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► **To cite this version:**

Guilherme Weber Sampaio de Melo, Ross Parnell-Turner, Robert Dziak, Deborah Smith, Marcia Maia, et al.. Uppermost mantle velocity beneath the Mid-Atlantic Ridge and transform faults in the equatorial Atlantic Ocean. *Bulletin of the Seismological Society of America*, 2020, 111 (2), pp.1067-1079. hal-03062624

**HAL Id: hal-03062624**

**<https://hal.science/hal-03062624v1>**

Submitted on 14 Dec 2020

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# Bulletin of the Seismological Society of America

## Uppermost mantle velocity beneath the Mid-Atlantic Ridge and transform faults in the equatorial Atlantic Ocean --Manuscript Draft--

<b>Manuscript Number:</b>	BSSA-D-20-00248R5
<b>Article Type:</b>	Article
<b>Section/Category:</b>	Regular Issue
<b>Full Title:</b>	Uppermost mantle velocity beneath the Mid-Atlantic Ridge and transform faults in the equatorial Atlantic Ocean
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<b>Abstract:</b>	<p>Seismic rays travelling just below the Moho provide insights into the thermal and compositional properties of the upper-mantle, and can be detected as Pn phases from regional earthquakes. Such phases are routinely identified in the continents, but in the oceans, detection of Pn phases is limited by a lack of long-term instrument deployments. We present estimates of upper-mantle velocity in the equatorial Atlantic Ocean from Pn arrivals beneath, and flanking, the Mid-Atlantic Ridge, and across several transform faults. We analyzed waveforms from 50 earthquakes with magnitude <math>MW &gt; 3.5</math>, recorded over 12 months in 2012–2013 by five autonomous hydrophones and a broadband seismograph located on the St. Peter and St. Paul archipelago. The resulting catalog of 152 ray paths allows us to resolve spatial variations in upper-mantle velocities, which are consistent with estimates from nearby wide-angle seismic experiments. We find relatively high velocities near to the St. Paul transform system (<math>\sim 8.4 \text{ km s}^{-1}</math>), compared to lower ridge-parallel velocities (<math>\sim 7.7 \text{ km s}^{-1}</math>). Hence, this method is able to resolve ridge-transform scale velocity variations. Ray paths in lithosphere younger than 10 Myr have mean velocities of <math>7.9 \pm 0.5 \text{ km s}^{-1}</math>, which is slightly lower than those sampling lithosphere older than 20 Myr (<math>8.1 \text{ km} \pm 0.3 \text{ s}^{-1}</math>). There is no apparent systematic relationship between velocity and ray azimuth, which could be due to thickened lithosphere or complex mantle upwelling, although uncertainties in our velocity estimates may obscure such patterns. We also do not find any correlation between Pn velocity and shear wave speeds from the global SL2013sv model at depths <math>&lt; 150 \text{ km}</math>. Our results demonstrate that data from long-term deployments of autonomous hydrophones can be used to obtain rare and insightful estimates of uppermost mantle velocities over hundreds of kilometers, in otherwise inaccessible parts of the deep oceans.</p>
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	John Collins Woods Hole Oceanographic Institution jcollins@whoi.edu Seismology expert with interest in passive methods
<b>Opposed Reviewers:</b>	
<b>Response to Reviewers:</b>	
<b>Additional Information:</b>	
<b>Question</b>	<b>Response</b>
<p><b>Key Point #1:</b> Three key points will be printed at the front of your manuscript so readers can get a quick overview. Please provide three COMPLETE sentences addressing the following: 1) state the problem you are addressing in a FULL sentence; 2) state your main conclusion(s) in a FULL sentence; and 3) state the broader implications of your findings in a FULL sentence. Each point must be 110 characters or less (including spaces).</p>	<p>Upper mantle velocity structure is important, but difficult and expensive to measure in remote ocean basins</p>
<b>Key Point #2:</b>	<p>Pn arrivals detected by autonomous hydrophones constrain upper mantle velocity in equatorial Atlantic Ocean</p>
<b>Key Point #3:</b>	<p>No azimuthal dependence of VPn, could be due to thickened lithosphere, or complex mantle upwelling</p>



# 1 **Uppermost mantle velocity beneath the Mid-Atlantic Ridge and transform** 2 **faults in the equatorial Atlantic Ocean**

3 Guilherme W. S. de Melo, Ross Parnell-Turner, Robert P. Dziak, Deborah K. Smith,  
4 Marcia Maia, Aderson F. do Nascimento, and Jean-Yves Royer

## 5 6 **Abstract**

7 Seismic rays travelling just below the Moho provide insights into the thermal and compositional  
8 properties of the upper-mantle, and can be detected as *Pn* phases from regional earthquakes.  
9 Such phases are routinely identified in the continents, but in the oceans, detection of *Pn* phases is  
10 limited by a lack of long-term instrument deployments. We present estimates of upper-mantle  
11 velocity in the equatorial Atlantic Ocean from *Pn* arrivals beneath, and flanking, the Mid-  
12 Atlantic Ridge, and across several transform faults. We analyzed waveforms from 50  
13 earthquakes with magnitude  $M_w > 3.5$ , recorded over 12 months in 2012–2013 by five  
14 autonomous hydrophones and a broadband seismograph located on the St. Peter and St. Paul  
15 archipelago. The resulting catalog of 152 ray paths allows us to resolve spatial variations in  
16 upper-mantle velocities, which are consistent with estimates from nearby wide-angle seismic  
17 experiments. We find relatively high velocities near to the St. Paul transform system ( $\sim 8.4 \text{ km s}^{-1}$ ),  
18 compared to lower ridge-parallel velocities ( $\sim 7.7 \text{ km s}^{-1}$ ). Hence, this method is able to  
19 resolve ridge-transform scale velocity variations. Ray paths in lithosphere younger than 10 Myr  
20 have mean velocities of  $7.9 \pm 0.5 \text{ km s}^{-1}$ , which is slightly lower than those sampling lithosphere  
21 older than 20 Myr ( $8.1 \text{ km} \pm 0.3 \text{ s}^{-1}$ ). There is no apparent systematic relationship between  
22 velocity and ray azimuth, which could be due to thickened lithosphere or complex mantle  
23 upwelling, although uncertainties in our velocity estimates may obscure such patterns. We also

24 do not find any correlation between  $P_n$  velocity and shear wave speeds from the global  
25 SL2013sv model at depths  $< 150$  km. Our results demonstrate that data from long-term  
26 deployments of autonomous hydrophones can be used to obtain rare and insightful estimates of  
27 uppermost mantle velocities over hundreds of kilometers, in otherwise inaccessible parts of the  
28 deep oceans.

29

### 30 **Introduction**

31 Seismic velocity measurements provide a useful tool for investigating spatial variations in upper-  
32 mantle properties, such as temperature and anisotropy, with implications for melt supply and  
33 mantle heterogeneity (e.g. Lin and Phipps Morgan, 1992; Dunn *et al.*, 2005). These  
34 measurements are relatively straightforward to obtain on the continents (e.g. Chulick and  
35 Mooney, 2002; Chulick *et al.*, 2013). However, it remains challenging and expensive to measure  
36 upper-mantle seismic velocity in the deep ocean, due to its remote location and difficulties in  
37 deploying long-term instruments on the seafloor.  $P_n$  phases are rays that are critically refracted  
38 at the Moho and propagate along the top of the uppermost mantle (e.g. Linehan, 1940;  
39 Brandsdottir and Menke, 1997). At the Mid-Atlantic Ridge (MAR) from  $10^\circ\text{N}$  to  $35^\circ\text{N}$ ,  $P_n$   
40 arrivals from 48 individual ray paths were recorded with hydrophones, and used to investigate  
41 upper-mantle velocities, giving a mean velocity of  $8.0 \pm 0.1$  km s<sup>-1</sup> (Dziak *et al.*, 2004). This  
42 velocity estimate was higher than that from nearby active source seismic experiments along the  
43 ridge axis ( $7.5\text{--}7.9$  km s<sup>-1</sup>; Canales *et al.*, 2000), probably due to the effects of younger and  
44 thinner oceanic lithosphere being sampled by the refraction profiles, and the effects of averaging  
45 velocities across all rays. Despite such advances, upper-mantle velocities in the deep oceans  
46 remain poorly constrained, and the potential for hydrophone-recorded  $P_n$  phases to resolve

47 spatial variations in upper-mantle velocity has not yet been sufficiently tested.

48         Here, we use *Pn* arrivals from regional earthquakes to constrain upper-mantle velocity in  
49 the equatorial Atlantic Ocean. Arrivals were recorded by a combination of five moored  
50 hydrophones and a single seismograph station installed on the St. Peter and St. Paul islets, giving  
51 152 ray paths that sample mantle conditions both on- and off-axis, and across the St. Paul  
52 transform system. Our study is coincident with several mantle velocity estimates from a wide-  
53 angle seismic experiment (Le Pichon *et al.*, 1965), and hence has the opportunity to validate  
54 spatial variations in velocity revealed by groups of similar ray paths.

55

### 56 ***Equatorial Atlantic Ocean***

57         In the equatorial Atlantic Ocean (10°N–5°S and 34°W–21°W), the MAR is offset by  
58 some of the longest transform faults on Earth, including the Strakhov, St. Paul, and Romanche  
59 transforms (Figure 1). The St. Paul transform system consists of four transform faults and three  
60 intra-transform ridge segments that accommodate an offset of 630 km. The northwest transform  
61 fault is currently undergoing transpression, giving rise to the ~200 km-long and ~30 km-wide  
62 Atobá ridge (Maia *et al.*, 2016), and also uplift of 1.5mm yr<sup>-1</sup> at the St. Peter and St. Paul islets  
63 (Campos *et al.*, 2010; Maia *et al.*, 2016). Other transforms in the system do not host topographic  
64 highs or an island related to transpression, and hence presumably are not experiencing uplift. In  
65 the three intervening spreading segments, seafloor spreading is slow, at ~16 mm yr<sup>-1</sup> average half  
66 rate (DeMets *et al.*, 2010). Faulting plays an important role in crustal accretion, and seismicity  
67 rates are relatively high, providing a useful tool to investigate the properties of the crust and  
68 upper mantle, as well as deformation at long-offset strike-slip systems (e.g. Francis *et al.*, 1978;  
69 Abercrombie and Ekstrom, 2001; de Melo and do Nascimento, 2018).

70 **Methods**

71 ***Waveform Data***

72 We analyzed *Pn* arrivals in waveform data recorded by a combination of five moored  
73 autonomous hydrophones and one land-based seismograph (Figure 2). The five autonomous  
74 hydrophone instruments were deployed during two separate experiments: stations EA2 and EA8  
75 were part of the Equatorial Atlantic (EA) array (Smith et al., 2012). Data were recorded at 16-bit  
76 resolution and a sampling rate of 250 Hz; for further details on these hydrophone instruments see  
77 Fox *et al.* (2001). Hydrophones H2, H4, H5 were deployed during the COLd Mantle Exhumation  
78 and Intra-transform Accretion experiment (COLMEIA; Maia *et al.*, 2014, 2016), and recorded  
79 data at 24 bit-resolution with a sampling rate of 240 Hz; for further instrument details see D’Eu  
80 *et al.*(2012). We also used waveform data recorded by a three-component broadband  
81 seismograph installed at the St. Peter and St. Paul Archipelago Scientific Station on the  
82 Belmonte islet (ASPSP; de Melo and do Nascimento, 2018). This station is operated by the  
83 Seismological Laboratory of Federal University of Rio Grande do Norte in cooperation with the  
84 Brazilian Navy. The sparse distribution and mixed instrument types we used means that data  
85 coverage is uneven, as shown in Figure 1b. Waveform data were examined for the time period  
86 from July 2012 to July 2013, with recording intervals dictated by technical challenges and vessel  
87 schedules (Figure 1b).

88

89 ***Pn analysis***

90 Prior to manually picking *Pn* arrivals, we applied a 6–20 Hz Butterworth bandpass filter  
91 to the hydrophone data in order to suppress unwanted noise. A bandpass filter with range 4–12  
92 Hz was applied prior to picking arrivals from the ASPSP seismograph, to suppress additional



93 microseism noise due to its island location. Based upon origin time, events were manually  
94 associated with earthquakes in the International Seismological Center Bulletin (ISC), yielding  
95 hypocenter locations, origin times, and magnitudes ranging from 3.5 to 5.4  $M_w$ . Earthquakes  
96 mostly occur due to strike-slip faulting along the Strakhov, St. Paul, and Romanche transform  
97 faults, with additional events due to extension along the intervening spreading ridge segments  
98 (Figure 2a). Example arrivals from three events are shown in Figures 3 and 4, highlighting the  
99 typical response to strike-slip and normal faulting earthquakes ranging in magnitude from 4.6 to  
100 5.3  $M_w$ .

101       Typical  $Pn$ -arrivals are emergent, and have low signal-to-noise ratio (SNR; noted in  
102 Figures 3 and 4), making pick identification challenging. Given the mixed nature of our network  
103 and often noisy arrivals, picks were made based on the onset of emergent energy combined with  
104 changes in SNR, waveform character and amplitude. The observation of linear move-out,  
105 consistent with upper mantle velocity, added confidence to our picks, since this moveout is  
106 evident across the hydrophone array stations due to wave propagation along the crust-mantle  
107 interface (see common-receiver plots in Supplementary Figures S1–S6).  $P$ -arrivals are easily  
108 distinguished from  $T$ -phase arrivals, which arrive much later than  $P$ -arrivals, are emergent in  
109 character, and are higher in amplitude than  $P$ -arrivals (see hydrophone H5 in Figure 4). The  
110 catalog of detected events is given in Table S1.

111       In order to further test whether the detected arrivals were  $Pn$  phases, we compared the  
112 observed travel times to those predicted by the global iasp91 velocity model (Kennett and  
113 Engdahl, 1991). For each source-receiver ray path, we calculated the predicted  $Pn$  arrival time  
114 using iasp91, with the addition of a station-dependent delay to account for the propagation time  
115 from seafloor to hydrophone. This delay (1.2–2.5 s, see Table 1) was estimated using the

116 hydrophone mooring cable length at each station, and the local water sound velocity estimated  
117 from the Global Ocean Sound Speed Profile Library (Barlow, 2019). The predicted  $Pn$  arrival  
118 times differ from the observed  $Pn$  arrivals by 5–10 s (Figures 3 and 4), a difference which arises  
119 since the iasp91 model contains a crustal layer that is much thicker (30 km) than that expected in  
120 the oceans (~6 km). Hence, the differences in observed and predicted  $Pn$  arrival time are  
121 probably dominated by this additional crustal layer thickness in the velocity model, plus  
122 earthquake location and origin time uncertainties. Although these differences are evident, the  
123 waveform character and linear move-out velocity give us confidence in our identification of  
124 these emergent phases as  $Pn$  arrivals.

125       ISC origin times were subtracted from the  $Pn$  arrival times to obtain travel times for each  
126 ray path (i.e. each event-station pair). We account for travel time in the oceanic crust by  
127 subtracting ray path distances and travel times for the portion of the path that travels through the  
128 crust, assuming that all events occurred at 10 km depth (ISC catalog), and that crustal thickness  
129 is uniformly 6.0 km with a crustal velocity of 6.5 km s<sup>-1</sup> (Christeson *et al.*, 2019). For each  
130 station, we then calculate the distance and travel time for the portion of the ray path that extends  
131 from an earthquake in the crust to the Moho, and back from the Moho to the receiver.  $Pn$  velocity  
132 is obtained by dividing the distance travelled in the mantle by the travel time in the mantle.  
133 Details of these corrections for each station are given in Table 1.

134       Our approach yielded 152  $Pn$  velocity estimates from the catalog of 50 regional  
135 earthquakes (Figure 5). Although epicentral distances range from 32 km to ~1095 km, all 50  
136 events were detected at nearly all available stations, implying that the detection threshold of the  
137 combined hydrophones and ASPSP station is at least  $M_w$  3.5. Since most stations were located  
138 either near to, or to the north of, the St. Paul fracture zone, our ray path coverage is more

139 comprehensive in the northern part of the study area. Ray paths sampling upper-mantle  
140 velocities to the south of the St. Paul fracture zone are restricted to events detected by  
141 hydrophone EA8, and those originating from four earthquakes located at the eastern end of the  
142 Romanche transform fault (Figure 5).

