $\pm$ 19	
	enamel	8.92 $\pm$ 0.25	1.164 $\pm$ 0.024	> 100	0.679 $\pm$ 0.037	119+13/-12	0.650				
Po2012-07	dentine	396.22 $\pm$ 13.70	1.062 $\pm$ 0.030	> 100	0.654 $\pm$ 0.035	114+12/-11	0.279	1376 $\pm$ 172	138 $\pm$ 17	75 $\pm$ 9	
	enamel	5.72 $\pm$ 0.24	1.473 $\pm$ 0.061	> 100	0.724 $\pm$ 0.035	127+13/-11	0.780				
Guado San Nicola (Bahain et al., 2014; Pereira et al., 2016)	SN1001	cement	40.76 $\pm$ 0.09	1.319 $\pm$ 0.004	> 100	1.133 $\pm$ 0.008	> 500	0.290	1053 $\pm$ 132	68 $\pm$ 8	12 $\pm$ 1
		dentine	103.67 $\pm$ 0.45	1.253 $\pm$ 0.003	> 100	0.950 $\pm$ 0.011	251 $\pm$ 12	0.407			
		enamel	2.45 $\pm$ 0.01	1.303 $\pm$ 0.004	42	0.897 $\pm$ 0.011	207 $\pm$ 8	0.796			
	SN1002	cement	108.95 $\pm$ 0.44	1.234 $\pm$ 0.004	> 100	0.840 $\pm$ 0.007	180 $\pm$ 5	0.352	1204 $\pm$ 151	67 $\pm$ 8	13 $\pm$ 2
		dentine	121.74 $\pm$ 0.45	1.248 $\pm$ 0.004	> 100	0.831 $\pm$ 0.009	175 $\pm$ 6	0.352			
		enamel	2.96 $\pm$ 0.01	1.281 $\pm$ 0.004	> 50	0.765 $\pm$ 0.008	145 $\pm$ 4	0.705			
SN0902	cement	38.99 $\pm$ 0.11	1.319 $\pm$ 0.003	> 50	1.167 $\pm$ 0.007	> 500	0.307	1273 $\pm$ 159	32 $\pm$ 4	80 $\pm$ 10	
	dentine	140.33 $\pm$ 0.43	1.204 $\pm$ 0.003	> 100	0.855 $\pm$ 0.005	190 $\pm$ 4	0.419				
enamel	3.38 $\pm$ 0.01	1.223 $\pm$ 0.005	100	0.786 $\pm$ 0.005	155 $\pm$ 3	0.662					
SN0906	dentine	67.27 $\pm$ 0.27	1.338 $\pm$ 0.006	> 100	1.048 $\pm$ 0.010	365 $\pm$ 31	0.085	1793 $\pm$ 224	48 $\pm$ 6	58 $\pm$ 7	
	enamel	1.35 $\pm$ 0.01	1.254 $\pm$ 0.006	7	1.114 $\pm$ 0.010	> 500	0.779				
SN1003	dentine	86.26 $\pm$ 0.21	1.328 $\pm$ 0.004	> 100	1.055 $\pm$ 0.009	384 $\pm$ 31	0.363	2328 $\pm$ 291	56 $\pm$ 7	28 $\pm$ 4	
	enamel	0.50 $\pm$ 0.01	1.230 $\pm$ 0.006	3	0.934 $\pm$ 0.012	241 $\pm$ 14	0.790				
SN1004	dentine	83.68 $\pm$ 0.38	1.318 $\pm$ 0.003	> 100	1.051 $\pm$ 0.010	379 $\pm$ 29	0.411	1896 $\pm$ 237	67 $\pm$ 8	40 $\pm$ 5	
	enamel	0.71 $\pm$ 0.01	1.402 $\pm$ 0.005	16	0.918 $\pm$ 0.011	213 $\pm$ 9	0.644				
Isernia la Pineta (Shao et al., 2011)	IS0901	cement	120.33 $\pm$ 0.77	1.327 $\pm$ 0.009	> 100	0.770 $\pm$ 0.018	146 $\pm$ 9	0.265	1098 $\pm$ 137	68 $\pm$ 9	108 $\pm$ 13
		dentine	225.20 $\pm$ 1.07	1.314 $\pm$ 0.005	> 100	0.720 $\pm$ 0.008	129 $\pm$ 3	0.255			
		enamel	3.78 $\pm$ 0.02	1.355 $\pm$ 0.001	26	0.673 $\pm$ 0.023	114 $\pm$ 9	0.647			
	IS0902	dentine	193.61 $\pm$ 0.95	1.276 $\pm$ 0.005	> 100	0.695 $\pm$ 0.010	122 $\pm$ 4	0.250	1269 $\pm$ 159	153 $\pm$ 19	88 $\pm$ 11
		enamel	3.25 $\pm$ 0.01	1.343 $\pm$ 0.001	12	0.691 $\pm$ 0.031	120 $\pm$ 2	0.479			
	IS0903	cement	216.03 $\pm$ 1.04	1.253 $\pm$ 0.010	> 50	0.621 $\pm$ 0.014	146 $\pm$ 9	0.182	1212 $\pm$ 151	115 $\pm$ 14	53 $\pm$ 7
dentine		237.29 $\pm$ 1.07	1.263 $\pm$ 0.005	> 100	0.574 $\pm$ 0.009	129 $\pm$ 3	0.238				
enamel	4.39 $\pm$ 0.01	1.320 $\pm$ 0.001	43	0.543 $\pm$ 0.016	82 $\pm$ 4	0.564					
IS0904	cement	126.93 $\pm$ 0.97	1.362 $\pm$ 0.013	22	0.901 $\pm$ 0.015	205 $\pm$ 13	0.173	1323 $\pm$ 165	243 $\pm$ 30	101 $\pm$ 13	
	dentine	201.84 $\pm$ 1.14	1.276 $\pm$ 0.005	> 100	0.686 $\pm$ 0.009	119 $\pm$ 4	0.239				
enamel	2.62 $\pm$ 0.01	1.343 $\pm$ 0.001	9	0.683 $\pm$ 0.037	118 $\pm$ 14	0.598					
Venosa Notarchirico Level supra $\alpha$ (this work)	VN1201	dentine	246.25 $\pm$ 1.77	1.153 $\pm$ 0.009	20	0.892 $\pm$ 0.010	217 $\pm$ 8	0.972	1254 $\pm$ 157	283 $\pm$ 35	88 $\pm$ 11
		enamel	3.19 $\pm$ 0.01	1.153 $\pm$ 0.001	> 50	0.532 $\pm$ 0.002	81 $\pm$ 1	1.000			
	VN1202	dentine	216.96 $\pm$ 1.68	1.164 $\pm$ 0.011	36	0.743 $\pm$ 0.012	141 $\pm$ 5	0.842	1469 $\pm$ 184	255 $\pm$ 32	58 $\pm$ 7
		enamel	11.10 $\pm$ 0.04	1.174 $\pm$ 0.006	35	0.638 $\pm$ 0.005	107 $\pm$ 1	1.000			
VN1203	cement	158.83 $\pm$ 1.23	1.246 $\pm$ 0.008	25	1.020 $\pm$ 0.013	342 $\pm$ 26	0.733	1105 $\pm$ 138	92 $\pm$ 12	124 $\pm$ 16	
	dentine	197.72 $\pm$ 2.32	1.307 $\pm$ 0.012	39	1.056 $\pm$ 0.017	398 $\pm$ 54	0.638				
enamel	7.65 $\pm$ 0.04	1.307 $\pm$ 0.007	20	0.965 $\pm$ 0.008	256 $\pm$ 8	0.515					
VN1204	dentine	140.80 $\pm$ 0.73	1.458 $\pm$ 0.008	> 50	1.078 $\pm$ 0.008	380 $\pm$ 21	0.679	1172 $\pm$ 146	190 $\pm$ 24	163 $\pm$ 20	
	enamel	5.77 $\pm$ 0.01	1.306 $\pm$ 0.003	> 50	1.140 $\pm$ 0.004	> 500	0.671				
Isoletta (Pereira et al., 2018)	ISOL1401	dentine	109.86 $\pm$ 9.97	1.281 $\pm$ 0.027	31	0.865 $\pm$ 0.036	190 +23/-19	0.865	1078 $\pm$ 100	98 $\pm$ 12	67 $\pm$ 8
enamel	1.65 $\pm$ 0.07	1.316 $\pm$ 0.063	> 100	0.953 $\pm$ 0.057	245 +76/-44	0.953					
Valle Giumentina (this work)	VG1501	dentine	53.67 $\pm$ 1.67	1.426 $\pm$ 0.038	> 50	1.078 $\pm$ 0.052	> 300	0.898	1063 $\pm$ 100	118 $\pm$ 15	109 $\pm$ 14
enamel	1.60 $\pm$ 0.07	1.316 $\pm$ 0.060	4	1.058 $\pm$ 0.061	> 300	0.364					

