

InterCarb: A community effort to improve inter-laboratory standardization of the carbonate clumped isotope thermometer using carbonate standards

Stefano M Bernasconi, Mathieu Daëron, Kristin D Bergmann, Magali Bonifacie, Anna Nele Meckler, Hagit Affek, Noah Anderson, David Bajnai, Eugeni Barkan, Emily Beverly, et al.

▶ To cite this version:

Stefano M Bernasconi, Mathieu Daëron, Kristin D Bergmann, Magali Bonifacie, Anna Nele Meckler, et al.. InterCarb: A community effort to improve inter-laboratory standardization of the carbonate clumped isotope thermometer using carbonate standards. 2021. hal-03094122

HAL Id: hal-03094122 https://hal.science/hal-03094122

Preprint submitted on 4 Jan 2021

HAL is a multi-disciplinary open access archive for the deposit and dissemination of scientific research documents, whether they are published or not. The documents may come from teaching and research institutions in France or abroad, or from public or private research centers.

L'archive ouverte pluridisciplinaire **HAL**, est destinée au dépôt et à la diffusion de documents scientifiques de niveau recherche, publiés ou non, émanant des établissements d'enseignement et de recherche français ou étrangers, des laboratoires publics ou privés.

1 InterCarb: A community effort to improve inter-laboratory

2 standardization of the carbonate clumped isotope

3 thermometer using carbonate standards

- 5 S.M. Bernasconi¹, M. Daëron², K.D. Bergmann³, M. Bonifacie⁴, A.N. Meckler⁵, H.P. Affek⁶, N.
- 6 Anderson³, D. Bajnai⁷, E. Barkan⁶, E. Beverly^{8†}, D. Blamart², L. Burgener⁹, D. Calmels^{4†}, C.
- 7 Chaduteau⁴, M. Clog¹⁰, B. Davidheiser-Kroll¹¹, A. Davies^{12††}, F. Dux^{13†}, J. Eiler¹⁴, B. Elliott¹⁵,
- 8 A. C. Fetrow¹¹, J. Fiebig⁷, S. Goldberg³, M. Hermoso^{4,23}, K.W. Huntington¹⁶, E. Hyland⁹, M.
- 9 Ingalls^{14†}, M. Jaggi¹, C.M. John¹², A. B. Jost³, S. Katz⁸, J. Kelson⁸, T. Kluge^{12†}, I.J. Kocken¹⁷,
- 10 A. Laskar 18 , T.J. Leutert 5† , D. Liang 18 , J. Lucarelli 15 , T. J. Mackey 3† , X. Mangenot 4,14 , N.
- 11 Meinicke⁵, S.E. Modestou⁵, I.A. Müller¹⁷, S. Murray¹⁹, A. Neary⁸, N. Packard⁸, B.H. Passey⁸, E.
- Pelletier⁸, S. Petersen⁸, A. Piasecki^{5††}, A. Schauer¹⁶, K.E. Snell¹¹, P.K. Swart²⁰, A. Tripati¹⁵, D.
- 13 Upadhyay¹⁵, T. Vennemann²¹, I. Winkelstern^{8††}, D. Yarian⁸, N. Yoshida²², N. Zhang²², M.
- 14 Ziegler¹⁷

4

- ¹Geological Institute, ETH Zürich, Sonneggstrasse 5, 8092 Zürich, Switzerland
- ²Laboratoire des Sciences du Climat et de l'Environnement, LSCE/IPSL, CEA-CNRS-UVSQ,
- 17 Université Paris-Saclay, Orme des Merisiers, F-91191 Gif-sur- Yvette, France
- ³Department of Earth, Atmospheric and Planetary Sciences, Massachusetts Institute of
 Technology, Cambridge, MA 02139, USA
- 20 ^{3†}now at Earth and Planetary Sciences, University of New Mexico, Albuquerque, New Mexico, USA
- ⁴Université de Paris, Institut de Physique du Globe de Paris, CNRS, F-75005 Paris, France
- 23 ^{4†} now at Geosciences Paris Sud (GEOPS), Université Paris-Saclay, CNRS, 91405 Orsay, France
- ⁵Bjerknes Centre for Climate Research and Department of Earth Science, University of Bergen,
- 25 Bergen, Norway
- 26 ^{5†}now at Max Planck Institute for Chemistry, Mainz, Germany; thomas.leutert@mpic.de
- 27 5^{††}now at Department of Earth Sciences, Dartmouth College, Hanover, NH, USA
- 28 ⁶Institute of Earth Sciences, Hebrew University of Jerusalem, Jerusalem, Israel
- ⁷Institute of Geosciences, Goethe University Frankfurt, Frankfurt am Main, Germany
- 30 ⁸Department of Earth and Environmental Sciences, University of Michigan, Ann Arbor,
- 31 Michigan, USA
- 32 8†now at Department of Earth and Atmospheric Sciences, University of Houston, Houston,
- 33 Texas, USA
- 34 8††now at Geology Department, Grand Valley State University, Allendale, Michigan, USA
- ⁹Marine, Earth and Atmospheric Sciences, North Carolina State University, Raleigh, USA
- 36 ¹⁰Scottish Universities Environmental Research Centre (SUERC), Glasgow, Scotland

¹¹University of Colorado Boulder, Boulder, Colorado, USA 37 ¹²Imperial College, London, UK 38 ^{12†}now at Karlsruher Institut für Technologie KIT 39 12††now at Stockholm University 40 ¹³School of Geography, University of Melbourne, Australia 41 ^{13†}now at School of Earth and Life Sciences, University of Wollongong, Australia 42 ¹⁴Geological and Planetary Sciences, California Institute of Technology, Pasadena, California, 43 44 USA ^{14†}now at Department of Geosciences, The Pennsylvania State University, University Park, PA, 45 46 ¹⁵Department of Earth, Planetary, and Space Sciences, University of California Los Angeles, Los 47 48 Angeles, CA, USA ¹⁶University of Washington, Seattle, Washington, USA 49 ¹⁷Department of Earth Sciences, University of Utrecht, Utrecht, Netherlands 50 ¹⁸Institute of Earth Sciences, Academia Sinica, Taipei, Taiwan 51 ¹⁹Macquarie University, Australia 52 ²⁰RSMAS/MGS University of Miami, Miami, FL, USA 53 ²¹Institute of Earth Surface Dynamics, University of Lausanne, Lausanne, Switzerland 54 ²²Earth-Life Science Institute, Tokyo Institute of Technology, Tokyo, Japan 55 ²³ Univ. Littoral Côte d'Opale, Univ. Lille, CNRS, Laboratoire d'Océanologie et de Géosciences 56 57 (UMR 8187 LOG), 62930 Wimereux, France. 58 59 Corresponding authors: Stefano M. Bernasconi (stefano.bernasconi@erdw.ethz.ch) Mathieu Daëron (daeron@lsce.ipsl.fr) 60 61 62 **Key Points:** The exclusive use of carbonate reference materials is a robust method for the 63 64 standardization of clumped isotope measurements Measurements using different acid temperatures, designs of preparation lines, and 65 mass spectrometers are statistically indistinguishable 66 We propose new consensus values for a set of 7 carbonate reference materials and 67 updated guidelines to report clumped isotope measurements 68

69

Abstract

70

71

72

73

74

75

76

77

78

79

80

81

82

83

84

85

86

87

88

Increased use and improved methodology of carbonate clumped isotope thermometry has greatly enhanced our ability to interrogate a suite of Earth-system processes. However, inter-laboratory discrepancies in quantifying carbonate clumped isotope (Δ_{47}) measurements persist, and their specific sources remain unclear. To address inter-laboratory differences, we first provide consensus values from the clumped isotope community for four carbonate standards relative to heated and equilibrated gases with 1,819 individual analyses from 10 laboratories. Then we analyzed the four carbonate standards along with three additional standards, spanning a broad range of δ^{47} and Δ_{47} compositions, for a total of 5,329 analyses on 25 individual mass spectrometers from 22 different laboratories. Treating three of the materials as known standards and the other four as unknowns, we find that the use of carbonate reference materials is a robust method for standardization that yields inter-laboratory discrepancies entirely consistent with inlaboratory analytical uncertainties. Carbonate reference materials, along with measurement and data processing practices described herein, provide the carbonate clumped isotope community with a robust approach to achieve inter-laboratory agreement as we continue to use and improve this powerful geochemical tool. We propose that carbonate clumped isotope data normalized to the carbonate reference materials described in this publication should be reported as Δ_{47} (I-CDES) for Intercarb-Carbon Dioxide Equilibrium Scale.

1 Introduction

89

90

91

92

93

94

95

96

97

98

99

100

101

102

103

104

105

106

107

108

109

110

111

112

113

114

115

116

117

118

119

Carbonate clumped isotope (Δ_{47}) thermometry is the most developed branch of the rapidly evolving field of clumped isotope geochemistry. Given the broad range of applications in Earth Sciences (e.g. Affek & Eiler, 2006; Eagle et al., 2010; Ferry et al., 2011; Ghosh, Adkins, et al., 2006; Ghosh, Garzione, et al., 2006; Grauel et al., 2013; Guo & Eiler, 2007; Huntington et al., 2011; Passey & Henkes, 2012, Dale et al., 2014, Mangenot et al., 2018, Veillard et al., 2019) and the improvement of analytical methods including automation (Bernasconi et al., 2013, 2018; Defliese & Lohmann, 2015; Dennis et al., 2011; Fiebig et al., 2019; Ghosh, Adkins, et al., 2006; He et al., 2012; Hu et al., 2014; Huntington et al., 2009; Meckler et al., 2014; Müller, Fernandez, et al., 2017; Passey et al., 2010; Petersen et al., 2019; Petersen et al., 2016; Schmid & Bernasconi, 2010, Adlan et al, 2020), the last 5-10 years have seen an increasing number of laboratories implementing this technique. The great potential of this thermometric method can only be fully exploited if the precision and accuracy of the data is sufficient to resolve differences of a few degrees in formation temperatures. In addition, like for any other geochemical method, widely available reference materials that match the sample matrices are necessary so that data can be robustly compared across laboratories (Meier-Augenstein & Schimmelmann, 2019). Currently the situation in the field of carbonate clumped isotope geochemistry is far from satisfactory. Published values for the ETH reference materials, the only carbonates that have been recently measured in many different laboratories worldwide, show a range of values for the same sample of up to 0.053 % (see Bernasconi et al. 2018, Thaler et al. 2020 for recent comparisons). This clearly calls for better standardization procedures to improve laboratory comparability.

The data normalization scheme currently used in clumped isotope geochemistry of carbonates in most laboratories is based on the comparison of the composition of the CO₂ liberated from carbonates by reaction with phosphoric acid with that of a set of CO₂ gases with different bulk and clumped isotope compositions (Dennis et al., 2011). These gases are prepared either by heating CO₂ at 1000 °C (heated gases; HG) or by CO₂ equilibration with water at low temperatures (equilibrated gases; EG, at *e.g.* 25 °C, 50 °C). By comparing the measured compositions with the theoretical predictions of the equilibrium thermodynamic abundance of multiply substituted isotopologues in heated and equilibrated gases (Wang et al. 2004 and

manuscript submitted to Geochemistry, Geophysics, Geosystems

updates in Petersen et al. 2019), the measurements are standardized to the scale that was named the "absolute reference frame" (ARF) by Dennis et al., (2011). In more recent publications the ARF is often referred to as the "Carbon Dioxide Equilibration Scale" (CDES), a terminology introduced by Passey and Henkes (2012). This approach was designed to allow different laboratories to link their measurements to an internationally recognized scale firmly anchored to theory using relatively easy and established laboratory protocols to produce CO₂ standard gases of known composition. Early comparisons of $\Delta_{47 \text{ CDES}}$ values for carbonates run in different laboratories and normalized with HG/EG normalization were promising (Dennis et al., 2011). While Bonifacie et al. (2017) reported similar $\Delta_{47 \text{ CDES}}$ values for nine dolomite samples covering a range of almost 0.4% measured at both at Caltech and IPGP laboratories with HG/EG normalization, Spooner et al. (2016) found that carbonate standardization improved agreement between data they obtained on samples ran both at Caltech and WHOI laboratories, compared to when they were using HG/EG normalization. Such recurrent cases of poor inter-laboratory reproducibility (see also Bernasconi et al., 2018 and Thaler et al., 2020) suggest that there is still unexplained differences in the results among laboratories (see Petersen et al. (2019) for a recent review).

Apart from preservation problems such as diagenesis and solid-state reordering, two known issues still limiting the reliability of this method to yield accurate temperature reconstructions are: [1] the lack of internationally recognized carbonate reference materials for a precise interlaboratory calibration, and [2] that published Δ_{47} -temperature calibrations produced in different laboratories have differed in both temperature dependence (slope), and absolute values (intercept). Possible reasons for the differences in slope and intercepts of the Δ_{47} temperature dependence have been widely discussed in the literature (e.g. Kluge et al, 2015; Bonifacie et al., 2017; Daëron et al., 2016; Fernandez et al., 2017; Katz et al., 2017; Kelson et al., 2017; Petersen et al., 2019; Schauer et al., 2016). Discrepancies have been attributed to analytical artefacts such as CO_2 -acid re-equilibration at different acid digestion temperatures (see Wacker et al. 2013; Swart et al., 2019, for a recent discussion). Other factors proposed to influence the calculated slope of the calibrations are the limitations of the datasets used in the individual studies, in particular in terms of the number of samples and replicates and of the temperature range covered by the available samples (Bonifacie et al., 2017; Fernandez et al., 2017). However, the

discrepancies in the intercept of the calibrations, for example between Kelson et al. (2017) and Peral et al. (2018), and a generally poor laboratory comparability remain problems that should be mitigated by using a more robust standardization method.

Petersen et al. (2019), in a recent effort to resolve differences in calibrations, compiled raw data of a number of published temperature calibrations and recalculated them all in a consistent way using the revised "IUPAC" correction parameters to correct for the ^{17}O abundance (Daëron et al., 2016; Schauer et al., 2016). The goal was to test whether data processing differences and/or the use of consistent but incorrect ^{17}O - correction parameters in the calculations were the root causes of inconsistencies. The result of this study was that differences among calibrations were reduced but not eliminated by the recalculation, implying that other factors must be responsible for the remaining discrepancies. These differences have pushed many laboratories to use laboratory-specific calibrations performed with the same analytical approach, as they at least partially take into consideration possible procedural differences (Petersen et al., 2019). However, if a laboratory changes analytical procedures or has not generated a robust in-house calibration, this approach is problematic. Achieving an inter-laboratory reproducibility at the level of accuracy necessary for meaningful interpretations of the observed variations is a requirement for Δ_{47} thermometry to reach its potential as a mature analytical method with broad acceptance and quantitative usefulness.

