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Sorption of rare earth elements on schwertmannite and their mobility in acid mine drainage treatments

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Abstract

Rare earth elements (REE) are currently considered critical raw materials due to their increasing use in modern technologies and their shortage of supply. Acid mine drainage (AMD) contains REE concentrations that are several orders of magnitude higher than those in continental and marine waters, and the sludge resulting from AMD treatment can become a supplementary source of REE. Schwertmannite, an Fe(III)-sulfate-hydroxide, is the most common mineral precipitated from AMD and is a main constituent of treatment sludge. The objective of this work was to study the mechanisms of REE retention in schwertmannite.

Suspensions of synthetic schwertmannite in sulfate solutions showed that Y and the lanthanides are effectively sorbed at pH values higher than 4.5, and sorption is complete at pH values higher than 6.5. The experimental partition coefficients clearly show a preferential enrichment of heavy REE (HREE) in the solid phase. Unlike the rest of the REE, Sc sorption occurred at a lower pH, from 3 to 5. The experimental results have been described with a non-electrostatic surface complexation model in which the aqueous complex MSO_4^+ exchanges with two H⁺ from the surface of schwertmannite, forming bidentate surface complexes, $(XO)_2MSO_4^-$. Sc sorption was also accurately predicted with the addition of two bidentate surface complexes, $(XO)_2MSO_4^-$ and $(XO)_2MOH$.

The model was successfully applied to describe the behavior of different REE observed in a column experiment imitating an AMD passive remediation system. Both basaluminite and schwertmannite were the main solid phases formed in the treatment. Although both minerals have similar sorption edges, Y and the lanthanides sorbed at pH values higher than 4.5, and sorption only occurred on basaluminite because schwertmannite forms at a pH below 4. In contrast, the Sc sorption edge extends from pH 3 to 5 and therefore can partially sorb on schwertmannite. As a practical consequence, REE preferentially accumulated in basaluminite residue of AMD neutralization systems, with only a minor but a significant fraction of Sc found in schwertmannite sludge.

Keywords

Scandium, yttrium, lanthanides, non-electrostatic model, bidentate surface complex, acid mine drainage, passive remediation, neutralization sludge

1 Introduction

The rare earth elements (REE) form the largest chemically coherent group in the periodic table. According to the International Union of Pure and Applied Chemistry, REE are formed by the lanthanide series elements yttrium and scandium. Their similar atomic structure and chemical properties, as well as the small differences among the elements in the group, convert their partition between water and minerals as tracers of geochemical processes (Henderson, 1984; Elderfield et al., 1990; Johanesson et al., 1999, among many others). Beyond their importance as geochemical tracers, in the last two decades, there has been an increasing interest in REE due to their high demand by modern industry. This demand, together with a relative shortage of supply, has motivated the search for alternative sources of REE raw materials worldwide (Hatch, 2012; Massari and Ruberti, 2013).

On the other hand, acid mine drainage (AMD) is one of the most important and widespread pollutants affecting watersheds and ecosystems (Nordstrom et al., 2015). However, AMD could also become a new and supplementary source for REE, including scandium, because the concentration of REE in AMD is two to three orders of magnitude higher than that in natural waters (Noack et al., 2014). In AMD, not only pH but also aqueous speciation are key factors for REE mobility. Thus, the pH increases to circumneutral values when AMD mixes with alkaline waters, leading to the precipitation of hydrous ferric/aluminum oxides and sulfate-hydroxides, which is accompanied by REE scavenging (Verplanck et al. 2004; Gammons et al., 2005 a, b; Ferreira da Silva et al., 2009). However, despite the overwhelming presence of REE as precipitates in AMD streams (Nordstrom, 2011), there is a lack of knowledge on the interaction mechanisms between schwertmannite and REE and their dependence on pH and sulfate concentration.

In addition to REE, AMD also contains high concentrations of toxic elements and represents a very important environmental threat that requires remediation. In mine operations, active AMD treatment is mainly based on its neutralization with alkaline reagents, usually lime, limestone, sodium carbonate or ammonia (Evangelou and Zhang, 1995; Coulton et al., 2003). This treatment produces high amounts of sludge formed by a mixture of Fe(III) and Al-sulfatehydroxides, disposal of which represents a major operating cost for companies and an important environmental concern (Ackman, 1982; Viadero et al., 2006). In contrast, passive remediation systems neutralize AMD by driving it through a permeable system filled with grained limestone (Heding et al., 1994; Ayora et al., 2013). Along the process, the created waste is represented by a sequential precipitation of mainly schwertmannite [Fe₈O₈(OH)₆SO₄] and basaluminite $[Al_4SO_4(OH)_{10}\cdot 5H_2O]$, clearly identified in two separated layers. Although scandium was not included, essentially all the lanthanide and yttrium content in the inflowing AMD were retained in the treatment, particularly in the basaluminite layer (Ayora et al., 2016). Consequently, the sorption of REEs on basaluminite and the dependence of this process on pH and sulfate availability has been recently studied (Lozano et al., 2019). However, despite being a major constituent of AMD treatment sludge, the interaction between schwertmannite and REEs remains unknown.