143

#### 144 ***Pn velocity uncertainty***

145         The two most significant potential sources of error in our analysis are hypocenter  
146 locations of events in the ISC Catalog, and *Pn* arrival time picks. We estimated hypocenter  
147 location (and hence epicentral distance) error to be  $\pm 10$  km, based upon ISC catalog location  
148 and typical error in global earthquake location (Lohman and Simons, 2005; Weston *et al.*, 2012).  
149 This hypocenter location error implicitly includes other uncertainties associated with ISC catalog  
150 locations, such as those caused by un-modeled three-dimensional velocity structure and picking  
151 errors, which result in trade-offs between origin time and location (Bondár and Storchak, 2011).  
152 Arrival time pick (and hence also travel time) errors were investigated by estimating SNR for  
153 each arrival via two methods, one using the amplitude ratio between peak signal and root mean  
154 square noise, and another via the ratio between the short time average amplitude and long time  
155 average amplitude (STA/LTA; Figure S7). We find that both SNR estimates are only weakly  
156 dependent on epicentral distance and magnitude, however we do observe station-dependent  
157 variations in the scatter in SNR. We quantify this scatter in terms of the standard deviation of  
158 SNR of arrivals for a particular station (Figure S7e), which likely is due to persistent local noise  
159 sources. Hence we estimated arrival time pick error based on the emergent character of arrivals  
160 and the standard deviation of SNR, with station-dependent errors defined as  $\pm 0.5$  s for EA2 and  
161 EA8;  $\pm 1.0$  s for H2, H4 and H5; and  $\pm 0.3$  s for ASPSP.

162 The total uncertainty in our velocity estimate,  $\delta v$ , was estimated by assuming that  
163 epicentral distance,  $d$ , and travel time,  $t$ , have errors that are uncorrelated and random. This  
164 assumption is valid since we attribute the main source of travel time error to uncertainty in  
165 picking of  $Pn$  arrivals (which in turn depends on waveform character and noise level), and the  
166 distance error is most significantly affected by error in earthquake location from the ISC catalog,  
167 which is assumed to be constant and hence is independent from hydrophone  $Pn$  pick error. We  
168 formally propagate the errors in  $d$  and  $t$ , as follows

$$169 \quad \delta v = v \sqrt{\left(\frac{\delta d}{d}\right)^2 + \left(\frac{\delta t}{t}\right)^2},$$

170 where  $\delta d$  is epicentral distance error, and  $\delta t$  is travel time error (e.g. Taylor, 1997).

171 Although receiver location uncertainty is negligible for the land station ASPSP (located  
172 with meter-scale accuracy via the Global Positioning System), there is potential location  
173 uncertainty for the moored hydrophones in our network. Moored hydrophone locations were  
174 obtained by acoustic triangulation between the mooring acoustic release and the deployment  
175 vessel soon after the moorings settled on the seafloor, within error of several meters. In order to  
176 account for the possibility of abnormally strong current motion, each instrument was fitted with a  
177 pressure and temperature logger below the floatation package, so that any significant hydrophone  
178 depth changes would be recorded (e.g. Fox *et al.*, 2001). Significant depth changes were not  
179 detected during depolyments, and thus we assume that the hydrophone location was constant  
180 during data collection, and hence hydrophone location uncertainty is less than 10 m.

181

182

183

184

## 185 **Results**

### 186 *Pn velocities*

187       The resulting 152 *Pn* ray paths (Figure 5b) and travel times (Figure 6) indicates upper-  
188 mantle velocities that vary considerably across the study area, with estimates ranging between  
189 7.2 and 11.1 km s<sup>-1</sup>, and uncertainties ranging from 0.1 to 1.9 km s<sup>-1</sup> (Table S2). Variability in  
190 reduced travel time increases with epicentral distance (Figure 6), although SNR does not show a  
191 similar trend (Figure S7). Hence the epicentral distance-dependent scatter in reduced travel time  
192 is likely due to variations in the depth of ray penetration (which increases with epicentral  
193 distance), and not due to increasing pick uncertainty. At the center of the study area there  
194 appears to be a longitudinal variation in *Pn* velocity, with events originating near the St. Paul  
195 transform system, and sampling adjacent lithosphere, having higher velocities than those from  
196 the adjacent spreading centers (Figure 5a). The best constrained estimate for sub-axis, ridge-  
197 parallel mantle velocity comes from ray paths that sample the portion of the spreading axis  
198 between the Strakhov fracture zone and stations near the St. Paul fracture zone (H2, H5 and  
199 ASPSP). Here, *Pn* travel times consistently imply relatively low velocities, with a mean of 7.7  
200 km s<sup>-1</sup>. Slightly higher velocities ranging between 7.8 and 8.2 km s<sup>-1</sup> are indicated by ray paths  
201 between hydrophone EA2 and the Strakhov fracture zone, oriented roughly parallel to a plate  
202 spreading flowline. Ray paths oriented southwest-northeast (azimuth ~060°), i.e. oblique to the  
203 spreading direction, between events on the St. Paul fracture zone and detected at hydrophone  
204 EA2, have some of the highest mantle velocities (between 7.6 and 8.5 km s<sup>-1</sup>) compared to other  
205 rays sampling areas unaffected by fracture zones. Velocity estimates in the vicinity of the St.  
206 Paul fracture zone itself (from transform faulting events detected by hydrophones H2, H4 and

207 H5, and ASPSP) show considerable variation, ranging from 8.0 to 9.1 km s<sup>-1</sup> and a mean of 8.4  
208 km s<sup>-1</sup>, and little apparent spatial consistency. Among these events, we encountered one of the  
209 highest *Pn* velocities (9.0 km ± 0.2 s<sup>-1</sup>) in this study, for a ray path oriented roughly parallel to  
210 the St. Paul transform fault (ray azimuth ~105°) between an event near the St. Paul islets and  
211 detected by hydrophone H4.

212 South of the St. Paul fracture zone, ray paths from events detected by hydrophone EA8  
213 showed considerable variation in upper-mantle velocity, which range from 7.2 to 9.0 km s<sup>-1</sup>. Ray  
214 paths originating from the spreading axis north of the St. Paul transform fault and trending ~170°  
215 towards EA8, have velocities of 7.3–8.1 km s<sup>-1</sup>, while ray paths from the St. Paul transform fault  
216 trending ~185° towards EA8 have consistently higher velocities of 7.6–9.1 km s<sup>-1</sup>.

217 Only 12 ray paths sampling the upper-mantle parallel and adjacent to the spreading axis  
218 between the southern extent of the St. Paul transform fault and the Romanche transform fault are  
219 available. This relatively poor coverage in ray paths in this area hinders our interpretation, where  
220 velocities range from 7.2 to 8.3 km s<sup>-1</sup>.

221

## 222 **Discussion**

### 223 *Upper-mantle velocity structure*

224 In general, rays originating from the St. Paul transform system have higher velocities than  
225 those originating from active spreading centers to the east and west (Figure 5a), probably due to  
226 cooler conditions at the Moho along the transform. Our estimates of upper-mantle *Pn* velocities  
227 broadly agree (within error) with *Pn* velocities from radially stratified velocity models such as  
228 PREM (Dziewonski and Anderson, 1981) and iasp91 (Figure 6; Kennett and Engdahl, 1991).  
229 Our *Pn* velocity estimates are also consistent with mantle velocity estimates from a series of

230 reversed wide-angle refraction seismic profiles (i.e. with multiple shot points giving overlapping  
231 coverage) collected in the equatorial Atlantic during R/V *Atlantis* cruise A180 (Figure 5b; Le  
232 Pichon *et al.*, 1965). The modal difference in velocity between refraction profiles from Le  
233 Pichon *et al.* (1965) and all intersecting ray paths is  $0.2 \text{ km s}^{-1}$  (see histogram in Figure 5c),  
234 although our *Pn* velocity estimates are consistently lower than those reported by Le Pichon *et al.*  
235 (1965), with a maximum disagreement of  $1.2 \text{ km s}^{-1}$ . A mantle velocity of  $8.30 \text{ km s}^{-1}$  was  
236 reported along profile A180-48, which is 283 km-long, and crosses the eastern side of the St.  
237 Paul transform fault (near  $\sim 26.3^\circ\text{W}$ ), trending northeast-southwest (Figure 5b). This velocity is  
238 consistent with that inferred from *Pn* ray paths with a similar orientation, originating from  
239 earthquakes on the St. Paul transform fault that were detected by hydrophone EA8. Ray paths  
240 that intersect profile A180-48 (at angles either perpendicular or oblique to the trend of the  
241 refraction profile) typically indicate lower upper-mantle velocities, ranging from  $7.3$  to  $8.1 \text{ km s}^{-1}$ ,  
242 with the exception of one anomalous ray path oriented parallel with the St. Paul transform fault  
243 with a velocity of  $9.0 \text{ km s}^{-1}$ . Refraction profiles A180-40 and -42 are oriented roughly east-  
244 west, are located  $\sim 100$  km north of the Romanche transform fault, and have velocities of  $8.03$   
245 and  $8.49 \text{ km s}^{-1}$ , respectively. Although there are only four *Pn* ray paths near to these profiles,  
246 with near-perpendicular orientation, they indicate velocities ranging from  $7.6$  to  $8.2 \text{ km s}^{-1}$ , and  
247 hence are in broad agreement with the refraction estimates. Our velocity estimates of  $7.6$  to  $8.2$   
248  $\text{km s}^{-1}$  are also in agreement with a velocity estimate of  $8.0 \text{ km s}^{-1}$  from an active source  
249 experiment near  $18^\circ\text{W}$  roughly perpendicular to the St. Paul fracture zone, which at this  
250 longitude separates 40 Myr old crust in the south from 70 Myr old crust in the north (Grove *et*  
251 *al.*, 2019). The general agreement between upper-mantle velocities from the refraction profiles

252 and our *Pn* arrivals validates our results, and implies that spatial trends observed in the study  
253 area are likely to be real.

254 Elsewhere along the MAR, between 10° to 40°N, a mean upper-mantle velocity of 8.0 km  
255  $\pm 0.1$  km s<sup>-1</sup> was estimated using a similar method to this study with *Pn* arrivals detected by an  
256 array of autonomous hydrophones (Dziak *et al.*, 2004). Ray paths used by Dziak *et al.*, (2004)  
257 often crossed the ridge axis, spanned a series of fracture zones, and extended onto older crust,  
258 which may explain the close agreement in results. This result suggests that off-axis and on-axis  
259 upper mantle characteristics are similar in the northern and equatorial Atlantic Ocean.

260 Near the Oceanographer transform fault on the MAR (~35°N), a two-dimensional  
261 tomographic inversion of wide-angle seismic refraction data suggests velocities of 7.4–7.8 km s<sup>-1</sup>  
262 (Canales *et al.*, 2000; Hooft *et al.*, 2000). These results agree within error with our estimates of  
263 *Pn* velocity from rays sampling on-axis upper-mantle to the north of the St. Paul transform fault  
264 (Figure 5b), which are typically 7.2–8.0 km s<sup>-1</sup>.

265

### 266 ***Upper-mantle velocity and plate age***

267 Seismic velocities in the upper-mantle near to the ridge axis, i.e. in young lithosphere, are  
268 expected to be lower than in off-axis areas, due to upwelling of hot material (e.g. Turcotte and  
269 Schubert, 2002). Following the removal of minor gridding artifacts associated with fracture zone  
270 traces, we used a global crustal age model (Müller *et al.*, 2008) to assign a mean crustal age  
271 along each ray path, for comparison with *Pn* velocity (Figure 7a).

272 Ray paths sampling lithosphere younger than 10 Myr show a wide range of velocities,  
273 with a mean of 7.9 km s<sup>-1</sup> and standard deviation of 0.5 km s<sup>-1</sup>. Twenty ray paths yield velocities  
274 less than 7.5 km s<sup>-1</sup>. *Pn* velocities for ray paths sampling lithosphere older than 20 Myr are



275 slightly higher, with a mean of  $8.1 \text{ km s}^{-1}$  and standard deviation of  $0.3 \text{ km s}^{-1}$ , while only two  
276 ray paths give velocities lower than  $7.5 \text{ km s}^{-1}$  (Figure 7a). Most rays cover a wide range of  
277 crustal ages, so this geometry, and our averaging approach, may smear the possible effects of  
278 lithospheric aging. The lack of rays travelling exclusively via older lithosphere may also obscure  
279 any progressive trend between upper-mantle velocity and crustal age. However, the tendency  
280 toward the inclusion of lower velocities in younger crust (Figure 7a) reflects the expected  
281 variation with respect to the zone of axial upwelling.

282

### 283 *Azimuthal Seismic Anisotropy*

284 Laboratory experiments have shown that the mantle can experience significant shear  
285 strain during corner flow at the ridge axis, leaving an anisotropic fabric in the lithospheric mantle  
286 as minerals (e.g. olivine) are aligned into a lattice preferred orientation (LPO; e.g. Zhang and  
287 Karato, 1995; Nicolas and Christensen, 2011). Anisotropy consistent with a LPO formed by two-  
288 dimensional mantle flow has been measured at some locations in the oceanic upper mantle, in  
289 particular at the fast-spreading East Pacific Rise (e.g. Raitt *et al.*, 1969; Lin *et al.*, 2016),  
290 however the strength of anisotropy varies widely, and debate remains about its origins (e.g. Mark  
291 *et al.*, 2019). Since isochrons in this region are fairly uniform (Figure 5),  $V_{Pn}$  anisotropy could  
292 be expected parallel to paleo-relative plate motion, although this assumption has been shown to  
293 not apply everywhere (VanderBeek and Toomey, 2017).

294 We investigated the dependence of mantle velocity with azimuth, and use epicentral  
295 distance as a proxy for depth of mantle penetration to group rays (Figure 7b). No discernable  
296 pattern is evident in rays grouped by epicentral distance, including those expected to sample  
297 deepest in the mantle with epicentral distances  $> 700 \text{ km}$  (blue lines in Figure 7c). Removing

298 rays with  $V_{Pn}$  error  $> 0.4 \text{ km s}^{-1}$  also does not resolve any azimuthal dependence (Figure 7d), nor  
299 does separating rays by mean crustal age (Figures 7e and 7f).

300         The apparent lack of such azimuthal dependence could be due to several reasons. First,  
301 azimuthal dependence may be too subtle to be resolved by our  $V_{Pn}$  estimates, given the  
302 uncertainties in hypocenter location and crustal thickness discussed above. Second, the slow  
303 spreading rate of the MAR ( $\sim 32 \text{ mm yr}^{-1}$  total rate; (DeMets *et al.*, 2010)) may result in a  
304 thickened lithosphere that is dominantly cooled by conduction, thus inhibiting corner flow (e.g.  
305 Sleep, 1975). As a result, deformation could be accommodated by faulting at depths of 5–10 km  
306 beneath the Moho, reducing the viscous strain in the mantle at these depths, and suppressing the  
307 anisotropy recorded in the mantle (e.g. Ribe, 1989). Observations of weaker or anomalous  
308 anisotropy elsewhere in the Atlantic Ocean are consistent with our findings (e.g. Gaherty *et al.*,  
309 2004; Dunn *et al.*, 2005). Third, complex, three-dimensional upwelling patterns near the ridge  
310 axis could result in anisotropy on relatively short wavelengths (Lin and Phipps Morgan, 1992),  
311 which would be smeared along our relatively long ray paths, and hence not be resolved.

312

### 313 ***Pn and surface wave velocity***

314         To explore the relationship between  $V_{Pn}$  and the thermal structure of the asthenospheric  
315 upper-mantle, we compared our velocity estimates with a global, vertically polarized shear speed  
316 model SL2013sv (Schaeffer and Lebedev, 2013). Our objective is to evaluate our observations  
317 of uppermost mantle properties in the context of deeper mantle properties. We do not aim to  
318 directly validate our  $V_{Pn}$  estimates via this comparison. This model was chosen because it is  
319 particularly sensitive to anomalies within the upper-mantle, and hence provides a window into  
320 the upper mantle structure directly beneath our  $Pn$  ray paths (Schaeffer and Lebedev, 2013). We

321 extracted values of vertically polarized tomographic shear velocity anomaly ( $\%dVs$ ) at 100 km  
322 intervals along each ray path, from slices through the SL2013sv model at depths of 25, 50, 75  
323 and 150 km. We then calculated the mean  $\%dVs$  along each ray path, at each depth interval  
324 (Figure 8). At 25 and 50 km depths, the effects of the ridge axis are evident, with higher  
325 velocities associated with ray paths travelling off-axis (detected by EA2 and EA8), and hence not  
326 sampling the relatively low-velocity axial region (Figures 8a and 8b). This effect is less  
327 pronounced at 75 km depth (Figure 8c), and is not apparent at 150 km depth, which presumably  
328 reflects sub-plate velocities. The lack of correlation between SL2013sv and  $Pn$  velocities at 150  
329 km suggests that our  $V_{Pn}$  estimates, sensitive to the velocity structure directly beneath the Moho,  
330 do not record deeper, larger-scale sub-plate (i.e. asthenospheric) processes and anomalies. Hence  
331 our observed  $V_{Pn}$  variability may instead arise due to local variations in melt supply, lithospheric  
332 thickness, or faulting.