While the definition of the CDES was a major milestone (Dennis et al., 2011), a known problem with this approach is that while the CO₂ standard gases equilibrated at known temperature (HG or EG) can be confidently used for correction of mass spectrometric fractionations/nonlinearities and for effects of the purification procedures, they cannot account for the effects of the phosphoric acid reaction on the composition of the produced CO₂. Among the factors responsible for discrepant calibrations and laboratory comparability, two important ones cannot be tested with a gas-based standardization: (1) the absolute value and temperature dependence of the phosphoric acid fractionation factor (see Petersen et al., 2019 for a recent compilation) and (2) possible CO₂ equilibration effects during acid digestion of the sample. Swart et al. (2019) presented evidence that equilibration of CO₂ with water or hot metal surfaces during phosphoric acid reaction and transfer of the CO₂ to the mass spectrometer could be a factor leading to the

183

184

185

186

187

188

189

190

191

192

193

194

195

196

197

198

199

200

201

202

203

204

205

206

207

208

209

210

211

212

manuscript submitted to Geochemistry, Geophysics, Geosystems

alteration of the apparent temperature dependence of clumped isotopes in carbonates and on the absolute value of calculated Δ_{47} . As many laboratories use custom built extraction lines with different designs and volumes of tubing and of acid vessels, these factors are impossible to precisely quantify for each laboratory and may further contribute to inter-laboratory discrepancies. We propose that these issues can be circumvented if carbonates, which undergo the same acid digestion as the samples, are used for normalization instead of or in addition to gases, consistent with the principle of identical treatment of sample and standards (Carter & Fry, 2013; Werner & Brand, 2001a). In addition, normalizing results to accepted carbonate reference material values, as is commonly done with conventional carbon and oxygen isotope analysis in carbonates, removes the requirement to precisely quantify acid fractionation factors at different temperatures (Bernasconi et al., 2018). A carbonate standardization approach was introduced by Schmid & Bernasconi (2010) and improved by Meckler et al. (2014), with the following benefits: (1) the use of carbonates can more easily be fully automated, eliminating time-consuming and possibly error-prone manual preparation of CO₂ standard gases (equilibrated at known temperature) by individual users on separate extraction lines; (2) in some automated systems designed for small sample sizes, the gases had to be measured through a different capillary than the carbonates with potential biases that would go unrecognized; and (3) in those systems the equilibrated and heated gases are measured at constant beam intensity, whereas the samples are measured with decreasing beams. These features argue in favor of carbonate standardization a priori, but it remains critical to assess a posteriori whether the results of this approach are as robust and accurate as expected, and whether they significantly improve the inter-laboratory reproducibility of Δ_{47} measurements. Discussions at the Sixth International Clumped Isotope Workshop (Paris, 2017) led to the present inter-laboratory comparison exercise (InterCarb) to evaluate the benefits and drawbacks of a carbonate-based standardization approach as an alternative to the use of gas standards. The primary goal of this study was to test whether the exclusive use of carbonate reference materials can minimize inter-laboratory discrepancies and provide an alternative to the measurement of heated and equilibrated gases for the entire community. This is particularly

important because of the increasing number of laboratories using commercial small-sample automated devices which cannot easily be standardized using the HG-EG approach. The InterCarb exercise also provides an opportunity to define the best community-derived absolute Δ_{47} values for the ETH standards of Meckler et al. (2014). Although these standards are already used in many laboratories, their current nominal Δ_{47} values are based on measurements from the ETH laboratory only. The InterCarb exercise can similarly establish community accepted values for other common carbonate reference materials, some of which have been in use for several years, in order to provide the community with a self-consistent set of carbonate reference materials with a broad range of bulk and clumped isotope compositions.

1.1. Nomenclature and data processing

Clumped isotope compositions are reported as an excess abundance of the CO₂ isotopologue of cardinal mass 47 (dominantly the isotopologues ¹³C¹⁸O¹⁶O) compared to a stochastic distribution according to the formula:

$$\Box_{47} = \Box^{47} / \Box^{47*} - 1$$

where R^{47} is the abundance of the minor isotopologue with mass 47 relative to the most abundant isotopologue with mass 44. The stochastic ratio R^{47*} is calculated using the measured abundance of 13 C and 18 O in the sample (Affek & Eiler, 2006). According to the IUPAC guidelines the formula does not include the factor 1000 (Coplen, 2011). Also we omit here the classically included terms involving R45* and R46*, which are assumed to be zero by definition when computing δ^{13} C and δ^{18} O, and in practice never exceed ± 0.00002 % in our calculations (Daeron et al. 2016). The measured abundance of isotopologues with m/z 47 in the sample with respect to the working gas (WG) in the mass spectrometer is reported in the traditional delta notation as:

$$\Box^{47} = \Box^{47} / \Box^{47}_{\Box\Box} - 1$$

The δ^{47} scale is a measure of the difference between the sample of interest and the WG of the specific instrument, therefore, it cannot be compared across laboratories. The same notation is used for masses 45, 46, 48 and 49.

240

241

242

243

244

245

246

247

248

249

250

251252

253

254

255

256

257

258

259

260

261

262

263

264

265

266

The CO₂ gas-based standardization scheme for clumped isotope thermometry in carbonates relies on a set of CO₂ standard gases with different bulk compositions (δ^{47}), preferably chosen by the user to encompass the δ^{47} values of unknown samples, that have been (1) heated at 1000°C to reach a near-stochastic distribution of all isotopologues, and (2) equilibrated with water at low temperatures to reach equilibrium enrichments in the mass-47 isotopologues (Dennis et al., 2011). The heated gases, having a stochastic distribution of the heavy isotopes among all isotopologues, define the zero point of the CDES scale ($\Delta_{47} = 0.0266$ %) and the equilibrated gases define a high point (e.g., at 25 °C Δ_{47} = 0.9196 %). The theoretical values linking measurements to theory were calculated by Wang et al. (2004), and revised by Petersen et al. (2019). A wide range in δ^{47} values of gases used for normalization is generally chosen to allow for accurate correction for an apparent dependence of Δ_{47} on δ^{47} which is caused by inaccurate pressure-dependent background corrections on the m/z 47 collector observed on many instruments (Bernasconi et al., 2013; He et al., 2012). The large range in Δ_{47} (i.e., 25 °C, 1000 °C), on the other hand, is necessary to correct for scale compression caused by processes of scrambling and molecule recombination in the source of the mass spectrometer or elsewhere in the sample preparation, transfer lines and/or the capillaries (Dennis et al., 2011, Swart et al. 2019). With properly chosen CO_2 standard gases with widely varying δ^{47} values it is possible to cover the entire range of natural carbonate compositions, avoiding extrapolations in the δ^{47} vs. Δ_{47} compositional space (Fig. 1). Note that with measurement errors (typically not better than 0.010 %) being relatively large compared to the natural compositional range (less than 0.5 %; Fig. 1), the large (0.9 ‰) difference in Δ_{47} of the CO₂ standard gases minimizes errors introduced by uncertainties resulting from the measurement of HG and EG.

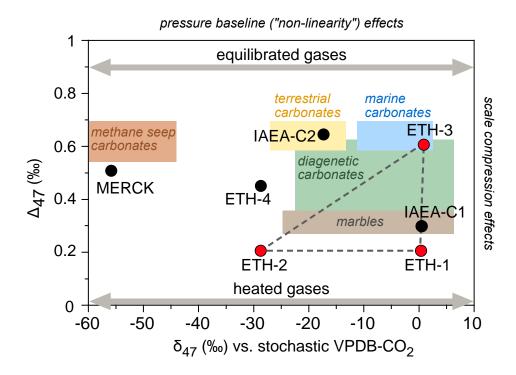


Figure 1. δ^{47} vs. Δ_{47} values of carbonate standards (Δ_{47} on the I-CDES scale proposed here) and heated and equilibrated gases in comparison to the compositional ranges of typical natural carbonates. The observed range in measured clumped isotope compositions in natural carbonates can be completely bracketed by heated and equilibrated CO₂ standard gases from which δ^{47} values have been chosen by the user. δ^{47} values for the anchor samples used in InterCarb (red) and the unknowns (black) are reported for a theoretical working gas with stochastic isotope distribution, derived from VPDB. Actual δ^{47} values will vary by laboratory depending on the composition of the working gas. Note the smaller achievable range in both δ^{47} and Δ_{47} values when using carbonate standards compared to heated and equilibrated gases, and the large extrapolation necessary for the determination of the composition for MERCK. Heated and equilibrated CO₂ standard gases have a larger Δ_{47} range, allowing for more robust stretching calculations with identical numbers of standard:sample analyses.

Meckler et al. (2014) attempted to achieve a similar framework as the CO₂ gas-based standardization but with carbonate standards. They described four carbonates that were developed at ETH Zürich to serve as replacements for HG-EGs and demonstrated that good long- and short-term reproducibility can be achieved using only carbonates for data correction. Bernasconi et. al (2018) discussed in detail these standards and postulated, based on a limited inter-laboratory dataset, that carbonate standardization should generally improve inter-laboratory

data comparability. This claim seems arguably strengthened by the results of Meinicke et al. (2020), Peral et al. (2018), Piasecki et al. (2019), Kele et al. (2015) as recalculated by Bernasconi et al. 2018), and Jautzy et al., (2020). The first three studies produced independent foraminiferabased the fourth a travertine and the fifth a synthetic carbonate-based Δ_{47} -temperature calibration anchored to the same set of carbonate standards, yielding statistically indistinguishable slopes and intercepts despite the use of independent sample sets and in the case of Peral et al. (2018), a different analytical system. In addition, a reanalysis of samples from five previous calibrations by Anderson et al. (2020) using carbonate standardization revealed no significant differences in temperature dependence of Δ_{47} between the different sample sets.

A possible limitation of carbonate standardization is that available carbonates have a smaller range in δ^{47} and, perhaps more importantly, a smaller range in Δ_{47} values than what is achievable with heated and equilibrated gases. In some specific cases, standardization procedures may require extrapolation to compositions that are not within the δ^{47} - Δ_{47} space created by carbonate standards (Fig. 1). In addition, the range of Δ_{47} values for carbonates is only on the order of 0.5 % between 0 and 1000 °C. The smaller range in Δ_{47} compared to HG-EG requires higher precision and thus high standard replication and a ~50:50 standard to sample replicate ratio to keep normalization errors small (see Daëron (this volume) and Kocken et al., 2019 for details).

1.2. InterCarb goals and design:

InterCarb was designed with the aim to carefully evaluate the potential of carbonates to serve as a standardization scheme that improves inter-laboratory agreement for 'unknown' carbonates both inside and outside of the δ^{47} - Δ_{47} space defined by the anchor samples (Fig. 2). The main questions posed are:

1. Is it possible to produce consistent carbonate clumped isotope measurements across laboratories using carbonate reference materials exclusively? In other words, does the observed inter-laboratory scatter in Δ_{47} values match that expected from intra-laboratory analytical precision?

- 317 2. How well does the carbonate standardization approach perform when extrapolating beyond the δ^{47} - Δ_{47} compositional space sampled by a set of carbonate reference materials?
- 320 3. Do carbonate reference materials fully correct effects arising from different reaction temperatures, sample preparation protocols, and analytical equipment?
- 322 4. Can we define a self-consistent set of widely available reference materials with community-agreed compositions accurately anchored to the CDES scale?
- 5. Does the use of carbonate reference materials for standardization improve the interlaboratory reproducibility compared to HG-EG standardization?

1.3 Approach:

Seven carbonate standards with a large range of δ^{47} and Δ_{47} values (Fig. 1) were distributed among participating laboratories and analyzed, treating three carbonates as "anchors" (whose Δ_{47} values are assigned *a priori*) and the remaining four as "unknowns" (whose Δ_{47} values are unknown, to be determined by comparison with the anchors). Due to their relatively widespread use in different laboratories, the three reference materials ETH-1, ETH-2 and ETH-3 (Meckler et al. 2014; Bernasconi et al. 2018) were chosen as anchors. They are still available today in relatively large quantities (>600 g), have been in use at ETH since 2013 and in many other laboratories for several years. Importantly, they have been thoroughly tested for homogeneity based on thousands of measurements in 80 to 150µg aliquot sizes in different laboratories and no changes in composition have been noticed at ETH in the 7 years they have been in use.

The "unknown" InterCarb reference materials were chosen to cover a wide natural range in δ^{47} and Δ_{47} values. These samples had to be available in large quantities, inexpensive, and if possible distributed by an organization with a long-term perspective in order to ensure future data quality and availability for the increasing number of laboratories globally.

346

347

348

349

350

351

352

353

354

355

356

357

358

359

360

361

362

363

364

365

366

367

368

369

370

371

372

373

manuscript submitted to Geochemistry, Geophysics, Geosystems

2 Materials and Methods 2.1. Sample description The anchor samples ETH-1 (Carrara marble heated at 600 °C), ETH-2 (synthetic carbonate heated at 600 °C) and ETH-3 (Upper cretaceous chalk) are described in detail in Bernasconi et al. (2018). IAEA-C1 (marble from Carrara, Italy) is distributed by the International Atomic Energy Agency (IAEA) as a mechanically crushed and milled product with grains ranging from 1.6 to 5 mm. All 50 g provided were ground and thoroughly homogenized in a ball mill at ETH Zürich to a grain size of less than 100 µm and transferred in 0.5 g aliquots to plastic vials for distribution. IAEA-C2 is a freshwater travertine from Bavaria distributed by IAEA as a powder which was treated identically to IAEA-C1. ETH-4 is a commercially available synthetic calcite (Riedel-De Haën; calcium carbonate Puriss. p.a.; Lot No. 30800) with intermediate formation temperature and the same bulk isotope composition as ETH-2 (see Bernasconi 2018 for details). MERCK (Catalog No. 1.02059.0050; lot no. B1164559 515) is an ultra-pure, commercially available synthetic calcite and was chosen for its very low δ^{13} C and δ^{18} O values, of approximately -41.7 % and -15.5 % (VPDB), respectively. This sample represents an extreme case of extrapolation from the δ^{47} - Δ_{47} space defined by the anchor materials (Fig. 2). 2.2 Instrumentation: The reported data were produced with a variety of preparation systems including custom built (13 laboratories) and commercial systems (11 laboratories; Protium MS IBEX, ThermoFisher Scientific Kiel IV device and Nu Instruments Nucarb). Reaction temperatures were generally 90

°C for "large-sample" custom preparation systems and 70 °C for the Kiel and the NuCarb. Four mass spectrometer types were used: Thermo Fisher Scientific MAT253 and 253Plus, Nu Instruments Perspective, and Elementar Isoprime 100. All participants contributed results they considered to be of a "publication-grade" quality, based on their existing quality-control procedures.