Previous laboratory experiments of sorption and coprecipitation of REE onto ferric and manganese oxides/hydroxides have previously demonstrated the key role of pH in the geochemical mobility of REE (De Carlo et al., 1998; Bau 1999; Ohta and Kawabe 2000a; Ohta

and Kawabe 2001; Quinn et al., 2004, 2006a). REE are removed from solution as the pH increases from 4 to 9. In addition to pH, aqueous speciation is also an important factor that has been widely studied in REE fractionation in marine environments. Thus, the formation of strong carbonate complexes with heavy REE (HREE) at neutral pH values prevents their sorption by ferric and manganese oxides. As a result, these solids show patterns with light REE (LREE) enrichment (Byrne and Kim, 1990; Koeppenkastrop and De Carlo, 1992; Sholkovitz, 1995; Byrne and Sholkovitz, 1996; Kawabe et al., 1999). More recently, some works have performed detailed sorption experiments and developed surface complexation models to describe the sorption of some lanthanides (Yb, La, Eu) onto hematite and amorphous ferric oxides/hydroxides (Marmier et al., 1997; Rabung et al., 1998; Marmier and Fromage, 1999; Quinn et al., 2006; Estes et al., 2013; Liu, 2017). REE scavenging is mainly explained by the sorption of free REE ions or their hydrolyzed species onto a solid surface.

The presence of sulfate is ubiquitous in AMD systems, usually with concentrations reaching up to several grams per liter. Rather than free ions or hydrolysis products, previous studies have highlighted the importance of aqueous sulfate complexed with REEs (Johanneson and Lyons, 1995; Gimeno et al., 2000; Verplanck et al., 2004). Similar to REE-carbonate speciation in seawater, sulfate complexes may explain REE fractionation in the solids precipitated from AMD environments. Therefore, due to its importance in AMD chemistry, the role of sulfate in REE sorption on schwertmannite should be quantitatively assessed.

Here, we present the results from REE sorption experiments on synthetic schwertmannite in a sulfate-rich AMD setting. A non-electrostatic surface complexation model was developed to describe the experimental data for the sorption of a complete data set of REE (the lanthanides, Y and Sc) at different pH values, sulfate concentrations and solid:liquid ratios. The model was then applied to predict the mobility of different REE during a passive AMD neutralization system, and the results are compared with those obtained from a column experiment.

2 Materials and methods

2.1 Experimental techniques

2.1.1 Structural characterization of synthetic schwertmannite

Synthetic schwertmannite was prepared following the procedure described by Cornell and Schwertmann (1996), where 10.8 g of FeCl₃·6H₂O and 3 g of Na₂SO₄ were mixed with 2 L of distilled water at 60°C. The suspension was stirred for 12 minutes at 60°C and then cooled to room temperature. The solid was dialyzed in a cellulose membrane against double-distilled water for 33 days. Once the water had conductivity less than 0.5 mS/cm, the precipitate was freeze-dried. Mineral purity was assessed by X-ray diffraction (XRD). The specific surface area was measured by the BET-N₂ sorption method (Micromeritics Gemini V analyzer; Micromeritics, IDAEA-CSIC).

2.1.2 REE sorption as a function of pH and sulfate concentration

The REE sorption on synthetic schwertmannite was measured as a function of pH at two different sulfate concentrations in two separate sets of batch experiments using suspensions of 1 g/L. Two stock solutions with a pH of 2.5 were prepared for each set. The first stock

solution consisted of REEs at a concentration of 1 mg/L each and 20 mM SO₄, whereas the second stock solution consisted of REEs at a concentration of 1 mg/L each and 2 mM SO₄. The solutions were prepared with an ICP standard mix of 16 elements, excluding Pm, and Na₂SO₄ (ACS reagent >99.9%, Merck). The stock solutions were intended to simulate acid mine waters with different sulfate concentrations. First, 10-mL aliquots of each stock solution were transferred to 50-mL centrifuge tubes, and the pH of each aliquot was adjusted to cover a pH range of 3-7 by adding small amounts of a 0.05 M NH₄OH solution. After adding 10 mg of schwertmannite to each aliquot, the suspensions were shaken for 6 hours at room temperature. Then, the suspensions were centrifuged for 15 minutes at 4500 rpm (4150 RCF), and the supernatant was filtered through 0.22- μ m nylon membranes and transferred to test tubes diluted fivefold with 1% HNO₃. The solids were dried at 40°C for 48 hours. The sorbed fractions were calculated as follows:

Sorbed fraction
$$= \frac{[\text{REE}]_{\text{initial}} - [\text{REE}]_{\text{final}}}{[\text{REE}]_{\text{initial}}}$$
(1)

where [REE] is the concentration of each REE. The initial concentration was measured in the stock solution and the final concentration after the experiment. The distribution coefficient, K_D (L/kg), for each REE was calculated from the [REE] sorbed onto the solid per the final [REE] remaining in solution and normalized per the solid:liquid ratio:

$$K_{D} = \frac{[REE]_{initial} - [REE]_{final}}{[REE]_{final}} \cdot \frac{V_{solution}}{M_{solid}}$$
(2)

2.1.3 Description of the column experiment and waste characterization

To obtain a complete vertical profile of REE distribution in the pore water and of the minerals precipitated along a passive AMD remediation system, a scaled column experiment was carried out. The column was built with transparent polymethyl methacrylate (9.6-cm inner diameter and 35 cm in height). A 3-cm layer of glass beads (3 mm in diameter) and a perforated drainpipe were installed at the bottom of the column. Lateral ports located at 3-cm intervals in the upper part of the column allowed pore water sampling. The column was filled with 20 cm of dispersed alkaline substrate consisting of 30 g of limestone sand (0.5-1 mm) and 300 g of pine-wood shavings (5-10 mm) (Rötting et al., 2008a). The role of the pine-wood shavings was to provide a porous matrix to allocate the minerals that precipitated inside the column and prevent clogging. The proportion of limestone was predesigned to be small in order to exhaust the reagent within a few weeks. The initial total porosity of the filling was determined to be 0.46, first by weighing the column filled with only the solid phase and then again with the column pores saturated with water.