279 **Table 4** – U-series and ESR data for the analysed teeth recovered from several Middle Pleistocene Italian sites.

280 U-series analyses were performed on each dental tissue in order to determine the U-uptake parameters necessary to the dose  
281 rate contributions and age calculations (see details in Shao et al., 2015b), either by alpha spectrometry (MNHN, Paris), or by  
282 Neptune Multi-Collector Inductively Coupled Plasma Mass Spectrometer (MC-ICPMS) (Nanjing Normal University, China).

Site	Samples	Tissue	D <sub>E</sub> (Gy)	U uptake parameters p (regular) or n ( <i>italics</i> )	D <sub>a</sub> α internal (μGy/a)	D <sub>a</sub> β (μGy/a)	D <sub>a</sub> (γ + cosm) (μGy/a)	D <sub>a</sub> total (μGy/a)	ESR/U-series (US or AU) ages (ka)	<sup>40</sup> Ar/ <sup>39</sup> Ar ages, ESR/U-series density plot and ESR/U-series isochron age estimates (ka)
La Polledrara di Cecanibbio (Pereira et al., 2017)	Po2012-01	dentine enamel	1665 ± 169	-0.360 ± 0.151 0.280 ± 0.248	1779 ± 868	708 ± 183	2716 ± 40	5203 ± 888	320 ± 44	<sup>40</sup> Ar/ <sup>39</sup> Ar age 324 ± 4
	Po2012-02	dentine enamel	1902 ± 105	-0.593 ± 0.027 0.221 ± 0.172	2691 ± 696	1150 ± 212	2716 ± 40	6557 ± 729	290 ± 28	
	Po2012-03	dentine enamel	1845 ± 141	0.208 ± 0.234 0.580 ± 0.287	2361 ± 487	1074 ± 52	2716 ± 40	6151 ± 702	300 ± 29	ESR/U-series density plot ages Two populations 196 ± 25 (n=2) 324 ± 25 (n=4)
	Po2012-04	dentine enamel	1718 ± 48	-0.683 ± 0.112 -0.538 ± 0.135	3008 ± 955	1902 ± 530	2716 ± 40	7626 ± 1093	183 ± 23	
	Po2012-05	dentine enamel	1421 ± 89	-0.076 ± 0.088 0.304 ± 0.114	1472 ± 314	1230 ± 200	2716 ± 40	5418 ± 375	317 ± 20	
	Po2012-06	dentine enamel	1500 ± 104	-0.128 ± 0.137 1.076 ± 0.263	1388 ± 528	863 ± 233	2716 ± 40	4968 ± 578	286 ± 28	ESR/U-series isochron age Two populations ≈ 153 (n=2) ≈ 321 (n=4)
	Po2012-07	dentine enamel	1248 ± 35	-0.641 ± 0.068 -0.489 ± 0.084	1566 ± 338	1928 ± 360	2716 ± 40	6211 ± 495	201 ± 15	
Guado San Nicola (Bahain et al., 2014; Pereira et al., 2016)	SN1001	cement dentine enamel	1029 ± 31	-0.0051 ± 0.0005 -0.0042 ± 0.0005 -0.0038 ± 0.0005	643 ± 154	1211 ± 211	1560 ± 150	3408 ± 301	302 ± 25	<sup>40</sup> Ar/ <sup>39</sup> Ar ages 378 ± 4 < X < 399 ± 7
	SN1002	cement dentine enamel	1153 ± 20	-0.549 ± 0.106 -0.518 ± 0.111 -0.250 ± 0.150	584 ± 233	1170 ± 329	1560 ± 150	3304 ± 430	349 ± 45	
	SN0902	cement dentine enamel	1193 ± 45	-0.0038 ± 0.0001 -0.0025 ± 0.0001 -0.0015 ± 0.0001	605 ± 160	1072 ± 223	1560 ± 150	3237 ± 313	369 ± 33	ESR/U-series density plot ages Two populations 311 ± 35 (n=3) 364 ± 38 (n=3)
	SN0906	dentine enamel	815 ± 26	-0.0042 ± 0.0001 -0.0044 ± 0.0001	438 ± 122	563 ± 96	1560 ± 150	2554 ± 216	319 ± 25	
	SN1003	dentine enamel	698 ± 74	-0.0037 ± 0.0001 -0.0033 ± 0.0001	128 ± 272	283 ± 126	1560 ± 150	1971 ± 335	354 ± 47	ESR/U-series isochron age Two populations Non calculable (n=3) ≈ 366 (n=3)
	SN1004	enamel	887 ± 45	-0.0035 ± 0.0004 -0.0028 ± 0.0004	204 ± 107	605 ± 172	1560 ± 150	2365 ± 227	373 ± 35	
Isernia la Pineta (Shao et al., 2011)	IS0901	cement dentine enamel	1645 ± 113	0.071 ± 0.167 0.385 ± 0.210 0.743 ± 0.256	448 ± 37	1264 ± 76	1910 ± 50	3622 ± 99	456 ± 54	<sup>40</sup> Ar/ <sup>39</sup> Ar age X < 584 ± 5
	IS0902	dentine enamel	1238 ± 61	0.323 ± 0.184 0.362 ± 0.187	401 ± 28	713 ± 43	1990 ± 50	3104 ± 71	400 ± 42	
	IS0903	cement dentine enamel	1360 ± 88	1.275 ± 0.343 1.796 ± 0.413 2.200 ± 0.466	284 ± 30	888 ± 66	1720 ± 50	2892 ± 88	471 ± 59	ESR/U-series density plot ages 435 ± 48 (n=3) ESR/U-series isochron age ≈ 585 (n=3)
	IS0904	cement dentine enamel	1455 ± 89	-0.551 ± 0.072 0.563 ± 0.206 0.599 ± 0.209	321 ± 23	1043 ± 42	1945 ± 50	3309 ± 69	441 ± 46	
Venosa Notarchirico (this work)	VN1201	dentine enamel	751 ± 20	-1.000 -0.151 ± 0.282	445 ± 127	2075 ± 546	1092 ± 55	3612	208 ± 32	<sup>40</sup> Ar/ <sup>39</sup> Ar 612 ± 4 < X < 658 ± 8
	VN1202	dentine enamel	1596 ± 77	0.479 ± 0.217 1.475 ± 0.344	1128 ± 327	701 ± 165	1092 ± 55	2921 ± 370	546 ± 64	
	VN1203	cement dentine enamel	2078 ± 79	-0.0044 ± 0.0006 -0.0045 ± 0.0006 -0.898 ± 0.024	2352 ± 487	3402 ± 502	1092 ± 55	6836 ± 702	304 ± 29	ESR/U-series density plot ages Non calculable ESR/U-series isochron age ≈ 134 (n=3)
	VN1204	dentine enamel	1889 ± 80	-0.0029 ± 0.0003 -0.0032 ± 0.0003	2026 ± 329	1350 ± 200	1092 ± 55	4477 ± 389	422 ± 32	
Isoletta (Pereira et al., 2018)	ISOL1401	dentine enamel	1151 ± 33	-0.578 ± 0.063 -0.0031 ± 0.0004	553 ± 125	1063 ± 194	1494 ± 150	3110 ± 275	370 ± 31	<sup>40</sup> Ar/ <sup>39</sup> Ar age 374 ± 4
Valle Giumentina (this work)	VG1501	dentine enamel	1577 ± 51	-0.0027 ± 0.0002 -0.0027 ± 0.0002	405 ± 110	800 ± 175	2307 ± 150	3512 ± 254	449 ± 29	<sup>40</sup> Ar/ <sup>39</sup> Ar age X < 455 ± 3