2.3 Clumped isotope compositions of the ETH anchor materials.

374

375

376

377

378

379

380

381

382

383

384

385

386

387

388

389

390

391

392

393

394

395

396

397

398

399

400

401

402

403

The clumped isotope compositions of the four ETH reference materials relative to the CO₂ reference frame CDES were first reassessed based on new data provided by 9 laboratories that also provided HG and EG data measured during the same sessions as the ETH reference materials. Data processing was performed with the same Python script that was used to process the carbonate data in order to avoid any differences in data processing (see section 2.4).

Although, strictly speaking, ¹³C-¹⁸O clumping in carbonate represents a mass-63 anomaly, the clumped isotope composition of carbonate minerals is reported as Δ_{47} i.e. as the mass-47 excess in the CO₂ produced by acid digestion of these minerals, including the respective temperaturedependent isotopic fractionation. As initially all reactions were carried out at 25°C (Ghosh et al. 2006), the Δ_{47} values have traditionally been reported for a 25°C acid temperature. With the advent of automated extraction lines, reaction temperatures have been raised to 70 or 90°C. To account for the temperature dependence of the acid fractionation factor (Guo et al., 2009, Petersen et al. 2019) and to project these results back to the original 25°C acid reaction, various acid temperature correction values have been reported over time, based on experimental observations and/or theoretical predictions. Given that here seven out of ten laboratories reacted carbonates at 90 °C, two at 70 °C, and only one at 25 °C, our redetermination of the Δ_{47} values of ETH-1/2/3/4 relative to the CDES projected to 25°C would substantially rely on the accuracy of these acid temperature corrections (which typically range between 60-90 ppm). Here we report instead the Δ_{47} values of CO₂ produced by reacting ETH-1/2/3/4 at 90 °C. With this choice the numerical effect of poorly known acid corrections is minimized, because the data from 70 °C and 25 °C reactions have relatively little influence on the final, error-weighted average Δ_{47} values (cf statistical weights in Fig. 2). We thus propose to break with tradition and

define the nominal Δ_{47} values of the anchor standards as those of CO₂ produced at 90 °C, providing the most robust relationship to the CDES.

2.4 Data processing, correction and error assessment.

 It should be stressed that the InterCarb experiment, by design, is not intended to grade the analytical "performance" of individual laboratories. Each participating laboratory (or mass spectrometer, in the case of laboratories with several instruments) was thus randomly assigned an anonymous identifying number. Within each laboratory, analyses were grouped in different analytical sessions defined by the participants themselves. An analytical session is generally defined by a time in which the behavior of the analytical system (preparation system, source tuning, backgrounds, isotope scrambling in the source) is considered to be similar. The database record of each analysis consists of a laboratory identifier; a session identifier; an analysis identifier; the name of the analyzed sample; the mass spectrometer model; the acid reaction temperature; the mass of reagent carbonate; and background-corrected δ^{45} , δ^{46} and δ^{47} values.

The only instrumental corrections to the raw data applied independently by each participating laboratory were background corrections ("Pressure Baseline Correction" or PBL) to the ion currents/voltages (Bernasconi et al., 2013; He et al., 2012). The PBL is strongly dependent on instrument design (it is not observed in some instruments) and configuration, and varies temporally depending on many factors. This correction, therefore, can only be carried out by each participating laboratory according to its own established procedures and monitoring.

To avoid artefacts arising from different calculation/standardization procedures, rounding errors, and ¹⁷O correction parameters, raw data from all laboratories were processed by a single Python script (http://doi.org/10.5281/zenodo.4314449) based on data reduction, standardization and error propagation methods described in detail in the companion paper (Daëron, this volume). Here we briefly summarize these calculations.

Session-averaged, background-corrected δ^{45} and δ^{46} values for each of the three anchor samples were first used to calculate the bulk isotope composition of the working gas used in each session, based on (a) previously reported $\delta^{13}C_{VPDB}$ and $\delta^{18}O_{VPDB}$ values of ETH-1, ETH-2, and ETH-3 (Bernasconi et al., 2018), (b) the IUPAC ¹⁷O correction parameters of Brand et al. (2010), and (c) a temperature-dependent oxygen-18 acid fractionation factor between CO_2 and calcite of Kim et al. (2015). This recalculation of working gas bulk compositions avoids (small) discrepancies potentially introduced by inaccuracies in the nominal compositions of the working gases.

Raw Δ_{47} values were computed according to:

$$\square_{47}^{\square\square\square} = \square^{47} / \square^{47*} - I$$

Where R^{47} is the measured ratio and R^{47*} the calculated stochastic ratio of mass 47 over mass 44 of CO_2 , assuming perfectly linear IRMS measurements and a stochastic working gas. Values are then normalized to "absolute" Δ_{47} values (noted $\Box_{47}^{\Box\Box\Box}$ in the equation below, and simply Δ_{47} thereafter) using session-specific relationships of the form:

448
$$\square_{47}^{raw} = a \square_{47}^{\square\square} + b \delta^{47} + c$$
449

For each session, the best-fit standardization parameters (a, b, c) are computed from an unweighted least-squares regression, treating \Box_{47}^{raw} as the response variable, only considering the three anchor samples ETH-1, ETH-2, and ETH-3. Note the advantage of this form over that in Dennis et al. 2011 is the ability to have three standards with distinct Δ_{47} values whilst being able to solve for b (compositional nonlinearity) (Daëron et al. 2016). Absolute Δ_{47} values are then computed for all replicates within that session. Standardization parameters for all sessions are listed in Table S1 (supplementary information).

In figure labels and in tables, measurement uncertainties are reported as 1SE of the mean, considering fully-propagated errors taking into account reference frame corrections. In Figures 2 and 4, different types of error bars reflect either these full analytical errors, or those only considering uncertainties in the analyses of a given sample. In both cases, the analytical error assigned to each individual raw Δ_{47} analysis is equal to the pooled "external" repeatability of raw

manuscript submitted to Geochemistry, Geophysics, Geosystems

 Δ_{47} measurements for all samples (anchors and unknowns) within each session. This treatment of error is a new approach that more fully accounts for error in both the sample measurement and reference frame.

3 Results and Discussion

3.1 Redetermination of nominal Δ_{47} values for the ETH standards relative to heated and equilibrated CO_2 gases

The weighted averages of the 4 standards (projected to 90 °C for the reactions at 25 and 70 °C using the acid temperature correction suggested by Petersen et al. 2019), comprising 619 analyses of the carbonate standards and 828 heated and equilibrated gases from 9 different laboratories, are reported in Table 1 and Fig 4. The large number of analyses and the appropriate consideration of the errors on the anchors (CO_2 gas analyses) distinguishes this effort from previous work, and allows a robust redetermination of the accepted values of the ETH reference materials with 1SE uncertainties of 2 ppm or less.

manuscript submitted to Geochemistry, Geophysics, Geosystems

	Laboratory	all	A	В	C	D	E	F	G	Н	I	J
	N of sessions	34	4	4	11	7	1	2	1	1	1	2
	N of H/E CO2	946	44	193	257	85	47	21	38	192	13	56
ETH-1	N of analyses	232	34	14	5	54	4	4	19	58	8	32
	$\Delta_{47}(90~^{\circ}\mathrm{C}$ acid)	0.2052	0.2016	0.1926	0.2108	0.1940	0.1601	0.2013	0.2143	0.1932	0.2183	0.2152
	± 1SE	0.0016	0.0046	0.0058	0.0069	0.0042	0.0245	0.0107	0.0032	0.0045	0.0109	0.0036
	Statistical weight		0.118	0.074	0.053	0.146	0.004	0.022	0.241	0.124	0.021	0.197
ETH-2	N of analyses	215	23	13	11	51	4	4	18	51	8	32
	Δ_{47} (90 °C acid)	0.2085	0.2077	0.1840	0.2225	0.1978	0.1374	0.1650	0.2141	0.1968	0.2172	0.2170
	± 1SE	0.0015	0.0047	0.0070	0.0046	0.0050	0.0233	0.0101	0.0029	0.0043	0.0154	0.0033
	Statistical weight		0.105	0.047	0.108	0.092	0.004	0.023	0.272	0.125	0.010	0.213
ЕТН-3	N of analyses	264	55	15	20	54	4	5	15	59	8	29
	$\Delta_{47}(90~^{\circ}\text{C acid})$	0.6132	0.6156	0.5975	0.6169	0.6102	0.5950	0.6143	0.6159	0.6094	0.6428	0.6124
	± 1SE	0.0014	0.0037	0.0056	0.0033	0.0038	0.0237	0.0099	0.0033	0.0042	0.0103	0.0035
	Statistical weight		0.140	0.062	0.175	0.134	0.003	0.020	0.179	0.110	0.018	0.158
ETH-4	N of analyses	162	10	12	5	55	4	4	12	47	7	6
	Δ_{47} (90 °C acid)	0.4505	0.4438	0.4230	0.4624	0.4506	0.4230	0.4454	0.4560	0.4414	0.4831	0.4646
	± 1SE	0.0018	0.0058	0.0071	0.0068	0.0049	0.0226	0.0095	0.0032	0.0042	0.0161	0.0057
	Statistical weight		0.093	0.064	0.068	0.133	0.006	0.035	0.314	0.177	0.012	0.097

Table 1. Newly determined nominal Δ_{47} values of the ETH standards projected to 90 °C acid reaction using acid correction factors of -0.088 ‰ and -0.022 ‰ for 25 °C and 70 °C reactions, respectively (Petersen et al. 2019). Reported standard errors represent analytical uncertainties associated both with reference frame errors (HG/EG) and carbonate sample reproducibility (Daëron, this volume).

When compared with the published values in Bernasconi et al. (2018) the average Δ_{47} values ETH-1 and ETH-2, projected back to 25 °C (+0.088 %), are respectively 0.035 and 0.040 % greater than the original values, whereas ETH-3 increases by 0.010 and ETH-4 by 0.031 %. Such positive offset of Δ_{47} values reported in Bernasconi et al. (2018) versus other laboratories has also been reported in Fiebig et al. (2019) and Thaler et al. (2020).

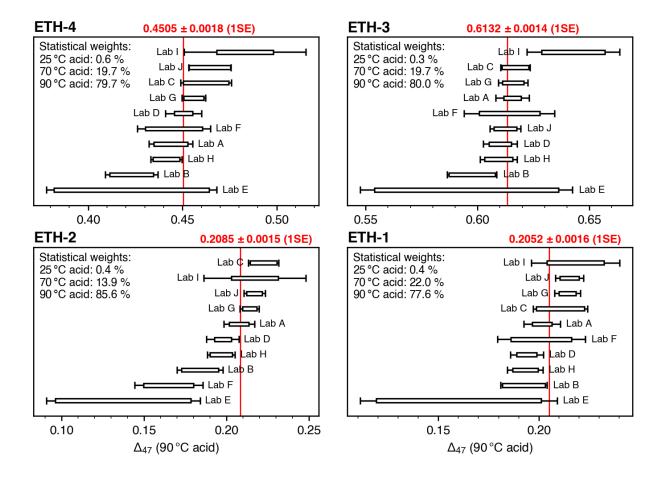


Figure 2. New determination of Δ_{47} values for the four ETH standards relative to the CDES using updated CO₂ equilibrium values. These measurements, using acid reaction temperatures of 90 °C, 70 °C or 25 °C, are projected to 90 °C using acid corrections of -0.088 ‰ and -0.022 ‰ for 25 °C and 70 °C reactions, respectively (Petersen et al. 2019). Error bars correspond to 95 % confidence limits taking into account fully propagated errors (i.e. taking into account errors in both unknown and anchor analyses). Boxes correspond to 95 % confidence limits not accounting for normalization errors (i.e. only taking into account errors in unknown analyses). Red numbers are the error-weighted average values (with statistical weights summarized in upper-left corners). Note different horizontal scales for the different samples.

The observation that these changes in nominal values decrease as Δ_{47} increases suggests a simple hypothesis to explain this discrepancy: in the original study of Meckler et al. (2014), the carbonate samples and the heated/equilibrated CO_2 gases experienced different analytical procedures. The HGs were measured as large samples at constant beam intensity through a

different capillary than the carbonates, which were measured using the microvolume and a decreasing beam). The potential effects of partial re-equilibration for the heated gases in the gas preparation line or in the capillaries of the mass spectrometer could be significant whereas it would be minuscule for the gases equilibrated at 25°, leading to an overestimation of Δ_{47} scale compression and thus of the stretching applied to the Δ_{47} scale towards theoretical values. The observed changes in apparent ETH-1 and ETH-2 Δ_{47} values may therefore simply reflect partial re-equilibration of heated gases at the time of measurements at ETH (and reported in Meckler et al., 2014), increasing their values in the original study by about 0.05 ‰ (Fig. 3).

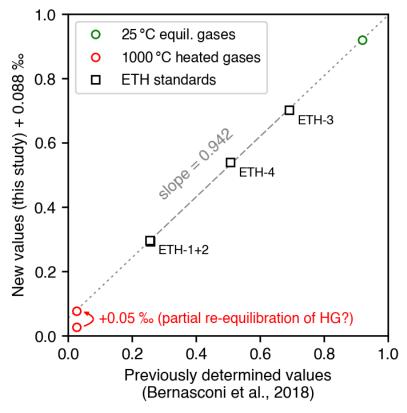


Figure 3. New nominal Δ_{47} values for the ETH standards compared to previously reported ones. The dashed gray line is a linear regression through the new versus old values of ETH-1/2/3/4, whose extrapolation coincides with 25 °C equilibrated CO₂ but not with heated gases. Apparent changes in the ETH-1/2/3/4 values thus scale linearly with the Δ_{47} difference between carbonate samples and 25 °C equilibrated CO₂, suggesting that Δ_{47} values of heated gases in the original study may have been biased by $\sim +0.05$ % through partial re-equilibration at room temperature between the quenching of heated CO₂ and its ionization in the IRMS source.

