

## Recent hydrological variability and flood events in Moroccan Middle-Atlas mountains: micro-scale investigation of lacustrine sediments

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# Methods

steep slopes and a cedar

forest.

## **1- Geochemistry**

(XRF) and *mineralogy* 

- X-ray fluorescence
- Itrax Core Scanner
- Molybdenum tube
- Resolution: 500 µm
- 15 s time exposure
- 30 kV and 25 mA
- X-ray diffraction XRD (Fig. 7)

## 2- Microsedimentology

- Thin sections (acetone exchange technique)
- Flatbed transparency scanner
- Optical microscopy
- SEM and energy dispersive spectroscopy (15 kV and 90 min
- acquisition time)



# Results Figure 5: Geochemistry (XRF intensity) VS lake level changes



# Recent hydrological variability and flood events in Moroccan Middle-Atlas mountains : micro-scale investigation of lacustrine sediments Jouve Guillaume (a,\*), Vidal L. (a), Adallal R. (a,b), Bard E. (a), Benkaddour A. (b), Simonneau A. (f), Sonzogni C. (a), Sylvestre F. (a), Tachikawa K. (a), Rostek F. (a), Viry E. (a) Viry E. (b), Simonneau A. (c) Viry E. (c) Viry E.

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## Scientific context and objectives

Since the 1990s, the Mediterranean basin undergoes an increase in extreme precipitation events and droughts likely to intensify in the XXI century (IPCC, 2013). Regional climate models indicate a strengthening of flood episodes at the end of the XXI century in Morocco (Tramblay et al, 2012). To understand the recent hydrological variability in North Africa, our study focuses on geochemical and microsedimentological analysis of a short sedimentary sequence from Azigza lake (Fig. 1; 2). This endoreic lake is located in the Middle Atlas karst system.

Limited data on past lake level changes during the last decades are provided by Gayral & Panouse (1954), Flower et al. (1989) and Flower & Foster (1992). To refine our knowledge of past hydrological changes in this region, the first objective is to reconstruct high and low lake levels throughout the last hundred years. The second objective is to detect and count flood events.

**Figure 6: Microsedimentology** 137Cs peak Results **Figure 4: Laser diffraction grain size b.** Facies 2: homogeneous sediments at depth 27 cm composed of autochthonous calcite and quartz **a. Facies 1**: sediment rich in wood and **Clayey silts** grains calcitic shells, with several erosive structures. Homogenous 300 µm grain size of 0.5-1 14-15 28-2 42-4 detrital particles (=limited changes in the sedimentation rate) Angulous => Age at the base silty-sized (64 cm depth) quartz 35 ~ year 1900 20 µm Particle diameter (µm) | Fine silts Calcitic shell and clave epithelial cells of Redox Κ Si (1 cm) (SEM-EDS map) 3000 200 250 2000 350 300 400 Overall sedimentological description: Unconsolidated, light brown to 2004 black, very thin bedded to laminated, mixing of autochthonous calcite and detrital clayey silts with few thin laminations of ostracod shells (SEM-EDS map) Ca Si 1994 **Discussion and perspectives** Results 1984 **Figure 7: Mineralogy (XRD)** Lake level changes during the past hundred years are recorded in the geochemistry and the microfacies of the 1974 13-1\_0-1 cm sedimentary sequence: - High lake level facies (Fig. 6a, Facies 1) is deposited when lake shorelines are closer to the vegetation line and steep slopes (Fig. 1). This facies is characterized by light brown sediments, less orga-1964 Quartz nic/more minerogenic (Fig. 5, 25 cm depth), with several erosive structures containing wood fragments and cal-Calcite citic shells of ostracods (Fig. 6a, Facies 1). Its geochemical signature is defined by higher Si, K, Fe and Ti that Pyrite 1954 indicates more detrital input. Since (1) Si covary with K (Fig. 5, PCA), and since (2) sands are poorly present Dolomite in the sediment (Fig. 4), we interpret the Si signal as indicator of the finest detrital fraction (clays and fine silts) Kaolinite 1944 Gypsum brought by superficial runoff (SEM-EDS images of silty quartz are available in Figure 6b, Facies 2). Flood Illite events are marked by Mn peaks, which is interpreted as manganese oxides precipitations under well-oxyge-1934 nated deep water after flood events. Facies 1 is deposited during periods of higher precipitations (Fig. 5). - Low lake level facies (Fig. 6b, Facies 2) is deposited when shorelines are close to smoother bank slopes, 1924 (Fig. 1). This facies is represented by homogeneous sediments composed of autochthonous calcite and quartz Position  $(2\theta)$  (Cobalt) grains, with substantial decreases in the XRF detrital proxies (Si, K, Ti and Fe, Fig. 5). Autochthonous calcite 1914 is not revealed in the CaCO3% and Ca-XRF signal because calcitic shells of ostracods are highly presents in Major peak calculations: Cal-Facies 1. Facies 2 is deposited during periods of lower precipitations (Fig. 5). cite: 45%, Quartz: 38%, Dolo-1904 These results demonstrate the high potential of Azigza lake to help understanding the past hydrological variabi-200 250 300 350 400 450 500 300 500 700 mite: 10%, Aragonite: 3%, Kaolity of the Middle-Atlas. Indeed, its water level and hydrosedimentary system are sensitive to rapid (floods), as Si Precipitation (mm) linite: 2%, Pyrite: 2%, Illite: 1%, well as «long»-term (dry and wet periods during several decades) changes in the precipitation regime. The two meters-long sedimentary sequence, recently retrieved from the deeper basin, would allow the reconstruc-5 years mean Gypse: 0.5% Low lake level High lake level (re-analyzed data at -5.5° W; 33.0° N, CRU) tion of the hydrological variability of Azigza lake for the past few hundred years.