283 **Table 5** – ESR/U-series data and ages for the analysed teeth recovered from several Middle Pleistocene Italian sites.

284 The equivalent doses D<sub>E</sub> were extrapolated from the obtained dose-response data sets using either The growth curves were built

285 for each sample using different fitting functions (single saturating exponential, SSE, according with Apers et al., 1981; double

286 saturating exponential function (DSE), according with Duval et al. (2009), or exponential plus linear function (E+L), according  
287 with Shao et al. (2015a). The equivalent dose corresponding to the function that best describes the experimental data and gives  
288 the best statistics is then used to calculate age.

289 The radioelement contents of sediment samples associated to each tooth were determined by in laboratory high resolution low  
290 background gamma spectrometry and *in situ* gamma measurements with TL Al<sub>2</sub>O<sub>3</sub> dosimeters at La Polledrara and Inspector  
291 1000 Canberra gamma spectrometer on the other sites. A sediment water content value of 15 ± 5 % was therefore assumed for  
292 the age calculations. The cosmic dose rate was estimated using the formula of Prescott and Hutton (1994). Were also used the  
293 following parameters: a k-value ( $\alpha$  efficiency) of 0.13 ± 0.02 (Grün and Katzenberger-Apel, 1994); water content of 0 wt% in the  
294 enamel and 7 wt% in the dentine and cementum; conversion contents-doses factors from Guérin et al.; (2011). For each dental  
295 tissue, Rn loss was estimated from both gamma and alpha/ICP measurements (Bahain et al., 1992). The beta dose contributions  
296 were corrected from the enamel part destroyed on each side of the enamel layer during the preparation process (according to  
297 Brennan et al., 1997).

298 ESR/U-series ages, different dose-rate contributions and U-uptake parameters were then calculated using either US model (Grün  
299 et al., 1988) or AU model (Shao et al., 2012). The “USESR”, “AUESR” and “combined ESR” computer programs were used for the  
300 age calculation in which the age uncertainty (1 $\sigma$ ) is calculated with Monte Carlo approach (Shao et al., 2014).

301 Isochron age estimates were newly determined in the present work except for La Polledrara samples (Pereira et al., 2017).