333

### 334 **Conclusions**

335 We used a network of five autonomous hydrophones and a broadband seismograph to  
336 detect  $Pn$  arrivals from regional earthquakes in the equatorial Atlantic Ocean over a period of  
337 ~12 months between 2012 and 2013. Our estimates of upper-mantle velocity from the travel  
338 times of 152  $Pn$  arrivals broadly agree (mostly within  $0.2 \text{ km s}^{-1}$ ) with those from nearby seismic  
339 refraction experiments.

340 We find that the upper-mantle near the St. Paul transform system has consistently high  
341 velocities ( $>8 \text{ km s}^{-1}$ ), compared to relatively low velocities ( $\sim 7.5 \text{ km s}^{-1}$ ) in the adjacent MAR  
342 spreading segments northwest of the transform. This spatial pattern is consistent with the notion  
343 that  $Pn$  ray paths sample lower velocity mantle near the ridge axis, and higher velocity material

344 near transforms, which are generally cooler, despite the presence of intra-transform spreading  
345 segments. We do not resolve any dependence between  $V_{Pn}$  and azimuth, which could either be  
346 due to observational uncertainty, or due to the combined effects of thickened lithosphere and  
347 more complex mantle upwelling patterns under slow-spreading conditions. We also do not find  
348 any correlation between  $V_{Pn}$  and vertically polarized shear speed from the global SL2013sv  
349 model, indicating that our method is not sensitive to properties of the asthenosphere. The close  
350 agreement between our results and those from seismic refraction experiments demonstrates that  
351 the relatively simple method of using sparse arrays of autonomous hydrophones to detect  $Pn$   
352 arrivals can be used to obtain accurate estimates of upper-mantle velocities. Hence, this method  
353 provides a useful complement to deployments of other seafloor instruments such as ocean  
354 bottom seismographs, in remote areas where direct observations are typically elusive.

355

### 356 **Data and Resources**

357 All  $Pn$  velocities obtained in this study using the hydrophones data of the COLMEIA/EA array  
358 (Smith *et al.*, 2012; Maia *et al.*, 2014) and the seismic records of the and ASPSP station (de Melo  
359 and do Nascimento., 2018), are presented in tables of Supplemental Material. Analysis and figure  
360 preparation were carried out using the Generic Mapping Tools version 5.4.5 (Wessel *et al.*,  
361 2013), Seismic Analysis Code (Helffrich *et al.*, 2013). Earthquake locations used in this work  
362 were obtained from the International Seismological Center Bulletin database at  
363 [www.isc.ac.uk/iscbulletin/search/bulletin/](http://www.isc.ac.uk/iscbulletin/search/bulletin/) (last accessed November 2019). The Global Centroid  
364 Moment Tensor Project database of Ekström *et al.* (2012) was searched  
365 using [www.globalcmt.org/CMTsearch.html](http://www.globalcmt.org/CMTsearch.html) (last accessed November 2019).

366

367

368

369 **Acknowledgements**

370 This research was supported by National Science Foundation grants EAR-1062238, EAR-  
371 1062165 and OCE-1839727, and by an InterRidge Student Fellowship for GWSdM. The  
372 COLMEIA expedition was funded by the French Ministry of Research through its grant to the  
373 French Oceanographic Fleet. COLMEIA hydrophone deployment was funded by LABEX MER  
374 grant “Actions à la mer”, and instruments were recovered with the help of the Brazilian Navy.  
375 AFdN thanks the support of the Brazilian Navy and CNPq for grants 392484441/2012-4 and  
376 303817/2014-3. This paper is NOAA/Pacific Marine Environmental Laboratory contribution  
377 number 5116. We thank Associate Editor, T. Brocher, B. VanderBeek, and an anonymous  
378 reviewer for their constructive input. Any opinion, findings, and conclusions or  
379 recommendations expressed in this material are those of the authors and do not necessarily  
380 reflect the views of the National Science Foundation.

381

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531 **Table 1.** Details of seismograph (S) and hydrophone (H) sensors used for *Pn* analysis. Sensor  
 532 depth is given below sea level (bsl); water delay is based upon cable length, and water/crust  
 533 corrections are applied to each *Pn* ray path individually.

534

Station name	Sensor type	Lat, °N	Lon, °E	Depth bsl, m	Cable length, m	Water delay, s	Crust path correction, km	Crust travel time correction, s
ASPSP	S	0.9169	-29.3459	-16	-	-	12.5	1.9
EA2	H	4.9907	-22.9931	800	3912	2.10	23.8	7.2
EA8	H	-2.5159	-29.2181	800	3242	2.54	23.0	6.5
H2	H	1.3297	-31.3445	700	2260	1.57	21.8	5.5
H4	H	0.4123	-24.6437	700	1860	1.24	21.3	5
H5	H	0.1552	-27.7875	700	3060	2.04	22.8	6.3

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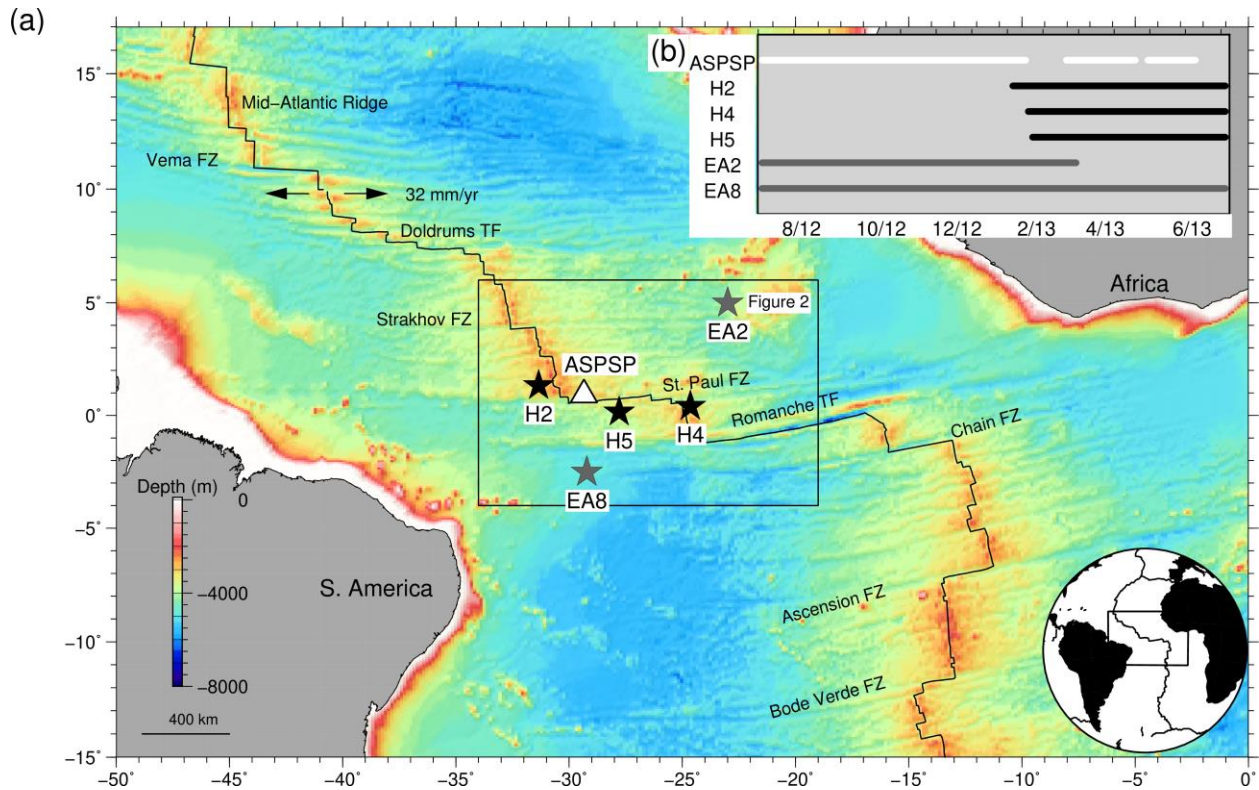
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617 **Figures**



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620 **Figure 1. a)** Regional bathymetric map of equatorial Atlantic ocean. White triangle shows

621 ASPSP seismograph station, located on St. Peter and St. Paul islets; black/gray stars are

622 COLMEIA / EA hydrophone networks, respectively (Smith *et al.*, 2012; Maia *et al.*, 2014); black

623 line is Mid-Atlantic Ridge, with selected transforms and half-spreading rate noted (arrows).

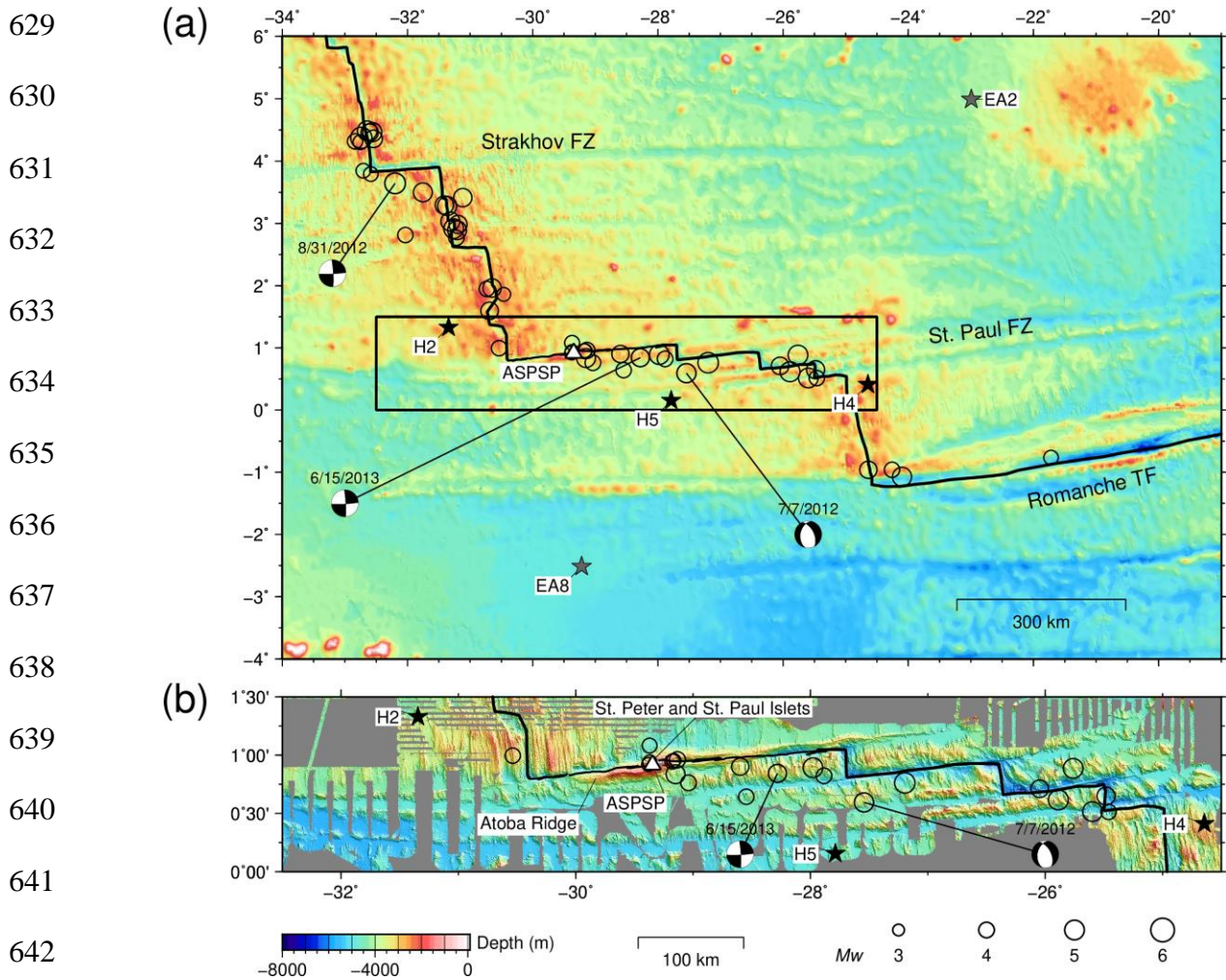
624 Black box shows location of Figure 2. **b)** Bars show instrument recording intervals: ASPSP

625 (white), COLMEIA (black), and EA (gray).

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629 **(a)**

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639 **(b)**

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645 **Figure 2. a)** Bathymetric map of equatorial Atlantic ocean. Black box shows location of (b).

646 Circles are earthquakes used in  $P_n$  analysis, scaled by  $M_w$ ; triangle shows ASPSP

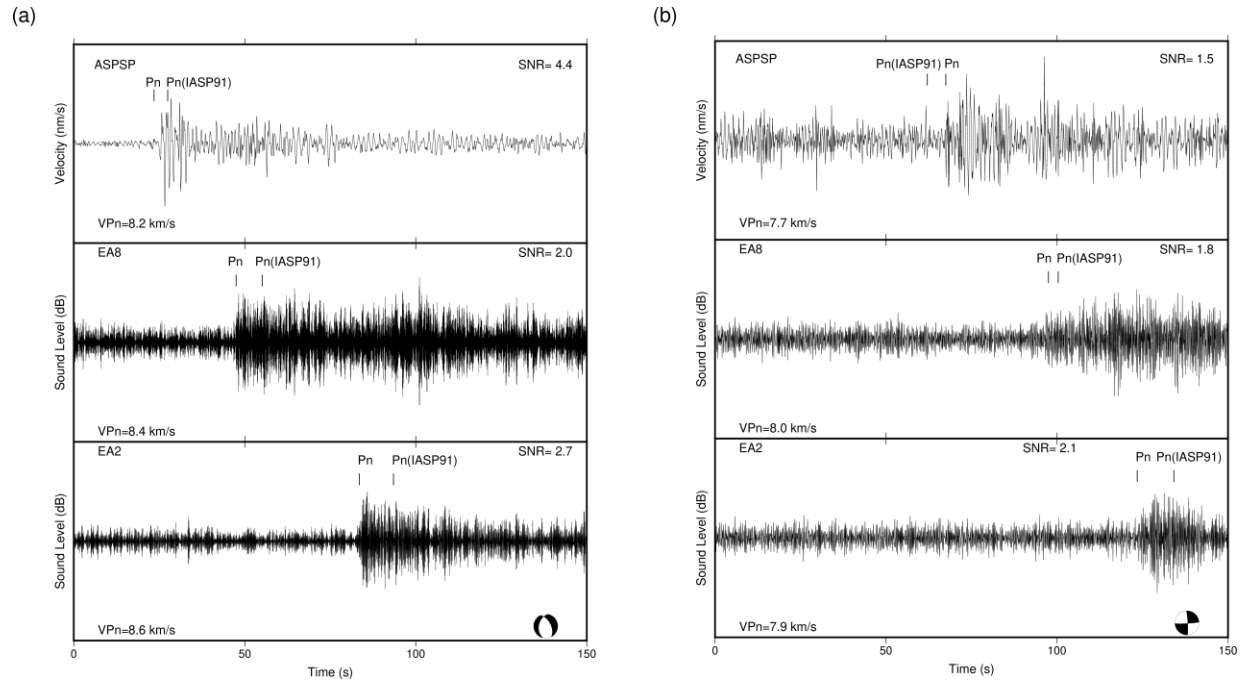
647 station; black/gray stars are COLMEIA / EA hydrophone networks, respectively (Smith *et al.*,

648 2012; Maia *et al.*, 2014); black line is Mid-Atlantic Ridge, with selected transforms labeled;

649 beach-balls are centroid moment tensors for three exemplar earthquakes (Ekström *et al.*, 2012),

650 waveforms shown in Figures 4 and 5. **b)** Bathymetric map showing details of St. Peter and St.