524

525

526

527

528

529

530

531

532

533

534

535

536

537

538

539

540

541

542

543

544

545

546

547

548

549

550

551

552

553

manuscript submitted to Geochemistry, Geophysics, Geosystems

It has been suggested previously that ETH-1 and ETH-2 should be indistinguishable in Δ_{47} and close to stochastic distribution (Müller, Violay, et al., 2017). This is because Δ_{47} values of ETH-1 and ETH-2, originally heated to 600 °C, were found to be higher by only around 0.006 % from the same carbonates heated at 1000 °C to achieve stochastic distribution of the isotopes. However, additional test measurements in multiple laboratories of samples heated at >1000 °C are necessary to confirm this observation. One laboratory (Lab F) did however observe a large difference in the value for ETH-1 and ETH-2, although their values of ETH-3 and ETH-4 are similar to other laboratories. The reason for these inconsistencies is probably due to the fact that ETH-1 was only measured four times with a limited number of HG/EG, and ETH-2 and ETH-4 were not measured in the same session. In addition, the labs with the smallest number of replicate measurements have uncertainties that are systematically larger (Table 1). These results highlight the importance of correction procedures in clumped isotope analysis. Sufficient replication of both standards and samples is critical and, if insufficient, offsets can arise when comparing results from different sessions. Due to these difficulties it is good practice to spread replicates of the same sample in different sessions over longer periods of time to obtain accurate results and follow a ~50:50 standard to sample replicate ratio. Based on the results above, the difference between the average of ETH1/2 and ETH-3 is reduced by 0.028 \,\times, thus leading to a compression of the scale by about 5.6 \times compared to the values reported by Bernasconi et al. (2018). As a consequence, the slopes of published temperature calibrations produced with carbonate standardization (Kele et al., 2015; Bernasconi et al., 2018; Jautzy et al., 2021; Meinicke et al., 2020; Peral et al., 2018; Piasecki et al., 2019) will become slightly shallower, with more positive y-intercepts. If Δ_{47} results from previous publications are also recalculated with the new standard values (see section 3.4), however, changes in calculated formation temperatures will be negligible. For this reason, when comparing data from publications using old accepted values of the ETH standards for standardization (either those published by Meckler et al. (2014) or those recalculated with the IUPAC parameters by Bernasconi et al. (2018)) to newer data, it is recommended to directly compare the reconstructed temperatures rather than recalculating Δ_{47} . Full recalculation of old measurements requires the

availability of the entire dataset including standards and the same correction procedures (e.g., averaging method) used in the original publications, and is described in Appendix 1.

3.2 InterCarb results

Results for the unknown carbonate samples were obtained from 26 mass spectrometers in 22 laboratories. Δ_{47} values of the 4 unknown samples were normalized to the new community-derived values of ETH-1, ETH-2, and ETH-3 (Table 1), averaged per individual analytical session and mass spectrometer (Tables 2 and S1). Sample mean values obtained in each mass spectrometer are shown in Figure 4. The details of each analytical session, including the number of samples and standards measured, the composition of the working standard, the scaling parameters and the reproducibilities of the individual sessions are given in the supplementary information (Table S1). Some laboratories reported data for only a subset of the unknown samples, and both replication level and analytical reproducibility vary from laboratory to laboratory (Table S1).

To clearly distinguish Δ_{47} values normalized to the carbon dioxide reference scale (CDES) using carbonates rather than heated and equilibrated gases, we propose the new acronym (I-CDES), short for InterCarb-CDES, to reflect the use of the proposed InterCarb reference materials for data standardization (see section 3.5 for more details).

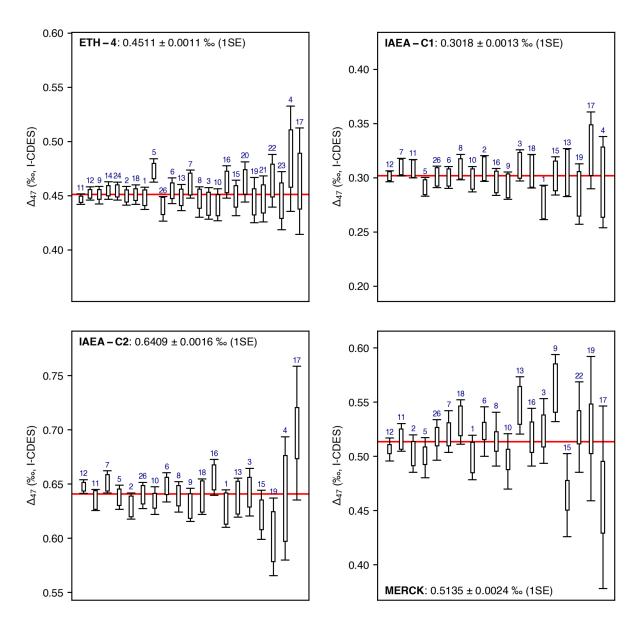


Figure 4. Final InterCarb results by laboratory. Error bars correspond to fully propagated 95 % confidence limits, taking into account errors in both unknown and anchor analyses. Boxes correspond to 95 % confidence limits not accounting for normalization errors (i.e., only taking into account errors in unknown analyses). Results are sorted by increasing analytical errors, and laboratories are identified by number. Overall error weighted average Δ_{47} values are displayed as solid red lines and reported in each panel. All plots have the same vertical scale.

Laboratories with stronger analytical constraints (i.e. better in-lab repeatability of Δ_{47} measurements and/or greater number of analyses) generally converge towards the overall mean

manuscript submitted to Geochemistry, Geophysics, Geosystems

value for each unknown sample. This observation suggests that the observed inter-laboratory variability is largely due to random errors that can be alleviated by replication, even for laboratories with relatively large analytical errors on individual measurements. It is also notable that fully propagated analytical errors (i.e., taking into account uncertainties in the standardization procedure) can be substantially larger than the errors based only on the uncertainty associated with unknown sample analyses, which is what is generally reported in the literature. The increase in error is also related to in-lab repeatability and the number of standards measured. In addition, error increases for unknown samples whose compositions lie outside the "anchor triangle" defined by ETH-1/2/3. This is illustrated by the increased scatter and errors associated with MERCK, the carbonate material farthest from the "anchor triangle", consistent with the models of Daëron (this volume, see also Kocken et al., 2019).

As seen in Table S1, there are stark differences in the total number of replicate analyses and the typical Δ_{47} reproducibility achieved in different laboratories. As a result, final uncertainties in the average Δ_{47} values of unknown samples vary considerably (Fig. 4). It is thus not very useful to quantify inter-laboratory reproducibility in terms of a single, overall scatter of Δ_{47} values: interlaboratory variability should be small among laboratories with small analytical uncertainties, and larger among laboratories with few replicate analyses and/or poor analytical repeatability. We may still, however, assess whether inter-lab discrepancies are significantly larger than expected from in-lab analytical uncertainties, i.e., whether we can detect the effects of hypothetical unrecognized sources of scatter beyond known analytical errors.

ETH-4			IAEA-C1	IAEA-C2		MERCK		
Lab	$\Delta_{47\text{(I-CDES)}} \text{ (\%0 ± 1SE)}$	N	$\Delta_{47 \text{(I-CDES)}} \text{ (\%o } \pm 1\text{SE)}$	N	$\Delta_{47 \text{(I-CDES)}} \text{ (\%o} \pm 1\text{SE)}$	N	$\Delta_{47 \text{(I-CDES)}} \text{ (\%o} \pm 1\text{SE)}$	N
1	0.4477 ± 0.0052	80	0.2773 ± 0.0080	19	0.6275 ± 0.0088	21	0.4991 ± 0.0105	22
2	0.4499 ± 0.0044	22	0.3086 ± 0.0060	6	0.6299 ± 0.0061	8	0.5025 ± 0.0089	6
3	0.4430 ± 0.0074	23	0.3114 ± 0.0073	17	0.6427 ± 0.0112	13	0.5235 ± 0.0152	11
4	0.4841 ± 0.0248	9	0.2959 ± 0.0215	6	0.6368 ± 0.0291	4		
5	0.4734 ± 0.0055	22	0.2916 ± 0.0044	18	0.6378 ± 0.0057	17	0.4987 ± 0.0094	14
6	0.4545 ± 0.0060	12	0.3004 ± 0.0051	12	0.6471 ± 0.0069	12	0.5229 ± 0.0116	12
7	0.4607 ± 0.0066	8	0.3099 ± 0.0042	16	0.6520 ± 0.0052	15	0.5231 ± 0.0098	8
8	0.4442 ± 0.0072	73	0.3099 ± 0.0060	67	0.6383 ± 0.0071	62	0.5159 ± 0.0127	70
9	0.4505 ± 0.0041	65	0.2926 ± 0.0064	8	0.6309 ± 0.0078	8	0.5630 ± 0.0158	2
10	0.4416 ± 0.0075	26	0.2987 ± 0.0060	33	0.6348 ± 0.0065	62	0.4954 ± 0.0130	31
11	0.4468 ± 0.0025	222	0.3085 ± 0.0043	31	0.6354 ± 0.0050	27	0.5175 ± 0.0066	25
12	0.4521 ± 0.0032	21	0.3015 ± 0.0026	20	0.6479 ± 0.0032	21	0.5064 ± 0.0054	19
13	0.4484 ± 0.0062	47	0.3048 ± 0.0113	6	0.6376 ± 0.0091	12	0.5470 ± 0.0135	9
14	0.4548 ± 0.0041	46						
15	0.4480 ± 0.0083	8	0.3016 ± 0.0090	4	0.6217 ± 0.0116	4	0.4642 ± 0.0195	4
16	0.4627 ± 0.0076	4	0.2962 ± 0.0063	4	0.6563 ± 0.0084	3	0.5176 ± 0.0136	2
17	0.4634 ± 0.0250	5	0.3254 ± 0.0181	6	0.6971 ± 0.0314	6	0.4623 ± 0.0429	3
18	0.4510 ± 0.0046	196	0.3060 ± 0.0079	26	0.6386 ± 0.0084	28	0.5317 ± 0.0104	31
19	0.4460 ± 0.0106	14	0.2851 ± 0.0142	5	0.6015 ± 0.0183	4	0.5256 ± 0.0339	4
20	0.4627 ± 0.0095	6						
21	0.4470 ± 0.0108	3						
22	0.4639 ± 0.0124	7					0.5269 ± 0.0213	7
23	0.4453 ± 0.0137	6						
24	0.4544 ± 0.0042	12						
26	0.4378 ± 0.0058	8	0.3008 ± 0.0051	6	0.6396 ± 0.0062	6	0.5152 ± 0.0095	6
w. avg	0.4511 ± 0.0011	945	0.3018 ± 0.0013	310	0.6409 ± 0.0016	333	0.5135 ± 0.0024	286

Table 2. Average Δ_{47} values (±1SE, fully propagated uncertainties) obtained by each participating laboratory. A full table of average Δ_{47} values for all analytical sessions is included in supplementary materials (Table S3).

616

617

618

619

620

621

622

623

624

625

626

627

628

629

630

631

632

633

634

635

636

637

638

639

640

641

642

manuscript submitted to Geochemistry, Geophysics, Geosystems

In order to do so, we compute the "number-of-sigma" deviation obtained by each laboratory for each unknown sample, relative to that sample's overall weighted average value. For example, the sigma-deviation for sample ETH-4 and Lab01 is equal to (0.4477 - 0.4511) / 0.0052 = -0.66 and that for MERCK and Lab13 is equal to (0.5470 - 0.5135) / 0.0135 = +2.48. If the analytical errors reported in Table 2 are reasonably accurate, we expect the population of sigma-deviations among all labs to be distributed as the canonical Gaussian distribution ($\mu = 0$; $\sigma = 1$), and we can test this prediction using established statistical methods such as a Kolmogorov-Smirnov test of normality (Massey, 1951). We carried out this test for two cases: only considering the error of sample replication (Fig. 5, upper row) and secondly including the normalization error, (i.e. the fully propagated error (Fig. 5, lower row). If we neglect uncertainties arising from standardization (the "allogenic" errors of Daëron, this volume), the sigma-deviations are no longer normally distributed (p = 0.003, upper-left panel). As shown in the lower-left panel of Figure 6, the distribution of sigma-deviations for all labs and all samples is statistically indistinguishable from the expected normal distribution (p = 0.19) when considering fully propagated analytical errors. Figure 6 also illustrates that neglecting standardization errors does not strongly affect the normality of sigma-deviations for ETH-4 and IAEA-C1, both of which have δ^{47} and Δ_{47} values within the range covered by the three anchor samples. By contrast, sigma-deviations for unknowns with "exotic" isotopic compositions (IAEA-C2 and especially MERCK) are only normally distributed if standardization uncertainties are correctly accounted for.

Based on these tests, we conclude that the inter-laboratory scatter observed in the InterCarb data set is neither smaller nor larger than expected from the analytical uncertainties computed within each laboratory, as long as standardization errors are taken into account. This important finding implies that, at least for the time being, we can rule out any systematic inter-laboratory discrepancies in carbonate-standardized Δ_{47} measurements, which constitutes an important milestone in the progress of clumped isotope measurement techniques.