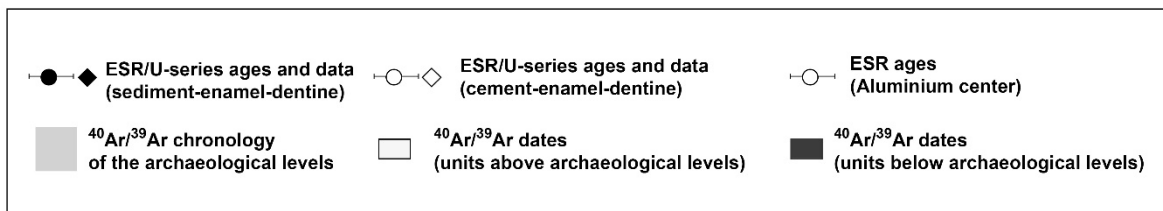
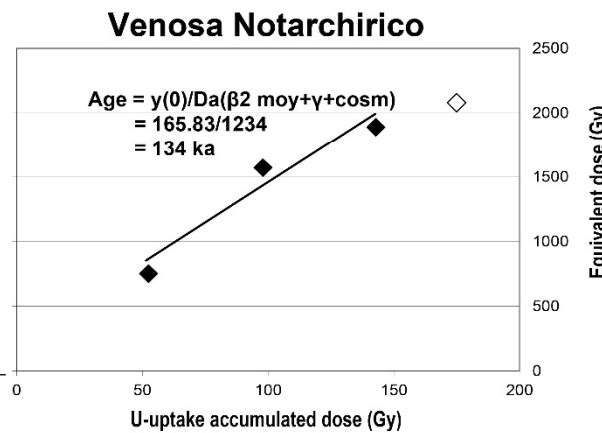
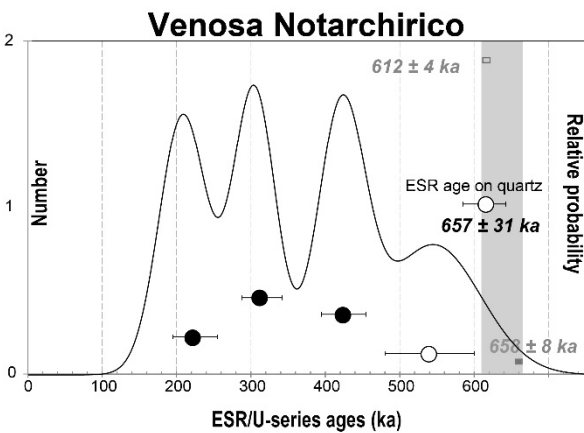
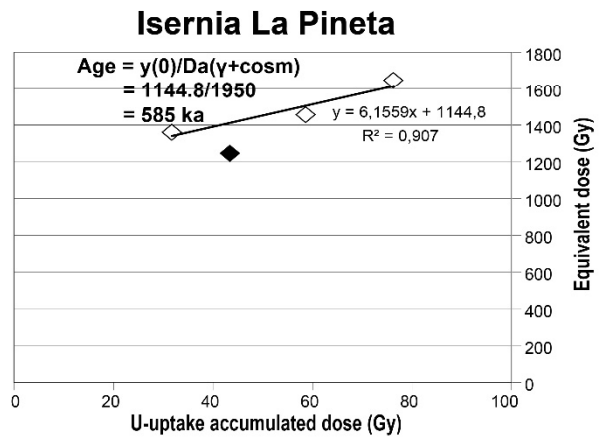
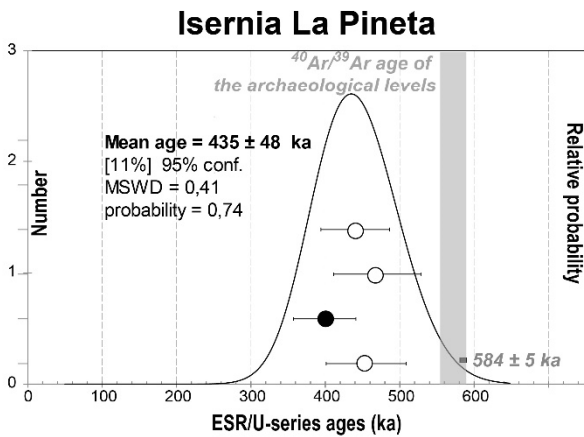
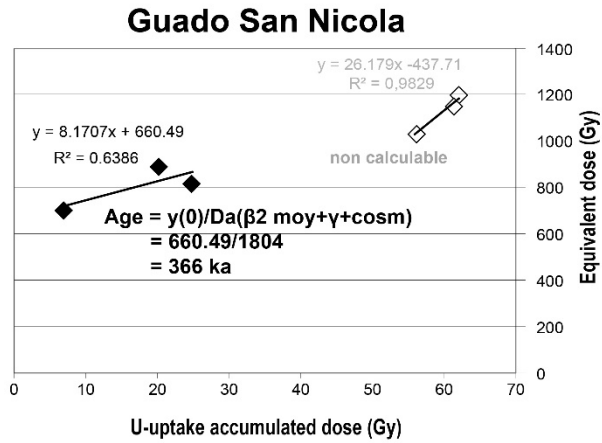
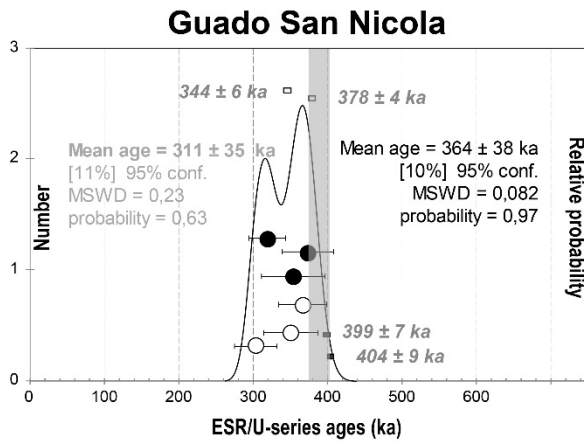
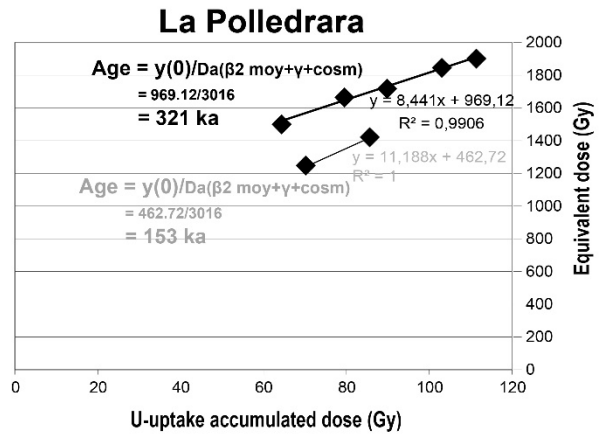
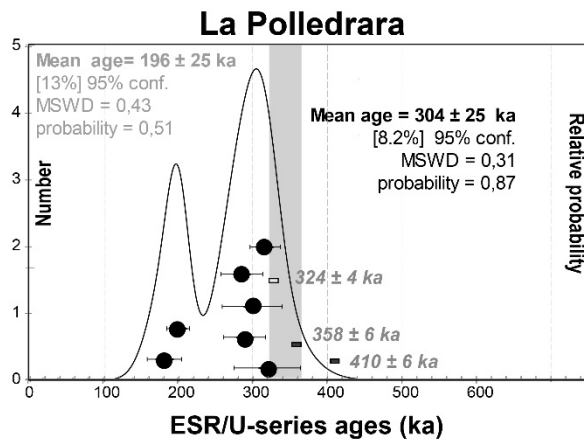
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307 **Figure 3** – Age density probability plots (left, built using Isoplot 3.0 software, Ludwig, 2003) and isochron plots (right) obtained  
308 for the different studied Middle Pleistocene sites of Central and Southern Italy with indication of the corresponding  $^{40}\text{Ar}/^{39}\text{Ar}$   
309 chronology. The U-uptake accumulated dose corresponds to the sum of the internal (enamel) and beta dose (corresponding to  
310 dentine and eventually cement) reconstructed for the considered sample all along its geological history from the modelled up-  
311 take parameters.