A sample of AMD from the Poderosa mine in the Iberian Pyrite Belt (SW Spain) was used as the REE-containing inflow to the column. The column was open to the atmosphere, and the input AMD was pumped into the top of the column with a peristaltic pump. The water flowed downward by gravity through the limestone substrate at a constant flow rate of 3.5 mL/min. This flow was similar to that of 2 L/s, which is the flow rate expected for a field-scale treatment of 200 m² of surface area to treat the discharge of the Poderosa mine. Water samples were collected every other week from the supernatant, and sampling points and their pH, redox potential, electrical conductivity, oxygen concentration and total alkalinity were measured in situ. The water was filtered through 0.22- μ m nylon filters and stored at 4°C until analysis. The

experiment lasted for 58 days and was stopped when the outflow showed a pH lower than 5, as an indication of reagent efficiency exhaustion.

Postmortem, the columns were drained, and the residue was divided into 2-cm slices and dried at room temperature. The major mineral phases present in the solid residue were identified by XRD. To investigate the partitioning of AI, Fe, and trace elements into the solid phases, a sequential extraction procedure adapted from Torres and Auleda (2013) was applied to the residue samples to obtain the following fractions: 1) the water-soluble fraction, extracted with deionized water; 2) calcite and basaluminite, extracted with 1 M ammonium acetate at pH 4; 3) low crystalline Fe(III)-oxyhydroxides and remaining basaluminite, extracted with 0.2 M ammonium oxalate at pH 3; 4) crystalline Fe(III)-oxides, extracted with 0.2 M ammonium oxalate at pH 3 in a water bath at 80°C for 2 hours; and 5) the residual fraction, extracted with concentrated HClO₄ and HNO₃ at 135°C. All the leachates resulting from the different extraction steps were filtered through 0.22-µm nylon syringe filters and stored at 4°C until analysis. Duplicates for 10% of the total samples were analyzed, showing a reproducibility of 90%. To validate the addition of masses obtained from all the extraction steps, a total digestion (performed according to step 5 of the protocol) of all the samples was conducted. The addition of sequential extractions was compared with the total extraction from the digestion, and only element recovery rates between 80% and 120% were considered suitable for discussion.

2.2 Analytical techniques

Measurements of pH from filtered aliquots were made with a Crison[®] glass electrode calibrated with buffer solutions of pH 2, 4, 7 and 9. Major cations were measured by ICP-AES (Thermo Scientific – iCAP 6500, radial acquisition), and REEs (La, Ce, Pr, Nd, Sm, Eu, Gd, Tb, Dy, Ho, Er, Tm, Yb, Lu, Y and Sc) were measured with ICP-MS (Perkin-Elmer[®]SciexElan 6000). The detection limits were 0.2 μ g/L for REEs and 0.1, 0.05, 0.1 and 0.1 mg/L for S, Fe, Al and Na, respectively. The analytical precision error was estimated to be approximately 2% for ICP-MS and ICP-AES measurements. Certified solutions (CPI International-CCV standard 1-solution A) were intercalated within the samples to check the analytical accuracy. Two AMD laboratory standards supplied by P. Verplank (USGS) were also analyzed for REE accuracy, yielding deviations lower than 4% of the recommended values, with the exception of Eu, which showed values 19% below the recommended value (Verplanck et al., 2001).

The XRD patterns of the solid phases after the experiments were obtained using a Bruker D8 A25 Advance Diffractometer θ - θ with CuK α 1 radiation. The diffractometer operated at 40 kV and 40 mA in a Bragg-Brentano geometry and was equipped with a linear LynxEye XE detector. The diffractograms were obtained from 4° to 60° in 2 θ with a step size of 0.015° and a counting time of 0.1 s.

2.3 Geochemical modeling

Chemical speciation of dissolved REE, saturation indexes of solid phases, surface complexation and 1D reactive transport modeling were calculated with the PHREEQC code (Parkhurst and Appelo, 1999) using the database Donnee Thermoddem_V1.10.dat (Blanc et al., 2012). The database has been updated with the stability constants at zero ionic strength and 25°C for the following inorganic complexes for the lanthanides and yttrium (M): MCO₃⁻¹, MCO₃⁻², MHCO₃⁺²

(Luo and Byrne, 2004); MSO_4^+ (Schijf and Byrne, 2004); MOH^{+2} (Klungness and Byrne, 2000); MOH_2^+ , MOH_3 (Lee and Byrne, 1992); MCI^{+2} (Luo and Byrne, 2001); MF^{+2} , MF_2^+ (Luo and Millero, 2004); and MNO_3^{+2} (Millero, 1992). Solubility products for REE M(OH)₃ and M(OH)_{3(am)} solids have been complementarily loaded from the LLNL database (Johnson et al., 1992) and Spahiu and Bruno (1995). Scandium aqueous complexes, $Sc(OH)^{+2}$, $Sc(OH)_2^+$, $Sc(OH)_3$, and $ScSO_4^+$, and the solubility product for $Sc(OH)_3$ by Wood and Samson (2006) were incorporated. The reactions and their equilibrium constants are compiled in Table S1 of the Supplementary Information.