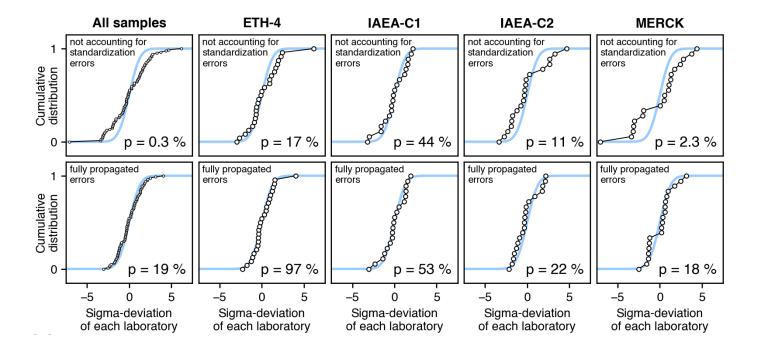


Figure 5. Kolmogorov–Smirnov tests of normality for the sigma-deviations obtained in each laboratory participating in the Intercarb effort (circular markers), either neglecting standardization uncertainty (upper row) or considering fully propagated analytical errors (lower row). Lower-right corner Kolmogorov–Smirnov p-values correspond to the null hypothesis that the sigma-deviations are normally distributed with a mean of zero and a standard deviation of 1. Blue lines correspond to the canonical Gaussian distribution ($\mu = 0$; $\sigma = 1$).

On demonstrating that we can fully account for inter-laboratory error using carbonate standardization, we revisit the results obtained for ETH1-4 using HG and EG (Fig. 2). Applying the same Kolmogorov–Smirnov test of normality yields unambiguous evidence that the inter-laboratory scatter observed here, using HG/EG standardization, is significantly greater than predicted from known analytical errors alone (p = 0.005, Fig. 6), contrary to the opposite finding for the InterCarb results, using carbonate-based standardization (p = 0.19, Fig. 5).

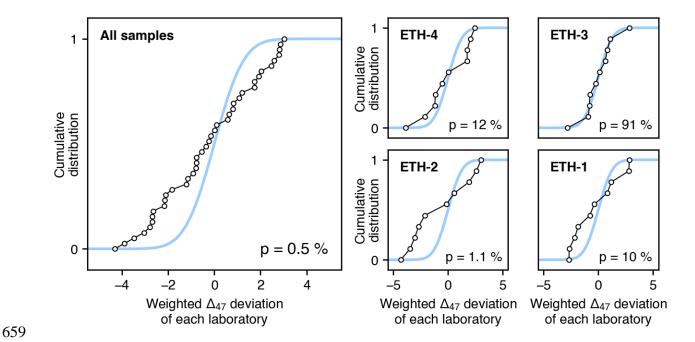


Figure 6. Kolmogorov–Smirnov tests of normality for the sigma-deviations, considering fully propagated analytical errors (accounting for uncertainties associated with conversion to the CDES reference frame), obtained in each laboratory participating in the ETH-1/2/3/4 determination using HG and EG (circular markers). Lower-right corner Kolmogorov–Smirnov p-values correspond to the null hypothesis that the sigma-deviations are normally distributed with a mean of zero and a standard deviation of 1. Blue lines correspond to the canonical Gaussian distribution ($\mu = 0$; $\sigma = 1$).

3.3 Effects of acid reaction temperature and IRMS models

Out of 25 sample preparation systems, 10 convert samples to CO_2 by acid reactions at 70 °C and 14 at 90 °C. To test for the possible effect of acid temperature, a commonly discussed cause for different slopes in the published temperature calibration curves (Came et al., 2014; Fernandez et al., 2014; Swart et al., 2019), we plot the Δ_{47} values of unknowns obtained by laboratories reacting at 70 °C vs. those obtained at 90 °C (Fig. 8, Table 3). Because acid fractionation effects equally affect anchors and unknowns , carbonate-standardized results can be compared directly without acid temperature correction. Δ_{47} values averaged by acid temperature are statistically indistinguishable for all of the unknowns. This implies that relative Δ_{47} differences between CO_2

evolved from different samples are independent of acid reaction temperature within the range of experimental conditions covered here, and for a very wide range of $\Delta_{47(I-CDES)}$ values spanning 0.302 % (marbles) to 0.641 % (carbonates formed at ambient temperatures).

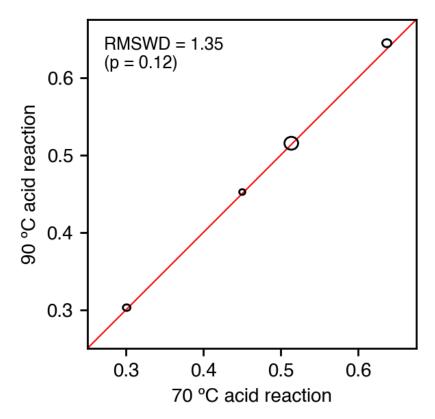


Figure 7. Error-weighted average Δ 47(I-CDES) values of unknowns obtained from acid reactions at 90 °C vs 70 °C. Solid black ellipses correspond to 95 % confidence limits. (see also Table 3).

	70 °C acid	90 °C acid	Difference (±1SE)
ETH-4	0.4501 ± 0.0016	0.4521 ± 0.0015	0.0020 ± 0.0022
IAEA-C1	0.3006 ± 0.0020	0.3026 ± 0.0017	0.0020 ± 0.0026
IAEA-C2	0.6369 ± 0.0024	0.6445 ± 0.0021	0.0076 ± 0.0032
MERCK	0.5134 ± 0.0036	0.5151 ± 0.0034	0.0017 ± 0.0049
average (all samples)			0.0033 ± 0.0017

Table 3. Error-weighted average $\Delta_{47(I\text{-}CDES)}$ values (±1SE) for each unknown as a function of acid reaction temperature (see also Fig. 7).

707

708

709

710711

712

713

The error-weighted results separated by mass spectrometer type and design of associated preparation lines, another postulated source of inter-lab disagreement (Swart et al., 2019), are shown in Figure 8 and Table 4. Out of the 22 participating laboratories, 8 use the Nu Perspective, 16 use versions of the Thermo MAT253, and one uses an Isoprime 100. Most results are statistically indistinguishable across instruments. Only IAEA-C2 yielded a significantly ($\geq 2\sigma$) higher mean Δ_{47} value ($\Delta\Delta_{47}$ of +0.0110 and +0.0081 % vs Nu perspective and MAT253, respectively) when analyzed with an Isoprime 100 versus other instrument models (but note that all of the Isoprime 100 data comes from a single laboratory). Inter-instrument differences averaged over all four samples (bottom row of Table 4) remain, however, indistinguishable from zero. Thus, any potential biases introduced by the use of different mass spectrometer models and/or the design of the preparation line which could cause partial equilibration of the produced CO₂ with the acid and/or heated metal surfaces (Swart et al.2019) are undetectable when using carbonate standardization. Sample sizes used for individual measurements ranged from 90-120 μg for the Kiel IV to ~500 μg for the NuCarb individual acid vial preparation systems, and to 3-12 mg for samples reacted in common acid bath custom-built extraction lines. The fact that small sample measurements are carried out at 70 °C and large ones at 90 °C, also suggests that there is no significant effect of sample sizes and variations in sample to acid ratios in these results.

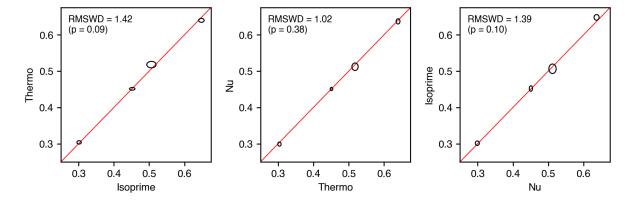


Figure 8. Error-weighted average $\Delta_{47(I-CDES)}$ values of unknowns obtained using different mass spectrometer types. Solid black ellipses correspond to 95 % confidence limits.

	MAT 253 vs Isoprime 100	Nu Perspective vs MAT 253	Isoprime 100 vs Nu Perspective
ETH-4	-0.0009 ± 0.0035	-0.0004 ± 0.0024	0.0013 ± 0.0036
IAEA-C1	0.0023 ± 0.0032	-0.0048 ± 0.0030	0.0025 ± 0.0035
IAEA-C2	-0.0081 ± 0.0039	-0.0029 ± 0.0037	0.0110 ± 0.0043
MERCK	0.0115 ± 0.0065	-0.0059 ± 0.0056	-0.0056 ± 0.0068
average (all samples)	0.0012 ± 0.0022	-0.0035 ± 0.0019	0.0023 ± 0.0024

Table 4. Error-weighted average $\Delta_{47(i\text{-}CDES)}$ values (±1SE) for each unknown as a function of mass spectrometer types.

3.4. Guidelines for improving data quality in the future.

The results of this inter-laboratory exercise support the use of carbonate standardization for clumped isotope measurements. When considering all laboratories, the standard deviation of laboratory averages for the 4 unknowns range from 0.011 ‰ to 0.018 ‰ for the samples requiring no to moderate extrapolation, to 0.025 ‰ for MERCK, an extreme case of extrapolation. The spread across laboratories is still relatively large, and of the same magnitude as those obtained by HG-EG normalization if we consider either (1) what has been reported on 4 carbonate standards for 4 laboratories in Dennis et al. (2011) or (2) the scatter in the values reported by the 10 laboratories that provided data for the re-determination of the accepted values of ETH-1 to ETH-3 in this study. However, based on this study, we can clearly state that the large scatter is dominated by random errors, and there are strategies to improve the repeatability within each laboratory.

The scatter in the InterCarb data (Fig. 4) is especially influenced by laboratories that show the largest errors in the individual sample reproducibility and have a significantly magnified normalization error induced by a small number of replicates of anchors. This observation underscores the necessity of increased replication, ideally in separate sessions, and of the measurement of a sufficient number of standards to produce data of the quality that is required for meaningful interpretations. Special care needs to be taken when reconstructing small temperature changes, such as for the reconstruction of climate change and/or the study of high

739

740

741

742

743

744

745

746

747

748

749

750

751

752

753

754

755

756

757

758

759

760

761

762

763

764

765

766

767

768

manuscript submitted to Geochemistry, Geophysics, Geosystems

temperature processes where the small temperature sensitivity of the clumped isotope thermometer requires highest analytical precision. We can, for example, define a target laboratory internal repeatability of 0.010 % (1SE) as a desirable goal with the currently available analytical systems. This arbitrary target number is within the shot-noise limits of modern IRMS instruments, and is necessary for applications in paleoclimate reconstructions, one of the main applications of clumped isotopes, as it corresponds to an uncertainty of approximately ± 3 °C. If for each sample we select only the results from the laboratories that have provided data with a combined error of less than 0.010 \% (1SE; Table 2), inter-laboratory standard deviation (1SD) becomes \leq 10 ppm for ETH-4 (N=22), IAEA-C1(N=15) and IAEA-C2 (N=13) and ≤ 0.015 % for MERCK (N=11; with SE< 0.0135). It can be noted that this does not significantly change the average value of the reference materials. Based on this example, it appears that with modern instrumentation from all manufacturers and with both custom built and commercially available systems used by the laboratories involved in this study, it is possible to reach this data quality and inter-laboratory consistency. The main factor to be taken into consideration is sufficient replication of both sample and standards (see Daëron, this volume, Bonifacie et al., 2017; Fernandez et al., 2017; Kocken et al., 2019). With standardization using carbonates, the principle of identical treatment of samples and standards (Werner & Brand, 2001b) is fulfilled, in contrast to the use of HG/EG standardization. Two important outcomes of this study are that acid reaction temperature and instrument and preparation line design are not a cause for differences among laboratories when standardization is based on carbonates. The lack of resolvable differences observed in our dataset, indicates that if preparation line differences affect Δ_{47} measurements, standardization with carbonates corrects any such effects whereas HG-EG standardization may fail to do so (Swart et al. 2019). Thus, when using carbonate standardization, these factors can be ignored provided the carbonate standards cover a large range in Δ_{47} . As the acid digestion conditions (e.g., reaction times, temperatures) vary with the carbonate mineralogy, possible effects on Δ_{47} might be mineral-specific. For this reason, it would be highly desirable to produce reference materials for other minerals such as dolomite, aragonite,

magnesite and siderite. For dolomite, three samples were proposed by Müller et al. (2019) as possible reference materials and are available upon request.

Standardization errors could be reduced to some extent by extending the range of bulk composition of the anchor samples (e.g., as illustrated by Fig. 1 of *Daëron*, this volume), especially when samples are measured that require significant extrapolation. A sample with an extreme bulk composition like MERCK would be a useful addition as an anchor, regardless of its Δ_{47} value. While with a three-anchor system, two heated standards for normalization are not strictly necessary, a "heated MERCK" anchor in combination with ETH-1 would furthermore allow verifying the PBL correction with greater confidence and with less replication than with ETH-2 (keeping in mind that small quadratic components to PBL correction might introduce a significant bias over a δ^{47} range of 60 ‰, e.g., Fig. 7 from He et al., [2012]).

3.5 Reporting data normalized to carbonates: definition of the "InterCarb" Carbon

Dioxide Equilibrium Scale (I-CDES)

There is a need for a community consensus on how to report clumped isotope measurements, both to promote data comparability and to reduce confusion stemming from the different scales used in the literature. Currently Δ_{47} data are reported for different temperatures of phosphoric acid digestion, mostly projected to 25 °C but also to 70 °C or 90 °C reactions, and generally labeled respectively as $\Delta_{47\text{CDES}25}$, $\Delta_{47\text{CDES}70}$ and $\Delta_{47\text{CDES}90}$, a terminology introduced in Bonifacie et al. (2017). In the literature, phosphoric acid correction factors used by different research groups to convert results from 90 to 25 °C reaction temperatures have varied between 0.069 % (Wacker et al., 2014) to 0.092 % (e.g., Bonifacie et al., 2017), thus representing a significant source of uncertainty and confusion.

The direct standardization to accepted values of solid phases, on the other hand, removes the need for a phosphoric acid correction, yielding results which are independent of the temperature at which the samples were reacted. In Intercarb this has only been tested for calcites, the

manuscript submitted to Geochemistry, Geophysics, Geosystems

mineralogy of all standards used here. Further studies are necessary in particular for dolomite and siderite, as these minerals require longer reaction times and there are contrasting findings in the literature on whether they require different phosphoric acid fractionations (Bonifacie et al., 2017; Müller et al. 2019, van Dijk et al., 2019; Petersen et al. 2019), possibly affecting the absolute values of samples with such mineralogies. Nevertheless, the use of the InterCarb reference carbonates, allows for a consistent correction of instrumental effects and normalization to the I-CDES, independent of the mineralogy. Possible effects related to the longer reaction time remain to be tested, and require intercomparison samples of different mineralogy. In any case for all calcites carbonate standardization eliminates different phosphoric acid correction factors as a source of uncertainty and provides a consistent framework to report Δ_{47} without uncertainties related to the reaction temperature.