### 313 La Polledrara di Cecanibbio

314 According with the  $^{40}\text{Ar}/^{39}\text{Ar}$  dates, the age of the teeth from La Polledrara di Cecanibbio upper archaeological  
315 level should be close to  $324 \pm 4$  ka, since the altered pumices extracted from this level are geologically contempora-  
316 neous of the animal death and associated archaeological remains. The age density plot obtained using the ESR/U-  
317 series results is bimodal and the ages are variably (1-43%) underestimated. It was interpreted in previous paper as  
318 reflecting a poor dosimetric reconstruction of the internal dose rate evolution, in relation with the high U-content  
319 measured in the different dental tissues of the analysed teeth (Pereira et al., 2017). However, it should be noticed  
320 for part of the teeth that the isochron curves provide generally an age close to the  $^{40}\text{Ar}/^{39}\text{Ar}$  estimates, especially for  
321 teeth without cement, that seems indicate that the measured external doses are well representative to the received  
322 historical dose. The isochron age (321 ka) obtained for 5 of the 7 analysed teeth, is similar than the oldest age mode  
323 obtained for the teeth of this site and is coherent with the  $^{40}\text{Ar}/^{39}\text{Ar}$  age of this specific volcanic event. On the other  
324 hand, the isochron plot built from the apparent younger teeth population provide severely underestimated date that  
325 could reflect the record of an event posterior to the archaeological sequence deposition, such as the opening of the  
326 present-day landscape, which would lead to a drastic change of the dosimetric environmental conditions for these  
327 teeth.

328  
329 Isoletta – At Isoletta, three layers were dated by  $^{40}\text{Ar}/^{39}\text{Ar}$  (Pereira et al., 2018). The age of the volcanic layer in  
330 which ESR 1 quartz sample was collected is  $402 \pm 4$  ka, while the level corresponding to ESR 4 sample, at the top of  
331 the sequence, gives an age of  $364 \pm 9$  ka.

332 The ESR results obtained for Isoletta sediments are consistent within  $1\sigma$  with this  $^{40}\text{Ar}/^{39}\text{Ar}$  chronology, except  
333 for ESR1 sample for which severe overestimation is observed whatever the center used, due probably to both an  
334 incomplete initial bleaching of these ESR centers and a poor dose rate evaluation (Voinchet et al., 2020). Two sam-  
335 ples (ESR4 and ESR3) provided similar ages for the three centers while the age obtained using Ti-H center is underes-

336 timated for ESR2 sample. Voinchet et al. (2020) suggest that Ti–H center cannot be used to determine reliable  $D_E$   
337 estimates higher than 300–400 Gy due to signal saturation, explaining such underestimation. Weighted mean ages of  
338  $442 \pm 58$  ka (ESR2,  $2\sigma$ , Full external error, MSWD = 0.65 and  $P = 0.42$ ),  $349 \pm 26$  ka (ESR 3,  $2\sigma$ , full external uncertain-  
339 ty, MSD = 2.08 and  $P = 0.059$ ) and  $396 \pm 41$  ka (ESR 4,  $2\sigma$ , full external error, MSWD = 0.27 and  $P = 0.76$ ) were ob-  
340 tained for these 3 sediments. The ESR/US age estimate obtained for Isoletta tooth sample,  $370 \pm 31$  ka, agrees with  
341 both ESR dates from overlying and underlying sediments and similar to the  $^{40}\text{Ar}/^{39}\text{Ar}$  age of this archaeological level,  
342  $374 \pm 4$  ka.

343 Lademagne - At Lademagne,  $^{40}\text{Ar}/^{39}\text{Ar}$  chronology places the age of the archaeological levels between  $404 \pm 6$   
344 ka (Lad inf) and  $388 \pm 6$  ka (Lad sup), providing hence a *terminus ante quem* for the level dated by ESR (Pereira et al.,  
345 2018). The ESR ages calculated from Ti-Li ( $398 \pm 51$  ka) and Al ( $406 \pm 51$  ka) signals are consistent, yielding a weighted  
346 mean age of  $402 \pm 71$  ka ( $2\sigma$ , Full external error, MSWD = 0.012 and  $P = 0.91$ ).

347 Guado San Nicola - The  $^{40}\text{Ar}/^{39}\text{Ar}$  data obtained chronologically place the human occupations between  $404 \pm 9$   
348 ka and  $378 \pm 4$  ka (Pereira et al., 2016), a result considered then quite surprising in relation with the occurrence of  
349 Mode 3 archaeological artefacts in these levels. Here again, the age density plots built from ESR/U-series results is  
350 bimodal and the age underestimation, 1-20% lower than the  $^{40}\text{Ar}/^{39}\text{Ar}$  ages, was interpreted in previous paper as  
351 related with the high U-content measured in the different dental tissues of the analysed teeth (Pereira et al., 2016).  
352 The isochron age derived from the teeth without cement, around 365 ka, is however close from the older ESR/U-  
353 series age population,  $364 \pm 38$  ka, and, taking into account uncertainties associated with this result, in agreement  
354 with the  $^{40}\text{Ar}/^{39}\text{Ar}$  age of the overlying volcanic deposit (S.U. A,  $378 \pm 4$  ka). In contrast, the isochron plot built from  
355 the teeth with cement population does not allow any age determination.

356 Valle Giumentina - The  $^{40}\text{Ar}/^{39}\text{Ar}$  chronology obtained for the Valle Giumentina sequence, previously supposed  
357 contemporaneous with the Rissian Glacial stage (MIS 6-8, Demangeot and Radmilli, 1953), indeed covers two gla-  
358 cial/interglacial cycles from MIS 15 to MIS 12 (see discussions in Nicoud et al., 2016 ; Villa et al., 2016 ; Limondin-  
359 Lozouet et al., 2017). The ESR/U-series age obtained for the tooth carried out from Acheulean ALB level,  $449 \pm 29$  ka,  
360 is undistinguishable from the  $^{40}\text{Ar}/^{39}\text{Ar}$  age of LABM level located just below,  $455 \pm 3$  ka.

361 Isernia La Pineta - A primary fallout tephra layer located just below the main archaeological level t3a was dated  
362 by  $^{40}\text{Ar}/^{39}\text{Ar}$  of  $584 \pm 5$  ka (Peretto et al., 2015), furnishing a maximum age for this level. The age underestimation of  
363 ESR/U-series data is here more severe, around 25% (mean age) lower than this date. The hypothesis of a change of  
364 the external dose in relation with late uptake history in the palaeontological remains was proposed previously to

365 explain this age underestimation (Falguères et al., 2007; Shao et al., 2011), but it should be underlined that the  
366 isochron age obtained for these teeth (585 ka) is in perfect agreement with the  $^{40}\text{Ar}/^{39}\text{Ar}$  date .