3 Results

3.1 Characterization of synthetic schwertmannite

The XRD patterns of the solid phase prior to and after the sorption experiments confirmed that the mineral was pure schwertmannite (Fig. S1, S2). The measured specific surface area of pure schwertmannite was 173 m²/g, close to the values reported by Antelo et al. (2012) and Regenspurg et al. (2004) (171 m²/g and 175-200 m²/g, respectively). The site density was calculated from the crystal structure proposed by Fernández-Martínez et al. (2010), accounting for the single coordinated sites per face area, yielding values of 2 sites/63.22 Å² for (1 0 0), 7.5 sites/113.77 Å² for (0 1 0) and 2 sites/64.95 Å² for (0 0 1) (Fig. S3). These calculations resulted in a total site density of 4.75 sites/nm², a value similar to 4.7 sites/nm² reported by Fukushi et al. (2004). The resulting total concentration of surface sites was therefore 1367 µmol/g.

3.2 REE sorption as a function of pH and sulfate concentration

Prior to sorption experiments, the kinetics of the sorption reactions were determined using 10 suspensions of 1 g/L schwertmannite and a solution of REE at a concentration of 1 mg/L each and 20 mM SO₄ at pH 5.5. The kinetics experiments showed a very fast initial sorption, which increased until a steady state was reached at 5 hours (Fig. S4). Similar equilibrium times were observed for ferric oxyhydroxide by Ohta and Kawabe (2001) and Koeppenksastrop and De Carlo (1992), with fast REE adsorption at approximately 4-6 hours. Therefore, sorption was considered almost complete after 6 hours of contact during the sorption experiments.

The chemical composition data of the initial and final solutions for the two sets of sorption experiments with 20 and 2 mM SO₄ are compiled in Tables S2 and S3, respectively. All solutions at the end of the sorption experiments were subsaturated with respect to the REE hydroxides (Tables S4, S5), and thus the hydroxide precipitation was discarded. The solutions were also subsaturated with respect to schwertmannite and were almost in equilibrium with respect to goethite. Schwertmannite exists in a metastable phase that transforms to goethite over weeks at different pH values (Schwertmann and Carlson, 2005; Acero et al., 2006; Burton et al., 2009). However, the presence of sulfate in solution delays this transformation (Regenspurg and Peiffer, 2005), and goethite was not detected by XRD in schwertmannite solids after the sorption experiments (Fig. S1, S2).

Sorption showed a high dependence on pH in the two sets of experiments, as the REE sorption increased with pH (Fig. 1). The sorption edges occurred within a range of 2.0 to 2.5 pH units and shifted to lower pH values for HREE than for LREE, showing that HREE sorption was

favored over LREE sorption at the same pH value. Similar pH dependence was observed by Bau (1999) and Quinn et al. (2006) in REE coprecipitation with amorphous ferric hydroxide and by Lozano et al. (2019) in REE sorption on basaluminite.

The pH at which sorption becomes significant depends on the element. The lanthanides and Y started to sorb at pH 5, whereas Sc started at pH 3.5, a value considerably lower than the rest of the REE. In the experiment with 20 mM SO₄ at pH 4.9, more than 95% of Sc was adsorbed on schwertmannite, whereas less than 20% of the remaining REE were adsorbed. Similar behavior was also observed for basaluminite (Lozano et al., 2019), although here Sc sorption occurred at a pH lower than 4, at which basaluminite is not stable. The preferential Sc sorption onto schwertmannite boosts the interest of Sc recovery from passive remediation waste, during which the schwertmannite layer precipitates are separated from the basaluminite layer. Based on this selective sorption, studies have focused on Sc separation and REE recovery during a two-step adsorption process with synthetic sorbents at pH 4 and 5, as recently proposed by Ramasamy et al. (2018).



Fig. 1. Sorption edges of selected REE on schwertmannite in solutions with 20 mM (A) and 2 mM (B) sulfate. Legend: Sc: circles, Y: squares, Lu: squares and La: diamonds. The error bars indicate the uncertainties in pH measurements (\pm 0.02) and in the sorbed fraction (from \pm 0.01 to \pm 0.07).

The distribution coefficients (K_D) for both sets of experiments were calculated as described in Eq. (2) and are shown in Fig. 2. The REE fractionation patterns showed four convex segments from La to Nd, from Nd to Gd, from Gd to Ho and from Ho to Lu, corresponding to the M-type tetrad effect (Masuda et al., 1987; Bau 1996). The K_D values increased by 3 orders of magnitude with pH increasing from 4.5 to 6.5, highlighting the dependence of REE sorption on pH. At pH values higher than 5, when sorption becomes effective, the K_D values clearly showed an enrichment of HREEs in the solid phase, which increased by 1 order of magnitude along the lanthanide series. Similar REE patterns with HREE enrichments appeared in free REEs that were coprecipitated with synthetic Fe(III)-hydroxides (De Carlo et al., 1998; Quinn et al., 2004). These results contrast with those described in seawater environments, where a preferential LREE enrichment in the Fe oxides has been observed. Stable aqueous carbonate complexes with HREE formed in marine environments to prevent their adsorption (Koeppenkastrop and

De Carlo, 1992; Kawabe et al., 1999b). In acid mine waters, however, sulfate complexes are the dominant aqueous species, while carbonate complexes are only relevant at pH values higher than 6.5.



Fig. 2. Log K_D values at different pH values showing the lanthanide M-type tetrad effect: (A) solutions with 20 mM SO₄; B) solutions with 2 mM SO₄.