For these reasons, we recommend that in the future, carbonate clumped isotope values should be reported relative to a carbonate reference frame uniquely defined by the absolute Δ_{47} values reported in Table 1 for ETH-1, ETH-2, and ETH-3. Unknown samples may be anchored to this reference frame either (a) directly by comparison to ETH-1/2/3, (b) indirectly by comparison with several of the four other carbonate standards used here (Figure 4), or (c) by comparison with a set of in-house laboratory standards whose composition is well-constrained relative to the materials reported here. To clearly distinguish this data normalization scheme from previous ones the denomination I-CDES (InterCarb - Carbon Dioxide Equilibration Scale) should be used, and, where applicable, the notation $\Delta_{47(I-CDES)}$. This approach is analogous to the change from the PDB to the VPDB scale, which was accomplished by assigning a consensus offset of +1.95 ‰ between the original PDB reference material and the NBS19 carbonate. This was subsequently, albeit temporarily, improved by defining a second anchor point with the L-SVEC Lithium carbonate standard (Coplen et al., 2006). We note that because the carbonate Δ_{47} values in Table 1 are firmly anchored to the CDES scale via HG/EG measurements in multiple laboratories, the two scales are in principle equivalent. However, I-CDES has three major advantages (1) it follows the principle of equal treatment of sample and standards (2)it removes uncertainties related to fractionation effects due to different acid reaction temperatures and designs of the preparation lines and (3) it is based on traceable, stable materials (calcium carbonates) that are widely available to interested laboratories. . What is more, the results

manuscript submitted to Geochemistry, Geophysics, Geosystems

summarized in Fig. 5 imply that I-CDES standardization yields consistent Δ_{47} values independent of laboratory and/or analytical protocols, so that inter-laboratory scatter is accurately predicted by the fully-propagated analytical uncertainties computed within each laboratory. By contrast, this does not always seem to be the case for gas-based standardization to the CDES reference frame (Fig. 6), suggesting the existence of poorly understood sources of inter-lab discrepancies in that approach. The broad availability of carbonate reference materials with widely varying bulk and clumped isotope compositions is an important step to help establish a worldwide equivalence among laboratories, and to help new laboratories establish and verify their analytical procedures.

The decision that $\Delta_{47\text{(I-CDES)}}$ values are reported for an acid reaction temperature of 90 °C implies that the I-CDES value of any given sample will be almost 0.1 % lower than $\Delta_{47\text{CDES25}}$ values

which is currently the most common, but not the sole, convention used to report clumped

isotopes. At first sight it may seem a disadvantage to lose the ability to intuitively compare new results to those obtained through different standardization approaches in the published literature.

However, it will make it immediately obvious that I-CDES-normalized values cannot directly be

compared to data standardized to the ETH-1/2/3/4 values reported by Bernasconi et al. (2018).

To directly compare previous carbonate-normalized Δ_{47} values and the already published calibration equations produced with ETH standard normalisation ((Kele et al., 2015; Bernasconi et al., 2018; Jautzy et al., 2020; Meinicke et al., 2020; Peral et al., 2018; Piasecki et al., 2019) to future data reported on the I-CDES scale, previous data have to be recalculated as described in

detail in the Appendix. We emphasize again, however, that the previously reported temperatures calculated from the original calibrations using the original ETH standard values are directly

comparable to temperatures calculated for samples normalized and calibrations recalculated to

the I-CDES. This is because only the nominal values of the standards have changed, and thus the

data normalization is internally consistent and traceable to the same solid standards.

In principle, I-CDES data are directly comparable to data produced by phosphoric acid reaction at 90°C using the HG/EG approach. With InterCarb, the nominal values of the ETH standards are robustly linked to the CDES, as they are now based on the average results of 10 laboratories, and are not only based on the values determined at ETH in 2013. Some laboratories may still

want to continue measuring HG and EG to keep established laboratory procedures and/or to cover ranges in bulk compositions that require large extrapolations. However, the results of InterCarb, and the discussions in the literature (e.g. Petersen et al. 2019) show that with the HG/EG approach there are still poorly understood interlaboratory discrepancies (as suggested by Fig. 6) which are absent in the carbonate-based normalization (see Fig 5). For this reason, it is of paramount importance that several of the InterCarb reference carbonates are incorporated in the laboratory procedures to ensure interlaboratory data compatibility. The use of matrix-matched reference materials is necessary so that delta values can be unambiguously compared on a likefor-like basis (see Meier-Augenstein and Schimmelmann, 2019 for a recent discussion). For laboratories using large sample common-acid bath methodologies and preferring gas-based data correction, we recommend that in addition to HG/EG a minimum of two of the InterCarb reference materials (or in-house standards with values calibrated to ETH standards) should be measured within the same analytical sessions as the samples and used in the data correction scheme. We recommend choosing two standards with a large difference in Δ_{47} , for example ETH-1 and ETH-3 or ETH-2 and IAEA-C2, depending on the bulk composition of the unknown samples (see Fig. 1). In the case of samples very low δ^{47} we additionally recommend the use of MERCK. The comparison of I-CDES data to older HG/EG normalized data, is in principle possible, with the caveat that in absence of measurements of widely available carbonates, a direct comparison remains uncertain. A community effort to robustly anchor the composition of legacy standards measured in the first laboratories, may alleviate this problem in the near future. For laboratories with commercially available single acid vial, small sample preparation systems we discourage the use of HG/EG, as the gases would most probably be measured under different conditions than the samples, an approach which is prone to error. In addition, InterCarb shows that carbonate-based standardization can provide robust and accurate data without the use of gases.

3.6. Outlook for future improvements

860

861

862

863

864

865

866

867

868

869

870

871

872

873

874

875

876

877

878

879

880

881

882

883

884

885

886

887

888

889

InterCarb has shown that with carbonate standardization it is possible to reach an inter-laboratory reproducibility that is as good as what can be reached within a single laboratory. The use of

carbonate standards allows for a better monitoring of the performance of the entire preparation system, including acid reaction and sample purification. Regular and systematic measurement of carbonate reference material distributed within a run or measurement interval is a prerequisite to correct for short- and long-term variations in the analytical systems. In this study, analyses were grouped in measurement intervals, and all data were processed assuming no short-term variation in the instrumentation. However, especially with "small sample approaches" relying on short (~30-45 min) measurements of many replicates, one can observe short-term variations (Bernasconi et al. 2018, Fig. 4). Thus a moving window correction with variable window size may be desirable in these cases.

Replicate sample measurements should be spread in time as much as possible. The number and distribution of standards in a run are also important parameters that can improve reproducibility and reduce errors. This has been discussed in detail by Kocken et al. (2019) who concluded that carbonate standards with bulk and clumped-isotope compositions similar to those of unknowns should be analyzed with greater frequency than the other anchors, while preserving a minimal level of replication for each anchor.

For InterCarb, all distributed aliquots of IAEA-C1, IAEA-C2 and MERCK originated from single bottles. The results suggest that after milling, these carbonates were homogeneous within these bottles, but we strongly recommend verifying that additional bottles purchased from IAEA and MERCK are identical to the ones tested here. Finally, as there is currently no consensus in the literature on whether different carbonate mineralogies exhibit the same phosphoric acid fractionation effects (e.g. Müller et al. 2017; van Djik et al. 2019; Bonifacie et al. 2017, Petersen et al. 2019) and on possible differences in the temperature-dependence of the phosphoric acid fractionation for dolomite (Defliese et al. 2015; Murray et al. 2016), there should be a community effort to identify and characterize reference materials for other carbonates such as aragonite and siderite, similar to the dolomite reference materials proposed by Müller et al. (2019).

4 Conclusions

920 •	This study demonstrates that carbonate-based standardization of clumped isotope
921	measurements solves many open questions so far limiting the application of
922	carbonate clumped isotope thermometry as a mature and reliable tool in Earth
923	sciences.
924 •	Carbonate-based standardization is robust in spite of the smaller range of isotopic
925	compositions compared to the heated/equilibrated gas approach, even when
926	extrapolation is necessary
927 •	Inter-laboratory discrepancies observed in this study among 22 laboratories are
928	not greater than those predicted from intra-laboratory analytical uncertainties.
929	We propose a set of 7 widely available calcite reference materials for
930	normalization of carbonate clumped isotope measurements with new community-
931	accepted values: the 4 ETH standards, two samples distributed by the IAEA (C1
932	and C2), and a synthetic carbonate produced by MERCK.
933	Standardization with ETH-1, ETH-2 and ETH-3 provides a robust framework for
934	converting carbonate clumped isotope data to the CDES using reference materials
935	matching the sample matrix and, thus, following the principle of equal treatment
936	of samples and standards.
937 •	Data standardized with this method should be reported as I-CDES to explicitly
938	signify that the data were normalized using these reference materials or a
939	combination of HG/EG and at least two of the reference carbonates.
•	Carbonate standardization removes the need to apply an acid digestion
941	fractionation factor, eliminating uncertainties due to poorly known acid
942	fractionation factors and different preparation systems and thus reduces
943	differences between laboratories.
•	The use of a reference material with an extreme bulk isotope composition such as
945	MERCK is recommended, as the large range of compositions reduces
946	standardization errors for samples with "exotic" bulk compositions (as previously
947	suggested by some authors but precisely quantified in Daëron, this volume).
948 •	Data expressed in the I-CDES are directly comparable to samples reacted at a
949	temperature of 90°C normalized to the CDES with HG/EG. However, we
950	emphasize that carbonate standardisation is preferred to pure HG/EG

normalization because it is based on traceable carbonate samples that can be measured in every laboratory. Reporting the measured compositions of carbonate reference materials together with the samples is the only way to ensure interlaboratory consistency.

- Robust standardization of clumped isotope measurements requires the analysis of
 a sufficient number of replicates of both samples and standard materials (either
 gases or carbonates) alongside unknowns to minimize error and obtain accurate
 measurements within a single laboratory.
- Replicate analyses should be distributed in time and ideally span different correction intervals
- When reporting data, error should include both replication error and reference frame error in proper statistical data treatment and reporting of uncertainties

Appendix:

A. Converting older, carbonate-anchored Δ_{47} values to the I-CDES

This section describes the steps necessary to convert existing Δ_{47} measurements to the I-CDES, provided that they were either standardized using carbonate anchors or analyzed simultaneously with several carbonate anchors. This mathematically exact approach is a simpler alternative to fully reprocessing the original raw data (with the caveat that this conversion will not provide, by itself, any error estimates).

By way of example, let us consider measurements originally standardized using ETH-1/2/3 with the nominal Δ_{47} values reported by Bernasconi et al. (2018). The "old" Δ_{47} values of these measurements are noted $^{\text{old}}\Delta_{47}$, and we wish to compute the "new" Δ_{47} values, noted $^{\text{new}}\Delta_{47}$, that would be obtained if the same data were standardized to the I-CDES.

Both old and new Δ_{47} values are derived from the same set of raw measurements using affine relationships of the form:

$$\Delta_{47}^{\text{old}} = x_1 + y_1 \delta^{47} + z_1 \Delta_{47}^{\text{raw}}$$

981

982 (A.2)
$$\Delta_{47}^{\text{new}} = x_2 + y_2 \delta^{47} + z_2 \Delta_{47}^{\text{raw}}$$

983

987

988

1000

1001

1002

1003

1004

1005

1006

1007

1008

984 We can rearrange the above equations to express Δ_{47}^{new} as function of δ^{47} and Δ_{47}^{old} :

becomes:

985 a
$$\Delta_{47}^{\text{new}} = a + b\delta^{47} + c\Delta_{47}^{\text{old}}$$

986

986 (A.3)
$$\Delta_{47}^{\text{new}} = a + b\delta^{47} + c\Delta_{47}^{\text{old}}$$

- Computing the numerical values of (a,b,c) is thus all that is required to compute Δ_{47}^{new} for any 989 sample whose $(\delta^{47}, \Delta_{47}^{\text{old}})$ values are known. This computation only requires knowing $(\delta^{47}, \Delta_{47}^{\text{old}})$ 990 , Δ_{47}^{new}) for three different anchors, for instance ETH-1/2/3. In matrix form, eq. (A.3) then 991
- 993 994 995

996
997 (A.5)
$$\begin{pmatrix} a \\ b \\ c \end{pmatrix} = \begin{pmatrix} 1 & \delta_{(\text{ETH1})}^{47} & \Delta_{47(\text{ETH1})}^{\text{old}} \\ 1 & \delta_{(\text{ETH2})}^{47} & \Delta_{47(\text{ETH2})}^{\text{old}} \\ 1 & \delta_{47(\text{ETH2})}^{47} & \Delta_{47(\text{ETH3})}^{\text{old}} \end{pmatrix}^{-1} \begin{pmatrix} \Delta_{47(\text{ETH1})}^{\text{new}} \\ \Delta_{47(\text{ETH2})}^{\text{new}} \\ \Delta_{47(\text{ETH3})}^{\text{new}} \end{pmatrix}$$

In this example, the Δ_{47}^{old} and Δ_{47}^{new} values are the old and new nominal values of ETH-1/2/3. If the δ^{47} values are defined as usual by reference to a working gas, (a,b,c) will vary when different working gases are used. However, the above equations remain valid if δ^{47} is defined instead by reference to a fixed, hypothetical CO_2 composition, e.g., stochastic VPDB- CO_2 ($\delta^{13}C_{VPDB} = 0$ %; $\delta^{18}O_{VSMOW} \approx 41.5$ %; $\Delta_{47} = 0$ %, $R^{47} = 4.834 \times 10 - 5$). In that case, numerical values of the parameters (a,b,c) can be determined once and then used to compute Δ_{47}^{new} for any unknown sample based only on its Δ_{47}^{old} and δ^{47} values (the latter being defined relative to VPDB-CO₂ and computed by applying an acid ¹⁸O/¹⁶O fractionation factor of 1.01025 to said sample).