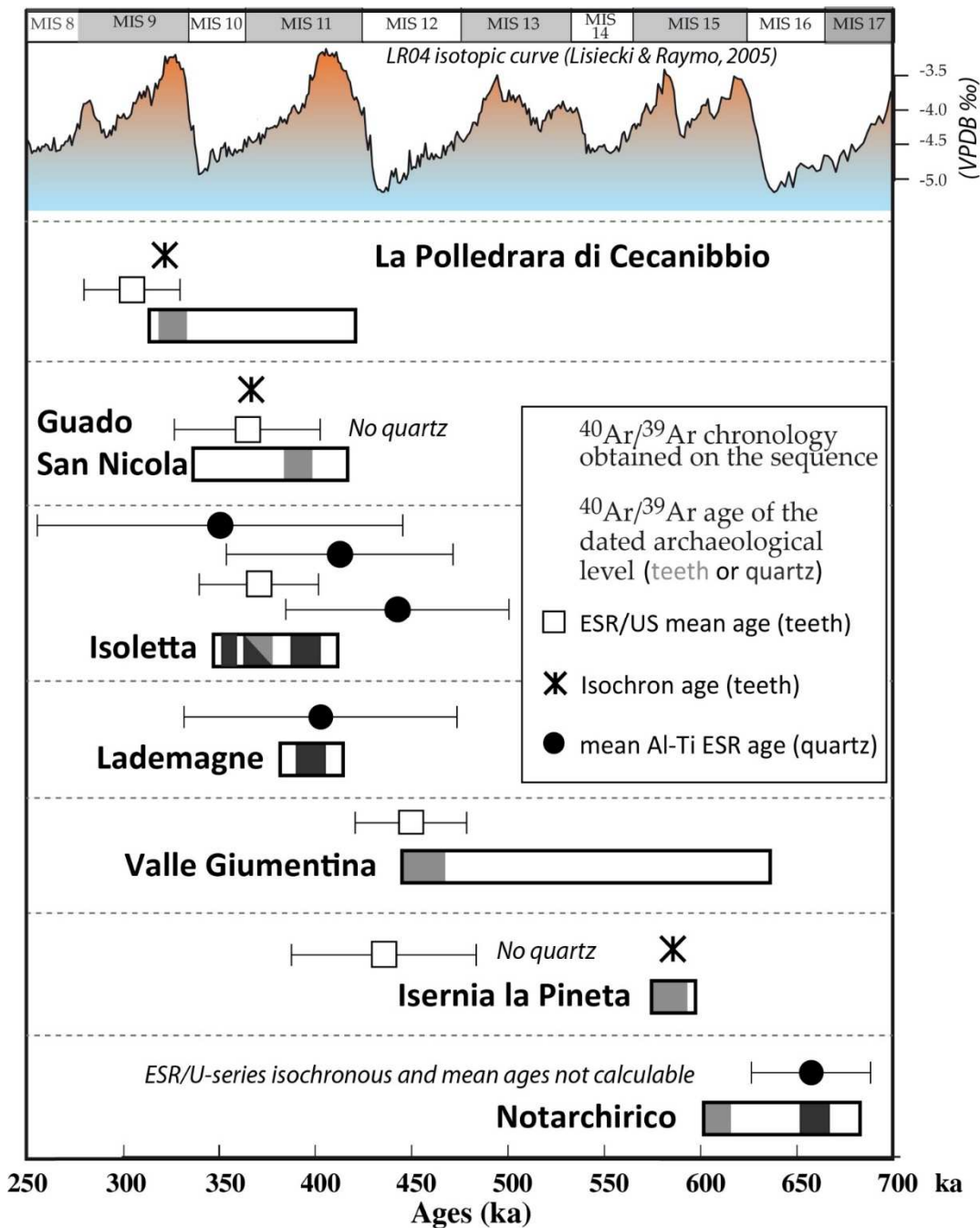
367 Notarchirico - The situation is different at Notarchirico. For the teeth recovered from the palaeoanthropological level  
368 supra  $\alpha$ , both the ESR/US ages (ranging from  $208 \pm 32$  ka to  $546 \pm 64$  ka), are so scattered that it precludes any mean  
369 age calculation. Isochron age estimate (134 ka) is much younger than the  $^{40}\text{Ar}/^{39}\text{Ar}$  age time range estimate, dated  
370 between  $612 \pm 4$  ka and  $659 \pm 9$  ka (Pereira et al., 2015). These results indicate that the teeth have either recently  
371 experienced an event avoiding ESR/U-series dating possibilities or that the palaeodosimetric reconstruction cannot  
372 be realized from present-day data. On the other hand, the ages derived from ESR analyses of Notarchirico sediments  
373 are quite coherent despite the fact than only the AI center was usable. The obtained ESR age ( $657 \pm 31$  ka) agrees at  
374  $1\sigma$  with the  $^{40}\text{Ar}/^{39}\text{Ar}$  age estimate for this level, constrained between 658 and 612 ka.

## 375

### 376 Discussion

377 The main purpose of the present paper was to discuss of the chronologies established by ESR and ESR/U-series  
378 methods on sediments and teeth sampled for various Middle Pleistocene sites of Central and Southern Italy, display-  
379 ing into their stratigraphic sequences also volcanic materials dated by  $^{40}\text{Ar}/^{39}\text{Ar}$  on single grains. The figure 4 summa-  
380 rizes the whole set of data obtained on the studied sites.

381  $^{40}\text{Ar}/^{39}\text{Ar}$  method was then used as reference method to provide independent chronological data on the studied  
382 archaeological sites.  $^{40}\text{Ar}/^{39}\text{Ar}$  results permit hence to place with confidence the studied sites in a single precise and  
383 coherent chronological framework. The site succession from the youngest to the older is the following: La Polledrara  
384 di Cekanibbio < Guado San Nicola  $\approx$  Isoletta  $\approx$  Lademagne < Valle Giumentina < Isernia la Pineta < Notarchirico. This  
385 succession is slightly different than the biochronological succession previously proposed for these sites and in which  
386 Isernia was supposed equivalent in age or slightly older than Notarchirico (Sardella et al., 2006; Palombo and Sar-  
387 della, 2007; Marra et al., 2014). From the archaeological point of view, this chronology ranges the deposition of the  
388 long stratigraphic sequence of Valle Giumentina from MIS 15 to MIS 12, while it was considered as coeval with the  
389 Rissian Glacial stage (MIS 6-8, Demangeot and Radmilli, 1953). Moreover  $^{40}\text{Ar}/^{39}\text{Ar}$  data indicate that the archaeolog-  
390 ical levels of Guado San Nicola, in which Mode 3 lithic industry was recovered, are contemporaneous with the MIS  
391 11 and 10 periods, placing the site among the oldest occurrences of this type of lithic technology (Peretto et al.,  
392 2014, 2016).