Bau (1999) observed an important positive Ce anomaly in hydrous ferric oxide precipitated with REE in the absence of sulfate that was attributed to the oxidation of Ce and the precipitation of CeO₂. This anomaly was not observed in the present experiments (Fig. 2), nor was it observed by Verplanck et al. (2004) in the precipitates obtained after neutralizing acid mine waters with high sulfate concentrations. This observation suggests that sulfate may inhibit Ce oxidation due to sulfate complexation.

According to its ionic radius, Y is commonly displayed between Dy and Ho, where it presents a negative anomaly (not shown in Fig. 2), as evidenced by the $K_D^{\ Y}/K_D^{\ Dy}$ and $K_D^{\ Y}/K_D^{\ Ho}$ ratios below unity (Table S6). The Y anomalies in amorphous Fe(III)-hydroxides were attributed to covalent interactions with the Fe solids (Quinn et al., 2004). The negative Gd and Yb anomalies at lower pH values in the experimental set with 20 mM SO₄ were not previously reported in the literature for similar experiments (Bau 1999; Quinn et al., 2006) and are attributed to analytical errors.

4 Discussion

4.1 Model description

In AMD environments, high sulfate concentrations are common, and dissolved sulfate strongly complexes with REE (Gimeno et al., 2000). An increase in sulfate concentration in solution leads to a higher proportion of sulfate complexes, which predominate for sulfate concentrations higher than 1 mM. Because MSO_4^+ is a predominant species in both experimental sets at 20 and 2 mM SO_4 , the sorption is similar, with only some differences between them. For the lanthanides and Y, other aqueous complexes, such as MOH^{2+} , MCO_3^+ and $M(CO_3)_2^-$, become important at low sulfate concentrations and pH values higher than 6.5 (Fig. 3A, B). Unlike Y and the lanthanides, Sc hydrolysis complexes are stronger and play a

significant role in Sc aqueous speciation at pH values higher than 4.6 (Fig. 3C, D). However, in contrast to the rest of the REE, carbonate complexes with Sc are not described due to the strong Sc affinity for hydrolysis products (Wood and Samson, 2006).



Fig. 3. Aqueous species distributions of Y and Sc (0.01 mM) with different pH values in solutions of 20 mM (A, C) or 2 mM (B, D) sulfate at 25° C and atmospheric CO₂ pressure.

Different surface complexation models have been commonly used to reproduce experimental REE sorption onto amorphous ferric hydroxide, hematite and goethite. Traditionally, REE sorption onto ferric, manganese and/or aluminum hydroxides/oxides has been explained by the process of free ion sorption onto mineral surface sites. Some models consider the electrostatic term, which accounts for the effect of surface charge on the sorption of free ions or hydrolysis products (Tochiyama et al., 1996; Rabung et al., 1998; Marmier et al., 1997; Marmier and Fromage, 1999; Quinn et al., 2006a). In contrast, a non-electrostatic model can be used when the chemical interaction of the adsorbed cations is stronger than the electrostatic forces present during sorption (Davis and Kent, 1990); therefore, the electrostatic term can be omitted. Here, a non-electrostatic model is proposed to describe REE sorption onto schwertmannite in an experimental setting. Similarly, some previous studies have proposed a non-electrostatic model to predict REE adsorption on basaltic rock (Tertre et al., 2008) and the sorption of Y and the lanthanides onto amorphous iron hydroxides (Rabung et al., 2000; Quinn et al., 2006) and onto basaluminite (Lozano et al., 2019).

Instead of developing a surface complexation model with free metal ions, the sorption of aqueous sulfate complex MSO_4^+ is proposed because sulfate complexes are predominant in the synthetic solutions prepared for these experiments (Fig. 3). In contrast, REE-carbonate aqueous species were not considered because the proportion of carbonate species is very low in the pH range and sulfate concentrations used in the experiments.

The sorption reaction consists of the exchange of the M^{2^+} aqueous complex, $(M^{2^+}$ accounting for MSO_4^+ or $Sc(OH)_2^+$) with n protons from n surface sites (XOH), reflecting the pH dependence observed in the experiments (section 3.2):

$$M^{z+} + nXOH = (XO)_n M^{z-n} + nH^+ K_M$$
 (3)

The equilibrium constant, K_M , of the surface complexation reaction for each REE aqueous complex would be

$$K_{M} = \frac{[(XO)_{n}M^{z-n}] \cdot a_{H^{+}}^{n}}{a_{M_{A}^{z+}} \cdot [XOH]^{n}}$$
(4)

where $[(XO)_n M^{z-n}]$ and [XOH] account for the mole fractions of the sorbed species ($[(XO)_n M^{z-n}]/[T_{XOH}]$) and of the free surface sites ($[XOH]/T_{XOH}$); $a_{M_4^{Z+}}$ and a_{H^+} represent the activities of the aqueous complex and proton, respectively. Considering the aqueous speciation described above, the total dissolved REE is assumed to be MSO_4^+ . A total surface site concentration, T_{XOH} , of $1.367 \cdot 10^{-3}$ mol/L was obtained from a schwertmannite site density of 4.7 sites/nm⁻² (section 3.1) and a solid:liquid ratio of 1 g/L. The equilibrium constants expressed in this work are equivalent to those referred to as K₃ in model 3 proposed by Wang and Giammar (2013), where the surface mole fraction is proposed to better describe the mass action expression for bidentate sorption. The concentration of free surface sites was calculated as the difference between the total surface site concentration and total REE sorbed:

$$[XOH] = T_{XOH} - \sum n \left[(XO)_n M^{z-n} \right]$$
(5)