1009	Importantly, this particular transformation applies to all data sets originally standardized in the
1010	reference frame defined by the "old" nominal values of ETH-1/2/3.
1011	
1012	For instance, the relationship linking the ETH-1/2/3 reference frame of Bernasconi et al. (2018)
1013	to the I-CDES is defined by:
1014	

1016 (A.6)

1015

1018

1020

1022

1023

1024

1025

1032

1034

1039

$$\begin{pmatrix} a \\ b \\ c \end{pmatrix} = \begin{pmatrix} 1 & 0.010 & 0.258 \\ 1 & -28.375 & 0.258 \\ 1 & 0.538 & 0.691 \end{pmatrix}^{-1} \begin{pmatrix} 0.2052 \\ 0.2085 \\ 0.6132 \end{pmatrix} = \begin{pmatrix} -0.037997 \\ -0.000182 \\ 0.942590 \end{pmatrix}$$

- 1019 Thus:
- 1021 (A.7) $\Delta_{47}^{\text{new}} = -0.037997 0.000182 \delta^{47} + 0.942590 \Delta_{47}^{\text{old}}$

- In this case, it is clear that the conversion from the old reference frame to the new one is not very sensitive to δ^{47} values: for unknown samples with δ^{47} values within ± 6 % of ETH-1 (or stochastic VPDB-CO₂), the effect of the second term in Eq. (A.7) is less than 1 ppm, and less
- than 3 ppm for unknowns within ± 18 ‰ of ETH-1. In such cases, the conversion may be
- simplified as a simpler affine transformation, akin to a more traditional two-point normalization
- 1031 (e.g., VSMOW-VSLAP standardization):
- 1033 (A.8) $\Delta_{47}^{\text{new}} = 0.942590 \ \Delta_{47}^{\text{old}} 0.037997$
- For instance, to convert the Δ_{47}^{old} value of ETH-4 reported by Bernasconi et al. (2018) to the I-
- 1036 CDES, we only need to know that $\delta^{47}_{(ETH4)} = -28.8$ % and $\Delta_{47}^{\text{old}} = 0.507 \pm 0.004$ %. The Δ_{47}^{new}
- value predicted by eq. (A.7) is then 0.445 ± 0.004 %, to be compared with the independently
- 1038 constrained values reported here in Table 1 (0.450 \pm 0.002 %) and Table 2 (0.451 \pm 0.001 %).
- 1040 The above computation could also be performed using any arbitrary set of three carbonate
- materials whose $(\delta^{47}, \Delta_{47}^{\text{old}}, \Delta_{47}^{\text{new}})$ values are known, provided that they span a wide enough
- range in δ^{47} and Δ_{47} . This is true even if the carbonates in question were not originally used to
- standardize the raw data, as would be the case for CO_2 -standardized measurements. In the case
- 1044 where only two suitable carbonate standards X and Y are available, an acceptable approach

would be to neglect effects (equivalent to setting the value of b to zero in eq. (A.3)) by solving the following equation:

1048 (A.9)
1049
$$\begin{pmatrix} a \\ c \end{pmatrix} = \begin{pmatrix} 1 & \Delta_{47(X)}^{\text{old}} \\ 1 & \Delta_{47(Y)}^{\text{old}} \end{pmatrix}^{-1} \begin{pmatrix} \Delta_{47(X)}^{\text{new}} \\ \Delta_{47(Y)}^{\text{new}} \end{pmatrix}$$
1050

1045

1046

1047

1051

1052

1053

1054

1055

1056

1057

1058

1059

1060

1061

1062

1063

1064

1065

1066

1067

1068

1069

1070

1071

1072

1073

1074

Acknowledgments, Samples, and Data

SMB acknowledges instrumentation funding from ETH Zürich and support from Swiss National Science Foundation grants 200021_143485, 200020_160046. MD acknowledges the clumpedisotope facility at LSCE is part of PANOPLY (Plateforme Analytique Géosciences Paris-Saclay) and was supported by the following institutions: Région Ile-de-France; Direction des Sciences de la Matière du Commissariat à l'Energie Atomique; Institut National des Sciences de l'Univers, Centre National de la Recherche Scientifique; Universtité de Versailles/Saint-Quentin-en-Yvelines. KDB and the MIT carbonate clumped isotope facility acknowledge support from the Packard Foundation, the Agouron Foundation and NASA Exobiology Grant 80NSSC19K0464. MB acknowledges the program Emergences Ville de Paris for early funding that allowed building her clumped isotope laboratory at Institut de Physique du Globe de Paris. MB also thanks IPGP and Université de Paris for support for the organization of the VIth Clumped Isotope Workshop in Paris in 2017. Funding was provided to CJ and the Imperial College Clumped Isotope Lab by Qatar Petroleum, Shell, and the Qatar Science and Technology Panel. Funding was provided to KH at University of Washington from NSF EAR grants 1933130, 1713275, and 1156134. TV acknowledges the support of the Swiss National Science Foundation grant 206021-164032. ANM acknowledges funding from the Trond Mohn Foundation and the European Research Council (ERC) under the European Union's Horizon 2020 research and innovation programme (grant agreement no. 638467). Funding was provided to the Tripati Lab at UCLA for these analyses from DOE BES grant DE-FG02-13ER16402. Funding was provided to the Hyland Lab at NCSU for these analyses from NSF EAR-FRES grant 1925973. Funding was provided to the Yoshida Lab at Tokyo Tech. for these analyses from JSPS grant JP17H06105. Funding was provided to the CLIMB Lab at the University of Bergen for these

1075 analyses by the European Research Council (ERC) under the European Union's Horizon 2020 1076 research and innovation programme (grant agreement No 638467) and by the Trond Mohn 1077 Foundation. Funding was provided to the UU Clumped Lab at Utrecht University by the 1078 Netherlands Research Organization (NWO) through VIDI grant 016.161.365. Funding was 1079 provided to the CUBES-SIL at CU Boulder by startup funds from the University of Colorado 1080 Boulder and NSF EAR grant 1524785. The UM SCIPP Lab was supported by startup funds from 1081 the University of Michigan. JR Kelson was supported by NSF PRF grant 1854873. 1082 Data 1083 All data analyzed as part of this study, along with all code used to process them, are available at 1084 https://github.com/mdaeron/InterCarb and http://doi.org/10.5281/zenodo.4314449. 1085 1086 **Author contributions** 1087 The community designed the study at the Paris 2017 International Clumped Isotope Workshop. 1088 S.M.B. prepared materials, gathered results and wrote the first draft of the manuscript. M.D. 1089 created all code and conducted all analysis of the results. S.M.B., M.D., K.D.B., N.M., and M.B. 1090 organized the effort and contributed to early drafts. All authors contributed to the analyses and curation of data shared in this study. All authors contributed to the final version of the 1091 1092 manuscript. 1093 1094 References 1095 1096 Adlan, Q., Davies, A.J., John, C.M., 2020. Effects of oxygen plasma ashing treatment on 1097 carbonate clumped isotopes, Rapid Communications in Mass Spectrometry, DOI: 1098 https://doi.org/10.1002/rcm.8802 1099 Affek, H. P., & Eiler, J. M. (2006). Abundance of mass 47 CO2 in urban air, car exhaust, 1100 and human breath. Geochimica et Cosmochimica Acta, 70(1), 1–12. 1101 https://doi.org/10.1016/j.gca.2005.08.021 1102 Anderson N.T., Kelson J.R., Kele S., Daëron M., Bonifacie M., Mackey T.J., John C.M., 1103 Huntington K.W., Bernasconi S.M. and Bergman K.D. (2020) Reanalyzing carbonate

1104	clumped isotope calibration materials using robust carbonate standardization. AGU
1105	Fall Meeting 2020 Abstracts
1106	Bernasconi, S. M., Hu, B., Wacker, U., Fiebig, J., Breitenbach, S. F. M., & Rutz, T. (2013).
1107	Background effects on Faraday collectors in gas-source mass spectrometry and
1108	implications for clumped isotope measurements. Rapid Communications in Mass
1109	Spectrometry, 27(5), 603-612. https://doi.org/10.1002/rcm.6490
1110	Bernasconi, S. M., Müller, I. A., Bergmann, K. D., Breitenbach, S. F. M. M., Fernandez, A.,
1111	Hodell, D. A., et al. (2018). Reducing uncertainties in carbonate clumped isotope
1112	analysis through consistent carbonate-based standardization. Geochemistry,
1113	Geophysics, Geosystems, 19(9), 1-20. https://doi.org/10.1029/2017GC007385
1114	Bonifacie, M., Calmels, D., Eiler, J. M., Horita, J., Chaduteau, C., Vasconcelos, C., et al.
1115	(2017). Calibration of the dolomite clumped isotope thermometer from 25 to 350 $^{\circ}$ C,
1116	and implications for a universal calibration for all (Ca, Mg, Fe)CO3 carbonates.
1117	Geochimica et Cosmochimica Acta, 200, 255–279.
1118	https://doi.org/10.1016/j.gca.2016.11.028
1119	Brand, W. a., Assonov, S. S., & Coplen, T. B. (2010). Correction for the 17O interference in
1120	$\delta(13C)$ measurements when analyzing CO2 with stable isotope mass spectrometry
1121	(IUPAC Technical Report). Pure and Applied Chemistry, 82(8), 1719-1733.
1122	https://doi.org/10.1351/PAC-REP-09-01-05
1123	Came, R. E., Brand, U., & Affek, H. P. (2014). Clumped isotope signatures in modern
1124	brachiopod carbonate. Chemical Geology, 377, 20-30.
1125	https://doi.org/10.1016/j.chemgeo.2014.04.004
1126	Carter, J. F., & Fry, B. (2013). Ensuring the reliability of stable isotope ratio data - Beyond
1127	the principle of identical treatment. Analytical and Bioanalytical Chemistry, 405(9),
1128	2799–2814. https://doi.org/10.1007/s00216-012-6551-0
1129	Daëron, M. (submitted) Full propagation of analytical uncertainties in Δ_{47} measurements.
1130	Submitted to Geochemistry, Geophysics, Geosystems (Companion Paper)
1131	Daëron, M., Blamart, D., Peral, M., & Affek, H. P. (2016). Absolute isotopic abundance
1132	ratios and the accuracy of Δ_{47} measurements. Chemical Geology, 442, 83–96.
1133	https://doi.org/10.1016/j.chemgeo.2016.08.014

1134	Dale, A., John, C. M., Mozley, P. S., Smalley, P. C., & Muggeridge, A. H. (2014). Time-
1135	capsule concretions: Unlocking burial diagenetic processes in the Mancos Shale using
1136	carbonate clumped isotopes. Earth and Planetary Science Letters, 394.
1137	https://doi.org/10.1016/j.epsl.2014.03.004
1138	Defliese, W. F., & Lohmann, K. C. (2015). Non-linear mixing effects on mass-47
1139	CO <inf>2</inf> clumped isotope thermometry: Patterns and implications. Rapid
1140	Communications in Mass Spectrometry, 29(9), 901–909.
1141	https://doi.org/10.1002/rcm.7175
1142	Defliese, W. F., Hren, M. T., & Lohmann, K. C. (2015). Compositional and temperature
1143	effects of phosphoric acid fractionation on D47 analysis and implications for
1144	discrepant calibrations. Chemical Geology, 396, 51-60.
1145	https://doi.org/10.1016/j.chemgeo.2014.12.018
1146	Dennis, K. J., Affek, H. P., Passey, B. H., Schrag, D. P., & Eiler, J. M. (2011). Defining an
1147	absolute reference frame for "clumped" isotope studies of CO2. Geochimica et
1148	Cosmochimica Acta, 75(22), 7117–7131. https://doi.org/10.1016/j.gca.2011.09.025
1149	Eagle, R. a, Schauble, E. a, Tripati, A. K., Tütken, T., Hulbert, R. C., & Eiler, J. M. (2010).
1150	Body temperatures of modern and extinct vertebrates from (13)C-(18)O bond
1151	abundances in bioapatite. Proceedings of the National Academy of Sciences of the
1152	United States of America, 107(23), 10377–10382.
1153	https://doi.org/10.1073/pnas.0911115107
1154	Fernandez, A., Tang, J., & Rosenheim, B. E. (2014). Siderite 'clumped' isotope
1155	thermometry: A new paleoclimate proxy for humid continental environments.
1156	Geochimica et Cosmochimica Acta, 126, 411–421.
1157	https://doi.org/10.1016/j.gca.2013.11.006
1158	Fernandez, A., Müller, I. A., Rodríguez-Sanz, L., van Dijk, J., Looser, N., & Bernasconi, S.
1159	M. (2017). A Reassessment of the Precision of Carbonate Clumped Isotope
1160	Measurements: Implications for Calibrations and Paleoclimate Reconstructions.
1161	Geochemistry, Geophysics, Geosystems, 18(12).
1162	https://doi.org/10.1002/2017GC007106

1163	Ferry, J. M., Passey, B. H., Vasconcelos, C., & Eiler, J. M. (2011). Formation of dolomite at
1164	40-80 ??C in the Latemar carbonate buildup, Dolomites, Italy, from clumped isotope
1165	thermometry. Geology, 39(6), 571-574. https://doi.org/10.1130/G31845.1
1166	Fiebig, J., Bajnai, D., Löffler, N., Methner, K., Krsnik, E., Mulch, A., & Hofmann, S.
1167	(2019). Combined high-precision Δ_{48} and Δ_{47} analysis of carbonates. <i>Chemical</i>
1168	Geology, 522 (Feb), 186–191. https://doi.org/10.1016/j.chemgeo.2019.05.019
1169	Ghosh, P., Adkins, J., Affek, H., Balta, B., Guo, W., Schauble, E. A., et al. (2006). 13C-
1170	18O bonds in carbonate minerals: A new kind of paleothermometer. Geochimica et
1171	Cosmochimica Acta, 70(6), 1439–1456. https://doi.org/10.1016/j.gca.2005.11.014
1172	Ghosh, P., Garzione, C., & Eiler, J. (2006). Rapid Uplift of the Alitplano Revealed Through
1173	13C-18O Bonds in Paleosol Carbonates. Science, 311(January), 511–515.
1174	https://doi.org/10.1126/science.1119365
1175	Grauel, A. L., Schmid, T. W., Hu, B., Bergami, C., Capotondi, L., Zhou, L., & Bernasconi,
1176	S. M. (2013). Calibration and application of the "clumped isotope" thermometer to
1177	foraminifera for high-resolution climate reconstructions. Geochimica et Cosmochimica
1178	Acta, 108, 125–140. https://doi.org/10.1016/j.gca.2012.12.049
1179	Guo, W., & Eiler, J. M. (2007). Temperatures of aqueous alteration and evidence for
1180	methane generation on the parent bodies of the CM chondrites. Geochimica et
1181	Cosmochimica Acta, 71(22), 5565–5575. https://doi.org/10.1016/j.gca.2007.07.029
1182	Guo, W., Mosenfelder, J. L., Goddard, W. A., & Eiler, J. M. (2009). Isotopic fractionations
1183	associated with phosphoric acid digestion of carbonate minerals: Insights from first-
1184	principles theoretical modeling and clumped isotope measurements. Geochimica et
1185	Cosmochimica Acta, 73(24), 7203-7225. https://doi.org/10.1016/j.gca.2009.05.071
1186	He, B., Olack, G. A., & Colman, A. S. (2012). Pressure baseline correction and high-
1187	precision CO2 clumped-isotope ($\Delta47$) measurements in bellows and micro-volume
1188	modes. Rapid Communications in Mass Spectrometry, 26(24), 2837–2853.
1189	https://doi.org/10.1002/rcm.6436
1190	Hu, B., Radke, J., Schlüter, H. J., Heine, F. T., Zhou, L., Bernasconi, S. M., et al. (2014). A
1191	modified procedure for gas-source isotope ratio mass spectrometry: The long-
1192	integration dual-inlet (LIDI) methodology and implications for clumped isotope