394  
395 **Figure 4** – ESR and ESR/U-series ages obtained for the different studied Middle Pleistocene sites of Central and Southern Italy  
396 with indication of the corresponding  $^{40}\text{Ar}/^{39}\text{Ar}$  chronology.  
397

398 The ESR ages obtained on the fluvial quartz grains are globally in agreement with these  $^{40}\text{Ar}/^{39}\text{Ar}$  chronologies on  
399 the studied sites. The results derived from Al and Ti-Li ESR centers are globally similar for the main part of the ana-  
400 lysed samples, while equivalent doses and ages obtained using Ti-H center are systematically underestimated, per-  
401 haps in relation with a faster saturation of these centers. Voinchet et al. (2020) recommend hence that Ti-H centers  
402 should not be used to date samples with equivalent doses greater than ca 300 Gy. These authors underline also that,  
403 if Al-center use can provide overestimated age estimate in case of incomplete initial bleaching, Ti-Li center meas-  
404 urement can be complicated for samples displaying poor signal-to-noise resolution (Voinchet et al., 2020).

405 On the other hand, the ages obtained from ESR/U-series results for the different studied sites in which several  
406 teeth were analysed appear quite systematically underestimated. This underestimation is slight for the main part of  
407 the teeth recovered from the two younger sites (< 450 ka, La Polledrara and Guado San Nicola), but more drastic for  
408 the whole set of teeth carried out from the two older ones (Isernia and Notarchirico) despite very different U-uptake  
409 histories (systematic late uptake at Isernia and early uptake or leaching at Notarchirico). The use of the isochron ap-  
410 proach attests however of the quite good reliability of the palaeodosimetric reconstruction for these different sites  
411 and the observed age underestimation could therefore be related to an underestimation of the equivalent dose. In  
412 such context, it should also be underlined that the results obtained for the single teeth from Isoletta and Valle Giu-  
413 mentina are in agreement with the  $^{40}\text{Ar}/^{39}\text{Ar}$  ages of the dated levels.

414 An explanation advanced to explain such age underestimation could be linked to the high U contents recorded in  
415 some of the analysed teeth enamels and associated high internal doses reconstruction. The possibility of an inverse  
416 correlation between the uranium concentration and ESR signal intensities in enamel displaying high U content was  
417 indeed postulated, considering that the crystal lattice of the enamel could locally be completely destroyed by alpha  
418 radiations emitted by uranium decay and so that the measured ESR signal corresponds only to low uranium area  
419 (Bahain et al., 1992). This assumption has however never been demonstrated and the results of the present study  
420 indicate that, if such effect could exist, it cannot explain the whole set of data obtained in the present study. Indeed,  
421 the association of high U content in enamel, low  $D_E$  value and leaching evidence in the different dental tissues (corre-  
422 sponding to high  $^{230}\text{Th}/^{234}\text{U}$  ratios) leads to more or less severe age underestimation.

423 Another possible explanation of such age underestimation is linked to the fact that ESR signal of enamel is indeed  
424 resulting of several  $\text{CO}_2^-$  radicals having different stabilities (Joannes Boyau and Grün, 2009). These radicals are indis-  
425 tinguishable from each other when working on powders as during the present study, but can be distinguished by  
426 analyses of enamel fragments (Joannes Boyau et al., 2010; see also discussion in Joannes Boyau, 2013 and references  
427 therein). Some of these radicals are unstable considering geological time range and are considered as responsible for  
428 an underestimation of up to 30% of  $D_E$  values (Joannes Boyau and Grün, 2011). The equivalent doses derived from  
429 T1-B2 peak-to-peak measurements could therefore be affected by this lower stability of these unstable  $\text{CO}_2^-$  radicals  
430 to a greater or lesser degree. It could be also dependant on the dose rate intensity, complicating for the moment the  
431 taking into account of this parameter. Systematic comparisons between powder and enamel fragment ESR studies  
432 could allow a better estimation of the impact of this poor stability but it necessitates specific ESR equipment and is  
433 very time consuming. As the isochron plots provide results in agreement for part of the studied teeth sets, it seems

434 that an reliable age estimate can be obtained by this way if the dose rate has not changed over the geological history  
435 of the site in question. Here also, such assumption will have to be confirmed by subsequent studies. It could be also  
436 interesting to rely powder and fragment studies as well to try to characterize the relative percentages of stable and  
437 unstable  $\text{CO}_2^-$  radicals in ESR powder spectra to correct for each sample a possible  $D_E$  underestimation. Unfortunate-  
438 ly, such studies are still scarce and the various methods proposed to such correction were not sufficiently checked  
439 with independent age control and systematic corrections seem thus to be premature (see details in Joannes Boyau,  
440 2013).

441 Despite these questions, the results obtained in the present study are encouraging and permitted to clarify the  
442 chronology of the early Middle Pleistocene prehistoric sites in Italy, ranging approximately between 300 and 700 ka.  
443 We are currently trying to apply to other site the same multi-method approach associating ESR and ESR/U-series  
444 with independent radiisotopic methods ( $^{40}\text{Ar}/^{39}\text{Ar}$ , but also U-series,  $^{14}\text{C}$  and other cosmonuclides) s, either more  
445 recent (late Middle Pleistocene and Upper Pleistocene) or older (Lower Pleistocene).

## 447 **Conclusions**

448 The ESR and combined ESR/U-series ages determined from analyses realized respectively on quartz and teeth  
449 carried out from several Middle Pleistocene Italian sites were systematically compared with  $^{40}\text{Ar}/^{39}\text{Ar}$  ages obtained  
450 on K-feldspars minerals derived from volcanic or fluvial deposits. The ESR/U-series ages seem quite systematically  
451 underestimated (from 1 to 66 % depending on the age of the sites) but the isochron age estimates obtained on three  
452 sites are in agreement with the corresponding  $^{40}\text{Ar}/^{39}\text{Ar}$  chronologies and indicate that the palaeodosimetric recon-  
453 struction is generally correct for the dated levels, with the notable exception of Notarchirico site. For the quartz ESR  
454 dating, the multi-center approach was used when possible. The ages determined both from Ti-Li and Al signals are  
455 generally in agreement with each other and consistent with the  $^{40}\text{Ar}/^{39}\text{Ar}$  ages, while the Ti-H age estimates appear  
456 as underestimated for samples associated to  $D_E$  values higher than 300 Gy, suggesting a saturation of the Ti-H traps  
457 in these samples (Voinchet et al., 2020).