Data from the experiment with 20 mM SO_4 were used to calculate the equilibrium constant. Thus, Eq. (4) was rearranged after taking logarithms, obtaining the linear expression shown in Eq. (6):

$$\log \frac{[(XO)_n M^{z-n}]}{a_{M^{z+1}}} - n \log[XOH] + (n-1) \log T_{XOH} = n pH + \log K_M$$
(6)

The amount of occupied sites with sorbed REEs was lower than 10% in the higher pH experiments and was decreased for lower pH values. Therefore, the concentration of free sites [XOH] can be approximated to that of total sites, T_{XOH} , and Eq. (6) can be simplified:

$$\log \frac{[(XO)_n M^{z-n}]}{a_{M^{z+1}}} - \log T_{XOH} = n pH + \log K_M$$
(7)

The experimental data were plotted according to Eq. (7), where the slope with respect to the pH indicated the denticity of the surface complex, and the equilibrium constant was obtained from the intercept. The slope varies from 1.4 to 2, increasing with decreasing ionic radius (Fig. 4). Therefore, bidentate surface complexes were assumed for simplicity, although a proportion

of monodentate complexes may also be present. The equilibrium constants were obtained as the intercept of a straight line of slope 2 by fitting the experimental data set of 20 mM SO₄, according to Eq. (7). Gnuplot v5.00 software (Williams and Kelley, 1986) was used for regression fitting. The error of the regression was larger and included the analytical error; therefore, it was assumed to be the error of the log K_M value.



Fig. 4. Regressions obtained from measured pH values and calculated activities from the final concentrations of the experimental data set of 20 mM SO₄. A, B, C) Regressions for Y, La and Lu, respectively, plotted as Eq. (7).

Sc aqueous speciation showed that in addition to $ScSO_4^+$, the species $ScOH^{2+}$ and $Sc(OH)_2^+$ are also relevant in the pH range studied (Fig. 3C, D). The linear regression of Eq. (7) for $ScSO_4^+$ experimental data in sorption with high sulfate concentration showed a slope close to 2 (Fig. 5A). Although the $Sc(OH)_2^+$ species became relevant from pH 5, the correlation with pH was very poor (not shown), and it was not considered. In contrast, $ScOH^{2+}$ showed a good linear regression, which resulted in a slope of 1.5 (Fig. 5B). Thus, similar to Y and to the lanthanides, $ScSO_4^+$ and $ScOH^{2+}$ were assumed to form bidentate surface complexes. Similar bidentate complexes were obtained by EXAFS fitting analysis of Lu sorbed onto ferrihydrite and Eu sorbed onto hematite (Dardenne et al., 2001; Estes et al., 2013). In contrast, REEs are found to form monodentate surface complexes in basaluminite (Lozano et al., 2019). Additionally, Rabung et al. (2006) obtained data showing that Lu formed monodentate surface complexes on γ -Al₂O₃ after fitting EXAFS spectra. This finding is probably because the Fe-O bond is longer than that of Al-O, providing more space to allocate bidentate surface complexes.



Fig. 5. Regressions obtained from experimental data of Sc plotted according to Eq. (7) for $ScSO_4^+$ (A) and $ScOH^{2+}$ (B) species.

The calculated log K_M values with the errors for the bidentate surface complexation reactions are displayed in Table 1. Equilibrium constants increased from -7.96 to -7.13 from La to Lu, indicating the higher affinity of HREEs than LREEs for the schwertmannite surface.

Element	Surface complex	Log K _M	Error
Sc	(XO) ₂ ScSO ₄	-5.27	0.07
	(XO)₂ScOH	-3.4	0.08
Y	(XO) ₂ YSO4 ⁻	-7.62	0.04
La	(XO) ₂ LaSO4 ⁻	-7.96	0.15
Ce	(XO) ₂ CeSO4 ⁻	-7.60	0.10
Pr	(XO) ₂ PrSO4 ⁻	-7.52	0.13
Nd	(XO) ₂ NdSO4 ⁻	-7.44	0.09
Sm	(XO) ₂ SmSO4 ⁻	-7.27	0.08
Eu	(XO) ₂ EuSO4 ⁻	-7.41	0.09
Gd	(XO) ₂ GdSO4 ⁻	-7.51	0.05
Tb	(XO) ₂ TbSO4 ⁻	-7.42	0.06
Dy	(XO) ₂ DySO4 ⁻	-7.27	0.03
Но	(XO) ₂ HoSO4 ⁻	-7.38	0.06
Er	(XO)₂ErSO4 ⁻	-7.39	0.05
Tm	(XO) ₂ TmSO4 ⁻	-7.17	0.08
Yb	(XO) ₂ YbSO4 ⁻	-7.26	0.04
Lu	(XO) ₂ LuSO4 ⁻	-7.13	0.09

Table 1. REE surface complexation constants for bidentate surface complexes obtained by fitting the experimental equation with a concentration of 20 mM SO₄, surface site density of 4.7 nm^{-2} , a specific surface area of $173 \text{ m}^2/\text{g}$ and a solid:liquid ratio of 1 g/L.

