

# Global patterns of biomass burning during the last glacial period

Anne-Laure Daniau

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1	Global patterns of biomass burning during the last glacial period
2	Anne-Laure DANIAU
3	Ms for PAGES newsletter
4	School of Geographical Sciences, University of Bristol, Bristol, UK
5	Email: al.daniau@bristol.ac.uk
6	
7	Sedimentary charcoal records covering part of the last glacial period were
8	synthesized to analyse the response of global biomass burning to rapid climate
9	changes during the so-called Dansgaard-Oeschger "cycles".
10	
11	The Global Palaeofire Working Group
12	(http://www.bridge.bris.ac.uk/projects/QUEST_IGBP_Global_Palaeo
13	fire_WG) has updated a database over 700 individual sedimentary charcoal records
14	worldwide (GCD_V2: Daniau et al., in prep). This database provides a powerful tool
15	for studying changes of biomass burning at global and regional scale (Power et al.,
16	2008; Power et al., in press). A focused study of fire records covering the last glacial
17	period (73.5-14.7 ka) allows examination of the response of biomass burning to the
18	rapid climate changes (within 10-200 years) of large magnitude that occur during
19	Dansgaard-Oeschger (D-O) cycle. D-O cycles are characterized in Greenland by a
20	marked warming followed by a cooling. D-O warming event refers here to the initial
21	rapid warming (Sánchez-Goñi and Harrison, in press). The Greenland Interstadial
22	(GI) corresponds to the D-O warming event followed by a first slow phase of cooling.
23	GI end by a precipitous cooling (refers here as D-O cooling event). Greenland Stadial
24	(GS) correspond to the final slow phase of cooling.
25	
26	
27	Methodology
28	Sixty-seven sites (11 marine and 56 terrestrial, Figure 1) which have records for some
29	part of the last glacial period were extracted from the database and used to analyse
30	changes in global biomass burning (Daniau et al., in press). Those records present a
31	broad range of quantification methods and units. For the majority, charcoal counts are

32 expressed as concentration/influx of number of particles or area measurements made

33 on pollen slide or using image analysis. Few records were studied using chemical 34 assay such as total organic fraction and black carbon concentration, or expressed as 35 Charcoal/Pollen ratio and percentage dry weight. The data were therefore 36 standardized to facilitate comparisons between sites and through time. Original 37 charcoal data (abundance or concentration) were transformed and rescaled to obtain a 38 common mean and variance for all sites, and detrended to remove orbital time-scale 39 variability. A Superposed- Epoch Analysis (SEA) was performed on the detrended 40 charcoal data to examine the pattern of fire around the time of key events, aligning the 41 charcoal deviations on the ages of the D-O warming and D-O cooling events (Figure 42 2). When the SEA summary curves fall outside a confidence band determined by 43 Monte Carlo methods, a significant or systematic response of the variable to some 44 externally determined event is inferred.

45

#### 46 Fire and temperature

47 Figure 2 illustrates SEAs for the NGRIP oxygen-isotope data and for the global 48 composite-curve of charcoal data to illustrate the nature of the abrupt climate changes 49 and the response of global biomass burning to the abrupt warming at the beginning of 50 the Greenland Interstadials (GI) and the most rapid interval of cooling that delimits 51 the beginning of Greenland Stadials (GS). A general increase in charcoal levels 52 during warming (Fig. 2A and B) and a decrease in charcoal levels during cooling 53 (Fig. 2C and D) suggest a strong correlation between biomass burning and 54 temperature. Biomass burning increases rapidly during D-O warming events, and 55 reaches a peak at nearly the same time as temperature. In contrast, biomass burning 56 decreases significantly at the onset of the GS and then returns to "background" levels 57 while temperatures remain low (Figs. 2C and D). The response of biomass burning to 58 the warming events is not the simple inverse of the response to cooling, suggesting a 59 nonlinear relationship between biomass burning and climate. The association between 60 rapid warming and increases in biomass burning seen during the glacial period has 61 also been observed during the last deglaciation at the termination of the Younger 62 Dryas chronozone (Marlon et al., 2009).