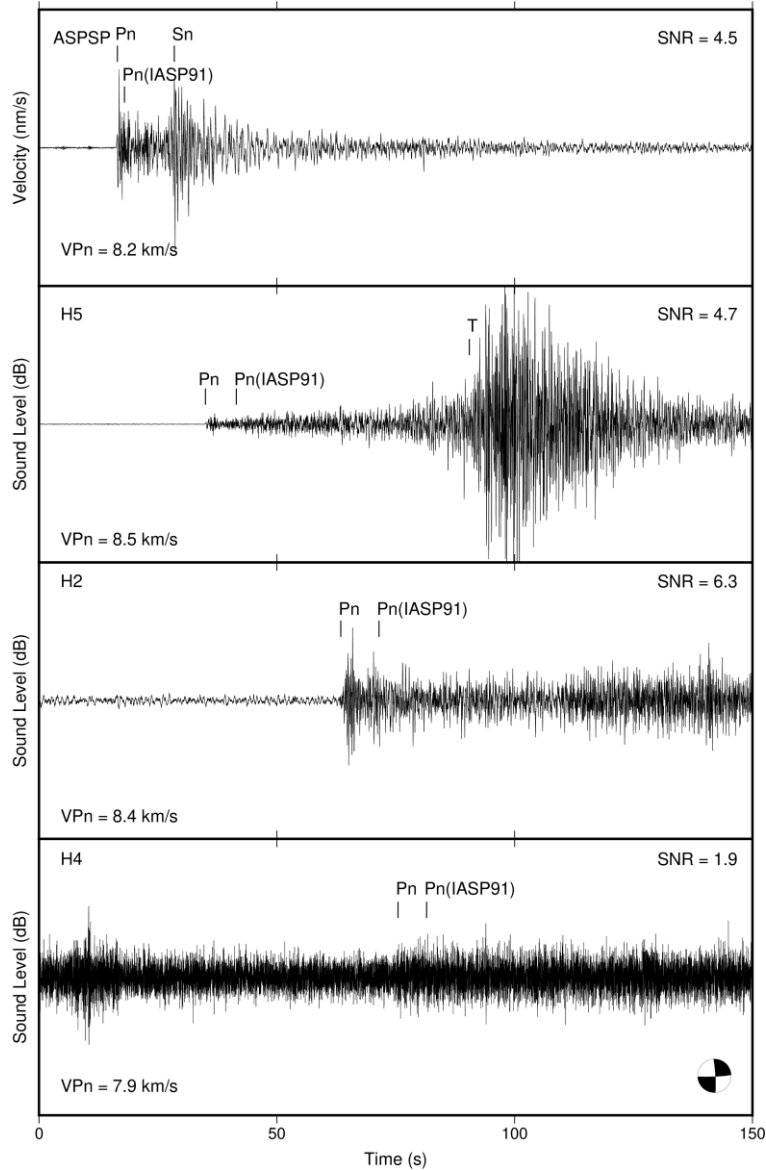
651 Paul fracture zone (from Udintsev *et al.*, 1996; Gasperini *et al.*, 1997; Maia *et al.*, 2016).



652

653 **Figure 3.** Example waveforms recorded by the ASPSP seismograph and EA array hydrophones,  
 654 with 4–12 Hz and 6–20 Hz Butterworth filters applied, respectively. **a)**  $M_w$  4.9 normal faulting  
 655 event on 7<sup>th</sup> July 2012, located on the St. Paul transform fault at 27.5°W. Picked  $P_n$  arrivals, and  
 656  $P_n$  arrivals predicted by iasp91 model are marked; beach-balls are centroid moment tensors  
 657 (Ekström et al., 2012);  $V_{P_n}$  and signal to noise ratio (SNR) noted for each station (this study),  
 658 SNR calculated STA/LTA. **b)**  $M_w$  5.3 strike-slip event on 31<sup>st</sup> August 2012, located on Strakhov  
 659 transform fault near 32.5°W.



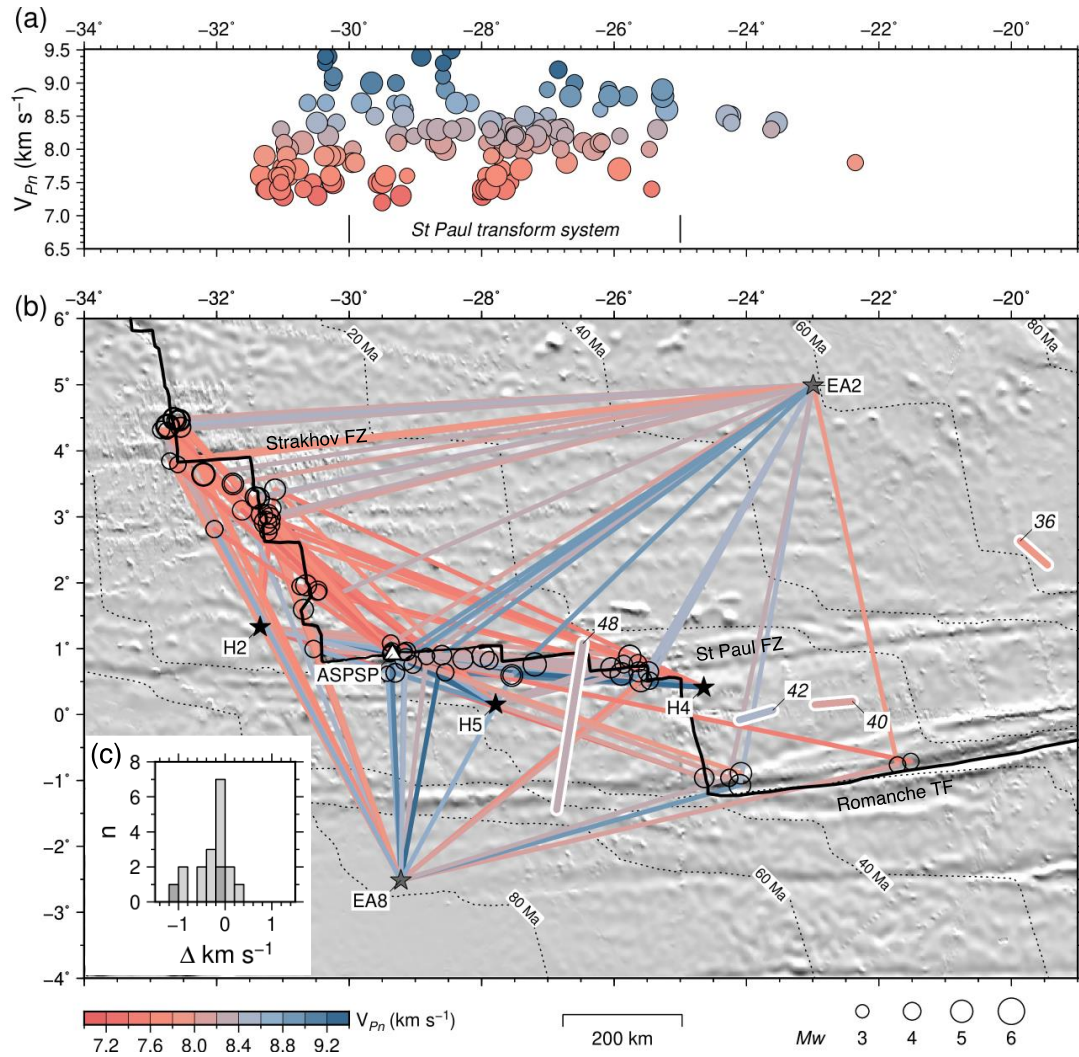


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662 **Figure 4.** Example of waveforms recorded by the ASPSP seismograph and COLMEIA  
 663 hydrophones, with 4–12 Hz and 6–20 Hz Butterworth filters applied, respectively, for mb 4.6  
 664 strike-slip event on 15<sup>th</sup> June 2013, located near St. Paul transform fault at 29.5°W. Picked  $P_n$   
 665 arrivals, and  $P_n$  arrivals predicted by iasp91 model are marked; beach-balls are centroid moment  
 666 tensors (Ekström et al., 2012);  $V_{Pn}$  and SNR noted for each station (this study).

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670 **Figure 5. a)**  $V_{Pn}$  plotted against mean longitude of ray path. Circle radius scaled by magnitude

671 of source event; colored by  $V_{Pn}$ ; St. Paul transform system marked by vertical bars. **b)** Shaded

672 relief map showing stations, earthquakes, and ray paths. Circles are earthquakes used in  $Pn$

673 analysis, scaled by  $M_w$ ; colored lines are ray paths shaded by  $Pn$  velocity; white triangle is

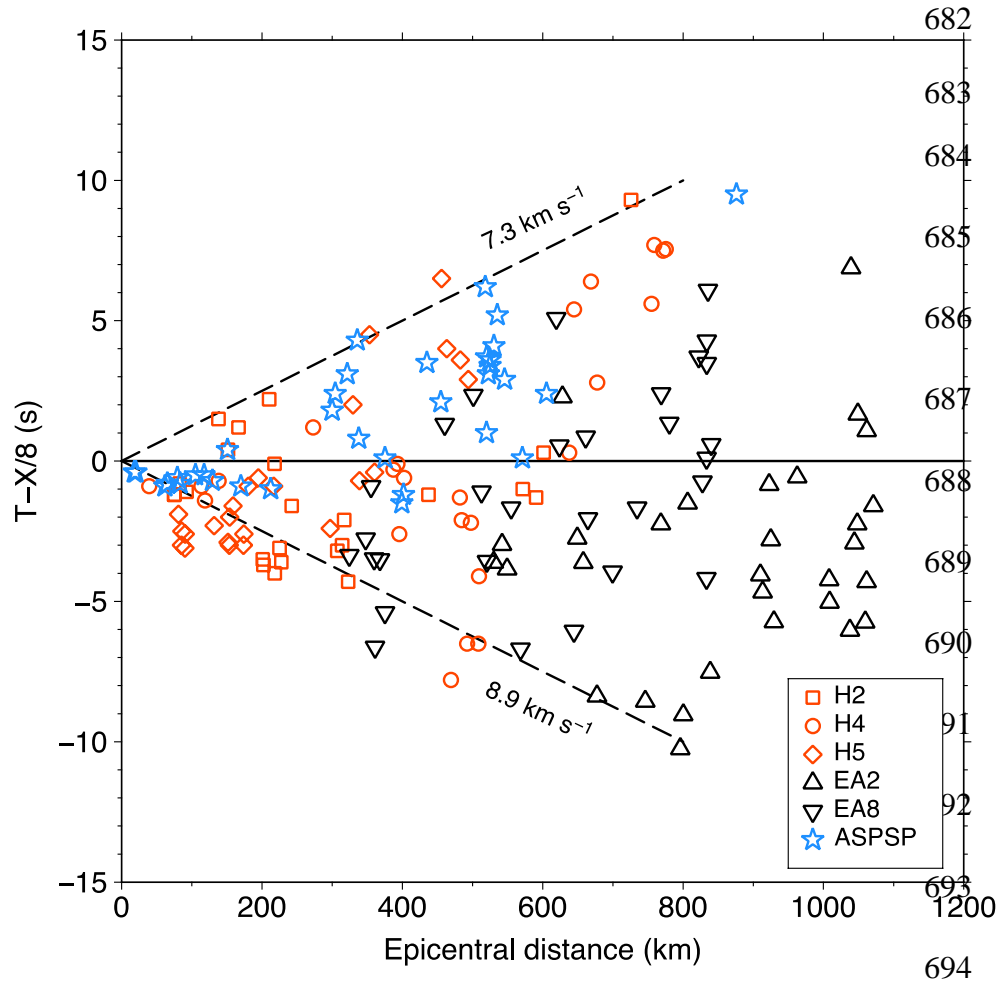
674 ASPSP station; black/gray stars are COLMEIA / EA hydrophone networks, respectively; thick

675 lines numbered 48, 42, 40, and 36 are seismic refraction profiles from cruise AT40-180 (Le

676 Pichon *et al.*, 1965), shaded by velocity; dotted lines are isochrons, modified from Müller *et al.*

677 (2008) to remove artifacts associated with fracture zone traces. **c)** Histogram of difference

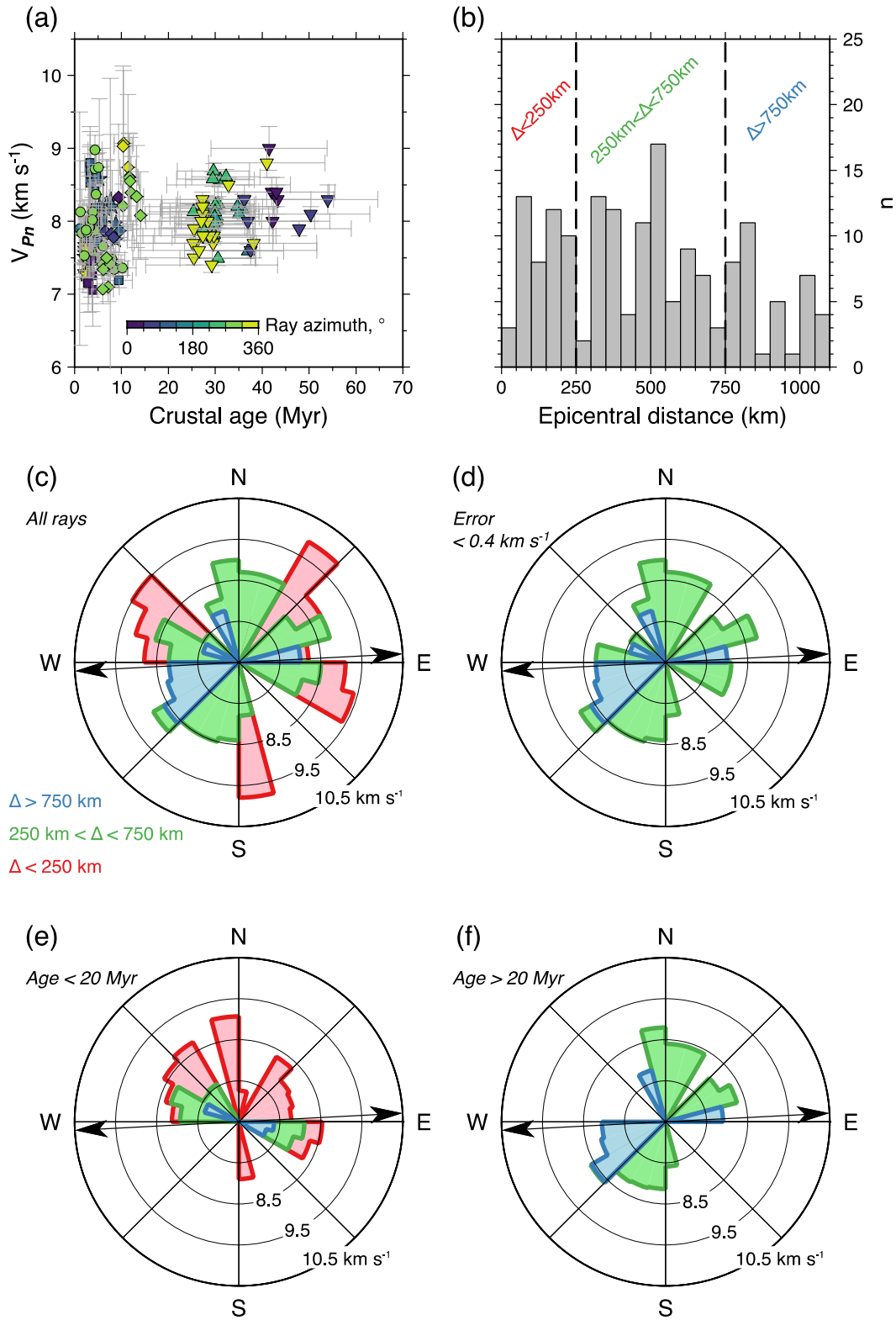
678 between velocity estimates from refraction experiment (Le Pichon *et al.*, 1965), and intersecting  
679 ray paths from this study; positive values indicate higher velocities estimated by refraction  
680 experiment; dark/light gray bars are velocities from profiles AT40-180 48 and 42, respectively.  
681



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696 **Figure 6.** Reduced travel time versus epicentral distance, plotted with a reduction velocity of 8  
 697  $\text{km s}^{-1}$ , approximately corresponding to velocity immediately below Moho from PREM and  
 698 iasp91 models (solid line; Dziewonski and Anderson, 1981; Kennett and Engdahl, 1991); dashed  
 699 lines show velocity bounds of 7.3 and 8.9  $\text{km s}^{-1}$ ; key shows recording station symbols.