458 These results are overall encouraging and pinpoints the necessity when possible to cross-check the methods as  
459 well to analyse several samples by dated level. At Notarchirico, both ESR/U-series and isochron ages are much  
460 younger than the expected age, by contrast with the ESR age which is in good agreement with the  $^{40}\text{Ar}/^{39}\text{Ar}$  ages.  
461 Lastly, at Isoletta, both ESR/U-series, ESR and  $^{40}\text{Ar}/^{39}\text{Ar}$  are in good agreement. These data indicate also that the

462 ESR/U-series and ESR chronologies established on Western European Palaeolithic sites when other independent  
463 methods are not available reflect rather well the geological ages of the dated Middle Pleistocene levels.

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671

672

673 **Figure caption**

674

675 **Figure 1** - Location of the studied Middle Pleistocene sites in Central and Southern Italy.

676

677 **Figure 2** – Simplified stratigraphic logs of the studied Middle Pleistocene sites with indication of the levels dated by  
678  $^{40}\text{Ar}/^{39}\text{Ar}$  and positions of the levels sampled for ESR and ESR/U-series studies.  $^{40}\text{Ar}/^{39}\text{Ar}$  ages were obtained directly  
679 on tephra levels (primary deposits) or on potassic feldspar grains extracted from fluvial sediments (secondary  
680 deposits), the youngest population corresponding then to a maximum age for the dated level.

681

682 **Figure 3** – Age density probability plots (left, built using Isoplot 3.0 software, Ludwig, 2003) and isochron plots (right)  
683 obtained for the different studied Middle Pleistocene sites of Central and Southern Italy with indication of the corre-  
684 sponding  $^{40}\text{Ar}/^{39}\text{Ar}$  chronology. The U-uptake accumulated dose corresponds to the sum of the internal (enamel) and  
685 beta dose (corresponding to dentine and eventually cement) reconstructed for the considered sample all along its  
686 geological history from the modelled uptake parameters.

687

688 **Figure 4** – ESR and ESR/U-series ages obtained for the different studied Middle Pleistocene sites of Central and  
689 Southern Italy with indication of the corresponding  $^{40}\text{Ar}/^{39}\text{Ar}$  chronology.

690

691 **Table caption**

692

693 **Table 1** – List and nature of the analysed teeth from several Middle Pleistocene Italian sites

694

695 **Table 2** –  $^{40}\text{Ar}/^{39}\text{Ar}$  ages obtained from volcanic minerals recovered from several Middle Pleistocene Italian sites.  
696 The indicated  $^{40}\text{Ar}/^{39}\text{Ar}$  age corresponds to the age of the youngest homogeneous potassic feldspars population.  
697 These ages were re-calculated in the present contribution according to the monitor flux standard ACs-2 at 1.1891 Ma  
698 (Niespolo et al., 2017) and the K total decay constant of Renne et al. (2011) from the original data published in the  
699 various referred papers. A homogeneous population is considered relevant when the weighted mean of these crys-



700 tals has the following statistical characteristics: MSWD < 1.5, Probability fit  $\geq 0.1$ . The weighted average ages were  
701 calculated using IsoPlot 3.0 (Ludwig, 2003) and given at 95% ( $2\sigma$ ) of confidence.

702  
703 **Table 3** – ESR data and ages for sediments recovered from several Middle Pleistocene Italian sites.

704 Equivalent doses ( $D_E$ ) were derived from the obtained intensity growth curves using an exponential + linear function  
705 for Al and Ti-Li centers (Duval et al., 2009; Voinchet et al., 2013) and by a single saturating exponential function for  
706 Ti-H from the six first points of the growth curves (Voinchet et al., 2020) with Microcal OriginPro 8 software, both  
707 with  $1/I^2$  weighting (according with Yokoyama et al., 1985). Age calculations were performed using the following  
708 parameters: dose-rate conversions factors from Guérin et al. (2011); a k-value of  $0.15 \pm 0.1$  (Laurent et al., 1998);  
709 alpha and beta attenuations from Brennan (2003) and Brennan et al. (1991); water attenuation formulae from Grün  
710 (1994); cosmic dose rate estimated from the Prescott and Hutton's equations (1994). The internal dose rate was  
711 considered as negligible because of the low contents of radionuclides usually found in quartz grains (Murray and  
712 Roberts 1997; Vandenbergue et al. 2008). ESR age estimates are given with one sigma error range.

713  
714 **Table 4** – U-series and ESR data for the analysed teeth recovered from several Middle Pleistocene Italian sites.

715 U-series analyses were performed on each dental tissue in order to determine the U-uptake parameters necessary to  
716 the dose rate contributions and age calculations (see details in Shao et al., 2015b), either by alpha spectrometry  
717 (MNHN, Paris), or by Neptune Multi-Collector Inductively Coupled Plasma Mass Spectrometer (MC-ICPMS) (Nanjing  
718 Normal University, China).

719  
720 **Table 5** – ESR/U-series data and ages for the analysed teeth recovered from several Middle Pleistocene Italian sites.

721 The equivalent doses  $D_E$  were extrapolated from the obtained dose-response data sets using either The growth  
722 curves were built for each sample using different fitting functions (single saturating exponential, SSE, according with  
723 Apers et al., 1981; double saturating exponential function (DSE), according with Duval et al. (2009), or exponential  
724 plus linear function (E+L), according with Shao et al. (2015a). The equivalent dose corresponding to the function that  
725 best describes the experimental data and gives the best statistics is then used to calculate age.

726 The radioelement contents of sediment samples associated to each tooth were determined by in laboratory high  
727 resolution low background gamma spectrometry and *in situ* gamma measurements with TL  $Al_2O_3$  dosimeters at La

728 Polledrara and Inspector 1000 Canberra gamma spectrometer on the other sites. A sediment water content value of  
729  $15 \pm 5 \%$  was therefore assumed for the age calculations. The cosmic dose rate was estimated using the formula of  
730 Prescott and Hutton (1994). Were also used the following parameters: a k-value ( $\alpha$  efficiency) of  $0.13 \pm 0.02$  (Grün  
731 and Katzenberger-Apel, 1994); water content of 0 wt% in the enamel and 7 wt% in the dentine and cementum; con-  
732 version contents-doses factors from Guérin et al.; (2011). For each dental tissue, Rn loss was estimated from both  
733 gamma and alpha/ICP measurements (Bahain et al., 1992). The beta dose contributions were corrected from the  
734 enamel part destroyed on each side of the enamel layer during the preparation process (according to Brennan et al.,  
735 1997).

736 ESR/U-series ages, different dose-rate contributions and U-uptake parameters were then calculated using either  
737 US model (Grün et al., 1988) or AU model (Shao et al., 2012). The “USESR”, “AUESR” and “combined ESR” computer  
738 programs were used for the age calculation in which the age uncertainty ( $1\sigma$ ) is calculated with Monte Carlo ap-  
739 proach (Shao et al., 2014).

740 Isochron age estimates were newly determined in the present work except for La Polledrara samples (Pereira et al., 2017).

741