63

#### 64 Fire regime and fuel

65 Changes in charcoal concentration have been described from two marine records off

66 the European margin and interpreted as changes in biomass burning caused by variations in fuel availability during D-O "cycles" (Daniau et al., 2007; Daniau et al., 67 68 2009). Recent syntheses of the glacial vegetation history of Europe, Japan and 69 adjacent parts of the Asian mainland, North America and the tropical regions of South 70 America and Africa (Fletcher et al., in press; Takahara et al., in press; Jiménez-71 Moreno et al., in press; Hessler et al., in press) suggest that these rapid climate 72 changes had an impact on vegetation globally. Warmings are generally accompanied 73 by shifts from open vegetation to forest or increases in tree abundance. The similarity 74 of variations in global levels of fire and regional vegetation patterns suggest that 75 climate-driven changes in vegetation and biomass productivity (fuel) caused changes 76 in global biomass burning during the last glacial period. During the D-O warming 77 events and GI, increased productivity led to increased fuel availability and promoted 78 fire. Conversely, decreased fire in response to rapid cooling was a consequence of 79 reduced fuel loads from lower productivity. Increase in biomass burning to levels 80 similar to pre-cooling fire regimes during GS may be caused by temperature-driven 81 forest die-back, which would create conditions for increased fire.

82

83

#### 84 Conclusion and outlook

85 This effort to synthesise charcoal records from the last glacial period allows us to 86 explore changes in global fire regimes during a period of high millennial-scale 87 climate variability. Biomass burning at the global scale apparently responded rapidly 88 to variations in temperature, as recorded in the Greenland ice core. Increases in 89 biomass burning were synchronous with warming whereas decreases in burning 90 followed by a return of fire to previous levels accompanied cooling. Changes in 91 vegetation productivity and fuel availability may explain the inferred variations in 92 biomass burning during D-O cycles. 93 Additional fire records however are needed to derive statistically robust results at a

regional scale and better examine fire and climate interactions. Global biomass
burning responds to temperature variation, but we also seek to explore the control of
seasonality of temperature and precipitation on fire, playing a role in shaping the
vegetation (type of biomes, abundance of fuel) but also on the flammability of fuel.
We intend to extend the database by adding worldwide long and high-resolution
marine and terrestrial charcoal records, in particular from poorly documented North

- 100 America, Africa and a large part of Eurasia. There is a great potential for the use of
- 101 long marine records to document regional fire where terrestrial records are not
- available. Characterisation of the role of temperature and precipitation on fire would
- 103 benefit from the close comparison between fire and well-documented high resolution
- 104 changes in vegetation during D-O rapid climate changes. This could be done in
- 105 particular in Europe where these rapid changes lead to an alternation between steppe-
- 106 like environments and different forest biomes.
- 107

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- 111 improved the manuscript.





## 115 **Figure 1:**

116 Location of marine and terrestrial sites with charcoal records covering all or part of

117 the last glacial period. Figure modified from Daniau et al., in press.



120 Figure 2:

Superposed Epoch Analysis (SEA) composites illustrating changes in biomass 121 122 burning associated with D-O warming event (A and B) and D-O cooling event (C and 123 D) with x-axis corresponding to time (before event on the left (negative values) and after event on the right (positive values)). The NGRIP  $\delta^{18}$ O SEA composite record for 124 125 the D-O cycles is shown for comparison. The NGRIP oxygen-isotope data (Wolff et 126 al., in press) were summarized using 20-year-wide bins (the resolution of the data), 127 and also using the same 80-year-wide bins used for charcoal data, for comparison 128 with the SEA of the charcoal data. The confidence intervals are given by Monte-Carlo simulation using the timing of the 20 D-O warming events and 19 D-O cooling 129 130 events. Figure modified from Daniau et al., in press. 131

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