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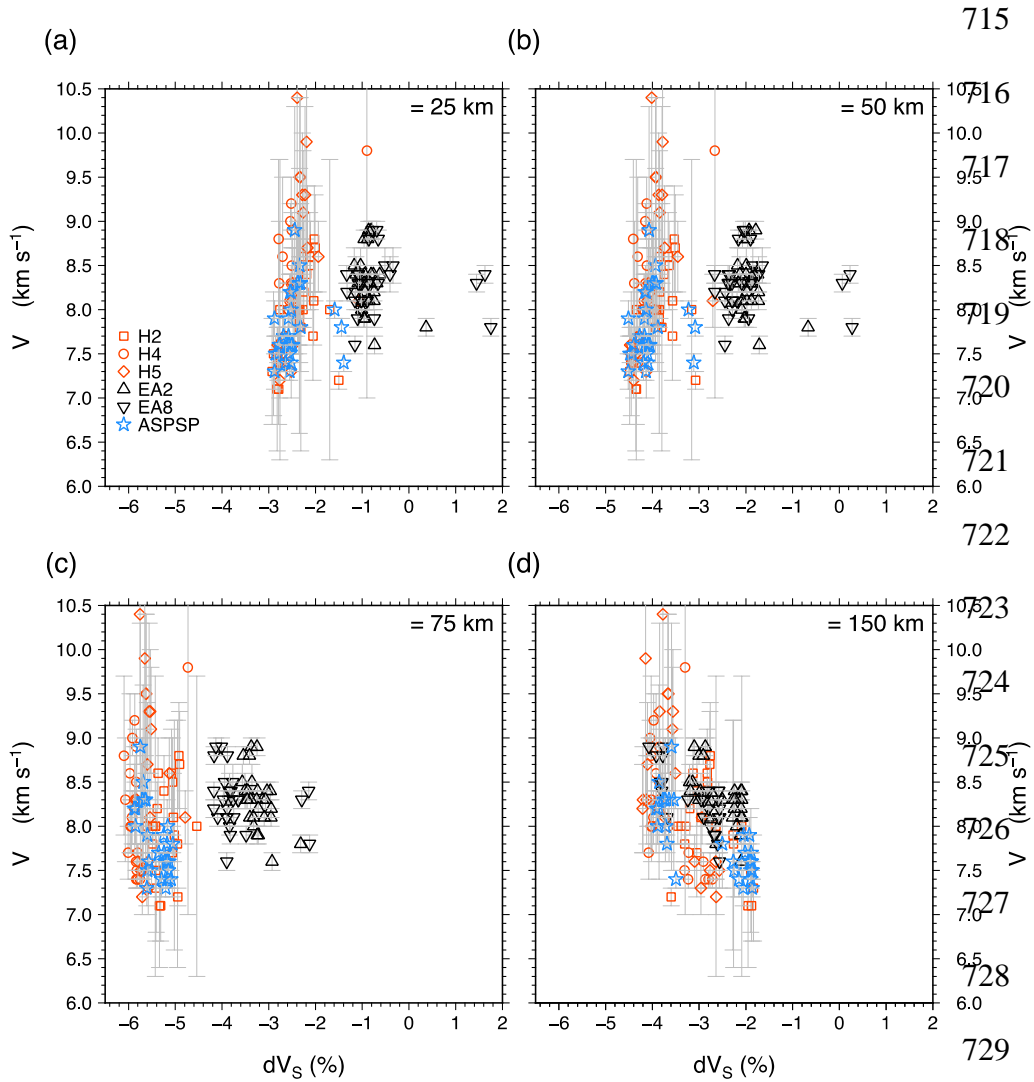


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703 **Figure 7. a)**  $V_{Pn}$  plotted against oceanic crustal age at epicentral location, colored by ray azimuth  
704 (crustal ages assigned from model of Müller *et al.*, 2008); key for station symbols given in  
705 Figure 6; horizontal error bars are  $2\sigma$  crustal age along ray path, vertical error bars are  $V_{Pn}$   
706 uncertainty described in text. **b)** Histogram of epicentral distances; dotted lines show cut-offs  
707 used to define categories in anisotropy analysis. **c)** Sector diagram showing  $V_{Pn}$  vs. azimuth for  
708 all rays; length of sectors scaled by median  $V_{Pn}$ , calculated in  $15^\circ$  bins, and colored by epicentral  
709 distance category; black arrows show plate spreading vector. **d)** Median  $V_{Pn}$  vs. azimuth for rays  
710 with  $V_{Pn}$  uncertainty  $< 0.4 \text{ km s}^{-1}$ , colored by epicentral distance category. **e)** Median  $V_{Pn}$  vs.  
711 azimuth for rays sampling crust  $< 20 \text{ Myr}$  in age, colored by epicentral distance category. **f)**  
712 Median  $V_{Pn}$  vs. azimuth for rays sampling crust  $> 20 \text{ Myr}$  in age, colored by epicentral distance  
713 category.

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730

731 **Figure 8. (a-d)** Relationship between  $V_{Pn}$  and vertically polarized tomographic shear velocity  
 732 anomaly at depths of 25, 50, 75 and 150 km, respectively, from global model SL2013sv  
 733 (Schaeffer and Lebedev, 2013).

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## **Supplemental Material for: Uppermost mantle velocity beneath the Mid-Atlantic Ridge and transform faults in the equatorial Atlantic ocean**

Guilherme W. S. de Melo, Ross Parnell-Turner, Robert P. Dziak, Deborah K. Smith, Marcia Maia, Aderson F. do Nascimento, and Jean-Yves Royer

### **Description of the Supplemental Material**

This supplemental material consists of one table, seven figures, accompanying narrative text and related references, and one Excel spreadsheet.

### **Supplemental Text**

Table S1 provides the catalog of 50 earthquakes used in  $P_n$  velocity analysis, and the Table S2 contains the parameters used for  $V_{P_n}$  estimates from individual source-receiver pairs. Figures S1 to S6 show common-receiver record sections for each station used in the study, which aid the identification of  $P_n$  phases in the waveform data. As documented elsewhere,  $P_n$  arrivals are often emergent and noisy (e.g. VanderBeek and Toomey, 2017), and hence difficult to identify when plotted as record sections, in particular when using data recorded by moored hydrophones. We also note that microseism noise obscures  $P_n$  arrivals recorded by the station ASPSP at epicentral distances  $> 350$  km, due to this station being located on the St. Peter and St. Paul islets (de Queiroz *et al.*, 2017). Figure S7 shows estimates of the signal to noise ratio (SNR), calculated for all arrivals, and used to illustrate uncertainties in arrival time picks.



## Supplemental Tables

Table S1. Catalog of 50 earthquakes used in  $Pn$  velocity analysis; magnitude estimates (de Melo and do Nascimento, 2018) hypocenters and origin times from ISC catalog. Note all depths are 10 km.

Table S2.  $V_{Pn}$  parameters for velocity estimate for each source-receiver pair, including origin times, hypocenter locations,  $Pn$  travel time,  $Pn$  distance, and  $Pn$  velocity estimates for the 50 earthquakes. They present the complete list of 152 raypaths.

## Supplemental Figures

Figures S1. Common-receiver record section for seismograph ASPSP on St Peter and St Paul islets; waveforms plotted with a 4–12 Hz Butterworth filter, amplitudes scaled to minimize overlap between adjacent traces; dashed solid/dashed lines show range of likely  $Pn$  velocities; colored triangles are  $Pn$  arrival picks.

Figures S2. Common-receiver record sections for hydrophone H2; waveforms plotted with a 6–20 Hz Butterworth filter, amplitudes scaled to minimize overlap between adjacent traces; dashed solid/dashed lines show range of likely  $Pn$  velocities; colored triangles are  $Pn$  arrival picks.

Figures S3. Common-receiver record sections for hydrophone H4; waveforms plotted with a 6–20 Hz Butterworth filter; amplitudes scaled to minimize overlap between adjacent traces; dashed solid/dashed lines show range of likely  $Pn$  velocities; colored triangles are  $Pn$  arrival picks.

Figures S4. Common-receiver record section for hydrophone H5; waveforms plotted with a 6–20 Hz Butterworth filter, amplitudes scaled to minimize overlap between adjacent traces; dashed solid/dashed lines show range of likely  $Pn$  velocities; colored triangles are  $Pn$  arrival picks.

Figures S5. Common-receiver record section for hydrophone EA2; waveforms plotted with a 6–20 Hz Butterworth filter, amplitudes scaled to minimize overlap between adjacent traces; dashed solid/dashed lines show range of likely  $Pn$  velocities; colored triangles are  $Pn$  arrival picks.

Figures S6. Common-receiver record section for hydrophone EA8; waveforms plotted with a 6–20 Hz Butterworth filter, amplitudes scaled to minimize overlap between adjacent traces; dashed solid/dashed lines show range of likely  $Pn$  velocities; colored triangles are  $Pn$  arrival picks.

Figure S7. Noise characterization of  $Pn$  arrivals. **a)** Signal to noise ratio estimated from the ratio between short time (1 s window) and long time (20 s window) average amplitudes ( $SNR_{STA/LTA}$ ), as a function of epicentral distance, key shows symbols used for stations. **b)** SNR estimated from ratio between the peak amplitude and the root mean square noise amplitude ( $SNR_{amp}$ ), as a function of epicentral distance. **c)**  $SNR_{STA/LTA}$  as a function of magnitude. **d)**  $SNR_{amp}$  as a function of magnitude. **e)**  $SNR_{STA/LTA}$  vs.  $SNR_{amp}$ , symbols with error bars are mean values of  $SNR$  for each station  $\pm 1$  standard deviation.

**Table S1.** Catalog of 50 earthquakes used in *Pn* velocity analysis; magnitude estimates (de Melo and do Nascimento, 2018) hypocenters and origin times from ISC catalog. Note all depths are 10 km.

Event	Date	Origin time, UTC	Lat, ° N	Long, ° W	Mag, Mw
1	Jul. 07. 2012	15:11:45	0.594	27.544	4.9
2	Jul. 09. 2012	15:40:22	0.836	29.149	4.8
3	Jul. 11. 2012	10:48:11	-0.961	24.259	3.9
4	Jul. 13. 2012	10:00:13	2.814	32.034	3.9
5	Jul. 14. 2012	06:47:17	3.281	31.366	4.9
6	Jul. 17. 2012	09:21:50	2.917	31.263	5.0
7	Jul. 18. 2012	05:43:30	2.984	31.193	4.6
8	Jul. 28. 2012	15:20:17	3.845	32.707	3.7
9	Jul. 28. 2012	15:23:42	4.306	32.751	3.8
10	Jul. 28. 2012	16:01:11	4.375	32.738	5.4
11	Jul. 28. 2012	16:02:33	4.459	32.604	4.4
12	Jul. 28. 2012	16:03:59	4.319	32.839	3.9
13	Jul. 28. 2012	16:12:38	3.793	32.585	3.7
14	Jul. 28. 2012	16:18:46	4.488	32.642	5.0
15	Aug. 09. 2012	08:59:23	-1.072	24.099	4.9
16	Aug. 16. 2012	08:04:57	0.514	25.461	3.9
17	Aug. 18. 2012	16:02:47	0.946	29.176	3.8
18	Aug. 22. 2012	10:19:55	4.464	32.562	4.9
19	Aug. 23. 2012	05:14:32	1.082	29.371	3.8
20	Aug. 31. 2012	00:35:35	3.644	32.199	5.3
21	Aug. 31. 2012	03:52:29	3.5	31.753	4.8
22	Sep. 19. 2012	02:26:33	-0.764	21.719	3.7
23	Sep. 23. 2012	06:29:39	1.597	30.689	4.5
24	Sep. 24. 2012	00:55:51	0.516	25.596	5.0
25	Oct. 26. 2012	14:57:30	0.901	28.599	4.4
26	Oct. 31. 2012	15:13:12	4.359	32.541	4.5
27	Nov. 11. 2012	08:02:28	3.296	31.411	4.7
28	Dec. 03. 2012	11:03:19	0.649	25.481	4.6
29	Feb. 26. 2013	06:29:28	0.762	29.039	4.0
30	Mar. 24. 2013	16:23:43	0.616	25.888	5.1
31	Apr. 01. 2013	20:01:10	0.892	27.978	5.0
32	Apr. 01. 2013	20:03:00	0.823	27.885	4.0
33	Apr. 03. 2013	05:29:36	0.761	27.197	5.2
34	Apr. 08. 2013	20:33:41	0.643	28.547	3.9
35	Apr. 09. 2013	03:07:04	3.415	31.118	4.7
36	Apr. 09. 2013	03:46:45	2.892	31.208	4.8
37	Apr. 14. 2013	04:28:40	2.769	31.218	3.9
38	Apr. 26. 2013	11:06:45	0.711	26.047	4.5
39	May. 06. 2013	21:15:49	0.94	29.37	3.6
40	May. 07. 2013	08:21:09	0.92	29.38	3.5
41	May. 21. 2013	00:51:04	-0.96	24.638	4.4
42	May. 28. 2013	22:32:39	0.96	29.134	4.1
43	May. 31. 2013	10:19:26	0.886	25.762	5.0
44	Jun. 12. 2013	03:54:05	0.955	29.165	3.6
45	Jun. 15. 2013	20:37:31	0.843	28.281	4.6
46	Jul. 20. 2013	01:59:52	1.963	30.65	4.7
47	Jul. 20. 2013	12:32:49	0.996	30.537	3.9
48	Jul. 21. 2013	14:54:12	1.949	30.734	3.9
49	Jul. 25. 2013	05:13:50	1.86	30.462	3.5
50	Aug. 09. 2013	01:26:16	3.032	31.316	4.8

All depths of these earthquakes are 10 km in the ISC catalog.

**Table S2.**  $V_{Pn}$  parameters for velocity estimate for each source-receiver pair, including origin times, hypocenter locations,  $Pn$  travel time,  $Pn$  distance, and  $Pn$  velocity estimates for the 50 earthquakes. They present the complete list of 152 raypaths.