The sorption edge for each REE was calculated with the PHREEQC software. As expected, the calculations match the experimental data set of 20 mM SO₄ (Fig. 6). The predicted and experimental sorption edges for the complete set of REE are shown in Fig. S5. With the exception of a slight overestimation for La, the model fit the concentration of the rest of the lanthanides and Y within the experimental error (Fig. S5). Regarding Sc, the good fit observed between the model-predicted outcome and the measured from experiments confirmed the existence of two different surface species. This result highlights the different behavior of Sc sorption with respect to the rest of the REE onto schwertmannite and is a consequence of its different aqueous speciation. On the atomic scale, the Sc ionic radius is 0.75 Å (Rudolph and Pye, 2000), which is considerably lower than that of Lu, the smallest radius of the lanthanide series (0.98 Å).



Fig. 6. Experimental (symbols) and calculated fractions (lines) of the sorption curves for Sc, Y, La and Lu for the experiments with 20 mM SO₄ (left column) and with 2 mM SO₄ (right column). Calculations at 25°C and atmospheric CO₂ pressure.

4.2 Model validation

The model constructed with the 20 mM SO₄ experimental data set was validated with three additional data sets: one set with 2 mM SO₄ and a solid:liquid ratio of 1 g/L, and another two sets with 20 mM SO₄ and solid:liquid ratios of 0.25 g/L and 2 g/L, respectively.

The experimental and predicted values for Y, La, Lu and Sc for 2 mM SO₄ and 1 g/L solid:liquid ratio are plotted in Fig. 6 (right column). The predicted and experimental values for the complete set of REE are shown in Fig. S6. Although with a slight underestimation at lower pH values, the modeled curves match the experimental sorption at lower sulfate concentrations within the experimental error, confirming the formation of bidentate surface complexes.

The comparisons between the modeled and experimental data of selected elements for 1, 2 and 0.25 g/L solid:liquid ratios are shown in Fig. 7, and the data for the complete set of REE are

plotted in Fig. S7. Despite a slight overprediction at higher pH values for the 0.25 g/L solid:liquid ratio, the sorbed fractions for the different suspension concentrations were successfully predicted by the model within the experimental errors.



Fig. 7. Experimental data (symbols) and predicted sorbed fraction (lines) for sorption in 20 mM SO₄ in solution and solid:liquid ratios of 2 g/L (orange diamonds), 1 g/L (blue circles) and 0.25 g/L (green triangles). Calculations at 25°C and atmospheric CO₂ pressure.

4.3 REE mobility in an AMD neutralization column

The objective of the column experiment was to obtain a data set of the REE concentrations in the pore water and in the solid phases using a depth-time grid. Only Sc, La and Y were compared due to their different sorption-pH behaviors; additionally, La and Y are representative of the LREE and HREE, respectively, and their analyzed values were always above the detection limits. The experimental data were compared with a 1D reactive transport model, including the REE sorption model on schwertmannite described above.



Fig. 8. View of the column after 28 days of the experiment: 1= supernatant AMD; 2= schwertmannite zone; 3= basaluminite+calcite zone.

The reactive transport model considered 0.25 mol/L of calcite being dissolved by an inflow of AMD (Table S7). As the acid water infiltrates, calcite dissolves according to the kinetic law described by Plummer et al. (1978) and included in the database of the PHREEQC code. The reactive surface area was calibrated to $0.23 \text{ m}^2/\text{L}$ in order to fit the observed pH profiles. As pH rose, first schwertmannite and later basaluminite and gypsum precipitated under local equilibrium conditions. The results of the calculations for major solutes in pore water and the amount of mineral dissolved/precipitated are plotted in Fig. 9A-F. Essentially, for a given time the pH increased with depth, and for a given depth the column was progressively invaded by more acidic water. A schwertmannite front appeared first, at lower pH values (pH 2.5), followed by a basaluminite front (pH> 3.8), with both fronts advancing with time. When calcite was exhausted, the acid water dissolved the basaluminite and the Al concentration in the pore water increased above the inflow concentration (Fig. 9C).

According to Fig. 6 and 7, lanthanides and Y only sorb at pH higher than 4. Since schwertmannite only formed at lower pH, sorption of these elements was only expected in basaluminite. Unlikely, Sc could also sorb in schwertmannite. The surface complexes and log K values described in Table 1 were used to describe the Sc sorption on schwertmannite, and those of Lozano et al. (2019) were used to calculate the lanthanide and Y sorption onto basaluminite. As initial conditions, the number of sorption sites was deduced from the properties of schwertmannite (4.7 sites/nm²; 173 m²/g, section 3.1) and basaluminite (4.6 sites/nm², 68 m²/g according to Lozano et al., 2019). The solid/liquid ratio for both minerals was calculated from their total amount formed (Fig. 9D, E) distributed in the column, and assuming a bulk solid density of 1.3 kg/L (including the wood shavings) and a porosity of 0.46.



Fig. 9. Depth profiles of experimental data (symbols) and the reactive transport model (lines) at four different stages of column operation: A) pH evolution; B, C) Fe and Al concentrations in pore water; D, E, F) mineral distributions; G, H, I) concentrations of selected REE in pore water.