1193	measurements. Rapid Communications in Mass Spectrometry, 28(13), 1413-1425.
1194	https://doi.org/10.1002/rcm.6909
1195	Huntington, K. W., Eiler, J. M., Affek, H. P., Guo, W., Bonifacie, M., Yeung, L. Y., et al.
1196	(2009). Methods and limitations of "clumped" CO2 isotope ($\Delta47$) analysis by gas-
1197	source isotope ratiomass spectrometry. Journal of Mass Spectrometry, 44(9), 1318-
1198	1329. https://doi.org/10.1002/jms.1614
1199	Huntington, K. W., Budd, D. A., Wernicke, B. P., & Eiler, J. M. (2011). Use of Clumped-
1200	Isotope Thermometry To Constrain the Crystallization Temperature of Diagenetic
1201	Calcite. Journal of Sedimentary Research, 81(9), 656–669.
1202	https://doi.org/10.2110/jsr.2011.51
1203	Jautzy, J. J., Savard, M, Dhillon, R, Bernasconi, S, Lavoie, D., & Smirnoff, A. (2021).
1204	Clumped isotope temperature calibration for calcite: Bridging theory and
1205	experimentation. Geochemical Perspective Letters, (2020), 36-41.
1206	https://doi.org/10.7185/geochemlet.2021
1207	Katz A., Bonifacie M., Hermoso M., Cartigny P., Calmels D. (2017) Laboratory-grown
1208	coccoliths exhibit no vital effect in clumped isotope (Δ_{47}) composition on a range of
1209	geologically relevant temperatures in Geochimica Cosmochimica Acta, 208, 335–353.
1210	https://doi.org/10.1016/j.gca.2017.02.025
1211	Kele, S., Breitenbach, S. F. M., Capezzuoli, E., Meckler, A. N., Ziegler, M., Millan, I. M.,
1212	Kluge, T., Deák, J., Hanselmann, K., John, C. M., Yan, H., Liu, Z., & Bernasconi, S.
1213	M. (2015). Temperature dependence of oxygen- and clumped isotope fractionation in
1214	carbonates: A study of travertines and tufas in the 6-95°C temperature range.
1215	Geochimica et Cosmochimica Acta, 168. https://doi.org/10.1016/j.gca.2015.06.032
1216	Kelson, J. R., Huntington, K. W., Schauer, A. J., Saenger, C., & Lechler, A. R. (2017).
1217	Toward a universal carbonate clumped isotope calibration: Diverse synthesis and
1218	preparatory methods suggest a single temperature relationship. Geochimica et
1219	Cosmochimica Acta, 197, 104–131. https://doi.org/10.1016/j.gca.2016.10.010
1220	Kim, S. T., Coplen, T. B., & Horita, J. (2015). Normalization of stable isotope data for
1221	carbonate minerals: Implementation of IUPAC guidelines. Geochimica et
1222	Cosmochimica Acta, 158, 276–289. https://doi.org/10.1016/j.gca.2015.02.011

1223	Kluge, T., John, C. M., Jourdan, AL., Davis, S., & Crawshaw, J. (2015). Laboratory
1224	calibration of the calcium carbonate clumped isotope thermometer in the 25-250°C
1225	temperature range. Geochimica et Cosmochimica Acta, 157, 213-227.
1226	https://doi.org/10.1016/j.gca.2015.02.028
1227	Kocken, I. J., Müller, I. A., & Ziegler, M. (2019). Optimizing the Use of Carbonate
1228	Standards to Minimize Uncertainties in Clumped Isotope Data. Geochemistry,
1229	Geophysics, Geosystems, 20(11), 5565-5577. https://doi.org/10.1029/2019GC008545
1230	Mangenot X., Gerdes A., Gasparrini M., Bonifacie M., Rouchon V. (2018). An emerging
1231	thermo-chronometer for carbonate-bearing rocks : $\Delta_{47}/(\text{U-Pb})$. Geology 46 (12), 1067-
1232	1070.
1233	Massey, F. J. (1951). The Kolmogorov-Smirnov Test for Goodness of Fit. Journal of the
1234	American Statistical Association, 46(253), 68–78.
1235	https://doi.org/10.1080/01621459.1951.10500769
1236	Meckler, A. N., Ziegler, M., Millan, M. I., Breitenbach, S. F. M., Bernasconi, S. M., Millán,
1237	M. I., et al. (2014). Long-term performance of the Kiel carbonate device with a new
1238	correction scheme for clumped isotope measurements. Rapid Communications in Mass
1239	Spectrometry, 28(15), 1705–1715. https://doi.org/10.1002/rcm.6949
1240	Meier-Augenstein, W., & Schimmelmann, A. (2019). A guide for proper utilisation of
1241	stable isotope reference materials *. Isotopes in Environmental and Health Studies,
1242	55(2), 113–128. https://doi.org/10.1080/10256016.2018.1538137
1243	Meinicke, N., Ho, S. L., Hannisdal, B., Nürnberg, D., Tripati, A., Schiebel, R., & Meckler,
1244	A. N. (2020). A robust calibration of the clumped isotopes to temperature relationship
1245	for foraminifers. Geochimica et Cosmochimica Acta, 270, 160-183.
1246	https://doi.org/10.1016/j.gca.2019.11.022
1247	Müller, I. A., Fernandez, A., Radke, J., Dijk, J. Van, Bowen, D., Schwieters, J., et al.
1248	(2017). Carbonate clumped isotope analyses with the long-integration dual-inlet (LIDI)
1249	workflow: scratching at the lower sample weight boundaries. Rapid Communications
1250	in Mass Spectrometry, 31(12), 1057–1066. https://doi.org/10.1002/rcm.7878
1251	Müller, I. A., Violay, M. E. S., Storck, JC. C., Fernandez, A., van Dijk, J., Madonna, C., &
1252	Bernasconi, S. M. (2017). Clumped isotope fractionation during phosphoric acid

1253	digestion of carbonates at 70 °C. Chemical Geology, 449, 1–14.
1254	https://doi.org/10.1016/j.chemgeo.2016.11.030
1255	Müller, I. A., Rodriguez-Blanco, J. D., Storck, JC., do Nascimento, G. S., Bontognali, T.
1256	R. R., Vasconcelos, C., et al. (2019). Calibration of the oxygen and clumped isotope
1257	thermometers for (proto-)dolomite based on synthetic and natural carbonates.
1258	Chemical Geology, 525(December 2018), 1–17.
1259	https://doi.org/10.1016/j.chemgeo.2019.07.014
1260	Murray, S. T., Arienzo, M. M., & Swart, P. K. (2016). Determining the Δ47 acid
1261	fractionation in dolomites. Geochimica et Cosmochimica Acta, 174, 42-53.
1262	https://doi.org/10.1016/j.gca.2015.10.029
1263	Passey, B. H., & Henkes, G. A. (2012). Carbonate clumped isotope bond reordering and
1264	geospeedometry. Earth and Planetary Science Letters, 351-352, 223-236.
1265	https://doi.org/10.1016/j.epsl.2012.07.021
1266	Passey, B. H., Levin, N. E., Cerling, T. E., BROWN, F. H., Eiler, J. M., & Turekian, K. K.
1267	(2010). High-temperature environments of human evolution in East Africa based on
1268	bond ordering in paleosol carbonates. Proceedings of the National Academy of
1269	Sciences of the United States of America, 107(25), 11245–11249.
1270	https://doi.org/10.1073/pnas.1001824107
1271	Peral, M., Daëron, M., Blamart, D., Bassinot, F., Dewilde, F., Smialkowski, N., et al.
1272	(2018). Updated calibration of the clumped isotope thermometer in planktonic and
1273	benthic foraminifera. Geochimica et Cosmochimica Acta, 239, 1-16.
1274	https://doi.org/10.1016/j.gca.2018.07.016
1275	Petersen, S. V., Defliese, W. F. F., Saenger, C., Daëron, M., Huntington, K. W. W., John, C.
1276	M. M., et al. (2019). Effects of Improved 17 O Correction on Inter-laboratory
1277	Agreement in Clumped Isotope Calibrations, Estimates of Mineral-Specific Offsets,
1278	and Temperature Dependence of Acid Digestion Fractionation. Geochemistry,
1279	Geophysics, Geosystems, 20, 3495-3519. https://doi.org/10.1029/2018GC008127
1280	Petersen, Sierra V., Winkelstern, I. Z., Lohmann, K. C., & Meyer, K. W. (2016). The effects
1281	of Porapak TM trap temperature on δ 18 O, δ 13 C, and Δ 47 values in preparing samples
1282	for clumped isotope analysis. Rapid Communications in Mass Spectrometry, 30(1),
1283	199–208. https://doi.org/10.1002/rcm.7438

1284	Piasecki, A., Bernasconi, S. M., Grauel, A., Hannisdal, B., Ho, S. L., Leutert, T. J., et al.
1285	(2019). Application of Clumped Isotope Thermometry to Benthic Foraminifera.
1286	Geochemistry, Geophysics, Geosystems, 2018GC007961.
1287	https://doi.org/10.1029/2018GC007961
1288	Schauer, A. J., Kelson, J., Saenger, C., & Huntington, K. W. (2016). Choice of 17O
1289	correction affects clumped isotope (??47) values of CO2 measured with mass
1290	spectrometry. Rapid Communications in Mass Spectrometry, 30(24), 2607-2616.
1291	https://doi.org/10.1002/rcm.7743
1292	Schmid, T. W., & Bernasconi, S. M. (2010). An automated method for "clumped-isotope"
1293	measurements on small carbonate samples. Rapid Communications in Mass
1294	Spectrometry, 24(14), 1955–1963. https://doi.org/10.1002/rcm.4598
1295	Spooner, P. T., Guo, W., Robinson, L. F., Thiagarajan, N., Hendry, K. R., Rosenheim, B.
1296	E., & Leng, M. J. (2016). Clumped isotope composition of cold-water corals: A role
1297	for vital effects? Geochimica et Cosmochimica Acta, 179, 123-141.
1298	https://doi.org/10.1016/j.gca.2016.01.023
1299	Swart, P. K., Murray, S. T., Staudigel, P. T., & Hodell, D. A. (2019). Oxygen Isotopic
1300	Exchange between CO 2 and Phosphoric Acid: Implications for the Measurement of
1301	Clumped Isotopes in Carbonates. Geochemistry, Geophysics, Geosystems, (1950),
1302	2019GC008209. https://doi.org/10.1029/2019GC008209
1303	Thaler C., Katz A., Bonifacie M., Ménez B., Ader M. Oxygen isotope composition of waters
1304	recorded in carbonates in strong clumped and oxygen isotopic disequilibrium (2020).
1305	Biogeosciences, 17, 1-14, 2020. DOI 10.5194/bg-17-1-2020
1306	van Dijk, J., Fernandez, A., Storck, J. C., White, T. S., Lever, M., Müller, I. A., et al.
1307	(2019). Experimental calibration of clumped isotopes in siderite between 8.5 and 62°C
1308	and its application as paleo-thermometer in paleosols. Geochimica et Cosmochimica
1309	Acta, 254, 1–20. https://doi.org/https://doi.org/10.1016/j.gca.2019.03.018
1310	Veillard, C. M. A., John, C. M., Krevor, S., & Najorka, J. (2019). Rock-buffered
1311	recrystallization of Marion Plateau dolomites at low temperature evidenced by
1312	clumped isotope thermometry and X-ray diffraction analysis. Geochimica et
1313	Cosmochimica Acta, 252, 190–212. https://doi.org/10.1016/J.GCA.2019.02.012

1314	Wacker, U., Fiebig, J., & Schoene, B. R. (2013). Clumped isotope analysis of carbonates:
1315	comparison of two different acid digestion techniques. Rapid Communications in Mass
1316	Spectrometry: RCM, 27(14), 1631-42. https://doi.org/10.1002/rcm.6609
1317	Wacker, U., Fiebig, J., Tödter, J., Schöne, B. R., Bahr, A., Friedrich, O., et al. (2014).
1318	Empirical calibration of the clumped isotope paleothermometer using calcites of
1319	various origins. Geochimica et Cosmochimica Acta, 141, 127-144.
1320	https://doi.org/10.1016/j.gca.2014.06.004
1321	Wang, Z., Schauble, E. A., & Eiler, J. M. (2004). Equilibrium thermodynamics of multiply
1322	substituted isotopologues of molecular gases. Geochimica et Cosmochimica Acta,
1323	68(23), 4779–4797. https://doi.org/10.1016/j.gca.2004.05.039
1324	Werner, R. A., & Brand, W. A. (2001a). Referencing strategies and techniques in stable
1325	isotope ratio analysis. Rapid Communications in Mass Spectrometry, 15(7), 501-519.
1326	https://doi.org/10.1002/rcm.258
1327	Werner, R. A., & Brand, W. A. (2001b). Referencing strategies and techniques in stable
1328	isotope ratio analysis. Rapid Communications in Mass Spectrometry, 15(7), 501-519.
1329	https://doi.org/10.1002/rcm.258