Station	Date (ISC)	Origin Time (ISC)	Lat (ISC)	Lon (ISC)	Dept h. km (ISC)	Mag (ISC)	Great Circle Distance. km	Pn Distance. km	Travel Time. s	Pn Travel Time. s	Total error. km/s	VPn. km/s
EA2	2012-07-07	15:11:45	0.59390	-27.5438	10.0	4.9	701.2	677.4	83.5	76.3	0.1	8.9
EA8	2012-07-07	15:11:45	0.59390	-27.5438	10.0	4.9	391.1	368.1	49.0	42.5	0.3	8.7
ASPSP	2012-07-07	15:11:45	0.59390	-27.5440	10.0	4.0	183.6	170.0	22.4	20.4	0.5	8.3
EA2	2012-07-09	15:40:22	0.83630	-29.1490	10.0	4.8	824.2	800.4	98.2	91.0	0.1	8.8
EA8	2012-07-09	15:40:22	0.83630	-29.1490	10.0	4.8	370.8	347.8	47.2	40.7	0.3	8.5
ASPSP	2012-07-11	10:48:11	-0.96100	-24.2591	10.0	3.9	584.7	571.1	73.5	71.5	0.1	8.0
EA2	2012-07-11	10:48:11	-0.96100	-24.2591	10.0	3.9	673.0	649.2	85.6	78.4	0.1	8.3
EA8	2012-07-11	10:48:11	-0.96100	-24.2591	10.0	3.9	577.9	554.9	74.2	67.7	0.2	8.2
EA2	2012-07-13	10:00:13	2.81410	-32.0339	10.0	3.9	1032.0	1008.2	129.0	121.8	0.1	8.3
EA8	2012-07-13	10:00:13	2.81410	-32.0339	10.0	3.9	667.5	644.5	81.0	74.5	0.1	8.7
ASPSP	2012-07-13	10:00:13	2.81410	-32.0339	10.0	3.9	335.1	321.5	45.3	43.3	0.2	7.4
EA2	2012-07-14	06:47:17	3.28140	-31.3666	10.0	4.9	948.7	924.9	120.0	112.8	0.1	8.2
EA8	2012-07-14	06:47:17	3.28140	-31.3666	10.0	4.9	684.2	661.2	90.0	83.5	0.1	7.9
ASPSP	2012-07-14	06:47:17	3.28140	-31.3666	10.0	4.4	349.2	335.6	48.2	46.2	0.2	7.3
EA2	2012-07-17	09:21:50	2.91710	-31.2627	10.0	5.0	946.5	922.7	121.7	114.5	0.1	8.1
EA8	2012-07-17	09:21:50	2.91710	-31.2627	10.0	5.0	642.4	619.4	89.0	82.5	0.1	7.5
ASPSP	2012-07-17	09:21:50	2.91710	-31.2627	10.0	4.5	313.6	300.0	41.3	39.3	0.3	7.6
EA2	2012-07-18	05:43:30	2.98460	-31.1929	10.0	4.6	937.2	913.4	116.7	109.5	0.1	8.3
EA8	2012-07-18	05:43:30	2.98460	-31.1929	10.0	4.6	646.7	623.7	85.0	78.5	0.1	7.9
ASPSP	2012-07-18	05:43:30	2.98460	-31.1929	10.0	4.3	318.0	304.4	42.5	40.5	0.3	7.5
EA2	2012-07-28	15:20:17	3.84550	-32.7071	10.0	3.7	1085.6	1061.8	141.0	133.8	0.1	7.9
EA8	2012-07-28	15:20:17	3.84550	-32.7071	10.0	3.7	803.4	780.4	105.4	98.9	0.1	7.9

ASPSP	2012-07-28	15:20:17	3.84550	-32.7071	10.0	3.7	544.1	530.5	72.4	70.4	0.1	7.5
ASPSP	2012-07-28	15:23:42	4.30570	-32.7514	10.0	3.8	538.6	525.0	71.2	69.2	0.1	7.6
EA2	2012-07-28	15:23:42	4.30570	-32.7514	10.0	3.8	1085.4	1061.6	135.6	128.4	0.1	8.3
EA8	2012-07-28	15:23:42	4.30570	-32.7514	10.0	3.8	850.6	827.6	109.2	102.7	0.1	8.1
ASPSP	2012-07-28	16:01:11	4.37560	-32.7379	10.0	4.8	548.9	535.3	74.1	72.1	0.1	7.4
EA2	2012-07-28	16:01:11	4.37560	-32.7379	10.0	5.4	1083.3	1059.5	133.9	126.7	0.1	8.4
EA8	2012-07-28	16:01:11	4.37560	-32.7379	10.0	5.4	856.7	833.7	115.0	108.5	0.1	7.7
ASPSP	2012-07-28	16:02:33	4.45970	-32.6045	10.0	4.4	533.7	520.1	70.7	68.7	0.1	7.6
EA2	2012-07-28	16:02:33	4.45970	-32.6045	10.0	4.4	1067.9	1044.1	134.8	127.6	0.1	8.2
EA8	2012-07-28	16:02:33	4.45970	-32.6045	10.0	4.4	858.4	835.4	117.0	110.5	0.1	7.6
EA2	2012-07-28	16:03:59	4.31910	-32.8387	10.0	3.9	1094.9	1071.1	139.5	132.3	0.1	8.1
EA8	2012-07-28	16:03:59	4.31910	-32.8387	10.0	3.9	856.4	833.4	106.5	100.0	0.1	8.3
ASPSP	2012-07-28	16:03:59	4.31910	-32.8387	10.0	3.9	538.6	525.0	71.0	69.0	0.1	7.6
EA2	2012-07-28	16:12:38	3.79280	-32.5852	10.0	3.7	1072.9	1049.1	140.0	132.8	0.1	7.9
EA8	2012-07-28	16:12:38	3.79280	-32.5852	10.0	3.7	791.8	768.8	105.0	98.5	0.1	7.8
ASPSP	2012-07-28	16:12:38	3.79280	-32.5852	10.0	3.7	536.3	522.7	70.5	68.5	0.1	7.6
EA2	2012-07-28	16:18:46	4.48820	-32.6422	10.0	5.0	1071.9	1048.1	136.0	128.8	0.1	8.1
EA8	2012-07-28	16:18:46	4.48820	-32.6422	10.0	5.0	863.1	840.1	112.1	105.6	0.1	8.0
ASPSP	2012-07-28	16:18:46	4.48820	-32.6422	10.0	4.6	531.8	518.2	73.0	71.0	0.1	7.3
ASPSP	2012-08-09	08:59:23	-1.07230	-24.0997	10.0	4.8	618.9	605.3	80.1	78.1	0.1	7.8
EA2	2012-08-09	08:59:23	-1.07230	-24.0997	10.0	4.9	681.6	657.8	85.8	78.6	0.1	8.4
EA8	2012-08-09	08:59:23	-1.07230	-24.0997	10.0	4.9	591.4	568.4	70.9	64.4	0.2	8.8
EA2	2012-08-16	08:04:57	0.51390	-25.4608	10.0	3.9	566.0	542.2	72.0	64.8	0.2	8.4
EA8	2012-08-16	08:04:57	0.51390	-25.4608	10.0	3.9	535.8	512.8	69.5	63.0	0.2	8.1
ASPSP	2012-08-16	08:04:57	0.51390	-25.4608	10.0	3.9	412.9	399.3	50.4	48.4	0.2	8.2
EA2	2012-08-18	16:02:47	0.94600	-29.1760	10.0	3.8	819.9	796.1	96.5	89.3	0.1	8.9
EA8	2012-08-18	16:02:47	0.94600	-29.1760	10.0	3.8	382.8	359.8	48.0	41.5	0.3	8.7
ASPSP	2012-08-22	10:19:55	4.46400	-32.5619	10.0	4.4	559.3	545.7	73.1	71.1	0.1	7.7
EA2	2012-08-22	10:19:55	4.46400	-32.5619	10.0	4.9	1063.2	1039.4	144.0	136.8	0.1	7.6
EA8	2012-08-22	10:19:55	4.46400	-32.5619	10.0	4.9	856.8	833.8	114.2	107.7	0.1	7.7
EA2	2012-08-23	05:14:32	1.08200	-29.3710	10.0	3.8	830.2	806.4	106.5	99.3	0.1	8.1

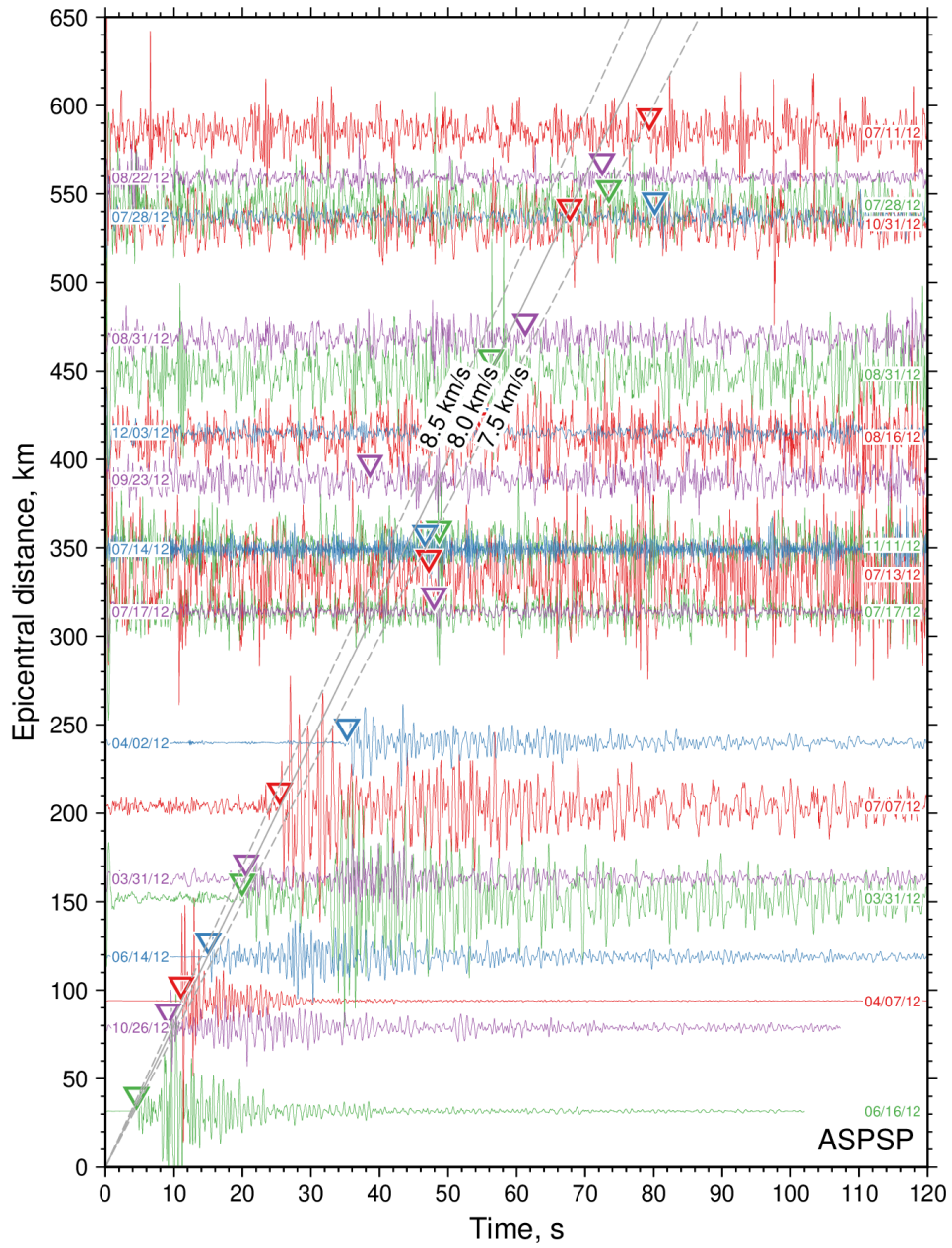
EA8	2012-08-23	05:14:32	1.08200	-29.3710	10.0	3.8	398.2	375.2	48.0	41.5	0.3	9.0
EA2	2012-08-31	00:35:35	3.64410	-32.1987	10.0	5.3	1032.6	1008.8	128.3	121.1	0.1	8.3
EA8	2012-08-31	00:35:35	3.64410	-32.1987	10.0	5.3	757.6	734.6	96.7	90.2	0.1	8.1
ASPSP	2012-08-31	00:35:35	3.64410	-32.1987	10.0	4.8	468.5	454.9	61.0	59.0	0.2	7.7
EA2	2012-08-31	03:52:29	3.50060	-31.7535	10.0	4.8	986.4	962.6	127.0	119.8	0.1	8.0
EA8	2012-08-31	03:52:29	3.50060	-31.7535	10.0	4.8	722.6	699.6	90.0	83.5	0.1	8.4
ASPSP	2012-08-31	03:52:29	3.50060	-31.7535	10.0	4.0	448.6	435.0	59.9	57.9	0.2	7.5
EA2	2012-09-19	02:26:33	-0.76440	-21.7191	10.0	3.7	652.0	628.2	88.0	80.8	0.1	7.8
EA8	2012-09-19	02:26:33	-0.76440	-21.7191	10.0	3.7	856.6	833.6	110.8	104.3	0.1	8.0
ASPSP	2012-09-19	02:26:33	-0.71000	-21.5190	10.0	3.7	889.6	876.0	121.0	119.0	0.1	7.4
EA2	2012-09-23	06:22:39	1.59680	-30.6888	10.0	4.5	933.9	910.1	116.9	109.7	0.1	8.3
EA8	2012-09-23	06:22:39	1.59680	-30.6888	10.0	4.5	483.3	460.3	65.4	58.9	0.2	7.8
ASPSP	2012-09-23	06:22:39	1.59680	-30.6888	10.0	4.3	164.6	151.0	21.3	19.3	0.5	7.8
ASPSP	2012-09-24	00:55:51	0.51650	-25.5963	10.0	4.0	388.8	375.2	49.0	47.0	0.2	8.0
EA2	2012-09-24	00:55:51	0.51650	-25.5963	10.0	5.0	573.2	549.4	72.0	64.8	0.2	8.5
EA8	2012-09-24	00:55:51	0.51650	-25.5963	10.0	5.0	524.3	501.3	71.5	65.0	0.2	7.7
EA2	2012-10-26	14:57:30	0.90120	-28.5988	10.0	4.4	769.9	746.1	91.9	84.7	0.1	8.8
EA8	2012-10-26	14:57:30	0.90120	-28.5988	10.0	4.4	384.0	361.0	45.0	38.5	0.3	9.4
ASPSP	2012-10-26	14:57:30	0.90120	-28.5988	10.0	4.4	78.7	65.1	9.3	7.3	1.4	8.9
ASPSP	2012-10-31	15:13:12	4.35950	-32.5410	10.0	4.4	533.4	519.8	68.0	66.0	0.2	7.9
EA2	2012-10-31	15:13:12	4.35950	-32.5410	10.0	4.5	1061.6	1037.8	130.9	123.7	0.1	8.4
EA8	2012-10-31	15:13:12	4.35950	-32.5410	10.0	4.5	845.4	822.4	113.0	106.5	0.1	7.7
ASPSP	2012-11-11	08:02:28	3.29600	-31.4112	10.0	4.5	351.4	337.8	45.0	43.0	0.2	7.9
EA2	2012-11-11	08:02:28	3.29600	-31.4112	10.0	4.7	953.2	929.4	117.7	110.5	0.1	8.4
EA8	2012-11-11	08:02:28	3.29600	-31.4112	10.0	4.7	687.4	664.4	87.5	81.0	0.1	8.2
ASPSP	2012-12-03	11:03:19	0.64900	-25.6250	10.0	4.5	415.2	401.6	51.0	49.0	0.2	8.2
EA2	2012-12-03	11:03:19	0.64940	-25.4813	10.0	4.6	554.0	530.2	69.9	62.7	0.2	8.5
EA8	2012-12-03	11:03:19	0.64940	-25.4813	10.0	4.6	543.6	520.6	68.0	61.5	0.2	8.5
H2	2013-02-26	06:29:28	0.76200	-29.0390	10.0	4.0	264.2	242.4	34.2	28.7	0.5	8.4
H4	2013-02-26	06:29:28	0.76200	-29.0390	10.0	4.0	490.8	469.5	55.9	50.9	0.3	9.2
H5	2013-02-26	06:29:28	0.76200	-29.0390	10.0	4.0	154.6	131.8	20.5	14.2	1.0	9.3