The differences in log K values between La and Y are small, and La and Y yield a similar distribution in pore water. In contrast, the different behavior of Sc with respect to La and Y is evident in both the measured and calculated pore water concentrations (Fig. 9 G, H, I). Thus, after two weeks, at 4 cm deep, the pH was 4.8, and Sc was completely sorbed, whereas La and Y, which had previously sorbed at an earlier time, were partially desorbing (notice their higher concentrations in the pore water than in the inflow). At 7 cm, the pH increased to 5.9, and the three REE remained sorbed. At deeper points, the pH was close to 6, and again, the three REE were still sorbed. After four weeks and at 4 cm deep, the pH was 2.6, and Sc was desorbing, whereas La and Y were already desorbed (their concentrations were equal to those in the inflow). At 7 cm deep, the pH was 4.1, and Sc was partially sorbed, as demonstrated by its concentration being lower than that in the inflow, and La and Y were desorbing, as shown again by their higher concentrations than those in the inflow. At 10 cm deep and with a pH of 5.7, Sc was entirely sorbed, whereas La and Y were still partially desorbing. After six weeks and at 4 cm deep, the pH remained at 2.5, and all three REE were completely desorbed. Meanwhile, at 7 and 10 cm deep with a pH of 3.1 and 3.4, respectively, Sc was still desorbing, whereas La and Y were practically desorbed. After 8 weeks, all three REE were desorbed at all recorded depths, with the exception of Sc, which was partially desorbed. At the end of the column (23 cm deep), the concentration of REE was very low throughout the experiment.

The column was dismantled after 58 days of functioning. The postmortem results of the column residue were in accordance with the mineral distribution predicted by the model. The presence of schwertmannite in the samples from the uppermost half of the column was marked by the amount of Fe recovered in steps II and III of the sequential extraction (Fig. 10). The presence of schwertmannite and minor goethite together with gypsum and residual calcite (data not presented) was confirmed by XRD. Although not identified on XRD due to its amorphous nature, the presence of basaluminite was evidenced by the amount of Al recovered in the second and third steps of the sequential extraction of samples of the lowermost half of the column. The presence of gypsum throughout the column was confirmed by the amount of Ca and S leached in extraction step I and the presence of calcite by the high amount of Ca present in extraction step II (data not presented). This solid phase distribution along the column is identical to that observed in other column experiments (Rötting et al., 2008a; Ayora et al., 2016) and those in two field scale treatments (Rötting et al., 2008b; Caraballo et al., 2009, 2011).

As evidenced in Fig. 10, a minor proportion of Sc and the rest of the REE was found in watersoluble step I, most likely linked to gypsum (Lin et al., 2019). However, most REE were found in steps II and III, which are representative of schwertmannite and basaluminite. However, some differences are evident in the depth distribution of Sc and the rest of the REE. More than 95% of the lanthanides and Y (ΣREY) match the Al distribution along the column depth, indicating that they are linked to basaluminite, as found in previous experiments (Ayora et al., 2016). However, up to 25% Sc was located in the first 6 cm of the column, matching the Fe distribution, and these results can be attributed to sorption on schwertmannite. According to the calculations, the pH value in the uppermost 10 cm was 2.5, predicting a complete desorption for La and Y and almost complete desorption for Sc. Therefore, the remaining REEs, and particularly Sc, analyzed in the upper 10 cm of solid residue indicate that desorption was



not entirely reversible. Indeed, desorption has also been reported to be a very slow process in experiments of Yb sorption on hematite and alumina (Marmier et al., 1997).

Fig. 10. Distribution of Fe, Sc, Al and Σ REY (except Sc) concentration in the solid phase. The numbers on the vertical axes indicate the depth (cm) in the column, and the colors indicate the different steps of sequential extraction (see text).

5 Conclusions

The experimental results confirm that the sorption of REE onto schwertmannite is highly dependent on pH. For a solid:liquid ratio of 1 g/L, the sorption of the lanthanide series and Y is significant from pH 5, and the REE were almost totally sorbed at pH 6.5, depending on the element. In contrast, the sorption of Y and the lanthanides were not affected by the amount of sulfate in solution, probably because MSO_4^+ is the overwhelming aqueous species above 1 mM SO_4 . Similar to other iron and aluminum oxides/hydroxides, preferential sorption was observed from La to Lu. Very differently, Sc sorption starts at pH 4.5, and more than 95% is retained at pH 5. The different behavior of Sc is explained by the adsorption of ScOH²⁺ in addition to ScSO₄⁺. Due to the higher importance of the hydrolyzed species, Sc sorption slightly decreased as the concentration of sulfate in solution increased.

The results of the sorption experiments can be explained by a non-electrostatic model in which MSO_4^+ (and $ScOH^{2+}$) exchanges with two surface protons, forming a bidentate surface complex. The value of the equilibrium constants of the surface complexation reactions were deduced from suspensions of 1 g/L schwertmannite in a solution with 20 mM SO_4 , and these constants were successfully applied to describe experiments with 2 mM SO_4 and with solid:liquid ratios of 2 and 0.25 g/L.

The sorption model has been applied to describe the REE behavior in a column simulating a passive AMD remediation system. As AMD infiltrates the column, calcite dissolves and schwertmannite precipitates. The progression of neutralization raises the pH above 4, and basaluminite forms. As a result, two fronts are formed: basaluminite progressing at a higher pH on the limestone column and schwertmannite behind, progressing at lower pH values. The model successfully predicts the analyses of REE in pore water and solid phases. Both the experiments and the model show different behaviors of Sc observed in the sorption experiments. Although schwertmannite and basaluminite have similar sorption edges, Y and the lanthanides only sorb at pH values higher than 4.5, and sorption only occurs on basaluminite because schwertmannite forms at pH values lower than 4. In contrast, the Sc sorption edge extends from pH 3 to 5 and therefore can partially sorb on schwertmannite. As a practical consequence, REE preferentially accumulate in the basaluminite residue of AMD neutralization systems. However, a minor but significant fraction of Sc can also be found in the schwertmannite sludge.

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