H2	2013-03-24	16:23:43	0.61600	-25.8880	10.0	5.1	612.4	590.6	78.0	72.5	0.2	8.1
H4	2013-03-24	16:23:43	0.61600	-25.8880	10.0	5.1	140.3	119.0	18.5	13.5	0.9	8.8
H5	2013-03-24	16:23:43	0.61600	-25.8880	10.0	5.1	217.5	194.7	30.0	23.7	0.5	8.2
ASPSP	2013-04-01	20:01:10	0.89200	-27.9790	10.0	5.0	131.0	117.4	16.2	14.2	0.7	8.3
H2	2013-04-01	20:01:10	0.89170	-27.9786	10.0	5.0	344.5	322.7	41.5	36.0	0.4	9.0
H4	2013-04-01	20:01:10	0.89170	-27.9786	10.0	5.0	408.5	387.2	53.1	48.1	0.3	8.0
H5	2013-04-01	20:01:10	0.89200	-27.9790	10.0	5.0	108.0	85.2	14.0	7.7	1.9	11.1
ASPSP	2013-04-01	20:03:00	0.82300	-27.8850	10.0	4.0	143.0	129.4	17.5	15.5	0.7	8.3
H2	2013-04-01	20:03:00	0.82300	-27.8850	10.0	4.0	338.8	317.0	43.0	37.5	0.3	8.5
H4	2013-04-01	20:03:00	0.82300	-27.8850	10.0	4.0	414.3	393.0	54.0	49.0	0.3	8.0
H5	2013-04-01	20:03:00	0.82300	-27.8850	10.0	4.0	113.0	90.2	14.5	8.2	1.8	11.0
ASPSP	2013-04-03	05:29:36	0.76100	-27.1970	10.0	5.2	226.0	212.4	27.6	25.6	0.4	8.3
H2	2013-04-03	05:29:36	0.76100	-27.1970	10.0	5.2	459.2	437.4	59.0	53.5	0.2	8.2
H4	2013-04-03	05:29:36	0.76100	-27.1970	10.0	5.2	294.2	272.9	40.3	35.3	0.3	7.7
H5	2013-04-03	05:29:36	0.91000	-27.1970	10.0	5.2	103.8	81.0	14.5	8.2	1.7	9.9
ASPSP	2013-04-08	20:33:41	0.64300	-28.5470	10.0	3.9	93.0	79.4	11.3	9.3	1.1	8.5
H2	2013-04-08	20:33:41	0.64300	-28.5470	10.0	3.9	239.6	217.8	32.6	27.1	0.5	8.0
H4	2013-04-08	20:33:41	0.64300	-28.5470	10.0	3.9	513.6	492.3	60.0	55.0	0.2	9.0
H5	2013-04-08	20:33:41	0.64300	-28.5470	10.0	3.9	181.5	158.7	24.5	18.2	0.7	8.7
H2	2013-04-09	03:07:04	3.41500	-31.1180	10.0	4.7	231.9	210.1	34.0	28.5	0.4	7.4
H4	2013-04-09	03:07:04	3.41500	-31.1180	10.0	4.7	793.0	771.7	109.0	104.0	0.1	7.4
H5	2013-04-09	03:07:04	3.41500	-31.1180	10.0	4.7	516.9	494.1	71.0	64.7	0.2	7.6
H2	2013-04-09	03:46:45	2.89200	-31.2080	10.0	4.8	173.4	151.6	24.8	19.3	0.7	7.9
H4	2013-04-09	03:46:45	2.89200	-31.2080	10.0	4.8	780.1	758.8	107.5	102.5	0.1	7.4
H5	2013-04-09	03:46:45	2.89200	-31.2080	10.0	4.8	486.2	463.4	68.3	62.0	0.2	7.5
H2	2013-04-14	04:28:40	2.76900	-31.2180	10.0	3.9	159.8	138.0	24.2	18.7	0.7	7.4
H4	2013-04-14	04:28:40	2.76900	-31.2180	10.0	3.9	776.6	755.3	105.0	100.0	0.1	7.6
H5	2013-04-14	04:28:40	2.76900	-31.2180	10.0	3.9	478.8	456.0	69.8	63.5	0.2	7.2
H2	2013-04-26	11:06:45	0.71100	-26.0470	10.0	4.5	593.6	571.8	76.0	70.5	0.2	8.1
H4	2013-04-26	11:06:45	0.71100	-26.0470	10.0	4.5	159.7	138.4	21.6	16.6	0.7	8.3
H5	2013-04-26	11:06:45	0.71100	-26.0470	10.0	4.5	203.2	180.4	28.0	21.7	0.6	8.3

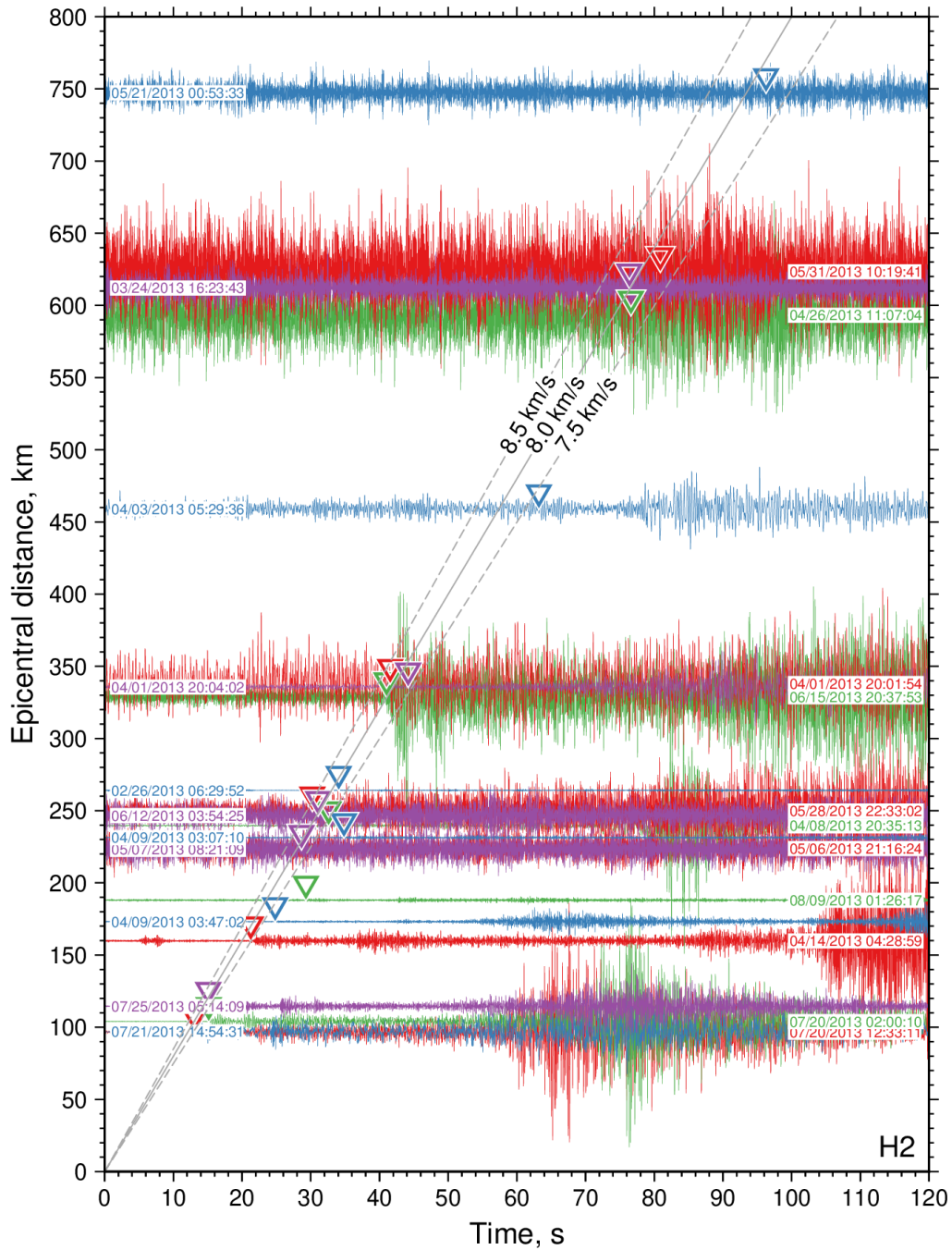
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H4	2013-05-06	21:15:49	0.94000	-29.3700	10.0	3.6	529.6	508.3	62.0	57.0	0.2	8.9
H5	2013-05-06	21:15:49	0.94000	-29.3700	10.0	3.6	196.4	173.6	25.0	18.7	0.7	9.3
H2	2013-05-07	08:21:09	0.92000	-29.3800	10.0	3.5	222.9	201.1	27.1	21.6	0.6	9.3
H4	2013-05-07	08:21:09	0.92000	-29.3800	10.0	3.5	530.6	509.3	64.6	59.6	0.2	8.5
H5	2013-05-07	08:21:09	0.92000	-29.3800	10.0	3.5	196.9	174.1	25.5	19.2	0.7	9.1
H2	2013-05-21	00:51:04	-0.96000	-24.6380	10.0	4.4	747.5	725.7	105.5	100.0	0.1	7.3
H4	2013-05-21	00:51:04	-0.96000	-24.6380	10.0	4.4	60.5	39.2	9.0	4.0	2.8	9.8
H5	2013-05-21	00:51:04	-0.96000	-24.6380	10.0	4.4	361.7	338.9	48.0	41.7	0.3	8.1
H2	2013-05-28	22:32:39	0.96000	-29.1340	10.0	4.1	249.4	227.6	30.4	24.9	0.5	9.1
H4	2013-05-28	22:32:39	0.96000	-29.1340	10.0	4.1	503.5	482.2	64.0	59.0	0.2	8.2
H5	2013-05-28	22:32:39	0.96000	-29.1340	10.0	4.1	174.3	151.5	22.3	16.0	0.9	9.5
H2	2013-05-31	10:19:26	0.88600	-25.7620	10.0	5.0	623.2	601.4	81.0	75.5	0.2	8.0
H4	2013-05-31	10:19:26	0.88600	-25.7620	10.0	5.0	135.1	113.8	18.3	13.3	0.9	8.6
H5	2013-05-31	10:19:26	0.88600	-25.7620	10.0	5.0	239.5	216.7	32.5	26.2	0.5	8.3
H2	2013-06-12	03:54:05	0.95500	-29.1650	10.0	3.6	246.8	225.0	30.5	25.0	0.5	9.0
H4	2013-06-12	03:54:05	0.95500	-29.1650	10.0	3.6	506.1	484.8	63.5	58.5	0.2	8.3
H5	2013-06-12	03:54:05	0.95500	-29.1650	10.0	3.6	176.2	153.4	22.5	16.2	0.9	9.5
ASPSP	2013-06-15	20:37:31	0.84300	-28.2810	10.0	4.6	118.8	105.2	14.7	12.7	0.8	8.3
H2	2013-06-15	20:37:31	0.84300	-28.2810	10.0	4.6	329.1	307.3	40.7	35.2	0.4	8.7
H4	2013-06-15	20:37:31	0.84300	-28.2810	10.0	4.6	423.7	402.4	54.7	49.7	0.3	8.1
H5	2013-06-15	20:37:31	0.84300	-28.2810	10.0	4.6	113.3	90.5	15.0	8.7	1.7	10.4
H2	2013-07-20	01:59:52	1.96300	-30.6500	10.0	4.7	104.4	82.6	15.0	9.5	1.4	8.7
H4	2013-07-20	01:59:52	1.96300	-30.6500	10.0	4.7	690.1	668.8	95.0	90.0	0.1	7.4
H5	2013-07-20	01:59:52	1.96300	-30.6500	10.0	4.7	376.1	353.3	55.0	48.7	0.3	7.3
H2	2013-07-20	12:32:49	0.99600	-30.5370	10.0	3.9	97.1	75.3	13.7	8.2	1.7	9.2
H4	2013-07-20	12:32:49	0.99600	-30.5370	10.0	3.9	659.2	637.9	85.0	80.0	0.2	8.0
H5	2013-07-20	12:32:49	0.99600	-30.5370	10.0	3.9	319.9	297.1	41.0	34.7	0.4	8.6
H2	2013-07-21	14:54:12	1.94900	-30.7340	10.0	3.9	96.5	74.7	13.6	8.1	1.7	9.2
H4	2013-07-21	14:54:12	1.94900	-30.7340	10.0	3.9	698.8	677.5	92.5	87.5	0.1	7.7
H5	2013-07-21	14:54:12	1.94900	-30.7340	10.0	3.9	383.3	360.5	51.0	44.7	0.3	8.1



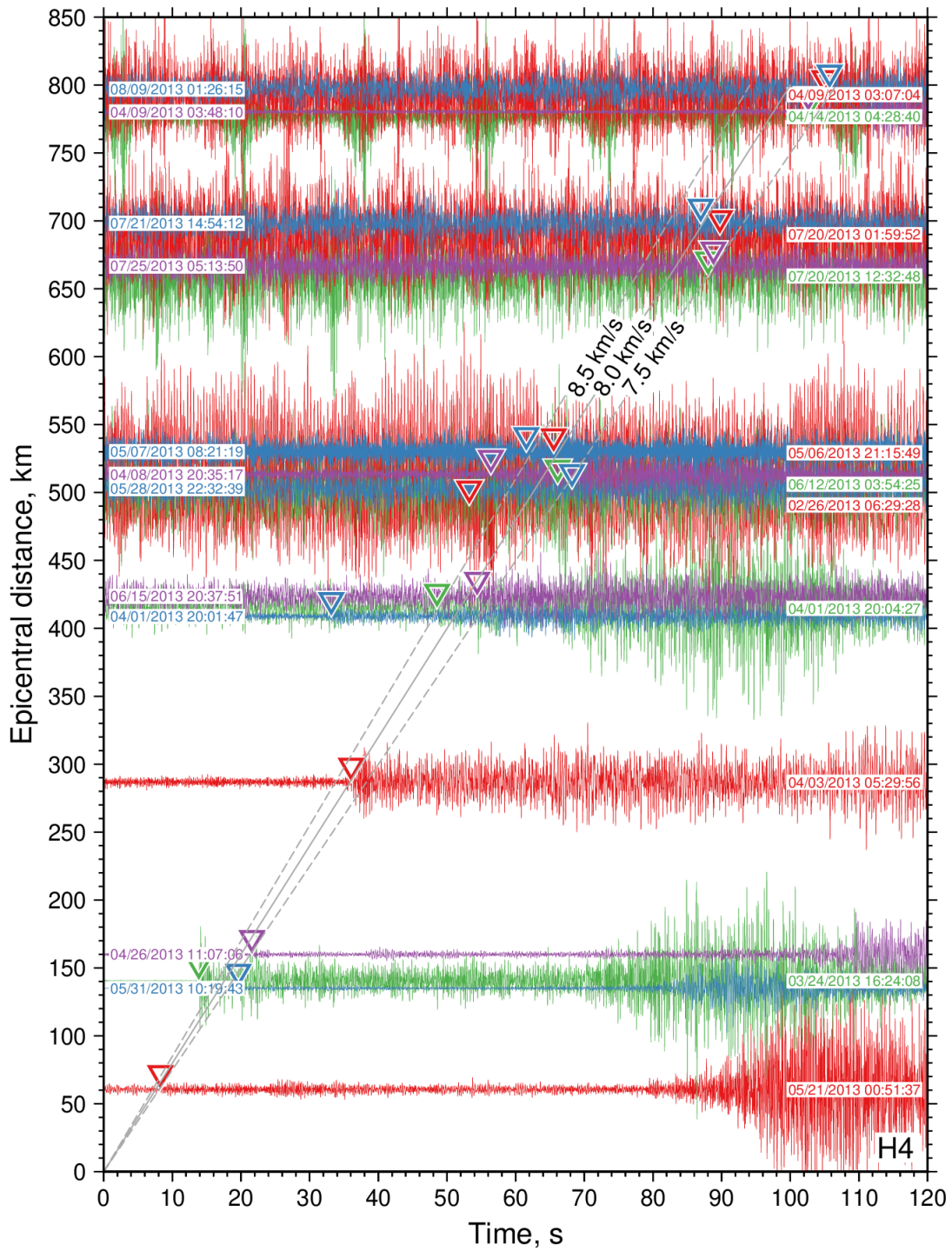
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H4	2013-07-25	05:13:50	1.86000	-30.4620	10.0	3.5	666.0	644.7	91.0	86.0	0.1	7.5
H5	2013-07-25	05:13:50	1.86000	-30.4620	10.0	3.5	352.3	329.5	49.5	43.2	0.3	7.6
H2	2013-08-09	01:26:16	3.03200	-31.3160	12.0	4.8	188.3	166.5	27.5	22.0	0.6	7.6
H4	2013-08-09	01:26:16	3.03200	-31.3160	12.0	4.8	796.9	775.6	109.5	104.5	0.1	7.4
H5	2013-08-09	01:26:16	3.03200	-31.3160	12.0	4.8	505.3	482.5	70.3	64.0	0.2	7.5



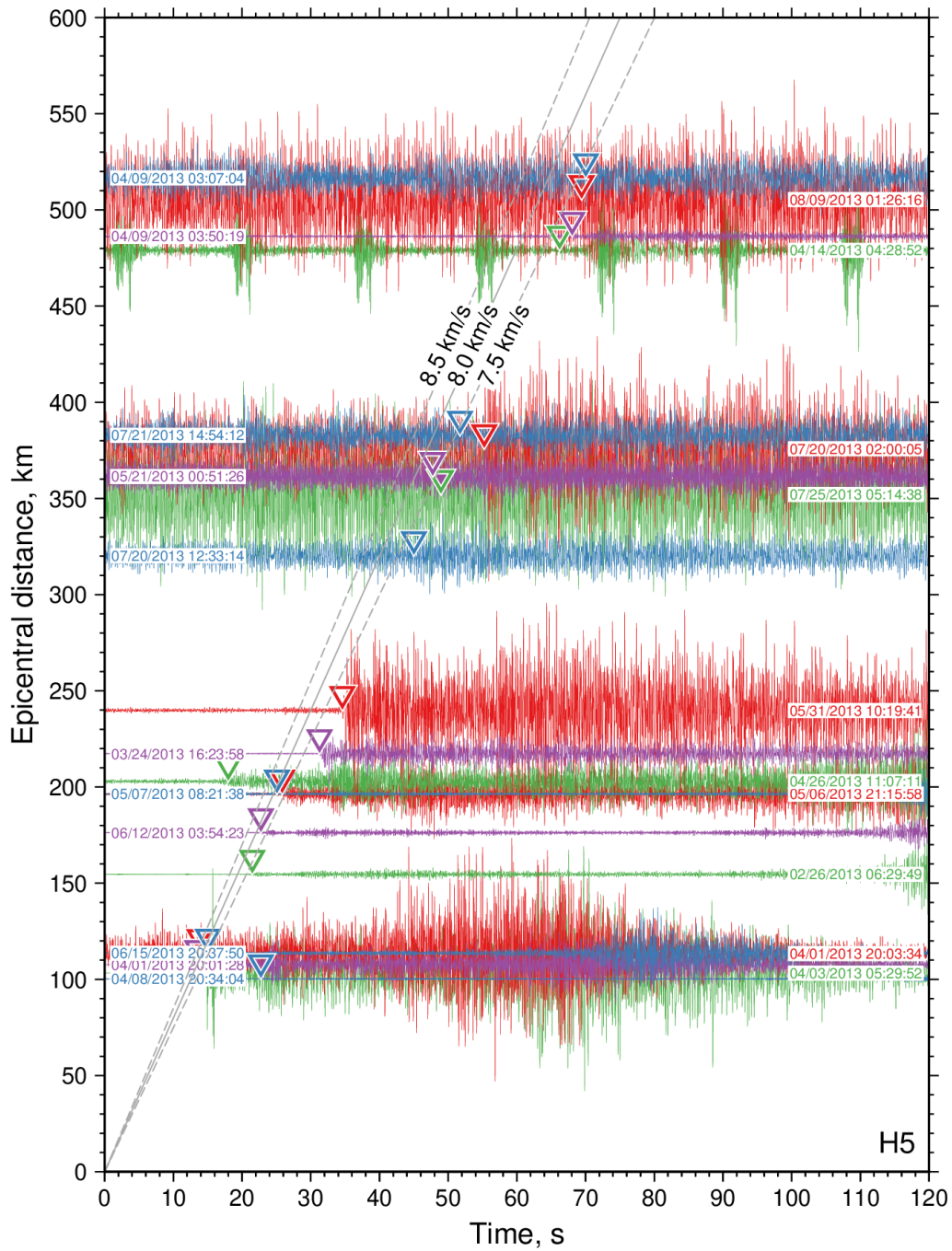
**Figures S1.** Common-receiver record section for seismograph ASPSP on St Peter and St Paul islets; waveforms plotted with a 4–12 Hz Butterworth filter, amplitudes scaled to minimize overlap between adjacent traces; dashed solid/dashed lines show range of likely  $P_n$  velocities; colored triangles are  $P_n$  arrival picks.



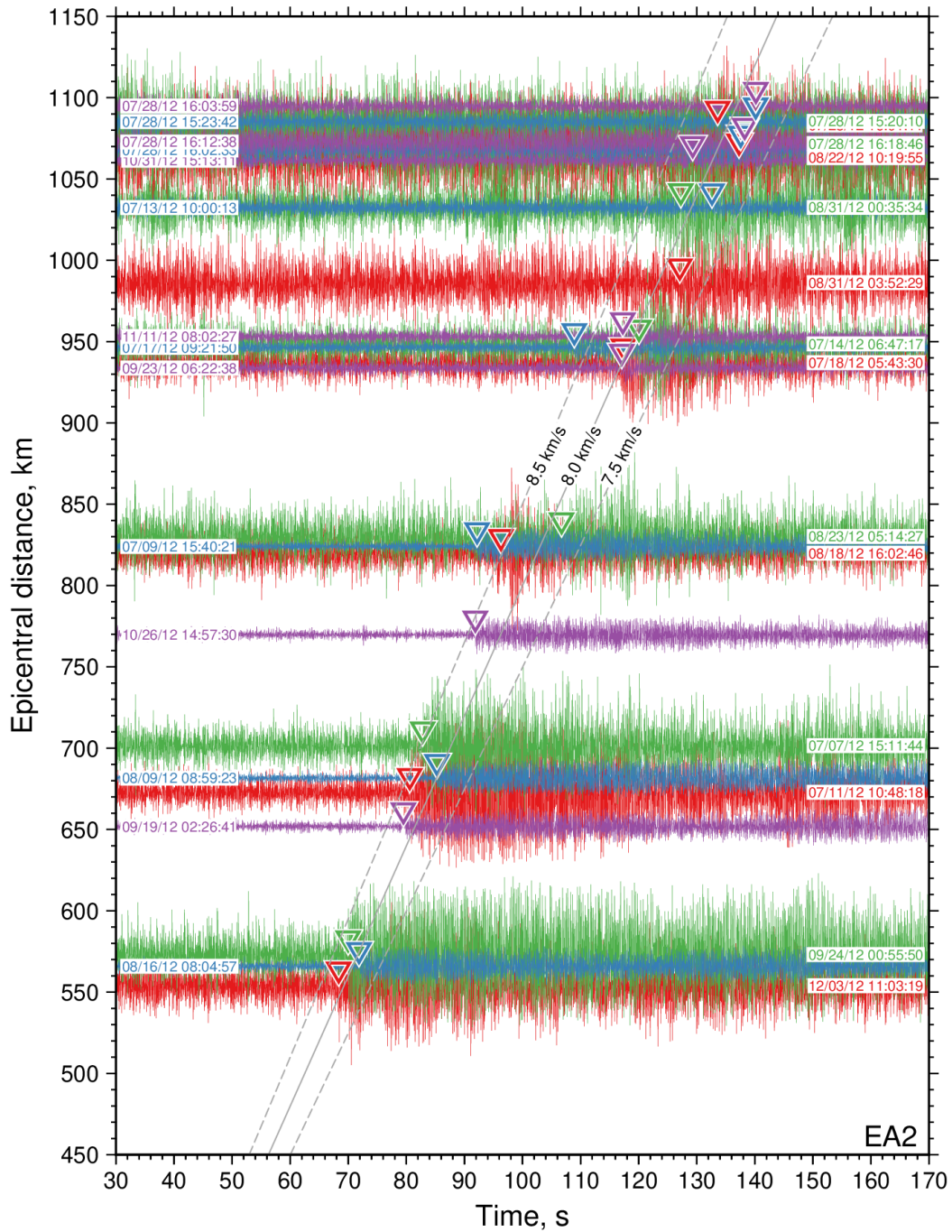
**Figures S2.** Common-receiver record sections for hydrophone H2; waveforms plotted with a 6–20 Hz Butterworth filter, amplitudes scaled to minimize overlap between adjacent traces; dashed solid/dashed lines show range of likely  $P_n$  velocities; colored triangles are  $P_n$  arrival picks.



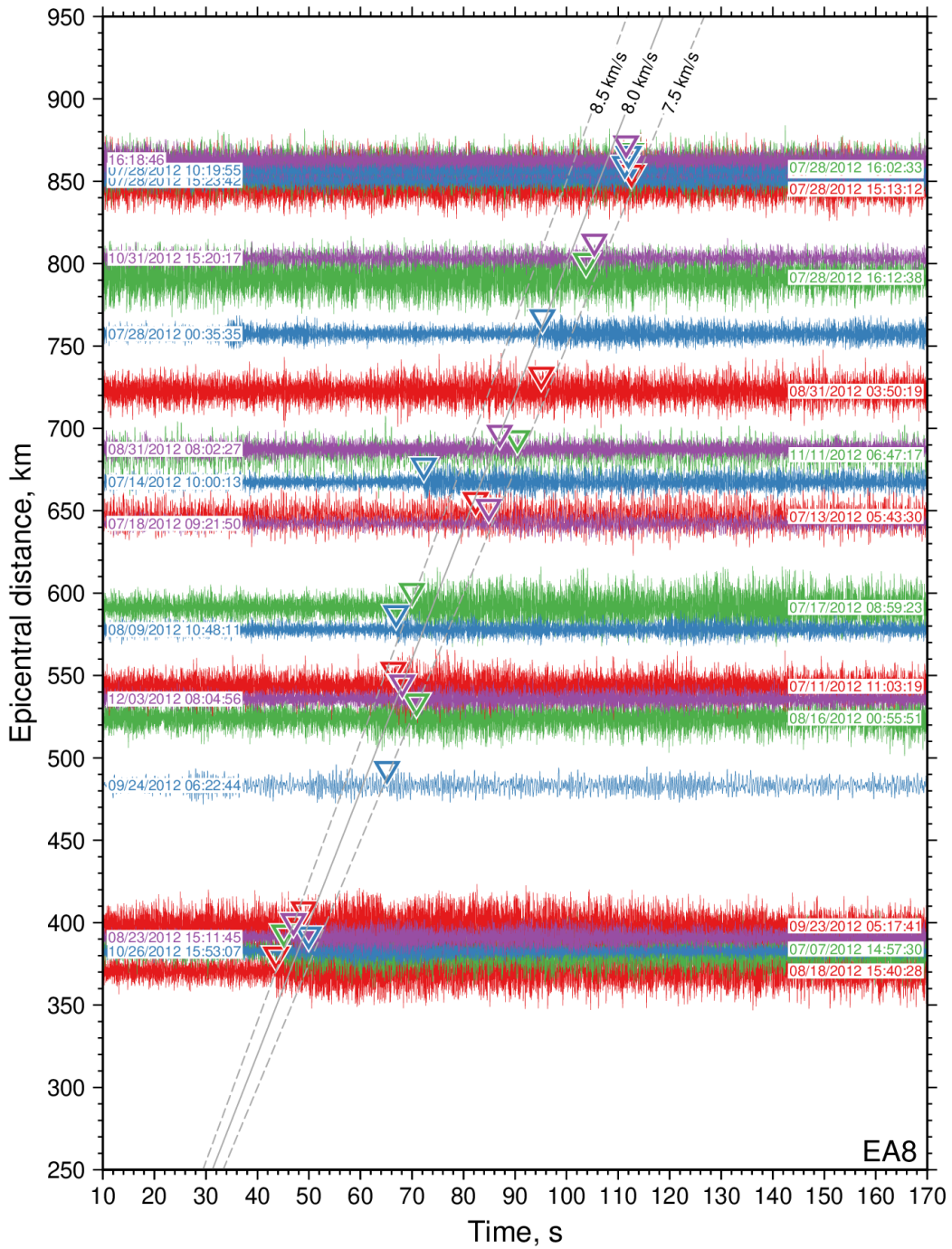
**Figures S3.** Common-receiver record sections for hydrophone H4; waveforms plotted with a 6–20 Hz Butterworth filter; amplitudes scaled to minimize overlap between adjacent traces; dashed solid/dashed lines show range of likely  $P_n$  velocities; colored triangles are  $P_n$  arrival picks.



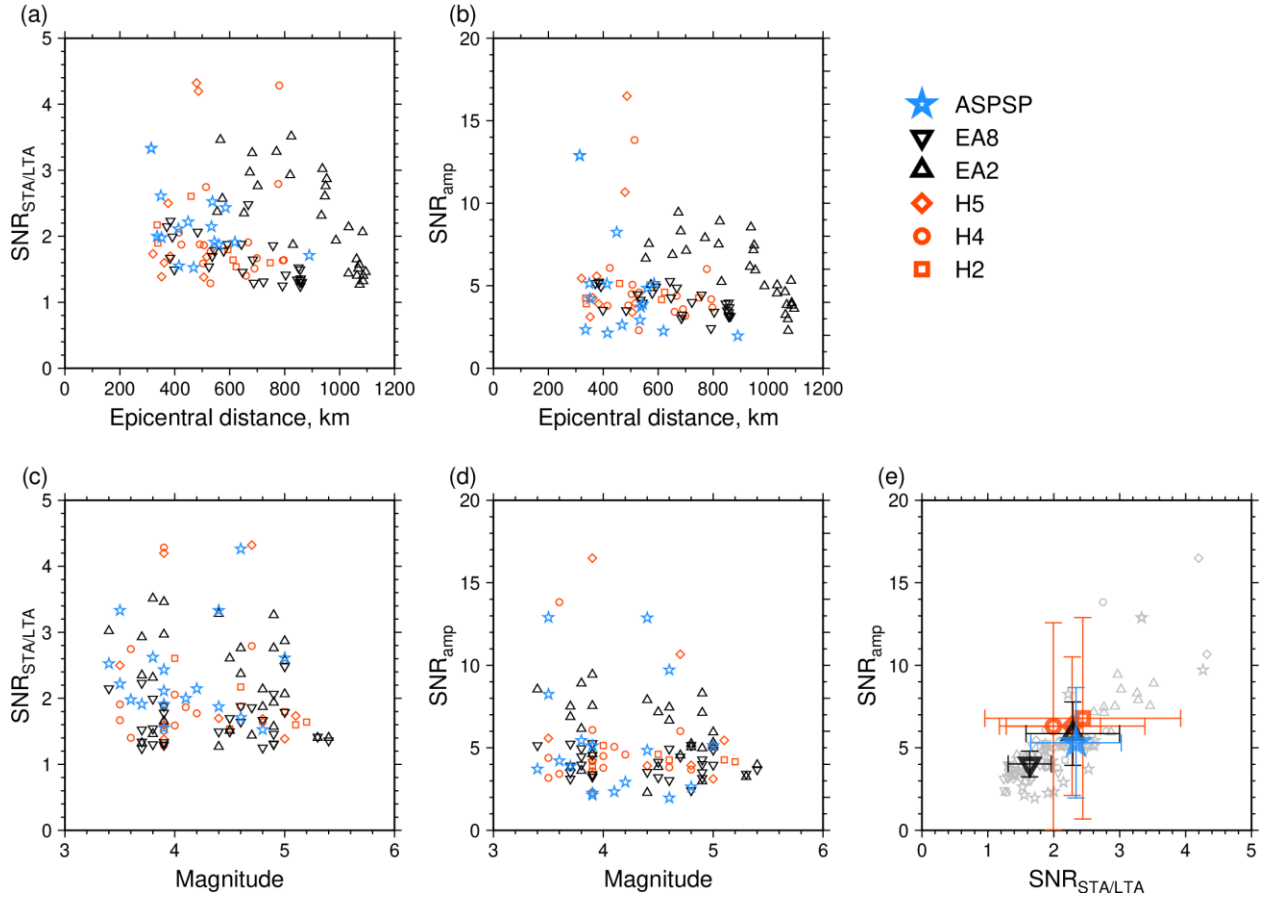
**Figures S4.** Common-receiver record section for hydrophone H5; waveforms plotted with a 6–20 Hz Butterworth filter, amplitudes scaled to minimize overlap between adjacent traces; dashed solid/dashed lines show range of likely  $P_n$  velocities; colored triangles are  $P_n$  arrival picks.



**Figures S5.** Common-receiver record section for hydrophone EA2; waveforms plotted with a 6–20 Hz Butterworth filter, amplitudes scaled to minimize overlap between adjacent traces; dashed solid/dashed lines show range of likely  $P_n$  velocities; colored triangles are  $P_n$  arrival picks.



**Figures S6.** Common-receiver record section for hydrophone EA8; waveforms plotted with a 6–20 Hz Butterworth filter, amplitudes scaled to minimize overlap between adjacent traces; dashed solid/dashed lines show range of likely  $P_n$  velocities; colored triangles are  $P_n$  arrival picks.



**Figure S7.** Noise characterization of  $P_n$  arrivals. **a)** Signal to noise ratio estimated from the ratio between short time (1 s window) and long time (20 s window) average amplitudes ( $SNR_{STA/LTA}$ ), as a function of epicentral distance, key shows symbols used for stations. **b)** SNR estimated from ratio between the peak amplitude and the root mean square noise amplitude ( $SNR_{amp}$ ), as a function of epicentral distance. **c)**  $SNR_{STA/LTA}$  as a function of magnitude. **d)**  $SNR_{amp}$  as a function of magnitude. **e)**  $SNR_{STA/LTA}$  vs.  $SNR_{amp}$ , symbols with error bars are mean values of SNR for each station  $\pm 1$  standard deviation.



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